# Vapour pressure deficit is the main driver of tree canopy conductance across biomes

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#### Abstract

We aim to identify the relative importance of vapour pressure deficit (VPD), soil water content (SWC) and photosynthetic photon flux density (PPFD) as drivers of tree canopy conductance, which is a key source of uncertainty for modelling vegetation responses under climate change. We use sap flow time series of 1858 trees in 122 sites from the SAPFLUXNET global database to obtain whole-tree canopy conductance (G). The coupling, defined as the percentage of variance (R2) of G explained by the three main hydrometeorological drivers (VPD, SWC and PPFD), was evaluated using linear mixed models. For each hydrometeorological driver we assess differences in coupling among biomes, and use multiple linear regression to explain R2 by climate, soil and vegetation structure. We found that in most areas tree canopy conductance is better explained by VPD than by SWC or PPFD. We also found that sites in drylands are less coupled to all three hydrometeorological couplings with G, with wetter climates, fine textured soils and tall vegetation being associated to tighter coupling. Differences across sites in the hydrometeorological coupling of tree canopy conductance may affect predictions of ecosystem dynamics under future climates, and should be accounted for explicitly in models.

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14	Key points											
15	• Vapour pressure deficit is the main driver of tree canopy conductance globally.											
16	• Tree canopy conductance dynamics are poorly explained by the main											
17	hydrometeorological drivers in drylands.											
18	• Tree canopy conductance is more tightly couplet to hydrometeorological drivers in											
19	wetter sites, fine textured soils and tall vegetation.											

20 Abstract

We aim to identify the relative importance of vapour pressure deficit (VPD), soil water 21 content (SWC) and photosynthetic photon flux density (PPFD) as drivers of tree canopy 22 conductance, which is a key source of uncertainty for modelling vegetation responses under 23 climate change. We use sap flow time series of 1858 trees in 122 sites from the 24 SAPFLUXNET global database to obtain whole-tree canopy conductance (*G*). The 25 26 coupling, defined as the percentage of variance  $(R^2)$  of G explained by the three main hydrometeorological drivers (VPD, SWC and PPFD), was evaluated using linear mixed 27 models. For each hydrometeorological driver we assess differences in coupling among 28 29 biomes, and use multiple linear regression to explain  $R^2$  by climate, soil and vegetation structure. We found that in most areas tree canopy conductance is better explained by VPD 30 than by SWC or PPFD. We also found that sites in drylands are less coupled to all three 31 hydrometeorological drivers than those in other biomes. Climate, soil and vegetation 32 structure were common controls of all three hydrometeorological couplings with G, with 33 34 wetter climates, fine textured soils and tall vegetation being associated to tighter coupling. Differences across sites in the hydrometeorological coupling of tree canopy conductance 35 may affect predictions of ecosystem dynamics under future climates, and should be 36 37 accounted for explicitly in models.

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39 Keywords

40 biome, global, radiation, sap flow, soil water content, transpiration

## 41 **1. Introduction**

Plants regulate transpiration in response to variation in hydrometeorological conditions. 42 However, despite decades of ecophysiological research measuring responses of leaf, plant 43 or ecosystem evaporative fluxes to atmospheric dryness, soil moisture and radiation 44 (Beerling, 2015), the relative importance of these drivers in determining plant controls on 45 transpiration at the global scale is still poorly known. It is important to disentangle the 46 47 biogeographical patterns of the individual dominant drivers of transpiration control, as such drivers are expected to show spatially heterogeneous dynamics with global change (Zhou et 48 *al.*, 2019). Thus, understanding their separate roles may help improve models to anticipate 49 50 climate change impacts on vegetation function and on global water and carbon cycles, and to disentangle land-atmosphere feedbacks (Massmann et al., 2019). 51

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53 Conductance to water vapour (*G*) derived from leaf, plant or ecosystem evaporative fluxes 54 has been frequently used to describe the dynamic control of transpiration by plants at 55 different organisational and temporal scales (Jarvis & McNaughton, 1986). At short timescales, this regulation is carried out via changes in stomatal aperture. Under low soil 56 water content (SWC) or high atmospheric water demand, which is often assessed using 57 58 atmospheric vapour pressure deficit (VPD), plants reduce *G* to avoid dangerous declines in water potentials preventing physiological damage and severe dehydration (Oren et al., 59 1999). In contrast, *G* responses to light (i.e. photosynthetic photon flux density, PPFD) are 60 61 linked to plant water use efficiency (WUE). Thus, plants would increase *G* with PPFD in order to optimize photosynthesis in relation to water loss (Sperry et al., 2016). In addition, 62 PPFD effects on *G* may be driven by the need to regulate leaf temperature under high 63

radiation levels (Fauset *et al.*, 2018). These responses have been assessed in multiple, single-64 site studies (Jarvis, 1976; Oren et al., 1999; Wang et al., 2020). However, the fact that these 65 studies frequently used different phenomenological models and model-fitting approaches 66 67 complicates synthesis efforts aimed at building a common understanding of the dynamics of G at broad spatial scales. In addition, most previous work focused on overall G68 sensitivity (e.g. Hoshika et al., 2018), not on the importance of the individual drivers (but 69 70 see for instance Bretfeld et al., 2018), hampering our understanding of which hydrometeorological drivers dominate *G* regulation globally. 71

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Large-scale syntheses of the relative importance of hydrometeorological drivers regulating 73 74 transpiration have been conducted using ecosystem evapotranspiration data. Novick et al. 75 (2016) compared the limiting effect of SWC and VPD across vegetation types and climates, 76 and found that limitation on ecosystem surface conductance to water vapour caused by 77 SWC increased with climatic dryness, but that VPD was higher than SWC limitation across 78 most mesic biomes. Similarly, Han et al. (2020) also reported an increased importance of 79 SWC with increasing ecosystem aridity, but instead found that net radiation was more relevant than VPD. Conversely, Zhao et al. (2019) identified that, globally, ecosystem 80 81 evapotranspiration was not primarily limited by hydrometeorological drivers, but by vegetation height, followed by SWC and PPFD. However, these results may not reflect the 82 relative importance of hydrometeorological drivers on tree transpiration regulation, since 83 84 partitioning transpiration from total evapotranspiration can be problematic and show 85 substantial variability across ecosystems (Berkelhammer et al., 2016). Here, we overcome 86

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database of plant-level transpiration from sap flow measurements (Poyatos *et al.*, 2021).

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In this study, we investigate the hydrometeorological coupling of tree-level canopy 89 conductance by quantifying the explanatory power  $(R^2)$  of individual hydrometeorological 90 drivers of *G* (VPD, SWC, PPFD). We also estimate the total predictive ability of a model 91 including all three drivers. We then examine how the hydrometeorological coupling of G92 93 differs across biomes as a function of climate, soil properties and vegetation structure. We 94 hypothesize differences in absolute and relative G coupling to the hydrometeorological 95 drivers across biomes as a result of specific environmental constraints, with tighter 96 coupling with VPD and SWC in drier biomes with higher exposure to drought stress. We 97 also expect that climate, soil and vegetation structure determine the coupling of G with 98 VPD, SWC and PPFD, with greater coupling in sites experiencing drier conditions and 99 marked climatic seasonality, in fine textured soils associated with lower soil water 100 availability, and in tall stands with low leaf areas that are expected to have tighter coupling 101 to *G* due to thinner canopy boundary layers (Peng *et al.*, 2019).

the limitations of ecosystem-scale approaches by taking advantage of the first global

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#### 2. Methods 103

#### 2.1. Sapflow and environmental data 104

105 We extracted 1858 time series of tree sap flow from the SAPFLUXNET database (Poyatos et al., 2021). These time series met our requirements for data quality (see filtering section 106 107 below), did not include any experimental treatment (Table S1) and corresponded to 130 108 species on 122 sites (Table S2). Sub-daily sap flow time series were obtained directly in

109	sap flux density units (SFD; $[cm^3 cm^{-2}_{Asw} h^{-1}]$ ) or, when sapwood area was not available, in
110	whole-tree sap flow units (SF; $[cm^3 h^{-1}]$ ; 24 out of 122 data-sets). In those latter cases, SF
111	time series were converted to SFD units by dividing SF data by an estimation of tree
112	sapwood area (Asw) using a global allometric relationship as a function of tree basal area
113	and functional type (i.e. angiosperm vs gymnosperm) as predictors ( $R^2 = 0.78$ ; n = 2262)
114	(Fig. S1). Sub-daily SFD time series were aggregated to daytime SFD values (i.e., 6 am to
115	6 pm solar time). Following Flo et al. (2019), sap flow time series measured with non-
116	calibrated heat dissipation sensors were corrected for bias in absolute SFD multiplying by a
117	constant factor (1.405).

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Similarly to SFD, we obtained VPD [kPa] and PPFD [ $\mu$ mol m<sup>-2</sup> s<sup>-1</sup>] time series for each 119 site from SAPFLUXNET on-site measurements, which were subsequently averaged to 120 daytime values. When PPFD data were not available in the datasets (12 out of 122 sites), 121 PPFD was calculated using the mean short-wave radiation between 6 am and 6 pm 122 extracted from the ERA5 re-analyses data base (Copernicus Climate Change Service (C3S), 123 124 2017) and then multiplying by 2.3 to transform it into PPFD. Soil water content (SWC; v/v) data were missing in 43% of the SAPFLUXNET datasets included in this study. To ensure 125 126 homogeneity across sites, we used SWC from the 15-30 cm soil depth layer obtained from the ERA5-land reanalysis dataset (Copernicus Climate Change Service (C3S), 2019) at 9x9 127 km resolution (see database validation in Flo et al., 2021). 128

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130 2.2. Data filtering

In order to minimize seasonal phenological changes in leaf area, we excluded all periods 131 between 15 days before the first daytime average temperature under 0°C and 30 days after 132 the last day with temperatures under 0°C, during the cold season of each site (similar to 133 134 Novick *et al.*, 2016). To prevent artefacts in whole-tree canopy conductance calculation (Ewers & Oren, 2000), we filtered out rainy days -days when SWC increased- and days 135 when average daytime VPD was under 0.3 kPa (Anderegg *et al.*, 2018). We also ensured a 136 sufficient range in hydrometeorological conditions by discarding sites with a total VPD 137 range below 0.5 kPa or SWC range below 0.05 m<sup>3</sup> m<sup>-3</sup>, and with PPFD maximum values 138 below 400  $\mu$ mol m<sup>-2</sup><sub>Asw</sub> s<sup>-1</sup>. 139

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## 141 2.3. Whole-tree canopy conductance calculation

142 To obtain  $G_s$ , we firstly transformed SFD units from  $[\text{cm}^3 \text{ cm}^{-2}_{Asw} \text{ h}^{-1}]$  to  $[\text{Kg m}^{-2}_{Asw} \text{ s}^{-1}]$  and 143 then we converted it to daytime tree canopy conductance per unit of sapwood area  $G_{Asw}$ 144  $[\text{mol m}^{-2}_{Asw} \text{ s}^{-1}]$  following Phillips & Oren (1998) and a unit transformation (eq.1).

145 
$$G_{Asw,j,i,k} = \frac{\left(115.8 + 0.4236 T_{j,i}\right) SFD_{j,i,k}}{VPD_{j,i}} \eta \frac{T_0}{\left(T_0 + T_{j,i}\right)} e^{-0.00012h_i}$$
(1)

146 Where  $SFD_{j,i,k}$  is the sap flux density value of each site (*j*), day (*i*), and tree (*k*);  $T_{j,i}$  [°C] is 147 the temperature,  $VPD_{j,i}$  [kPa] is the daytime vapour pressure deficit,  $\eta$  equals 44.6 mol m<sup>-3</sup>, 148  $T_0$  is 273 K, and *h* [m] is the altitude of each site. For two sites where *h* values were not 149 available, it was extracted from The Shuttle Radar Topography Mission (Earth Resources 150 Observation And Science (EROS) Center, 2017).

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#### 152 *2.4. Hydrometeorological coupling quantification*

We define hydrometeorological coupling as the coefficient of determination ( $R^2$ ) of simple 153 and multiple linear mixed models of VPD, SWC and PPFD explaining *G*<sub>Asw</sub> at the site-tree 154 level. High  $R^2$  levels imply high predictive power of hydrometeorological drivers over  $G_{Asw}$ . 155 We fitted uni-variate models for each site using  $G_{Asw}$  as response variable and the neperian 156 logarithm of each driver as predictor (Fig. S2). Similarly, we also fitted additive, multiple 157 regression models of site-level  $G_{Asw}$  as a function of the logarithm of all three 158 159 hydrometeorological drivers (FULL model). The hierarchical structure of species and trees within sites was taken into account using linear mixed models, implemented with the *lmer* 160 161 function of the 'lme4' R package (Bates et al., 2015). When sites had more than one tree per species and more than one species (54 out of 122 sites), random intercept and slopes 162 parameters were fitted for species, and random intercept parameters for trees nested into 163 species. When models did not converge, the random structure was simplified and only the 164 random intercept for trees was considered (33 out of 54 sites). When sites had just one 165 166 species and multiple trees (67 out of 122 sites), we fitted a random intercept for trees. When a site had multiple species and just one tree per species (1 out of 122), random intercept and 167 168 slopes were fitted for species.

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Since we were interested in the overall coupling of all the individuals at a site, hydrometeorological coupling was set as the conditional  $R^2$  of the models (i.e.  $R^2_{VPD}$ ,  $R^2_{SWC}$ ,  $R^2_{PPFD}$ ,  $R^2_{FULL}$ ) (Table S3), calculated with the 'MuMIn' R package (Bartoń, 2020). We fitted simple and multiple regression models instead of more sophisticated non-linear models to reduce complexity and gain generalizability across the data sets.

Three alternative sets of models were fitted to ensure consistency of the estimated  $R^2_{VPD}$ , 176  $R^{2}_{SWC}$ ,  $R^{2}_{PPFD}$  and  $R^{2}_{FULL}$  values (Fig. S3). Firstly, we checked for issues related to 177 unbalanced distributions of *G*<sub>Asw</sub> throughout the range of VPD, SWC or PPFD. To do that, 178 179 we repeated the same models as above but using binned data (binned data models). Specifically, we calculated the average of  $G_{Asw}$  measurements comprised in 0.2 kPa VPD 180 intervals, five site-specific SWC intervals and 250  $\mu$ mol m<sup>-2</sup><sub>Asw</sub> s<sup>-1</sup> PPFD intervals. For each 181 summarized *G*<sub>Asw</sub> we defined a specific VPD, SWC and PPFD value as the average of the 182 data inside each bin. Secondly, to avoid possible artefacts due to the different sample size at 183 each site, models were repeated by randomly sampling 10 days per tree. We implemented a 184 185 bootstrapping approach with 100 repetitions and coupling values were calculated as the median of the R<sup>2</sup>'s distributions of each model (sampled data models). Finally, the third 186 alternative implied obtaining coupling estimates by means of more flexible generalized 187 additive models (GAM), as implemented in the bam function of the 'mgcv' R package 188 (Wood, 2011). These GAM models were fitted using all the data available per site and 189 specifying the same random structure as the simple linear models described above. The 190 alternative R<sup>2</sup>'s coupling metrics obtained with the binned data, the sampled data and the 191 gam models were all very similar to the R<sup>2</sup>'s coupling metrics from the linear models using 192 193 all data (Fig. S3), and hence the latter were used in all the following analyses.

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## 195 2.5. Biome classification and plot-level bioclimatic data

The estimates of  $G_{Asw}$  hydrometeorological coupling were complemented with site-level data on climate, soil properties and vegetation structure. These data were either directly obtained from the metadata associated to each SAPFLUXNET dataset or from additional data sources. We took from SAPFLUXNET the biome corresponding to each site –obtained from Whittaker diagrams using Chelsa Climate databases (Karger *et al.*, 2017) (Fig. S4)–
and carried out an exhaustive quality control to reassign site biomes as indicated by
SAPFLUXNET datasets contributors (Table S2). Biomes were simplified into 5 groups;
drylands (DRY), woodlands (WOOD), temperate forest (TEMP), boreal forests (BOR) and
tropical forests (TROP) (Table S3 and Fig. S4).

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For each site, we extracted climate information from global rasters (Fig. S5). We used 206 monthly mean precipitation, monthly maximum temperature and monthly minimum 207 208 temperature rasters for the period 1979 to 2013 from the Chelsa Climate databases (Karger et al., 2017), to estimate monthly potential evapotranspiration (mPET), annual potential 209 210 evapotranspiration (PET) and mean annual precipitation (MAP) using the 'envirem' R 211 package (Title & Bemmels, 2018). Then, we calculated MAP over PET (PPET) as a water availability index, and the standard deviation of the monthly differences between mean 212 213 precipitation and mPET (P-PETsd) as an index of seasonality in water availability. Relevant soil parameters were obtained from in situ SAPFLUXNET data and 214 complemented with SoilGrids 2.0 (Hengl *et al.*, 2017) when data were not available in 215 SAPFLUXNET (Table S4). We used the proportion of sand and clay particles in the fine 216 217 earth fraction [%], the total nitrogen [g kg<sup>-1</sup>] and the depth to bedrock (up to 200 cm) to characterize soils. We used bedrock depth because of its ecological relevance, but results 218 for this variable should be considered with caution due to its particularly high variability at 219 220 fine spatial scales. Stand height was available in SAPFLUXNET for most sites. When this was not the case, information was completed using the average tree height of the 221 corresponding site (again from SAPFLUXNET, 3 out of 122 sites) or when both were 222

absent it was extracted from the Global 1 km Forest Canopy Height raster (Simard *et al.*,
2011) (3 out of 122 sites) (Table S4). When site LAI was not available from
SAPFLUXNET (37 out of 122 sites), it was estimated as the average of the 95th percentile
of the period 2010 to 2016 of the MCD15A3H.006 MODIS Leaf Area Index product
(0.5x0.5 km grid) (Myneni, 2015), calculated using Google Earth Engine (Gorelick *et al.*,
2017) (Table S4).

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## 230 2.6. Statistical Analyses

In order to test whether the hydrometeorological coupling of  $G_{Asw}$  varies across biomes, we fitted weighted regressions using the modelled  $R^2_{VPD}$ ,  $R^2_{SWC}$ ,  $R^2_{PPFD}$  and  $R^2_{FULL}$  as response variables and biome as explanatory variable (fixed factor). The number of tree-days with SFD measurements in each site was used as a weighting variable. Similarly, we also tested the significance of cross-biome differences between paired hydrometeorological couplings (e.g. difference between VPD and SWC coupling,  $R^2_{VPD} - R^2_{SWC}$ ) using the same model structure.

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We further explained the biogeographical patterns in the hydrometeorological coupling across sites as a function of climate, soil properties and vegetation structure. We fitted four multiple weighted regression models with  $R^2_{VPD}$ ,  $R^2_{SWC}$ ,  $R^2_{PPFD}$  and  $R^2_{FULL}$  as response variables and log(PPET), log(P-PETsd), soil % clay, soil total nitrogen, soil bedrock depth, stand height and LAI as bioclimatic predictors (Fig. S5). We also used the number of treedays of each site as weighting variable. Sand percentage was not included due to a high correlation with soil % clay (r = -0.73). A stepwise model selection process based on minimising AIC was applied. We checked for normality and homoscedasticity of residuals
in all models, and we also checked for multicollinearity by quantifying Variance Inflation
Factors (VIF) using the 'performance' R package (Lüdecke *et al.*, 2020). These models were
combined with the rasters of bioclimatic data (at a uniform resolution of 9x9 km), to predict
and map global patterns of *G* hydrometeorological coupling to VPD, SWC and PPFD.

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We also assessed the relative importance of each hydrometeorological driver by extracting 252 the marginal partial  $R^2$  of each hydrometeorological variable in the FULL model. These 253 partial *R*<sup>2</sup> were calculated using 'r2beta' function of the R 'r2glmm' package (Jaeger, 2017) 254 and relativized by the sum of the three partial  $R^2$  (relative  $R^2$ ). These relative  $R^2$  values can 255 256 be interpreted as the relative importance of each hydrometeorological variable in explaining 257 daily variations in canopy conductance. Then, similarly as above, we fitted three multiple weighted regression models using the estimated relative  $R^2$  as response variables and the 258 259 bioclimatic variables as predictors. The resulting models were used to project the relative 260 importance of each hydrometeorological driver globally. All statistical analyses were performed in R 4.0.5 (R Core Team, 2021). 261

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### 263 **3. Results**

We found large differences in the coupling of  $G_{Asw}$  ( $R^2$  coupling metric) to each of the individual hydrometeorological drivers globally and among biomes (Fig. 1 and Table 1). We observed that  $G_{Asw}$  was predominantly coupled to VPD across biomes, whereas coupling to SWC and PPFD was comparatively less important (Fig. 1). The difference between the coupling to VPD and to the other two drivers was significant in all cases 269 except for SWC on woodlands and for both SWC and PPFD in boreal biomes (Table 1). The coupling to SWC was higher for TEMP and particularly WOOD biomes whereas 270 PPFD tended to dominate in TROP biome (marginally significant effect) (Table 1). The 271 outcomes of the linear models show a significantly higher VPD and SWC coupling for 272 TEMP, TROP and also WOOD biomes than for the DRY biome (Table. 1). The PPFD 273 coupling was also lowest for the DRY biome and was significantly higher for TEMP, BOR 274 275 and TROP biomes; the *G*<sub>Asw</sub> coupling to PPFD was also significantly lower in the WOOD biome compared to TEMP and TROP biomes (Table 1). The DRY biome was the one in 276 which all three drivers collectively (FULL model) explained less variability in  $G_{Asw}$ . 277

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In the models explaining the biogeographical patterns of  $G_{Asw}$  hydrometeorological 279 coupling (which explained 30-52% of the variance), soil and vegetation structure variables 280 were identified as common controls on the  $G_{Asw}$  hydrometeorological coupling (Table 2). In 281 particular, soil clay %, and stand height were selected for all three hydrometeorological 282 drivers (i.e. VPD, SWC and PPFD) and the FULL model, with tighter coupling always 283 associated to fine textured soils and taller vegetation. In addition, higher soil nitrogen 284 concentrations had a positive effect on VPD, PPFD and overall (FULL model) coupling. 285 286 Log(PPET) and bedrock depth were only selected for the  $R^{2}_{SWC}$  and the FULL models, in which lower climatic water availability and deeper soils were associated with looser 287 coupling (Table 2). Higher LAI was associated with lower VPD, SWC and overall (FULL 288 model) coupling, although the effect was only significant for the FULL model. Seasonality 289 in water availability (P-PETsd) was not included in any of the models. 290

292 When predictions of  $G_{Asw}$  coupling to each of the hydrometeorological drivers were mapped at the global scale, spatial patterns differed substantially among drivers (Fig. 2). G<sub>Asw</sub> 293 coupling to VPD was higher than ca. 50% almost everywhere except for some sub-tropical 294 regions. The regulation of tree water fluxes at high northern latitudes (above 50° N) and in 295 tropical regions was highly coupled to VPD, SWC and PPFD (Fig. 2). In contrast, trees 296 living in subtropical regions tended to be less coupled (Fig. 2), consistent with the lower 297 298 coupling to individual drivers in WOOD and particularly DRY biomes (Table 1). When considering the relative importance (partial  $R^2$ ) of each of the three variables in driving 299 300 canopy conductance (Fig. 3), temperate regions, drylands and savannas were typically 301 coupled to VPD, many boreal areas were coupled to PPFD, and tropical regions tended to be coupled to both VPD and PPFD (Fig. 3). Although a relevant role of SWC in 302 conjunction with VPD was identified in some regions (e.g. SW Asia, Mediterranean basin), 303 SWC did not consistently emerge as a regionally important, dominant predictor of  $G_{Asw}$ . 304

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### 306 **4. Discussion**

This study provides the first examination of the importance of VPD, SWC and PPFD as the main hydrometeorological drivers of tree canopy conductance (*G*) at the global level. All sites presented some degree of coupling to the hydrometeorological drivers considered, although there was substantial variability in the magnitude of this coupling (i.e.  $R^2_{VPD}$ ,  $R^2_{SWC}$ ,  $R^2_{PPFD}$ ). We demonstrate that *G* is predominantly coupled to VPD in most biomes, while *G* regulation caused by SWC and PPFD is generally comparably lower.

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Our results clearly identify vapour pressure deficit as the major regulator of *G* globally. This is consistent with recent reports showing that VPD limits vegetation growth at the global scale (Babst *et al.*, 2019; Yuan *et al.*, 2019), but contradicts other studies focusing on the controls of primary productivity (Jung *et al.*, 2017; Liu *et al.*, 2020) that find a global dominant role of SWC. However, caution is needed when comparing our results with these remote sensing and ecosystem-level studies, as our sample size is much smaller (and possibly spatially biased), and our approach focuses at the plant-level and uses actual transpiration data. In addition, we consider the effect of radiation availability, which has been rarely assessed in this type of studies.

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324 We did not observe the hypothesized increase in coupling to VPD and SWC in drier biomes, although, as expected, coupling to VPD and SWC was tighter than coupling to 325 326 PPFD in temperate and particularly woodland biomes. Interestingly, the importance of 327 SWC decreased in DRY biome, even if actual sensitivity to SWC was high (Fig. S2). This result contrasts to those found at the ecosystem level, showing that drier sites present larger 328 SWC control over evapotranspiration than wetter ones (Novick et al., 2016). This opposite 329 330 result between transpiration and evapotranspiration drivers importance suggests that the bare soil and the understory contribution to ecosystem surface conductance may be large 331 (Li et al., 2019) and strongly driven by SWC in DRY biome (Scott et al., 2021). 332

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The low coupling of *G* to all three hydrometeorological drivers in drylands is intriguing and may be related to the diversity of water use strategies in water limited systems, which range from drought-deciduousness to deep rooting or high hydraulic safety (e.g. Ackerly, 2004; Jacobsen *et al.*, 2007). Deep roots reaching the groundwater, for instance, could allow sufficient water supply to uncouple transpiration from hydrometeorological drivers and 339 specifically from shallow SWC (Barbeta & Peñuelas, 2017). At the other extreme, exposure to low water potentials results in early stomatal closure (Martin-StPaul et al., 2017), 340 effectively disconnecting transpiration from hydrometeorological drivers for long periods 341 342 of the year. Memory effects (Ogle *et al.*, 2015) are also likely to be more common in waterlimited systems, which may result in more complex responses of transpiration to individual 343 hydrometeorological drivers. In addition, in water-limited regions SWC can show strong 344 345 seasonal interactions with VPD (Zhou et al., 2019) and PPFD (Boese et al., 2019) (see also Figs. S6-S8), which could produce compound drought effects that would complicate 346 347 disentangling the coupling of transpiration to individual drivers. Finally, it should also be 348 noted that we focus here on relatively tall woody vegetation, as this is the one likely to be measured with sap flow sensors (Poyatos et al., 2021), and hence our analysis excludes 349 extremely arid sites likely to be totally driven by water availability (compare the grey areas 350 in our Fig. 3 with Fig. 1 in Running et al. (2004)). 351

Beyond general biome effects, differences in coupling among sites were explained by 352 353 differences in soil, stand structure, and to a lesser extent by climate. Besides the lower 354 coupling with SWC in areas with less climatic water availability, consistent with the results discussed in the previous paragraph, our results show the key importance of soil 355 356 characteristics, particularly texture, in explaining variability in *G* coupling. Trees increase the coupling to all three hydrometeorological drivers under high clay content (finer 357 texture). This result is to be expected since plants in fine textured soils would effectively 358 359 experience lower water potentials more gradually as SWC decreases, than plants in more coarsely textured soils that will experience a threshold-like transition from high to low 360 water potentials (Hillel, 1998). Hence, sandy soils dry rapidly and trees end up spending a 361

large fraction of their time at very low soil water potential, reducing the coupling. In
addition, our results indicate that deeper soils (higher bedrock depth) were associated to
lower coupling to SWC, consistent with the notion that access to deep water may uncouple
transpiration from shallow SWC (Barbeta & Peñuelas, 2017).

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Vegetation height and LAI are also important drivers of vegetation transpiration coupling. 367 VPD, SWC and PPFD limitations to *G* were typically higher in taller trees, consistent with 368 previous studies (Boese et al., 2019; Zhao et al., 2019). Tree height is associated with 369 productive areas with high resource availability, including water. In general, taller trees 370 371 have higher water transport efficiencies and lower resistance to embolism (Liu *et al.*, 2019; 372 Flo *et al.*, 2021). These traits are associated with acquisitive water use strategies and a tighter 373 stomatal control of transpiration (Klein, 2014). A similar argument can be used to explain 374 tighter coupling to VPD and radiation in areas with high soil nitrogen concentrations, as the 375 latter have been related to increased stomatal conductance (Maire et al., 2015) and greater 376 degree of stomatal control under drought (Ewers et al., 1999). Taller canopies are also more 377 aerodynamically rough and, therefore, show higher VPD coupling due to higher levels of leaf surface VPD. Interestingly, once accounting for the effect of vegetation height LAI had 378 379 a negative effect on transpiration coupling to VPD. We associate this result to the fact that higher LAI is related to lush canopy structures that would have a significant proportion of 380 the leaves effectively decoupled from the atmosphere (Zhang et al., 2016). 381

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383 Differences in coupling among sites should also reflect different water use strategies in the 384 corresponding communities. These differences are reflected in part in the climatic, soil, and

structural differences we studied. However, independently of coupling strength, these relationships had a relatively large proportion of unexplained variance, which could be related to contrasting water use strategies coexisting in the same biomes and even in the same sites (Anderegg *et al.*, 2018; Denham *et al.*, 2021). This implies that species traits should be accounted for if we aim to understand the fine-scale distribution of *G* responses (Flo *et al.*, 2021) and predict ecosystem-level responses to environmental variation.

391

392 In conclusion, we found that VPD is the main hydrometeorological driver of tree canopy 393 conductance globally but we also showed that VPD coupling did not increase in warmer sites, as found in ecosystem-level studies (Novick et al., 2016). The role of VPD in driving 394 395 transpiration regulation will likely be larger in a warmer world, given the generalised 396 increases in projected VPD (Grossiord et al., 2020). Our results indicate clear differences in 397 hydrometeorological couplings among biomes and under different environmental contexts, which likely underlie observed differences in the dynamics of vegetation water use, tree 398 399 growth and ecosystem production. Importantly, the low hydrometeorological coupling 400 observed in drylands suggests that models simulating vegetation-atmosphere fluxes may 401 fail in these regions unless additional processes are considered (Pan et al., 2020). An explicit 402 consideration of plant water transport (Anderegg & Venturas, 2020) and of differences in plant water use strategies (Flo et al., 2021) appear essential to characterize and model the 403 effects of VPD, SWC and PPFD on tree water use and their variability in space and time. 404

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414	
415	Data availability statement
416	The data that support the findings of this study are openly available at Flo, V. (2021).
417	Vapour pressure deficit is the main driver of tree canopy conductance across biomes (4.0)
418	[Data set]. Zenodo. https://doi.org/10.5281/zenodo.5517725
419	The code used to obtain the results of the study can be found at
420	https://github.com/vflo/drivers_importance
421	
422	5. References

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Table 1. Analysis of variance testing differences among biomes in the coupling (conditional 423  $R^{2}$ 's from mixed models) of tree-level water conductance ( $G_{Asw}$ ) to each of the main 424 hydrometeorological drivers: vapour pressure deficit ( $R^2_{VPD}$ ), soil water content ( $R^2_{SWC}$ ), 425 radiation  $(R^{2}_{PPFD})$  and the complete model including all drivers  $(R^{2}_{FULL})$ . The table shows the 426 427 mean coupling obtained across all sites in each biome. We also show the means of the paired differences between individual hydrometeorological couplings 428 and the corresponding statistical significance. DRY: dry and desert biomes; WOOD: woodlands 429 and shrublands; TEMP: temperate biomes; BOR: boreal and tundra; TROP: tropical and 430 subtropical biomes. Different superscript letters indicate significant (p < 0.05) Tukey tests 431 of paired differences between biomes. Asterisks indicate statistically significant differences 432 from zero for the paired differences between hydrometeorological couplings. 433

Biome	$R^2_{VPD}$		$R^{2}_{SWC}$		$R^{2}_{PPFD}$		$R^{2}_{FULL}$		$R^2_{VPD}$ - $R^2_{S}$	SWC	$R^2_{VPD}$ - $R^2_F$	PFD	$R^{2}_{SWC}$ - $I$	$R^{2}_{PPFD}$	Number of sites
DRY	0.31	A	0.183	А	0.189	А	0.389	А	0.127***	BC	0.121***	AB	-0.006	А	7
WOOD	0.436	В	0.412	В	0.281	AB	0.619	В	0.024	A	0.155***	В	0.131**	** B	29
TEMP	0.461	В	0.389	В	0.358	С	0.544	В	0.072***	AB	0.103***	А	0.031*	А	70
BOR	0.575	ABC	0.45	AB	0.481	BC	0.603	AB	0.124	ABC	0.093	AB	-0.031	AB	8
TROP	0.601	С	0.4	В	0.457	С	0.627	В	0.201***	С	0.144***	AB	-0.057.	A	8

434 Statistical significant levels: "." p<0.1 ; "\*" p<0.05; "\*\*" p<0.01; "\*\*\*" p<0.001

Table 2. Parameters of the models explaining  $G_{Asw}$  coupling to VPD, SWC, PPFD and to all three hydrometeorological drivers ( $R^2_{VPD}$ ,  $R^2_{SWC}$ ,  $R^2_{PPFD}$  and  $R^2_{FULL}$ , respectively) as a function of climatic, soil and stand structure variables. log(PPET): logarithm of precipitation over potential evapotranspiration [% log(mm mm<sup>-1</sup>)<sup>-1</sup>]; log(P-PETsd): logarithm of the standard deviation of the difference between precipitation and potential evapotranspiration [% log(mm)<sup>-1</sup>]; Clay percentage [% %<sup>-1</sup>clay]; Total Nitrogen [% (Kg g<sup>-1</sup>)<sup>-1</sup>]; Bedrock depth [% cm<sup>-1</sup>]; Stand Height [% m<sup>-1</sup>]; LAI: leaf area index [% (m<sup>2</sup> m<sup>-2</sup>)<sup>-1</sup>]. NI means that the variable was not included in the model after model selection. The R<sup>2</sup> of each multiple regression is also shown.

		Climate		Soil			Vegetation		
Variable	Intercept [%]	log(PPET)	log(P-PETsd)	Clay	Total Nitrogen	Bedrock depth	Stand Height	LAI	R <sup>2</sup>
$R^2_{VPD}$	25.927 ***	NI	NI	0.314 ***	3.522 ***	NA	0.746 ***	-1.524.	0.297
$R^{2}_{SWC}$	61.591 ***	11.692 ***	NI	0.475 ***	NI	-0.174 *	0.429 ***	-1.206 ns	0.521
$R^{2}_{PPFD}$	10.741 ***	NI	NI	0.230**	2.873 ***	NI	0.762 ***	NI	0.365
$R^{2}_{FULL}$	80.342 ***	5.932 *	NI	0.470***	1.849 *	-0.198 *	0.595 ***	-1.961 *	0.351

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Statistical significant levels: "." p<0.1 ; "\*" p<0.05; "\*\*" p<0.01; "\*\*\*" p<0.001; ns not significant.



Figure 1. Bi-variate and uni-variate distributions of the coupling of  $G_{Asw}$  to the hydrometeorological drivers studied: vapour pressure deficit ( $R^2_{VPD}$ ), soil water content ( $R^2_{SWC}$ ) and radiation ( $R^2_{PPFD}$ ) for different biomes. Points correspond to site-level modelled conditional  $R^2$  values. Colours represent different biomes, DRY: dry and desert biomes; WOOD: woodlands and shrublands; TEMP: temperate biomes; BOR: boreal and tundra; TROP: tropical and subtropical biomes. Dashed black line shows 1:1 relation.



Figure 2. Global projection of  $G_{Asw}$  coupling to VPD, SWC and PPFD ( $R^2_{VPD}$ ,  $R^2_{SWC}$  and  $R^2_{PPFD}$ , respectively), obtained from regression models of each coupling as a function of climatic, soil and stand structure variables (left panels). Right panels show projected Standard Error of the corresponding model.



447 Figure 3. Relative importance (partial  $R^2$ ) of the three hydrometeorological drivers of transpiration regulation calculated from the complete (FULL) model, and projected at the 448 449 global scale using linear models with climate, soil and vegetation structural variables as explanatory variables. Grid values were calculated using the 'tricolore' package (Schöley & 450 451 Kashnitsky, 2020) for each cell as the relative value of the projections of the relative 452 importance of each hydrometeorological variable. Colour gradient indicate the relative importance of the three hydrometeorological constraints. Light grey colour are deserts or 453 non-forested areas. % VPD: vapour pressure deficit relative importance. % SWC: soil water 454 content relative importance. % PPFD: photosynthetic photon flux density relative 455 importance. Points indicate locations of study sites. 456

Supporting Information for

#### Vapour pressure deficit is the main driver of tree canopy conductance across biomes

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### Contents of this file

Figures S1 to S8

Tables S1 to S4



Figure S1: SAPFLUXNET global scaling relationship between basal area and sapwood area. Basal area and sapwood area are both in cm<sup>2</sup>. Shaded areas are 95% model confidence interval.



Figure S2: Log relationships of the three environmental variables estimated with the FULL model (VPD + SWC + PPFD) and grouped by biome. Coloured lines are biome-averaged models calculated from LMM predictions with  $G_{\text{Asw}}$  as response variable and the neperian logarithm of the environmental constrains as explanatory variables. Dashed line shows standard error of the average models calculated with bootstrap prediction using 100 simulations.



Figure S3: Comparison of biomes  $R_{FULL}^2$  (VPD + SWC + PPFD),  $R_{VPD}^2$  (VPD),  $R_{SWC}^2$  (SWC) and  $R_{PPFD}^2$  (PPFD) calculated with the four modelling approaches: using linear mixed models on all data, binned data and sampled data and, using GAM models on all data. Dots are mean values for each biome and error bars represent two standard deviations. There were no intra-biome significant differences among models.



Figure S4: Bioclimatic distribution of the SAPFLUXNET datasets used in the study. Points show the different datasets in a Whittaker diagram showing the classification of the aggregated biomes used in the study.



Figure S5: Global projection of climatic, soil and stand structure variables. log(PPET): logarithm of precipitation over potential evapotraspiration  $[log(mm mm^{-1})]$ ;  $log(P-PET_{sd})$ : logarithm of the standard deviation of the difference between precipitation and potential evapotranspiration [log(mm)]; Clay: percentage of clay in the soil; Total N: total nitrogen in the soil  $[g kg^{-1}]$ ; Bedrock [cm]; Stand height [m]; LAI: leaf area index  $[m^2 m^{-2}]$ . Total N values above 5 g kg<sup>-1</sup> were truncated.



Figure S6: Correlations at the site level between VPD and SWC. R values are Pearson correlations. Different colors indica and desert biomes; WOOD: woodlands and shrublands; TEMP: temperate biomes; BOR: boreal and tundra; TROP: tropica



Figure S7: Correlations at the site level between VPD and PPFD. R values are Pearson correlations. Different colors indicand desert biomes; WOOD: woodlands and shrublands; TEMP: temperate biomes; BOR: boreal and tundra; TROP: tropica



Figure S8: Correlations at the site level between PPFD and SWC. R values are Pearson correlations. Different colors indic and desert biomes; WOOD: woodlands and shrublands; TEMP: temperate biomes; BOR: boreal and tundra; TROP: tropica

Plot treatment
NA
None
Control
control
Ambient Control
Control - Unthinned
natural conditions
Reference
1Premortality
2premortality
distructive sampling
Girdling early successional
Pre-thinning
Before thinning
Before Thinning
non thinned
none (periodict thinning every 5-6 years 20 to 25% of basal area)
Radiation Level
AMBIENT CO2 FACE rings
fertilization at plantation
AcaciaMonoculture
MixtureEucalyptusAndAcacia
EucalyptusMonoculture
Pre Irrigation

Table S1: SAPFLUXNET stand treatments included in the this study.

Site code	Latitude	Longitude	Biome	# Tree-days	# Species	$\#~{\rm Trees}$
AUS_CAN_ST1_EUC	-37.58	149.17	WOOD	500	1	12
AUS_CAN_ST2_MIX	-37.58	149.17	WOOD	1077	2	22
AUS_CAN_ST3_ACA	-37.58	149.17	WOOD	609	1	12
AUS_CAR_THI_CON	-38.38	146.68	TEMP	69	1	3
AUS_ELL_UNB	-36.78	146.58	TEMP	140	1	2
AUS_MAR_UBD	-37.69	145.56	TEMP	32	2	2
AUS_MAR_UBW	-37.89	145.57	TEMP	121	3	5
AUS_WOM	-37.42	144.09	TEMP	4454	2	11
AUT_PAT_FOR	47.21	11.45	BOR	286	1	3
AUT_PAT_KRU	47.21	11.45	BOR	105	1	2
AUT_PAT_TRE	47.21	11.45	BOR	133	1	3
BRA_CAM	-22.69	-45.52	TROP*	89	1	5
BRA_CAX_CON	-1.79	-51.43	TROP	2406	8	15
CAN_TUR_P39_PRE	42.71	-80.36	TEMP	2225	1	18
CAN_TUR_P74 CUN_ABC_CWS	42.71	-80.35	TEMP	10586	1	16
CHN_ARG_GWS	41.38	89.94	DRY	300	1	2
CHN_HOR_AFF	42.72	122.37	DBY	2780	1	10
CHN_YIN_SII	42.40	80.72	TROD	122	1	0
CZE DIL DIL	10.39	-84.03	TEMD*	1200	17	20
CZE_BIL_BIL	49.20	16.09	TEMP*	400	1	6
CZE_KRI_KRI	49.52	16.75	TEMP*	450	2	17
CZE BAL BAL	40.00	16.55	TEMP*	360	1	6
CZE SOB SOB	40.25	16.60	TEMP*	1727	1	6
CZE_SOB_SOB	49.04	17.07	TEMP	348	1	8
CZE UTE BPO	49.28	16.65	TEMP*	456	1	6
DELL HIN OAK	53.33	13.19	TEMP*	864	1	8
DEU HIN TER	53.33	13.19	TEMP*	1954	2	16
DEU MER BEE NON	49.27	7.81	TEMP	841	1	8
DEU MER DOU NON	49.27	7.81	TEMP	895	1	7
DEU MER MIX NON	49.27	7.81	TEMP	1945	2	17
DEU STE 2P3	53.10	13.00	TEMP*	1228	1	10
DEU STE 4P5	53.10	13.00	TEMP*	402	1	10
ESP ALT ARM	40.78	-2.33	WOOD	8306	3	15
ESP ALT HUE	40.79	-2.29	WOOD	3698	2	8
ESP_ALT_TRI	40.80	-2.23	WOOD	5411	2	12
ESP_CAN	41.43	2.07	WOOD	6871	4	21
ESP_GUA_VAL	40.90	-4.03	WOOD	3424	1	24
ESP_LAS	28.31	-16.57	WOOD	4406	1	10
ESP_MAJ_MAI	39.94	-5.77	WOOD	2833	1	6
ESP_MON_SIE_NAT	41.12	-3.50	WOOD	2587	3	20
ESP_RIN	40.60	-6.02	WOOD	770	1	8
ESP_RON_PIL	36.69	-5.02	TEMP	4114	2	12
ESP_TIL_MIX	41.33	1.01	WOOD	15699	2	32
ESP_TIL_OAK	41.33	1.01	WOOD	2381	1	10
ESP_TIL_PIN	41.33	1.01	WOOD	1976	1	9
ESP_VAL_BAR	42.20	1.82	WOOD	1394	1	12
ESP_VAL_SOR	42.20	1.81	WOOD	1943	1	13
ESP_YUN_C1	36.72	-4.97	WOOD	2935	1	6
ESP_YUN_C2	36.72	-4.97	WOOD	830	1	6
FIN_HYY_SME	61.85	24.29	TEMP	10	1	1
FIN_PET	69.49	27.23	BOR*	216	1	7
FRA_FON	48.48	2.78	TEMP*	720	1	3
FRA_HES_HE1_NON	48.67	7.06	TEMP	1273	1	10
FRA_HES_HE2_NON	48.67	7.06	TEMP	4167	1	10
FRA_PUE	43.74	3.60	WOOD	23566	1	25
GBR_ABE_PLO	56.62	-3.80	TEMP	692	1	15
GBR_DEV_CON	56.03	-3.72	TEMP*	215	1	4
GBR_GUI_STI	57.27	-4.82	TEMP	834	1	15

Table S2: SAPFLUXNET sites included in the study. Biome was estimated using a Whittaker diagram. \*Indicates that the biome was manually adjusted and confirmed by SAPFLUXNET contributors.

Site code	Latitude	Longitude	Biome	# Tree-days	# Species	$\#~{\rm Trees}$
GBR_GUI_ST2	57.27	-4.82	TEMP	621	1	9
GBR GUI ST3	57.27	-4.82	TEMP	444	1	8
GUF_GUY_GUY	5.28	-52.92	TROP	710	6	6
GUF_GUY_ST2	5.28	-52.91	TROP	885	7	11
GUF NOU PET	4.08	-52.68	TROP	923	10	22
HUN SIK	47.93	20.44	WOOD	550	2	4
ISR YAT YAT	31.34	35.05	DRY	15766	1	24
ITA FEI S17	46.69	10.61	TEMP	378	1	6
ITA KAE S20	46.70	10.61	BOR	586	1	6
ITA MUN	46.68	10.58	TEMP*	885	1	6
ITA REN	46.59	11.43	TEMP	577	3	8
ITA RUN N20	46.70	10.64	BOR	766	2	8
MEX COR YP	19.49	-97.04	TEMP	113	1	7
NLD LOO	52.17	5.74	TEMP*	3033	1	6
NLD SPE DOU	52.25	5.69	TEMP*	150	1	3
NZL HUA HUA	-36.80	174.49	TEMP	1107	1	6
PRT LEZ ARN	38.83	-8.82	WOOD	1764	1	4
PRT MIT	38.54	-8.00	WOOD	1510	1	4
PRT PIN	38.25	-8.76	WOOD	2991	2	20
RUS CHE Y4	68.74	161.41	BOR	587	1	11
BUS FYO	56.46	32.92	TEMP	2338	3	17
RUS POG VAR	56.36	92.95	TEMP*	1290	3	9
SEN SOU PRE	16.34	-15.43	DBY	1706	1	3
SWE NOR ST1 BEF	60.09	17.48	TEMP*	843	2	22
SWE NOD ST2	60.00	17.48	TEMD*	104	2	6
SWE_NOR_ST2	60.09	17.48	TEMP*	1105	2	37
SWE NOR ST5 REF	60.08	17.48	TEMP*	1255	2	25
SWE SKO MIN	58.26	12.15	TEMP	1072	1	20
SWE_SKO_MIN	60.12	17.84	TEMP*	455	1	12
SWE_SKI_381	60.10	17.82	TEMP*	1046	2	12
SWE SVA MIX NON	64.26	10.77	TEMP	1540	2	20
THA KHU	15.27	103.08	TROP	1278	1	20 6
USA BNZ BLA	64.70	-148.32	BOB*	1705	1	6
USA_CHE_ASP	45.94	-00.27	TEMP	2005	Ê	142
USA CHE MAD	45.05	-00.27	TEMP	3333	2	145
USA DUK HAD	36.08	-90.20	TEMP	601	6	22
USA UIL UE2	30.98	-19.09	TEMP	021	5	33
USA_INM	30.22	-76.60	TEMP	1648	6	0
USA MOR SE	20.22	-86.41	TEMP	680	4	6
USA NWH	24.59	-00.41	TEMP	525	3	10
USA ODN ST1 AMP	25.00	-91.20	TEMP	210	1	8
USA DAD FFD	35.90	-04.33	TEMP	726	1	8
USA_FAR_FER	30.80	-70.07	TROP	19190	1	80
USA_FER_FER USA_PIS_P04_AMB	30.21	-03.07	DRV	12120	2	10
USA DIS DOS AMB	24.20	-106.53	DRV	10754	2	10
USA DIS DI2 AMB	24.20	-106.53	DRV	19598	2	10
USA_SIL_OAK_1PR	20.02	-74.60	TEMP	2204	4	18
USA_SIL_OAK_IPR	39.92	-74.00	TEMP	6700	4	22
USA_SIL_OAK_2FK	39.92	-74.00	TEMP	1799	48 E	22
USA_SWI_SER	24.11	-70.00	TEMP	056	3	30
USA SVI. HL1	46.94	-80.25	TEMP	8620	2	10
USA SVI UIO	46.24	-80.25	TEMP	3679		20
USA TND	96.47	-09.00	TEMP	1940	4	20
USA_TND	30.47	-04.70	TEMP	1240	41 E	0
USA_TND	30.97	-04.28	TEMP	12/1	0 F	9
USA_IMP_CON	30.90 45.56	-84.29	TEMP	1342	0	9
USA_UMB_CUR	45.50	-04.71	TEMP	10175	0	57
USA_UMB_GIK	45.00	-84.70	TEMP	19170	41 E	16
USA_WIL_WOI	20.02	-90.09	TEMP	1300	0	010
USA_WVP	39.00	-19.09	1 EAVIP	989	0	0

Table S2: SAPFLUXNET sites included in the study. Biome was estimated using a Whittaker diagram. \*Indicates that the biome was manually adjusted and confirmed by SAPFLUXNET contributors. (continued)

Site code	Latitude	Longitude	Biome	# Tree-days	# Species	$\#~{\rm Trees}$
ZAF_FRA_FRA	-33.88	19.06	WOOD	563	1	3
ZAF_RAD	-34.08	19.11	WOOD	660	1	3
ZAF_SOU_SOU	-34.09	19.09	WOOD	424	1	2
ZAF_WEL_SOR	-33.48	18.96	WOOD*	538	1	3

Table S2: SAPFLUXNET sites included in the study. Biome was estimated using a Whittaker diagram. \*Indicates that the biome was manually adjusted and confirmed by SAPFLUXNET contributors. (continued)

Original biome name	Study biome group
Desert	DRY
Temperate grassland desert	DRY
Subtropical desert	DRY
Woodland/shrubland	WOOD
Temperate forest	TEMP
Boreal forest	BOR
Tundra	BOR
Tropical rainforest	TROP
Tropical seasonal forest/savanna	TROP

Table S3: Table of equivalence between Whittaker biomes and the groups of biomes used in the study.

Table S4: Summary table of site level  $R_{VPD}^2$ ,  $R_{SWC}^2$ ,  $R_{PPFD}^2$ , climate, soil properties and vegetation structure data. PPET is in [mm mm<sup>-1</sup>], P-PET<sub>sd</sub> is in [mm], Clay and Sand are in [%], Total N is in [g kg<sup>-1</sup>], Stand height is in [m], LAI is in [m<sup>2</sup><sub>leaves</sub> m<sup>2</sup><sub>soil</sub>]. Letters show data source: a = SAPFLUXNET, b = Global rasters, c = SAPFLUXNET plant height.

				RD		(PED		6.0				
Shecode	RUPD	RESWO	Ropped	Reling	Relinip	Relinp	PPET	P-PEL	Class	Sand	Total	Bedro
AUS_CAN_ST1_EUC	0.77	0.49	0.60	0.66	0.34	0.00	1.23	47.52	26.30 b	45.10 b	1.02	184
AUS_CAN_ST2_MIX	0.83	0.62	0.72	0.83	0.17	0.00	1.23	47.52	26.30 b	45.10 b	1.02	184
AUS_CAN_ST3_ACA	0.83	0.69	0.75	0.84	0.15	0.02	1.23	47.52	26.30 b	45.10 b	1.0.2	184
AUS_CAR_THI_CON	0.41	0.00	0.07	0.81	0.05	0.14	1.36	49.01	27.20 b	44.30 b	2.34	111
AUS_MAD_UDD	0.86	0.46	0.76	0.98	0.00	0.02	1.08	67.10	26.70 b	48.50 D	1.95	63
AUS_MAR_UBD	0.81	0.23	0.37	0.82	0.15	0.04	1.35	70.37	20.00 b	44.00 D	1.90	89
AUS_MAR_OBW	0.90	0.78	0.81	0.89	0.00	0.11	1.21	60.38	27.90 D	43.90 D	2.00	173
AUT DAT FOR	0.77	0.32	0.64	0.19	0.00	0.20	2.17	16.78	20.90 B	60.00 a	2.04	180
AUT PAT KRU	0.43	0.05	0.04	0.84	0.00	0.00	2.17	16.78	5.00 a	60.00 a	3.94	180
AUT PAT TRE	0.56	0.27	0.20	0.70	0.30	0.00	2.17	16.78	5.00 a	60.00 a	3.94	180
BRA CAM	0.84	0.70	0.70	0.65	0.27	0.07	1.66	88.82	27.60 h	52.00 h	2.26	200
BRA CAX CON	0.75	0.68	0.68	0.73	0.00	0.27	1.90	122.90	8.00 a	79.00 a	1.45	197
CAN TUR P39 PRE	0.49	0.33	0.31	0.72	0.09	0.19	1.39	42.08	1.00 a	98.00 a	1.58	200
CAN TUR P74	0.21	0.32	0.16	0.31	0.33	0.35	1.39	41.87	1.00 a	98.00 a	1.60	200
CHN_ARG_GWS	0.45	0.34	0.35	0.52	0.45	0.03	0.01	63.51	17.70 b	46.00 b	0.70	172
CHN_HOR_AFF	0.37	0.33	0.32	0.51	0.47	0.01	0.59	31.24	8.00 a	83.00 a	1.00	200
CHN_YIN_ST1	0.46	0.46	0.44	0.48	0.41	0.11	0.19	35.09	20.80 b	32.90 b	2.41	148
CRI_TAM_TOW	0.68	0.68	0.67	0.47	0.15	0.37	3.57	159.99	36.10 b	34.70 b	2.75	200
CZE_BIL_BIL	0.53	0.53	0.41	0.52	0.20	0.29	0.71	28.98	29.60 b	27.40 b	1.91	200
CZE_KRT_KRT	0.61	0.50	0.30	0.62	0.04	0.34	0.85	27.00	26.00 b	27.40 b	2.10	200
CZE_LAN	0.74	0.73	0.69	0.67	0.07	0.26	0.66	37.49	17.80 a	71.80 a	2.46	200
CZE_RAJ_RAJ	0.35	0.36	0.39	0.37	0.03	0.60	0.99	26.14	21.80 b	33.90 b	1.96	200
CZE_SOB_SOB	0.40	0.43	0.13	0.49	0.30	0.21	0.71	28.98	29.60 b	27.40 b	1.91	200
CZE_STI	0.52	0.36	0.43	0.69	0.23	0.09	1.13	27.10	34.20 a	47.60 a	1.65	200
CZE_UTE_BPO	0.64	0.67	0.53	0.41	0.17	0.42	0.75	29.86	26.70 b	23.80 b	2.71	200
DEU_HIN_OAK	0.36	0.18	0.29	0.93	0.06	0.01	0.95	35.10	17.90 b	49.90 b	2.4.2	200
DEU_HIN_TER	0.22	0.17	0.18	0.74	0.02	0.24	0.95	35.10	18.00 b	50.50 b	2.05	200
DEU_MER_BEE_NON	0.38	0.25	0.27	0.80	0.05	0.15	1.48	47.33	4.00 a	71.00 a	2.56	200
DEU_MER_DOU_NON	0.38	0.25	0.18	0.58	0.29	0.14	1.48	41.33	4.00 a	71.00 a 71.00 a	2.50	200
DEU_MER_MIX_NON	0.33	0.21	0.25	0.83	0.01	0.14	0.00	41.55	2.50 a	02.50 a	2.00	200
DEU STE 4P5	0.45	0.28	0.23	0.64	0.32	0.03	0.90	37.45	2.50 a	92.50 a	3.28	200
ESP ALT ARM	0.44	0.39	0.29	0.77	0.19	0.04	0.66	65.80	21.90 b	41.50 h	1.27	187
ESP ALT HUE	0.42	0.27	0.22	0.77	0.00	0.23	0.51	63.08	21.60 b	35.90 b	1.46	200
ESP ALT TRI	0.48	0.42	0.26	0.67	0.21	0.13	0.57	63.48	21.00 b	40.00 b	1.31	196
ESP CAN	0.51	0.44	0.35	0.62	0.09	0.28	0.94	46.91	32.90 b	28.30 b	1.76	179
ESP GUA VAL	0.50	0.30	0.25	0.71	0.00	0.29	0.68	69.09	24.80 b	40.90 b	1.27	200
ESP LAS	0.24	0.26	0.07	0.43	0.45	0.13	1.63	37.89	1.00 a	70.00 a	1.65	197
ESP_MAJ_MAI	0.53	0.47	0.28	0.69	0.15	0.16	0.76	97.45	9.00 a	80.00 a	1.18	200
ESP_MON_SIE_NAT	0.38	0.33	0.40	0.49	0.05	0.45	0.62	63.05	20.80 b	41.90 b	1.4.5	200
ESP_RIN	0.82	0.61	0.60	0.96	0.00	0.04	0.85	76.30	15.00 a	9.00 a	2.17	200
ESP_RON_PIL	0.34	0.26	0.16	0.59	0.11	0.31	1.05	93.66	18.00 a	30.00 a	1.86	200
ESP_TIL_MIX	0.40	0.42	0.30	0.43	0.17	0.40	0.77	48.10	20.00 a	60.00 a	1.44	162
ESP_TIL_OAK	0.16	0.37	0.18	0.29	0.35	0.37	0.77	48.10	20.00 a	60.00 a	1.44	162
ESP TIL PIN	0.20	0.39	0.06	0.28	0.55	0.17	0.79	48.10	20.00 a	60.00 a	1.78	188

Table S4: Summary table of site level  $R^2_{\rm VPD}, R^2_{\rm SWC}, R^2_{\rm PPFD},$  climate, soil properties and vegetation structure data. PPET is in [mm mm^{-1}], P-PET\_{\rm sd} is in [mm], Clay and Sand are in [%], Total N is in [g kg^{-1}], Stand height is in [m], LAI is in [m^2\_{leavest} m^2\_{\rm soil}]. Letters show data source: a = SAPFLUXNET, b = Global rasters, c = SAPFLUXNET plant height. (continued)

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SHEDDE	RUPD	RENC	Reper	Relimp	Relinip	Relimp	PPET	8. P.B.	Class	Sand	TOBIL	Bedro
ESP_VAL_BAR	0.56	0.24	0.27	0.87	0.00	0.13	0.70	34.07	32.63 a	9.81 a	1.94	200
ESP_VAL_SOR	0.50	0.32	0.26	0.66	0.10	0.24	0.78	32.15	20.00 a	60.00 a	2.04	200
ESP_YUN_C1	0.29	0.44	0.17	0.17	0.67	0.16	0.83	93.65	29.00 a	22.00 a	1.37	197
ESP_YUN_C2	0.27	0.61	0.24	0.22	0.43	0.35	0.78	91.33	29.00 a	22.00 a	1.37	188
FIN_HYY_SME	0.51	0.01	0.10	0.61	0.10	0.29	1.20	38.32	6.50 a	37.00 a	1.67	200
FIN_PET	0.63	0.63	0.58	0.40	0.60	0.01	1.13	26.34	7.30 b	60.80 b	5.08	200
FRA_FON	0.68	0.62	0.64	0.72	0.22	0.06	0.89	45.10	19.00 a	37.00 a	1.26	200
FRA_HES_HE1_NON	0.42	0.53	0.43	0.41	0.26	0.33	1.31	47.72	25.00 a	8.00 a	1.41	200
FRA_HES_HE2_NON	0.25	0.35	0.12	0.45	0.27	0.28	1.31	47.72	25.00 a	8.00 a	1.41	200
CDD ADD DLO	0.40	0.47	0.27	0.46	0.27	0.27	1.27	70.16	39.00 a	26.00 a	1.69	195
GBR_ABE_PLO	0.28	0.26	0.21	0.51	0.23	0.25	1.92	47.48	10.00 a	60.00 a	3.70	179
GBR_DEV_CON	0.88	0.48	0.62	0.93	0.04	0.04	2.10	44.38	14.80 D	50.90 D	3.44	200
GBR_GUL ST1	0.80	0.18	0.11	0.58	0.01	0.41	3.19	08.11	3.70 D	80.40 b	14.26	197
GBR_GUI_S12	0.59	0.55	0.47	0.59	0.06	0.35	3.19	08.11	3.70 D	80.40 b	14.20	197
GBR_GUI_ST3	0.82	0.81	0.77	0.66	0.03	0.31	3.19	08.11	3.70 D	80.40 D	14.20	197
CUP_GUY_ST2	0.90	0.90	0.93	0.82	0.07	0.11	2.88	135.18	43.00 a 43.00 a	48.00 a	1.03	200
GUF_GUY_S12	0.79	0.78	0.75	1.00	0.49	0.04	3.02	141.34	43.20 a 50.20 a	47.80 a	1.00	200
HUN SIK	0.81	0.35	0.71	0.02	0.00	0.00	2.09	30.64	39.20 a 30.40 b	33.20 a 44.00 b	2.22	200
ISP VAT VAT	0.19	0.30	0.48	0.33	0.01	0.00	0.10	82.42	28.00 2	21.00 5	0.71	178
TA FEL S17	0.54	0.32	0.20	0.68	0.13	0.10	1.08	22.07	8 00 a	76.00 a	3.1.1	117
ITA KAE S20	0.62	0.47	0.47	0.72	0.06	0.22	1.05	22.97	17.00 a	50.00 a	3.64	121
ITA MUN	0.56	0.39	0.42	0.74	0.22	0.04	0.80	29.87	7 00 a	55.00 a	193	188
ITA BEN	0.79	0.73	0.74	0.99	0.00	0.01	1.61	12.59	17 70 b	47.90 h	2.73	143
ITA BUN N20	0.78	0.71	0.71	0.88	0.00	0.12	1.39	15.28	14.00 a	54.00 a	3.33	123
MEX COR YP	0.68	0.23	0.31	0.79	0.05	0.16	1.42	81.77	22.20 b	46.40 b	2.94	200
NLD LOO	0.18	0.12	0.12	0.79	0.00	0.21	1.33	41.67	1.00 a	99.00 a	2.61	200
NLD SPE DOU	0.75	0.59	0.63	0.85	0.03	0.12	1.42	39.80	4.80 b	80.70 b	1.62	200
NZL HUA HUA	0.65	0.61	0.59	0.70	0.15	0.15	2.62	42.52	71.20 a	13.20 a	1.73	200
PRT LEZ ARN	0.59	0.28	0.27	0.78	0.01	0.21	0.72	77.42	5.04 a	90.38 a	1.52	200
PRT MIT	0.66	0.55	0.32	0.61	0.33	0.06	0.51	80.80	16.10 b	64.50 b	1.33	200
PRT PIN	0.65	0.55	0.38	0.65	0.32	0.02	0.76	74.76	16.60 b	61.20 b	1.26	200
RUS CHE Y4	0.32	0.22	0.24	0.94	0.00	0.06	0.62	34.23	21.10 b	23.20 b	4.96	200
RUS_FYO	0.61	0.56	0.54	0.81	0.02	0.17	1.24	30.87	18.20 b	48.80 b	3.77	198
RUS_POG_VAR	0.70	0.50	0.58	0.81	0.00	0.19	0.70	33.02	28.60 b	37.50 b	2.64	200
SEN_SOU_PRE	0.64	0.40	0.22	0.66	0.34	0.01	0.13	43.94	6.00 a	90.00 a	0.23	200
SWE_NOR_ST1_BEF	0.74	0.65	0.64	0.61	0.28	0.10	1.07	36.70	5.80 a	58.60 a	2.63	185
SWE_NOR_ST2	0.32	0.29	0.23	0.55	0.02	0.44	1.07	36.70	5.80 a	58.60 a	2.63	185
SWE_NOR_ST3	0.55	0.59	0.56	0.48	0.22	0.30	1.07	36.70	5.80 a	58.60 a	2.63	185
SWE_NOR_ST5_REF	0.53	0.56	0.53	0.48	0.20	0.32	1.07	36.55	19.20 b	43.50 b	2.83	190
SWE_SKO_MIN	0.71	0.70	0.67	0.52	0.04	0.44	1.60	45.86	17.30 b	52.00 b	2.48	133
SWE_SKY_38Y	0.34	0.43	0.37	0.05	0.94	0.00	1.39	33.61	21.70 b	43.80 b	3.93	184
SWE_SKY_68Y	0.35	0.53	0.34	0.05	0.83	0.11	1.30	33.80	18.90 b	46.50 b	4.15	184
SWE_SVA_MIX_NON	0.65	0.54	0.57	0.78	0.22	0.00	1.33	34.34	0.50 a	92.50 a	1.67	200
THA_KHU	0.50	0.41	0.41	0.78	0.21	0.01	0.83	84.24	10.00 a	65.00 a	0.75	200
USA_BNZ_BLA	0.52	0.37	0.46	0.68	0.21	0.10	0.69	33.86	10.30 b	36.80 b	2.57	200
USA CHE ASP	0.67	0.32	0.30	0.91	0.03	0.06	1.23	20.06	12.00 a	74.00 a	1.5.2	200

Table S4: Summary table of site level  $R_{VPD}^2$ ,  $R_{SWC}^2$ ,  $R_{PPFD}^2$ , climate, soil properties and vegetation structure data. PPET is in [mm mm<sup>-1</sup>], P-PET<sub>ad</sub> is in [mm], Clay and Sand are in [%], Total N is in [g kg<sup>-1</sup>], Stand height is in [m], LAI is in [m<sub>lenves</sub> m<sub>soil</sub><sup>2</sup>]. Letters show data source: a = SAPFLUXNET, b = Global rasters, c = SAPFLUXNET plant height. *(continued)* 

				.85	)	C ARE	Ş	15 ad				
SHEDDE	RUPD	REN	Reper	Relimp	Reling	Relimp	PPET	P. PET	Class	Sand	Total	Bedro
USA_CHE_MAP	0.57	0.52	0.52	0.76	0.06	0.18	1.22	19.85	6.63 a	$59.31 \ a$	2.54	200
USA_DUK_HAR	0.72	0.62	0.67	0.91	0.02	0.07	1.12	41.33	33.90 b	31.00 b	0.76	200
USA_HIL_HF2	0.75	0.71	0.74	0.61	0.00	0.39	1.14	37.46	26.00 a	43.00 a	0.71	200
USA_INM	0.40	0.35	0.38	0.56	0.00	0.44	1.18	39.20	26.70 b	8.00 b	1.05	200
USA_MOR_SF	0.65	0.58	0.49	0.90	0.09	0.01	1.18	39.20	30.00 a	10.00 a	1.05	200
USA_NWH	0.88	0.85	0.73	0.80	0.04	0.15	1.05	59.56	36.70 b	4.90 b	0.80	200
USA_ORN_ST1_AMB	0.63	0.62	0.53	0.51	0.10	0.39	1.14	61.36	24.00 a	21.00 a	0.85	200
USA_PAR_FER	0.43	0.17	0.22	0.69	0.04	0.27	1.32	25.96	10.00 a	60.00 a	1.75	200
USA_PER_PER	0.55	0.30	0.37	0.85	0.01	0.14	1.32	34.41	3.40 b	89.20 b	6.13	200
USA_PJS_P04_AMB	0.31	0.09	0.26	0.74	0.03	0.23	0.25	49.32	6.00 a	52.00 a	0.82	186
USA_PJS_P08_AMB	0.33	0.10	0.22	0.94	0.03	0.03	0.25	49.32	3.00 a	49.00 a	0.82	186
USA_PJS_P12_AMB	0.28	0.14	0.15	0.63	0.37	0.00	0.25	49.32	6.00 a	54.00 a	0.82	186
USA_SIL_OAK_1PR	0.38	0.40	0.36	0.29	0.58	0.13	1.36	38.70	1.00 a	98.00 a	0.74	200
USA_SIL_OAK_2PR	0.39	0.32	0.36	0.92	0.06	0.01	1.36	38.70	1.00 a	98.00 a	0.74	200
USA_SMI_SER	0.51	0.37	0.37	0.67	0.32	0.01	1.03	40.05	28.70 b	30.90 b	0.82	200
USA_SWH	0.86	0.63	0.50	0.94	0.01	0.05	1.09	62.29	43.10 b	6.30 b	0.69	200
USA_SYL_HL1	0.46	0.36	0.38	0.97	0.03	0.00	1.27	25.01	8.90 b	51.00 b	1.41	200
USA_SYL_HL2	0.47	0.46	0.46	0.55	0.01	0.44	1.27	25.01	8.90 b	51.00 b	1.41	200
USA_TNB	0.25	0.27	0.24	0.41	0.12	0.46	1.39	48.33	21.60 b	34.90 b	0.84	200
USA_TNO	0.40	0.40	0.38	0.50	0.22	0.28	1.41	60.02	29.60 b	30.20 b	0.83	200
USA_TNP	0.33	0.36	0.31	0.40	0.37	0.23	1.41	61.60	31.60 b	26.60 b	0.81	200
USA_UMB_CON	0.51	0.39	0.38	0.79	0.03	0.18	1.30	30.60	1.00 a	92.00 a	2.02	200
USA_UMB_GIR	0.43	0.35	0.34	0.78	0.05	0.17	1.25	30.69	1.00 a	92.00 a	2.49	200
USA_WIL_WC1	0.33	0.18	0.16	0.82	0.16	0.02	1.19	20.23	6.90 b	53.20 b	1.01	200
USA_WVF	0.28	0.26	0.25	0.59	0.06	0.36	1.63	30.35	24.90 b	29.90 b	1.37	200
ZAF_FRA_FRA	0.43	0.08	0.14	0.92	0.00	0.08	0.90	99.17	20.00 b	69.90 b	0.95	200
ZAF_RAD	0.46	0.42	0.40	0.57	0.10	0.33	0.95	82.73	21.30 b	61.40 b	1.18	200
ZAF_SOU_SOU	0.39	0.22	0.21	0.59	0.06	0.35	0.97	86.39	23.00 b	61.90 b	1.13	200
ZAF_WEL_SOR	0.62	0.33	0.34	0.63	0.08	0.29	0.50	79.71	20.00 a	60.00 a	0.81	179