# Identifying efficient ensemble perturbations for initializing subseasonal-to-seasonal prediction

Jonathan Demaeyer<sup>1</sup>, Stephen Gregory Penny<sup>2</sup>, and Stephane Vannitsem<sup>1</sup>

<sup>1</sup>Royal Meteorological Institute of Belgium

 $^2\mathrm{Cooperative}$  Institute for Research in Environmental Sciences at the University of Colorado Boulder

November 22, 2022

#### Abstract

The prediction of the weather at subseasonal-to-seasonal (S2S) timescales is dependent on both initial and boundary conditions. An open question is how to best initialize a relatively small-sized ensemble of numerical model integrations to produce reliable forecasts at these timescales. Reliability in this case means that the statistical properties of the ensemble forecast are consistent with the actual uncertainties about the future state of the geophysical system under investigation. In the present work, a method is introduced to construct initial conditions that produce reliable ensemble forecasts by projecting onto the eigenfunctions of the Koopman or the Perron-Frobenius operators, which describe the time-evolution of observables and probability distributions of the system dynamics, respectively. These eigenfunctions can be approximated from data by using the Dynamic Mode Decomposition (DMD) algorithm. The effectiveness of this approach is illustrated in the framework of a low-order oceanatmosphere model exhibiting multiple characteristic timescales, and is compared to other ensemble initialization methods based on the Empirical Orthogonal Functions (EOFs) of the model trajectory and on the backward and covariant Lyapunov vectors of the model dynamics. Projecting initial conditions onto a subset of the Koopman or Perron-Frobenius eigenfunctions that are characterized by time scales with fast-decaying oscillations is found to produce highly reliable forecasts at all lead times investigated, ranging from one week to two months. Reliable forecasts are also obtained with the adjoint covariant Lyapunov vectors, which are the eigenfunctions of the Koopman operator in the tangent space. The advantages of these different methods are discussed.

# Identifying efficient ensemble perturbations for initializing subseasonal-to-seasonal prediction

# Jonathan Demaeyer<sup>1</sup>, Stephen G. Penny<sup>2,3</sup>, and Stéphane Vannitsem<sup>1</sup>

<sup>1</sup>Royal Meteorological Institute of Belgium, Brussels, Belgium
<sup>2</sup>Cooperative Institute for Research in Environmental Sciences, University of Colorado Boulder
<sup>3</sup>NOAA Physical Sciences Laboratory, Boulder, Colorado

## Key Points:

1

2

3

4 5 6

7

8	•	Several methods for initializing ensemble forecasts with long lead times are tested
9		in the context of an ocean-atmosphere coupled model
10	•	The methods providing the most reliable ensembles are the adjoint Lyapunov vec-
11		tors and the adjoint modes of the Dynamic Mode Decomposition
12	•	These are related to the eigenfunctions of the Koopman and Perron-Frobenius op-
13		erators of the system

 $Corresponding \ author: \ Jonathan \ Demaeyer, \ {\tt jodemaey@meteo.be}$ 

#### 14 Abstract

The prediction of the weather at subseasonal-to-seasonal (S2S) timescales is dependent 15 on both initial and boundary conditions. An open question is how to best initialize a rel-16 atively small-sized ensemble of numerical model integrations to produce reliable forecasts 17 at these timescales. Reliability in this case means that the statistical properties of the 18 ensemble forecast are consistent with the actual uncertainties about the future state of 19 the geophysical system under investigation. In the present work, a method is introduced 20 to construct initial conditions that produce reliable ensemble forecasts by projecting onto 21 the eigenfunctions of the Koopman or the Perron-Frobenius operators, which describe 22 the time-evolution of observables and probability distributions of the system dynamics, 23 respectively. These eigenfunctions can be approximated from data by using the Dynamic 24 Mode Decomposition (DMD) algorithm. The effectiveness of this approach is illustrated 25 in the framework of a low-order ocean-atmosphere model exhibiting multiple character-26 istic timescales, and is compared to other ensemble initialization methods based on the 27 Empirical Orthogonal Functions (EOFs) of the model trajectory and on the backward 28 and covariant Lyapunov vectors of the model dynamics. Projecting initial conditions onto 29 a subset of the Koopman or Perron-Frobenius eigenfunctions that are characterized by 30 time scales with fast-decaying oscillations is found to produce highly reliable forecasts 31 at all lead times investigated, ranging from one week to two months. Reliable forecasts 32 33 are also obtained with the adjoint covariant Lyapunov vectors, which are the eigenfunctions of the Koopman operator in the tangent space. The advantages of these different 34 methods are discussed. 35

#### <sup>36</sup> Plain Language Summary

Weather forecasts often reach their limit of predictability at one to two weeks. In 37 order to extend forecast skill beyond this two week limit, the weather prediction com-38 munity has begun transitioning to the use of coupled models that include both atmo-39 sphere and ocean dynamics, with the slower ocean dynamics enabling an extended fore-40 cast horizon. Due to uncertainties in the accuracy of the initial conditions and the model 41 itself, such forecasts must be probabilistic. The primary approach for probabilistic weather 42 prediction is to generate ensemble forecasts that integrate multiple copies of the model 43 started from slightly different initial conditions. Here we show that the method used to 44 determine the ensemble of initial conditions has a significant impact on the probabilis-45 tic forecast skill at horizons ranging from a few weeks to a few months. We show that 46 many of the existing techniques used for short forecasts are suboptimal for longer fore-47 cast horizons. We introduce a new perspective and corresponding techniques that per-48 mit the initialization of these ensemble forecasts using information that is intrinsic to 49 the nature of the evolution of the coupled system dynamics, and present data-driven meth-50 ods that allow this information to be estimated directly from historical data. 51

#### 52 1 Introduction

Long-term forecasts of the atmosphere at sub-seasonal, seasonal, and decadal time 53 scales are affected by both the choice of the initial condition and the slow evolution of 54 surface boundary conditions. This multi-timescale forecasting is a key target of the at-55 mospheric and climate communities (Vitart et al., 2017; Cassou et al., 2018). Forecast 56 error grows quickly at increasing lead times due to the instability properties of weather 57 dynamics. For this reason, a probabilistic approach is necessary in order to isolate the 58 multiple possible outcomes of a set of forecasts. Since the 1990's, such an approach has 59 been developed in many operational weather prediction centers based on using multi-60 ple numerical integrations of the models starting from slightly different initial conditions. 61 This approach is known as ensemble forecasting (Buizza, 2019; Kalnay, 2019). Differ-62 ent perturbation techniques have been designed for initializing ensemble weather fore-63

casts, of which the most popular are the singular vectors (Molteni et al., 1996), the bred
modes (Toth & Kalnay, 1997), and perturbed observations applied within data assimilation systems (Buizza et al., 2005; Kleist & Ide, 2015). Buizza et al. (2005) noted that
the accuracy of initial conditions is just as important as the accuracy of the forecast models for generating reliable ensemble forecasts.

As operational centers expanded their focus to include longer forecast horizons, the 69 same perturbation approaches were also used for sub-seasonal to decadal predictions. How-70 ever, at timescales beyond the limits of predictability for the atmosphere, coupled Earth 71 72 system models must be used. This introduces the additional difficulty of building appropriate perturbations for the different components of such multi-scale systems (O'Kane 73 et al., 2019). Several approaches consisting of appropriately tuning the bred modes to 74 capture the long time scales of the dynamics have for instance been proposed (Peña & 75 Kalnay, 2004; S.-C. Yang et al., 2008; O'Kane et al., 2019), and the use of backward Lya-76 punov vectors (BLV), closely related to the bred modes, have been used to build reliable 77 ensemble forecasts in idealized scenarios (Vannitsem & Duan, 2020). 78

In the present work, we address this ensemble initialization problem by consider-79 ing tools coming from the probabilistic description of dynamical systems and finding their 80 roots in the conservation of the number of trajectories in phase space described by the 81 Liouville equation (Gaspard, 2005; Nicolis & Nicolis, 2012). The evolution operator as-82 sociated with this equation is known as the Perron-Frobenius operator (Lasota & Mackey, 83 2008), sometimes also called the *transfer* operator. It has been used as a theoretical framework to describe *probabilistic forecasting* (Ehrendorfer, 2006; Giannakis, 2019), i.e. fore-85 casting based on the time evolution of a probability distribution, and for which the en-86 semble forecasting methods provide approximations. The adjoint of the Perron-Frobenius 87 operator, known as the Koopman operator, has become popular to describe the dynam-88 ics of observables on attractors (Mezić, 2013; Susuki et al., 2016; Arbabi & Mezić, 2017; 89 Santos Gutiérrez et al., 2021), due to the fact that when operating on functional spaces 90 it is a linear operator, an *observable* being defined as any function mapping the system 91 state to some real or complex value. A trade-off, however, in converting the nonlinear 92 dynamics to a linear representation is that the Koopman operator generally acts on an 93 infinite dimensional space, but as we will see, methods exist to obtain finite-dimensional 94 approximate representations of these operators. The computation of the spectrum of these 95 operators has also been considered in order to study bifurcations in low- and high-dimensional 96 systems (Tantet, Lucarini, & Dijkstra, 2018; Tantet, Lucarini, Lunkeit, & Dijkstra, 2018). 97 The eigenvalues and eigenfunctions of these operators can then be obtained in the func-98 tional spaces, and provide the key building blocks of the dynamics of the probability density and observables. These are precisely the quantities that are used in the present work 100 to generate the ensemble forecasts initialization, as they constitute generic features of 101 the dynamics of the probability density. 102

The eigenfunctions of the Koopman operator can be approximated using Dynamic 103 Mode Decomposition (DMD) (Rowley et al., 2009; Tu et al., 2014). The DMD approach 104 is a rediscovery of the Linear Inverse Model (LIM), which was developed first within the 105 seasonal prediction community (Penland, 1989; Penland & Magorian, 1993; Penland & 106 Sardeshmukh, 1995). For computational efficiency, and due to the large volume of data 107 involved, LIMs are typically formed using data projected to the space of Empirical Or-108 thogonal Functions (EOFs) and then truncated. As such, a mathematical equivalence 109 between this form of the LIM and the projected DMD was noted by Tu et al. (2014). 110 The LIM approach is now being used experimentally for seasonal forecasts by the US 111 National Oceanographic and Atmospheric Administration (NOAA) Climate Prediction 112 Center (CPC) (Wang et al., 2021). Since its rediscovery by the fluid mechanics commu-113 nity under the name of DMD (Schmid, 2010), many new algorithms, theoretical results, 114 and variants have been developed that have advanced understanding of this approach. 115

<sup>116</sup> Most notably, perhaps, is the connection between DMD and the Koopman operator (Rowley <sup>117</sup> et al., 2009).

The usefulness of the Perron-Frobenius and Koopman operators for producing en-118 semble forecasts will be analyzed in a reduced order coupled ocean-atmosphere model, 119 previously demonstrated for a similar purpose by Vannitsem and Duan (2020). The model 120 will be briefly described in Section 2. The experimental setup will be then presented in 121 Section 3. In Section 4, the different bases onto which the perturbed initial conditions 122 are projected will be presented: first, the EOFs that are often used in initializing climate 123 models and their ensemble integrations, e.g. Polkova et al. (2019); second, the Lyapunov 124 vector approach used by Vannitsem and Duan (2020), which is closely related to the bred 125 modes and ensemble Kalman filters; and third, the eigenfunctions of the Koopman and 126 Perron-Frobenius operators. The Koopman and Perron-Frobenius operators are exten-127 sively discussed, as important clarifications on their link with DMD is needed. Section 128 5 describes the specific choices of bases used for the experiments. Section 6 presents the 129 results of experiments using the aforementioned bases to initialize ensemble forecasts. 130 It will be shown that the eigenfunctions of the Koopman and Perron-Frobenius opera-131 tors are indeed the most efficient tools for producing reliable ensemble forecasts in such 132 multiscale systems. Finally, conclusions are drawn in Section 7. 133

#### <sup>134</sup> 2 The coupled ocean-atmosphere model

Experiments are conducted with a coupled ocean-atmosphere model that was first 135 introduced by Vannitsem et al. (2015), and was further generalized by De Cruz et al. (2016) 136 and Demaeyer et al. (2020). It consists of a two-layer quasi-geostrophic atmospheric model 137 coupled both thermally and mechanically to a shallow-water oceanic component on a beta 138 plane. The coupling between the ocean and the atmosphere includes the wind stress and 139 heat exchanges. The fields of the model are defined on a rectangular domain with the 140 zonal and meridional coordinates x and y being restricted to  $0 \le x \le 2\pi L/n$  and  $0 \le x \le 2\pi L/n$ 141  $y \leq \pi L$ , where n is the aspect ratio of the domain and L is the characteristic spatial 142 scale. The atmospheric fields are defined in a zonally periodic channel with no-flux bound-143 ary conditions in the meridional direction, i.e. if  $\psi$  is such an atmospheric field then  $\partial \psi / \partial x \equiv$ 144 0 at  $y = 0, \pi L$ . The oceanic fields are defined on a closed basin, with no flux through 145 the boundaries. 146

The model fields include the atmospheric barotropic  $\psi_{a}$  and baroclinic streamfunctions  $\theta_{a}$ , and the ocean streamfunction  $\psi_{o}$  and the temperature field  $\theta_{o}$ . These fields are expanded in series of Fourier modes  $F_{i}(x, y)$  for the atmosphere and  $\phi_{i}(x, y)$  for the ocean, both respecting the prescribed boundary conditions:

$$\begin{split} \psi_{\mathbf{a}}(x,y) &= \sum_{i=1}^{n_{\mathbf{a}}} \psi_{\mathbf{a},i} \, F_i(x,y) \\ \theta_{\mathbf{a}}(x,y) &= \sum_{i=1}^{n_{\mathbf{a}}} \theta_{\mathbf{a},i} \, F_i(x,y) \\ \psi_{\mathbf{o}}(x,y) &= \sum_{i=1}^{n_{\mathbf{o}}} \psi_{\mathbf{o},i} \, \phi_i(x,y) \\ \theta_{\mathbf{o}}(x,y) &= \sum_{i=1}^{n_{\mathbf{o}}} \theta_{\mathbf{o},i} \, \phi_i(x,y) \end{split}$$

After projecting the partial differential equations (PDEs) of the model on the Fourier modes, one obtains a set of ordinary differential equations (ODEs) governing the time

evolution of the coefficients  $\psi_{a,i}$ ,  $\theta_{a,i}$ ,  $\psi_{o,i}$  and  $\theta_{o,i}$ :

$$\begin{aligned} \dot{\boldsymbol{x}} &= \boldsymbol{f}(\boldsymbol{x}) \tag{1} \\ \boldsymbol{x} &= [\psi_{\mathrm{a},1}, \dots, \psi_{\mathrm{a},n_{\mathrm{a}}}, \theta_{\mathrm{a},1}, \dots, \theta_{\mathrm{a},n_{\mathrm{a}}}, \\ \psi_{\mathrm{o},1}, \dots, \psi_{\mathrm{o},n_{\mathrm{o}}}, \theta_{\mathrm{o},1}, \dots, \theta_{\mathrm{o},n_{\mathrm{o}}}]^{\mathsf{T}} \end{aligned}$$

where <sup>T</sup> denotes the matrix transposition operation. These coefficients thus form the set of the model state variables and the equation above allows one to simulate the physical system using numerical integration. In the present study, we consider the so-called *VDDG* model configuration first defined by Vannitsem et al. (2015), with the atmospheric and the oceanic fields each being expanded into a series of  $n_a = 10$  and  $n_o = 8$  selected modes, respectively, leading to a system with d = 36 dimensions<sup>1</sup>.

A critical parameter of the model is the friction coefficient C between the ocean 153 and the atmosphere. Indeed, it was shown by Vannitsem et al. (2015) that the strength 154 of the wind stress controls the presence and the amplitude of a low-frequency variabil-155 ity (LFV) typically found in the real atmosphere at midlatitude. Following Vannitsem 156 (2017) and Vannitsem and Duan (2020), we shall consider two cases: one with weak LFV 157  $(C = 0.01 \text{ kg m}^{-2} \text{ s}^{-1})$  and another with much more pronounced LFV (C = 0.016158 kg m<sup>-2</sup> s<sup>-1</sup>). Solutions of the models for both cases are depicted in Figure 1, where the 159 difference in the amplitude of LFV between the left and right panels is clear. The vari-160 ables shown in this figure are the coefficients corresponding to the first mode of the baro-161 clinic streamfunction and the second mode of the ocean temperature field, each sampled 162 every  $\Delta t = 10$  nondimensional model timeunits (MTU), corresponding to 1.1215 days. 163 The former mode is related to the meridional temperature gradient in the system, while 164 the second corresponds to a dominant double-gyre signal in the ocean. 165

These two cases will allow us to highlight how the different methods of initialization that we consider perform in different settings, with different timescales and different correlation structures between the components being involved.

#### <sup>169</sup> 3 Experiment design

The focus of this study is initialization methods for ensemble forecasts. To this end, the long reference runs depicted in Figure 1 were computed to serve as the 'truth' in our experiments. We select N points of the reference runs, denoted  $\boldsymbol{x}_n(0)$ , as initial conditions to produce N ensemble forecasts with the VDDG model, using the same parameters as the reference runs. To ensure that the experiments are initialized from a state close to the true trajectory, but with an ensemble mean state that is not precisely equal to the truth, we first obtain the initial conditions  $\boldsymbol{x}_n^{\text{ctrl}}$  of a deterministic control forecast by perturbing the N points of the reference 'truth' by a random perturbation  $\delta \boldsymbol{x}_0^{\text{ctrl}}$ sampled from a uniform distribution  $\rho_0^{\text{pert}}$ :

$$\boldsymbol{x}_n^{\text{ctrl}}(0) = \boldsymbol{x}_n(0) + \boldsymbol{\delta} \boldsymbol{x}_0^{\text{ctrl}}.$$
(2)

An ensemble is then generated by perturbing the control run initial conditions with a set of M-1 perturbations  $\delta \boldsymbol{x}_0^m$  drawn from a distribution  $\tilde{\rho}_0^{\text{pert}}$ , which is taken to be the same distribution as that used to obtain the control:  $\tilde{\rho}_0^{\text{pert}} \equiv \rho_0^{\text{pert}}$ . The initial conditions of the ensemble are thus:

$$\boldsymbol{y}_{m,n}(0) = \boldsymbol{x}_n^{\text{ctrl}}(0) + \boldsymbol{\delta} \boldsymbol{x}_0^m \quad , \quad m = 1, \dots, M - 1.$$
 (3)

Including the control run, i.e.  $\delta x_0^M = 0$ , this provides a reference *perfect* ensemble of M members. In the present study, as in Vannitsem and Duan (2020), a uniform distri-

 $<sup>^{1}</sup>$  In the following, the letter d will always refer to the dimension of the dynamical system.

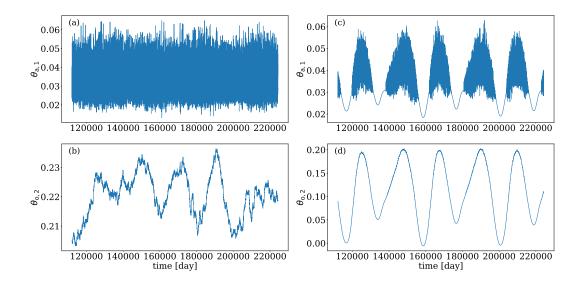


Figure 1. Time evolution of a trajectory on the attractor for: (a),(b) the case without low-frequency variability and (c),(d) the case with low-frequency variability. For this latter case, it shows that the presence or absence of atmospheric instability and fast variability is linked to the low-frequency variation of the meridional temperature gradient in the ocean.

bution defined on the interval  $[-\varepsilon/2, \varepsilon/2]$  with  $\varepsilon = 10^{-6}$  was used to perturb each component of the system's state vector. Additional computations done with Gaussian distributions with the same standard deviation did not show any differences in the results of the study.

Due to the high dimensionality of more realistic applications, and the cost involved in integrating long model forecasts of this size, ensemble perturbations must be built from a relatively small subspace of the original system. While this distribution could be sampled randomly, we seek a set of initial conditions that can efficiently reproduce the true error growth characteristics of the coupled system dynamics. The question remains as to what bases are most efficient for initializing a reliable ensemble forecast, and how can those bases be practically constructed in a realistic setting.

We examine reduced-size ensembles constructed using linear projections of the 'per-183 fect' ensemble initial conditions onto various bases forming subspaces with rank less than 184 M, and compare these to the perfect ensemble as a benchmark. The specific bases that 185 we use will be detailed in the next section. By design, the ensemble perturbations ob-186 tained by projection cannot be more reliable than the original reference ensemble. How-187 ever, we will show that depending on the modes and subspaces selected as a basis, these 188 reduced-size ensembles can achieve similar performance to the full-rank 'perfect' ensem-189 ble. 190

To determine whether the ensemble forecasts generated from the projected initial conditions are reliable, the mean square error (MSE) of the ensemble mean and the variance of the ensemble (the square of the ensemble spread) are computed at each lead time  $\tau$  of the ensemble forecasts as:

$$MSE(\tau) = \frac{1}{N} \sum_{n=1}^{N} \| \boldsymbol{x}_{n}(\tau) - \bar{\boldsymbol{y}}_{n}(\tau) \|^{2}$$
(4)

Spread<sup>2</sup>(
$$\tau$$
) =  $\frac{1}{N} \sum_{n=1}^{N} \frac{1}{M-1} \sum_{m=1}^{M} \| \boldsymbol{y}_{m,n}(\tau) - \bar{\boldsymbol{y}}_{n}(\tau) \|^{2}$  (5)

where

197

$$\bar{\boldsymbol{y}}_{n}(\tau) = \frac{1}{M} \sum_{m=1}^{M} \boldsymbol{y}_{m,n}(\tau)$$
(6)

is the ensemble mean over the members  $m{y}_{m.n}( au)$  of the  $n^{\mathrm{th}}$  ensemble forecast and  $m{x}_n( au)$ 191 is the corresponding reference solution. Finally,  $\|\cdot\|^2$  is the L<sup>2</sup>-norm. If the Spread<sup>2</sup> and 192 the MSE are close to one another, indicating that the estimated error is close to the true 193 error, then the ensemble forecast is considered reliable (Leutbecher & Palmer, 2008). The 194 results based on these measures are presented in the supplementary material. 195

An alternative measure of reliability of the ensemble forecasts can also be assessed by considering the proper ignorance (or logarithmic) score (Roulston & Smith, 2002):

$$\mathcal{I}[\rho_{\tau}^{\text{ens}}] = -\ln \rho_{\tau}^{\text{ens}}(\boldsymbol{x}_n(\tau) | \boldsymbol{x}_n^{\text{ctrl}}(0)).$$
(7)

Applying the ignorance score to a Gaussian, one obtains the related proper two-moment skill score derived by Dawid and Sebastiani (1999). As such, and regardless of whether the distributions being considered is Gaussian or not (Leutbecher, 2019), the Dawid-Sebastiani Score (DSS) provides an evaluation of the quality of the first and second moments of the forecast distribution estimated by the ensemble, with respect to the true moments. The bias-free univariate DSS for the  $n^{\text{th}}$  ensemble forecast and the  $i^{\text{th}}$  variable of the system can be written as (Siegert et al., 2019):

$$DSS_{n,i}(\tau) = \frac{1}{2} \log(2\pi) + \frac{1}{2} \log \sigma_{n,i}^2(\tau) + \frac{1}{2} \frac{M-3}{M-1} \left( \bar{y}_{n,i}(\tau) - x_{n,i}(\tau) \right)^2 \middle/ \sigma_{n,i}^2(\tau),$$
(8)

where  $\sigma_{n,i}^2$  is an estimator of the *i*<sup>th</sup> variable ensemble variance:

$$\sigma_{n,i}^2(\tau) = \frac{1}{M-1} \sum_{m=1}^M |y_{m,n,i}(\tau) - \bar{y}_{n,i}(\tau)|^2.$$
(9)

This score can then be averaged over the N realizations:

$$DSS_i(\tau) = \frac{1}{N} \sum_{n=1}^{N} DSS_{n,i}(\tau).$$
(10)

The lower the DSS score, the more reliable the ensemble forecasts are for this particu-196

lar variable. In particular, the DSS score has been used to characterize the ensembles reliability in the study done by Vannitsem and Duan (2020).

198

#### 4 Initialization methods for ensemble forecasts 199

We now discuss the different bases onto which the set of ensemble perturbations will be projected. Assume that a basis comprises k vectors of dimension d arranged as columns of the matrix  $\mathsf{B} \in \mathbb{C}^{d \times k}$ . We can construct the projection operator onto this basis as,

$$\Pi = \mathsf{B}(\mathsf{B}^*\,\mathsf{B})^{-1}\,\mathsf{B}^*.\tag{11}$$

If B is unitary, this reduces to  $\Pi = BB^*$  (Meyer, 2000). Assuming that  $\Pi \in \mathbb{R}^{d \times d}$ , if one considers the ensemble of M-1 perturbations  $\delta x_0^m$  of the control initial conditions, then the projection of the perturbations onto the subspace spanned by B is given by:

$$\delta \boldsymbol{x}_0^{\prime m} = \Pi \, \delta \boldsymbol{x}_0^m. \tag{12}$$

The resulting perturbations are used to initialize ensemble forecasts with the initial conditions:

$$\boldsymbol{y}_{m,n}'(0) = \boldsymbol{x}_n^{\text{ctrl}}(0) + \boldsymbol{\delta}\boldsymbol{x}_0'^m \quad , \quad m = 1, \dots, M$$
(13)

#### in the experiments discussed in Section 6.

207

Let us now detail the various basis vectors considered and the subspaces that they span, namely the EOFs, the backward (BLVs) and covariant (CLVs) Lyapunov vectors, and the Koopman and Perron-Frobenius eigenfunctions determined using DMD. Unlike the other basis vectors used, the Lyapunov vectors are time-dependent, defined locally at each point of the reference trajectory, and are related to the stability of the local linearized dynamics.

#### 4.1 Empirical Orthogonal Functions

The EOFs of the dataset are obtained using a Principal Component Analysis (PCA), which decomposes the data into a set of orthogonal basis functions and time-dependent coefficients. These orthogonal patterns can be obtained directly by singular value decomposition (SVD) of the data matrix, or by computing the eigenvectors of the data covariance matrix (Wilks, 2011).

Assuming that the dataset is represented by the matrix  $X = [x_0 \dots x_{K-1}]$ , with  $X \in \mathbb{R}^{d \times K}$ . The columns of X are the system states  $x_k = \Phi^{t_k}(x_0)$  at times  $t_k = k\Delta t$  where  $\Phi^t$  is the flow of the system (1):  $x(t) = \Phi^t(x(0))$ , The EOFs are the column vectors of U as determined by the PCA:

$$\mathsf{T} = \mathsf{U}^* \bar{\mathsf{X}} \tag{14}$$

where  $\bar{X} = X - \langle X \rangle_k$  is the matrix of system states with zero empirical time mean, and U is a matrix whose columns are the orthogonal eigenvectors of the matrix  $\bar{X}\bar{X}^*$  which is proportional to the covariance matrix of the system, and T is the time-series of the coefficients of the decomposition. The eigenvalues of the matrix  $\bar{X}\bar{X}^*$  are related to the variance of the data projected onto the corresponding mode. The amplitude of the eigenvalues comparatively to the others then provide the 'fraction of explained variance' by a given EOF.

The EOFs can alternatively be obtained by SVD of  $\bar{X}$ :

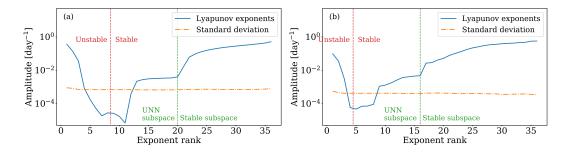
$$\bar{\mathsf{X}} = \mathsf{U}\mathsf{\Sigma}\mathsf{V}^* \tag{15}$$

where U and V are two unitary square matrices and  $\Sigma$  is diagonal, containing the singular values of  $\bar{X}$ . The matrix U contains the EOFs of  $\bar{X}$  since  $\bar{X}\bar{X}^* = U\Sigma\Sigma^*U^*$ , and the PCA time-series of coefficients can be represented as  $T = U^*\bar{X} = \Sigma V^*$ .

#### 4.2 The Lyapunov Vectors

We next consider the backward Lyapunov vectors (BLVs), the covariant Lyapunov 224 vectors (CLVs), and the adjoint CLVs as basis vectors B in Eq. (11). The Lyapunov vec-225 tors are locally defined in the tangent space of the trajectory of the model, and give in-226 formation about the stability therein. For instance, Osedelets has shown that the tan-227 gent space can be decomposed into a set of nested subspaces  $S_k^-$  that are invariant un-228 der the tangent linear model dynamics (V. I. Oseledets, 1968; V. Oseledets, 2008). Ar-229 bitrary  $k\mbox{-}\mathrm{volumes}$  defined in the tangent space converge to the subspace  $S_k^-$  under the 230 action of the tangent flow. These subspaces are spanned by the BLVs  $\varphi_i^-$ :  $S_k^- = \text{Span}\{\varphi_i^-|i=$ 231  $1, \ldots, k$ . The BLVs are thus related to the asymptotic properties of volumes in the tan-232 gent space, i.e. to how volumes contract or expand in the tangent space. The Lyapunov 233 exponents characterize the time-average expansion and contraction rates of these vol-234 umes over the entire attractor. 235

The CLVs  $\varphi_i$  are defined as stability directions in the tangent space that are covariant under the application of the tangent linear model dynamics. The tangent linear flow maps a CLV at one time to the same CLV at a later time, but multiplied by a *stretch*-



Absolute value of the Lyapunov exponents (the LEs are expressed in  $day^{-1}$ ), along Figure 2. with the one standard-deviation uncertainty, for: (a) the case with weak low-frequency variability and (b) the case with strong low-frequency variability. The approximate

separation between the positive (unstable) and negative (stable) exponents is depicted by a vertical dashed red line.

ing factor defined over the same timescale as the tangent linear mapping, which indi-239 cates the local stability of this CLV (Kuptsov & Parlitz, 2012). Finally, the adjoint CLVs 240  $\tilde{\varphi}_i$  are vectors that are covariant as well, but with respect to the flow of the adjoint model. 241 See Appendix A for more details. 242

To determine the sets B of basis vectors that we will consider in the experiments, 243 it is useful to consider the Lyapunov spectra  $\sigma_i$  (depicted in Figure 2). These exponents 244 have been estimated by averaging the local stretching rate  $^{2}\chi_{i}$  along the trajectories de-245 picted in Figure 1 with a bootstrap algorithm (Efron & Tibshirani, 1993) to increase its 246 statistical significance. The standard deviation of the time series used to compute the 247 averaged Lyapunov exponent is also shown. 248

A chaotic dynamical system generally has positive (unstable) and negative (sta-249 ble) exponents, along with a single zero-valued exponent that corresponds to the direc-250 tion of flow of the system trajectory. For the coupled atmosphere-ocean system, how-251 ever, because the magnitude of many of the near-zero exponents is smaller than the stan-252 dard deviation of the time series itself, it is difficult to precisely identify the zero-valued 253 Lyapunov exponent that separates the stable and unstable directions in the spectra (Vannitsem 254 & Lucarini, 2016; S. Penny et al., 2019). This is true for both model configurations (weak 255 and strong LFV). 256

The BLVs have been computed with the Benettin algorithm (Benettin et al., 1980), 257 while the CLVs and their adjoint have been computed by the method seeking to find the 258 intersection of the subspaces spanned by the BLVs and the Forward Lyapunov Vectors 259 (FLVs) (Legras & Vautard, 1996), see Appendix A. To this end, the FLVs have also been 260 computed using the Benettin algorithm. 261

262

## 4.3 The Koopman (KM) and Perron-Frobenius (PF) eigenfunctions

263

# 4.3.1 Koopman and the Perron-Frobenius operators

The Koopman operator provides a means of representing a finite-dimensional nonlinear system as an infinite-dimensional linear system by 'lifting' the underlying state space to a set of observables. The Koopman operator  $\mathcal{K}^{\tau}$  acts upon an observable  $g(\boldsymbol{x})$ of the system state  $\boldsymbol{x}$  as,

$$\mathcal{K}^{\tau} g(\boldsymbol{x}) = g\left(\boldsymbol{\Phi}^{\tau}(\boldsymbol{x})\right) \tag{16}$$

 $<sup>^2</sup>$  See Eq. (A12) in Appendix A.

where the mapping  $\mathbf{\Phi}^{\tau}$  describes the flow of the system (1) such that  $\mathbf{x}(t+\tau) = \mathbf{\Phi}^{\tau}(\mathbf{x}(t))$ .

While the Koopman operator governs the time evolution of observables of the system, its adjoint, the Perron-Frobenius (or transfer) operator  $\mathcal{P}^{\tau}$ , governs the time evolution of the probability density  $\rho_t$ . The probability density given at any lead time  $\tau$  is thus,

$$\rho_{t+\tau} = \mathcal{P}^{\tau} \,\rho_t. \tag{17}$$

The Koopman and Perron-Frobenius operators can both be used to determine the evolution of the expected value of an observable. Indeed, if we consider the expected value of an observable g, for a given distribution  $\rho_t$  at time t, to be defined as,

$$\langle g \rangle_t = \int g(\boldsymbol{x}) \, \rho_t(\boldsymbol{x}) \, \mathrm{d}\boldsymbol{x}$$
 (18)

and the inner product is defined as,

$$\langle a,b\rangle = \int a^*(\boldsymbol{x}) \, b(\boldsymbol{x}) \, \mathrm{d}\boldsymbol{x},$$
 (19)

then for a real-valued scalar observable g, we have  $\langle g \rangle_t = \langle g, \rho_t \rangle = \langle g, \mathcal{P}^t \rho_0 \rangle = \langle \mathcal{K}^t g, \rho_0 \rangle$ . Note that we have used the fact that the observable is real, i.e.  $g(\boldsymbol{x})^* = g(\boldsymbol{x})$ , and that  $\mathcal{K}^t$  is the adjoint of  $\mathcal{P}^t$ .

For the remainder, for the sake of simplicity, we shall assume that the spectra of the Koopman and Perron-Frobenius operators are discrete.<sup>3</sup> Importantly, a vector-valued observable  $\boldsymbol{g}$  can then be decomposed using the eigenfunctions  $\phi_i$  of the Koopman operator

$$\boldsymbol{g}(\boldsymbol{x}) = \sum_{i=1}^{\infty} \boldsymbol{c}_i^{\text{KM}} \phi_i(\boldsymbol{x})$$
(20)

and the application of the Koopman operator can thus be decomposed into a set of eigenvalues  $\lambda_i$ , eigenfunctions  $\phi_i$ , and modes (coefficients)  $c_i^{\text{KM}}$  as,

$$\mathcal{K}^{\tau}\boldsymbol{g}(\boldsymbol{x}) = \sum_{i=1}^{\infty} \boldsymbol{c}_{i}^{\mathrm{KM}} \lambda_{i}(\tau) \phi_{i}(\boldsymbol{x}).$$
(21)

This indicates that the propagation of an observable due to the Koopman operator can be represented as a superposition of oscillating stretching/contracting factors applied to the Koopman eigenfunctions. A challenge in translating the use of this Koopman operator to practical applications is the determination of an appropriate truncation of this infinite series.

Similarly for the Perron-Frobenius operator, a probability density  $\rho$  defined over the phase space can be expanded in terms of its eigenfunctions  $\psi_i$ :

$$\rho(\boldsymbol{x}) = \sum_{i=1}^{\infty} c_i^{\text{PF}} \psi_i(\boldsymbol{x})$$
(22)

and its time evolution is then also decomposable in term of a set of eigenvalues  $\lambda_i$ , eigenfunctions  $\psi_i$ , and coefficients  $c_i^{\text{PF}}$ :

$$\mathcal{P}^{\tau}\rho(\boldsymbol{x}) = \sum_{i=1}^{\infty} c_i^{\mathrm{PF}} \lambda_i^*(\tau) \psi_i(\boldsymbol{x}).$$
(23)

 $<sup>^{3}</sup>$  However, special care must be taken when the system is chaotic, as it may include degenerate eigenvalues (Jordan blocks) and continuous parts (Gaspard et al., 1995; Arbabi & Mezić, 2017; Mezić, 2020). See also the conclusion (Section 7).

The eigenfunctions of the Koopman and Perron-Frobenius operators are biorthonormal to one another  $\langle \phi_i, \psi_j \rangle = \delta_{i,j}$ , and therefore, the Koopman modes of a given observable  $\boldsymbol{g}$  can be determined using the Perron-Frobenius eigenfunctions:  $\boldsymbol{c}_i^{\text{KM}} = \langle \psi_i, \boldsymbol{g} \rangle$ , where the inner product is applied component-wise.

The time evolution of the expected value of the observable can then be given more simply as

$$\langle \boldsymbol{g} \rangle_t = \langle \boldsymbol{g}, \mathcal{P}^t \rho_0 \rangle$$

$$= \sum_{i=1}^{\infty} \langle \boldsymbol{g}, \lambda_i^*(t) \psi_i c_i^{\mathrm{PF}} \rangle$$

$$= \sum_{i=1}^{\infty} \lambda_i^*(t) c_i^{\mathrm{PF}} \left( \boldsymbol{c}_i^{\mathrm{KM}} \right)^*$$

$$(24)$$

From now on, to present numerical methods to approximate the Koopman and Perron-Frobenius eigenfunctions decompositions, we will consider a set of realizations  $\boldsymbol{g}_t = \boldsymbol{g}(\boldsymbol{x}_t)$  of a vector-valued observable  $\boldsymbol{g}$  of dimension P evaluated over the system states  $\boldsymbol{x}_t$  that are assumed to satisfy,

$$\boldsymbol{g}_{t+\tau} = \mathcal{K}^{\tau} \boldsymbol{g}_t. \tag{25}$$

If this time evolution is repeated sequentially with a fixed lead time  $\tau = \Delta t$ , then it constitutes thus a dataset of K + 1 input-output pairs  $(\boldsymbol{g}_k, \boldsymbol{g}_{k+1}), k = 0, 1, \dots, K$  of the operator:

$$\boldsymbol{g}_{k+1} = \mathbb{K}\boldsymbol{g}_k \quad , \qquad \mathbb{K} \equiv \mathcal{K}^{\Delta t}. \tag{26}$$

For example, if the observables  $g_k$  are the states of the system (1) depicted on Figure 1, then  $\Delta t = 10$  MTU, corresponding roughly to one day.

279

## 4.3.2 The Dynamic Mode Decomposition (DMD) algorithm

The DMD algorithm is a data-driven approach that provides a linear decomposi-280 tion of a given signal of input-output pairs into a set of spatial patterns called dynamic 281 modes that are modulated by a damping or growing oscillating factor. The approach was 282 first developed in the climate community under the name LIM (Penland, 1989; Penland 283 & Sardeshmukh, 1995), with its corresponding linear normal modes, as an extension of 284 the Principle Oscillation Patterns (POP) technique of Hasselmann (1988) and Von Storch 285 et al. (1988). It was later rediscovered in the fluid mechanics community by Schmid (2010) as an extension of the proper orthogonal decomposition (POD) (Berkooz et al., 1993). 287 In the framework of dynamical systems like (1), DMD has been identified as an algorithm 288 to approximate the Koopman operator  $\mathcal{K}^{\tau}$  (Rowley et al., 2009; Williams et al., 2015) 289 and the Perron-Frobenius operator  $P^{\tau}$  (Klus et al., 2016). 290

The DMD algorithm identifies two sets of vectors, the *adjoint* DMD modes and the *exact* DMD modes (Tu et al., 2014). The former are approximately related to the eigenfunctions of the Koopman operator, while the latter are related to the Koopman modes.

For this purpose, the input-output pairs  $(\boldsymbol{g}_k, \boldsymbol{g}_{k+1})$  are stacked as the columns of two matrices  $\mathsf{X} = [\boldsymbol{g}_0 \dots \boldsymbol{g}_{K-1}]$  and  $\mathsf{Y} = [\boldsymbol{g}_1 \dots \boldsymbol{g}_K]$ . When the time steps are evenly partitioned, this is simply a repeated representation of the dataset, offset by one timestep. It is assumed that a matrix  $\mathsf{M}^{\mathrm{DMD}}$  exists that approximates the operator  $\mathcal{K}^{\tau}$  so that,

$$\mathbf{Y} = \mathbf{M}^{\mathrm{DMD}} \mathbf{X},\tag{27}$$

and thus,

$$\mathsf{M}^{\mathrm{DMD}} = \mathsf{Y} \mathsf{X}^+, \tag{28}$$

where  $\mathsf{X}^+$  is the pseudoinverse of  $\mathsf{X}.$  Alternatively, the matrix  $\mathsf{M}^{\mathrm{DMD}}$  is sometimes written

$$\mathsf{M}^{\mathrm{DMD}} = \mathsf{A} \,\mathsf{G}^+,\tag{29}$$

where  $A = Y X^*$  and  $G = X X^*$  (Klus et al., 2018). The matrix  $M^{DMD}$  approximates the operator K in the least-squares sense. The eigenvalues and the right eigenvectors of  $M^{DMD}$  are called the DMD eigenvalues and DMD modes of the data. In practice, the eigendecomposition of  $M^{DMD}$  can be performed with the SVD (Tu et al., 2014), or using the Arnoldi algorithm (Rowley et al., 2009). The SVD is computed as  $X = U\Sigma V^*$ . A truncated form can be defined to permit a reduced dimension form of  $M^{DMD}$ . In that case, the equation (28) can be transformed as,

$$\tilde{\mathsf{M}}^{\mathrm{DMD}} = \mathsf{U}^* \mathsf{M}^{\mathrm{DMD}} \mathsf{U} = \mathsf{U}^* \mathsf{Y} \mathsf{V} \Sigma^{-1}. \tag{30}$$

The nonzero eigenvalues  $\lambda_i^{\text{DMD}}$  of  $\mathsf{M}^{\text{DMD}}$  are the same as those of  $\tilde{\mathsf{M}}^{\text{DMD}}$ . The right eigenvectors  $\tilde{\boldsymbol{v}}_i$  of  $\tilde{\mathsf{M}}^{\text{DMD}}$  can be used to recover the corresponding right eigenvectors  $\boldsymbol{v}_i = \frac{1}{\lambda_i^{\text{DMD}}}\mathsf{YV}\Sigma^{-1}\tilde{\boldsymbol{v}}_i$  of  $\mathsf{M}^{\text{DMD}}$ . The left eigenvectors  $\tilde{\boldsymbol{w}}_i$  of  $\tilde{\mathsf{M}}^{\text{DMD}}$  can be used to recover the left eigenvectors  $\boldsymbol{w}_i = \mathsf{U}\,\tilde{\boldsymbol{w}}_i$  of  $\mathsf{M}^{\text{DMD}}$ , satisfying the biorthonormality condition,

$$\boldsymbol{w}_{i}^{*}\boldsymbol{v}_{j} = \tilde{\boldsymbol{w}}_{i}^{*}\boldsymbol{\mathsf{U}}^{*} \frac{1}{\lambda_{j}^{\mathrm{DMD}}}\,\boldsymbol{\mathsf{Y}}\boldsymbol{\mathsf{V}}\boldsymbol{\Sigma}^{-1}\,\tilde{\boldsymbol{v}}_{j} = \tilde{\boldsymbol{w}}_{i}^{*}\,\frac{1}{\lambda_{j}^{\mathrm{DMD}}}\tilde{\boldsymbol{\mathsf{M}}}^{\mathrm{DMD}}\,\tilde{\boldsymbol{v}}_{j} = \tilde{\boldsymbol{w}}_{i}^{*}\,\tilde{\boldsymbol{v}}_{j} = \delta_{i,j}$$
(31)

where we have assumed that the left and right eigenvectors of  $M^{\text{DMD}}$  are scaled in order to form biorthonormal bases. The left eigenvectors  $\boldsymbol{w}_i$  are called the *adjoint* DMD modes, while the right eigenvectors  $\boldsymbol{v}_i$  are called the *exact* DMD modes (Tu et al., 2014).

The left eigenvectors  $\boldsymbol{w}_i$  of  $\mathsf{M}^{\mathrm{DMD}}$  can be used to produce approximations of the Koopman (KM) eigenfunctions. As shown by (Williams et al., 2015), if the eigenfunctions of the Koopman operator are approximated as

$$\phi_i(\boldsymbol{x}) \approx \boldsymbol{w}_i^* \, \boldsymbol{g}(\boldsymbol{x}). \tag{32}$$

then any observable h can be decomposed according to

$$\boldsymbol{h}(\boldsymbol{x}) = \sum_{i=1}^{P} \boldsymbol{c}_{i}^{\text{DMD}} \, \boldsymbol{w}_{i}^{*} \, \boldsymbol{g}(\boldsymbol{x})$$
(33)

and applying the operator  $\mathbb{K}$ , we recover a time-discretized vector-valued approximation of Eq. (21):

$$\mathbb{K}\boldsymbol{h}(\boldsymbol{x}) = \sum_{i=1}^{P} \boldsymbol{c}_{i}^{\text{DMD}} \boldsymbol{w}_{i}^{*} \mathbb{K}\boldsymbol{g}(\boldsymbol{x})$$

$$\approx \sum_{i=1}^{P} \boldsymbol{c}_{i}^{\text{DMD}} \boldsymbol{w}_{i}^{*} \mathbb{M}^{\text{DMD}} \boldsymbol{g}(\boldsymbol{x})$$

$$= \sum_{i=1}^{P} \boldsymbol{c}_{i}^{\text{DMD}} \lambda_{i}^{\text{DMD}} \boldsymbol{w}_{i}^{*} \boldsymbol{g}(\boldsymbol{x})$$
(34)

In particular, if the observable g is the identity (g(x) = x), then the decomposition (33) is analogous to a one-term Taylor expansion of h (Williams et al., 2015). In the remainder, references to the KM eigenfunctions are made under the assumption that they are approximated using this procedure, and due to Eq. (32), it may refer equivalently to the eigenfunctions  $\phi_i$  or the left eigenvectors  $w_i$ .

Finally, the coefficients  $c_i^{\text{DMD}}$  are provided by the right eigenvectors  $v_i$ , i.e. the *DMD* modes approximating the Koopman modes  $c_i^{\text{KM}}$  (see for instance a trivial example in the Section 4.3.4 below where  $c_i^{\text{DMD}} = v_i$ ).

#### 4.3.3 The Perron-Frobenius mode decomposition

Since the Perron-Frobenius operator is the adjoint of the Koopman operator, it is also possible to obtain a finite dimensional representation of the former with this relation, as shown by Klus et al. (2016). Because we are working in the space of observables, we can access the eigenfunctions of the Perron-Frobenius operator using the adjoint property with the Koopman operator (e.g. using the inner product (19)). The finite dimensional representation of the Perron-Frobenius operator is given by,

$$\mathsf{M}^{\mathrm{PFMD}} = \mathsf{A}^{\mathsf{T}}(\mathsf{G}^{+})^{\mathsf{T}}.$$
(35)

where again  $A = Y X^*$  and  $G = X X^*$ .

312

322

Similarly as for the DMD decomposition and the Koopman operator, a distribution  $\rho(\boldsymbol{x})$  can be decomposed on the left eigenvectors  $\boldsymbol{\omega}_i$  of  $\mathsf{M}^{\mathrm{PFMD}}$  as

$$\rho(\boldsymbol{x}) = \sum_{i=1}^{P} c_i^{\text{PFMD}} \,\boldsymbol{\omega}_i^* \,\boldsymbol{g}(\boldsymbol{x})$$
(36)

where the eigenfunctions  $\psi_i$  of the Perron-Frobenius operator are thus approximated as,

$$\psi_i(\boldsymbol{x}) \approx \boldsymbol{\omega}_i^* \, \boldsymbol{g}(\boldsymbol{x}). \tag{37}$$

The decomposition (36) of the densities is thus a time-discretized approximation of Eq. (22), and we call it a Perron-Frobenius mode decomposition (PFMD). In the following, references to the PF eigenfunctions are made under the assumption that they are approximated using this procedure, and may refer equivalently to the eigenfunctions  $\psi_i$  or the left eigenvectors  $\omega_i$ .

Finally, we note that the Perron-Frobenius operator being considered here is defined with respect to the invariant distribution of the system (Klus et al., 2018), since the matrix M<sup>PFMD</sup> is constructed from a long trajectory of the system dynamics.

#### 4.3.4 Projections of ensemble distributions

From now on, we assume that the observable g used to obtain the representations 323  $M^{\text{DMD}}$  and  $M^{\text{PFMD}}$  is the identity: g(x) = x. For instance, for a dynamical system, 324 the datasets X and Y considered thus consists of observed states of the system. For a given 325 ensemble of initial condition perturbations  $\delta x_0^m$  of the state of the system (1), these can 326 be projected onto a subset of the KM left eigenvectors  $w_i$  or onto a subset of the PF left 327 eigenvectors  $\omega_i$ . In the first case, it decomposes the perturbations - viewed as local ob-328 servables - onto selected (approximate) eigenfunctions of the Koopman operator of the 329 system. This subset of eigenfunctions allows one to (approximately) reduce the action 330 of the Koopman operator on a given invariant subspace of this operator, which is char-331 acterized by the left eigenvectors  $\boldsymbol{w}_i$  and the 'timescales'  $\lambda_i^{\text{DMD}}$ . The propagation of the 332 projected ensemble of initial conditions by system (1) is then assumed to be equivalent 333 to the action of the Koopman operator  $\mathcal{K}^t \delta x_0^m$  restricted on this invariant subspace. 334

Let's be more precise about these projections: For a given observable h evaluated on a perturbed state  $x + \delta x$ , we have:

$$h(x + \delta x) \approx h(x) + \nabla_x h \,\delta x \tag{38}$$

The second term is a local approximation of the observable h around the unperturbed state x, and whose time evolution is well represented by the DMD decomposition. If the observable h is the identity (h(x) = x), we have naturally  $\nabla_x h = I$  where I is the identity matrix, and  $h(x + \delta x) = h(x) + h(\delta x)$ . Therefore, one can decompose the perturbation according to Eq. (33) to get:

$$\boldsymbol{h}(\boldsymbol{\delta x}) = \mathsf{C}^{\mathrm{DMD}} \,\mathsf{W}^* \,\boldsymbol{g}(\boldsymbol{\delta x}) = \mathsf{C}^{\mathrm{DMD}} \,\mathsf{W}^* \,\boldsymbol{\delta x} \tag{39}$$

where W is the column matrix of left eigenvectors  $\boldsymbol{w}_i$  of  $\mathsf{M}^{\mathrm{DMD}}$ . Since the observable  $\boldsymbol{h}$  is now the identity we have - due to the biorthonormality relationship (31) - that the matrix  $\mathsf{C}^{\mathrm{DMD}}$  is given by  $\mathsf{C}^{\mathrm{DMD}} = \mathsf{V}$  where  $\mathsf{V}$  is the column matrix of right eigenvectors  $\boldsymbol{v}_i$  of  $\mathsf{M}^{\mathrm{DMD}}$ . According to Eq. (32), the decomposition above is a decomposition in terms of the (approximated) eigenfunctions  $\phi_i(\delta \boldsymbol{x}) \approx (\mathsf{W}^* \delta \boldsymbol{x})_i$ . Projecting the perturbation  $\delta \boldsymbol{x}$  onto a subset of KM eigenfunctions is thus equivalent to making the expansion above according to a partial choice  $\mathsf{W}'^* \delta \boldsymbol{x}$  of eigenfunctions, where  $\mathsf{W}'$  is a column matrix composed of a choice of columns from the matrix W, i.e. a choice amongst the left eigenvectors of  $\mathsf{M}^{\mathrm{DMD}}$ . The projected perturbation is thus developed as:

$$\delta \boldsymbol{x}' = \mathsf{V}' \,\mathsf{W}'^* \,\delta \boldsymbol{x} \tag{40}$$

where V' are the right eigenvectors biorthonormal to the left eigenvectors W'. Identifying B = W' in Eq. (11) for the projector  $\Pi$ , we get

$$\Pi = W' (W'^* W')^{-1} W'^* \tag{41}$$

and using the fact that  $W'^* V' = I$ , where I is the identity matrix, we have  $V' = W' (W'^* W')^{-1}$ . Therefore, Eq. (40) is exactly the projected perturbation  $\delta x' = \prod \delta x$  obtained with the projector (11).

Similarly, in the case where the perturbations are projected onto a selected subset of the (approximated) eigenfunctions  $\psi_i(\delta x) \approx \omega_i^* g(\delta x)$  of the Perron-Frobenius operator of the system, because this subset forms an invariant subspace of the Perron-Frobenius operator, one can assume that the propagation of the projected ensemble of initial conditions with system (1) is equivalent to the action of the Perron-Frobenius operator on the projection of the distribution  $\rho^{\text{ens}}$  of the ensemble.

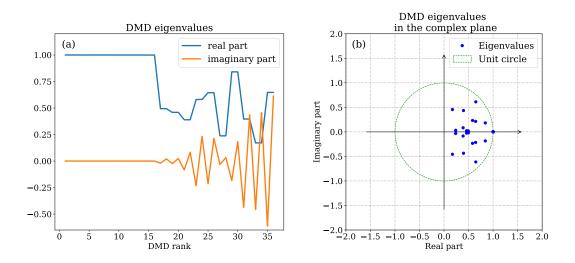
344

345

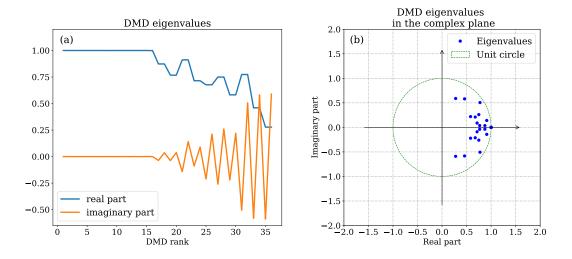
## 4.3.5 The Koopman and Perron-Frobenius eigenfunctions of the coupled ocean-atmosphere model

To study the dynamic modes in the coupled ocean-atmosphere system, the KM eigen-346 functions have been estimated using the data of the reference trajectories depicted in Fig-347 ure 1 sampled every  $\Delta t = 10$  MTU (roughly every day), using the SVD method described 348 in Section 4.3.2. The results are shown in Figure 3 for the weak LFV case, and in Fig-349 ure 4 for the strong LFV case. In both cases, we note that there are 16 eigenvalues in 350 the vicinity of the point 1 + 0 i in the complex plane. These eigenvalues correspond to 351 very slow decaying and oscillating KM eigenfunctions, describing the LFV signal in the 352 system. The remaining eigenvalues are related to faster decaying oscillations. The am-353 plitude of each component of the KM eigenfunctions is shown in Figure 5. Each KM eigen-354 function is a complex-valued vector, and is paired with another KM eigenfunction that 355 is its complex conjugate (except for the presence of real eigenvalues), each correspond-356 ing to complex conjugate eigenvalues. For this reason, Figure 5 shows both the real and 357 imaginary parts of the KM eigenfunctions every two columns. A clear distinction can 358 be made between the slow decaying KM eigenfunctions and the others. Indeed, the slow-359 decaying KM eigenfunctions (1 to 16) involve both the ocean streamfunction variables 360 (variables 21 to 28) and temperature variables (variables 29 to 36), with a predominance 361 of the streamfunction variables. The fast-decaying KM eigenfunctions (17 to 36) involve 362 the ocean streamfunction variables with a coupling to the atmospheric variables (vari-363 ables 1 to 20), and a far weaker coupling to the ocean temperature variables. 364

The PF eigenfunctions have been obtained by directly computing the eigenvectors of the finite dimensional representation of the Perron-Frobenius operator. They possess the same spectrum of eigenvalues as the KM eigenfunctions (see the Supplementary Materials), and while being different, they share the same global slow-fast organization as the KM eigenfunctions (see Figure 5).



**Figure 3.** Koopman eigenvalues estimated using DMD for the case without low-frequency variability.



**Figure 4.** Koopman eigenvalues estimated using DMD for the case with low-frequency variability.

Figure 5. Comparison of the averaged energy of the Lyapunov vectors components with the EOF, KM, and PF eigenfunctions patterns for: (a) the case with a weak low-frequency variability, (b) the case with a strong low-frequency variability.

## <sup>370</sup> 5 Selected bases for experiments

Finally, we choose a speci c set of bases using the methods described above. Each of the methods determine a set of basis vectors that de ne the entire state space. We will split each of these into subspaces onto which to project the Y `perfect' ensemble perturbations. The projected perturbations hence obtained will be used to obtain the ensemble initial conditions of each experiment in the next section, according to the formula (13).

In Figure 6, we show the estimated percent explained variance for the EOFs of the 377 ocean-atmosphere coupled quasi-geostrophic system. The EOF modes are shown in Fig-378 ure 5, with each mode independently normalized to unit magnitude. The leading EOFs 379 explaining most of the variance are related to the ocean temperature and the atmospheric 380 streamfunction variables. We note that the last 8 EOFs, while explaining very little of 381 the total variance, have a qualitatively di erent pattern than the other modes, with a 382 dominant component along the ocean streamfunction. Therefore, the di erent bases 383 of EOFs that we have selected for the experiments are the following: 384

385	The rst 12 EOFs, $U_{1:12}$ , which account for the most signi cant part of the vari-
386	ability
387	The last 8 EOFs, U <sub>29:36</sub> , which have a qualitatively di erent pattern from the oth-
388	ers
389	The remaining 16 EOFs, U <sub>13:28</sub> , which display a more uniform distribution across
390	the di erent model elds
	We shall computed arbitrarily divide the CLVs, adjaint CLVs, PLVs into 2 parts:
391	We shall somewhat arbitrarily divide the CLVs, adjoint CLVs, BLVs into 2 parts:

Figure 7. DSSS skill score summed over components at di erent lead times for the case without low-frequency variability. The lower the DSSS score, the better. The methods are sorted by increasing total score value over all four components.

ensemble forecasts. Moreover, a decomposition of the observables in terms of the CLVs
 and the adjoint CLVs on the tangent linear space yields a similar structure as the one
 of the DMD decomposition, as shown in Appendix A, Section A3.

Finally, we note that while the UNN adjoint CLVs yield reliable forecasts, the KM 547 and PF eigenfunctions are similar in terms of performance but are much simpler and more 548 straightforward to compute using the DMD algorithm. While the computation of the 549 CLVs typically requires the integration of the tangent linear model over long time pe-550 riods, both forward and reverse in time, the KM and PF eigenfunctions can be computed 551 from data produced either by numerical simulations, observational analysis products, or 552 reanalysis products, requiring only an e cient algorithm to perform the SVD decom-553 position. 554

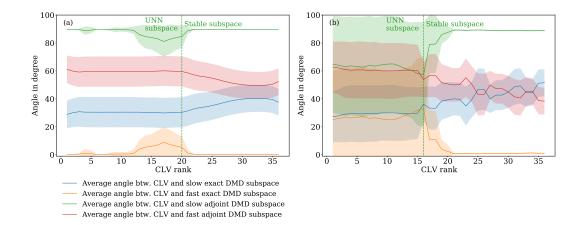


Figure 11. Averaged angle in degrees between the Covariant Lyapunov Vectors (CLVs) and the Dynamic Modes (DMDs), for: (a) the case without low-frequency variability and (b) the case with low-frequency variability. The one standard deviation intervals are depicted by the shaded area. The slow and fast exact DMD subspaces are spanned by the right eigenvectors  $v_i$ (the DMD modes), for respectively  $i \in \{1, \ldots, 16\}$  and  $i \in \{17, \ldots, 36\}$ , while the slow and fast adjoint DMD subspaces are spanned by the left eigenvectors  $w_i$  (the KM eigenfunctions), again for respectively  $i \in \{1, \ldots, 16\}$  and  $i \in \{17, \ldots, 36\}$ . See Section 5 for an explanation of the slow-fast separation on the modes and eigenfunctions. Note that due to the biorthormality relationship (31) between the vectors  $v_i$  and  $w_i$ , the slow exact DMD subspace is orthogonal to the fast adjoint DMD subspace, while the fast exact DMD subspace is orthogonal to the slow adjoint DMD subspace. The separation between the CLVs belonging to the UNN and stable subspace is depicted by a vertical dashed line.

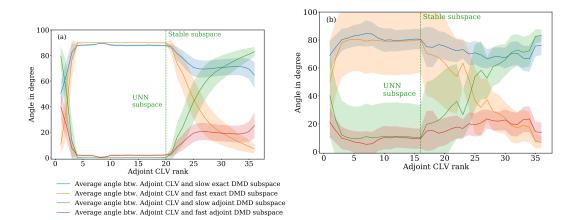


Figure 12. Averaged angle in degrees between the adjoint Covariant Lyapunov Vectors (adjoint CLVs) and the Dynamic Modes (DMDs), for: (a) the case without low-frequency variability and (b) the case with low-frequency variability. The one standard deviation intervals are depicted by the shaded area. The slow and fast exact DMD subspaces are spanned by the right eigenvectors  $\mathbf{v}_i$  (the DMD modes), for respectively  $i \in \{1, \ldots, 16\}$  and  $i \in \{17, \ldots, 36\}$ , while the slow and fast adjoint DMD subspaces are spanned by the left eigenvectors  $\mathbf{w}_i$  (the KM eigenfunctions), again for respectively  $i \in \{1, \ldots, 16\}$  and  $i \in \{17, \ldots, 36\}$ . See Section 5 for an explanation of the slow-fast separation on the modes and eigenfunctions. Note that due to the biorthormality relationship (31) between the vectors  $\mathbf{v}_i$  and  $\mathbf{w}_i$ , the slow exact DMD subspace is orthogonal to the fast adjoint DMD subspace. The separation between the adjoint CLVs belonging to the UNN and stable subspace is depicted by a vertical dashed line.

## 555 7 Conclusions

In this work, the impact of the choice of the initial perturbations on ensemble fore-556 casts of coupled ocean-atmosphere systems has been addressed by investigating a reduced-557 order coupled model. Different types of perturbations have been selected, including tra-558 ditional approaches like the Empirical Orthogonal Functions and the Lyapunov vectors, 559 but also novel approaches based on the Dynamic Mode Decomposition which has been 560 noted in recent years as a reasonable computational approximation of the modes and eigen-561 functions of the Koopman and Perron-Frobenius operators (Rowley et al., 2009; Tu et 562 al., 2014). After a detailed analysis of the different definitions of the DMD modes that 563 are found in the literature, their use as a basis for initializing ensemble forecasts was ex-564 plored. 565

A key result is that projecting initial perturbations onto the fast-decaying KM eigen-566 functions and PF eigenfunctions – which refer here to linear approximations of the eigenfunctions of the Koopman and Perron-Frobenius operators – provides reliable ensemble 568 forecasts in the system at hand and at the considered lead times. This further suggests 569 that these eigenfunctions are essential for providing reliable ensemble forecasts. More-570 over, they seems to be less sensitive to the model's regime and local predictability than 571 other methods. Another important result is the usefulness of the adjoint CLVs, which 572 can be seen as eigenfunctions of the Koopman operator in the tangent space of the sys-573 tem trajectory. The adjoint CLVs also provide reliable ensemble forecasts. A key differ-574 ence between the KM and PK eigenfunctions and the adjoint CLVs lies in the fact that 575 the former are defined globally over the attractor of the system, while the latter are lo-576 cal properties of the flow. In an operational setting, the adjoint CLVs would therefore 577 be quite difficult to compute. On the other hand, it is straightforward to compute an 578 estimate of the KM and PF eigenfunctions directly from data with the DMD method, 579 which provides significant flexibility in their computation and use. 580

This thought experiment should now be expanded in a more realistic setting by in-581 vestigating the use of these tools in intermediate order climate models. In this frame-582 work, a first research question is related to the validity of the DMD-estimated KM spec-583 trum of the systems being considered: In the present considered system, spectra that are 584 clearly identifiable and separable were found. However, it is known that chaotic systems 585 possess complicated spectra (Arbabi & Mezić, 2017; Mezić, 2020) which contains con-586 tinuous components. These complicated spectra might hamper the application of the present 587 method to real datasets or to high-resolution models, the DMD analysis providing too 588 few relevant patterns to work with. This will have to be investigated, notably in systems 589 where the dimension is too high to apply the DMD method directly, and have thus to 590 be reduced first. 591

Another research question concerns the other sources of uncertainty affecting the 592 ensemble forecasts. As previously noted, systematic errors in the forecast model share 593 roughly equal importance with the specification of initial conditions in producing accu-594 rate and reliable forecasts. In an operational setting, it is important to take these sys-595 tematic model errors into account. A possible path forward is to evaluate the projection 596 of assumed model errors onto the KM or PF eigenfunctions, and randomly perturbing 597 the model in that direction. This question will be explored in the future in the context 598 of the current model. 599

Finally, a few important steps toward an operational implementation of the DMD approach are still needed: First to investigate the impact of data assimilation on the statistics of the initial error and their projections on the KM and PF eigenfunctions, and second to compare the DMD approach to the singular vector techniques that are often used for ensemble initialization and for the propagation of the error covariances, e.g. (Ehrendorfer & Tribbia, 1997). These steps are planned in a future investigation.

#### 606 Acknowledgments

625

An earlier version of this manuscript benefited from insightful suggestions from Tom Hamill. J.D. and S.V. acknowledge partial support from ROADMAP, a coordinated JPI-Climate/JPI-

<sup>609</sup> Oceans project, financed by the Belgian Science Policy under contract B2/20E/P1/ROADMAP.

S.G.P. acknowledges support from the Office of Naval Research (ONR) grants N00014-

<sup>611</sup> 19-1-2522 and N00014-20-1-2580 and from National Oceanographic and Atmospheric Ad-

ministration (NOAA) grants NA18NWS4680048, NA19NES4320002, NA20OAR4600277,

and NA20NWS4680053. S.G.P. would like to thank Tom Hamill for discussions on S2S

<sub>614</sub> prediction, and Cecile Penland for discussions on the LIM.

615 Code availability: The code used to obtain the trajectories and Lyapunov vectors of

the *VDDG* model is qgs (Demaeyer et al., 2020). It is available at https://github.com/ Climdyn/qgs and on Zenodo (Demaeyer, 2021). A version of the computation notebooks

suitable for publication is in preparation and will be released by the end of the reviewprocess.

## Appendix A Lyapunovs vectors (BLVs, CLVs, and their adjoints)

In dynamical systems described by a set of ODEs like (1), vectors can be defined to describe the local linear stability around its solutions. These vectors can be obtained by considering the linearization of Eq. (1) around such a solution  $\boldsymbol{x}(\tau)$ :

$$\dot{\boldsymbol{\delta x}}(\tau) = \left. \frac{\partial \boldsymbol{f}}{\partial \boldsymbol{x}} \right|_{\boldsymbol{x}(\tau)} \, \boldsymbol{\delta x}(\tau) \tag{A1}$$

where  $\partial f / \partial x$  is the Jacobian matrix of f. The solution of the linearized equation can be formally written as

$$\boldsymbol{\delta x}(t) = \mathsf{M}(t, t_0) \, \boldsymbol{\delta x}_0 \qquad , \quad \boldsymbol{\delta x}_0 = \boldsymbol{\delta x}(t_0) \tag{A2}$$

where  $M(t, t_0) = \nabla_{\boldsymbol{x}(t_0)} \phi^{t-t_0}$  is the fundamental matrix of solutions of the system (A1), i.e. the Jacobian matrix of the flow  $\phi^{t-t_0}$  of (1), and is thus identified with the linear propagator that propagates the perturbations in the tangent space of  $\boldsymbol{x}(\tau)$  between the times  $t_0$  and t.

## A1 Osedelets splitting of the tangent space

The Osedelets theorem (V. I. Oseledets, 1968; V. Oseledets, 2008) states that the term  $(\mathsf{M}(t,t_0)\mathsf{M}(t,t_0)^{\mathsf{T}})^{1/(2(t-t_0))}$  is well defined in the limit  $t_0 \to \infty$ . Its eigenvectors and the logarithm of its eigenvalues are respectively the Backward Lyapunov Vectors (BLVs)  $\varphi_i^-(t)$  at the time t and the Lyapunov exponents  $\sigma_i$  of the system. The set of the Lyapunov exponents is sometimes called the Lyapunov spectrum and is assumed here to be sorted in decreasing order. The vectors  $\varphi_i^-(t)$  are orthogonal and span a set of subspaces

$$S_{j}^{-}(t) = \operatorname{span}\{\varphi_{i}^{-}(t)|i = 1, 2, \dots, j\}$$
,  $j = 1, \dots, d$ 
(A3)

toward which any *j*-arbitrary volume  $V_j(t_0)$  (dim  $V_j(t_0) = j$ ) defined at a time  $t_0$  in the far past converges under the action of the propagator<sup>6</sup>:

$$\lim_{t_0 \to -\infty} \mathsf{M}(t, t_0) V_j(t_0) \subset S_j^-(t).$$
(A4)

By construction, we have  $S_1^- \subset S_2^- \subset \ldots \subset S_{d-1}^- \subset S_d^-$  which is called a Osedelets splitting of the tangent space at the time t (Kuptsov & Parlitz, 2012). The BLVs thus

 $<sup>^{6}</sup>$  For the sake of simplicity, we present here the case where there are no degenerate Lyapunov exponents in the spectrum. The general case is presented in Kuptsov and Parlitz (2012).

span and describe volumes of the tangent space that are reached asymptotically at a given time by arbitrary volumes defined in the far past, and are thus preserved under the tangent flow

$$\mathsf{M}(t, t_0) S_i^-(t_0) = S_i^-(t). \tag{A5}$$

Similarly, one can take the limit of the matrix  $(\mathsf{M}(t,t_0)^{\mathsf{T}}\mathsf{M}(t,t_0))^{1/(2(t-t_0))}$  for  $t \to \infty$  and its eigenvectors are the Forward Lyapunov Vectors (FLVs)  $\varphi_i^+(t)$ . Its eigenvalues are also the Lyapunov exponents  $\sigma_i$ . The vectors  $\varphi_i^+(t)$  are orthogonal as well and span a set of subspaces

$$S_{j}^{+}(t) = \operatorname{span}\{\varphi_{i}^{+}(t)|i = j, j + 1, \dots, d\}$$

$$, \quad j = 1, \dots, d$$
(A6)

toward which any arbitrary *j*-volume  $V_j(t)$  (dim  $V_j(t) = j$ ) defined at a time *t* in the far future converges under the action of the time-reversed propagator:

$$\lim_{t \to \infty} \mathsf{M}(t_0, t) V_j(t) \subset S_j^+(t_0).$$
(A7)

By construction, we have the sequence  $S_d^+ \subset S_{d-1}^+ \subset \ldots \subset S_2^+ \subset S_1^+$  which forms another Osedelets splitting of the tangent space at the time  $t_0$ . The FLVs thus span and describe volumes of the tangent space that are reached asymptotically<sup>7</sup> at a given time by arbitrary volumes defined in the far future. These volumes are thus preserved under the time-reversed tangent flow

$$\mathsf{M}(t_0, t)S_i^+(t) = S_i^+(t_0). \tag{A8}$$

#### A2 Covariant Lyapunov Vectors and their adjoint

The Covariant Lyapunov Vectors (CLVs) are vectors  $\varphi_i$  such that when the linear propagator M is applied to them, one obtains

$$\mathsf{M}(t,t_0)\,\varphi_i(t_0) = \Lambda_i(t,t_0)\,\varphi_i(t). \tag{A9}$$

and the linearized dynamics (A2) transports the CLVs from a time  $t_0$  onto the CLVs at time t further downstream the trajectory  $\boldsymbol{x}(\tau)$  by multiplying by a *stretching* factor  $\Lambda_i$  (Gaspard, 2005; Kuptsov & Parlitz, 2012). The CLVs can thus be shown to be the solutions of the equation

$$\dot{\boldsymbol{\varphi}}_{i}(\tau) = \left. \frac{\partial \boldsymbol{f}}{\partial \boldsymbol{x}} \right|_{\boldsymbol{x}(\tau)} \, \boldsymbol{\varphi}_{i}(\tau) - \chi_{i}(\tau) \, \boldsymbol{\varphi}_{i}(\tau) \tag{A10}$$

with

$$\Lambda_i(t, t_0) = \exp\left\{\int_{t_0}^t \chi_i(\tau) \,\mathrm{d}\tau\right\}$$
(A11)

where  $\chi_i(\tau)$  is the *local* stretching rate at time  $\tau$ . The global Lyapunov exponents of the system are recovered in the limit as  $t \to \infty$ ,

$$\sigma_i = \lim_{t \to \infty} \frac{1}{t} \ln |\Lambda_i(t, t_0)| = \lim_{t \to \infty} \frac{1}{t} \int_{t_0}^t \chi_i(\tau) \,\mathrm{d}\tau.$$
(A12)

By definition, each CLV lies at the intersection between the Osedelets subspaces  $S_j^-$  and  $S_j^+$  (Eckmann & Ruelle, 1985),

$$\boldsymbol{\varphi}_j(t) \in S_j^+(t) \cap S_j^-(t). \tag{A13}$$

 $<sup>^{7}</sup>$  Under the evolution of the time-reversed tangent flow.

The linear propagator M can be decomposed in terms of the CLVs  $\varphi_i$  and their corresponding stretching factors  $\Lambda_i$  as

$$\mathsf{M}(t,t_0) = \sum_{i=1}^{d} \varphi_i(t) \Lambda_i(t,t_0) \tilde{\varphi}_i^{\mathsf{T}}(t_0)$$
(A14)

where the vectors  $\tilde{\varphi}_i$  are the adjoint Covariant Lyapunov Vectors satisfying the biorthonormality relation with the CLVs:

$$\tilde{\varphi}_i^{\mathsf{T}} \varphi_j = \delta_{i,j} \tag{A15}$$

at any point of the phase space of the system (Gaspard, 2005). The adjoint CLVs are solutions of the adjoint of Eq. (A10),

$$\dot{\tilde{\varphi}}_{i}(\tau) = \left. \frac{\partial \boldsymbol{f}}{\partial \boldsymbol{x}}^{\mathsf{T}} \right|_{\boldsymbol{x}(\tau)} \tilde{\varphi}_{i}(\tau) - \chi_{i}(\tau) \,\tilde{\varphi}_{i}(\tau) \tag{A16}$$

and are covariant with respect to the adjoint dynamics,

$$\mathsf{G}(t,t_0)\,\tilde{\boldsymbol{\varphi}}_i(t_0) = \Lambda_i^{-1}(t,t_0)\,\tilde{\boldsymbol{\varphi}}_i(t) \tag{A17}$$

with  $G(t,t_0) = (M(t,t_0)^{-1})^T$ , but they are multiplied by the inverse of the stretching factor (Kuptsov & Parlitz, 2012). Note that both Eqs. (A9) and (A17) are time-reversible, with the property imposed by Eq. (A11) that the stretching factors are inverted upon time-reversal,  $\Lambda_i(t,t_0) = \Lambda_i^{-1}(t_0,t)$ . By definition, each adjoint CLV lies at the intersection between the adjoint Osedelets subspaces  $H_j^+$  and  $H_j^-$  (Eckmann & Ruelle, 1985),

$$\tilde{\boldsymbol{\varphi}}_j(t) \in H_j^+(t) \cap H_j^-(t). \tag{A18}$$

which are preserved under the adjoint tangent flow.

$$G(t, t_0)H_j^+(t_0) = H_j^+(t),$$
(A19)

$$\mathsf{G}(t,t_0)H_i^-(t_0) = H_i^-(t). \tag{A20}$$

These subspaces form Osedelets splittings of the tangent space as well,

$$H_i^+(t) = \text{span}\{\varphi_i^+(t)|i=1,2,\dots,j\}$$
 (A21)

$$H_j^-(t) = \operatorname{span}\{\varphi_i^-(t)|i=j, j+1, \dots, d\}$$

$$\forall j = 1, \dots, d.$$
(A22)

To summarize, the BLVs and FLVs can be interpreted as orthonormal basis of vectors defining volumes covariant with the dynamics, while the CLVs and adjoint CLVs define directions in the tangent space that are covariant with the dynamics.

#### 630

#### A3 Koopman operator of the tangent flow

For a given observable g of a system like (1), the time-evolution starting at time  $t_0 = 0$  of the observables in the neighborhood of a given state  $x_0$  can be approximated by

$$\mathcal{K}^{t}g(\boldsymbol{x}_{0} + \boldsymbol{\delta}\boldsymbol{x}_{0}) = g\left(\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0} + \boldsymbol{\delta}\boldsymbol{x}_{0})\right)$$
  

$$\approx g\left(\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0}) + \left(\boldsymbol{\nabla}_{\boldsymbol{x}_{0}}\boldsymbol{\phi}^{t}\right)\boldsymbol{\delta}\boldsymbol{x}_{0}\right)$$
  

$$\approx g\left(\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0})\right) + \left(\boldsymbol{\nabla}_{\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0})}g\right)^{\mathsf{T}}\left(\boldsymbol{\nabla}_{\boldsymbol{x}_{0}}\boldsymbol{\phi}^{t}\right)\boldsymbol{\delta}\boldsymbol{x}_{0}$$
  

$$= g\left(\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0})\right) + \left(\boldsymbol{\nabla}_{\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0})}g\right)^{\mathsf{T}}\mathsf{M}(t,0)\boldsymbol{\delta}\boldsymbol{x}_{0} \qquad (A23)$$

On the other hand, one can naturally define a Koopman operator  $\mathcal{T}_{\boldsymbol{x}_0}^t$  on the tangent linear space of a given trajectory  $\boldsymbol{\phi}^t(\boldsymbol{x}_0)$ , its expression being

$$\mathcal{T}_{\boldsymbol{x}_0}^t \,\bar{g}(\boldsymbol{\delta}\boldsymbol{x}) = \bar{g}(\bar{\boldsymbol{\phi}}_{\boldsymbol{x}_0}^t(\boldsymbol{\delta}\boldsymbol{x})) = \bar{g}(\mathsf{M}(t,0)\,\boldsymbol{\delta}\boldsymbol{x}) \tag{A24}$$

where  $\bar{\phi}_{x_0}^t$  and  $\bar{g}$  are respectively the flow and an observable defined on the tangent linear space. The action of the Koopman operator  $\mathcal{K}^t$  of the system (1) approximated by Eq. (A23) in a neighborhood of  $x_0$  can thus be rewritten

$$\mathcal{K}^{t}g(\boldsymbol{x}_{0} + \boldsymbol{\delta}\boldsymbol{x}_{0}) \approx \mathcal{K}^{t}g(\boldsymbol{x}_{0}) + \mathcal{T}^{t}_{\boldsymbol{x}_{0}}\bar{g}(\boldsymbol{\delta}\boldsymbol{x}_{0})$$
(A25)

with  $\bar{g}(\boldsymbol{\delta x}) = \left(\boldsymbol{\nabla}_{\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0})}g\right)^{\mathsf{T}} \boldsymbol{\delta x}$ . Using the decomposition (A14), we get then:

$$\mathcal{T}_{\boldsymbol{x}_{0}}^{t} \bar{g}(\boldsymbol{\delta}\boldsymbol{x}_{0}) = \sum_{i=1}^{d} \left( \boldsymbol{\nabla}_{\boldsymbol{\phi}^{t}(\boldsymbol{x}_{0})} g \right)^{\mathsf{T}} \boldsymbol{\varphi}_{i}(t) \Lambda_{i}(t,0) \tilde{\boldsymbol{\varphi}}_{i}^{\mathsf{T}}(0) \boldsymbol{\delta}\boldsymbol{x}_{0}$$
(A26)

and from this equation, one can see that similarly to the DMD left eigenvectors for the Koopman operator  $\mathcal{K}^t$  presented in section 4.3.2, the adjoint CLVs provide an analogy<sup>8</sup> for the "eigenfunctions" of the first-order Koopman operator  $\delta \mathcal{K}^t$ , whose representation is provided by the linear propagator M. Indeed, if one considers the functions

$$\phi_i^{\mathrm{TL}}(\boldsymbol{\delta x}, t) = \tilde{\boldsymbol{\varphi}}_i^{\mathsf{T}}(t) \, \boldsymbol{\delta x},\tag{A27}$$

it is straightforward, using Eq. (A14), that

$$\mathcal{T}_{\boldsymbol{x}_{0}}^{s} \phi_{i}^{\mathrm{TL}}(\boldsymbol{\delta}\boldsymbol{x}, t) = \tilde{\boldsymbol{\varphi}}_{i}^{\mathsf{T}}(t) \mathsf{M}(t, s) \boldsymbol{\delta}\boldsymbol{x}$$
$$= \Lambda_{i}(t, s) \tilde{\boldsymbol{\varphi}}_{i}^{\mathsf{T}}(s) \boldsymbol{\delta}\boldsymbol{x}$$
$$= \Lambda_{i}(t, s) \phi_{i}^{\mathrm{TL}}(\boldsymbol{\delta}\boldsymbol{x}, s)$$
(A28)

On the other hand, the CLVs span the space of the Koopman modes of the operator  $\delta \mathcal{K}^t$ , and one can rewrite Eq. (A26) as

$$\mathcal{T}_{\boldsymbol{x}_{0}}^{t} \bar{g}(\boldsymbol{\delta}\boldsymbol{x}_{0}) = \sum_{i=1}^{d} c_{i}^{\mathrm{TL}}(t) \Lambda_{i}(t,0) \phi_{i}^{\mathrm{TL}}(\boldsymbol{\delta}\boldsymbol{x}_{0},0)$$
(A29)

which is analogous to Eqs. (21) and (34). However, note that since the time-evolution in the tangent space is given by a non-autonomous system (A1), both the functions (A27) and modes  $c_i^{\text{TL}}(t) = \left(\boldsymbol{\nabla}_{\boldsymbol{x}(t)}g\right)^{\mathsf{T}} \boldsymbol{\varphi}_i(t)$  of this decomposition are time-dependent.

Finally, due to the similarity between Eq. (A29) and Eq. (34), the discussion in Section 4.3.4 about ensemble projections remains appropriate here. Projecting an ensemble of initial conditions  $\delta x_0^m$  on subspaces spanned by the adjoint CLVs and propagating them is tantamount to projecting on invariant subspaces of the Koopman operator  $\mathcal{T}_{x_0}^t$ .

#### 639 **References**

- Arbabi, H., & Mezić, I. (2017). Study of dynamics in post-transient flows using
   Koopman mode decomposition. *Physical Review Fluids*, 2(12), 124402.
- Benettin, G., Galgani, L., Giorgilli, A., & Strelcyn, J.-M. (1980). Lyapunov characteristic exponents for smooth dynamical systems and for Hamiltonian systems;
  a method for computing all of them. Part 1: Theory. *Meccanica*, 15(1), 9–20.
- Berkooz, G., Holmes, P., & Lumley, J. L. (1993). The proper orthogonal decomposition in the analysis of turbulent flows. Annual review of fluid mechanics, 25(1), 539–575.
- Bocquet, M., & Carrassi, A. (2017). Four-dimensional ensemble variational data
   assimilation and the unstable subspace. Tellus A: Dynamic Meteorology and
   Oceanography, 69(1), 1304504. doi: 10.1080/16000870.2017.1304504

 $<sup>^{8}</sup>$  In particular, compare Eq. (A27) with Eq. (32).

651	Bocquet, M., Gurumoorthy, K. S., Apte, A., Carrassi, A., Grudzien, C., & Jones,
652	C. K. R. T. (2017). Degenerate Kalman Filter Error Covariances and Their
653	Convergence onto the Unstable Subspace. SIAM/ASA Journal on Uncertainty
654	Quantification, 5(1), 304–333. doi: 10.1137/16M1068712
655	Buizza, R. (2019). Introduction to the special issue on "25 years of ensemble fore-
656	casting". Quarterly Journal of the Royal Meteorological Society, 145(S1), 1-11.
657	doi: https://doi.org/10.1002/qj.3370
	Buizza, R., Houtekamer, P. L., Pellerin, G., Toth, Z., Zhu, Y., & Wei, M. (2005).
658 659	A Comparison of the ECMWF, MSC, and NCEP Global Ensemble Pre-
660	diction Systems. Monthly Weather Review, 133(5), 1076 - 1097. doi:
661	10.1175/MWR2905.1
662	Carrassi, A., Bocquet, M., Demaeyer, J., Gruzien, C., Raanes, P., & Vannitsem, S.
663	(2021). Data assimilation for chaotic dynamics. In S. K. Park & L. Xu (Eds.),
664	Data Assimilation for Atmospheric, Oceanic and Hydrologic Applications (Vol.
665	<i>IV</i> ) (p. in print). Springer International Publishing, Switzerland.
	Cassou, C., Kushnir, Y., Hawkins, E., Pirani, A., Kucharski, F., Kang, IS., &
666	Caltabiano, N. (2018). Decadal climate variability and predictability: Chal-
667	
668	lenges and opportunities. Bulletin of the American Meteorological Society,
669	99(3), 479 - 490. doi: 10.1175/BAMS-D-16-0286.1
670	Dawid, A. P., & Sebastiani, P. (1999). Coherent dispersion criteria for optimal ex-
671	perimental design. Annals of Statistics, 65–81.
672	De Cruz, L., Demaeyer, J., & Vannitsem, S. (2016). The Modular Arbitrary-Order
673	Ocean-Atmosphere Model: MAOOAM v1.0. Geoscientific Model Development,
674	9(8), 2793-2808.
675	Demaeyer, J. (2021, July). qgs version 0.2.4 release. Zenodo. Retrieved from
676	https://doi.org/10.5281/zenodo.5081145 doi: 10.5281/zenodo.5081145
677	Demaeyer, J., De Cruz, L., & Vannitsem, S. (2020). qgs: A flexible Python frame-
678	work of reduced-order multiscale climate models. Journal of Open Source Soft-
679	ware, 5(56), 2597. doi: 10.21105/joss.02597
680	Demaeyer, J., & Vannitsem, S. (2017). Stochastic parametrization of subgrid-
681	scale processes in coupled ocean–atmosphere systems: benefits and limitations
682	of response theory. Quarterly Journal of the Royal Meteorological Society,
683	143(703), 881–896.
684	Eckmann, JP., & Ruelle, D. (1985). Ergodic theory of chaos and strange attractors.
685	Reviews of Modern Physics, 57(3), 617.
686	Efron, B., & Tibshirani, R. (1993). An introduction to the bootstrap. Chapman and
	Hall, London.
687	
688	Ehrendorfer, M. (2006). The Liouville equation and atmospheric predictability.
689	In T. Palmer & R. Hagedorn (Eds.), <i>Predictability of weather and climate</i>
690	(p. 59–98). Cambridge University Press. doi: 10.1017/CBO9780511617652
691	.005
692	Ehrendorfer, M., & Tribbia, J. J. (1997). Optimal prediction of forecast error covari-
693	ances through singular vectors. Journal of the Atmospheric Sciences, $54(2)$ ,
694	286 - 313. doi: $10.1175/1520-0469(1997)054(0286:OPOFEC)2.0.CO;2$
695	Gaspard, P. (2005). Chaos, scattering and statistical mechanics (Vol. 9). Cambridge
696	University Press.
697	Gaspard, P., Nicolis, G., Provata, A., & Tasaki, S. (1995). Spectral signature of the
698	pitchfork bifurcation: Liouville equation approach. Physical Review $E, 51(1),$
699	74.
700	Giannakis, D. (2019). Data-driven spectral decomposition and forecasting of ergodic
701	dynamical systems. Applied and Computational Harmonic Analysis, 47(2),
702	
	Hasselmann, K. (1988). PIPs and POPs: The reduction of complex dynamical
703 704	systems using principal interaction and oscillation patterns. Journal of Geo-

systems using principal interaction and oscillation patterns. Journal of Geo physical Research: Atmospheres, 93(D9), 11015-11021. doi: https://doi.org/10

706	.1029/JD093iD09p11015
707	Kalnay, E. (2019). Historical perspective: earlier ensembles and forecasting fore-
708	cast skill. Quarterly Journal of the Royal Meteorological Society, 145(S1), 25-
709	34. doi: https://doi.org/10.1002/qj.3595
710	Kleist, D. T., & Ide, K. (2015). An OSSE-based evaluation of hybrid variational-
711	ensemble data assimilation for the NCEP GFS. Part II: 4DEnVar and hybrid
712	variants. Monthly Weather Review, 143(2), 452–470.
	Klus, S., Koltai, P., & Schütte, C. (2016). On the numerical approximation of the
713	Perron-Frobenius and Koopman operator. Journal of Computational Dynam-
714	
715	$ics, \beta(1), 51.$
716	Klus, S., Nüske, F., Koltai, P., Wu, H., Kevrekidis, I., Schütte, C., & Noé, F. (2018).
717	Data-driven model reduction and transfer operator approximation. Journal of
718	Nonlinear Science, 28(3), 985–1010.
719	Kuptsov, P. V., & Parlitz, U. (2012). Theory and computation of covariant lyapunov
720	vectors. Journal of nonlinear science, $22(5)$ , $727-762$ .
721	Lasota, A., & Mackey, M. C. (2008). Probabilistic properties of deterministic sys-
722	<i>tems</i> . Cambridge university press.
723	Legras, B., & Vautard, R. (1996). A guide to Liapunov vectors. In Proceedings 1995
724	ECMWF seminar on predictability (Vol. 1, pp. 143–156).
725	Leutbecher, M. (2019). Ensemble size: How suboptimal is less than infinity? Quar-
726	terly Journal of the Royal Meteorological Society, 145, 107–128.
727	Leutbecher, M., & Palmer, T. N. (2008). Ensemble forecasting. Journal of computa-
728	tional physics, $227(7)$ , $3515-3539$ .
729	Meyer, C. D. (2000). Matrix analysis and applied linear algebra (Vol. 71). Siam.
730	Mezić, I. (2013). Analysis of fluid flows via spectral properties of the Koopman oper-
731	ator. Annual Review of Fluid Mechanics, 45, 357–378.
732	Mezić, I. (2020). Spectrum of the Koopman operator, spectral expansions in func-
733	tional spaces, and state-space geometry. Journal of Nonlinear Science, $30(5)$ ,
734	2091-2145.
735	Molteni, F., Buizza, R., Palmer, T., & Petroliagis, T. (1996). The ECMWF en-
736	semble prediction system: methodology and validation. Quarterly Journal of
737	the Royal Meteorological Society, 122, 73-119. doi: https://doi.org/10.1002/qj
738	.49712
739	Nicolis, G., & Nicolis, C. (2012). Foundations of complex systems: emergence, infor-
740	mation and predicition. World Scientific.
741	Oseledets, V. (2008). Oseledets theorem. Scholarpedia, 3(1), 1846. (revision
742	#142085) doi: 10.4249/scholarpedia.1846
743	Oseledets, V. I. (1968). A multiplicative ergodic theorem. Lyapunov characteristic
744	numbers for dynamical systems. Trans. Moscow Math. Soc., 19, 179–210.
745	O'Kane, T. J., Sandery, P. A., Monselesan, D. P., Sakov, P., Chamberlain, M. A.,
746	Matear, R. J., Stevens, L. (2019). Coupled data assimilation and ensem-
747	ble initialization with application to multiyear ENSO prediction. Journal of
748	Climate, 32(4), 997–1024.
749	Peña, M., & Kalnay, E. (2004). Separating fast and slow modes in coupled chaotic
750	systems. Nonlinear Processes in Geophysics, 11(3), 319–327. doi: 10.5194/npg
751	-11-319-2004
752	Penland, C. (1989). Random forcing and forecasting using principal oscillation pat-
753	tern analysis. Monthly Weather Review, 117(10), 2165–2185.
754	Penland, C., & Magorian, T. (1993). Prediction of Niño 3 sea surface temperatures
755	using linear inverse modeling. Journal of Climate, 6(6), 1067 - 1076. doi: 10
756	.1175/1520-0442(1993)006(1067:PONSST)2.0.CO;2
757	Penland, C., & Sardeshmukh, P. D. (1995). The optimal growth of tropical sea sur-
758	face temperature anomalies. Journal of Climate, 8(8), 1999 - 2024. doi: 10
759	.1175/1520-0442(1995)008(1999:TOGOTS)2.0.CO;2
760	Penny, S., Bach, E., Bhargava, K., Chang, CC., Da, C., Sun, L., & Yoshida, T.
	· · · · · · · · · · · · · · · · · · ·

761	(2019). Strongly coupled data assimilation in multiscale media: Experiments
762	using a quasi-geostrophic coupled model. Journal of Advances in Modeling
763	Earth Systems, 11(6), 1803–1829.
764	Penny, S. G. (2017). Mathematical foundations of hybrid data assimilation from a
765	synchronization perspective. Chaos: An Interdisciplinary Journal of Nonlinear
766	Science, 27(12), 126801. doi: 10.1063/1.5001819
767	Polkova, I., Köhl, A., & Stammer, D. (2019). Climate-mode initialization for decadal
768	climate predictions. Climate Dynamics, 53, 7097-7111.
769	Roulston, M. S., & Smith, L. A. (2002). Evaluating probabilistic forecasts using in-
770	formation theory. Monthly Weather Review, 130(6), 1653–1660.
771	Rowley, C. W., Mezić, I., Bagheri, S., Schlatter, P., & Henningson, D. (2009). Spec-
772	tral analysis of nonlinear flows. Journal of Fluid Mechanics, 641(1), 115–127.
773	Santos Gutiérrez, M., Lucarini, V., Chekroun, M. D., & Ghil, M. (2021). Reduced-
774	order models for coupled dynamical systems: Data-driven methods and the
775	Koopman operator. Chaos: An Interdisciplinary Journal of Nonlinear Science,
776	31(5), 053116.
777	Schmid, P. J. (2010). Dynamic mode decomposition of numerical and experimental
778	data. Journal of fluid mechanics, 656, 5–28.
779	Siegert, S., Ferro, C. A., Stephenson, D. B., & Leutbecher, M. (2019). The
780	ensemble-adjusted Ignorance Score for forecasts issued as normal distributions.
781	Quarterly Journal of the Royal Meteorological Society, 145, 129–139.
782	Susuki, Y., Mezic, I., Raak, F., & Hikihara, T. (2016). Applied Koopman operator
783	theory for power systems technology. Nonlinear Theory and Its Applications,
784	$IEICE, \ 7(4), \ 430-459.$
785	Tantet, A., Lucarini, V., & Dijkstra, H. A. (2018). Resonances in a chaotic attractor
786	crisis of the Lorenz flow. Journal of Statistical Physics, $170(3)$ , 584–616.
787	Tantet, A., Lucarini, V., Lunkeit, F., & Dijkstra, H. A. (2018). Crisis of the chaotic
788	attractor of a climate model: a transfer operator approach. Nonlinearity,
789	31(5), 2221.
790	Tondeur, M., Carrassi, A., Vannitsem, S., & Bocquet, M. (2020). On temporal scale
791	separation in coupled data assimilation with the ensemble Kalman filter. Jour-
792	nal of Statistical Physics, 1–25.
793	Toth, Z., & Kalnay, E. (1997). Ensemble Forecasting at NCEP and the Breeding
794	Method. Monthly Weather Review, 125(12), 3297-3319.
795	Tu, J. H., Rowley, C. W., Luchtenburg, D. M., Brunton, S. L., & Kutz, J. N. (2014).
796	On Dynamic Mode Decomposition: Theory And Applications. Journal of $C_{\text{computational Dynamics}}$ $1(2)$ 201 421
797	Computational Dynamics, $1(2)$ , $391-421$ .
798	Vannitsem, S. (2017). Predictability of large-scale atmospheric motions: Lyapunov exponents and error dynamics. <i>Chaos: An Interdisciplinary Journal of Nonlin</i> -
799	ear Science, 27(3), 032101.
800	Vannitsem, S., Demaeyer, J., De Cruz, L., & Ghil, M. (2015). Low-frequency vari-
801	ability and heat transport in a low-order nonlinear coupled ocean-atmosphere
802 803	model. Physica D: Nonlinear Phenomena, 309, 71–85.
	Vannitsem, S., & Duan, W. (2020). On the use of near-neutral Backward Lya-
804 805	punov Vectors to get reliable ensemble forecasts in coupled ocean–atmosphere
806	systems. Climate Dynamics, 55, 1125–1139.
807	Vannitsem, S., & Lucarini, V. (2016). Statistical and dynamical properties of covari-
808	ant lyapunov vectors in a coupled atmosphere-ocean model—multiscale effects,
809	geometric degeneracy, and error dynamics. Journal of Physics A: Mathematical
810	and Theoretical, $49(22)$ , 224001.
811	Vitart, F., Ardilouze, C., Bonet, A., Brookshaw, A., Chen, M., Codorean, C.,
812	Zhang, L. (2017). The Subseasonal to Seasonal (S2S) Prediction Project
813	Database. Bulletin of the American Meteorological Society, 98(1), 163 - 173.
814	doi: 10.1175/BAMS-D-16-0017.1
815	Von Storch, H., Bruns, T., Fischer-Bruns, I., & Hasselmann, K. (1988). Principal

816	oscillation pattern analysis of the 30-to 60-day oscillation in general circulation
817	model equatorial troposphere. Journal of Geophysical Research: Atmospheres,
818	93(D9), 11022-11036.
819	Wang, H., Lillo, S., & Newman, M. (2021). LIM Week 2-4 Forecast. Retrieved from
820	https://ftp.cpc.ncep.noaa.gov/hwang/LIM/
821	Wilks, D. (2011). Statistical methods in the atmospheric sciences. Elsevier.
822	Williams, M. O., Kevrekidis, I. G., & Rowley, C. W. (2015). A data-driven approx-
823	imation of the Koopman operator: Extending dynamic mode decomposition.
824	Journal of Nonlinear Science, 25(6), 1307–1346.
825	Xu, M., & Paul, M. R. (2016). Covariant Lyapunov vectors of chaotic Rayleigh-
826	Bénard convection. Physical Review $E$ , $93(6)$ , $062208$ .
827	Yang, Hl., Takeuchi, K. A., Ginelli, F., Chaté, H., & Radons, G. (2009). Hyper-
828	bolicity and the effective dimension of spatially extended dissipative systems.
829	Physical review letters, 102(7), 074102.
830	Yang, SC., Kalnay, E., Cai, M., & Rienecker, M. M. (2008). Bred Vectors
831	and Tropical Pacific Forecast Errors in the NASA Coupled General Cir-
832	culation Model. Monthly Weather Review, 136(4), 1305 - 1326. doi:
833	10.1175/2007MWR2118.1

# Supplementary Material : Additional Figures with the MSE, Spread and DSSS as a function of the forecast lead time, and the PFMD spectrum

Jonathan Demaeyer<sup>1</sup>, Stephen G. Penny<sup>2,3</sup> and Stéphane Vannitsem<sup>1</sup>

<sup>1</sup>Institut Royal Météorologique de Belgique, Avenue Circulaire, 3, 1180 Brussels, Belgium <sup>2</sup>Cooperative Institute for Research in Environmental Sciences, University of Colorado Boulder <sup>3</sup>NOAA Physical Sciences Laboratory, Boulder, Colorado

# Introduction

MSE, Spread and DSSS as a function of the forecast lead time: In this supplementary note, we show some figures depicting the time-evolution of the scores as a function of the forecast lead time. To recall first the definition of these scores, let's consider a dynamical system

$$\dot{\boldsymbol{x}} = \boldsymbol{f}(\boldsymbol{x}) \tag{1}$$

and a set of *n* ensemble forecasts  $\boldsymbol{y}_{m,n}(t)$ , m = 1, ..., M performed with it, *M* being the size of the ensembles. If  $\boldsymbol{x}_n(t)$  is the "truth" corresponding to the *n*<sup>th</sup> forecast, then the MSE and the Spread of the forecasts are defined as

$$MSE(\tau) = \frac{1}{N} \sum_{n=1}^{N} \|\boldsymbol{x}_{n}(\tau) - \bar{\boldsymbol{y}}_{n}(\tau)\|^{2}$$
$$Spread^{2}(\tau) = \frac{1}{N} \sum_{n=1}^{N} \frac{1}{M-1} \sum_{m=1}^{M} \|\boldsymbol{y}_{m,n}(\tau) - \bar{\boldsymbol{y}}_{n}(\tau)\|^{2}$$

where

$$\bar{\boldsymbol{y}}_n(\tau) = \frac{1}{M} \sum_{m=1}^M \, \boldsymbol{y}_{m,n}(\tau)$$

is the ensemble mean over the members  $\boldsymbol{y}_{m,n}(\tau)$  of the  $n^{\text{th}}$  ensemble forecast. f the Spread<sup>2</sup> and the MSE are close to one another, indicating that the estimated error is close to the true error, then the ensemble forecast is considered reliable [Leutbecher and Palmer, 2008].

The bias-free univariate DSS for the  $n^{\text{th}}$  ensemble forecast and the  $i^{\text{th}}$  variable of the system can be written as [Siegert et al., 2019]:

$$DSS_{n,i}(\tau) = \frac{1}{2} \log(2\pi) + \frac{1}{2} \log \sigma_{n,i}^2(\tau) + \frac{1}{2} \frac{M-3}{M-1} \left( \bar{y}_{n,i}(\tau) - x_{n,i}(\tau) \right)^2 / \sigma_{n,i}^2(\tau),$$

where  $\sigma_{n,i}^2$  is an estimator of the *i*<sup>th</sup> variable ensemble variance:

$$\sigma_{n,i}^2(\tau) = \frac{1}{M-1} \sum_{m=1}^M |y_{m,n,i}(\tau) - \bar{y}_{n,i}(\tau)|^2.$$

This score can then be averaged over the N realizations:

$$DSS_i(\tau) = \frac{1}{N} \sum_{n=1}^{N} DSS_{n,i}(\tau).$$

The lower the DSS score, the more reliable the ensemble forecasts are for this particular variable. In the context of the MAOOAM-VDDG ocean-atmosphere model considered in the paper, the Dawid-Sebastiani Score (DSS) can be aggregated per component of the system:

$$DSS_{\psi_{a}}(\tau) = \sum_{i=1}^{n_{a}} DSS_{\psi_{a,i}}(\tau)$$
$$DSS_{\theta_{a}}(\tau) = \sum_{i=1}^{n_{a}} DSS_{\theta_{a,i}}(\tau)$$
$$DSS_{\psi_{o}}(\tau) = \sum_{i=1}^{n_{o}} DSS_{\psi_{o,i}}(\tau)$$
$$DSS_{\theta_{o}}(\tau) = \sum_{i=1}^{n_{o}} DSS_{\theta_{o,i}}(\tau).$$

where  $\psi_a$  and  $\theta_a$  are respectively the streamfunction and temperature of the atmosphere, while  $\psi_o$  and  $\theta_o$  are respectively the streamfunction and temperature of the ocean.

Finally, considering several methods to obtain the ensemble forecasts, these aggregated score can be compared to *perfect* ensemble forecasts with the Dawid-Sebastiani Skill Score (DSSS) that we defined as:

$$DSSS_{\psi_{a}}^{\text{method}}(\tau) = 1 - \frac{DSS_{\psi_{a}}^{\text{method}}(\tau)}{DSS_{\psi_{a}}^{\text{perfect}}(\tau)}$$
$$DSSS_{\theta_{a}}^{\text{method}}(\tau) = 1 - \frac{DSS_{\theta_{a}}^{\text{method}}(\tau)}{DSS_{\theta_{a}}^{\text{perfect}}(\tau)}$$
$$DSSS_{\psi_{o}}^{\text{method}}(\tau) = 1 - \frac{DSS_{\psi_{o}}^{\text{method}}(\tau)}{DSS_{\psi_{o}}^{\text{perfect}}(\tau)}$$
$$DSSS_{\theta_{o}}^{\text{method}}(\tau) = 1 - \frac{DSS_{\psi_{o}}^{\text{method}}(\tau)}{DSS_{\theta_{o}}^{\text{perfect}}(\tau)}$$

The smaller the DSSS, the better. A value of zero indicates that the considered method matches the perfect ensemble reliability. On the other, a negative value of the DSSS would indicate that the method outperforms the perfect one.

We consider in this supplementary the two different model configurations mentioned in the paper, i.e. one with a weak low-frequency variability (LFV), and one with a strong LFV. In the latter case, we distinguish between two different regions of the attractor: a chaotic region for  $\theta_{0,2} > 0.12$  and a more "quiet" region for  $\theta_{0,2} < 0.08$ .

**PFMD spectra:** We also plot the PFMD<sup>1</sup> spectra, to show that they are the same as the one obtained with DMD and depicted in the paper.

More precisely, considering two collections of states of the dynamical system (1)  $X = [x_0 \dots x_{K-1}]$  and  $Y = [x_1 \dots x_K]$ , the PFMD representation of the Perron-Frobenius operator is given by

$$\mathsf{M}^{\mathrm{PFMD}} = \mathsf{A}^{\mathsf{T}}(\mathsf{G}^{+})^{\mathsf{T}}.$$
(2)

where  $A = Y X^*$  and  $G = X X^*$ . The eigenvalues of the matrix  $M^{\text{PFMD}}$  form then the above-mentioned spectrum.

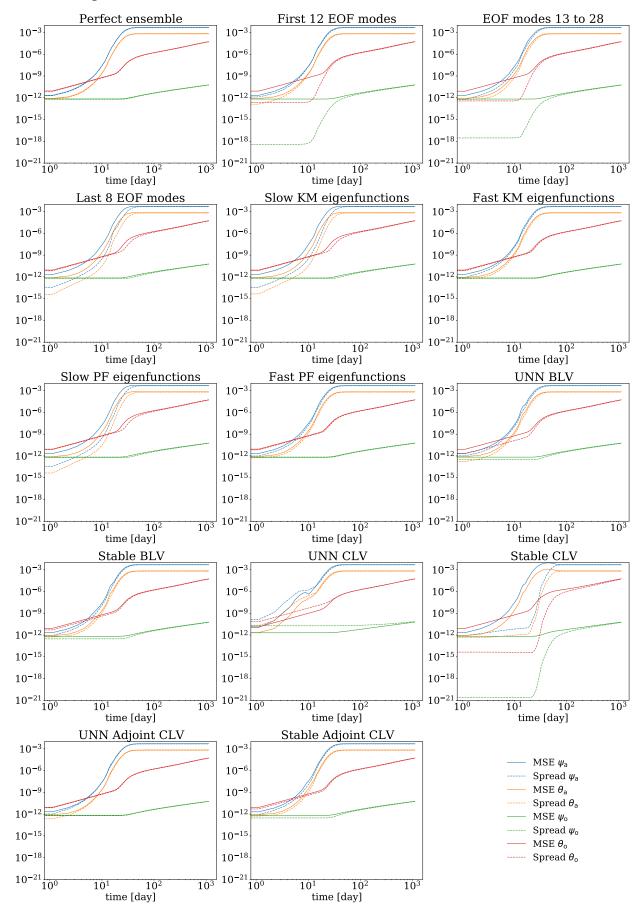
<sup>&</sup>lt;sup>1</sup>PFMD for Perron-Frobenius Modes Decomposition.

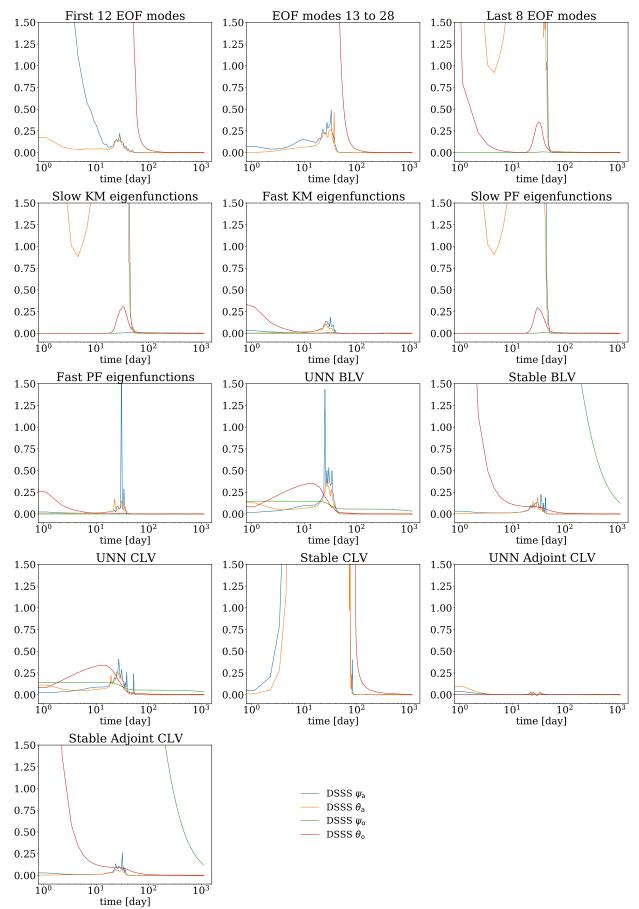
# Contents

1	Experiment the weak LFV         1.1       MSE and Spread as a function of the forecast lead time         1.2       DSSS as a function of the forecast lead time	
<b>2</b>	Experiment the strong LFV	7
	2.1 Case where $\theta_{0,2} > 0.12$	7
	2.1.1 MSE and Spread as a function of the forecast lead time	
	2.1.2 DSSS as a function of the forecast lead time	
	2.2 Case where $\theta_{o,2} < 0.08$	9
	2.2.1 MSE and Spread as a function of the forecast lead time	
	$2.2.2$ DSSS as a function of the forecast lead time $\ldots \ldots \ldots$	
3	PFMD spectra	
	3.1 Experiment without LFV	11
	3.2 Experiment with LFV	

# 1 Experiment the weak LFV

# 1.1 MSE and Spread as a function of the forecast lead time

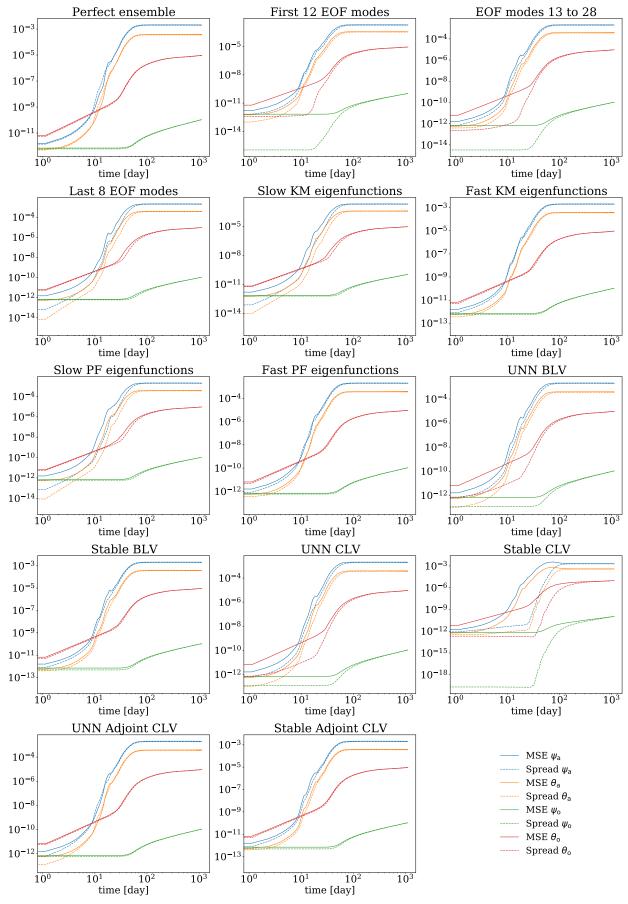


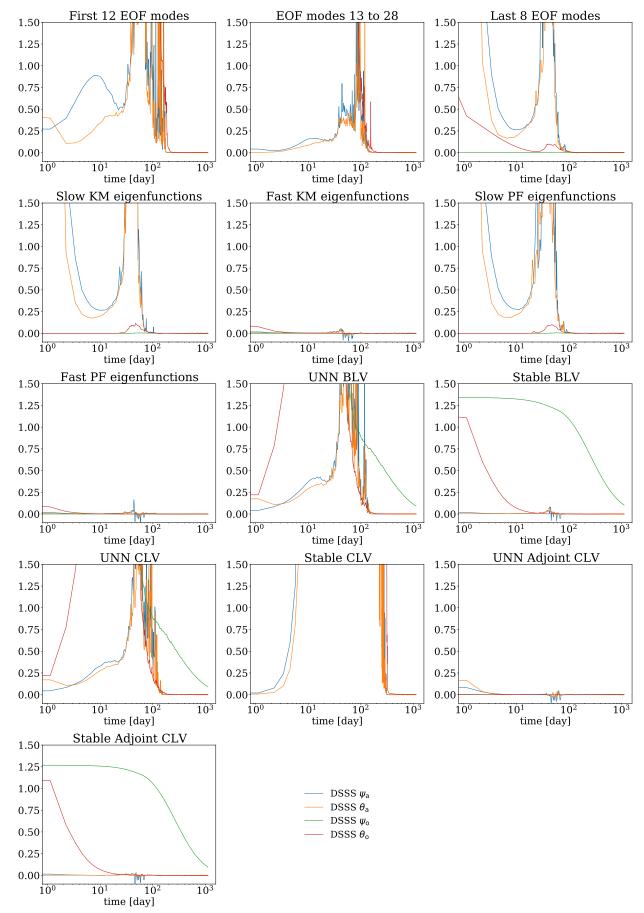


# 2 Experiment the strong LFV

**2.1** Case where  $\theta_{0,2} > 0.12$ 

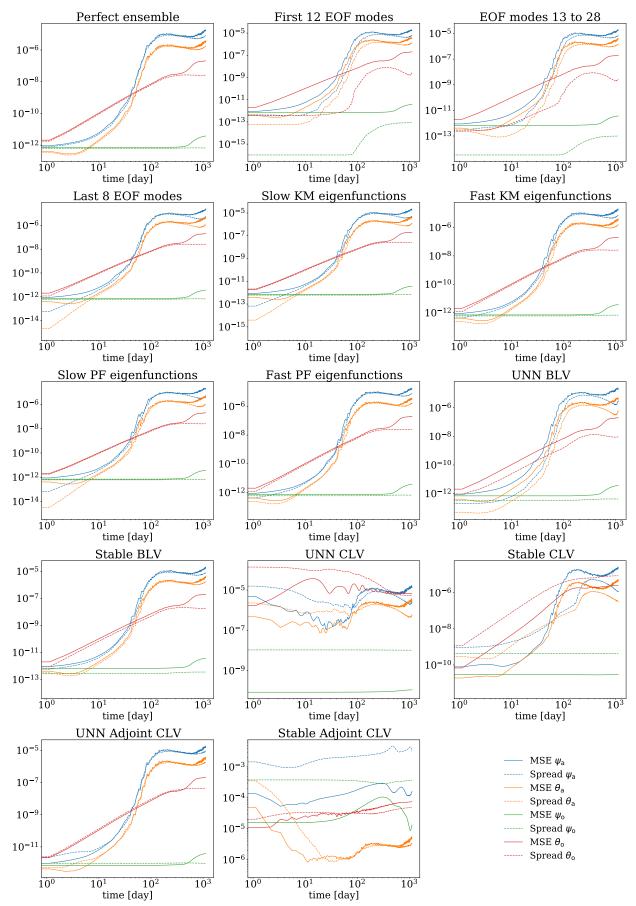
# 2.1.1 MSE and Spread as a function of the forecast lead time

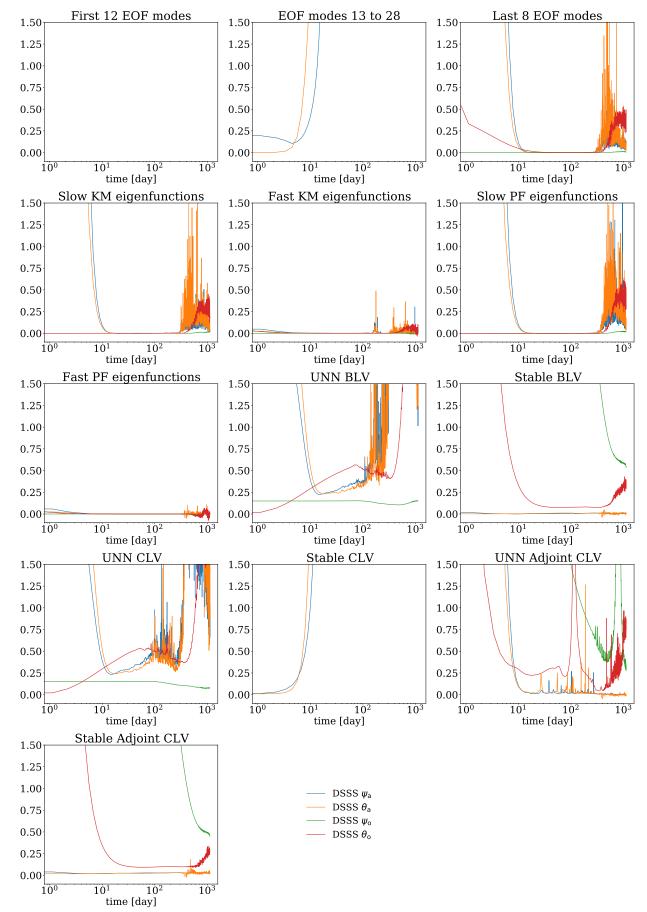




# 2.2 Case where $\theta_{o,2} < 0.08$

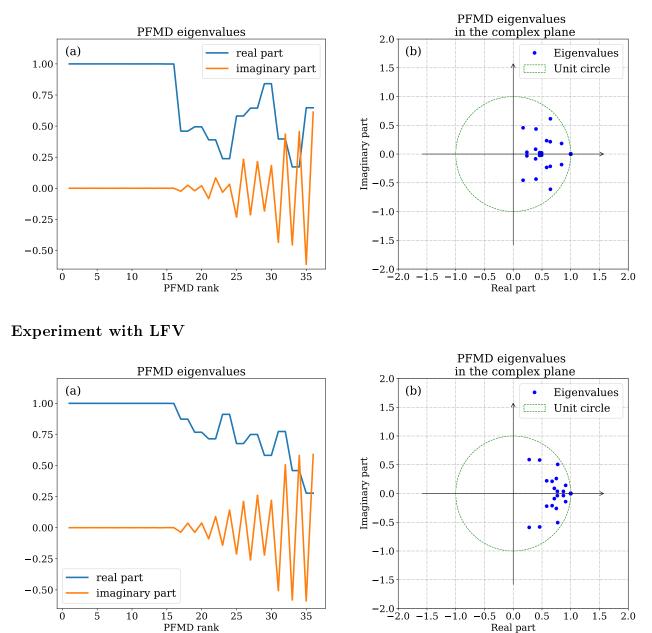
# 2.2.1 MSE and Spread as a function of the forecast lead time





# 3 PFMD spectra

# 3.1 Experiment without LFV



# References

3.2

[Leutbecher and Palmer, 2008] Leutbecher, M. and Palmer, T. N. (2008). Ensemble forecasting. Journal of computational physics, 227(7):3515-3539.

[Siegert et al., 2019] Siegert, S., Ferro, C. A., Stephenson, D. B., and Leutbecher, M. (2019). The ensemble-adjusted Ignorance Score for forecasts issued as normal distributions. *Quarterly Journal of the Royal Meteorological* Society, 145:129–139.