Ensemble-Based Assimilation of Satellite All-Sky Microwave Radiances Improves Intensity and Rainfall Predictions for Hurricane Harvey (2017) 2 3

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Abstract

Ensemble-based data assimilation of radar observations across inner-core regions of tropical cyclones (TCs) in tandem with satellite all-sky infrared radiances across the TC domain improves TC track and intensity forecasts. This study further investigates potential enhancements in TC track, intensity, and rainfall forecasts via assimilation of all-sky microwave radiances using Hurricane Harvey (2017) as an example. Assimilating GPM constellation all-sky microwave radiances in addition to GOES-16 all-sky infrared radiances reduces the forecast errors in the TC track, rapid intensification, and peak intensity compared to assimilating all-sky infrared radiances alone, including a 24-hour increase in forecast lead-time for rapid intensification. Assimilating all-sky microwave radiances also improves Harvey's hydrometeor fields, which leads to improved forecasts of rainfall after Harvey's landfall. This study indicates that avenues exist for producing more accurate forecasts for TCs using available yet underutilized data, leading to better warnings of and preparedness for TC-associated hazards in the future.

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- 20
- 21 Key Points:
- Satellite all-sky infrared and microwave radiances are assimilated to assess their impacts on forecasts for Hurricane Harvey.
- Along with infrared radiances, microwave radiances improve the track and intensity
 forecasts for Harvey.
- Microwave radiance assimilation leads to better analyses of the hydrometeor fields and
 more accurate rainfall forecasts.

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41 Plain Language Summary

42 Track, intensity, and rainfall are fundamental elements of all forecasts and warnings associated with tropical cyclones (TCs). Over the last few decades, the forecast community has significantly 43 improved TC track forecasts. Notable improvements in TC intensity forecasts have recently been 44 achieved using high-resolution models and remote-sensing observations over the inner-core region 45 of TCs. This study builds on these earlier efforts by investigating the impacts of utilizing 46 microwave observations on the forecast accuracy of TC track, intensity, and rainfall. Because 47 microwave radiances are sensitive to water vapor, liquid water, and ice, using these observations 48 in TC computer forecasts is expected to improve estimates of the liquid water and ice within TCs, 49 which can then lead to better rainfall forecasts. These expectations are borne out in our study's 50 tests with Hurricane Harvey. These results indicate that incorporating currently available yet 51 underutilized observations into TC computer forecasts can further improve warnings of, and 52 53 preparedness for, TC-associated hazards in the future.

54 **1 Introduction**

Tropical cyclones (TCs; see Appendix A for a complete list of acronyms) are among the most devastating natural disasters in the tropics and mid-latitudes. They make for a triple-threat of wind damage, surge inundation, and inland/freshwater flooding, the last of which is a leading cause of fatalities in the United States from TCs (Rappaport 2014). Accurate predictions of TCs are valuable to society because these predictions facilitate targeted and efficient preparations for mitigating the loss of life and property.

While forecasts of TC track and intensity have been continually improving over recent 61 decades (e.g., DeMaria et al. 2014, Cangialosi et al. 2020), one important remaining challenge is 62 the accurate prediction of hazardous TC precipitation (Kidder et al. 2005). Hazardous TC 63 precipitation events are difficult to predict because such events often result from the hard-to-64 predict TC rain bands [e.g., Hurricane Harvey (2017); Blake and Zelinsky, 2018] and long-distance 65 interactions (Galarneau et al. 2010, Meng and Zhang 2012). The forecast challenges associated 66 with the inner (e.g., Montgomery and Kallenbach 1997, Wang 2002) and outer (e.g., Diercks and 67 Anthes 1976, Chow et al. 2002) spiral rain bands are multi-faceted: spiral rain bands' existence, 68 intensity, storm-relative location, and small-scale structures are difficult to forecast accurately. 69 70 Consequently, rainfall forecasts, such as from the Weather Prediction Center (WPC), often cover

a broad area and come with an expected range of rain accumulations tagged with footnotes of
 possible localized extreme values.

Some of the most important observations of TCs over the ocean are satellite infrared (IR) 73 and microwave (MW) brightness temperatures (BTs; used interchangeably with radiance 74 hereafter). IR sensors onboard geostationary satellites provide seamless, high-spatiotemporal-75 resolution BTs of the tropics and the subtropics. They are sensitive to the absorption and emission 76 of IR radiation associated with water vapor and hydrometeors, hence provide information on cloud 77 locations, cloud-top heights, and atmospheric moisture in cloud-free regions. IR BTs are also one 78 of the critical components of the Dvorak technique for estimating TC intensity (Dvorak 1975; 79 Velden et al. 2006). While MW BTs are much less sensitive to cloud particles, they are sensitive 80 to the absorption and scattering of MW radiation associated with larger precipitation-related 81 hydrometeors. Therefore, passive MW BTs are often used in assessing TC structure and intensity 82 83 and contributing to operational products from the National Hurricane Center (NHC) that include information on low- and mid-level circulations of pre-TC disturbances that would otherwise be 84 obscured by the outflow anvil clouds of deep convection, and secondary eyewalls and potential 85 eyewall replacement cycles for mature TCs. 86

87 While IR and MW BTs are heavily used in the qualitative assessment of TCs, they are still underutilized in operational global and regional models for TC prediction (Geer et al. 2018, 88 Gustafsson et al. 2018). Recently, studies examining the ensemble-based assimilation of all-sky 89 (i.e., both clear-sky and cloud-affected) IR BTs into regional models have demonstrated its 90 potential in improving TC forecasts (Minamide and Zhang 2018, Honda et al. 2018, Zhang et al. 91 2019, Hartman et al. 2021). However, IR BTs contain little direct information on precipitation that 92 93 may exist below opaque cloud tops. For these conditions, techniques like the ensemble Kalman filter (EnKF) rely on ensemble covariances to update the model state underneath the cloud tops. 94 Unfortunately, these covariances are sometimes erroneous because of the limited ensemble size 95 (Zhang et al. 2021a, b). 96

On the other hand, MW BTs are able to reflect the distributions of hydrometeors 97 underneath the cloud tops, providing information in regions that are unobservable for the IR BTs. 98 Recent demonstrations of realistic correlations between all-sky MW BTs and TC intensity and 99 structure (Zhang et al. 2021c) motivate studying the potential benefits of simultaneously 100 assimilating all-sky MW BTs and all-sky IR BTs for the analysis and prediction of TCs. In this 101 work, we employ Hurricane Harvey (2017) as a case study. This study expands upon recent efforts 102 in employing ensemble-based assimilation of all-sky MW BTs for TCs (e.g., Wu et al. 2019; 103 Sieron 2020; Kim et al. 2020; Christophersen et al. 2021; Xu et al. 2021) by examining the impacts 104 of all-sky MW BTs on TC's track, intensity, and rainfall forecasts. 105

106 2 Methodology

For this study, we utilized the PSU WRF-EnKF data assimilation and forecast system (Zhang and Weng 2015; Weng and Zhang 2012, 2016; Zhang et al. 2009, 2011, 2016; Chen and Zhang 2019; Chan et al. 2020). The system configuration largely follows previous studies by Zhang et al. (2019) and Minamide et al. (2020), except that we adopted the Thompson (2008) microphysics scheme. Following Sieron et al. (2017, 2018), non-spherical ice-hydrometeor scattering properties consistent with the microphysics are included to realistically simulate the 113 MW BTs. AOEI (Minamide and Zhang 2017; for both IR and MW BTs) and ABEI (Minamide 114 and Zhang 2019; for IR BTs only) are applied to mitigate the deleterious impacts of strong 115 nonlinearities in the assimilation of all-sky BTs.

Because multiple studies have demonstrated that all-sky IR BT assimilation improves 116 forecasts of TC track and intensity (e.g., Zhang et al. 2016, 2019; Honda et al. 2018; Minamide 117 and Zhang 2018; Minamide et al. 2020; Hartman et al. 2021), the baseline experiment for this 118 study assimilates conventional surface and upper-air observations from the GTS, TC center 119 pressure information from TCVitals, and hourly all-sky IR BTs from channel 8 (6.2-µm) of the 120 GOES-16 ABI. This experiment is called "IR-only" hereafter. BTs from ABI's channel 8 are 121 mostly sensitive to moisture in the upper-troposphere in clear-sky regions, and our group has had 122 success assimilating them in many previous TC studies (Minamide and Zhang 2017, 2018, 2019; 123 Zhang et al. 2019; Minamide et al. 2020; Hartman et al. 2021). 124

The benefits of assimilating all-sky MW BTs are evaluated through an experiment that 125 assimilates all-sky MW BTs from the GPM constellation sensors (Hou et al. 2014; Skofronick-126 Jackson et al. 2017; see Appendix B for a complete list of assimilated channels) in addition to all 127 observations assimilated in the IR-only experiment. This second experiment is called "IR+MW" 128 hereafter. We used GPM constellation sensors' BTs in this study because they underwent extensive 129 quality control and cross-calibration. MW BTs from two channels are assimilated: the ~19 GHz 130 vertically polarized low-frequency channel ("the LF channel" hereafter; only assimilated over the 131 ocean because of uncertainties in modeled land emissivity) and the 183.31±6.6 GHz high-132 133 frequency channel ("the HF channel" hereafter; assimilated everywhere because surface contributions at this frequency are negligible for our purposes). These two channels were selected 134 for a litany of reasons (Sieron 2020): they are sensitive to liquid (the LF channel) and ice (the HF 135 channel) water contents, have the best one-to-one correspondence between water content and 136 137 changes against clear-sky BTs, have less sensitivity to non-water-content atmospheric/surface properties, have high climatological agreements between observed and simulated BTs for 138 139 precipitating regions in the EnKF priors, and have the highest frequency of occurrence across all sensors in the observing system. Of the channels in the 183-GHz family, the \pm 6.6-GHz channel is 140 chosen because its clear-sky weighting function peaks in the lower troposphere, making it 141 complementary with the ABI channel 8 IR BT whose weighting function peaks at higher altitudes 142 143 (Zhang et al. 2021c). Channels around 89 GHz are used for those sensors that do not have a channel near 183 GHz. 144

We initialize both IR-only and IR+MW experiments at 0000 UTC 22 August with 60 ensemble members that contain random perturbations generated by WRFDA and performe cycling EnKF data assimilation from 1200 UTC 22 August to 0000 UTC 25 August. Deterministic forecasts out to 0000 UTC 27 August are produced from the EnKF analysis mean every 6 hours, starting from 1800 UTC 22 August. 23 out of the 61 EnKF cycles assimilates all-sky MW radiances, 17 of which include MW BTs from both LF and HF channels and the remaining 6 cycles include only HF channel BTs.

152 **3 Results**

We first examine how the analysis-to-observation fits change from the IR-only experiment to the IR+MW experiment. We then compare the forecast performances of the two experiments in terms of their forecasts of TC Harvey's track, intensity, and rainfall amount after landfall.





Figure 1. (first column) Observed and (second and third columns) simulated BTs from the EnKF analysis ensemble mean at (a)–(i) 1200 UTC 22 August and (j)–(r) 0900 UTC 24 August for (a–c, j–l) ABI channel 8, (d–f, m–o) the MW LF channel, and (g–i, p–r) the MW HF channel.

160 3.1 Comparison of EnKF analyses

We first compare simulated IR and MW BTs from the analyses from the first EnKF cycle 161 (1200 UTC 22 August) against the assimilated observations (Figs. 1a-i), which qualitatively reveal 162 the changes with the assimilation of these observations. Both IR-only and IR+MW experiments 163 show simulated IR BTs that are qualitatively similar to the observations (Figs. 1a-c). More 164 importantly, while both experiments overestimate the coverage of the cold cloud tops within the 165 domain, the overestimation is milder for the IR+MW experiment (Fig. 1c). Furthermore, near the 166 tip of the Yucatan Peninsula, the IR+MW analysis better captured the warm LF MW BTs (Figs. 167 1d,f) and the cold HF MW BTs values (Figs. 1g,i) than the IR-only analysis (Figs 1e, h). These 168 differences in MW BTs suggest that the IR+MW analysis better captured the abundant liquid and 169 ice hydrometeors in that region. Since both experiments have identical priors at this first cycle, the 170 differences in their analyses at this time are solely associated with the assimilation of the MW BTs. 171 172 The first cycle's results thus indicate that the inclusion of MW observations can improve the analyzed hydrometeor fields. It is also worth noting that the match between the IR+MW analysis 173 and the observations is noticeably better than found in the previous studies of Wu et al. (2019). 174 We attribute this improvement to the microphysics-consistent non-spherical ice-particle scattering 175 tables developed for CRTM by Sieron et al. (2017, 2018) and the use of AOEI (Minamide and 176 Zhang 2017). 177

178 We also compared the two experiments' analyses against the IR and MW observations shortly after the onset of Harvey's rapid intensification (RI). Figures 1j-r show the observed and 179 180 simulated BTs at 0900 UTC 24 August, which is the first EnKF cycle with available MW BTs after the onset of Harvey's RI, and 8 hours after the most recent cycle that included MW BT 181 182 assimilation. At this point, clouds and rainband structures that are typical of TCs are apparent in both the IR and MW observations (Figs. 1j,m,p). The cumulative effects of the cycling EnKF 183 resulted in close matches between both experiments' simulated IR BTs (Figs. 1k,l) and the 184 observations (Fig. 1j). However, both experiments' analyses noticeably underestimated the 185 amount and areal extent of the liquid hydrometeors, indicated by the cooler-than-observed warm 186 LF MW BTs. Systematic cold biases in both experiments for the LF MW channel is beyond the 187 scope of this study but needs further investigation, and may be related to biases in the microphysics 188 scheme, as the Thompson et al. (2008) microphysics scheme is known to underpredict rainwater 189 190 (e.g., Conrick and Mass 2019).

The inclusion of the MW observations also improved the analysis in terms of the HF MW 191 channel. According to Figure 1q, the IR-only analysis exhibits a cold center that matches 192 reasonably with the observations but fails to capture the secondary cold centers to the northeast 193 and southeast of the TC center. These missing two features are associated with intense outer 194 rainbands (Fig. 1q). With the assimilation of all-sky MW radiances, these missing rainbands are 195 better captured (Fig. 1r). The primary rainband that extends southward from the TC center is 196 particularly well-represented in IR+MW. This implies that the addition of MW observations to 197 data assimilation improves the analyzed rainbands. 198

In summary, the addition of MW observations resulted in analysis improvements for both the IR and MW observations. These BT improvements indicate improvements to the analyzed structure and distribution of hydrometeors of Harvey. Next, we examine how these improvements impact Harvey's track, intensity, and rainfall forecasts.

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Figure 2. Analyses and forecasts of (first column) track, (second column) minimum sea-level pressure, and (third column) maximum surface wind speed for the (first row) IR-only and (second row) IR+MW experiments. (third row) Errors in the forecasts verified against NHC's best-track analysis.

209 3.2 Comparison of deterministic forecasts

Figure 2 shows the analyses and forecasts of Hurricane Harvey's track and intensity for the 210 IR-only and IR+MW experiments, as well as associated forecast errors with respect to the forecast 211 lead time. Both the IR-only and IR+MW experiments predict the track with reasonable accuracy, 212 213 especially for forecasts that are initialized relatively late. Additionally, the westward biases in the 1800 UTC August 22 forecast and the eastward biases in the three forecasts from 0000 UTC to 214 1200 UTC August 23 of the IR-only experiment (Fig. 2a) are noticeably reduced in the IR+MW 215 forecasts (Fig. 2d). Although reduced errors in these forecasts are diluted after averaging across 216 all 10 forecasts, the track forecast errors in the IR+MW experiment are slightly smaller, overall, 217 than in the IR-only forecasts beyond 72 h (Fig. 2g), although it is not statistically significant at 218 219 95% confidence level using a Wilcoxon signed-rank test (Wilks 2011).





Figure 3. (a, c) 700-hPa and (b, d) 850-hPa horizontal winds (barbs) and wind speeds (shading) 221 222 from the EnKF analyses of the (a, b) IR-only and (c, d) IR+MW experiment averaged every 6 hours from 1800 UTC 22 August through 1200 UTC 23 August. (e) Track of the reconnaissance 223 flight (grey) with the colored section showing SFMR surface wind speeds from 1045 UTC to 1115 224 UTC 23 August; the red star marks Harvey's center using NHC best track data. (f) Comparisons 225 of SFMR-retrieved wind speeds from 1045 UTC to 1115 UTC 23 August with those from the IR-226 only and IR+MW experiment EnKF analyses at 1100 UTC 23 August. (The numbers within the 227 228 legend represent RMSEs between the SFMR-retrieved wind speeds and those from the EnKF analysis.) 229

The forecast errors for intensity, in terms of either minimum sea-level pressure or 230 maximum surface wind speed, are also reduced when MW BTs are assimilated. There is a clear 231 bifurcation in the IR-only forecasts (Figs. 2b,c): forecasts initialized before 0000 UTC 24 August 232 are not able to capture the RI of Harvey, whereas the forecasts initialized after 0600 UTC 24 233 August do. The period from 0000 UTC to 0600 UTC 24 August is when the convection starts to 234 235 become more organized (figure not shown), contributing to the RI of Harvey shortly thereafter. For the IR-only experiment, the lack of direct information on TC organization within the IR BTs 236 may have hindered or delayed the RI of Harvey in the IR-only forecasts originating from times 237 before 0000 UTC 24 August. 238

239 The addition of MW observations resulted in forecasts that captured the RI of Harvey, even those forecasts that are initialized within 24 hours of the start of the cycling EnKF (Figs. 2e,f). 240 Furthermore, the assimilation of MW observations also resulted in forecasts with smaller mean 241 absolute errors in intensity, with the largest error reductions around 40% at 60-h forecast lead times 242 (Figs. 2h,i; statistically significant at 95% confidence level between 42 to 78 hours for minimum 243 sea-level pressure and 48 to 60 hours for maximum surface wind speed). These forecast intensity 244 improvements, especially in the early forecasts initialized before the observed RI of Harvey, likely 245 result from changes in the TC's structures introduced by all-sky MW BT assimilation. The initial 246 conditions for the first four forecasts from the IR+MW experiment have higher wind speeds 247 associated with stronger cyclonic circulation in the lower troposphere (Figs. 3c,d) compared with 248 those of the IR-only experiment (Figs. 3a,b). The higher wind speeds in the IR+MW experiment 249 better match SFMR-retrieved surface wind speed (not assimilated) from a reconnaissance flight 250 that covered the northeast quadrant of Harvey (Figs. 3e,f). A stronger cyclonic circulation in the 251 IR+MW experiment likely enabled this experiment to produce more accurate forecasts of the onset 252 of Harvey's RI than the IR-only experiment. 253

The assimilation of all-sky MW BTs also improves Harvey's rainfall forecasts. Figure 4 254 shows the accumulated rainfall forecasts from both experiments for the period from 0000 UTC 26 255 August through 0000 UTC 27 August, along with Stage-IV rainfall estimates (Lin and Mitchell 256 2005). The Stage-IV estimates reveal intense rainfall near Harvey's center as well as in the 257 rainband to the northeast of the center (Fig. 4a). Both intense rainfall regions contributed to 258 widespread flash flooding. To compare the performance of the two experiments, Equitable Threat 259 Scores (ETS; Wilks 2011) were calculated for a range of verification rainfall thresholds and 260 aggregated across all 10 forecasts. The ETS values (Fig. 4b) reveal that the IR+MW experiment 261 forecasts have more accurate rainfall predictions than the IR-only experiment forecasts at all 262 verification rainfall thresholds, ranging from almost +0.07 greater for the 5-mm threshold to more 263 than +0.04 greater for the 100-mm threshold. 264

Differences between rainfall amount forecasts and Stage-IV estimates for the two 265 experiments at two different times are also presented in Fig. 4. The 0000 UTC 23 August IR-only 266 experiment forecasts are characterized by noticeable track forecast errors (Fig. 3a); therefore, a 267 dipole structure is visible in its differences with the Stage-IV estimates (Fig. 4c). With the track 268 forecast errors reduced, the dipole structure disappears in the IR+MW experiment forecasts (Fig. 269 4e). Moreover, the severe underestimation of rainfall outside the core region in the southwest and 270 northwest quadrants relative to the core in the IR-only experiment forecasts (Fig. 4c) is greatly 271 reduced in the IR+MW experiment forecasts (Fig. 4e). This is likely the result of better analyses 272 of the TC rainbands (e.g., Fig. 2), leading to an RMSE reduction from 63.96 mm to 48.77 mm. For 273

the 0000 UTC 25 August forecasts for which both experiments have small track errors, the IR+MW experiment forecast still outperforms the IR-only experiment forecast with smaller biases, especially for the outer rainbands to the northeast over Houston. These smaller biases again led to more accurate rainfall amounts overall (Figs. 4d,f). These results show that assimilating all-sky MW BTs leads to substantial improvements in the accuracy of rainfall prediction during the landfall of TC Harvey.



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Figure 4. (a) Stage-IV total rainfall estimates accumulated from 0000 UTC 26 August through 0000 UTC 27 August. (b) Equitable Threat Scores (ETS) with different thresholds on rainfall amount from 0000 UTC 26 August through 0000 UTC 27 August for the predicted rainfall averaged over all the forecasts. Forecast minus observed rainfall amount differences from the (c,) 0000 UTC 23 August and (d,f) 0000 UTC 25 August forecasts for the (c,d) IR-only and (e, f) IR+MW experiments for rainfall amounts accumulated from 0000 UTC 26 August through 0000 UTC 27 August.

288 4 Concluding remarks

This study reveals the value of assimilating all-sky MW BTs from low-Earth-orbiting 289 satellites for improving the prediction of TC track, intensity, and precipitation through a case study 290 of Hurricane Harvey (2017). This work builds upon recent successes in improving TC prediction 291 through ensemble-based assimilation of all-sky IR BTs from geostationary satellites. Cloud-top 292 information from the IR BTs in combination with information on the hydrometeors beneath the 293 cloud tops from the MW BTs leads to better estimates of Harvey's structure. These improvements 294 from assimilating all-sky MW BT lead to more accurate track and intensity forecasts and earlier 295 accurate predictions of Harvey's RI, especially when the TC circulation was not yet well 296 established. In addition, better representation of Harvey's structure following MW assimilation 297 resulted in better rainfall forecasts after Harvey's landfall. 298

This is the first study to demonstrate improvements in track, intensity, and rainfall forecasts for a TC via assimilation of all-sky MW BTs in an ensemble-based convection-permitting data assimilation system. The influence of MW assimilation on TC prediction also depends upon AOEI, ABEI, and implementation of microphysics-consistent ice-particle scattering properties based on non-spherical ice particles.

Many challenges remain in the effective assimilation of all-sky MW BTs in support of 304 predicting TCs and their associated hazards. Appropriate adaptive bias correction and localization 305 for all-sky BT assimilation remain unresolved challenges. Comparisons of the low-frequency and 306 high-frequency MW channel BTs from different analyses suggest that the performance of 307 assimilating all-sky MW BTs using multiple channels depends on the choice of microphysics 308 schemes, which will eventually impact the performance of the subsequent forecasts. Therefore, in 309 order to better assimilate all-sky multi-channel MW BTs, there is a pressing need to develop 310 microphysics schemes that more realistically simulate hydrometeors and/or observation operators 311 that account for the uncertainties in microphysical processes. Nevertheless, our study demonstrates 312 that, despite model, observation, and data assimilation deficiencies, there are benefits from the 313 assimilation of the currently underutilized all-sky MW BTs for the prediction of TCs and their 314 associated hazards. 315

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331 Data Availability

All observations and global model analyses and forecasts are downloadable from their publicly available archives. The EnKF analyses and deterministic forecasts produced by the IRonly and MW+IR experiments are available at http://hfip.psu.edu/yuz31/Zhangetal2021GRL/.

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507	Appendix			
508	Appendix A. List of acronyms			
509	ABEI	 Adaptive Background Error Inflation 		
510	ABI	 Advanced Baseline Imager 		
511	AMSR2	 Advanced Microwave Scanning Radiometer 2 		
512	AOEI	 Adaptive Observation Error Inflation 		
513	ARW	 Advanced Research WRF model 		
514	ATMS	 Advanced Technology Microwave Sounder 		
515	BT	– Brightness temperature		
516	CRTM	- Community Radiative Transfer Model		
517	DMSP	– Defense Meteorological Satellite Program		
518	EnKF	– Ensemble Kalman filter		
519	ETS	– Equitable Threat Score		
520	GCOM-W1	– Global Change Observation Mission 1st - Water		
521	GMI	– GPM Microwave Imager		
522	GPM	- Global Precipitation Measurement project		
523	GOES	- Geostationary Operational Environmental Satellite		
524	IR	– Infrared		
525	MHS	- Microwave Humidity Sounder		
526	MW	- Microwave		
527	NHC	– National Hurricane Center		
528	NOAA	- National Oceanic and Atmospheric Administration		
529	PBL	– Planetary boundary layer		
530	PSU	– The Pennsylvania State University		
531	RI	– Rapid intensification		
532	RMSE	– Root-mean-square error		
533	RRTMG	- Rapid Radiative Transfer Model for Global Circulation Model		
534	TC	– Tropical cyclone		
535	SAPHIR	- Sounder for Probing Vertical Profiles of Humidity		
536	SFMR	- Stepped-Frequency Microwave Radiometer		
537	SSM/I	 Special Sensor Microwave/Imager 		
538	SSMIS	- Special Sensor Microwave Imager/Sounder		
539	Suomi NPP	– Suomi National Polar-orbiting Partnership		

- 540WPC– Weather Prediction Center541WRF– Weather Research and Forecasting model542WRFDA– WRF Data Assimilation system
- 543 YSU Yonsei University

Sensor	Satellite	LF Channel	HF Channel
AMSR2	GCOM-W1	7 (18.7 GHz)	13 (89.0 GHz)
ATMS	Suomi NPP		18 (183.31±7.0 GHz)
GMI	GPM Core Observatory	3 (18.7 GHz)	13 (183.31±7.0 GHz)
MHS	NOAA-18		5 (190.31 GHz)
SAPHIR	Megha-Tropiques		5 (183.31±6.6 GHz)
SSM/I	DMSP-F15	1 (19.35 GHz)	6 (85 GHz)
SSMIS	DMSP-F16, F17, F18	13 (19.35 GHz)	9 (183.31±6.6 GHz)

Appendix B. Assimilated channels from the GPM constellation sensors.