

# Prevalence of updip rupture propagation in interplate earthquakes along the Japan Trench

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## Abstract

The development of seafloor seismic observations facilitates reliable estimation of the rupture directivities of offshore earthquakes. We used seismic waveforms obtained by a new seafloor seismic network (S-net) and onland stations to systematically examine the rupture directivities of interplate earthquakes along the Japan trench. We estimated the rupture directions of 206 (M<sub>w</sub> 3.5-5) events, most of which occurred near the base of the seismogenic zone. We found that most earthquake ruptures (>~80 %) were directional, primarily propagating in the updip direction. This tendency cannot be explained by the effect of the bimaterial interface. The prevalence of updip rupture in the data suggests that deep, steady creep and upward fluid migration along the plate interface affected earthquake ruptures in the subduction zone. The updip ruptures redistributed the accumulated shear stress from the base of the seismogenic zone to the shallow large seismic patches. Furthermore, the updip ruptures may open up ways for deeper fluids to migrate further upward along the plate interface. Both the stress redistribution and the upward fluid migration facilitate the occurrence of shallow megathrust earthquakes.

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1 **Prevalence of updip rupture propagation in interplate earthquakes along the Japan Trench**

2

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19 **Abstract (175 words)**

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25 earthquake ruptures (>80 %) were directional, primarily propagating in the updip direction. This  
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28 interface affected earthquake ruptures in the subduction zone. The updip ruptures redistributed the  
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30 patches. Furthermore, the updip ruptures may open up ways for deeper fluids to migrate further  
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32 facilitate the occurrence of shallow megathrust earthquakes.

33

34 **Keywords**

35 Rupture directivity, Interplate earthquake, Japan trench, Deep creep, Fluid migration, Subduction  
36 zone

37

## 38 **1. Introduction**

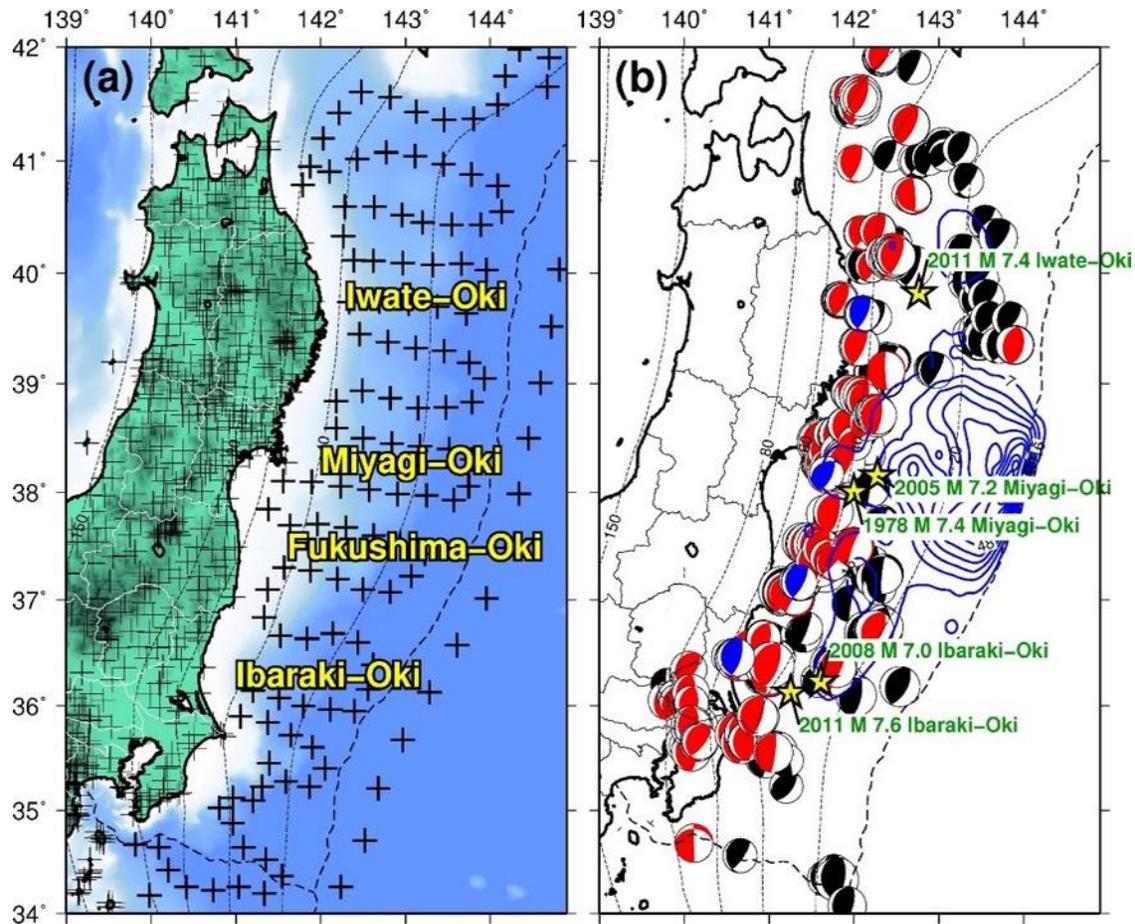
39 A physical understanding of earthquakes involves not only the static characteristics of the  
40 rupture, but also the dynamics. Earthquake ruptures are often approximated by symmetrical  
41 expansions of circular fault patches (e.g., Sato and Hirasawa, 1973). However, according to  
42 systematic investigations, many earthquake ruptures propagate asymmetrically (McGuire et al.,  
43 2002; Chounet et al., 2018; Yoshida, 2019). Rupture directivity is a fundamental characteristic of  
44 earthquake growth and is essential for improving our current physical understanding of  
45 earthquakes. Besides, rupture directivity provides information regarding the stress redistribution  
46 process that occurs along the fault.

47 The effect of material contrast across a fault interface (bimaterial effect) has received  
48 attention for its role in understanding earthquake rupture direction (Weertman, 1980; Shi and  
49 Ben-Zion, 2006). When an in-plane subshear rupture propagates in the slip direction of the  
50 compliant side, a dynamic reduction in normal stress occurs, leading to a lower frictional  
51 resistance and hence a higher rupture velocity in this direction. The bimaterial effect leads to a  
52 hypothesis that a preferable rupture direction may exist, controlled by the bimaterial effect which  
53 is independent of the earthquake size. Some reports have concluded that the observed dominant  
54 rupture direction of small earthquakes on the San Andreas fault is consistent with the prediction  
55 derived from the bimaterial effect (e.g., Lengline and Got, 2010; Kane et al., 2013). However,  
56 others argue that effects other than the bimaterial effect, such as spatial variations in background  
57 stress, play more crucial roles in earthquake rupture directivity (Harris and Day, 2005; Kane et  
58 al., 2013). It is important to assess the above hypothesis by incorporating reliable data from  
59 different tectonic settings as this hypothesis is involved in predicting the rupture directions of  
60 future earthquakes. Currently, the rupture characteristics of interplate earthquakes have been

61 examined in detail, mostly for a small number of large earthquakes in subduction zones  
62 (Yamanaka and Kikuchi, 2004; Ye et al., 2016). However, the rupture characteristics of a large  
63 amount of small- to moderate-sized earthquakes have not been thoroughly studied, despite their  
64 potential to improve the significance level of the direction estimates by increasing the sample  
65 size. The reliability of the rupture-direction measurement of small earthquakes, however, highly  
66 depends on the azimuthal coverage of observations. Poor azimuthal coverage makes the reliable  
67 estimation difficult. Interplate earthquakes in subduction zones mostly occur beneath the ocean,  
68 where seismic stations are usually scarce, causing a paucity of data.

69 The present study systematically estimates the rupture directivity of interplate earthquakes  
70 (Mw 3.5–5) along the Japan trench to investigate earthquake rupture dynamics. Many large  
71 earthquakes have occurred along the Japan trench and caused disasters in Japan, including the  
72 2011 M 9 Tohoku-Oki earthquake. Recently, Seafloor observation network for earthquakes and  
73 tsunamis along the Japan Trench (S-net), was installed along the Japan trench by the National  
74 Research Institute for Earth Science and Disaster Resilience (NIED) (Fig. 1a; Aoi et al., 2020).  
75 Seismic waveform data from September 2016 onwards are available online  
76 (<https://www.seafloor.bosai.go.jp/>). Fig. 1 (a) shows the distribution of S-net stations,  
77 surrounding the epicenters of the recent interplate earthquakes. Waveform records from the S-net  
78 and onland stations provide unique observations with good azimuthal coverage for interplate  
79 earthquakes occurring beneath the ocean. It is thereby possible to systematically examine the  
80 rupture directivity of small earthquakes along the Japan trench.

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83

84 **Fig. 1.** The distribution of earthquakes and seismic stations in the study area. (a) Distribution of  
 85 onland and S-net seismic stations. The thick and thin cross indicates S-net stations and onland  
 86 stations, respectively. (b) Distribution of earthquake focal mechanisms is indicated by beach balls.  
 87 Red beach balls indicate earthquakes for which rupture parameters were estimated, while black  
 88 beach balls indicate earthquakes for which rupture parameters were not estimated. Blue beach balls  
 89 indicate earthquakes for which the results are shown in Fig. 2. The contour lines in blue denote the  
 90 coseismic slip distribution of the 2011 Tohoku-Oki earthquake determined by Iinuma et al. (2012).  
 91 Yellow stars represent the five  $M > 7$  interplate earthquakes (1978 M 7.4 Miyagi-Oki, 2005 M 7.2  
 92 Miyagi-Oki, 2008 M 7.0 Ibaraki-Oki, 2011 M 7.4 Iwate-Oki, and 2011 M 7.6 Ibaraki-Oki

93 earthquakes) that occurred near the earthquakes analyzed in this study. Dotted contours indicate  
94 the depth to the upper plate interface of the subducting Pacific plate, from Nakajima et al. (2009)  
95 and Kita et al. (2010).

96

## 97 **2. Data and Methods**

98 This study incorporates waveforms obtained from the seismic network composed of stations  
99 from the Japan Meteorological Agency (JMA), national universities, Hi-net (NIED, 2019a), F-  
100 net (NIED, 2019b), V-net (NIED, 2019c), and S-net (NIED, 2019d), as shown in Fig. 1(a),  
101 after removing the instrument responses. Initially, interplate earthquakes were selected (section  
102 2.1) and the apparent moment rate functions (AMRFs) were estimated, based on the empirical  
103 Green's function (EGF) (section 2.2). Next, the rupture direction was estimated by fitting the  
104 simple rupture model described by Haskell (1964) to the duration of the AMRFs (section 2.3).

105

### 106 **2.1. Selection of interplate earthquakes**

107 The F-net moment tensor catalog  
108 (<https://www.fnet.bosai.go.jp/event/dreger.php?LANG=en>) and JMA unified earthquake catalog  
109 ([https://www.data.jma.go.jp/svd/eqev/data/bulletin/hypo\\_e.html](https://www.data.jma.go.jp/svd/eqev/data/bulletin/hypo_e.html)) were used to select the interplate  
110 earthquakes. The F-net moment tensor catalog lists the moment tensors of earthquakes with  $M_w$   
111 greater than about 3.5 in and around Japan, while the JMA unified catalog includes location data  
112 for earthquakes with smaller magnitudes. The data period extended from August 2016 to  
113 December 2019, for which the S-net station seismograms were also available.

114 In total, 420 interplate events (target events) were selected using the F-net moment tensors  
115 based on the criteria used by Asano et al. (2011) and Hasegawa et al. (2012). The requirements

116 were as follows: a rake angle of  $> 0^\circ$ , a focal mechanism 3-D rotation angle (Kagan, 1991) relative  
117 to that of a reference interplate earthquake (strike:  $195^\circ$ , dip:  $15^\circ$ , rake:  $90^\circ$ ) of  $< 35^\circ$ , and a depth  
118 separation of  $< 20$  km between the centroid and the plate interface (Nakajima et al., 2009). To  
119 estimate the directivity of the target events, 660 earthquakes were selected and their waveform  
120 data were used for waveform deconvolution (EGF events). These EGF events satisfied the  
121 following two criteria: (1) the hypocentral distance from the target event was  $< 3.0$  km according  
122 to the JMA unified catalog, and (2) the magnitude was 1–2 magnitudes smaller than the target  
123 earthquake. Each target earthquake may have multiple EGF events. Note that the hypocenter  
124 locations of the offshore earthquakes had relatively large estimation errors, and thus, earthquakes  
125 located far from the target earthquake may sometimes be selected as EGF events. However, these  
126 were removed during the waveform deconvolution procedure, as described in section 2.2.

127

## 128 **2.2. Determination of the apparent moment rate function and source duration**

129 For waveform deconvolution, we followed the procedure of Yoshida (2019), which adopted  
130 the method of Ligorria and Ammon (1999). The transverse components of direct S-waves were  
131 used. The orientations of the S-net seismometers determined by Takagi et al. (2019) were used.  
132 We first applied a low pass filter to the waveforms of the target and the EGF events. The cut-off  
133 frequency ( $f_1$ ) should be higher than the source corner frequency of the target earthquake. Based  
134 on the source model of Sato & Hirasawa (1973), the source corner frequency of S-wave is related  
135 to the stress drop  $\Delta\sigma$  as follows:

$$136 \quad f_{cs}(\Delta\sigma) = kV_s \left( \frac{16 \Delta\sigma}{7 M_0} \right)^{\frac{1}{3}} \quad (1)$$

137 Where,  $V_s$  is the S-wave velocity and  $M_0$  is the seismic moment. We assumed  $k = 1.9$

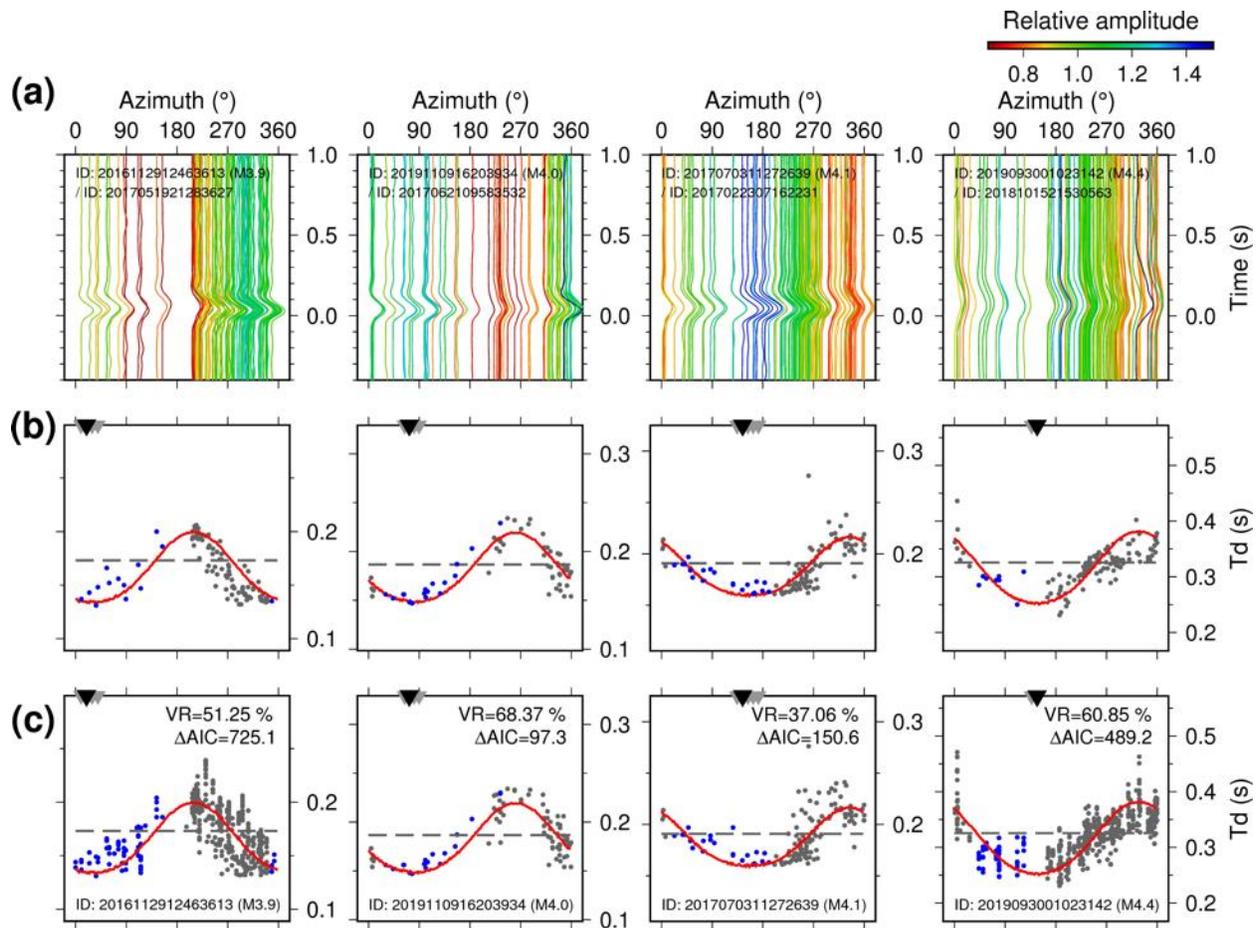
138 corresponding to the situation where the rupture speed is 0.9 times  $V_s$ . The typical value of stress  
139 drop,  $\Delta\sigma$  of interplate earthquakes is 3 MPa (Kanamori & Anderson, 1975). We used  $f_1 =$   
140  $f_{cS}(\Delta\sigma = 50 \text{ MPa})$ . Thus,  $f_1$  is sufficiently higher than the typical source corner frequency of  
141 interplate earthquakes  $f_{cS}(\Delta\sigma = 3 \text{ MPa})$ .

142 The AMRFs were then derived by deconvoluting the waveforms of the target events using  
143 the EGFs with a positive value constraint. Synthetic waveforms for the target events can be  
144 obtained by convoluting the derived AMRFs with the EGFs. When a synthetic waveform does not  
145 reproduce the observed waveform adequately, it implies that the EGF was inappropriate for the  
146 observed waveform. Waveform deconvolution thereby naturally removes such inappropriate EGFs  
147 (Abercrombie, 2014). Synthetic waveform which reproduced less than 80% of the observed  
148 waveform measured in the variance reduction was discarded. If the number of AMRFs from a  
149 single EGF event were less than eight, we discarded the AMRFs from that event. The corner  
150 frequency (Andrews, 2013)  $f_c$  for each AMRF was computed and the apparent source durations  
151 were obtained  $T_d = \frac{1}{2f_c}$ .

152 Figs. 2 and S1 denote examples of the azimuthal distributions of the AMRFs and the source  
153 durations. AMRFs were obtained from more than 15 different stations for 206 of the initial 420  
154 target earthquakes (Fig. S2), and their rupture directivities were examined in detail. The rupture  
155 directions of the remaining 214 earthquakes could not be estimated owing to an insufficient  
156 number of stations with available AMRFs. For the 214 events, the minimum distances of the EGF  
157 events tended to be larger ( $> 1.5 \text{ km}$ ) than those of the other 206 events (0–1.5 km) (Fig. S2). This  
158 tendency implies that the scarcity of AMRFs for the 214 events was caused by the absence of an  
159 appropriate EGF event, probably because the candidate EGF events were too far to assume the  
160 same path effect. This natural selection of analyzed events does not bias the statistical tendency of

161 the rupture directions because the presence or absence of nearby ( $< 1.5$  km) small events is not  
 162 related to the rupture direction.

163



164

165

166 **Fig. 2.** (a) Examples of the AMRF azimuthal distributions for events shown in Fig. 1(b) from north  
 167 to south. The different colors indicate the relative maximum amplitude as per the scale provided  
 168 on the top. (b), (c) Examples of AMRF corner frequencies from a single EGF event and those  
 169 compiled from multiple EGF events, respectively. Azimuth is measured in degrees, clockwise from  
 170 the north. Blue dots indicate the results from S-net stations, and gray dots indicate those from  
 171 onland stations. Black triangles indicate the best-fit azimuth of rupture propagation based on the

172 unilateral rupture model. Gray triangles indicate the results from bootstrap re-sampling. The red  
 173 curve and the dashed gray line indicate the computed azimuthal dependences of the corner  
 174 frequency from the best-fit unilateral model and the non-directional model, respectively. Event IDs  
 175 (from the JMA catalog) are shown in the lower-left corner of each frame with the moment  
 176 magnitude (from the F-net catalog).

177

### 178 **2.3. Determination of rupture directivity**

179 To estimate the rupture directivity, we followed the procedure described by Yoshida et al.  
 180 (2019), which uses the 1-D unilateral rupture model (Haskell, 1964) to compute the apparent  
 181 rupture duration:

$$182 \quad T_{d0}^{\text{model}} = T_0(1 - \eta \vec{R} \cdot \vec{U}). \quad (2)$$

183 Where,  $T_0$  is the actual rupture duration,  $\eta$  is the speed of the unilateral rupture divided by  $V_S$ , and  
 184  $\vec{R}$  and  $\vec{U}$  represent the unit vectors of the rupture propagation and the ray direction at the source,  
 185 respectively. The shape of a theoretical AMRF is assumed to be a symmetrical triangle with the  
 186 base of  $T_{d0}^{\text{model}}$ . We applied the same low pass filter to the synthetic triangle and obtained the  
 187 duration  $T_d^{\text{model}}$  as the same way used for observed AMRFs.

188 We grid-searched for the combination of  $\vec{R}$ ,  $T_0$ , and  $\eta$  that best explained the obtained  
 189 apparent source durations.  $\vec{R}$  was searched on the west-dipping nodal planes of focal mechanisms  
 190 that corresponded to the plate boundary. The evaluation function is as follows:

$$191 \quad \text{Var}_{\text{uni}} = \frac{\sum_{i=1}^n (T_d^{\text{model}}(T_0, \eta, \vec{R}, \vec{U}_i) - T_{d_i})^2}{n} \quad (3)$$

192 where  $T_{d_i}$  and  $\vec{U}_i$  are  $T_d$  and  $\vec{U}$ , respectively, of the  $i$ -th AMRFs, and  $n$  is the number of the  
 193 AMRFs.  $\vec{U}_i$  was computed based on the 1-D velocity model of Ueno et al. (2002). We changed the

194 apparent speed ratio  $\eta$  (0–1.0) by dividing the interval by 0.05, and the rupture propagation  
 195 direction (10–360°) that determined  $\vec{R}$  by dividing the interval by 10°. The rupture duration  $T_0$  was  
 196 searched from 100 points, evenly dividing the interval on a logarithmic scale between 0.1 and ten  
 197 times the mean value of  $T_{d_i}$ . Ruptures may propagate faster than the S-wave velocity in the Mode  
 198 II direction (Andrews, 1976). However, we considered the likelihood of  $\eta > 1$  to be low because  
 199  $Var_{uni}$  increased beyond  $\eta = 1$  for almost all cases in the dataset used in this study.

200 The residual of Eq. (3) was compared with that of the constant source duration  $\overline{T_d}$  case (we  
 201 referred to the model as the “non-directional model”).

$$202 \quad Var_{\text{mean}} = \frac{\sum_{i=1}^n (\overline{T_d} - T_{d_i})^2}{n} \quad (4)$$

$$203 \quad VR = 100 \% \times \left( 1 - \frac{Var_{\text{uni}}}{Var_{\text{mean}}} \right) \quad (5)$$

204 If a unilateral rupture more suitably models the duration data,  $VR$  will approach 100%.  $\overline{T_d}$  was  
 205 determined during the above grid-search procedure (best-fit value when  $\eta = 0$ ).

206 We computed the Akaike Information Criterion (AIC) (Akaike, 1974) by assuming that  
 207 measurement errors in the apparent duration followed a Gaussian distribution, as  $AIC = n \ln 2\pi +$   
 208  $n \ln Var + 2(m + 1)$ .  $Var$  and  $m$  are the mean squared residual and the number of model  
 209 parameters, respectively.  $m = 1$  and  $m = 3$  was assumed for the non-directional model and the  
 210 unilateral rupture model, respectively. We computed the difference  $\Delta AIC = AIC_{\text{mean}} - AIC_{\text{uni}}$ ,  
 211 where  $AIC_{\text{mean}}$  and  $AIC_{\text{uni}}$  are the AICs of the non-directional model and the unilateral rupture  
 212 model, respectively.

213 We estimated the uncertainty range of the directivity parameters for each earthquake by  
 214 conducting the above procedure on 1,000 simulated datasets based on bootstrap re-sampling of the  
 215 duration data. We determined the 95% confidence intervals of  $\vec{R}$ ,  $T_0$ , and  $\eta$ . Figs. 2(b) and (c)

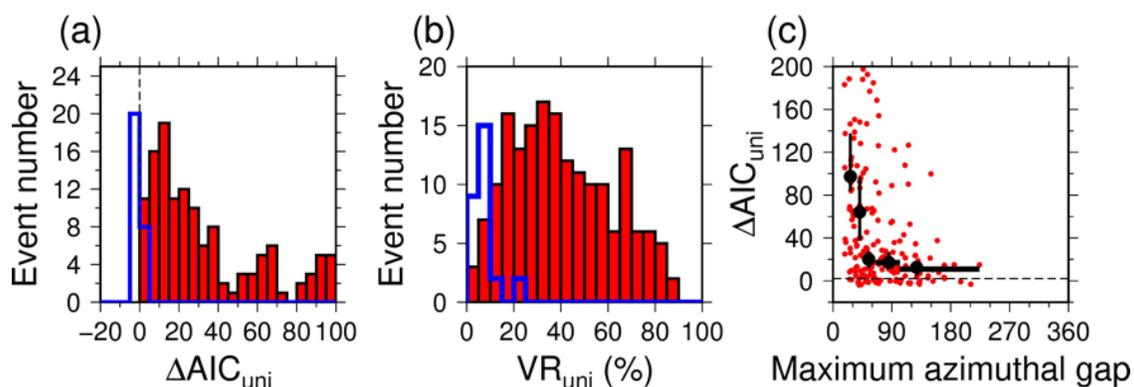
216 include examples of the best-fit values and 95% confidence intervals of the rupture azimuths  
217 determined using this method.

218

### 219 3. Results

220 We obtained the rupture parameters for 206 earthquakes (Mw 3.5–5), for which the  $\Delta AICs$   
221 are shown in Fig. 3 (a) by their frequency distribution.  $\Delta AIC$  was positive in 186 out of the 206  
222 earthquakes, indicating that 186 earthquakes were better modeled by the unilateral rupture model  
223 than by the non-directional model.  $\Delta AIC > 2$  in 178 earthquakes; that is, the unilateral rupture  
224 model was a better model for these earthquakes based on the 95% confidence interval (Yoshida,  
225 2019). For the 178 events with  $\Delta AIC > 2$ ,  $VR$  was larger than 25% for  $\sim 75\%$  of the events, and  
226  $VR$  was larger than 40% for  $\sim 40\%$  of the events (Fig. 3b). The mean  $\Delta AIC$  decreased with an  
227 increasing maximum azimuthal gap of the duration data (Fig. 3c), implying that  $\Delta AIC < 0$  does  
228 not necessarily reflect a non-directional event, but is due to limited data available for resolving the  
229 directivity effect.

230



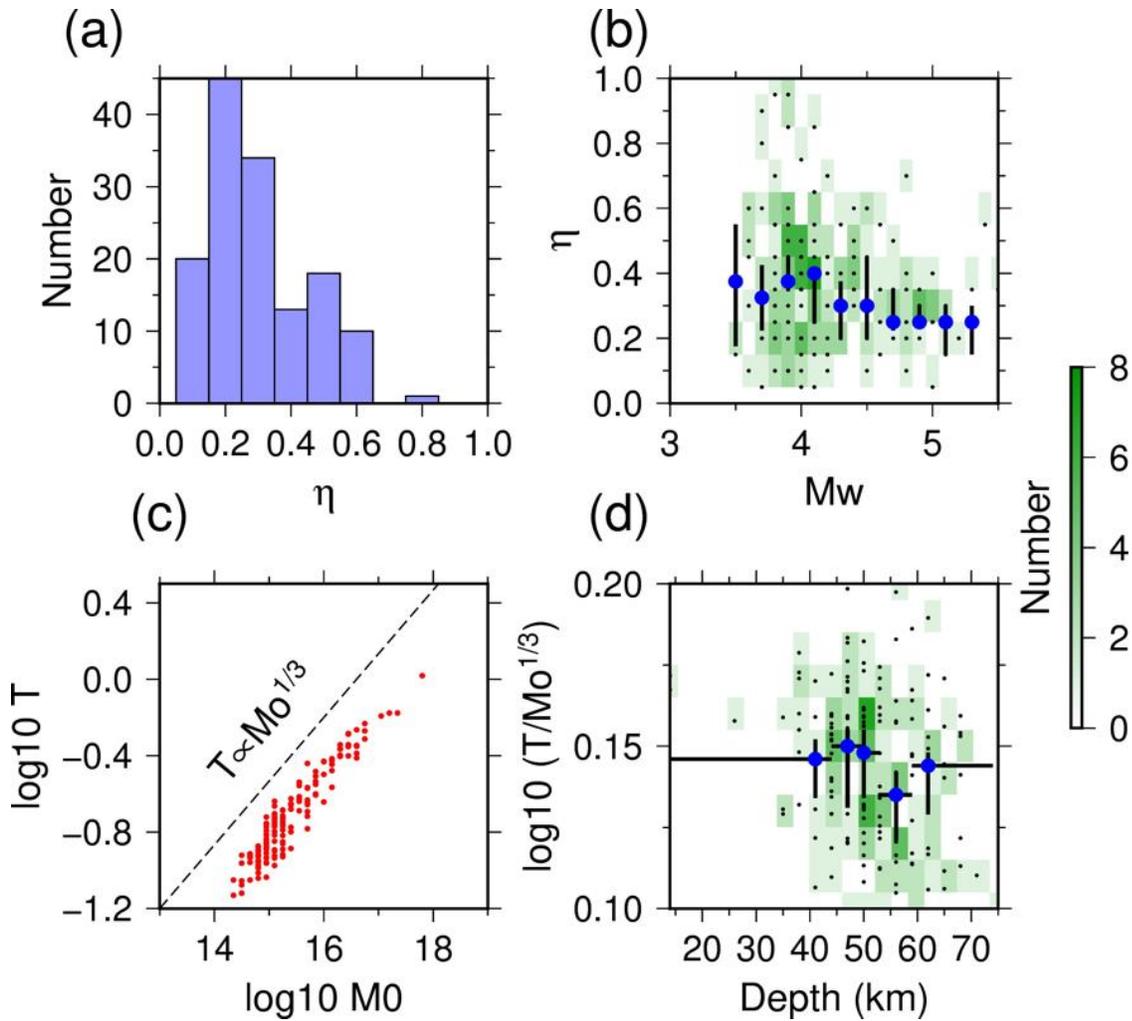
231  
232 **Fig. 3.** Results of applying the unilateral rupture model to the dataset used in this study. (a) and  
233 (b): Frequency distributions of  $\Delta AIC$  and  $VR$ , respectively. The red bars denote the results with  
234  $\Delta AIC > 2$  and blue bars denote those with  $\Delta AIC < 2$  (including negative values). (c): Relationship

235 between  $\Delta AIC$  and the maximum azimuthal gap of the corner frequency data used to estimate the  
236 rupture parameters. Vertical lines indicate 95% confidence intervals of the median value based on  
237 bootstrap re-sampling in each bin with the same number of events ( $n=41$ ). Horizontal lines indicate  
238 the ranges of the maximum azimuthal gap in each bin.

239

240 Fig. 4 shows the obtained rupture parameters for events with  $\Delta AIC > 2$ . The estimated value  
241 of  $\eta$  is shown in Figs. 4(a) and (b) when the 95 % confidence interval was less than 0.2.  $\eta$  mostly  
242 ranged from 0.1 to 0.6 (Fig. 4a).  $\eta$  was variable, and the number of samples for the larger  
243 earthquakes were smaller than others, but may have had a tendency to decrease with the moment  
244 magnitude (Fig. 4b). Source duration  $T_0$  mostly ranged from 0.1 to 1 s and increased by the power  
245 of 1/3 of the seismic moment (Fig. 4c). This relationship is consistent with the self-similarity of  
246 earthquakes (Aki, 1967). No clear dependence of the characteristic source duration  $T_0/M_0^{\frac{1}{3}}$  on the  
247 centroid depth was obtained (Fig. 4d).

248



249

250

251 **Fig. 4.** Derived rupture parameters. Only results with  $\Delta AIC > 2$  are shown. (a): Frequency

252 distribution of  $\eta$ . (b): Relationship between  $\eta$  and the moment magnitude. (c): Relationship

253 between seismic moment  $M_0$  and source duration  $T_0$ . (d): Relationship between the characteristic

254 source duration  $T_0/M_0^{1/3}$  and the centroid depth. Blue circles and vertical lines in (b) and (d)

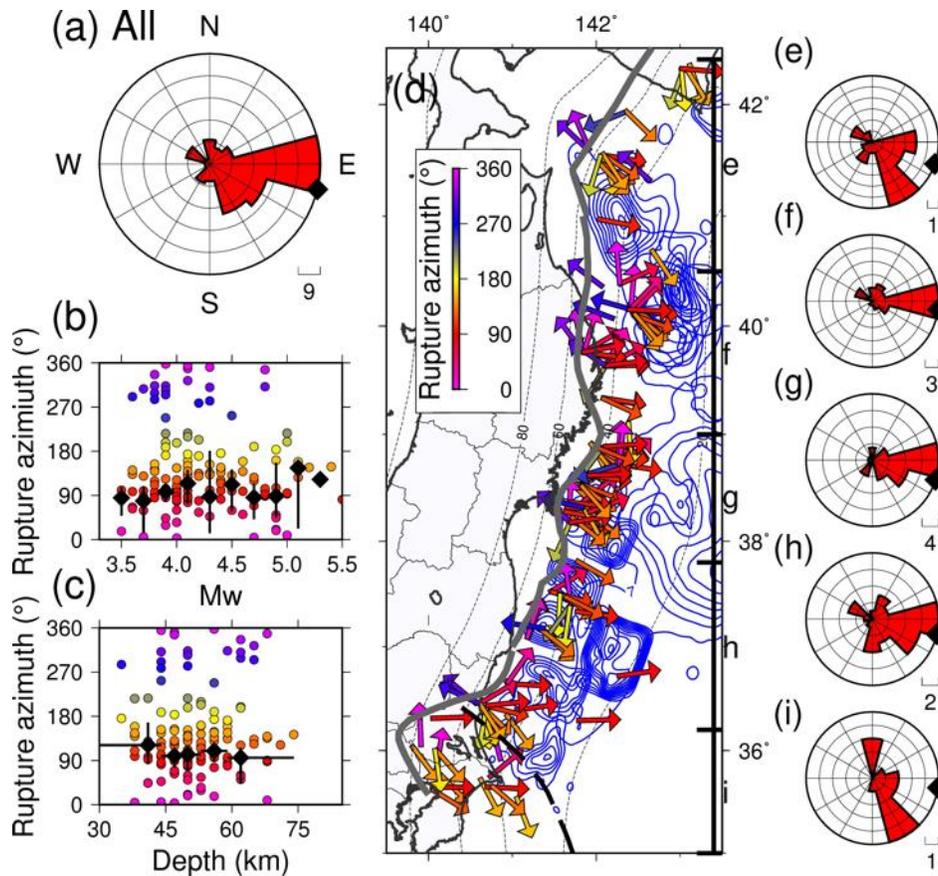
255 indicate mean values and 90% confidence intervals based on bootstrap re-sampling in each bin

256 with the same number of events, respectively. Horizontal lines indicate the depth ranges in the bins.

257 The color scale indicates the number of samples in (b) and (d).

258

259 A clear tendency for the rupture propagation direction was observed. Fig. 5 shows the  
260 rupture propagation directions of the 148 earthquakes that satisfied the following two conditions:  
261 the 95 % confidence interval of the rupture propagation direction was less than  $45^\circ$  and  $\Delta AIC >$   
262 2. The frequency distribution (Fig. 5a) indicates that ruptures of most of the analyzed earthquakes  
263 ( $\sim 80\%$ ) were oriented eastward rather than westward. The ruptures of most of the interplate  
264 earthquakes along the Japan trench propagated in the updip direction. This tendency appears to  
265 be independent of earthquake magnitude (Fig. 5b) and centroid depth (Fig. 5c), and barely  
266 changes, even if different criteria were adopted for showing the results. Fig. S3 shows 178 results  
267 with a 95% confidence interval for a rupture azimuth of less than  $90^\circ$  regardless of  $\Delta AIC$  and Fig.  
268 S4 shows 114 results with a 95% confidence interval for a rupture azimuth of less than  $45^\circ$  and  
269  $\Delta AIC > 20$ .  
270



271

272

273 **Fig. 5.** Estimated rupture propagation directions. Results are indicated for events with a 95%

274 confidence interval of less than  $45^\circ$  and  $\Delta AIC > 2$ . (a): Frequency distribution of the rupture

275 direction azimuth shown in the form of a rose diagram. Black colored diamond indicates the updip

276 direction obtained from the mean values of the focal mechanisms. (b) and (c): Relationships

277 between the rupture direction azimuths and moment magnitude and centroid depth, respectively,

278 colored according to the rupture azimuth scale in (d). Diamonds and vertical lines indicate the

279 mean values and the 90 % confidential intervals based on bootstrap re-sampling, respectively, in

280 each bin with the same number of events. (d): Map view showing the rupture direction azimuths.

281 Blue contours show slip areas of large earthquakes from Nagai et al. (2001), Yamanaka and

282 Kikuchi (2004), Inuma et al. (2012), and Kubo and Nishikawa (2020). The dashed curve indicates

283 the northeastern limit of the Philippine Sea Plate (Uchida et al., 2009). The gray curve denotes the  
284 depth limit of interplate earthquakes on the Pacific plate (Igarashi et al., 2001; Uchida et al., 2009;  
285 Kita et al., 2010). (e)-(i): Frequency distributions of the rupture direction azimuths in the five  
286 latitude ranges as shown in (d).

287

288 The predominance of updip rupture was most comprehensible in the Iwate-Oki and Miyagi-  
289 Oki regions (Figs. 5f and g). The same tendency was also observed in the Fukushima and Ibaraki-  
290 Oki regions (Fig. 6h), which are located to the north of the northeastern limit of the Philippine  
291 Sea plate. In the northernmost region and the southernmost region (south of the northeastern limit  
292 of the Philippine Sea plate), the dominant rupture directions are SSE (Figs. 5e and i). Although  
293 this SSE direction is roughly consistent with the predominance of updip rupture, the fault-strike  
294 component is large.

295

#### 296 **4. Discussion**

297 The results indicate that ruptures of most of the analyzed earthquakes primarily propagated  
298 in the updip direction. This tendency was noticed along the Japan trench at depths of 35–60 km.  
299 This systematic tendency suggests that rupture propagation is not governed by a random process  
300 or by the rupture history but instead it is due to permanent structures present in the system. This  
301 systematic tendency is difficult to explain using small-scale heterogeneities in stress, friction,  
302 material properties, and fault geometry. Large-scale properties (much larger than the fault size of  
303 an individual earthquake) are required to explain this widespread tendency. The effect of a  
304 bimaterial fault interface affects the earthquake rupture directions, where two blocks with  
305 different elastic properties are located across a fault (Weertman, 1980; Shi and Ben-Zion, 2006).

306 This model predicts that Mode-II subshear rupture tends to propagate in the direction of motion  
307 of the compliant side of the fault, which is supported by observations on the San Andreas Fault  
308 (Lengliné and Got, 2011) and laboratory experiments (Anooshehpour and Brune, 1999). This  
309 study used interplate earthquakes that were mostly deeper than 30 km. In this depth range, the  
310 footwall material (oceanic crust) is more compliant than the hanging wall material, because the  
311 hanging wall is the mantle in this depth range (Ito et al., 2004). Therefore, the bimaterial effect  
312 predicts that earthquake ruptures will tend to propagate westward, which is the opposite of what  
313 we observed here. That is, the bimaterial effect alone cannot explain the results of this study.

314 Determining the cause of the rupture direction is difficult, but the depth dependencies of  
315 stress, friction, and pore pressure probably play essential roles in rupture dynamics in the Japan  
316 trench. A similar prevalence for updip rupture propagation has been reported for large interplate  
317 earthquakes in subduction zones (Kato and Seno, 2003; Chounet et al., 2018) and intraplate  
318 earthquakes near the base of the seismogenic zone (Sibson, 1982). One possible cause for the  
319 prevalence of updip rupture is a steady creep that occurs along the deeper portion of the plate  
320 boundary. The earthquakes analyzed in this study were mainly located near the lower limit of the  
321 seismogenic zone (gray curve in Fig. 5), beneath which the plate boundary dislocates aseismically.  
322 The contrast in slip velocity between the creeping region and the locked seismogenic patches  
323 increases the shear stress along the deeper edge of the locked patches. This facilitates rupture  
324 initiation near the deeper edge and propagation in the updip direction, given that the nucleation  
325 size is sufficiently smaller than the final rupture area (Kato and Seno, 2003). Numerical modeling  
326 of earthquake cycles has shown that some small earthquakes can nucleate near the rheological  
327 transition boundary and propagate in a non-preferred direction predicted by the bimaterial effect  
328 (Erickson and Day, 2016). Upward fluid flow along the plate boundary may play a similar role.

329 Fluids dehydrated from the subducting slab are suggested to migrate upward along the plate  
330 boundary because of the permeability anisotropy (Kawano et al., 2013; Sano et al., 2014). Higher  
331 pore pressures result in a higher reduction of fault strength along the deeper portion of the plate  
332 boundary. This may contribute to rupture initiation near the deeper edge of the locked patch and  
333 propagation in the updip direction. The updip ruptures may open up ways for fluids to migrate  
334 further upward along the plate interface.

335 The seismic patches of large earthquakes ( $M > 7$ ) are located in the updip region of the small  
336 earthquakes analyzed in this study (Fig. 5d). The updip ruptures of small earthquakes may  
337 redistribute the shear stress accumulated near the base of the seismogenic zone to the shallow  
338 seismic patches of large interplate earthquakes, thereby facilitating the occurrence of large  
339 earthquakes. Some of these small earthquakes may represent partial ruptures of the seismic patches  
340 of large earthquakes.

341 The dominant rupture direction has a significant fault-strike component in the northernmost  
342 (Fig. 5e) and the southernmost (Fig. 5i) regions. Since the trench axis bends near the northernmost  
343 region, the bending-related stress and structure may affect the rupture directions of earthquakes in  
344 this region. For the southernmost region, the Philippine Sea Plate subducts from the south, making  
345 the situation there different from the other areas. The seismic coupling substantially changes across  
346 the northern limit of the Philippines Sea plate (Uchida et al., 2009). The difference in material  
347 properties and the state of plate coupling may affect the rupture directivity in this region.

348 The predominance of updip rupture obtained in this study appears to be the opposite of the  
349 downdip ruptures reported for some large earthquakes ( $M > 7$ ) that occurred near the earthquakes  
350 analyzed in this study. Fig. 1(b) shows the hypocenters of the  $M \sim 7$  interplate earthquakes near the  
351 earthquakes analyzed in this study. In the 2011 ( $M 7.4$ ) Iwate-Oki and 2011 ( $M 7.6$ ) Ibaraki-Oki

352 earthquakes, the ruptures propagated in the updip direction (Kubo et al., 2013; Kubo and  
353 Nishikawa, 2020) were similar to the results presented herein. However, in the 1978 ( $M$  7.4) and  
354 2005 ( $M$  7.2) Miyagi-Oki, and the 2008 ( $M$  7.0) Ibaraki-Oki earthquakes, the ruptures were  
355 estimated to have propagated in the downdip direction (Wu et al., 2008; Takiguchi et al., 2011).  
356 The bimaterial effect, which facilitates downdip rupture along the deeper portion of the  
357 seismogenic zone, may be more dominant for larger earthquakes because the dynamic changes in  
358 normal stress increase with propagation distance (Shi and Ben-Zion, 2006). The competition  
359 between the bimaterial effect and other effects may cause diversity in the rupture directions of  
360 large earthquakes. The values of  $\eta$  estimated in this study ranged from 0.1 to 0.6, which are  
361 significantly smaller than the typical range of  $V_r/V_s$  (0.6–0.9) (Geller, 1976).  $\eta$  corresponds to the  
362 actual value of  $V_r/V_s$  only when the rupture is completely unilateral; otherwise,  $\eta$  should be smaller  
363 than the actual  $V_r/V_s$  (Abercrombie et al., 2017). This is because the fit of the simple unilateral  
364 model cannot distinguish between slow and bilateral ruptures. The low values of  $\eta$  obtained in this  
365 study suggests that the earthquake ruptures along the Japan trench were directional, but not  
366 completely unilateral.  $\eta$  was variable, but appeared to decrease slightly with earthquake size from  
367  $M_w$  3.5 to 5 (Fig. 4b), which may reflect an enhanced bimaterial effect with magnitude that acts  
368 against the effect of deeper creep.

369         Alternatively, the apparent partial discrepancy in some large earthquake rupture directions  
370 may be due to temporal changes in the preferable rupture direction. Such a temporal change was  
371 reported after the 2004 Parkfield earthquake (Kane et al., 2003). Along the Tohoku-Oki plate  
372 boundary, the stress and frictional state were significantly perturbed by the 2011 ( $M$  9) Tohoku-  
373 Oki earthquake. The results obtained in this study are for the period after the 2011 ( $M$  9) Tohoku-  
374 Oki earthquake, while the three earthquakes with downdip ruptures (1978 and 2005 Miyagi-Oki

375 and 2008 Ibaraki-Oki earthquakes) occurred before the 2011 Tohoku-Oki earthquake. In the  
376 Miyagi-Oki region, interplate earthquakes occurred in many offshore areas before the Tohoku-  
377 Oki earthquake, after which the offshore area was quiet and interplate earthquakes only occurred  
378 near the downdip limit of interplate earthquakes. This was probably due to the stress release and  
379 locking in the offshore region (e.g., Asano et al. 2011, Uchida and Matsuzawa, 2013). The 2011  
380 Tohoku-Oki earthquake may have affected the preferential rupture direction of subsequent  
381 earthquakes, although it is not clear whether the effect of the Tohoku-Oki earthquake can explain  
382 the prevalence of updip rupture throughout the region. Unfortunately, the temporal changes in  
383 rupture directivity cannot be examined before and after the 2011 Tohoku-Oki earthquake using  
384 S-net data, as S-net was installed after the 2011 Tohoku-Oki earthquake. Fig. 3(c) shows the  
385 difficulty of estimating the rupture propagation direction using solely onland stations. However,  
386 the earthquake rupture directivities may be constrained solely by onland stations for few ideal  
387 cases (Figs. 2 and S1). To some extent, it may be possible to constrain the rupture directivities of  
388 small earthquakes along the Japan trench using only onland stations exclusively and assess the  
389 above hypothesis for future research.

390 The results presented herein indicate that the ruptures of most deep interplate earthquakes  
391 along the Japan trench had significant directivities. The prevalence of asymmetrical rupture has  
392 also been obtained for inland intraplate earthquakes in Japan (Yoshida, 2019). McGuire (2002)  
393 conducted global surveys of large earthquakes and concluded that unilateral rupture was  
394 predominant. Most studies of small to moderate-sized earthquakes along the Japan trench have  
395 been conducted mainly by assuming a non-directional model, including estimates of the  
396 earthquake stress drop (e.g., Yamada et al., 2021). However, the results of this study suggest that  
397 this assumption may be inappropriate for many interplate earthquakes. This issue may be reduced

398 by incorporating the earthquake rupture directivity in the estimation method (Yoshida, 2019;  
399 Yoshida et al., 2019). The prevalence of asymmetrical rupture may enable the acquisition of  
400 detailed earthquake information that cannot be obtained if the directivity effects are ignored (e.g.,  
401 fault orientation).

402

## 403 **5. Conclusions**

404 Prevalence of updip ruptures in small earthquakes was observed near the base of the  
405 seismogenic zone along the Japan trench. This updip rupture propagation tendency cannot be  
406 explained by the effect of a bimaterial interface; instead, the results suggest that deep, steady creep  
407 largely affects earthquake ruptures in subduction zones, although upward fluid flow along the plate  
408 boundary may also contribute to the ruptures. The updip ruptures redistribute the shear stress  
409 accumulated near the base of the seismogenic zone to shallow large seismic patches. Also, the  
410 updip ruptures may open up ways for fluids to migrate further upward along the plate interface.  
411 They may facilitate the occurrence of shallow megathrust earthquakes. A similar updip rupture  
412 propagation tendency was reported for large intraplate earthquakes near the base of the  
413 seismogenic zone. However, this study unambiguously demonstrates a similar trend in many small  
414 to moderate-sized interplate earthquakes obtained from both land-based and offshore seismic  
415 networks.

416 The present study could scarcely examine shallow interplate earthquakes ( $z < 30$  km) that  
417 occurred far from the land, due to the difficulty in finding appropriate EGF events for the  
418 earthquakes. The stress release by the Tohoku-Oki earthquake reduces seismicity there. Studies in  
419 the future may be able to examine the rupture characteristics of shallow interplate earthquakes by  
420 detecting more offshore earthquakes and improving the location accuracy based on the seafloor

421 observation data.

422

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