Explaining the Variability in High-Frequency Nitrate Export Patterns Using Long-Term Hydrological Event Classification

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Abstract

Runoff events play an important role for nitrate export from catchments, but the variability of nutrient export patterns between events and catchments is high and the dominant drivers remain difficult to disentangle. Here, we rigorously asses if detailed knowledge on runoff event characteristics can help to explain this variability. To this end, we conducted a long-term (1955 - 2018) event classification using hydro-meteorological data, including soil moisture, snowmelt and the temporal organization of rainfall, in six neighboring mesoscale catchments with contrasting land use types. We related these event characteristics to nitrate export patterns from high-frequency nitrate concentration monitoring (2013 - 2017) using concentration-discharge relationships. Our results show that small rainfall-induced events with dry antecedent conditions exported lowest nitrate concentrations and loads but exhibited highly variable concentration-discharge relationships. We explain this by a low fraction of active flow paths, revealing the spatial heterogeneity of nitrate sources within the catchments and by an increased impact of biogeochemical retention processes. In contrast, large rainfall or snowmelt-induced events exported highest nitrate concentrations and loads and converged to similar chemostatic export patterns across all catchments, without exhibiting source limitation. We explain these homogenous export patterns by high catchment wetness that activated a high number of flow paths. Long-term hydrometeorological data indicated an increase of events with dry antecedent conditions in summer and decreased snow-influenced events. These trends will likely continue and lead to an increased nitrate concentration variability during low-flow seasons and to changes in the timing of largest nitrate export peaks during high-flow seasons.

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19	Key Points:
20 21	• Hydrological event classification enables transferability of dominant drivers of event- scale nitrate export patterns across catchments
22 23	• Small events with low antecedent wetness exported low nitrate concentrations with highly variable CQ relationships
24 25 26	• Large events with high antecedent wetness exported high nitrate concentrations and loads with chemostatic patterns across all catchments

27 Abstract

28 Runoff events play an important role for nitrate export from catchments, but the 29 variability of nutrient export patterns between events and catchments is high and the dominant 30 drivers remain difficult to disentangle. Here, we rigorously asses if detailed knowledge on runoff event characteristics can help to explain this variability. To this end, we conducted a long-term 31 32 (1955 - 2018) event classification using hydro-meteorological data, including soil moisture, 33 snowmelt and the temporal organization of rainfall, in six neighboring mesoscale catchments 34 with contrasting land use types. We related these event characteristics to nitrate export patterns 35 from high-frequency nitrate concentration monitoring (2013 - 2017) using concentration-36 discharge relationships. Our results show that small rainfall-induced events with dry antecedent 37 conditions exported lowest nitrate concentrations and loads but exhibited highly variable 38 concentration-discharge relationships. We explain this by a low fraction of active flow paths, 39 revealing the spatial heterogeneity of nitrate sources within the catchments and by an increased 40 impact of biogeochemical retention processes. In contrast, large rainfall or snowmelt-induced 41 events exported highest nitrate concentrations and loads and converged to similar chemostatic 42 export patterns across all catchments, without exhibiting source limitation. We explain these 43 homogenous export patterns by high catchment wetness that activated a high number of flow 44 paths. Long-term hydro-meteorological data indicated an increase of events with dry antecedent 45 conditions in summer and decreased snow-influenced events. These trends will likely continue and lead to an increased nitrate concentration variability during low-flow seasons and to changes 46 47 in the timing of largest nitrate export peaks during high-flow seasons.

48

Plain Language Summary

49 Runoff events play an important role for nitrate exports from catchments. However, the 50 response of nitrate export to runoff events is highly variable and therefore difficult to understand. 51 Here, we classified runoff events according to their inducing precipitation and antecedent soil 52 moisture and related those event characteristics to nitrate export patterns. Our results show that 53 small summer and autumn events exported lowest nitrate concentrations and loads with highly 54 variable patterns, such as increasing or decreasing nitrate concentrations. We explain this 55 variability with nitrate mobilization being restricted to near-stream areas with variable nitrate 56 availability and by an increased impact of biogeochemically controlled nitrate retention. In 57 contrast, larger winter and spring events exported highest nitrate concentrations and loads. These 58 events showed only a small increase of nitrate concentrations compared to discharge, so that 59 discharge dominated overall nitrate loads. This was similar in all catchments, which we explain 60 by high catchment wetness connecting all nitrate sources within a catchment to the stream. Long-61 term trends indicate a decrease of soil moisture in summer and a decrease of snow-influenced 62 events. These trends might cause an increasing variability in nitrate concentrations during summer and change the timing of large nitrate export peaks during winter and spring. 63

64 **1 Introduction**

High riverine nitrate concentrations and loads from diffuse agricultural sources threaten
drinking water quality and the health of freshwater as well as marine ecosystems (Carpenter et
al., 1998; Elser, 2011; Mekonnen & Hoekstra, 2020). In this context, runoff events play a
dominant role for the mobilization and transport of nitrate from catchments to the downstream
receiving water resources (Blaen et al., 2017; Inamdar et al., 2006; Ockenden et al., 2016).
Climate change is predicted to change the frequency and characteristics of such runoff events,

and these changes are in turn predicted to significantly alter water quality and nutrient export

72 (IPCC, 2018; Marshall & Randhir, 2008; Sebestyen et al., 2009; Trang et al., 2017; Wagena et

al., 2018). Therefore, an in-depth understanding of nitrate mobilization and transport during

74 runoff events under different hydro-meteorological conditions is needed to better predict and

75 mitigate water quality deteriorations.

76 Hydro-meteorological data on a high temporal resolution (i.e., daily) has been readily 77 available for many decades and allows for a robust characterization of catchment hydrologic 78 functioning during runoff events on the long term (Tarasova et al., 2020; Kirchner et al. 2004). 79 Those analyses showed that with different antecedent wetness conditions, different flow paths 80 within a catchment can become activated that connect different catchment areas with the stream 81 network (Jencso et al., 2009). For dry antecedent conditions, typically only a smaller fraction of 82 the catchment area is connected to the stream network, often via deeper subsurface flow paths, 83 which deliver older water with longer transit times. In contrast, during wet antecedent conditions, 84 additional shallower and faster flow paths become activated and transport younger water (i.e., 85 with shorter transit times) also from more distant locations to the stream (Jencso et al., 2009; 86 Kumar et al., 2020; J. Yang et al., 2018). Moreover, in a temperate climate, runoff events can be 87 generated by precipitation events of different nature, such as pure rainfall or snowmelt (Tarasova 88 et al., 2020). In such climates, rain-on-snow events (i.e., snowmelt in concurrence with rainfall 89 and high antecedent soil moisture) often form the largest runoff events of the year and can 90 activate all or most of the available flow paths within a catchment (Berghuijs et al., 2019; Jencso 91 et al., 2009; Stieglitz et al., 2003).

92 It is most likely that the spatiotemporal variability in the hydrological land-to-stream 93 connectivity causes different responses in nutrient mobilization and transport as well (Stieglitz et 94 al., 2003). With the advent of high frequency measurements for nitrate and other nutrient 95 concentrations, we can know measure water quality at the same temporal resolution as water 96 quantity to analyze in detail the impact of runoff event characteristics on nitrate export patterns 97 (Kirchner et al., 2004; Rode et al., 2016a). Here we refer to runoff event characteristics as all 98 related hydro-meteorological characteristics including antecedent conditions, the characteristics 99 of the inducing precipitation event and the characteristics of the runoff event as such (for 100 example peak discharge). To maintain consistency in the use of terms with previous studies (i.e., 101 Musolff et al., 2015, 2021; Tarasova et al., 2020), we use the terms "discharge" and "runoff" 102 synonymously, referring to the total volumetric water flow rate in the stream at a gauging point.

103 Several studies took advantage of the high frequency measurements and conducted 104 detailed analysis on nutrient mobilization and transport during runoff events and confirmed the 105 important role of runoff events for nutrient export (e.g., Blaen et al., 2017; Burns et al., 2019; 106 Rose et al., 2018, Fovet et al., 2018, Knapp et al., 2020). For example, Casson et al., (2010) and 107 Pellerin et al. (2012) showed that large rain-on-snow events can account for a disproportional 108 amount of annually exported large nitrate loads. These studies also revealed large inter-event and 109 inter-catchment variability of nutrient export dynamics. For example, Blaen et al. (2017) found a 110 positive correlation between antecedent wetness and event nitrate concentrations in a catchment 111 with mixed agricultural and forested land use. On the contrary, Knapp et al. (2020) found a 112 negative correlation between antecedent wetness and event nitrate concentrations in a small 113 mountainous catchment that is covered by forest and meadows. These examples show that nitrate 114 export patterns can be related to event characteristics, such as the contribution of meltwater or

antecedent conditions, whereas this relationship can be highly variable between catchments of different configurations e.g., with regrads to land use and nitrate availability.

117 A common tool to reveal the relevant sources and flow paths for nitrate transport under 118 changing hydrological conditions are concentration-discharge (CQ) relationships, which 119 represent the directional relationship between concentrations and (e.g., Bowes et al., 2015; 120 Musolff et al., 2021; Vaughan et al., 2017). A negative slope of the CQ relationship can indicate 121 high base flow concentrations that are diluted by newly activated discharge zones (Bowes et al., 122 2015) or a depletion of nutrient sources (Vaughan et al., 2017). A positive CQ slope can indicate 123 the additional activation of more shallow and younger (Musolff et al., 2015) or more distant 124 nutrient source zones (Bowes et al., 2015). A chemostatic pattern is instead described by a CQ 125 slope close to zero (Godsey et al., 2009; Musolff et al., 2015; Thompson et al., 2011) and is 126 mainly attributed to ubiquitous and uniformly distributed N sources in agricultural catchments 127 (Basu et al., 2010). The CQ slope is not necessarily consistent across time scales and can thus 128 reveal complementary information on nutrient export during single runoff events compared to 129 CQ relationships across seasons that integrate several events (e.g., Godsey et al., 2019; Knapp et 130 al., 2020; Minaudo et al., 2019; Musolff et al., 2021). Nevertheless, a rigorous assessment of 131 how much of the inter-event variability of nutrient export patterns can be explained by a more 132 thorough understanding of runoff event characteristics is still missing. Furthermore, studies that 133 relate hydrological runoff events with nutrient transport are generally limited to relatively short 134 time periods, which is often not more than one or two years (e.g., Bieroza et al., 2018; Bowes et 135 al., 2015; Carey et al., 2014). Therefore, it mostly remains unclear if analyzed event 136 characteristics are representative for the long-term catchment behaviour and if runoff event 137 characteristics might change over longer time scales.

138 In this study, we explore to what extend runoff event characteristics govern nitrate export 139 during and across runoff events in different catchments. To this end, we related event 140 characteristics such as the wetness state of a catchment, the nature of an inducing precipitation 141 event and the temporal organization of rainfall to nitrate concentrations and loads. For this we 142 used a 5-year period of high frequency water quality and hydro-meteorological data from six 143 mesoscale Central European catchments with different land use settings. We classified runoff 144 events according to their different hydro-meteorological conditions (Tarasova et al., 2020) and 145 utilized CQ relationships to infer the relevant flow paths and source areas for nitrate transport 146 and mobilization. We then combined the findings from such analysis with the changes in event 147 characteristics and catchment state conditions over past decades obtained from long-term daily 148 hydro-meteorological time-series to identify possible trends in the long-term runoff event 149 characteristics that could impact nitrate export dynamics in the future.

- 150 **2 Materials and Methods**
- 151 2.1 Study area

Event characteristics and nitrate export patterns were analyzed in six sub-catchments of the Bode River catchment (Figure 1), which is an intensively monitored catchment within the network of the TERestrial Environmental Observatories (TERENO, Wollschläger et al., 2017). Warme Bode (WB), Rappbode (RB) and Hassel (HS) are part of the Rappbode Reservoir Observatory (Rinke et al., 2013), whereas Silberhütte (SH), Meisdorf (MD) and Hausneindorf (HD) are three subsequent gauging stations of the nested Selke River catchment. All six

- 158 catchments are located in the Harz Mountains and the Harz foreland in Saxony-Anhalt, Germany
- 159 (Figure 1). They have contrasting characteristics in regard to their size, land use, elevation and
- 160 mean annual precipitation (Table 1).



162 **Figure 1**. Map of the study site, showing all six mesoscale catchments (WB, RB, HS, SH, MD

- 163 and HD) with their respective their land use.
- 164 **Table 1.** Characteristics of the six studied mesoscale catchments within the Bode River165 catchment.

Catchmont	Area	Mean annual precipitation	Mean annual temperature	Land use and land cover				Elevation	Mean
Catenniem				Agriculture	Forest	Urban	Other	range	Slope
	[km²]	[mm yr ⁻¹]	[°C]	[%]	[%]	[%]	[%]	[m.a.s.l.]	[%]
WB	101.1	1111.9	6.6	5.9	90.2	2.9	1.0	429 - 957	7.7
RB	39.1	969.3	7.1	19.5	74.7	4.1	1.7	454 - 636	6.8
HS	42.0	820.9	7.0	59.8	35.6	4.5	0.1	436 - 604	4.8
SH	102.5	726.6	6.7	34.6	62.2	3.2	0.0	335 - 597	6.9
MD	178.6	693.0	7.2	23.2	73.6	3.1	0.1	196 - 597	8.4
HD	460.1	589.4	8.1	54.8	36.7	6.1	2.4	68 - 597	4.9

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- 167 2.2 Data
- 168 2.2.1 Long-term daily data

Long-term daily data (Figure S1–S6) were used to classify runoff events and to analyse potential trends in these. Daily discharge data was provided by the State Office of Flood

170 potential fields in filese. Daily discharge data was provided by the state office of Floor 171 Protection and Water Management of Saxony-Anhalt (LHW) and calculated to specific

- 172 discharge [mm d⁻¹]. In all catchments except HS and HD, discharge data is available from 1955
- 173 until 2018. In HS and HD, discharge data records started in 1968 and 1980, respectively, and

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174 lasted until 2018. Daily precipitation data over these time periods were provided by Germany's

175 National Meteorological Service (Deutscher Wetterdienst, DWD) as interpolated station data at a

spatial resolution of 1 km² (REGNIE; Rauthe et al., 2013). Daily average temperatures were

- interpolated to a 4 km grid from the DWD stations by External Drift Kriging using elevation as
- an explanatory variable (Zink et al., 2017). Daily soil moisture and snow water equivalent were according the measure level $M_{\rm eff}$ is the measure of $M_{\rm eff}$ is the measure of M_{\rm eff} is the measure of $M_{\rm eff}$ is the measure of M
- 179 calculated using the mesoscale Hydrological Model (mHM, Kumar et al., 2013; Samaniego et al. 2010; Zink et al. 2017)
- 180 al., 2010; Zink et al., 2017).
- 1812.2.2 High-frequency hourly data

182 High-frequency hourly data were used to analayse exported nitrate loads and concentration-discharge relationships within and between runoff events from 2013 to 2017. 183 184 Discharge data at a temporal resolution of 15 minutes were provided by the LHW, which we 185 aggregated to hourly values [mm h⁻¹]. Similar to nitrate concentration data (see below), a moving 186 average was applied over a 5-hour window to smooth the data. Hourly precipitation data as 187 reprocessed radar data were provided by the DWD with precipitation amounts adjusted to station 188 observations and a spatial resolution of 1 km² (RADOLAN; Winterrath et al., 2017). Due to a 189 lack of hourly temperature data, we reconstructed those from the daily data by using hourly 190 weights based on month-specific sine functions obtained from long-term minimum and 191 maximum temperatures to resemble the diurnal cycle of temperature. Hourly snow water 192 equivalent and soil moisture data was simulated using mHM (Zink et al., 2017).

193 Nitrate concentration data between 2013 and 2017 was collected via TRIOS ProPS-UV 194 sensors at 15 min. intervals (Kong et al., 2019; Rode, et al., 2016a), which we aggregated to 195 hourly averages. Data from the WB catchment were previously published by Kong et al. (2019) 196 and Musolff et al. (2021), data from the three Selke catchments (SH, MD and HD) were 197 previously published by Rode et al. (2016b), X. Yang et al. (2018), Winter et al. (2021) and 198 Musolff et al. (2021). For the processing of the nitrate concentration data we refer to the 199 references above and to our supporting information (Text S1). Briefly, raw data was restricted to a realistic range $(0 - 100 \text{ mg L}^{-1})$, outliers were removed, a moving average was applied over a 200 201 5-hour window to smooth the data and concentrations were calibrated against grab samples analyzed in the lab (R² 0.80 - 0.91, Figute S7 and S8). Note that the LHW gauging station at HS, 202 203 where long-term and high-frequency discharge data were measured, is located upstream of the 204 measurement point for concentration data, thus delineating a catchment size of around 29 km² for 205 discharge data compared to 42 km² for measured concentration data (Figure 1, Table 1). 206 Nevertheless, area-specific discharge data [mm h⁻¹] from upstream and downstream 207 measurement points (available downstream between 2013 and 2014) showed a good agreement 208 in their temporal dynamics with a R² of 0.88 and in absolute values with a small percentage bias 209 of -3.0% (Figure S9). We thus found area-specific discharge from the upstream station data to be 210 suitable for further analysis at the downstream station.

211

2.3 Runoff event identification and classification

Runoff events were separated and classified using the recently developed approach by Tarasova et al. (2018, 2020), which allows for an automated separation and classification of runoff events. The approach is explained in detail in the cited studies and is, therefore, only briefly described here. As a first step, events from daily long-term and high-frequency data were identified using an automated event separation approach from Tarasova et al. (2018). Then, we adopted the event classification framework from Tarasova et al. (2020), developed for daily data

- resolution (Figure 2). Each runoff event was classified by the characteristics of the inducing
- precipitation event (Figure 2a, Layer 1) and the pre-event wetness state of the catchment (Figure
 2a, Layer 2). The nature of precipitation events was identified by the ratio of meltwater volume
- 220 2a, Layer 2). The nature of precipitation events was identified by the ratio of meltwater volume 221 (M_{vol}) and total precipitation volume (i.e., sum of rainfall and snowmelt, P_{vol}). Using a threshold
- of $M_{vol}/P_{vol} = 0.05$ (Figure 2a), events were classified as *Rain* or *Snow*-influenced events. The
- temporal organization of precipitation was characterized by means of the temporal coefficient of
- 224 variation of the precipitation rate (CV_{temp}) and by the ratio between the maximum precipitation
- rate during an event and precipitation volume (P_{max}/P_{vol}). Events with a $CV_{temp} > 1$ and
- 226 $P_{max}/P_{vol} > 0.5$ were defined as *intensity-dominated* and all other events as *volume-dominated*
- (Figure 2a). Third, the wetness state of a catchment was characterized by means of antecedent
- soil moisture (SM_{ant}). Using a threshold of maximum κ , with κ representing the catchmentspecific curvature of the non-linear relationship between event runoff coefficients and soil
- moisture (Tarasova et al., 2020), events were classified as *Wet* od *Dry* events (Figure 2a). In
- total, this classification resulted in five event classes (Figure 2b): i) snow-influenced events
- 232 (*Snow*), ii) rain-induced events that were *volume dominated* and occurred under wet antecedent
- soil moisture conditions (*Rain-Wet-Vol*), iii) rain-induced events that occurred under wet
- antecedent conditions and were *intensity dominated* (*Rain-Wet-Int*), iv) rain-induced events that
- 235 occurred under dry antecedent conditions and were *volume dominated* (*Rain-Dry-Vol*) and v)
- rain-induced events that occurred under dry antecedent conditions and were *intensity dominated*
- 237 (*Rain-Dry-Int*). To assure comparability of event classes between two datasets of different
- resolution, we classified events using the daily time series (1955 2018) and then assigned the
- respective classes to the corresponding events from the hourly time series (2013 2017).



240

Figure 2. a) Event characteristics and thresholds for the classification of events. Threshold for the wetness state of the catchments is defined by the maximum of κ , which represents the catchment-specific curvature of the nonlinear relationship between event runoff coefficients and soil moisture. b) The resulting event classes. Modified from Tarasova et al. (2020, CC BY 4.0).

245 2.4 Long-term trends in event characteristics

We used the non-parametric Mann-Kendall test (Kendall, 1998; Mann, 1945) to detect monotonic trends in the continuous event characteristics and event classes with a significance level of 5%. We considered the following continuous event characteristics: i) antecedent soil moisture (SM_{ant}), ii) the ratio of meltwater volume and precipitation volume (M_{vol}/P_{vol}) and iii) 250 the ratio of maximum precipitation and precipitation volume (P_{max}/P_{vol}), which is an indicator for

intensity- or *volume-dominated* events, respectively (Figure 2a). To reduce inter-event variability

- between those characteristics, we calculated seasonal averages for each year and analyzed these
- for seasonal long-term trends. Similarly, we analyzed seasonal trends in the annual contribution and total number of i) *Snow* events (vs. *Rain* events), ii) *Rain-Dry* (vs. *Rain-Wet*) events and iii)
- 255 *Intensity-dominated* (vs. *Volume-dominated*) events.
- 256 2.5 Nitrate export
- 257 2.5.1 Descriptors of nitrate export

To characterize nitrate transport from the high-frequency data, we chose four descriptors for the event scale: i) median nitrate concentration (C_{med}), ii) average loads per event in kg ha⁻¹ yr⁻¹ (this unit was chosen for a better comparison between catchments and events of different duration), iii) inter-event CQ slopes and iv) event-specific CQ slopes. Event-specific CQ slopes were assessed by fitting the parameter *b* from the following power-law relationship after Godsey et al. (2009) and Musolff et al. (2015) to the data of the individual events:

264 C(

$$C(t) = aQ(t)^b \tag{1}$$

where C(t) represents the time series of nitrate concentrations during a specific event in 265 266 mg L⁻¹, Q(t) represents the time series of discharge in mm h⁻¹, and a and b represent the intercept and linear slope of the CQ relationship in the log-log space. A parameter b < 0 describes a 267 268 negative CO slope, i.e., decreasing concentrations with increasing discharge and therefore a 269 dilution pattern. A parameter b > 0 describes a negative CQ slope, i.e., increasing concentrations 270 with increasing discharge and therefore an accretion pattern. Both scenarios are accounted for as 271 chemodynamic patterns (Godsey et al., 2009; Musolff et al., 2015, 2017). If parameter b is close 272 to zero, there is no clear directional relationship. This pattern can be described as chemostatic 273 under the assumption that the coefficient of variation of concentrations is much smaller than that 274 of discharge (Godsey et al., 2009; Musolff et al., 2015, 2017). Similar to the event-specific CQ 275 slope, we analyzed the CQ relationship across all events within each catchment (i.e., the inter-276 event CQ slope) using the power law model from eq. 1 with C_{med} and Q_{med} of each event instead 277 of C(t) and Q(t) within each specific event.

278 2.5.2 Statist

2.5.2 Statistical analysis

279 All computations and statistical analyses were conducted in R (R Core Team, 2020). We 280 used the non-parametric Kruskal-Wallis test (Kruskal & Wallis, 1952) to test for differences in 281 loads, C_{med}, Q_{med} and the CQ slope between event classes and the Pairwise Wilcoxon Test 282 (Wilcoxon, 1945) with Holms correction for multiple comparisons (Holm, 1979) to test for 283 differences in-between the event classes, both at the significance level of 5%. In order to test the 284 impact of event classes on the inter-event CQ slope, we tested the simple linear $\ln(C_{med}) - \ln(Q_{med})$ 285 regression against a linear regression model that includes event classes and their interactions 286 with $ln(Q_{med})$. Both models were compared via the sample-size corrected Akaike Information 287 Criterion (AICc; Akaike, 1973; Hurvich & Tsai, 1989; Sugiura, 1978). If accounting for event 288 classes led to a substantial improvement (i.e., AICc decreased at least by 2, similar to Marinos et 289 al., 2020) their impact was regarded as considerable. Otherwise the added value from event 290 classes compared to a simple CQ model was negligible for nitrate export estimations.

3 Results

292

3.1 Long-term and high frequency runoff event characteristics

293 In total we identified and classified 5872 events over the long-term period (on average 294 14.5 - 19.0 events per catchment and year) and 388 events over the high-frequency time period 295 (on average 9.6 - 16.2 events per catchment and year). Event classes generally differed more 296 strongly between seasons than between catchments (Figure 3). In both long-term and high-297 frequency event classes, winter was dominated by Snow and Rain-Wet-Vol events. Spring 298 showed the greatest variability of event classes and was the season with the highest percentage of 299 Rain-Wet-Vol and Rain-Wet-Int events. Summer and autumn were dominated by Rain-Dry-Vol 300 events and *Rain-Dry-Int* events. Differences between catchments reflect a decreasing percentage 301 of Snow events and an increasing percentage of Rain-Dry-Vol and Rain-Dry-Int events from west 302 to east (WB to HD catchment; Figure 3). Across all seasons of the long-term event classes, more 303 than half of all events in the six catchments were classified as Rain-Dry events (from 51.2% in 304 WB to 61.9% in HD). Around a fifth up to a quarter of all observed events were classified as 305 Snow events (from 17.7% in MD to 25.0% in WB), and the proportion of Rain-Wet events ranged between 14.0% in HS and 23.9% in WB. Rain-induced events were more frequently 306 307 volume- than intensity-dominated.

308Runoff events had an average duration of 11.2 ± 9.1 days for long-term data and $11.2 \pm$ 3098.9 days during the period of high-frequency data, with considerable differences between event310classes (Figure S10). Intensity-dominated events (Rain-Dry-Int and Rain-Wet-Int) were shortest,311lasting in average 5.5 ± 4.7 days and 8.6 ± 6.5 days for long-term data, respectively, followed by312Rain-Dry-Vol events that lasted in average, 10.0 ± 8.1 days. Rain-Wet-Vol and Snow events were

313 the longest events, lasting in average 14.4 ± 8.8 and 18.2 ± 10.2 days.

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314 315 **Figure 3.** Seasonal absolute frequency of event classes (y axis) in the six study catchments,

316 showing a) long-term changes between 8-year periods from 1955 until 2018 and b) the period of

317 high-frequency data from 2013 until 2017. Periods in panel a) with no or very few events in HS

318 and HD indicate missing discharge data.

319

3.1.1 Long-term trends and changes in event characteristics

320 In agreement with increasing temperature due to climate change (IPCC, 2013), Mann 321 Kendall trends analysis indicated a decrease in the number and proportion of *Rain-Wet* events in 322 summer, which was significant in half of the catchments (WB, SH and MD; Table S1). This 323 decrease goes along with a significant decrease in antecedent soil moisture in spring and/or summer in WB, RB and MD catchments. Moreover, the number and/or proportion of snow 324 325 events decreased significantly in spring in WB, RB, SH and MD catchments (Table S1). These 326 trends go along with a significant decrease in the proportion of meltwater volume per event 327 (M_{vol}/P_{vol}) in winter and spring in all catchments except HS. Only one catchment (WB) showed a 328 significant decrease in the number and proportion of *intensity*- vs. *volume-dominated* events, 329 which occured during summer. In contrast, the RB, SH and MD catchments showed a significant 330 increase in P_{max}/P_{vol} during summer, but no significant trend in the total number or proportion of 331 intensity- or volume-dominated events (Table S1).

332

3.2 Nitrate export during runoff events

333 Runoff event and nitrate export characteristics differed between catchments. Event runoff 334 decreased roughly from west to east, along the precipitation gradient (Table 1) with highest average Q_{med} in the WB catchment (1.1 mm h⁻¹) and lowest average Q_{med} in the HD catchment 335 336 (0.3 mm h⁻¹). Nitrate export during runoff events varied across catchments in line with their land 337 use patterns (Table 1). Catchments with the highest percentage of agricultural land use and 338 lowest percentage of forests exported in average highest C_{med} (HS and HD with 2.5 mg L⁻¹ and 339 2.3 mg L^{-1}), followed by mixed land use catchments (MD and SH with 1.6 mg L^{-1} and 1.5 mg L^{-1} 340 ¹) and lowest average C_{med} was observed in the dominantly forested catchments RB and WB (0.6 341 mg L^{-1} and 0.7 mg L^{-1}).

342 3.2.1 Nitrate loads

343 Runoff events had a prominent role for annual nitrate export. The cumulative duration of 344 all identified events from the high-frequency data was on average 39.6% (30.8% - 48.1% 345 depending on the catchment) of the analyzed time period (2013 - 2017), while they accounted on 346 average for 51.2% (44.8% - 63.3%) of all nitrate loads (Text S2). In relation to catchment area 347 (Table. 1), the HS catchment transported highest median nitrate loads across all event classes (5.5 kg ha⁻¹ yr⁻¹) followed by HD (1.8 kg ha⁻¹ yr⁻¹), MD (1.7 kg ha⁻¹ yr⁻¹), WB (1.6 kg ha⁻¹ yr⁻¹), 348 349 SH (1.4 kg ha⁻¹ yr⁻¹) and lowest median loads were exported from RB catchment (1.0 kg ha⁻¹ yr⁻¹) ¹). Between event classes, lowest loads were transported during *Rain-Dry-Int* and *Rain-Dry-Vol* 350 351 events, which were responsible for around 25.6% (14.8% in MD up to 41.6% in HD) of all 352 event-driven loads. Highest loads were transported during Rain-Wet and Snow events, which 353 were responsible for around 74.4% (58.4% - 85.2%) of all event-driven loads (Figure 4). Kruskal 354 Wallis test showed significant differences in exported nitrate loads between the event classes in 355 all catchments. Results of the pairwise Wilcoxon Test indicated that these differences are mainly 356 driven by the differences between Rain-Dry-Int or Rain-Dry-Vol events and Rain-Wet-Vol or 357 Snow events, whereas no significant difference between Rain-Dry-Vol and Rain-Dry-Int events, 358 nor between Rain-Wet-Vol and Snow events were detected. Rain-Wet-Int events were generally 359 too low in their frequency (n = 1-7) to be compared reliably (Figure 4).



360

Figure 4. Nitrate loads (on logarithmic scale) transported during runoff events, divided into
 catchments and runoff event classes.

363

3.2.2 Inter-event concentration-discharge relatioships

364 The inter-event CQ relationship is characterized by the slope between $\ln(C_{med})$ and 365 ln(Q_{med}) of all events within one catchment. It shows consistently positive CQ relationships in 366 the log-log space, indicating that C_{med} increased with Q_{med} but with a different slope depending on the catchment (Figure 5). In line with transported loads (Figure 4), Rain-Dry-Vol and Rain-367 368 Dry-Int events are mainly located in the lower part of the CQ relationship, representing small 369 events (low Q_{med}) with low concentrations (low C_{med}) that occur mainly during summer and 370 autumn (Figure 3). Rain-Wet-Vol and Snow events that occurred mainly in winter and spring 371 (Figure 3) are located on the upper part of the CQ relationship, showing the highest Q_{med} and 372 C_{med} (Figure 5). Additionally, some *Rain-Dry-Vol* events are located at the upper end of the CO relationship. These events occurred mainly during autumn and often extended into the winter 373 374 period with higher Qmed and Cmed. Rain-Wet-Int events occurred only occasionally and represent 375 mainly smaller events in winter and spring with medium C_{med} and Q_{med}, plotting in between the 376 other event classes.

377 The inter-event CQ relationship could account for most of the variance in C_{med} with a R² 378 varying between 0.51 and 0.91 (Figure 5). Except for the SH catchment, adding information on 379 event classes to the model did not improve its performance in terms of AICc compared to a 380 simple CQ model. This indicates that Q_{med} was the most powerful predictor of C_{med} and no or 381 only a small part of additional variance was explained by the event classes themselves. In the SH 382 catchment, event classes clearly improved the linear regression model (AICc decreased by 13.6 383 units). While no clear differences between Rain-Dry-Vol and Rain-Dry-Int events nor between 384 *Rain-Wet-Vol* and *Snow* events are visible (Figure S11), the main event class differences in SH 385 was a higher intercept of Rain-Wet-Vol and Snow events compared to Rain-Dry-Int and Rain-386 Dry-Vol events.



Figure 5. Median discharge (Q_{med}) and nitrate concentrations (C_{med}) for each runoff event with log-scale x- and y-axis, separated into the six catchments (a-f). Colors indicate the five different event classes, grey lines show the linear relationship between $\ln(C_{med})$ and $\ln(Q_{med})$ and colored lines show individual linear relationships between $\ln(C_{med})$ and $\ln(Q_{med})$ for each event class, only shown when event classes clearly improved the linear regression model (differences in AICc > 2; d).

394

3.2.3 Event-specific concentration-discharge relationships

395 Across all catchments, most events (72.4%) were characterized by a positive event-396 specific CQ slope, indicating an increase of nitrate concentrations with increasing discharge (Figure 6). We found that event-specific CQ slopes in all catchments showed a large variability 397 398 between small events (low Q_{med}), whereas CQ slopes for larger events (higher Q_{med}) collapse to a 399 slightly positive CQ slope that is roughly between 0.1 and 0.3, close to a chemostatic pattern 400 (Figure 6a). Some catchment-specific differences can be observed between CQ slopes during 401 small events (Figure 6b). The more forested and pristine catchments dominantly showed positive 402 CQ slopes, whereas the agriculturally dominated catchments HS and HD tended towards close-403 to-zero or negative CQ slopes.





Figure 6. Event-specific CQ slopes (slope of nitrate concentrations and discharge in the log-log space) against the specific median discharge (Q_{med}) of each event shown for all catchments in one plot (a) and with logarithmic x-axis and separated by catchments to visualize differences in events with low Q_{med} between catchments (b). Colors of dots indicate the five event classes, dot sizes indicate the event load. Grey-shaded areas indicate event-specific CQ slopes close to zero (between -0.1 and 0.1).

411 Rain-Dry-Int and Rain-Dry-Vol events cause most of the variability between event-412 specific CQ slopes (Figure 6). These two event classes are distinguished by the temporal 413 organization of the inducing rainfall, being either *intensity*- or *volume-dominated*. To assess 414 whether the difference in temporal organization of rainfall events explains any additional 415 variability in event-specific CO slopes of small events, we compared both classes using the Kruskal-Wallis test (Kruskal & Wallis, 1952). We found significantly higher event-specific CQ 416 417 slopes for Rain-Dry-Int events compared to Rain-Dry-Vol events in half of the catchments 418 (Figure 7; WB, SH and MD), all of them showing >60% forest cover (Table 1). Moreover, the 419 median of event-specific CQ slopes from Rain-Dry-Int events was higher compared to Rain-Dry-420 Vol events in all six catchments. In contrast, no significant difference in C_{med} and Q_{med} between 421 Rain-Dry-Int and Rain-Dry-Vol events could be detected (Figure S12), except for the SH catchment, where C_{med} and Q_{med} were significantly higher during Rain-Dry-Vol events. 422





423 424 Figure 7. Event-specific CQ slopes (slope between nitrate concentrations and discharge in the 425 log-log space) for all six study catchments and the two event classes Rain-Dry-Int and Rain-Dry-426 Vol, representing rain-induced runoff events that occurred under dry antecedent soil moisture 427 conditions with either *intensity*- or *volume-dominated* rainfall. Orange asteriscs in the header 428 indicate significant (p < 0.05) differences between the event classes for a particular catchment.

429 Grey-shaded areas indicate event-specific CQ slopes close to zero (between -0.1 and 0.1).

430 **4** Discussion

431

4.1 Impact of runoff event characteristics on nitrate export

432 The main aim of this study was to understand if and how runoff event characteristics, as 433 characterized in the typology of events, relate to nitrate export during and across runoff events. 434 Our results show that in general nitrate export across events is strongly driven by event size (in 435 regard to runoff) with a pronounced seasonality. Large events in winter and spring exported 436 highest nitrate concentrations and loads and small events during summer and autumn exported 437 lowest concentrations and loads. The variability of event-specific CQ slopes decreased with increasing event size, indicating differences in the dominant drivers for nutrient export between 438 439 low and high-flow seasons. In the following, we discuss the impact of runoff event 440 characteristics on nitrate export during these different conditions in more detail. Further, we 441 discuss how well the analyzed runoff event characteristics from the high frequency period (2013 442 - 2017) represent the long-term runoff event characteristics and implications of trends in the 443 long-term runoff event characteristics.

444 4.1.1 Small runoff events

445 Small and often relatively short rain-induced events occurred mainly during summer and 446 autumn and coincided with dry antecedent soil moisture conditions, classified as Rain-Dry-Int or 447 *Rain-Dry-Vol.* These events exported lowest nitrate concentrations and loads and showed highly 448 variable event-specific CQ slopes. We can explain these relatively low nitrate loadings by a 449 decreased hydrological connectivity (i.e., a low fraction of active flow paths) with lower 450 antecedent soil moisture during summer and autumn, compared to winter and spring. As a result, 451 only nitrate sources in close proximity to the stream network and from sources connected via deeper groundwater flow paths are connected to the stream network (Musolff et al., 2015; 452 Stieglitz et al., 2003; J. Yang et al., 2018). These flow paths during the dry period are generally 453

454 characterized by longer transit times and thus enable more nitrate uptake and removal via
455 denitrification (Ebeling et al., 2020; Ehrhardt et al., 2019; Kumar et al., 2020; Nguyen et al.,

- 456 2021). Moreover, with higher temperatures during summer and autumn, nitrate uptake and
- 457 removal increases, especially in streams and in the riparian zones (Baird et al., 1995; Lutz et al.,
- 458 2020; Rode et al., 2016b). Due to lower discharge during summer and autumn, also the relative
- 459 role of those processes increases (Knapp et al., 2020).

460 The seasonal differences in discharge and concentrations shaped the inter-event CQ 461 relationship that was positive across catchments, reflecting higher concentrations during high 462 flow in winter and spring and lower concentrations during low-flow conditions in summer and 463 autumn. At the event level, CQ slopes can reveal complementary information on the underlying 464 processes of nitrate mobilization and transport within seasons (Godsey et al., 2019; Musolff et 465 al., 2021; Winter et al., 2021). We found large variability in event-specific CQ slopes during 466 small events, reflecting a high variability of mobilization and transport processes. We explain 467 this high variability by an increased relevance of different environmental factors such as i) 468 catchment characteristics and the spatial distribution and connectivity of N sources within a 469 catchment, ii) riparian and in-stream biogeochemical processes and iii) the spatial and temporal 470 organization of the inducing precipitation event. In regard to catchment characteristics, we could 471 show that the more forest-dominated catchments (WB, RB, SH, MD, Figure 1) showed mainly 472 positive event-specific CQ slopes during small events, whereas the more agriculturally 473 dominated catchments (HS and HD) tended towards negative event-specific CQ slopes (Figure 474 6b). The dilution patterns in the agricultural catchments can be explained by relatively high base-475 flow concentrations (reflecting high N input from fertilization) and the spatial distribution of N 476 sources within these catchments. For example, Musolff et al. (2021) argued that due to large 477 buffer strips (100 m) in the HS catchment, there are no or only few nitrate sources in the riparian 478 zones. Hence, small events that activate only proximate flow paths from this area could cause the 479 observed dilution pattern. In the HD catchment, Winter et al. (2021) found that a 480 disproportionally large part of event runoff is generated in the upstream area that is mainly 481 covered by forests and thus exports lower nitrate concentrations. Runoff from this area can thus 482 dilute higher concentrations in base flow, which are largely generated by groundwater from the 483 downstream agricultural areas. Hence, the preferential mobilization from certain areas of lower 484 N availability, here riparian zones or upstream areas, can cause a dilution pattern in catchments 485 with an overall high N input. In contrast, the dominant accretion pattern in the more forested 486 catchments might be explained by a flushing of proximate shallow nitrate sources, likely from 487 the upper soil layers of the riparian zones as also suggested by Musolff et al. (2021) for the WB 488 catchment and a sub-catchment of RB. In regard to in-stream and near-stream processing, several 489 studies argued that biogeochemical processes such as nitrate uptake and denitrification in-stream 490 or in the riparian zones have a stronger relative role for nitrate export during small events (e.g., 491 Knapp et al., 2020; Marinos et al., 2020; Moatar et al., 2017). Hence, variability in these 492 processes through, for example, varying instream temperature (Rode, et al., 2016a) or in the 493 riparian zone partly due to stream water infiltration (Lutz et al., 2020; Nogueira et al., 2021) 494 might be responsible for the observed higher variability between event-specific CQ slopes. This 495 is supported by a study from Heathwaite & Bieroza (2021), who found that nutrient export 496 dynamics during small events can be considerably influenced by diurnal cycling.

By separating runoff event classes into *intensity*- and *volume-dominated* precipitation, we
could show that the impact of the temporal organization of precipitation can explain another part
of the variability in mobilization patterns during small runoff events. *Intensity-dominated* events

500 (Rain-Dry-Int) showed higher event-specific CQ slopes compared to volume dominated events 501 (*Rain-Dry-Vol*) in half of the catchments. Those catchments comprise forested or mixed land use 502 and showed overall positive event-specific CQ slopes for both Rain-Dry-Int and Rain-Dry-Vol 503 events (Figure 1, Figure 7). Both event classes are rain-induced with dry antecedent conditions. 504 During Rain-Dry-Int events however, runoff is generated by a shorter and rather intense rainfall, 505 whereas during Rain-Dry-Vol events, the duration of rainfall is typically longer with a lower 506 ratio of the maximum precipitation rate compared to the total precipitation volume (Tarasova et 507 al., 2020). As argued further above, nitrate mobilized during small runoff events in those forested 508 catchments may mainly stem from shallow and proximate N sources (Musolff et al., 2021). One 509 possible explanation for the difference in event-specific CQ slopes might be that relatively short 510 but intensive runoff events preferentially activate proximate and shallow flow paths and mobilize 511 those shallow N sources. This mobilization then causes an increase in nitrate concentrations that 512 is reflected by the positive event-specific CQ slope. Longer volume-dominated events might 513 create a higher, yet delayed hydrological connectivity with more distant sources than those near-514 stream N source zones, which is reflected in a decreasing event-specific CQ slope. As such, CQ 515 slopes during volume-dominated events approximate more chemostatic patterns and show a

516 higher similarity with larger runoff events under wet antecedent conditions (see section 4.1.2).

517 4.1.2 Large

4.1.2 Large runoff events

518 In contrast to small runoff events in summer and autumn, runoff events in winter and 519 spring were either snow-influenced (*Snow*) or rain-induced (< 5% snowmelt) and generated by 520 volume-dominated precipitation under wet antecedent conditions (Rain-Wet-Vol, Figure 2, Figure 521 3). These two event classes were found to be the largest runoff events in regard to median 522 discharge (Q_{med}) and caused the highest nitrate concentrations and loads (Figure 4, Figure 5). 523 Approximately three quarters of all event-driven loads were exported during Snow and Rain-524 Wet-Vol events, which is in agreement with other studies that reported exceptionally high nitrate 525 export during large rain-on-snow events (Crossman et al., 2016; Koenig et al., 2017; Sebestyen 526 et al., 2009; Seybold et al., 2019). These results underline the important role of Snow and Rain-527 Wet-Vol event classes for nitrate export and show that missing information on the winter period, 528 which is often the case (e.g., Blaen et al., 2017; Carey et al., 2014; Knapp et al., 2020; Wollheim 529 et al., 2017), can lead to a lack of information about the most relevant events for the export of 530 nitrate loads.

531 In temperate climates, rain-on-snow events often form the largest runoff events of the 532 year due to the cumulative effect of rainfall and additional input from snowmelt (Casson et al., 533 2014; Pellerin et al., 2012). Nevertheless, we identified Rain-Wet-Vol events (not influenced by 534 snowmelt) that caused comparable or even higher Q_{med}, especially in the Selke catchment 535 (Figure 5d-f). Those events transported comparably high nitrate loads (Figure 4), fell on the same 536 or a very similar inter-event CQ slope (Figure 5) and showed similar event-specific CQ slopes 537 (Figure 6) as the Snow events. This indicates that both event classes, Snow and Rain-Wet-Vol, 538 activate the same or very similar N sources within a catchment, despite their differences in the 539 meltwater fraction.

540 Similar to Stieglitz et al. (2003), we argue that large runoff events during winter and
541 spring can activate all relevant nitrate sources within a catchment, including distant sources
542 (Bowes et al., 2015) and shallow and younger N sources (Fovet et al., 2018; Musolff et al., 2015,
543 2017; J. Yang et al., 2018). During winter and spring, discharge and antecedent soil moisture are

544 generally higher, which leads to more active flow paths compared to summer and autumn 545 (Stieglitz et al., 2003; J. Yang et al., 2018). Together with a reduced N demand of ecosystems in 546 winter and spring compared to summer (Baird et al., 1995; Rode et al., 2016b), this flow path 547 activation in a highly saturated catchment can explain the high nitrate concentrations and loads 548 observed in the studied catchments (Figure 4, Figure 5). Moreover, it can explain the low 549 variability in event specific-CQ slopes (Figure 6), because if all flow paths are activated, no 550 changes in nitrate mobilization through bypassing or activation of additional N sources can 551 occur.

552 Remarkably, the event-specific CQ slopes during large events did not show any signs of 553 dilution (Figure 6). Other studies have reported such dilution pattern during precipitation events 554 across the whole year including large events, which might indicate source depletion (Kincaid et 555 al., 2020; Vaughan et al., 2017) or high base flow concentrations from deeper groundwater that 556 are diluted by lower concentrated water from newly activated zones (Fovet et al., 2018; Rose et 557 al., 2018). Here, we reported consistently slightly positive CQ slopes (roughly 0.1 - 0.3) that 558 reflect a milder increase of concentrations compared to that of discharge, indicating increasingly 559 chemostatic export patterns with increasing event runoff. This is further supported by the fact 560 that the event-specific coefficient of variation of concentrations is much smaller than that of 561 discharge (Musolff et al., 2015) with a median ratio of 0.28 for large events ($O_{med} > 1 \text{ mm h}^{-1}$). 562 These patterns provide strong evidence for a transport rather than a source limitation of N in all 563 six catchments (Basu et al., 2010), even in the forest dominated catchments, which is alarming in 564 terms of water quality. While fertilization in the agricultural catchments (HS and HD) is likely 565 the main N source (Ehrhardt et al., 2019; Winter et al., 2021), there is reported evidence for high atmospheric N deposition over the Harz Mountains (Kuhr et al., 2014; Winter et al., 2021) that 566 567 could explain part of those non-depleting N sources in the more forested catchments. 568 Additionally, Ohte et al., (2004) and Sebestyen et al. (2009), showed that atmospheric N stored 569 in the snowpack can considerably contribute to nitrate export during snow-influenced events in a 570 forested catchment. Nevertheless, the strikingly similar nitrate export patterns during Snow and 571 Rain-Wet-Vol events with comparable event size hints at similar N sources for both event classes and thus not at the melting snowpack as key source. Hence, we hypothesize that a large part of 572 573 those N sources stem from legacy N stores in the soils, either from the previous summer or from 574 N accumulation over longer time periods (Dupas et al., 2020; Van Meter et al., 2016).

4.2 Long-term trends of event characteristics and their implications for nutrient exportpatterns

577 We analyzed nitrate export patterns for a 5-year period of high-frequency nitrate 578 concentration data (2013 - 2017), which is not sufficient to estimate any long-term trends. The 579 analyzed long-term runoff event characteristics from daily data, however, allowed us to detect 580 those trends and discuss their possible impact on nitrate export patterns.

We found a decrease of soil moisture in summer, which aligned along a decrease of wet compared to dry events. This is in agreement with increasing summer temperatures over Europe (Briffa et al., 2009; IPCC, 2013) and other studies that report a decreasing contribution of summer precipitation (Szwed, 2019) and an increased risk for summer droughts in large parts of Europe (Hari et al., 2020; Pal et al., 2004). Here, we found that runoff events generated during dry catchment conditions are associated with smaller event size (in regard to runoff), proportionally lower nitrate concentrations and loads and a higher variability in CQ slopes, 588 compared to wet conditions. Therefore, possibly drier antecedent conditions resulting from

- 589 increasing future temperatures (IPCC, 2018; Pal et al., 2004) might lead to a decrease in nitrate
- 590 export in summer periods but also to a higher variability in concentrations, due to more variable
- and partly higher event-specific CQ slopes. Additionally, rewetting after especially dry summer
- 592 periods was often reported to cause disproportionally high nitrate export peaks, which can cause 593 severe water quality deteriorations and increase the inter-annual variability of nitrate
- 595 severe water quality deteriorations and increase the inter-annual variability of intrate 594 concentrations (Jarvie et al., 2003; Morecroft et al., 2000; Mosley, 2015; Oborne et al., 1980). In
- summary, we see evidence for an increased variability of nitrate concentrations and export
- 596 dynamics with increasingly dry conditions in summer and autumn.

597 In addition to the increasingly dry summer conditions, we found a decrease in the 598 contribution and number of snow-influenced events (Snow) as well as a decrease in the 599 proportion of meltwater during winter and spring. These events exported highest nitrate loads; 600 hence a decrease of large nitrate export peaks could be expected, which was also reported by 601 Sebestyen et al. (2009) for a mountainous forested catchment. However, winter precipitation is 602 predicted to substantially increase in most of Europe (Stahl et al., 2010). The resulting larger 603 rain-induced events could potentially counterbalance the decrease of Snow events and trigger 604 similarly high event runoff and nitrate export, as observed in the SH, MD and HD catchments 605 (Figure 5d-f). As such, nitrate export peaks would not be restricted to the melting period but to 606 the entire high flow season in winter and early spring. Additionally, several studies predict that 607 an earlier start of snowmelt due to increasing temperatures causes a time shift of discharge and 608 nitrate export peaks towards earlier in the year (Clow, 2010; IPCC, 2014; Sebestyen et al., 2009). 609 In summary, we do not see clear evidence for a change in nitrate loading during large winter and 610 spring events but we do see evidence for a change in the timing of nitrate export peaks.

611

4.3 How representative are the obtained results for these and other catchments?

612 Classification of long-term series of runoff events in this study allowed for a consistent 613 characterization of typical hydro-meteorological and catchment-state conditions, their 614 seasonality, and temporal changes (i.e., trends) in their composition. To our knowledge, this 615 placement of short-term nitrate export dynamics into a larger context of long-term runoff event 616 characteristics has never been conducted before. Runoff event classes from the shorter and more 617 recent high frequency period (with available nitrate concentration data) deviated from the long-618 term average runoff event classes mainly in their proportion of *Rain-Dry* events (which mainly 619 increased in summer) and in their proportion of snow-influenced events (which mainly decreased 620 in spring). These deviations can help us understand possible trajectories of runoff event 621 characteristics and nitrate export in the future. Furthermore, these long-term runoff event 622 characteristics allow us to embed the observed catchments into a larger group of catchments with 623 very similar runoff event characteristics, classified by Tarasova et al. (2020). The six studied 624 catchments match well with the clusters that characterize runoff events in the Central Uplands of 625 Germany (including the Harz Mountains where this study is located) and in the Alpine Foreland 626 (Tarasova et al., 2020). Over a time period from 1979 - 2002, runoff events in these clusters were typically dominated by *Rain-Dry* events, while approximately 15% – 25% were snow-influenced 627 628 events (Snow) and volume-dominated rainfall prevailed over intensity-dominated rainfall 629 (Tarasova et al., 2020). This is well in line with our results that include more recent years (until 630 2018) and show >50% Rain-Dry events, 18% – 25% Snow events and volume-dominated rainfall 631 prevailing over *intensity-dominated* rainfall. Basesd on this, we argue that our observed runoff 632 event classes are representative for many upland areas and forelands of higher mountain ranges

633 in a temperate climate. Nevertheless, to get a representative picture of nitrate export during those

- runoff events, one needs to consider that export also depends on additional factors, such as the
- amount and distribution of N sources within a catchment, which are strongly driven by land use
- and its spatial distribution (Dupas et al., 2019; Musolff et al., 2017) among other things.
- 637 However, by analyzing the impact of these representative runoff event characteristics on nitrate
- export in six catchments that span a significant range of different land use types and othercharacteristics (Table 1), we are confident that the presented results are generally transferable to
- 640 other upland areas and mountain forelands in a temperate climate.
- Including an extended set of hydro-meteorological variables into our analysis enabled us to disentangle a large part of the variability in nitrate export patterns and to create results that are better transferable to other catchments and time periods. A hydrological classification can thus be seen as one prerequisite for creating transferrable results in both space and time to better compare the partly contradicting results between different studies (e.g., Knapp et al., 2020; Koenig et al., 2017; Rose et al., 2018; Vaughan et al., 2017; Winter et al., 2021) and to create a more coherent picture of the processes that shape nitrate export dynamics.

648 **5 Conclusions**

649 In this study, we analyzed the impact of runoff event characteristics on high-frequency 650 nitrate export in six contrasting mesoscale catchments. We used long-term runoff event 651 characteristics to embed the relationship between event runoff and nitrate export into a larger 652 hydrological description of events and catchments. This framework allowed us to identify 653 potential long-term trends in nitrate export and their implications under a changing climate. We 654 found that nitrate export differed substantially between runoff events with different 655 characteristics, and strong drivers being event size and a pronounced seasonality. With our 656 findings, we argue that the variability and timing of nitrate export is likely to change with a 657 changing frequency of event types that is driven by future global warming i.e., projected changes 658 in temperatures and other hydro-meteorological conditions.

659 Lowest nitrate concentrations and loads were transported during small rain-induced 660 events with dry antecedent soil moisture (Rain-Dry-Int and Rain-Dry-Vol), which occurred 661 mainly during summer and autumn. These lower nitrate loadings, compared to high-flow 662 seasons, can be explained by a small fraction of active flow paths, longer residence times, and 663 higher nitrate uptake and denitrification rates during the vegetation period. Additionally, we 664 found an increasing variability of event-specific CQ slopes with decreasing event size. We 665 explain this high variability by an increased relevance of different environmental factors for nitrate export dynamics, such as the horizontal and vertical distribution of nitrate sources and 666 667 their connectivity to the streams, as well as the spatial and temporal distribution of precipitation (i.e., volume- or intensity-dominated) and biogeochemical processes in-stream and in the riparian 668 669 zone. Consequently, more frequent dry spells will likely lead to more variable and less 670 predictable water quality in rivers and streams.

In contrast, highest nitrate concentrations and loads were exported during large
snowmelt-induced (*Snow*) or *volume-dominated* rain-induced events under wet antecedent
conditions (*Rain-Wet-Vol*), which occurred mainly during winter and spring. Nitrate
mobilization, represented by event-specific CQ slopes, was surprisingly homogenous among
large events across all catchments, showing a relatively small increase of nitrate concentrations
compared to discharge (approximately chemostatic), which we explain by the activation of all

677 relevant flow path within a catchment that facilitate the land-to-stream connection of all relevant

- N sources. As classes for large events, i.e., *Snow* and *Rain-Wet-Vol*, showed a very similar
- 679 nitrate export behavior, we suggest that not the meltwater fraction, but instead other common
- 680 characteristics such as event size and catchment saturation are the main drivers of nitrate export 681 during large runoff events. No dilution patterns (negative event-specific CQ slope) were
- 682 observed for those events; hence, even forest-dominated catchments showed no sign of N source
- 683 depletion, which could be a warning sign for future water quality trends. Increasing temperatures
- 684 might cause a change in the timing of large nitrate export peaks within the high flow season, but
- 685 we could not find evidence for a change in the amount of nitrate export, because declining
- snowfall (and consequently snow-influenced events) could potentially be compensated for byincreasing winter rainfall.

688 Runoff event characteristics in this study are generic and hence comparable between 689 catchments. Therefore we argue that they are also representative for other upland or foreland 690 areas in temperate climates. Covering a range of different catchment characteristics, e.g., 691 dominantly forested versus mainly agricultural land cover, allowed us to analyse various 692 catchment configurations and the respective event-driven nitrate export patterns and thus to 693 represent a range of possible generic relationships between runoff event types and nitrate export. 694 The potential of a hydrological event classification to create transferable results should be further 695 exploited to analyze event-driven nitrate export also in catchments from other areas, ideally in 696 larger-scale comparisons across a wide range of catchment characteristics, N input histories and 697 climatic conditions. Establishing robust relationships between runoff event characteristics and 698 nitrate export, and relating them to long-term trends in runoff event characteristics would be an 699 informative tool for understanding possible directions of future changes in water quality.

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708 Data availability

- 709 The raw discharge data can be freely obtained from the State Office of Flood Protection
- and Water Quality of Saxony-Anhalt (LHW) under https://gld-sa.dhi-wasy.de/GLD-Portal/. The
- 711 raw meteorological data sets can be freely obtained from Germanys National Meteorological
- 712 Service (Deutscher Wetterddienst, DWD) under
- 713 ftp://opendata.dwd.de/climate_environment/CDC/grids_germany/daily/regnie/ (daily
- 714 precipitation) and
- $715 \qquad ftp://opendata.dwd.de/climate_environment/CDC/grids_germany/hourly/radolan/reproc/2017_00$
- 716 2/ (hourly precipitation). Gridded products based on Zink et al. (2017) are available from
- 717 https://www.ufz.de/index.php?en=41160. Raw nitrate concentration data are archieved in the
- TERENO data base and are available upon request through the TERENO-Portal

(www.tereno.net/ddp). All runoff event characterisics from the long-term and from the high-frequency data will be available in hydroshare upon acceptance of the manuscript.

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