

Structural Nexus and Divide: the Western Tauern Window, Eastern Alps

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Abstract

The western Tauern Window of the Eastern Alps constitutes a mature transpressive wrench zone generated by Oligocene-Miocene convergence between the Dolomites Indenter and the European foreland. As a major structural divide, it translates convergence-parallel north-south shortening into convergence-perpendicular eastward lateral extrusion. About 7000 foliation, lineation, fold-axis, axial-plane, shear-zone, and shear-band measurements at ~1800 sites detail the structural geometry and evolution during indentation and transpression, modeled by numerical and analogue-material experiments. The results outline the western Tauern Window as an orogen-scale zone of localized deformation, consisting of tight, upright folds, reactivated by a sinistral shear-zone network. It connects the Giudicarie Belt southwest of the indenter salient with the Salzach-Ennstal-Mariazell-Puchberg Fault (SEMP) in the northeast. This transpressive zone has a sigmoidal vortex geometry, decoupling the Ötztal Basement in the west from the extruding wedges of the Eastern Alps in the east. Shortening was successively absorbed by nappe stacking, upright folding, and dome formation and then maintained by transpressional shearing that led to lateral extrusion, probably promoted by a change in the lithospheric configuration. The western part of the SEMP likely rotated [?]20deg clockwise around a pole approximately halfway along its entire length during the indentation.

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Structural Nexus and Divide: the Western Tauern Window, Eastern Alps

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Key Points:

- The western Tauern Window forms a mature transpressive wrench zone with a sigmoidal vortex geometry
- The western Tauern Window links the Giudicarie Belt to the SEMP Fault, translating orogen-perpendicular shortening into -parallel extension
- The western Tauern Window decouples the Ötztal Basement in the west from the extruding wedges of the Eastern Alps in the east

Abstract

The western Tauern Window of the Eastern Alps constitutes a mature transpressive wrench zone generated by Oligocene-Miocene convergence between the Dolomites Indenter and the European foreland. As a major structural divide, it translates convergence-parallel north-south shortening into convergence-perpendicular eastward lateral extrusion. About 7000 foliation, lineation, fold-axis, axial-plane, shear-zone, and shear-band measurements at ~1800 sites detail the structural geometry and evolution during indentation and transpression, modeled by numerical and analogue-material experiments. The results outline the western Tauern Window as an orogen-scale zone of localized deformation, consisting of tight, upright folds, reactivated by a sinistral shear-zone network. It connects the Giudicarie Belt southwest of the indenter salient with the Salzach-Ennstal-Mariazell-Puchberg Fault (SEMP) in the northeast. This transpressive zone has a sigmoidal vortex geometry, decoupling the Ötztal Basement in the west from the extruding wedges of the Eastern Alps in the east.

Shortening was successively absorbed by nappe stacking, upright folding, and dome formation and then maintained by transpressional shearing that led to lateral extrusion, probably promoted by a change in the lithospheric configuration. The western part of the SEMP likely rotated 20° clockwise around a pole approximately halfway along its entire length during the indentation.

Plain Language Summary

The Alps constitute one of Earth's best-studied continent-continent collision orogens. There, the Apulian plate, a continental sliver of the African plate, started to collide with the European plate ~66 million years ago. In the north, the Apulian plate has a triangular shape, the Dolomites Indenter. The Tauern Window at the tip of the Dolomites Indenter deformed stronger than its neighboring areas, exposing deeply-buried European-plate rocks that are surrounded by Apulian-plate upper crustal rocks. We studied the deformation within the western Tauern Window and found that it connects two major fault zones, i.e., the Giudicarie Belt in the southwest that accommodates orogen-perpendicular shortening, and the SEMP Fault in the northeast that accompanies orogen-parallel extension. Based on kinematic models and field data, the western SEMP Fault likely rotated clockwise 20°.

1 Introduction

Salients or oroclines (e.g., Bandar Abbas, Pamir-Hindu Kush, Hazare-Kashmir, Namche Barwa, Nanga Parbat), where promontories advance into weak crust, are studied to deduce the kinematics of indenter tectonics and the associated displacements in the lithosphere (e.g., Critelli & Garzanti, 1994; Kufner et al., 2016; Molinario et al., 2004). Commonly observed processes are strain partitioning and localization, decoupling between tectonic blocks, deep-crustal exhumation, transpression, transtension, and slab breakoff (e.g., Chopin et al., 2012; Dewey et al., 1998; Fossen et al., 2013; Frehner, 2016; Kufner et al., 2016; Reiter et al., 2011; Reiter et al., 2018; Robin & Cruden, 1994). The structural evolution at indenter fronts has been modeled by numerical and analogue-material experiments, but controversies on the boundary conditions, reference systems and directions, and displacement amounts and rates in orogens hinder the transfer of their results to the natural systems (Burg & Podladchikov, 1999; Ratschbacher et al., 1991a, 1991b; Reiter et al., 2011; Rosenberg et al., 2004, 2007; Schueller & Davy, 2008). In addition, pre-existing anisotropies, common in orogens, may induce a structural evolution that deviates from the predicted one. Another difficulty is the quantification of shortening at the indenter fronts. Shortening in transpressional zones is often overestimated, as the wrench component of deformation remains underconsidered due to the lateral material movements out of the cross sections used to quantify the shortening (Tikoff & Peterson, 1998). To overcome this problem, we herein derive the kinematic vorticity number W_k to estimate the amounts of convergence-parallel shortening (contraction) and fold-hinge-parallel extension (stretching) based on the convergence angle γ , mea-

sured between the convergence direction and the active shear plane (Supporting Information Figure S1); this reflects the ratio between the simple and pure shear components of the bulk deformation (Fossen & Tikoff, 1993; Tikoff & Fossen, 1993).

We studied the western Tauern Window (TW) in front of the Dolomites Indenter in the European Alps (Figure 1a). The structural maps outline areas of dominate transpression and subordinate transtension, for example outlined by non-cylindrical folds and transecting foliations. In the field, we took foliation, lineation, fold axis, fold axial plane, shear zone, and shear band measurements. Based on their orientation, we defined five structural domains (A to E). We focus on the spatial distribution and relative temporal evolution of these structures and calculate the dihedral angles (section 4 Methods and section 5 Results) between the mean values of two distinct structures that formed in kinematic succession. We use the rotational direction of the dihedral angles (being clockwise or counterclockwise) and the kinematic vorticity number W_k to (1) deduce the kinematic connections between major fault systems, (2) discriminate between wrench-dominated and pure-shear-dominated transpressional domains, and (3) quantify the amounts of shortening and extension within the transpressional folds of the western TW. We review the existing hypotheses on the structural development of the TW and, based on our findings, modify them regarding the role of indentation, lateral extrusion, and fault linkage and rotation. Finally, we discuss the coupling between major fault zones, the decoupling of tectonic blocks, and complement the ongoing discussion about fault linkage and absolute displacements in the western TW by developing the Proto-SEMP hypothesis.

2 Geological Setting

The TW (Figure 1a) is the largest tectonic window of the Eastern Alps (Frisch, 1976), where lower plate rocks have been exhumed from >35-km-depth in the Cenozoic (e.g., Kurz et al., 2008; Selverstone & Spear, 1985; Spear & Franz, 1986). In this east-west elongate structural and metamorphic dome, the arcuate dome axis coincides with rocks recording the peak temperatures of high-grade Barrovian metamorphism (Grundmann & Morteani, 1985; Rosenberg et al., 2018; Selverstone, 1985; Schneider et al., 2015) and the youngest cooling ages (Bertrand et al., 2017; Luth & Willingshofer, 2008; Most, 2003; Rosenberg & Berger, 2009). We subdivided the TW into the eastern, central, and western sub-domes; the eastern and western sub-domes are characterized by elongate and concentric metamorphic isograd trajectories (Grundmann & Morteani, 1985; Scharf et al., 2013b) and intersecting and bordering shear zones (Figure 1a; e.g., Linzer et al., 2002; Neubauer et al., 1999; Rosenberg et al., 2004; Scharf et al., 2013a; Schneider et al., 2013). Framing shear zones are absent in the central dome; there, only structurally high nappes are exposed (e.g., Bousquet et al., 2012a; Schmid et al., 2013), partly folded into a map-scale sheet fold (Groß et al., 2020).

Two end-member models are debated for the formation of the TW. One empha-

sizes orogen-parallel, large-scale extension, exhuming the rocks of the TW along low-angle normal faults (e.g., Frisch et al., 2000; Neubauer et al., 1999), the other upright folding associated with erosional denudation (e.g., Cornelius, 1940; Laubscher, 1990). Yet, most authors agree on the activity of two kinematically-coupled processes, namely shortening due to the northward displacement of the Dolomites Indenter and eastward lateral material transport, associated with orogen-parallel extension; Ratschbacher et al. (1991a, 1991b) defined these as lateral extrusion, encompassing extension collapse, i.e., gravitational spreading of thickened and uplifted crust in an orogen, and tectonic escape. However, the amounts of shortening, displacement, and extension remain controversial (Frisch et al., 2000; Fügenschuh et al., 2012; Linzer et al., 2002; Luth et al., 2013; Neubauer et al., 1999; Pomella et al., 2011, 2012; Rosenberg & Berger, 2009; Rosenberg & Garcia, 2011, 2012; Scharf et al., 2013a; Schmid et al., 2013; Wolff et al., 2020). We contribute to this discussion with a compilation of structural measurements in the western TW to distinguish structural domains, and to assess the traces and interconnections of major shear zones.

2.1 Tectonostratigraphic Framework

Mesozoic ocean-derived and ophiolite-bearing units of the Glockner Nappe and Matreier Zone and siliciclastic rocks of the European distal margin were imbricated during European continental margin subduction (Figure 1a). The deepest structural units of the TW consist of Europe-derived, late-Variscan intrusions and their para-autochthonous, pre-Variscan host rocks covered by post-Variscan metasediments (Figure 1a), forming the Venediger Duplex, an antiformal stack generated by north-vergent nappe emplacement during the Adria-Europe collision. Above the intervening ocean-derived units of the Glockner Nappe and Matreier Zone, the hanging, Adria-derived Austroalpine nappes include the Meran-Mauls Basement, Ötztal Basement, Northern Calcareous Alps, Niedere Tauern, Gurktal Alps, and the DAV Basement (Figure 1a). All these units were thrusted over the European continental-margin units with the units in the TW shortened due to nappe stacking (D_1) and upright folding (D_2) during the Alpine collision in the Eocene to Oligocene. This resulted in a thickened European crust (e.g., Frisch, 1976; Lammerer et al., 2008; Lammerer & Weger, 1998; Rosenberg et al., 2014, 2018; Schmid et al., 2013) with a declining thickness east of the TW (Spada et al., 2013). In the late stage of the orogenic evolution, in the Oligocene-Miocene, east-west extension and strike-slip motion (D_3) have accompanied shortening.

2.2 Structural Framework

The western TW consists of the (D_2) upright Ahorn-, Tuxer-, and Zillertaler antiforms that fold the Early Alpine (D_1) foliation of the Venediger Duplex (Figure 1a; Schmid et al., 2013; Lammerer & Weger, 1998). Estimates of the amount of shortening accommodated by these folds vary between ~60 km (Rosenberg et al., 2015), ~49 km (Rosenberg & Berger, 2009), and ~32 km (Schmid et al.,

2013), based on line-length balancing and assuming constant nappe thicknesses. The tight folds strike ENE, show sub-horizontal, doubly-plunging fold axes and sub-vertical axial planes; they are often associated with axial-plane foliations, and have up to \sim 17 km amplitudes (Figure 1b). In contrast, the upright folds in the central and eastern TW strike east to ESE, are open and flat-topped (box-type) with amplitudes of up to \sim 10 km, and lack axial-plane foliations (Figures 1c to 1d). These structural differences point to an abrupt along-strike change in the way shortening has been accommodated (Figure 2a).

The Brenner Fault at the western margin of the TW (Figure 1a) is subdivided in this study into the Brenner Mylonites (BM) and the Silltal Fault (SF, Figure 2b; Behrmann, 1988; Schmidgigg, 1953; Selverstone, 1988). The top-west, normal-slip Brenner Fault detaches the Ötztal Basement from the underlying Glockner nappe. The steep limbs of the antiforms and the cores of the tight synforms in the westernmost TW are overprinted by sinistral shear zones, striking sub-parallel to their axial planes. From north to south, the map-scale sinistral shear zones are the Ahorn (ASZ), Greiner (GSZ), and Ahrntal (AhSZ) Shear Zones, each \sim 1 to 3-km-wide and tens of kilometers long (Figures 1a and 2b). In addition, several 10s to 100s of meters wide sinistral shear zones, e.g., the Tuxer (TSZ) and the Olperer (OSZ) Shear Zones, occur.

2.3 Surrounding Fault Zones

2.3.1 Giudicarie Belt The sinistral transpressive Giudicarie Belt (GB, Figure 1a) forms the western margin of the Dolomites Indenter, a \sim 77-km-long, N30°E-striking fault system, that offsets the ESE-striking Periadriatic Fault System. Displacement estimates along the GB vary between 15 to 30 km (e.g., Müller et al., 2001; Picotti et al., 1995; Viola et al., 2001) and \sim 70 km (e.g., Laubscher, 1988; Ratschbacher et al., 1991b; Werling, 1992) due to differing assumptions on its pre-indentation orientation and the amount of shortening accommodated at its southwestern tail. The former group of authors assumes an originally bent Periadriatic Fault System and therefore argues for moderate displacement along the GB, whereas the latter group assumes a straight, pre-indentation Periadriatic Fault System, requiring a larger displacement. Based on balanced cross-sections across the Southern Alps (Nussbaum, 2000; Schönborn, 1992), we favor an originally straight Periadriatic Fault System, hence a sinistral displacement of \sim 70 km along the GB. The GB includes the Pejo (PeF) and the Northern (NGF) and Southern Giudicarie Faults (SGF), and terminates in the north in the Jaufen (JF), Meran-Mauls (MMF), and Passeier Faults (PaF); the latter delimit, together with the southwestern corner of the TW, the N55°E-striking Meran-Mauls Basement (Figure 1a) (Pomella et al., 2011, 2012). The PaF cuts the JF and the MMF (Müller et al., 2001; Pomella et al., 2011). The thrust geometries inside the Dolomites Indenter exhibit an arcuate shape, indicating north-south (\pm 10°) shortening when applying the bow-and-arrow rule (e.g., Elliot, 1976) to its major thrusts. GNSS measurements of the present-day Alpine kinematics indicate a counterclockwise rotation of Adria

against Europe (Caporali & Martin, 2000; Le Breton et al., 2017).

Seismic tomography outlined two lithospheric slabs whose lateral terminations spatially coincide with the GB. The eastern slab was interpreted as the Adriatic plate subducted northward down to 250 km depth; the western slab, likely part of the European plate, features incipient slab break-off (Kästle et al., 2020; Lippitsch et al., 2003; Mitterbauer et al., 2011; Zhao et al., 2016). This discontinuity in the lithospheric structure coincides with a change in seismic anisotropy (Bokelmann et al., 2013; Hetényi et al., 2018; Qorbani et al., 2015). Earthquakes with a magnitude M_L 5 (Richter scale) and hypocentral depths of 20 km have strike-slip focal mechanisms along the GB (Reiter et al., 2018).

2.3.2 Salzach-Ennstal-Mariazell-Puchberg Fault The sinistral transpressive (west) to transtensive (east) Salzach-Ennstal-Mariazell-Puchberg Fault (SEMP, Figure 1a) strikes N70°E from Mittersill along the northern margin of the TW to the Vienna Basin (e.g., Cole et al., 2007; Frost et al., 2009; Linzer et al., 1995). Currently, the SEMP is a structural divide between the seismically active Northern Calcareous Alps in the north and the low-seismicity TW in the south (Reiter et al., 2018). Piercing points imply ~60 km displacement along the SEMP (Linzer et al., 2002), equivalent to ~56 km of east-west extension. Analogue-material models (Ratschbacher et al., 1991a; Rosenberg et al., 2004), simulating the indentation of the Dolomites Intender into the Eastern Alps, suggest that the sinistral faults along the northern boundary of the extruding wedges, e.g., the Proto-SEMP (chapter 6.5), initiated in the earliest stages of indentation and then rotated clockwise during progressive indentation. For the west-central SEMP, Wang and Neubauer (1998) inferred six evolution stages from paleostress analysis. Two involve predominantly strike-slip faulting and four are dominated by fault-normal compression; the maximum compressive stress axes rotated counterclockwise from NE to NW over time. Throughout the evolution of the SEMP, its northeastern tail has been confined to the southeastern salient of the European Platform, the southeastern tip of the Bohemian Massif (Ratschbacher et al., 1991b). At present, the southeastern tip of the Bohemian spur is located below the subparallel Mur-Mürz Fault (Figure 1a) at the Semmering Pass, close to Gloggnitz (16°00'East, 47°40'North; van Gelder et al., 2020). Geophysical data outline its subsurface extent (Grad et al., 2009; Baroň et al., 2019; Reinecker & Lenhardt, 1999) and suggest that the Bohemian spur is still acting as a buttress.

2.3.3 Pustertal-Gailtal Fault The dextral, N100°E-striking Pustertal-Gailtal Fault (PGF), a segment of the Periadriatic Fault System, divides the Eastern from the Southern Alps (Figure 1a, Polinski & Eisbacher, 1992) and defines the southern margin of the extruding wedges of the Eastern Alps (Ratschbacher et al., 1991b). Estimates of the Oligo-Miocene displacement along the PGF are indefinite. Frisch et al. (1998) and Laubscher (1988) argued for 100–150 km, pre-Miocene dextral displacement along the entire Periadriatic Fault System based on apparent map-view offsets and palinspastic

reconstructions. The dextral, NE-dipping Sprechenstein-Mauls Fault (SMF, Figure 2b), the youngest and westernmost, purely brittle segment of the PGF, offsets the western ductile tail of the PGF by ~2 km (Bistacchi et al., 2010). Earthquakes with focal mechanisms M_L 4 and hypocentral depths 10 km indicate dextral strike-slip along the PGF (Reiter et al., 2018).

2.3.4 Brenner Mylonites and Silltal Fault First estimates of east-west extension across the west-dipping BM varied between ~9 and 14 km (Behrmann, 1988) and “several to tens” of kilometers (Selverstone, 1988). Recent estimates are ~70 km (Fügenschuh et al., 1997), 4 to 12 km (Rosenberg & Garcia, 2011, 2012), 35 ± 10 km (Wolff et al., 2020), and ~44 km (Fügenschuh et al., 2012). The throw across the BM varies continuously along strike (Rosenberg & Garcia, 2011, 2012), with a maximum of ~17 km in the area of the Brenner Pass along the hinge of the Tuxer Antiform (Figure 1a); there, the BM are thickest (Axen et al., 1995; Behrmann, 1988; Selverstone, 1988). Towards the northern and southern tails of the BM and SF, the throw decreases to ~5 km, which cannot be explained by folding of the footwall (Rosenberg & Garcia, 2011). Reassessment (Fügenschuh et al., 2012; Rosenberg & Garcia, 2011, 2012) showed that extension was overestimated because exhumation was entirely attributed to normal faulting, ignoring the contributions of folding and erosion (Axen et al., 1995; Fügenschuh et al., 1997; Selverstone et al., 1995; Wolff et al., 2020). Following structural arguments (Rosenberg & Garcia, 2011, 2012) and using the average 29° dip of the BM obtained in this study (chapter 5.7.5 Extensional Shear Zones), we calculated 10 km extension across the BM, in agreement with the first estimates (Behrmann, 1988; Selverstone, 1988).

The moderately west-dipping SF strikes ~30 km from Innsbruck to Sterzing (Figure 2b and 2c). North of Steinach (Figure 2c), syn- and antithetic normal faults occur in the hanging wall of the SF (Figures 1a and 2b; Reiser, 2010; Reiter et al., 2018). The seismically active Brenner-Inntal Transfer zone connects the SF with the Inntal Fault System (Figure 1a) with transtensive and transpressive faults in the hanging and footwall of the SF, respectively (Ortner et al., 2006). The hypocentral depths of the M_L 4 earthquakes are between 5 and 15 km (Reiter et al., 2018).

2.3.5 Tauern-Northern-Boundary Fault

The Tauern-Northern-Boundary Fault (TNBF) between the Glockner Nappe and the Lower Austroalpine nappes in the northwestern TW was interpreted to link the BM to the SEMP (Töchterle et al., 2011), to transform extensional into strike-slip displacements (Fügenschuh et al., 2012), and to have a throw of ~3 km, deduced from the TRANSALP seismic section (Lammerer et al., 2008). The TNBF is an outcrop-scale, 50-m-thick, poorly exposed fault zone with varying shear sense indicators, perhaps with dominant top-north reverse and minor sinistral displacements (Bergmeister, 2010; Töchterle et al., 2011).

3 Deformation Fabrics

In the following, we present new and published structural data from the western TW (Figure 2) (Angel & Staber, 1950; Becker, 1993; Exner, 1956, 1962a, 1962b, 1979, 1980, 1983, 1989; Frank & Pestal, 2008; Frank et al., 1987; Giese, 2004; Höck et al., 1994; Karl et al., 1979; Lichtenheld, 2004; Moser, 2006; Rocken-schaub & Nowotny, 2009, 2011). These data highlight a structural difference between the central and eastern and the western TW. Based on the mean orientation of the structures, we subdivided the western TW into five structural domains (Figure 2c, Domains A to E) and discriminated three deformation phases (D_1 to D_3), associated with foliations S_1 to S_3 .

S_1 foliations formed prior to the upright folds (D_2 ; Figure 2d). Although of possibly varying ages, S_1 are associated with D_1 nappe stacking and Eocene/Oligocene metamorphism (Bousquet et al., 2012b; Glodny et al., 2005; Nagel et al., 2013; Schmid et al., 2013). The widespread sub-vertical axial-plane foliations (S_{2a}) of the D_2 tight and upright folds strike ENE and cut the folded nappe contacts in the tight D_2 synforms of the western TW (Figure 2d and 2c). The micro- and meso-scale D_2 folds are disharmonic passive shear folds, with thickened hinges, thinned limbs (measured perpendicular to bedding), and Z- and S-shaped second-order folds. Along the margins of the western TW and in the Austroalpine nappes, flexural-slip folds dominate. Along the western margin of the TW, the anticline hinges plunge west, indicated by the circular strike of the nappe contacts with outward pointing bedding. There, the folded S_1 are cut by or form composite foliations with moderately west-dipping S_{2b} foliations, part of the normal-slip shear zones along the BF. S_{2a} and S_{2b} are spatially separated, occurring NE and NW of the indenter salient, respectively (S_{2a} and S_{2b} appear as solid and dashed trajectories in Figure 2d).

D_2 folds and S_2 foliations are overprinted by D_3 km- to meter-scale, subvertical strike-slip and extensional shear zones (Figure 2b). Major NNE- to ENE-striking, sinistral (C_{sin}) shear zones and minor ESE- to ENE-striking, dextral (C_{dex}) ones occur (Pennacchioni and Mancktelow, 2007; Bistacchi et al., 2010). The west-dipping extensional shear zones (C_{nor}) along the western margin of the TW (Figure 2b) formed contemporaneously with the strike-slip shear zones because of mutual overprinting relationships (Axen et al., 1995; Behrmann, 1988; Fügenschuh et al., 1997; Selverstone, 1988). All D_3 shear zones are mylonitic, often brittle-ductile, have higher mica content and smaller grain size than the host rocks, and show clear shear-sense indicators (e.g., Mancktelow & Pennacchioni, 2005; Pennacchioni & Mancktelow, 2007; Schneider et al., 2013; Selverstone, 1993).

4 Methods

We first summarize the geometry, kinematics, and metamorphism associated with the major shear zones of the western TW (Figure 2b). Most of these structures have been described (e.g., Behrmann, 1988; Behrmann & Frisch, 1990;

Lammerer and Weger 1998; Neubauer et al., 1999; Müller et al., 2001; Pomella et al., 2011; Reicherter et al., 1993; Schneider et al., 2013; Spiess, 1995; Steffen & Selverstone, 2006; Viola et al., 2001), but their extent and location differ in the literature. Therefore, we reassess the structures and the shear-zone pattern based on our new structural data and a literature compilation. Then, we present and discuss the field measurements and mean orientation values of specific structures shown in equal area, lower hemisphere stereoplots (created by Stereo32 version 0.9; Röller & Trepmann, 2008). When the data show a preferred orientation and number >15 , we contoured them using a cosine exponential equation. If lineations and foliations are shown together, we contoured only the foliations. Mean values were deduced graphically from the data maxima. The individual field measurements of the Domains A to E are listed in the Supporting Information Dataset S1. Rotation of these structures during transpression led to deactivation and replacement by new structures representing the prevailing deformation field during the progressive deformation. The rotational sense of the successively formed structures is opposed to the material paths. We calculated dihedral angles between the mean values of structures, which formed in the prevailing deformational field, and that we infer to have formed in kinematic succession. This illustrates and quantifies their transpressional character and yielded different rotational senses for Domains A to E during transpression (Borradaile, 1987; Treagus and Treagus, 1981). The algebraic sign is positive (blue arrows) or negative (purple arrows) when the rotation from one mean value to the succeeding one is clockwise or counterclockwise, respectively (Figure S1).

5 Results and Interpretation

5.1 Ahorn Shear Zone

The ~2.5-km-wide and ~50-km-long, ~ENE-striking Ahorn Shear Zone (ASZ) is a transpressive mylonitic belt located along the northern limb of the western TW (Rosenberg & Schneider, 2008), extending the SEMP westward into deeper crustal levels (Figure 2b; Cole et al., 2007). The ASZ shows sinistral-oblique, south-side-up kinematics with a sub-vertical to steeply SSE-dipping mylonitic foliation and a shallowly west-plunging stretching lineation (Rosenberg & Schneider, 2008). The ASZ transitions into the SEMP in the area of Mittersill and Krimml where sinistral, brittle-ductile and ductile shear zones strike ~NNE and turn abruptly into the east-striking SEMP (Figure 2b and 2c; Cole et al., 2007; Reicherter et al., 1993). The metamorphic grade of the mylonites increases westward along-strike and southward perpendicular to the ASZ strike from 300°C to 400–500°C based on the quartz and feldspar recrystallization mechanisms and the stability of biotite (Schneider et al., 2013).

5.2 Tuxer Shear Zones

Lammerer & Weger (1998) described two sinistral shear zones of >40 km length within the Tuxer Antiform (Salzach-Riffler Fault), striking ESE from the Olperer

and Riffel summits in the west to Krimml in the east (Figure 2b and 2c). At Lammerer & Weger's (1998) locations, we mapped an interconnected network of outcrop-scale, 1- to 10-m-thick, lower amphibolite-facies, sinistral shear zones, the Tuxer Shear Zones (TSZ, Figure 2b). The TSZ strike NE to east, dip sub-vertically to steeply south, and connect to the ASZ and the Greiner Shear Zone (see chapter 5.4) at their northeastern and southwestern tails, respectively (Figure 2b).

5.3 Olperer Shear Zones

We divided Lammerer & Weger's (1998) shear zones in the Tuxer Antiform into the TSZ (section 5.2) and the Olperer Shear Zones (OSZ) because of their different orientations and kinematics. In agreement with Ebner et al. (2004), the OSZ constitute two shear zones of 10 to 15 km length, NE to north strike, and moderate to sub-vertical NW to west dip. Sinistral and top-NW normal shear occurred within these ~300-m-thick zones at their western tails. Feldspar ductility and biotite stability indicate syn-kinematic flow at amphibolite-facies conditions.

5.4 Greiner Shear Zone

The ~1-km-thick and ~50-km-long Greiner Shear Zone (GSZ) cuts all D₁ nappe contacts of the western TW and attains a maximum thickness in the Glockner Nappe (Figure 2b and 2e; Selverstone, 1993; Steffen et al., 2001). It strikes east to ENE, is sub-vertical, and has a sub-horizontal stretching lineation. Kinematic indicators imply sinistral (Behrmann & Frisch, 1990; De Vecchi & Baggio, 1982) and dextral (Barnes et al., 2004) shear. Metamorphism in the GSZ mylonites increases eastward along strike from greenschist to amphibolite facies, controlled by the west-plunging Greiner Synform (Figure 1a; e.g., Selverstone, 1985; Selverstone & Spear, 1985).

5.5 Ahrntal Shear Zone

The ~2-km-wide and ~70-km-long transpressive Ahrntal Shear Zone (AhSZ) at the southern limb of the western TW strikes NE to ENE, dips sub-vertically to steeply NNW, and shows sinistral and minor dextral and north-side-up shear (Figure 2b; Behrmann & Frisch, 1990; Reicherter et al., 1993; Schneider et al., 2009). The AhSZ cuts all D₁ nappe contacts and shows predominantly shallowly east-plunging stretching lineations. Metamorphism increases eastwards along strike from greenschist to lower amphibolite facies and northward, perpendicular to the strike, from 300°C to 400–500°C based on quartz and feldspar recrystallization mechanisms, biotite stability, and epidote composition (Kitzig, 2010; Reicherter et al., 1993; Wollnik, 2012). Hence, the metamorphic trends are complementary to the ASZ. Key outcrops in the Vals Valley, south of the Grabspitze summit (Figure 2c), show upright, open to tight folds in the north progressing into isoclinal and rootless folds with axial-plane foliations in the south, overprinted by sinistral shear zones (Figure 3a to 3d).

5.6 Jaufen Fault and Meran-Mauls Basement

The ~200- to 300-m-wide, >15-km-long JF strikes ~ENE; its ductile-brittle mylonites formed at lower greenschist-facies conditions (Müller et al., 2001; Spiess, 1995, 2001; Viola et al., 2001). The JF splits into two branches, one with top-west normal shear in the area of Gasteig (Viola et al., 2001), and one with sinistral-transpressive shear at the Penserjoch road NW of Elzenbaum (Figures 2b and 4; Müller et al., 2001). Our structural data, collected in the Meran-Mauls Basement and summarized in chapter 5.7 as Domain C, indicate that the upright folds of the western TW can be traced into the Austroalpine nappes (Figures 2c to 2d and 4).

5.7 Structural Elements of the Western Tauern Window

5.7.1 S_1 Foliation In Domains A to C (Figure 2c), the folded S_1 strike ENE and dip moderately to steeply NNW but locally SSE (Figure 3a and 3b), indicating an upright to slightly south-vergent sub-dome structure. The sub-vertical -circle of S_1 in Domain A traces a sub-horizontal, WSW-trending fold axis, whereas the -circles of S_1 in Domains B and C indicate gently ENE- and WSW-plunging fold axes, respectively (Figure 5a to 5c; -circle: great circle that fits best the distribution of poles to a folded surface). S_1 and S_2 form a composite foliation along the steep limbs of the D_2 antiforms. A crenulation cleavage, comprising the folded S_1 cut by the axial-plane S_{2a} foliation (Figure 3a and 3b), dominates the fold hinges.

5.7.2 D_2 Fold Axes and Axial Planes The D_2 axial planes (AP_2) of Domains A and B strike ENE, whereas those of Domain C strike NE, sub-parallel to the northwestern margin of the Dolomites Indenter (Figure 5k to 5m). In a cylindrical fold, the -circle and the mean axial plane should be orthogonal to each other and the -axis should be the fold axis (Ramsay & Huber, 1987, p. 334). However, the values for $_1$ (mean AP_2 – -axis; i.e., angle between the strike of the mean D_2 axial planes and the pole to the -circle) in Domains A and B show a -5° (counterclockwise) deviation from being orthogonal, whereas the $_1$ value in Domain C indicates a $+11^\circ$ (clockwise) deviation (Figures 5a to 5c, 5k to 5m, and S1a to S1c, S1g in the supporting information). The D_2 fold axes (FA_2) of Domain A are sub-horizontal and trend ENE, whereas the fold axes in Domains B and C plunge moderately ENE and SW, respectively (Figure 5a to 5c and 5f to 5h). Further, the $_2$ values (mean FA_2 – -axis) indicate a deviation of -8° and -6° in Domains B and A, in contrast to a deviation of $+8^\circ$ in Domain C (Figures 5f to 5h, 5k to 5m, and S1). The orientation of successively formed structures (deformation fields) during the D_2 folding of S_1 , indicated by $_1$ and $_2$, thus rotated counterclockwise in Domains A and B and clockwise in Domain C; the corresponding material paths rotated opposite to the deformation field, i.e., counterclockwise in Domain C and clockwise in Domains A and B (chapter 6.4).

The FA_2 – -axis deviations ($_1$ and $_2$) in Domains A, B, and C indicate non-

cylindrical folding in the western TW, which is diagnostic for transpressional or transtensional folding (Dewey et al., 1998; Ramsay & Huber, 1987; Tikoff & Peterson, 1998; Treagus & Treagus, 1981). The opposite rotations in Domains A and B respectively Domain C (Figures 5a to 5c, 5k to 5m, and S1) reflect the different strikes of the indenter margins causing dextral transpressive folds in domain C and sinistral ones in domains A and B (Figure S1). The different strikes are governed by the orientation of indenter margins and their angle to the convergence direction.

5.7.3 S_2 Foliation and L_2 Stretching Lineation In Domain A, the subvertical S_{2a} axial-plane foliations strike ENE, whereas in Domains B and C, they strike ENE–NE, i.e., slightly more to the north, and dip oppositely, SE and NW, respectively (Figure 6a to 6c). The angle $_3$ (mean AP₂–mean S_{2a}) describes the transection of the axial planes (AP₂) by the axial plane foliation (S_{2a}); S_{2a} deviates counterclockwise from AP₂ in Domains A and B and clockwise in Domain C (Figures 5k to 5m, 6a to 6c, and S1). Equally, the $_4$ values (mean FA₂–mean S_{2a}) in Domains A to C deviate clockwise in Domain C and counterclockwise in Domains A and B. According to Borradaile (1978), the $_3$ and $_4$ values obtained from cleavage-transected folds indicate dextral transpressive shear for Domains A and B, and sinistral one for Domain C. However, Treagus and Treagus (1981) showed that the relationship between cleavage-transected folds and transpressive shear sense rather depends on the initial orientation of the convergence direction to the shear zone boundary. Accordingly, the dihedral angles $_1$ to $_4$ indicate clockwise rotation of the deformation field in Domain C west of the indenter salient and counterclockwise rotation in Domains A and B east of the indenter salient during the same transpressive event. The dihedral angles $_5$ (-axis– S_{2a}) of Domains A to C are -6°, -4°, and -5°. The L_2 stretching lineations are sub-horizontal in Domain B (Figure 6b), sub-horizontal to gently WSW-, rarely ENE-plunging in Domain A, and mostly shallowly WSW-plunging in Domain C, with an increasing WSW-plunge along-strike from Domain B to C (Figure 6a to 6c).

5.7.4 Transcurrent Shear Zones In Domains A, B, and C, the values for $_6$ (mean S_{2a} –mean C_{\sin}) are -7°, -15°, and -18°, respectively; these are typical but especially for Domain A relatively small angles for sinistral S-C fabrics (Figures 6a to 6c, 6f to 6h, and S1; Passchier & Trouw, 2005). The values for $_7$ (mean C_{\sin} –mean C'_{\sin}) are -20°, -16°, and -18°, indicating typical C-C' angular relationships in all Domains (Figures 6f to 6h, 6k to 6m, and S1). The relatively small $_6$ angle in Domain A and the typical $_7$ angles suggest a mature shear zone system formed by prolonged activity. In Domain A, the sinistral shear zones (C_{\sin}) are sub-vertical and strike ~ENE; in Domain B and C, they strike NE with opposing dips, steeply SE and NW, respectively (Figure 6g to 6h). Stretching lineations (L_{\sin}) are sub-horizontal in Domain B, plunge shallowly SW in Domain A, and moderately SW to WNW in Domain C (Figure 6f to 6h). In Domains A, C, and D, conjugate dextral shear zones (C_{dex}) are associated

with variably-oriented stretching lineations (L_{dex} , Figure 7a, 7e, and 7i). The mean strike of C_{dex} changes from ENE to E to SE across the Domains C, A, and D, forming an arc around the indenter salient.

5.7.5 Extensional Shear Zones Extensional shear zones (C_{nor}) in Domains A and E show a large orientation scatter, but are, in general, sub-parallel to the S_{2b} foliation in Domain E (Figure 7c, 7m, and 7o). Based on the spatial distribution of S_{2b} (Figure 2d), the west-plunging FA_2 (Figures 2c and 5j), and the westward increasing plunge of L_2 (Figure 6e), we infer that the C_{nor} nucleated along the hinges and limbs of the west-plunging D_2 folds, ultimately forming the BM. In the areas of the OSZ and the JF (Figure 2b), outcrop-scale shear zones show sinistral-oblique, top-west normal kinematics, i.e., transtension, connecting the BM to the sinistral shear zones of the western TW. The extensional shear zones dip moderately NW in Domain A and gently NE in Domain C, where they are associated with shallowly NW-plunging stretching lineations (Figure 7c and 7g), parallel the Sprechenstein-Mauls Fault (Bistacchi et al., 2010).

5.7.6 Shear Bands Sinistral shear bands (C'_{sin}) in all five Domains (A to E) have a similar appearance, strike in average N34°E, and dip steeply to sub-vertically, indicating a homogeneous deformation field in and around the western TW (Figure 6k to 6o). Assuming that shear bands form at angles of 15–35° to the shear-zone boundary (Passchier and Trouw, 2005), the MMF (N55°E) with an angle of 21° to the average C'_{sin} strike may constitute the structure that is closest to the shear zone boundary during the C'_{sin} formation. The dextral shear bands (C'_{dex}) in Domains A, C, and D strike east; in Domain E, they strike ESE (Figure 7b, 7f, 7j, and 7l). Moderately west-dipping extensional shear bands (C'_{nor}) are common in Domains E and C, where they parallel the BM (Axen et al., 1995); they are rare in Domain A (Figure 7d, 7h, and 7n).

5.7.7 Structural elements in Domain D: Incipient Extruding Wedge Structural elements like the S_2 foliation (S_{2a} and S_{2b}) and the D_3 shear zones are less developed in Domain D than in Domains A to C. The folded S_1 predominantly dip gently to steeply south, i.e., opposite to those in Domains A to C (Figure 5a to 5d). The sub-horizontal FA_2 trend east (Figure 5d and 5i) and the sub-vertical AP_2 strike east (Figure 5n). Towards the west, the D_2 folds tighten and axial-plane foliations become frequent. The rarely observed, sub-vertical S_{2a} strikes east and the related L_2 lineation plunges gently east (Figure 6d), i.e., opposite to Domain C (Figure 6c). The sub-vertical sinistral shear zones (C_{sin}) strike ENE, show gently west-plunging L_{sin} (Figure 6i); the sub-vertical dextral shear zones (C_{dex}) strike east (Figure 7j). Both the C_{sin} and C_{dex} are more eastward oriented than those in Domains A to C.

5.7.8 Structural elements in Domain E: Western Margin of the Tauern Window The major difference between Domain E and the other Domains is

that none of its structural elements is sub-vertical. The folded S_1 dip mainly NNE, indicating south-verging folds (Brandner et al., 2008; Rosenberg & Garcia, 2011); the FA_2 plunge gently west, and the AP_2 dip moderately north (Figure 5e, 5j, and 5o), similar to the D_2 folds in Domain D (Figure 5i and 5n). The S_{2a} axial-plane foliations strike ENE and mainly dip moderately NNW; the associated stretching lineations plunge gently west (Figure 6e). In Domain E and along its transition to Domain A, the shallowly WNW-dipping S_{2b} dominate the structural grain (Figures 5e and 7m o), following the west-plunging hinge of the Tuxer Antiform. We observed only few NE-striking sinistral and ESE-striking dextral shear zones, but numerous, top-west normal shear zones, paralleling the BM, with down-dip stretching lineations (Figures 6j, 7k, and 7m) (e.g., Axen et al., 1995). In our field study, we could not find convincing evidence for the existence of the TNBF (Bergmeister, 2010; Fügenschuh et al., 2012; Töchterle et al., 2011).

A key outcrop in St. Jodok (Figure 2c) exhibits NNW-dipping normal shear zones (C_{nor}) associated with D_3 folds with gently NNW-dipping axial planes (AP_3) (Figure 8a to 8c). The D_3 folds refold the tight, south-verging D_2 folds and are accompanied by subvertical, NE-striking tension gashes (Figure 8d). These relationships indicate that the normal shear zones along the western margin of the TW post-date the D_2 folds. Additionally, the C_{nor} and AP_3 follow the arcuate strike of the folded S_1 along the northern limb of the Tuxer Antiform, delimiting the BM to the north, and connecting them to the ENE-striking sinistral-transtensive OSZ. To the south, the BM are delimited by the sinistral-transtensive JF.

5.8 Estimates of the Amounts of Shortening, Stretching, and Sinistral Strike-Slip

A critical parameter for estimating the amounts of shortening, stretching, and sinistral strike-slip is the movement direction of the Dolomites Indenter; the latter has been assumed to be NNW to N (e.g., Laubscher, 1988; Scharf et al., 2013; Schmid et al., 2013). Based on the analogue-material modeling results of Rosenberg et al., (2007) and the interpretation of the PaF, NGF, and SGF as strike-slip faults formed in a wrench-dominated transpressional regime, we argue for a convergence direction of N20°E; we include a ±10° uncertainty in our calculations (Figure S3). In section 6.3 (Transpressional Hypothesis), we detail the geometry of the western TW as transpressive zone and anticipate the structural orientation of its center, Domain A, as being representative for this zone. The angle between the mean strike (N60°E) of the sinistral shear zones (C_{sin}) in domain A and the inferred N20°E convergence direction is 40°. This indicates pure-shear-dominated transpression with a kinematic vorticity number W_k of 0.22, estimated using the graph proposed by Tikoff & Peterson (1998, Figure S3). Hence, 22% of the N30°E-displacement of ~70 km along the GF, which caused the shortening in the TW in front of the Dolomites Indenter (amount of contraction) (Laubscher, 1988), can be attributed to the hinge

parallel stretch and 78% to folding (true amount of contraction). According to Tikoff & Peterson (1998) and using the $\pm 10^\circ$ uncertainty in the convergence direction, we calculated $55 +6/-9$ km of shortening accommodated by folds and $15 +9/-6$ km hinge parallel stretch (Figure S3). Our result for shortening is larger than the estimate of ~ 32 km (Schmid et al., 2013) but similar to the estimates of ~ 49 km (Rosenberg and Berger, 2009) and ~ 60 km (Rosenberg et al., 2015) based on balanced cross sections. According to Tikoff & Peterson (1998), the amount of hinge parallel stretch equals the wrench component in transpressive systems. Therefore, sinistral strike-slip together with the hinge parallel stretch in the western TW is $15 +9/-6$ km, together contributing $13 +8/-5$ km to E-W extension.

6 Discussion

In the following, we discuss three hypotheses for the localized exhumation in the western TW. We also discuss the rotation of the Proto-SEMP during transpression, and wrench zone formation across the Eastern Alps.

6.1 Transtension Hypothesis

The TW was interpreted as a pull-apart structure in a transtensional setting (e.g., Genser & Neubauer, 1989; Neubauer et al., 1999). Strain modeling shows that upright open folds can form during transtension (Fossen et al., 2013). Such experiments indicate that the axial planes of transtensional folds strike perpendicular or at a high angle to the extension direction (Rey et al., 2011), and rotate toward the divergence vector (\vec{v}), that is the vector sum of the simple and pure shear components in transtensive (and transpressive) strain fields (Fossen et al., 2013). In sinistral transtension, the orientation of the divergence vector, approximated by the transtensional fold axes, has a clockwise deviation from the shear zone boundary (Fossen et al., 2013, their Figure 2a).

We used the mean strike ($N60^\circ E$) of C_{\sin} in Domain A as approximating the shear zone boundary during the D_2 and D_3 deformation. The mean trend ($N57^\circ E$) of the FA_2 fold axes is sub-parallel to the shear zone boundary in Domain C (Figure 5h), and deviates 14° and 8° clockwise in Domains A and B, respectively (Figures 5f and 5g). These fold axes–shear zone boundary relationships exclude the formation of D_2 folds as transtensional folds (Fossen et al., 2013). In addition, the geometry of transtensional folds is different from those observed in the western TW, where D_2 fold hinges are sub-parallel, not sub-perpendicular to the extension direction. The axial planes of D_2 folds in the western TW strike at high angles to the BM and are therefore kinematically unsuitable to accommodate large amounts of east-west extension. Only small-scale D_3 folds in Domain E have an appropriate orientation to accommodate east-west extension and could be interpreted as collapse folds (Froitzheim, 1992; Fossen et al., 2013).

6.2 Extensional Unroofing Hypothesis

Exhumation of the western TW by extensional unroofing requires that extension along the BM and SF are transferred to transcurrent sinistral and dextral strike-slip faults at the northern and southern tails of the Brenner Fault, respectively (Fügenschuh et al., 1997; Scharf et al., 2013a; Schmid et al., 2013). Our fieldwork shows that neither a sinistral strike-slip fault at the northern boundary of the western TW nor a km-scale, dextral strike-slip fault at the southern tail of the BM exists. The dextral PGF nucleates at the salient of the Dolomites Indenter that is located ~10 km ESE of the Brenner Fault. We traced the westernmost occurrence of the dextral PGF to the Valler Jöchl that is ~13 km away from the Brenner Fault. In spite of the widespread occurrence of sinistral shear zones throughout the western TW and the Meran-Mauls Basement (Figures 2b and 6f to 6i), Domain E lacks such structures (Figures 2b and 6j). The TNBF is a small-scale structure (section 2.3.5). The northwestern margin of the TW is characterized by south-verging, asymmetric F_2 folds (Figures 2c, 5e, 5j, and 5o) and the folded S_1 of the northern limbs are reactivated/overprinted by S_{2a} and S_{2b} (Figures 6e and 7o). The brittle SMF that possibly connects to the BM at the southwestern margin of the TW (Bistacchi et al., 2010), has no ductile precursor. Hence, it never was part of the pre-indentation Periadriatic Fault System. We predominantly found sinistral and normal-sense shear indicators in the field. Dextral shear indicators along the Meran-Mauls Fault document the pre-existence of dextral strike-slip along the Periadriatic Fault System.

6.3 Transpression Hypothesis

We interpret the large sinistral shear zone network within the western TW as a first-order structure linking the GB to the SEMP, transferring indentation into lateral extrusion. The mean orientation of the structural elements within Domains A to C show four along-strike trends from SW to NE in the western TW (Figures 2b to 2d). First, the S_{2a} foliations and C_{\sin} strike NE at the lateral terminations of the western TW (Domains B and C) and ENE in its center (Domain A, Figures 6a to 6c, 6f to 6h, and 9a). Second, these foliations and shear zones show a steepening of the average dip from the lateral terminations (Domains B and C) to the center (Domain A), and third, an along-strike flip in dip direction (Figures 6a to 6c, 6f to 6h, and 9a). Fourth, the fold axes are doubly plunging towards the lateral terminations of the western TW (Domains B and C) and sub-horizontal in its center (Domain A, Figures 5a to 5c, and 5f to 5h). This sigmoidal foliation pattern resembles that of a transpression zone in a sinistral strike-slip system, e.g., a restraining bend or a strike-slip duplex, in which Domain A is the transpression zone and Domains B and C approximate the bordering wrench zones (Dewey et al., 1998; Sanderson & Marchini, 1984; Woodcock & Fischer, 1986). The vortex of the foliation is located in the central Domain A (Figure 9b), as described for transpressive systems (Robin & Cruden, 1994).

Generally, the stretching lineations are shallowly plunging but scatter between

sub-horizontal and steeply westward plunging along the axis of the western TW. The variable bearing and plunge agree with numerical model results of stretching lineations formed in transpression (Robin & Cruden, 1994). Further, the fold amplitudes decrease and the fold wavelengths narrows from the central Domain A towards the outer Domains B and C, indicating non-cylindrical (conical) folds; such style is predicted by numerical models of transpressional folds (Frehner, 2016).

The dihedral angles α_1 to α_4 in Domains A to C, which mainly describe the geometric evolution of the D_2 folds, indicate a deviation of the mean strike/trend of the different structures formed in kinematic succession that is consistently clockwise in Domain C and counterclockwise in Domains A and B (Figures S1a to c, S1g). This indicates counterclockwise material-plane rotations in Domain C and clockwise rotations in Domains A and B and can be explained by the orientations of the two margins of the Dolomites Indenter, parallel to the MMF and PGF, causing different deformation fields east and west of the indenter salient, e.g., bending the preexisting S_1 foliations. The dihedral angles α_5 to α_7 , which mainly describe the geometric evolution of the D_2 shear zones, indicate exclusively counterclockwise deviations during sinistral shearing in all three Domains (A to C; Figures S1d to f, and S1g). The mean strike of the C'_{sin} , sub-parallel to the PaF, NGF and SGF, is constant in all domains, indicating a similar, late-stage deformation field related to the Dolomites Indenter (Figures 6k to 6o). As the sinistral shear zones (C_{sin}) in Domains A to C reactivate the AP_2 axial plane foliation, we argue that they are kinematically linked to the D_2 folds and can be interpreted as part of the transpressive zone (Woodcock & Fischer, 1986; Dewey et al., 1998). This transpressive zone initiated during non-cylindrical folding during indentation and progressed by the reactivation of the S_{22} as slip planes during sinistral transpression. The switch in the rotational sense of the structures from opposing directions during non-cylindrical folding in Domain C compared to Domains A and B into a uniform rotation sense during shearing in all three Domains supports the two-stage evolution within a constant deformation field.

6.4 Corner Effect, Partitioning, and Decoupling

The sinistral transpressive zone of the western TW established a structural connection between the GB and the SEMP with a kinematic continuity between NNE-ward indentation and eastward extrusion. Indentation induced a transpressive zone with upright folding and hinge-parallel stretching that transitioned into deformation governed by strike-slip shear zones. Outside the transpressive zone, conjugate sinistral and dextral strike-slip faults formed in Domain D along the NNE-facing indenter margin, and extensional shear zones formed in Domain E. Partitioning of NNE-SSW shortening caused a divergence of material flow away from the indenter salient within Domains D and E (cf. Ježek et al., 2002; Reiter et al., 2011).

Kinematically, the BM, JF, and OSZ (Figure 2b) decouple the Ötztal Basement

from the western TW by transtensional to extensional deformation west of the transpressive zone. East of the indenter salient conjugate shear zones, i.e., the sinistral AhSZ, DAV, SpSZ and the dextral PGF, Iseltal, and Mölltal faults (Figure 1a) and other small-scale, conjugate faults in Domain D decouple the eastward extruding wedge from the transpressive zone. Hence, the transpressive zone is not part of the eastward extruding wedge, instead it comprises a restraining-bend structural junction between the GB and the SEMP (Figure 9b).

Analog-material and numerical experiments showed that within sinistral transpressive zones, transpressive fold axes initiate parallel to the maximum instantaneous stretching axis and with increasing strain, rotate clockwise and stay parallel to the long axis of the strain ellipse (Frehner, 2016; Ratschbacher et al., 1991a, 1991b; Rosenberg et al., 2004, 2007). The mean orientation of fold axes in the western TW, especially in Domain A, form an angle of $\sim 44^\circ$ to the GB, suggesting—at first glance—a rather juvenile transpressive structure, where transpressional folds only initiated and started to rotate clockwise towards the indenter margin (Fossen et al., 2013; Frehner, 2016; James & Watkinson, 1994; Tikoff & Peterson, 1998; Treagus & Treagus, 1981). However, the western TW combines transpression due to the connection of two fault zones (GB and SEMP) and the indentation imposed by the Dolomites Indenter. Numerical indentation models (Ježek et al., 2002) show that the deformation pattern in front of an indenter changes little through time, supporting the interpretation of the western TW as a rather mature sinistral transpressive zone that has clear borders in Domains D and E and connects the GB to the SEMP (Ježek et al., 2002).

6.4.1 Decoupling to the West Rocks north of the GB (Figure 1a; Ötztal Basement) were shortened little because of two geometric aspects. First, the angle of 10 to 25° between the N 20° E convergence direction and the N 30° E strike of the PaF, SGF, and NGF caused a rather oblique deformation field west of the indenter margin compared to the nearly orthogonal shortening in the northern margin of the indenter (the PGF; Figure 9b). Second, the Dolomites Indenter experienced a net eastward movement, which promoted extension in front of the indenter salient. The FA₂ and AP₂ in Domain C strike NE, subparallel to the northwestern margin of the Dolomites Indenter, and rotated counterclockwise during ongoing shortening, indicated by consistently positive dihedral angles γ_1 to γ_4 (Figure S1). The FA₂ and AP₂ in Domains A and B strike ENE and rotated clockwise during ongoing shortening indicated by exclusively negative dihedral angles γ_1 to γ_4 (Figure S1). We argue that this opposing rotation of the material planes/lines NW and NE of the indenter salient caused additional, about convergence-perpendicular extension, necking the D₂ folds (Figure 9b, stereoplots of the D₂ folds). The S_{2b} foliation is restricted to the west-plunging dome axis and the BM (Domain E). Together with the westward increasing steepness of the L₂ stretching lineation within the transpressive zone, it likely is a structural expression of this local extension. The S_{2b} foliation pre-dates the top-west extensional shear zones/bands (C_{nor}, C'_{nor}), that detach and decouple

the Ötztal Basement from the transpressive zone. The OSZ (Ebner et al., 2004; Lammerer & Weger, 1998), structures observed at St. Jodok (Figure 8), and partly the JF (Müller et al., 2001; Viola et al., 2001), are transtensive, transferring sinistral displacement of the western TW into east-west extension along the BM (Figure 2b). Extensional shear zones within the TW in the area of Steinach (Figures 2b and 2c) dip NNW (Figure 8a), tracing the BM in a convex arc northeast-ward (Fügenschuh et al., 1997; Töchterle et al., 2011). Therefore, the ENE-striking normal shear zones in the area of St. Jodok separate the thick BM in the south from the brittle SF in the north (Behrmann, 1988; Fügenschuh et al., 1997; Reiser, 2010; Reiter et al., 2018; Schmidegg, 1953). Finally, the NNE-SSW shortening caused a component of eastward motion of the indenter, promoting maximum east-west extension in front of the indenter tip (Rosenberg et al., 2004, 2007). Hence, Domain E marks a decoupling zone between the shortened western TW and the Ötztal Basement (Figure 9b). The BM are coupled with and the SF together with syn- and antithetic faults in its hanging wall are decoupled from the transpressive zone of the western TW (Figure 9b). Thus, how to explain the extension along the SF? We speculate that it is part of a late-stage releasing band connecting the sinistral shear along the IF in the north with that along JF and PaF in the south; these would involve up to a few kms of late brittle normal faulting along the BF.

6.4.2 Decoupling to the East The AhSZ bounds the central TW in the west that lacks high-amplitude folds, subvertical foliations, and transcurrent shear zones (Figures 2a to 2e). Accordingly, the exhumation level is less deep than in the western TW and large areas expose rocks of the structurally high units. North of the PGF (Figure 1a), the convergence in Domain D was mainly accommodated by upright D_2 folds; their axial planes (AP_2) and the rarely developed S_{2a} have an angle of $\sim 20^\circ$ to the PGF and the latter fade out eastward (Figure 5n). The folds are overprinted by conjugate sinistral and dextral shear/fault zones implying NNE-SSW shortening and eastward material transfer away from the indenter salient (Figure 9b). Domain D, bound by the AhSZ in the north and the PGF in the south, forms the western tip of an eastward extruding wedge (Ratschbacher et al., 1991b).

The ENE-striking transpressive zone of the western TW (Domain A) turns into a NE-strike in Domain B. There, the number of D_2 folds is reduced to two major antiforms, indicating less localized shortening. Several NE-striking sinistral shear zones delimit the transpressive zone to the east from the central TW, where it enters the SEMP with a sudden change in strike (Figures 2b, 2d, 2e; Cole et al., 2007; Rosenberg & Schneider, 2008). We argue that Domain B marks the area where the distributed transpressive deformation expressed by non-cylindric folding and sinistral shearing of the western TW was transferred into more localized sinistral shear zones south of Mittersill (e.g., RSZ, Figure 2b; Cole et al., 2007) and ultimately to the brittle-ductile SEMP (Frost et al., 2011). The fault/shear zones south and east of the western TW that connect to the SEMP and the PGF decouple the extruding central and eastern TW

from the stationary western TW (Linzer et al., 2002; Peresson & Decker, 1997). Based on the following observations, we interpret that shortening inside the western TW is at maximum and gradual decreases eastward, and strike-slip motion and extension is minimum in the western TW and gradually increases eastward. First, the location of the western TW—at the tip of the indenter—accommodated maximum shortening and minimum lateral translation. Second, east of the western TW, shortening is localized along the SEMP and PGF, given the large displacements of both faults east of the western TW. Third, the SEMP and numerous conjugate strike-slip faults south and east of the TW nucleate east of the western TW. Forth, the SEMP shows a succession from transpression in the west to transtension in the east along strike.

6.5 Proto-SEMP Hypothesis

In accordance with the results of the analogue-material models (Ratschbacher et al., 1991a, 1991b; Rosenberg et al., 2007), we propose that the N70°E-striking SEMP had a more northerly initial strike and nucleated northeast of the indenter tip, i.e., it never was connected to the GB in the first place (Figures 9b and S4). We present two approaches to restore its orientation to the time of early indentation (Proto-SEMP). The first approach (Proto-SEMP_1) is a counterclockwise back-rotation of the present-day western SEMP around a vertical rotation pole close to Liezen (Figure 1a), using the 55-km NNE-SSW shortening and 13-km eastward extension suggested in this study for the western TW. In this restauration, the Proto-SEMP_1 had an initial N50°E strike, subparallel to the present-day strike of the Meran-Mauls Basement, and rotated 20° clockwise together with the sub-parallel IF and MMzF (Figures 1a and S4). Palinspastic restorations of Fodor (1995) and Linzer et al. (1995) also imply a N50°E strike of the Proto-SEMP. The Northern Calcareous Alps in between rotated counterclockwise along antithetic Riedel faults, deduced from paleomagnetic analyses of intramontane basins interpreted as domino-block rotation (Haubold et al., 1999; Márton et al., 2000).

The second approach uses the results of paleostress analysis along the western SEMP; these yielded six deformational events, two (D_3 and D_5) involved mostly strike-slip deformation, four mainly thrusting (Wang & Neubauer, 1999). Against the authors' interpretation, we kept the orientation of the principal stress axes constant and rotated the present-day SEMP 55° counterclockwise towards a Proto-SEMP_2 orientation of N15°E (Figure 9b), using the rotations derived by Wang & Neubauer (1999); we consider this rotation value an extreme. Both approaches support our hypothesis of a clockwise rotating SEMP during indentation under a stable convergence direction (Figures 9c and S4).

6.6 Wrench Zone Hypothesis

The TW consists of upright folds, delimited by conjugate transpressive shear zones (Ratschbacher et al., 1991a, b; Rosenberg et al., 2004; 2007), forming an arcuate culmination in front of the Dolomites Indenter. This arcuate dome of the

TW with a triangular crustal block south of it (Figure 1a) is consistent with the majority of indentation models (Schueler & Davy, 2008, for a review), which show localized shortening and conjugate shear in front of a triangle pointing away from the indenter margin. However, the western TW accumulated more deformation as the central and eastern TW, indicating a different evolution in the late stage. We defined the Domains A to C as a mature sinistral transpressive zone, connecting the initially separated GB and the SEMP via a restraining bend (Figure 9b). The western TW forms the western margin of the stable triangle in front of the Dolomites Indenter. In an early stage of indentation, the western TW developed as a right step-over between the independently existing GB and the Proto-SEMP with localized high-strain shortening accommodated by upright folds (Figure S4); it evolved over time into a zone dominated by sinistral wrenching. We argue that this transpressive zone together with the GB and the SEMP constitutes a sinistral wrench zone translating Dolomites indentation into lateral extrusion. By that, it decouples the Ötztal Basement in the west, experiencing normal shear/faulting along the Brenner Fault and sinistral transtension along the OSZ and the JF, and the central and eastern TW in the east, experiencing eastward extrusion along conjugate shear zones, e.g., the PGF, SEMP and minor strike-slip faults in between (Ratschbacher et al., 1991a, 1991b; Rosenberg et al., 2007).

7 Conclusions

The western Tauern Window shows a structural evolution distinct from the central and eastern Tauern Window. Tight, upright folds of high-amplitude, subvertical axial plane foliations, and sinistral shear zones indicate absorption of orogen-perpendicular shortening. The western Tauern Window constitutes a right step-over connecting the sinistral transpressive Giudicarie Belt, subparallel to the indentation direction of the Dolomites Indenter, with the sinistral SEMP Fault, a major fault zone related to the eastward lateral extrusion of the central and eastern Eastern Alps. The ~70 km indentation by the Dolomites Indenter has been manifested within the western Tauern Window by upright folds and sinistral shear zones, accommodating $55 +6/-9$ km of shortening, and $15 +9/-6$ km hinge parallel stretch and likewise $15 +9/-6$ km sinistral shear zones, resulting in $13 +7/-5$ km of east-west extension. Farther east, the SEMP Fault has accommodated ~56 km of eastward lateral extrusion. At the salient of the Dolomites Indenter, the western Tauern Window forms a relatively stationary structural divide, decoupling the western Eastern Alps along the top-west, normal-slip Brenner Fault from the central and eastern Eastern Alps that extrude eastward. In the central Tauern Window, north-south shortening decreases in comparison to the western Tauern Window, as seen from the preservation of higher tectonostratigraphic units, the lower amplitude of folds, and the lack of axial plane foliations. In turn, the amount of lateral extrusion increases, as indicated by the increasing abundance of conjugated strike-slip shear/fault zones within the Austroalpine nappes south and east of the Tauern Window. The structures and the inferred displacements within the western Tauern Win-

dow are consistent with exhumation dominated by erosional denudation during transpressive folding. We interpret the western Tauern Window as a transpressive, right step-over between the sinistral Giudicarie Belt and a sub-parallel Proto-SEMP, which was pinned by the West-European Platform at the southern salient of the Bohemian massif. During indentation, the western Proto-SEMP progressively rotated clockwise around a vertical axis close to Liezen, attaining its present day orientation.

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This data set contains 6965 structural data from 1774 outcrops in the western Tauern Window, collected during July to August, 2007, June to September, 2008, June to September, 2009, July to October, 2010, using a Breithaupt "GEKOM" Stratum Compass, provided by the Freie Universität Berlin, Germany. The raw data, plotted in the lower hemisphere and equal area plots of Figures 5a to 5o, 6a to 6o, and 7a to 7o are given in the Supporting Information Dataset S1. The Supplementary Information Figures 1a to g, 2, and 3 contains the different dihedral angles between the mean values of the respective structural elements of the Domains A to C. The georeferenced raw data are available in the open access research data base OpARA of the Technische Universität Dresden and Technische Universität Bergakademie Freiberg.

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10 Figure Captions

Figure 1a. Simplified tectonostratigraphic map of the Tauern Window and surrounding units modified after (Bousquet et al., 2012b). **b. to d.** Structural sections of the western, central, and eastern sub-dome of the Tauern Window simplified after Brander (2013) and Schmid et al. (2013).

Figure 2a. Foliation map of the Tauern Window shows the main foliation (undifferentiated). Compiled literature and new data (see paragraph 2 for references). **b. to e.** Close ups of the western Tauern Window based on (Bigi et al., 1994) showing **b.** Traces of the macro-scale sinistral shear zones; **c.** Outcrops of this study, definition of the five structural Domains (A to E), and traces of the structural sections (Figure 4); **d.** Interpreted traces of the S_{2a} axial-plane (solid lines) and the S_{2b} foliations (dashed lines); **e.** Interpreted traces of the F_2 upright folds and the tectonostratigraphic units.

Figure 3. Field photographs along the Vals valley from north to south. **a.** Folded S_1 upright, tight folds, Width of the photo is approximately 4 m; **b.** Rootless S_1 folds and newly formed sub-vertical S_{2a} axial-plane foliation; **c.** S_{2a} axial-plane foliation and localized sinistral shear zones (C_{\sin}), having an additional north-side-up component; **d.** Top view on sub-vertical S_{2a} axial-plane foliation and localized sinistral shear bands C'_{\sin} .

Figure 4. SE-trending structural cross section running from Rosskopf to Onsberg (see Figure 2a for trace). Lithological units based on the Geologische Karte 1:50.000, Blatt 175 Sterzing.

Figure 5. Stereoplots of contoured data in ten contouring intervals, number of data, maximum, minimum, mean densities (), cosine exponent, and mean values are given in the bottom-left columns. **a. to e.** S_1 foliation, great circles indicate -circles. **f. to j.** D_2 fold axes (FA_2). **k. to o.** D_2 axial planes (AP_2).

Figure 6. Stereoplots of contoured data in ten contouring intervals, number of data, maximum, minimum, mean densities (), cosine exponent, and mean values are given in the bottom-left columns and the lower right corner. **a. to e.** S_{2a} axial-plane foliations and L_2 lineations, great circles indicate S_{2a} mean values. **f. to j.** Sinistral shear zones (C_{\sin}) and stretching lineations (L_{\sin}), great circles indicate C_{\sin} mean values. **k. to o.** Sinistral shear bands (C'_{\sin}), great circles indicate C'_{\sin} mean values.

Figure 7. Stereoplots of contoured data in ten contouring intervals, number of data, maximum, minimum, mean densities (), cosine exponent, and mean values are given in the bottom-left columns and the lower right corners. **a., e., k., j.** Dextral shear zones (C_{dex}) and stretching lineations (L_{dex}), great circles indicate C_{dex} mean values. **b., f., l., i.** Dextral shear bands (C'_{dex}), great circles indicate C'_{dex} mean values. **c., g., m.** Extensional shear zones (C_{nor}) and stretching lineations (L_{nor}), great circles indicate C_{nor} mean values. **d., h., n.** Extensional shear bands (C'_{nor}), great circles indicate C'_{nor} mean values. **o.** Composite foliations (S_{2b}), great circles indicate S_{2b} mean value.

Figure 8. a. Sketch of a key outcrop near St. Jodok (outcrop length ~100 m) and stereoplots; from left to right: S_1 foliations, AP_2 axial planes, C_{nor} extensional shear zones, and AP_3 collapse folds. **b. to d.** South-verging, tight D_2 folds (F_2) are sheared by normal-sense shear zones along their limbs. Collapse folds (F_3) and tension gashes overprint steep F_2 folds.

Figure 9. a. Block model of the western Tauern Window. Structural map and sections are based on Brandner (2013) and Schmid et al. (2013). Map is displaced and view is parallel to the wrench zone (~NE) to illustrate the along strike variation of the sinistral shear zones (red lines). **b.** Structural summary of the wrench zone highlighted in red. Mean great circles and mean vectors based on data shown in Figures 5, 6, and 7. Sectors with upper-right-to-lower-left stripes indicate the assumed N20°E ($\pm 10^\circ$) convergence direction of the Dolomites Indenter with respect to the major faults forming the wrench zone (GF, MMF, and Proto-SEMP). Sectors with upper-left-to-lower-right stripes indicate the conver-

gence vector of the respective shear zone boundary (-20°), in which transpression is wrench-dominated. For the MMF and the western TW, the convergence angles are 35° and 40° , respectively, indicating pure-shear-dominated transpression (Tikoff and Peterson, 1998). For the GF, the Proto-SEMP₁, and the Proto-SEMP₂, the convergence vectors are towards N 10° E, N 30° E and N 15° E, respectively, indicating mostly wrench-dominated transpression (Tikoff and Peterson, 1998). **c.** Structural reinterpretation of the deformation stages of Wang and Neubauer (1998), left: Wang and Neubauer's (1998) interpretation of counterclockwise rotation of principle stress axes around a fixed SEMP, right: reinterpretation in this study assumes a fixed ($\sim N$) direction of principle stresses and a $\sim 55^\circ$ clockwise rotation of the SEMP into the Proto-SEMP orientation.

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain D			Domain A		
161	88	S ₂	263	53	C _{sin}	4	21	C' _{nor}	94	18	L ₂	318	32	AP ₃
175	73	S ₂	254	65	C _{sin}	293	53	C' _{nor}	107	37	L ₂	300	22	AP ₃
171	79	S ₂	304	80	C _{sin}	273	43	C' _{nor}	110	11	L ₂	314	29	AP ₃
163	45	S ₂	149	82	C _{sin}	276	45	C' _{nor}	126	12	L ₂	285	45	AP ₃
335	80	S ₂	121	83	C _{sin}	226	22	C' _{nor}	132	34	L ₂	288	33	AP ₃
152	50	S ₂	129	73	C _{sin}	244	28	C' _{nor}	135	28	L ₂	246	52	AP ₃
150	89	S ₂	307	77	C _{sin}	252	43	C' _{nor}	3	45	L ₂	272	45	AP ₃
147	77	S ₂	287	89	C _{sin}	334	80	C' _{nor}	281	19	L _{sin}	283	41	AP ₃
150	86	S ₂	304	40	C _{sin}	332	78	C' _{nor}	284	60	L _{sin}	92	60	AP ₃
146	71	S ₂	304	69	C _{sin}	270	57	C' _{nor}	235	34	L _{sin}	268	69	AP ₃
350	75	S ₂	271	76	C _{sin}	260	70	C' _{nor}	256	28	L _{sin}	253	65	AP ₃
337	63	S ₂	308	54	C _{sin}	208	40	C' _{nor}	251	34	L _{sin}	274	72	AP ₃
341	62	S ₂	306	59	C _{sin}	219	53	C' _{nor}				89	87	AP ₃
189	79	S ₂	302	76	C _{sin}	228	53	C' _{nor}				310	24	AP ₃
349	81	S ₂	295	65	C _{sin}	344	70	C' _{nor}				312	24	AP ₃
337	61	S ₂	304	54	C _{sin}	333	66	C' _{nor}				241	28	AP ₃
353	68	S ₂	320	66	C _{sin}	283	64	C' _{nor}				236	44	AP ₃
329	50	S ₂	295	65	C _{sin}	263	38	C' _{nor}				222	49	AP ₃
322	68	S ₂	311	64	C _{sin}	248	54	C' _{nor}				321	17	AP ₃
351	49	S ₂	328	62	C _{sin}	250	63	C' _{nor}				314	15	AP ₃
335	55	S ₂	345	85	C _{sin}	271	38	C' _{nor}				342	25	AP ₃
322	45	S ₂	156	81	C _{sin}	239	45	C' _{nor}				338	35	AP ₃
326	70	S ₂	140	76	C _{sin}	264	74	C' _{nor}				302	10	AP ₃
134	87	S ₂	348	52	C _{sin}	9	70	C' _{nor}				37	30	FA ₁
326	83	S ₂	348	42	C _{sin}	305	65	C' _{nor}				51	22	FA ₁
333	74	S ₂	297	80	C _{sin}	239	75	C' _{nor}				78	37	FA ₁
154	88	S ₂	132	83	C _{sin}	247	63	C' _{nor}				86	24	FA ₁
144	67	S ₂	143	80	C _{sin}	267	67	C' _{nor}				242	7	FA ₁
329	70	S ₂	309	87	C _{sin}	266	61	C' _{nor}				224	31	FA ₁
331	85	S ₂	324	85	C _{sin}	257	61	C' _{nor}				242	28	FA ₁
162	85	S ₂	153	89	C _{sin}	24	66	C' _{nor}				224	31	FA ₁
330	80	S ₂	156	90	C _{sin}	5	90	C' _{nor}				250	10	FA ₁
344	66	S ₂	146	85	C _{sin}	258	58	C' _{nor}				254	14	FA ₁
331	90	S ₂	149	84	C _{sin}	117	72	C' _{rev}				274	2	FA ₁
338	86	S ₂	340	81	C _{sin}	118	69	C' _{rev}				277	11	FA ₁
149	84	S ₂	328	82	C _{sin}	72	42	C' _{rev}				71	4	FA ₁
334	89	S ₂	336	86	C _{sin}	67	42	C' _{rev}				245	5	FA ₁
160	85	S ₂	342	85	C _{sin}	301	60	C' _{rev}				85	7	FA ₁
150	81	S ₂	113	70	C _{sin}	307	57	C' _{rev}				74	4	FA ₁
332	69	S ₂	100	80	C _{sin}	303	50	C' _{rev}				96	7	FA ₁
349	84	S ₂	333	85	C _{sin}	320	30	C' _{rev}				92	5	FA ₁
347	82	S ₂	327	88	C _{sin}	325	75	C' _{rev}				254	7	FA ₁
166	85	S ₂	286	55	C _{sin}	304	70	C' _{rev}				259	0	FA ₁
173	83	S ₂	328	82	C _{sin}	292	65	C' _{rev}				110	24	FA ₁
351	68	S ₂	331	90	C _{sin}	271	51	C' _{rev}				102	18	FA ₁

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
166	86	S ₂	154	73	C _{sin}	68	77	C' _{rev}	71	59	L ₂	100	15	FA ₁
350	88	S ₂	142	72	C _{sin}	58	64	C' _{rev}	74	69	L ₂	69	7	FA ₂
355	80	S ₂	332	76	C _{sin}	201	69	C' _{rev}	73	67	L ₂	251	3	FA ₂
329	71	S ₂	312	82	C _{sin}	224	77	C' _{rev}	70	24	L ₂	227	26	FA ₂
329	61	S ₂	313	84	C _{sin}	205	70	C' _{rev}	64	48	L ₂	240	14	FA ₂
328	72	S ₂	148	65	C _{sin}	184	73	C' _{rev}	72	41	L ₂	255	4	FA ₂
357	73	S ₂	139	58	C _{sin}	240	90	C' _{rev}	68	46	L ₂	249	2	FA ₂
355	77	S ₂	163	72	C _{sin}	306	88	C' _{rev}	84	45	L ₂	72	16	FA ₂
344	62	S ₂	166	65	C _{sin}	305	75	C' _{rev}	90	36	L ₂	84	24	FA ₂
340	78	S ₂	343	86	C _{sin}	294	42	C' _{rev}	90	34	L ₂	250	5	FA ₂
175	75	S ₂	158	80	C _{sin}	311	36	C' _{rev}	93	51	L ₂	78	6	FA ₂
183	85	S ₂	346	89	C _{sin}	333	55	C' _{rev}	79	55	L ₂	228	8	FA ₂
7	67	S ₂	167	88	C _{sin}	334	50	C' _{rev}	98	57	L ₂	81	12	FA ₂
2	59	S ₂	339	86	C _{sin}	27	65	C' _{rev}	275	45	L ₂	264	24	FA ₂
334	60	S ₂	345	86	C _{sin}	20	70	C' _{rev}	270	35	L ₂	254	30	FA ₂
348	64	S ₂	139	80	C _{sin}	18	67	C' _{sin}	256	23	L ₂	254	16	FA ₂
0	59	S ₂	148	70	C _{sin}	119	72	C' _{sin}	254	46	L ₂	256	19	FA ₂
332	87	S ₂	130	67	C _{sin}	114	71	C' _{sin}	262	62	L ₂	255	16	FA ₂
341	81	S ₂	146	55	C _{sin}	126	85	C' _{sin}	96	34	L ₂	93	29	FA ₂
332	80	S ₂	152	75	C _{sin}	113	77	C' _{sin}	68	39	L ₂	93	30	FA ₂
332	85	S ₂	155	87	C _{sin}	277	89	C' _{sin}	89	45	L ₂	251	5	FA ₂
0	70	S ₂	147	74	C _{sin}	298	85	C' _{sin}	96	47	L ₂	71	1	FA ₂
359	71	S ₂	313	77	C _{sin}	316	55	C' _{sin}	50	45	L ₂	63	10	FA ₂
344	52	S ₂	310	74	C _{sin}	122	75	C' _{sin}	52	43	L ₂	72	20	FA ₂
347	56	S ₂	314	72	C _{sin}	306	78	C' _{sin}	58	40	L ₂	74	20	FA ₂
1	70	S ₂	304	90	C _{sin}	324	70	C' _{sin}	54	30	L ₂	65	26	FA ₂
358	65	S ₂	301	80	C _{sin}	254	82	C' _{sin}	78	35	L ₂	72	20	FA ₂
181	84	S ₂	289	89	C _{sin}	270	69	C' _{sin}	64	29	L ₂	69	28	FA ₂
337	86	S ₂	111	88	C _{sin}	257	70	C' _{sin}	60	52	L ₂	14	66	FA ₂
343	65	S ₂	334	81	C _{sin}	292	77	C' _{sin}	72	5	L ₂	16	50	FA ₂
184	71	S ₂	331	77	C _{sin}	124	70	C' _{sin}	12	64	L ₂	66	20	FA ₂
1	89	S ₂	161	86	C _{sin}	117	76	C' _{sin}	251	47	L ₂	85	13	FA ₂
167	82	S ₂	316	90	C _{sin}	302	90	C' _{sin}	237	39	L ₂	272	32	FA ₂
169	88	S ₂	315	84	C _{sin}	300	72	C' _{sin}	225	37	L ₂	253	5	FA ₂
9	90	S ₂	321	66	C _{sin}	274	77	C' _{sin}	233	38	L ₂	245	12	FA ₂
171	75	S ₂	122	78	C _{sin}	77	88	C' _{sin}	249	25	L ₂	245	15	FA ₂
178	82	S ₂	315	86	C _{sin}	261	87	C' _{sin}	247	25	L ₂	79	5	FA ₂
191	86	S ₂	150	90	C _{sin}	126	80	C' _{sin}	242	19	L ₂	69	4	FA ₂
8	84	S ₂	140	84	C _{sin}	123	84	C' _{sin}	243	22	L ₂	76	0	FA ₂
188	83	S ₂	321	83	C _{sin}	136	81	C' _{sin}	244	24	L ₂	234	10	FA ₂
345	69	S ₂	305	78	C _{sin}	121	79	C' _{sin}	247	24	L ₂	242	9	FA ₂
341	75	S ₂	304	77	C _{sin}	283	83	C' _{sin}	31	43	L ₂	269	24	FA ₂
178	79	S ₂	349	67	C _{sin}	284	57	C' _{sin}	51	29	L ₂	244	22	FA ₂
178	83	S ₂	352	62	C _{sin}	260	77	C' _{sin}	50	16	L ₂	244	11	FA ₂
336	64	S ₂	145	75	C _{sin}	265	77	C' _{sin}	242	26	L ₂	222	59	FA ₂

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
334	72	S ₂	140	77	C _{sin}	110	45	C' _{sin}	243	25	L ₂	240	80	FA ₂
342	82	S ₂	310	42	C _{sin}	112	44	C' _{sin}	247	52	L ₂	257	1	FA ₂
357	58	S ₂	295	72	C _{sin}	106	89	C' _{sin}	248	47	L ₂	256	2	FA ₂
351	81	S ₂	139	84	C _{sin}	105	90	C' _{sin}	93	28	L ₂	244	3	FA ₂
346	77	S ₂	329	84	C _{sin}	237	83	C' _{sin}	97	27	L ₂	90	11	FA ₂
346	67	S ₂	332	85	C _{sin}	149	74	C' _{sin}	98	38	L ₂	240	9	FA ₂
335	90	S ₂	332	89	C _{sin}	105	76	C' _{sin}	234	38	L ₂	55	3	FA ₂
334	90	S ₂	246	78	C _{sin}	108	75	C' _{sin}	231	46	L ₂	62	1	FA ₂
334	78	S ₂	264	85	C _{sin}	107	74	C' _{sin}	237	18	L ₂	40	14	FA ₂
333	76	S ₂	312	62	C _{sin}	100	86	C' _{sin}	234	20	L ₂	219	11	FA ₂
339	80	S ₂	347	66	C _{sin}	301	84	C' _{sin}	235	17	L ₂	218	19	FA ₂
349	73	S ₂	349	63	C _{sin}	277	65	C' _{sin}	39	15	L ₂	225	26	FA ₂
342	55	S ₂	110	74	C _{sin}	275	72	C' _{sin}	55	10	L ₂	70	15	FA ₂
349	80	S ₂	307	88	C _{sin}	90	85	C' _{sin}	62	3	L ₂	70	21	FA ₂
334	50	S ₂	303	78	C _{sin}	86	73	C' _{sin}	61	9	L ₂	239	22	FA ₂
345	69	S ₂	91	72	C _{sin}	306	68	C' _{sin}	81	34	L ₂	75	16	FA ₂
324	78	S ₂	320	75	C _{sin}	306	80	C' _{sin}	67	22	L ₂	56	1	FA ₂
161	78	S ₂	305	90	C _{sin}	233	90	C' _{sin}	74	22	L ₂	237	5	FA ₂
335	88	S ₂	308	87	C _{sin}	42	89	C' _{sin}	88	23	L ₂	265	20	FA ₂
159	82	S ₂	342	68	C _{sin}	246	85	C' _{sin}	94	38	L ₂	254	5	FA ₂
159	74	S ₂	141	88	C _{sin}	225	78	C' _{sin}	152	74	L ₂	266	2	FA ₂
339	87	S ₂	332	84	C _{sin}	306	82	C' _{sin}	86	59	L ₂	265	4	FA ₂
333	80	S ₂	167	88	C _{sin}	309	90	C' _{sin}	55	27	L ₂	257	17	FA ₂
342	72	S ₂	329	84	C _{sin}	129	86	C' _{sin}	87	27	L ₂	264	21	FA ₂
339	71	S ₂	335	82	C _{sin}	280	58	C' _{sin}	56	36	L ₂	260	7	FA ₂
344	79	S ₂	129	83	C _{sin}	271	68	C' _{sin}	62	32	L ₂	263	11	FA ₂
349	89	S ₂	134	85	C _{sin}	106	81	C' _{sin}	88	25	L ₂	281	54	FA ₂
334	69	S ₂	167	86	C _{sin}	116	84	C' _{sin}	84	28	L ₂	285	3	FA ₂
338	74	S ₂	171	86	C _{sin}	114	81	C' _{sin}	83	48	L ₂	84	26	FA ₂
336	85	S ₂	134	40	C _{sin}	99	75	C' _{sin}	77	33	L ₂	95	17	FA ₂
329	80	S ₂	132	60	C _{sin}	115	77	C' _{sin}	85	12	L ₂	88	32	FA ₂
330	85	S ₂	155	75	C _{sin}	298	88	C' _{sin}	82	8	L ₂	89	31	FA ₂
299	69	S ₂	325	79	C _{sin}	310	80	C' _{sin}	168	77	L ₂	90	35	FA ₂
332	76	S ₂	149	83	C _{sin}	118	85	C' _{sin}	222	58	L ₂	90	12	FA ₂
332	77	S ₂	328	86	C _{sin}	305	87	C' _{sin}	232	42	L ₂	89	18	FA ₂
346	88	S ₂	149	86	C _{sin}	141	78	C' _{sin}	107	63	L ₂	263	33	FA ₂
336	88	S ₂	313	81	C _{sin}	82	68	C' _{sin}	73	53	L ₂	245	20	FA ₂
152	70	S ₂	165	82	C _{sin}	93	80	C' _{sin}	81	31	L ₂	112	5	FA ₂
153	81	S ₂	356	88	C _{sin}	289	70	C' _{sin}	83	28	L ₂	111	1	FA ₂
160	81	S ₂	166	70	C _{sin}	98	86	C' _{sin}	84	46	L ₂	102	32	FA ₂
150	83	S ₂	165	71	C _{sin}	90	89	C' _{sin}	89	44	L ₂	94	23	FA ₂
156	80	S ₂	345	82	C _{sin}	96	80	C' _{sin}	246	42	L ₂	84	18	FA ₂
156	81	S ₂	346	78	C _{sin}	291	81	C' _{sin}	243	58	L ₂	86	22	FA ₂
146	82	S ₂	350	87	C _{sin}	286	87	C' _{sin}	240	52	L ₂	281	5	FA ₂
322	88	S ₂	343	73	C _{sin}	315	77	C' _{sin}	246	56	L ₂	340	14	FA ₂

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
333	80	S ₂	353	80	C _{sin}	308	82	C' _{sin}	235	52	L ₂	330	24	FA ₂
3387	80	S ₂	0	80	C _{sin}	319	87	C' _{sin}	250	49	L ₂	329	32	FA ₂
159	75	S ₂	181	89	C _{sin}	274	77	C' _{sin}	252	55	L ₂	285	10	FA ₂
160	74	S ₂	170	86	C _{sin}	283	79	C' _{sin}	263	56	L ₂	270	1	FA ₂
325	84	S ₂	161	68	C _{sin}	247	63	C' _{sin}	253	56	L ₂	71	10	FA ₂
326	87	S ₂	116	64	C _{sin}	267	67	C' _{sin}	262	55	L ₂	70	16	FA ₂
157	79	S ₂	127	78	C _{sin}	266	61	C' _{sin}	237	55	L ₂	94	23	FA ₂
157	66	S ₂	119	52	C _{sin}	257	61	C' _{sin}	239	54	L ₂	89	29	FA ₂
295	65	S ₂	285	82	C _{sin}	306	88	C' _{sin}	247	36	L ₂	77	34	FA ₂
298	67	S ₂	125	62	C _{sin}	305	75	C' _{sin}	242	47	L ₂	287	8	FA ₂
144	83	S ₂	112	65	C _{sin}	147	20	AP ₁	269	75	L ₂	284	2	FA ₂
331	90	S ₂	314	80	C _{sin}	195	56	AP ₁	248	63	L ₂	91	5	FA ₂
335	90	S ₂	318	67	C _{sin}	204	54	AP ₁	241	61	L ₂	96	15	FA ₂
144	69	S ₂	306	66	C _{sin}	286	44	AP ₁	75	37	L ₂	78	4	FA ₂
146	76	S ₂	331	71	C _{sin}	305	34	AP ₁	91	42	L ₂	79	9	FA ₂
327	87	S ₂	333	84	C _{sin}	306	38	AP ₁	75	52	L ₂	268	6	FA ₂
149	88	S ₂	343	88	C _{sin}	37	67	AP ₁	75	28	L ₂	92	32	FA ₂
142	88	S ₂	324	84	C _{sin}	123	27	AP ₁	252	4	L ₂	45	45	FA ₂
153	89	S ₂	336	83	C _{sin}	172	45	AP ₁	251	19	L ₂	89	22	FA ₂
144	90	S ₂	324	80	C _{sin}	210	29	AP ₁	250	27	L ₂	65	34	FA ₂
148	90	S ₂	330	60	C _{sin}	153	30	AP ₁	68	13	L ₂	230	39	FA ₂
151	90	S ₂	310	90	C _{sin}	140	63	AP ₁	65	15	L ₂	236	18	FA ₂
142	90	S ₂	286	89	C _{sin}	97	14	AP ₁	53	9	L ₂	275	62	FA ₂
330	85	S ₂	138	88	C _{sin}	73	52	AP ₁	59	20	L ₂	64	17	FA ₂
336	80	S ₂	330	52	C _{sin}	70	50	AP ₁	59	11	L ₂	72	8	FA ₂
154	86	S ₂	316	56	C _{sin}	291	16	AP ₁	47	20	L ₂	53	12	FA ₂
141	84	S ₂	313	67	C _{sin}	265	18	AP ₁	15	38	L ₂	52	29	FA ₂
155	87	S ₂	123	55	C _{sin}	303	10	AP ₁	40	27	L ₂	95	12	FA ₂
330	70	S ₂	267	83	C _{sin}	216	23	AP ₁	34	28	L ₂	99	22	FA ₂
332	78	S ₂	295	81	C _{sin}	197	45	AP ₁	81	28	L ₂	59	32	FA ₂
153	80	S ₂	330	51	C _{sin}	89	0	AP ₁	81	20	L ₂	92	4	FA ₂
149	85	S ₂	332	84	C _{sin}	228	14	AP ₁	89	10	L ₂	248	13	FA ₃
336	83	S ₂	330	82	C _{sin}	279	28	AP ₁	68	24	L ₂	250	13	FA ₃
334	75	S ₂	308	60	C _{sin}	274	54	AP ₁	78	14	L ₂	253	2	FA ₃
333	75	S ₂	309	66	C _{sin}	22	26	AP ₁	75	25	L ₂	243	54	FA ₃
333	75	S ₂	314	51	C _{sin}	33	29	AP ₁	85	24	L ₂	251	32	FA ₃
335	84	S ₂	335	62	C _{sin}	334	36	AP ₁	77	19	L ₂	256	42	FA ₃
336	75	S ₂	318	58	C _{sin}	239	11	AP ₁	68	24	L ₂	252	12	FA ₃
336	85	S ₂	326	64	C _{sin}	1	26	AP ₁	71	20	L ₂	85	36	FA ₃
156	87	S ₂	358	46	C _{sin}	353	15	AP ₁	69	27	L ₂	266	35	FA ₃
161	89	S ₂	334	39	C _{sin}	274	29	AP ₁	71	29	L ₂	264	28	FA ₃
157	87	S ₂	319	81	C _{sin}	308	85	AP ₂	80	23	L ₂	242	43	FA ₃
331	90	S ₂	297	69	C _{sin}	300	86	AP ₂	67	32	L ₂	235	31	FA ₃
334	86	S ₂	142	78	C _{sin}	351	48	AP ₂	78	18	L ₂	25	35	FA ₃
335	86	S ₂	132	86	C _{sin}	24	64	AP ₂	76	33	L ₂	82	47	FA ₃

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
333	88	S ₂	154	78	C _{sin}	323	67	AP ₂	76	22	L ₂	212	52	FA ₃
336	89	S ₂	151	77	C _{sin}	341	82	AP ₂	256	15	L ₂	217	60	FA ₃
325	70	S ₂	128	72	C _{sin}	144	52	AP ₂	257	8	L ₂	231	59	FA ₃
339	82	S ₂	132	76	C _{sin}	130	57	AP ₂	248	4	L ₂	153	85	FA ₃
339	82	S ₂	139	72	C _{sin}	90	90	AP ₂	256	21	L ₂	271	17	FA ₃
161	88	S ₂	124	60	C _{sin}	329	65	AP ₂	266	22	L ₂	274	15	FA ₃
357	79	S ₂	121	79	C _{sin}	326	84	AP ₂	252	39	L ₂	298	18	FA ₃
126	47	S ₂	115	80	C _{sin}	312	73	AP ₂	253	14	L ₂	252	4	FA ₃
135	58	S ₂	127	77	C _{sin}	302	72	AP ₂	261	33	L ₂	258	11	FA ₃
160	69	S ₂	118	70	C _{sin}	309	65	AP ₂	253	11	L ₂	267	8	FA ₃
156	73	S ₂	303	85	C _{sin}	204	73	AP ₂	254	18	L ₂	82	11	L ₁
135	50	S ₂	307	90	C _{sin}	166	69	AP ₂	242	8	L ₂	285	7	L ₁
150	57	S ₂	312	90	C _{sin}	171	69	AP ₂	263	5	L ₂	272	3	L ₁
137	78	S ₂	309	78	C _{sin}	17	9	AP ₂	220	30	L ₂	258	18	L ₁
139	76	S ₂	313	83	C _{sin}	330	15	AP ₂	232	10	L ₂	259	20	L ₁
144	64	S ₂	337	67	C _{sin}	95	37	AP ₂	237	20	L ₂	265	13	L ₁
145	57	S ₂	344	70	C _{sin}	68	22	AP ₂	249	27	L ₂	263	20	L ₁
132	51	S ₂	356	74	C _{sin}	110	74	AP ₂	234	4	L ₂	65	4	L ₁
142	55	S ₂	107	61	C _{sin}	93	85	AP ₂	229	15	L ₂	252	8	L ₁
133	68	S ₂	300	78	C _{sin}	280	75	AP ₂	230	20	L ₂	261	11	L ₁
137	64	S ₂	304	85	C _{sin}	258	86	AP ₂	241	8	L ₂	239	34	L ₁
157	77	S ₂	332	67	C _{sin}	266	83	AP ₂	243	5	L ₂	255	21	L ₁
156	72	S ₂	322	67	C _{sin}	294	79	AP ₂	250	21	L ₂	255	32	L ₁
344	58	S ₂	319	55	C _{sin}	299	72	AP ₂	236	17	L ₂	260	32	L ₁
328	52	S ₂	321	71	C _{sin}	284	64	AP ₂	229	29	L ₂	248	21	L ₁
160	88	S ₂	323	67	C _{sin}	282	67	AP ₂	42	18	L ₂	262	19	L ₁
336	88	S ₂	168	72	C _{sin}	277	45	AP ₂	52	10	L ₂	261	6	L ₁
332	86	S ₂	324	48	C _{sin}	317	50	AP ₂	232	10	L ₂	249	21	L ₁
148	86	S ₂	309	35	C _{sin}	320	70	AP ₂	224	8	L ₂	246	2	L ₁
0	72	S ₂	336	60	C _{sin}	333	52	AP ₂	60	7	L ₂	254	15	L ₁
345	85	S ₂	338	65	C _{sin}	333	69	AP ₂	51	5	L ₂	249	46	L ₁
352	64	S ₂	342	45	C _{sin}	325	82	AP ₂	250	14	L ₂	235	41	L ₁
21	65	S ₂	355	78	C _{sin}	343	90	AP ₂	242	9	L ₂	221	24	L ₁
10	68	S ₂	349	82	C _{sin}	131	37	AP ₂	236	12	L ₂	55	8	L ₁
172	60	S ₂	353	85	C _{sin}	133	34	AP ₂	228	4	L ₂	60	4	L ₁
173	60	S ₂	349	84	C _{sin}	117	60	AP ₂	238	6	L ₂	254	4	L ₁
315	80	S ₂	339	86	C _{sin}	243	72	AP ₂	90	13	L ₂	235	18	L ₁
323	73	S ₂	353	85	C _{sin}	270	88	AP ₂	275	5	L ₂	232	11	L ₁
326	90	S ₂	354	87	C _{sin}	276	76	AP ₂	82	16	L ₂	83	20	L ₁
137	90	S ₂	343	84	C _{sin}	90	70	AP ₂	82	5	L ₂	88	21	L ₁
158	84	S ₂	342	77	C _{sin}	80	70	AP ₂	246	46	L ₃	245	6	L ₁
160	82	S ₂	325	63	C _{sin}	91	70	AP ₂	242	44	L ₃	255	14	L ₁
157	82	S ₂	326	57	C _{sin}	345	73	AP ₂	242	39	L ₃	257	12	L ₁
162	87	S ₂	339	50	C _{sin}	345	73	AP ₂	235	44	L ₃	232	2	L ₁
159	82	S ₂	342	77	C _{sin}	326	87	AP ₂	273	3	L ₃	208	52	L ₁

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
185	83	S ₂	329	66	C _{sin}	324	82	AP ₂	250	23	L _{dex}	226	25	L ₁
340	90	S ₂	294	80	C _{sin}	329	80	AP ₂	271	16	L _{dex}	232	24	L ₁
349	86	S ₂	300	68	C _{sin}	146	84	AP ₂	261	58	L _{dex}	236	24	L ₁
339	86	S ₂	286	62	C _{sin}	330	78	AP ₂	109	30	L _{dex}	236	28	L ₁
340	87	S ₂	292	60	C _{sin}	93	80	AP ₂	242	24	L _{dex}	264	19	L ₁
344	89	S ₂	320	87	C _{sin}	321	88	AP ₂	216	15	L _{dex}	264	22	L ₁
159	83	S ₂	145	86	C _{sin}	145	89	AP ₂	238	39	L _{dex}	273	32	L ₁
340	82	S ₂	346	84	C _{sin}	146	87	AP ₂	78	45	L _{dex}	251	19	L ₁
337	84	S ₂	308	86	C _{sin}	37	76	AP ₂	78	20	L _{dex}	255	17	L ₁
171	80	S ₂	351	82	C _{sin}	312	45	AP ₂	105	26	L _{dex}	62	4	L ₁
169	86	S ₂	156	90	C _{sin}	345	52	AP ₂	273	11	L _{dex}	241	17	L ₁
169	88	S ₂	321	86	C _{sin}	260	86	AP ₂	21	54	L _{dex}	250	14	L ₁
172	85	S ₂	331	89	C _{sin}	311	74	AP ₂	23	59	L _{dex}	246	10	L ₁
164	87	S ₂	312	83	C _{sin}	312	72	AP ₂	262	30	L _{rev}	248	19	L ₁
166	86	S ₂	310	80	C _{sin}	308	70	AP ₂	268	35	L _{rev}	57	4	L ₁
335	84	S ₂	316	83	C _{sin}	316	75	AP ₂	66	6	L _{rev}	262	20	L ₁
338	65	S ₂	301	77	C _{sin}	356	86	AP ₂	230	25	L _{rev}	89	20	L ₁
335	58	S ₂	306	79	C _{sin}	180	34	AP ₂	49	5	L _{rev}	218	10	L ₁
317	63	S ₂	317	77	C _{sin}	167	26	AP ₂	48	14	L _{rev}	85	28	L ₁
326	70	S ₂	309	85	C _{sin}	305	78	AP ₂	245	10	L _{rev}	265	23	L ₁
343	82	S ₂	317	86	C _{sin}	321	85	AP ₂	265	19	L _{sin}	109	11	L ₁
335	80	S ₂	313	80	C _{sin}	357	76	AP ₂	259	14	L _{sin}	271	5	L ₁
332	79	S ₂	317	76	C _{sin}	352	62	AP ₂	264	17	L _{sin}	76	5	L ₁
324	78	S ₂	318	70	C _{sin}	218	58	AP ₂	260	28	L _{sin}	251	9	L ₁
168	66	S ₂	288	88	C _{sin}	252	82	AP ₂	265	20	L _{sin}	82	8	L ₁
161	63	S ₂	303	86	C _{sin}	269	73	AP ₂	163	64	L _{sin}	264	1	L ₁
154	70	S ₂	306	88	C _{sin}	267	72	AP ₂	247	32	L _{sin}	272	5	L ₁
145	65	S ₂	304	86	C _{sin}	246	88	AP ₂	243	32	L _{sin}	106	3	L ₁
180	70	S ₂	301	85	C _{sin}	245	90	AP ₂	234	17	L _{sin}	72	4	L ₁
176	78	S ₂	301	90	C _{sin}	75	83	AP ₂	236	8	L _{sin}	82	5	L ₁
162	86	S ₂	125	85	C _{sin}	83	68	AP ₂	226	10	L _{sin}	261	7	L ₁
323	83	S ₂	114	88	C _{sin}	72	73	AP ₂	223	3	L _{sin}	258	8	L ₁
345	87	S ₂	321	90	C _{sin}	284	36	AP ₂	252	17	L _{sin}	76	4	L ₁
140	87	S ₂	140	78	C _{sin}	293	44	AP ₂	251	21	L _{sin}	243	3	L ₁
133	88	S ₂	106	80	C _{sin}	290	40	AP ₂	228	18	L _{sin}	70	1	L ₁
131	84	S ₂	114	69	C _{sin}	353	70	AP ₂	231	15	L _{sin}	72	4	L ₁
0	82	S ₂	124	90	C _{sin}	150	76	AP ₂	228	27	L _{sin}	261	21	L ₁
153	74	S ₂	136	87	C _{sin}	357	68	AP ₂	228	22	L _{sin}	267	24	L ₁
152	67	S ₂	305	78	C _{sin}	12	53	AP ₂	230	21	L _{sin}	243	8	L ₁
355	80	S ₂	317	75	C _{sin}	282	50	AP ₂	222	22	L _{sin}	224	14	L ₁
331	86	S ₂	149	84	C _{sin}	273	46	AP ₂	236	12	L _{sin}	59	8	L ₁
340	80	S ₂	150	84	C _{sin}	213	60	AP ₂	233	37	L _{sin}	38	24	L ₁
350	80	S ₂	145	80	C _{sin}	208	56	AP ₂	234	37	L _{sin}	232	1	L ₁
159	90	S ₂	144	78	C _{sin}	337	28	AP ₂	244	12	L _{sin}	248	5	L ₁
148	89	S ₂	130	61	C _{sin}	23	46	AP ₂	243	17	L _{sin}	270	14	L ₁

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
335	81	S ₂	94	64	C _{sin}	2	85	AP ₂	240	18	L _{sin}	218	38	L ₁
343	74	S ₂	114	72	C _{sin}	346	77	AP ₂	232	8	L _{sin}	160	67	L ₁
336	89	S ₂	117	70	C _{sin}	111	33	AP ₂	232	8	L _{sin}	220	49	L ₁
339	78	S ₂	129	68	C _{sin}	233	56	AP ₂	233	9	L _{sin}	215	13	L ₁
152	84	S ₂	108	54	C _{sin}	223	78	AP ₂	230	10	L _{sin}	219	21	L ₁
173	78	S ₂	115	64	C _{sin}	194	71	AP ₂	235	14	L _{sin}	233	20	L ₁
142	83	S ₂	136	82	C _{sin}	298	84	AP ₂	226	12	L _{sin}	234	12	L ₁
314	85	S ₂	130	75	C _{sin}	139	78	AP ₂	251	7	L _{sin}	250	38	L ₁
330	69	S ₂	140	80	C _{sin}	299	75	AP ₂	260	24	L _{sin}	256	12	L ₁
341	37	S ₂	148	90	C _{sin}	129	65	AP ₂	77	14	L _{sin}	76	11	L ₁
331	67	S ₂	135	79	C _{sin}	117	75	AP ₂	42	1	L _{sin}	250	9	L ₁
339	62	S ₂	133	75	C _{sin}	104	72	AP ₂	42	15	L _{sin}	222	3	L ₁
341	54	S ₂	325	89	C _{sin}	171	42	AP ₂	229	29	L _{sin}	266	7	L ₁
298	65	S ₂	330	87	C _{sin}	343	51	AP ₂	222	24	L _{sin}	260	6	L ₁
346	62	S ₂	150	89	C _{sin}	338	49	AP ₂	239	16	L _{sin}	248	18	L ₁
323	74	S ₂	338	64	C _{sin}	32	55	AP ₂	217	26	L _{sin}	226	4	L ₁
320	77	S ₂	277	90	C _{sin}	45	61	AP ₂	203	7	L _{sin}	269	14	L ₁
312	81	S ₂	122	82	C _{sin}	279	10	AP ₂	262	28	L _{sin}	248	17	L ₁
323	85	S ₂	292	83	C _{sin}	279	32	AP ₂	244	24	L _{sin}	329	22	L ₁
319	70	S ₂	282	72	C _{sin}	338	28	AP ₂	262	21	L _{sin}	240	22	L ₁
324	73	S ₂	276	67	C _{sin}	332	32	AP ₂	42	37	L _{sin}	330	24	L ₁
321	60	S ₂	275	70	C _{sin}	340	22	AP ₂	249	12	L _{sin}	264	10	L ₁
323	78	S ₂	130	72	C _{sin}	326	40	AP ₂	253	5	L _{sin}	261	3	L ₁
169	72	S ₂	135	70	C _{sin}	131	78	AP ₂	256	11	L _{sin}	82	2	L ₁
141	90	S ₂	132	85	C _{sin}	111	75	AP ₂	58	2	L _{sin}	81	24	L ₁
149	84	S ₂	297	83	C _{sin}	133	74	AP ₂	79	19	L _{sin}	96	28	L ₁
141	89	S ₂	299	80	C _{sin}	116	67	AP ₂	82	22	L _{sin}	85	32	L ₁
313	74	S ₂	304	82	C _{sin}	101	90	AP ₂	261	15	L _{sin}	76	36	L ₁
327	63	S ₂	113	80	C _{sin}	310	90	AP ₂	256	17	L _{sin}	43	20	L ₁
332	63	S ₂	133	71	C _{sin}	319	90	AP ₂	66	6	L _{sin}	44	17	L ₁
325	68	S ₂	306	74	C _{sin}	151	68	AP ₂	230	25	L _{sin}	64	7	L ₁
319	65	S ₂	104	87	C _{sin}	150	76	AP ₂	49	5	L _{sin}	58	4	L ₁
306	61	S ₂	289	74	C _{sin}	326	46	AP ₂	48	14	L _{sin}	252	24	L ₁
267	74	S ₂	296	84	C _{sin}	349	39	AP ₂	245	10	L _{sin}	248	22	L ₁
295	65	S ₂	297	68	C _{sin}	175	45	AP ₂	69	32	L _{sin}	252	12	L ₁
302	67	S ₂	319	85	C _{sin}	181	61	AP ₂	86	30	L _{sin}	275	22	L ₁
294	66	S ₂	139	82	C _{sin}	175	59	AP ₂	83	31	L _{sin}	275	26	L ₁
305	53	S ₂	305	72	C _{sin}	202	84	AP ₂	254	24	L _{sin}	264	17	L ₁
314	73	S ₂	127	81	C _{sin}	340	82	AP ₂	46	38	L _{sin}	256	4	L ₁
288	50	S ₂	294	89	C _{sin}	359	70	AP ₂	29	58	L _{sin}	260	3	L ₁
310	62	S ₂	241	71	C _{sin}	5	87	AP ₂	247	37	L _{sin}	259	8	L ₁
137	62	S ₂	106	69	C _{sin}	5	72	AP ₂	242	37	L _{sin}	256	12	L ₁
143	63	S ₂	150	83	C _{sin}	348	69	AP ₂	93	29	L _{sin}	240	45	L ₁
137	54	S ₂	144	90	C _{sin}	200	32	AP ₂	234	12	L _{sin}	238	32	L ₁
141	55	S ₂	118	72	C _{sin}	220	32	AP ₂	26	44	L _{sin}	222	52	L ₁

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A			Domain A		
143	50	S ₂	135	83	C _{sin}	210	88	AP ₂	37	36	L _{sin}	242	40	L ₁
149	74	S ₂	136	82	C _{sin}	197	5	AP ₂	41	35	L _{sin}	228	38	L ₁
141	68	S ₂	314	88	C _{sin}	11	64	AP ₂	40	25	L _{sin}	241	20	L ₁
121	90	S ₂	143	78	C _{sin}	13	66	AP ₂	33	37	L _{sin}	239	21	L ₁
149	61	S ₂	143	89	C _{sin}	24	48	AP ₂	71	34	L _{sin}	240	28	L ₁
138	70	S ₂	308	84	C _{sin}	21	65	AP ₂	72	56	L _{sin}	107	11	L ₁
155	67	S ₂	137	82	C _{sin}	29	86	AP ₂	75	14	L _{sin}	267	13	L ₁
151	75	S ₂	78	37	AP ₁	323	87	AP ₂	77	31	L _{sin}	267	7	L ₁
142	75	S ₂	359	64	AP ₁	340	84	AP ₂	264	13	L _{sin}	100	29	L ₁
145	80	S ₂	22	70	AP ₁	304	40	AP ₂	277	3	L _{sin}	105	8	L ₁
142	80	S ₂	131	36	AP ₁	77	75	AP ₃	287	33	L _{sin}	0	76	L ₁
150	54	S ₂	45	26	AP ₁	88	81	AP ₃				72	7	L ₁
129	53	S ₂	104	42	AP ₁	86	77	AP ₃				232	18	L ₁
81	50	S ₂	155	43	AP ₁	90	90	AP ₃				4	90	L ₁
337	75	S ₂	146	34	AP ₁	80	79	AP ₃				209	40	L ₁
347	88	S ₂	73	56	AP ₁	95	78	AP ₃				257	10	L ₁
342	82	S ₂	83	29	AP ₁	251	49	L ₁				255	24	L ₂
353	86	S ₂	10	54	AP ₁	263	48	L ₁				252	16	L ₂
341	88	S ₂	353	61	AP ₁	290	41	L ₁				254	12	L ₂
347	87	S ₂	277	32	AP ₁	345	17	L ₁				259	6	L ₂
321	84	S ₂	198	58	AP ₁	341	39	L ₁				214	15	L ₂
172	90	S ₂	177	44	AP ₁	8	47	L ₁				239	12	L ₂
333	75	S ₂	173	82	AP ₁	357	43	L ₁				265	3	L ₂
3	82	S ₂	166	72	AP ₁	320	58	L ₁				260	7	L ₂
336	84	S ₂	351	56	AP ₁	232	24	L ₁				255	13	L ₂
357	80	S ₂	328	24	AP ₁	42	1	L ₁				260	6	L ₂
339	84	S ₂	327	28	AP ₁	25	17	L ₁				262	3	L ₂
171	84	S ₂	298	24	AP ₁	29	32	L ₁				80	18	L ₂
171	87	S ₂	166	15	AP ₁	36	27	L ₁				254	4	L ₂
167	86	S ₂	12	33	AP ₁	10	16	L ₁				265	1	L ₂
354	86	S ₂	321	45	AP ₁	240	6	L ₁				260	26	L ₂
170	86	S ₂	335	45	AP ₁	56	20	L ₁				254	23	L ₂
155	81	S ₂	146	9	AP ₁	58	13	L ₁				276	24	L ₂
161	75	S ₂	173	33	AP ₁	63	2	L ₁				251	14	L ₂
162	75	S ₂	165	43	AP ₁	62	13	L ₁				251	16	L ₂
150	73	S ₂	188	38	AP ₁	112	40	L ₁				64	6	L ₂
151	73	S ₂	169	28	AP ₁	109	14	L ₁				264	21	L ₂
159	76	S ₂	153	60	AP ₁	108	24	L ₁				241	13	L ₂
163	82	S ₂	168	64	AP ₁	268	35	L ₁				256	5	L ₂
149	74	S ₂	170	37	AP ₁	279	24	L ₁				258	1	L ₂
152	72	S ₂	22	71	AP ₁	289	28	L ₁				262	17	L ₂
165	84	S ₂	133	18	AP ₁	195	36	L ₁				102	6	L ₂
162	83	S ₂	338	72	AP ₂	196	31	L ₁				258	23	L ₂
171	79	S ₂	332	74	AP ₂	3	9	L ₁				240	16	L ₂
163	81	S ₂	344	80	AP ₂	18	25	L ₁				258	19	L ₂

Supporting Table 1: raw data of structural measurements

Domain A			Domain A			Domain C			Domain A		
158	81	S ₂	335	89	AP ₂	355	9	L ₁	256	17	L ₂
176	85	S ₂	346	83	AP ₂	248	22	L ₂	260	7	L ₂
168	83	S ₂	169	89	AP ₂	248	10	L ₂	251	8	L ₂
180	81	S ₂	11	69	AP ₂	243	37	L ₂	60	11	L ₂
178	79	S ₂	11	60	AP ₂	252	32	L ₂	276	6	L ₂
191	75	S ₂	153	80	AP ₂	251	5	L ₂	269	3	L ₂
310	75	S ₂	160	82	AP ₂	245	21	L ₂	231	17	L ₂
320	81	S ₂	331	70	AP ₂	250	22	L ₂	254	29	L ₂
320	76	S ₂	338	65	AP ₂	244	25	L ₂	252	14	L ₂
176	67	S ₂	344	80	AP ₂	255	23	L ₂	262	10	L ₂
170	87	S ₂	335	79	AP ₂	231	23	L ₂	240	6	L ₂
168	86	S ₂	6	77	AP ₂	241	24	L ₂	261	5	L ₂
156	90	S ₂	354	83	AP ₂	299	72	L ₂	252	4	L ₂
152	83	S ₂	160	88	AP ₂	292	72	L ₂	75	2	L ₂
152	87	S ₂	342	65	AP ₂	240	25	L ₂	253	24	L ₂
156	81	S ₂	157	68	AP ₂	244	22	L ₂	258	24	L ₂
158	82	S ₂	165	72	AP ₂	238	45	L ₂	81	24	L ₂
144	87	S ₂	345	84	AP ₂	260	64	L ₂	249	9	L ₂
152	71	S ₂	340	75	AP ₂	266	58	L ₂	238	16	L ₂
159	74	S ₂	325	86	AP ₂	260	61	L ₂	258	13	L ₂
149	72	S ₂	339	77	AP ₂	266	56	L ₂	250	8	L ₂
336	84	S ₂	332	71	AP ₂	304	60	L ₂	264	22	L ₂
335	88	S ₂	190	71	AP ₂	300	64	L ₂	264	16	L ₂
169	83	S ₂	335	90	AP ₂	306	64	L ₂	253	11	L ₂
159	83	S ₂	156	75	AP ₂	262	50	L ₂	75	1	L ₂
149	88	S ₂	164	78	AP ₂	262	24	L ₂	79	14	L ₂
346	85	S ₂	337	81	AP ₂	250	28	L ₂	64	12	L ₂
328	85	S ₂	187	50	AP ₂	253	29	L ₂	272	55	L ₂
156	81	S ₂	167	82	AP ₂	262	26	L ₂	252	25	L ₂
148	85	S ₂	339	80	AP ₂	249	25	L ₂	253	30	L ₂
340	75	S ₂	131	88	AP ₂	257	31	L ₂	260	41	L ₂
340	78	S ₂	135	80	AP ₂	232	36	L ₂	266	43	L ₂
326	84	S ₂	159	81	AP ₂	274	17	L ₂	226	35	L ₂
315	80	S ₂	342	62	AP ₂	262	3	L ₂	244	20	L ₂
351	72	S ₂	4	62	AP ₂	250	7	L ₂	253	16	L ₂
343	84	S ₂	176	74	AP ₂	34	39	L ₂	228	22	L ₂
146	84	S ₂	5	81	AP ₂	13	45	L ₂	253	28	L ₂
150	87	S ₂	342	73	AP ₂	31	46	L ₂	51	9	L ₂
323	85	S ₂	344	43	AP ₂	46	44	L ₂	256	9	L ₂
168	76	S ₂	341	41	AP ₂	37	49	L ₂	251	6	L ₂
158	80	S ₂	170	53	AP ₂	247	3	L ₂	245	5	L ₂
163	77	S ₂	145	82	AP ₂	253	16	L ₂	246	2	L ₂
170	80	S ₂	298	78	AP ₂	263	9	L ₂	242	0	L ₂
152	78	S ₂	147	48	AP ₂	292	18	L ₂	247	4	L ₂
148	75	S ₂	163	71	AP ₂	288	9	L ₂	249	19	L ₂

Supporting Table 1: raw data of structural measurements

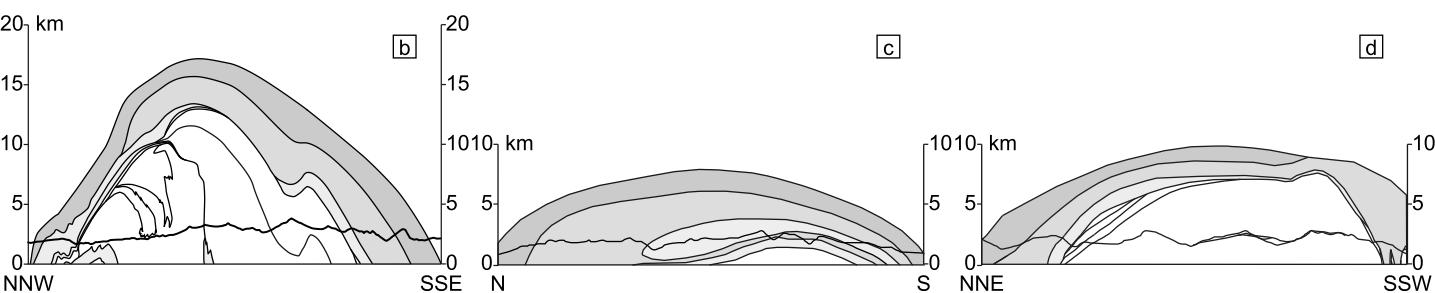
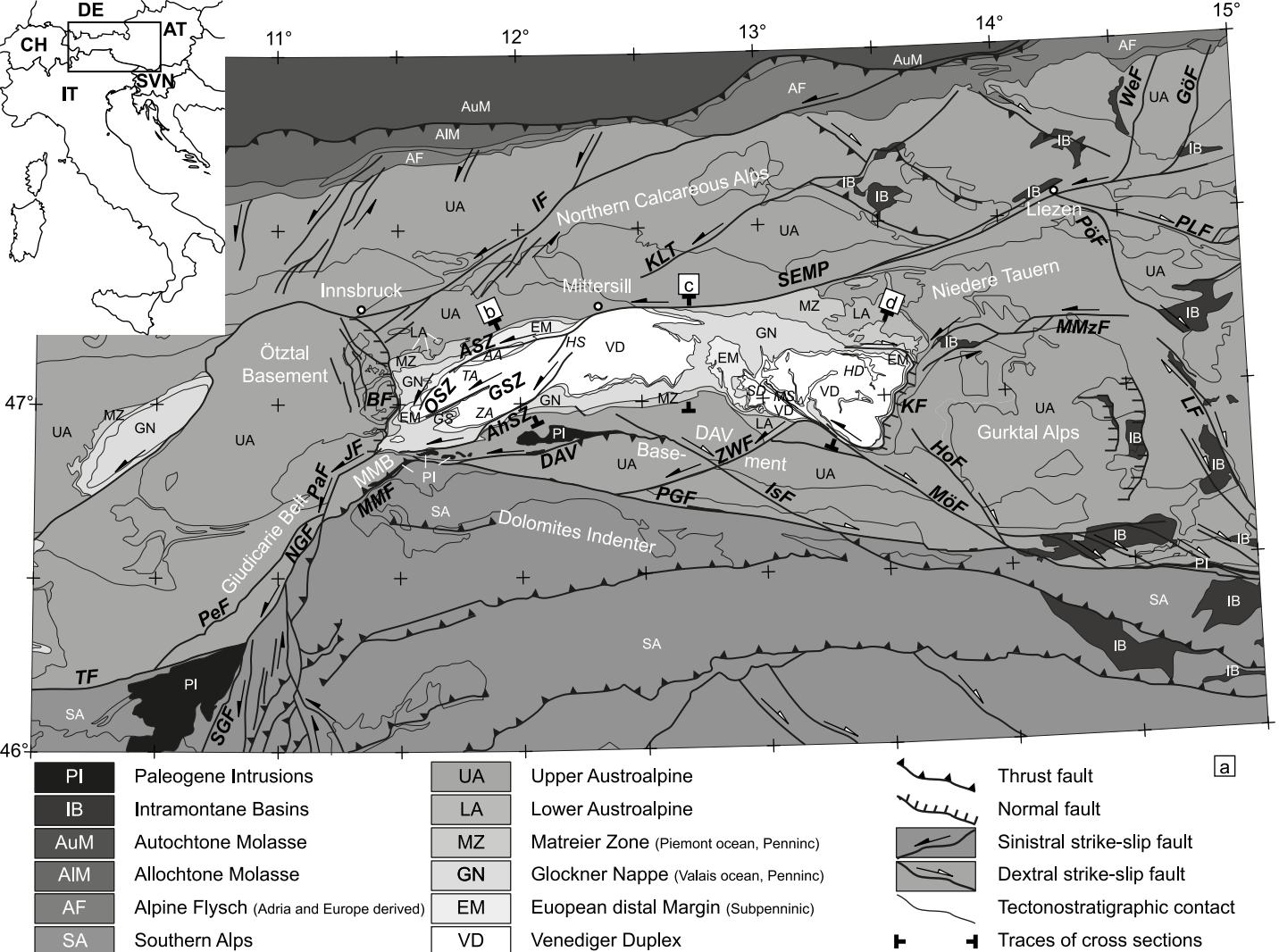
Domain A			Domain A			Domain C			Domain A		
152	75	S ₂	155	80	AP ₂	289	11	L ₂	255	20	L ₂
140	90	S ₂	329	77	AP ₂	249	19	L _{dex}	248	6	L ₂
148	90	S ₂	333	63	AP ₂	266	0	L _{dex}	254	14	L ₂
146	90	S ₂	317	80	AP ₂	313	7	L _{nor}	260	18	L ₂
314	49	S ₂	135	75	AP ₂	315	12	L _{nor}	252	14	L ₂
328	47	S ₂	136	84	AP ₂	314	8	L _{nor}	246	8	L ₂
330	46	S ₂	327	73	AP ₂	316	20	L _{nor}	243	7	L ₂
340	84	S ₂	156	87	AP ₂	310	16	L _{nor}	242	11	L ₂
161	90	S ₂	159	90	AP ₂	240	30	L _{sin}	233	37	L ₂
337	65	S ₂	332	84	AP ₂	246	26	L _{sin}	258	4	L ₂
344	73	S ₂	332	54	AP ₂	306	45	L _{sin}	261	11	L ₂
129	76	S ₂	349	56	AP ₂	292	44	L _{sin}	90	29	L ₂
145	78	S ₂	13	46	AP ₂	310	48	L _{sin}	92	22	L ₂
177	81	S ₂	140	75	AP ₂	283	46	L _{sin}	279	2	L ₂
151	71	S ₂	144	66	AP ₂	225	26	L _{sin}	87	11	L ₂
160	69	S ₂	356	51	AP ₂	287	39	L _{sin}	266	4	L ₂
163	80	S ₂	343	63	AP ₂	285	52	L _{sin}	175	2	L ₂
152	75	S ₂	343	53	AP ₂	305	52	L _{sin}	85	3	L ₂
165	84	S ₂	352	35	AP ₂	271	65	L _{sin}	94	0	L ₂
163	81	S ₂	348	55	AP ₂	269	39	L _{sin}	269	4	L ₂
153	77	S ₂	358	62	AP ₂	259	27	L _{sin}	275	6	L ₂
160	86	S ₂	342	54	AP ₂	266	3	FA ₁	270	4	L ₂
157	85	S ₂	348	52	AP ₂	275	19	FA ₁	266	5	L ₂
150	75	S ₂	346	59	AP ₂	343	10	FA ₁	268	4	L ₂
348	74	S ₂	19	76	AP ₂	33	16	FA ₁	265	4	L ₂
343	67	S ₂	10	76	AP ₂	37	11	FA ₁	88	28	L ₂
346	69	S ₂	10	80	AP ₂	130	12	FA ₁	90	44	L ₂
339	79	S ₂	9	89	AP ₂	108	22	FA ₁	91	32	L ₂
343	76	S ₂	11	90	AP ₂	107	26	FA ₁	91	25	L ₂
341	80	S ₂	181	81	AP ₂	195	23	FA ₁	83	22	L ₂
350	75	S ₂	181	82	AP ₂	68	24	FA ₁	94	34	L ₂
356	77	S ₂	163	81	AP ₂	125	54	FA ₁	100	27	L ₂
354	75	S ₂	178	74	AP ₂	175	7	FA ₁	92	21	L ₂
348	90	S ₂	348	89	AP ₂	293	8	FA ₁	67	27	L ₂
353	90	S ₂	348	90	AP ₂	250	10	FA ₁	278	6	L ₂
349	87	S ₂	26	25	AP ₂	274	15	FA ₁	62	12	L ₂
349	85	S ₂	80	38	AP ₂	91	4	FA ₁	69	12	L ₂
344	87	S ₂	46	39	AP ₂	158	15	FA ₁	75	20	L ₂
352	84	S ₂	12	86	AP ₂	254	6	FA ₁	64	2	L ₂
343	85	S ₂	6	82	AP ₂	203	4	FA ₁	45	7	L ₂
342	72	S ₂	1	67	AP ₂	285	22	FA ₁	61	17	L ₂
350	74	S ₂	357	73	AP ₂	96	15	FA ₁	74	5	L ₂
174	83	S ₂	336	65	AP ₂	267	20	FA ₁	62	28	L ₂
308	78	S ₂	349	86	AP ₂	265	12	FA ₁	234	3	L ₂
348	86	S ₂	349	62	AP ₂	261	24	FA ₁	93	28	L ₂

Supporting Table 1: raw data of structural measurements

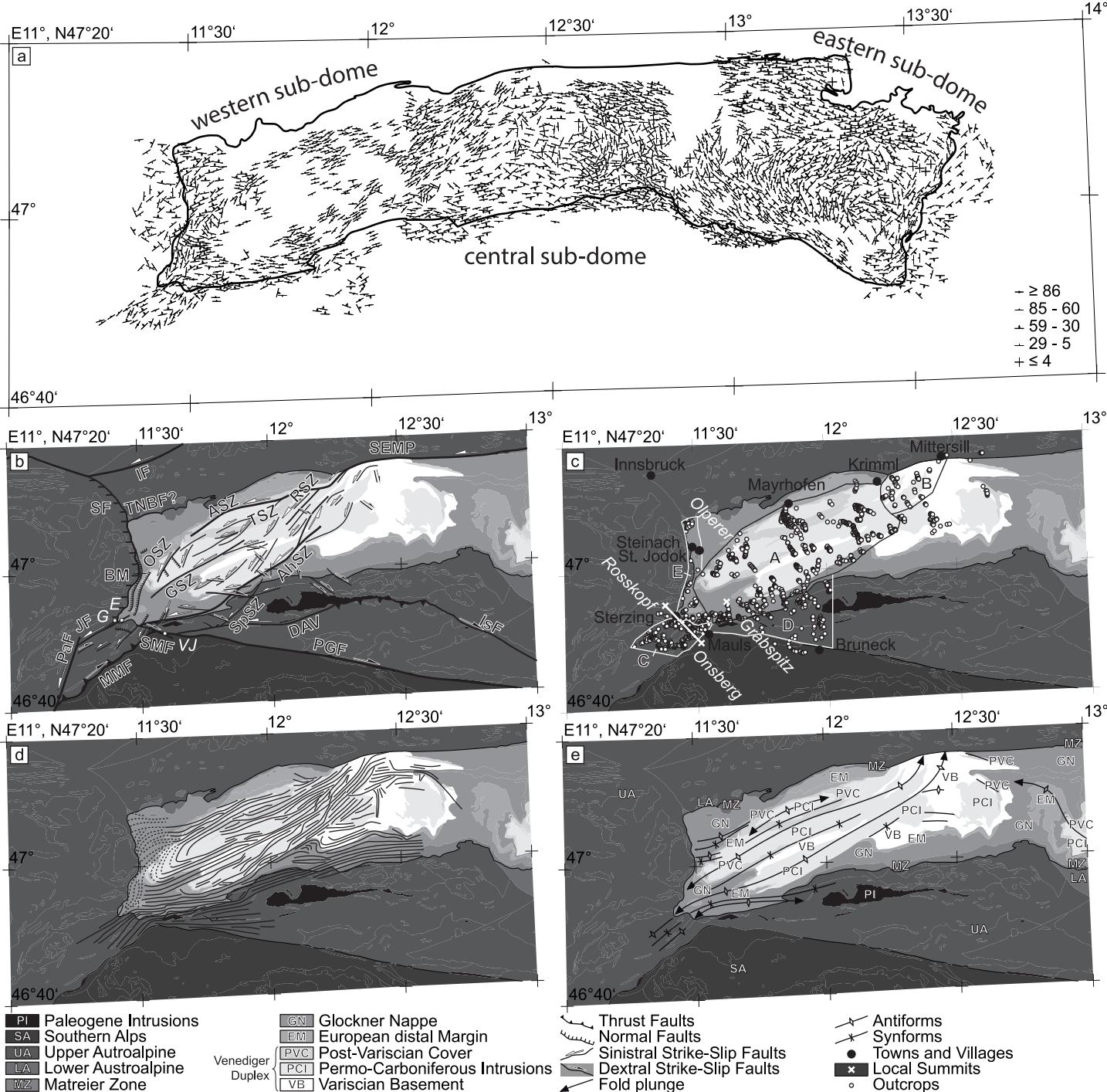
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324	59	S ₂	159	82	AP ₂	258	13	FA ₁	76	35	L ₂
330	72	S ₂	163	78	AP ₂	9	65	FA ₂	70	26	L ₂
334	65	S ₂	342	84	AP ₂	30	60	FA ₂	90	25	L ₂
345	64	S ₂	154	76	AP ₂	225	26	FA ₂	262	28	L ₂
351	60	S ₂	340	88	AP ₂	214	23	FA ₂	258	25	L ₂
306	87	S ₂	328	67	AP ₂	226	41	FA ₂	232	12	L ₂
127	67	S ₂	344	84	AP ₂	205	13	FA ₂	49	11	L ₂
332	70	S ₂	186	84	AP ₂	51	4	FA ₂	226	24	L ₂
339	76	S ₂	182	78	AP ₂	264	13	FA ₂	235	10	L ₂
322	83	S ₂	149	80	AP ₂	76	2	FA ₂	251	30	L ₂
323	90	S ₂	167	82	AP ₂	77	7	FA ₂	176	68	L ₂
320	90	S ₂	348	86	AP ₂	29	0	FA ₂	244	16	L ₂
153	84	S ₂	346	80	AP ₂	34	14	FA ₂	242	37	L ₂
151	78	S ₂	354	79	AP ₂	40	23	FA ₂	236	30	L ₂
328	72	S ₂	8	81	AP ₂	26	18	FA ₂	237	48	L ₂
324	78	S ₂	13	86	AP ₂	26	10	FA ₂	241	30	L ₂
328	75	S ₂	13	85	AP ₂	28	32	FA ₂	252	14	L ₂
311	81	S ₂	9	66	AP ₂	13	33	FA ₂	254	12	L ₂
127	53	S ₂	344	72	AP ₂	357	12	FA ₂	254	14	L ₂
129	62	S ₂	181	86	AP ₂	2	13	FA ₂	261	20	L ₂
147	75	S ₂	352	86	AP ₂	189	13	FA ₂	242	12	L ₂
144	82	S ₂	355	83	AP ₂	190	20	FA ₂	242	4	L ₂
155	61	S ₂	1	85	AP ₂	13	17	FA ₂	274	2	L ₂
153	61	S ₂	5	88	AP ₂	12	14	FA ₂	94	14	L ₂
316	87	S ₂	173	88	AP ₂	237	4	FA ₂	58	53	L ₂
134	85	S ₂	180	90	AP ₂	258	40	FA ₂	250	7	L ₂
316	87	S ₂	2	87	AP ₂	80	15	FA ₂	252	13	L ₂
134	85	S ₂	134	73	AP ₂	227	16	FA ₂	242	8	L ₂
126	69	S ₂	356	87	AP ₂	238	20	FA ₂	254	9	L ₂
127	69	S ₂	176	85	AP ₂	190	20	FA ₂	254	0	L ₂
133	80	S ₂	174	81	AP ₂	194	15	FA ₂	250	0	L ₂
135	62	S ₂	168	81	AP ₂	164	8	FA ₂	245	1	L ₂
132	58	S ₂	176	84	AP ₂	164	8	FA ₂	76	15	L ₂
149	77	S ₂	6	83	AP ₂	56	8	FA ₂	82	18	L ₂
150	80	S ₂	355	86	AP ₂	242	29	FA ₂	280	6	L ₂
324	75	S ₂	170	84	AP ₂	241	40	FA ₂	78	39	L ₂
332	75	S ₂	176	86	AP ₂	241	30	FA ₂	51	41	L ₂
314	79	S ₂	359	70	AP ₂	243	13	FA ₂	98	26	L ₂
313	84	S ₂	4	76	AP ₂	72	29	FA ₂	91	27	L ₂
154	88	S ₂	313	13	AP ₂	318	40	FA ₂	88	25	L ₂
335	67	S ₂				243	13	FA ₂	83	23	L ₂
338	75	S ₂				333	36	FA ₂	272	2	L ₂
329	69	S ₂				230	9	FA ₂	83	34	L ₂
176	67	S ₂				243	9	FA ₂	87	30	L ₂
321	87	S ₂				230	60	FA ₂	96	26	L ₂

Supporting Table 1: raw data of structural measurements

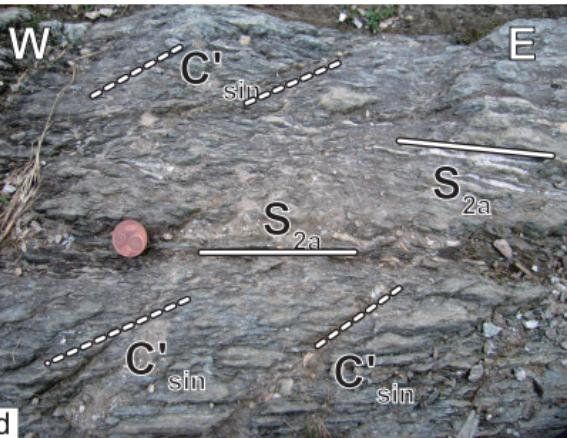
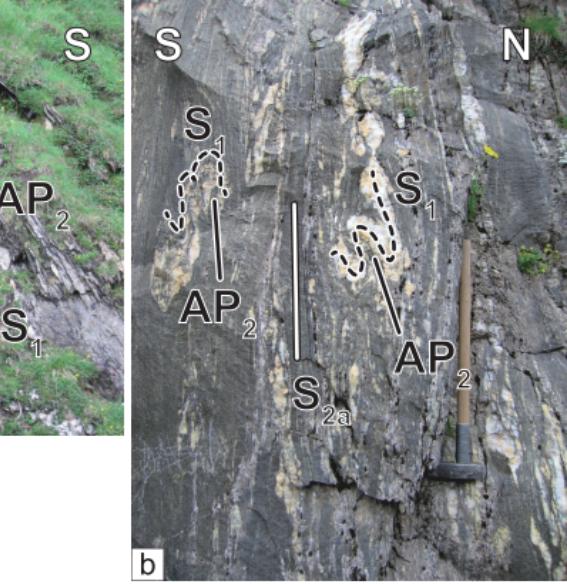
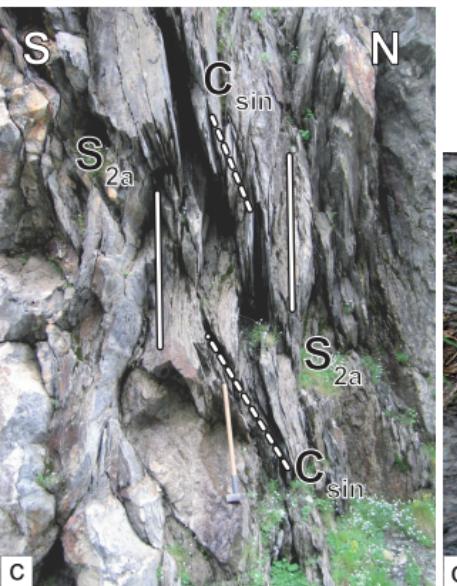
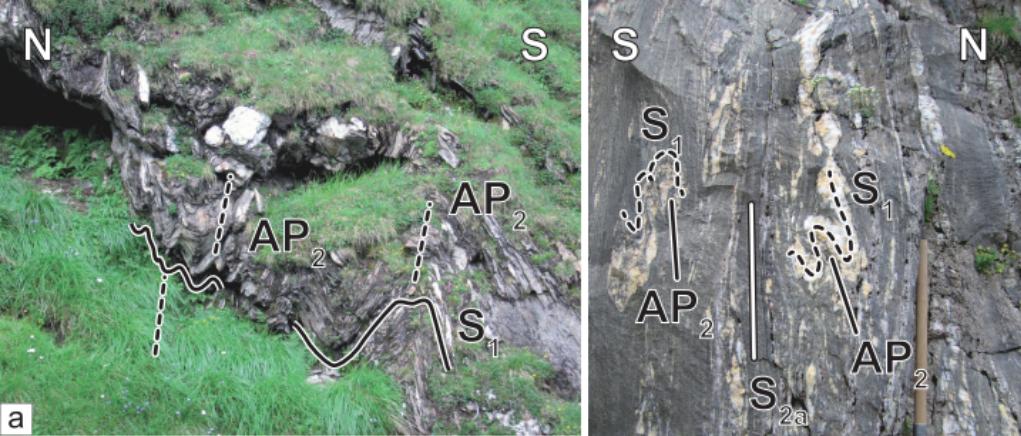
Domain A			Domain C			Domain A		
329	85	S ₂	240	45	FA ₂	98	27	L ₂
300	57	S ₂	225	18	FA ₂	92	35	L ₂
328	77	S ₂	240	14	FA ₂	69	87	L ₂
328	85	S ₂	293	42	FA ₂	94	25	L ₂
322	86	S ₂	339	28	FA ₂	90	34	L ₂
316	84	S ₂	351	20	FA ₂	94	24	L ₂
318	68	S ₂	329	5	FA ₂	96	9	L ₂
324	73	S ₂	346	21	FA ₂	101	14	L ₂
325	85	S ₂	77	34	FA ₂	92	30	L ₂
328	89	S ₂	269	68	FA ₂	87	14	L ₂
152	87	S ₂	267	60	FA ₂	119	34	L ₂
306	81	S ₂	295	14	FA ₂	81	24	L ₂
323	81	S ₂	101	8	FA ₂	87	7	L ₂
166	78	S ₂	290	53	FA ₂	94	35	L ₂
171	75	S ₂	101	43	FA ₂	92	17	L ₂
346	80	S ₂	238	30	FA ₂	92	35	L ₂
176	77	S ₂	47	17	FA ₂	95	17	L ₂
160	73	S ₂	194	8	FA ₂	94	41	L ₂
149	88	S ₂	194	59	FA ₂	91	17	L ₂
148	87	S ₂	197	27	FA ₂	92	26	L ₂
164	86	S ₂	237	21	FA ₂	96	3	L ₂
191	78	S ₂	122	18	FA ₂	112	29	L ₂
170	79	S ₂	124	23	FA ₂	86	16	L ₂
178	76	S ₂	19	12	FA ₂	86	26	L ₂
182	74	S ₂	13	26	FA ₂	88	12	L ₂
185	72	S ₂	20	4	FA ₂	87	12	L ₂
180	76	S ₂	19	28	FA ₂	250	21	L ₂
191	72	S ₂	8	31	FA ₂	251	21	L ₂
177	87	S ₂	47	14	FA ₂	240	40	L ₂
346	86	S ₂	59	18	FA ₂	234	35	L ₂
340	86	S ₂	35	25	FA ₂	237	41	L ₂
318	70	S ₂	57	4	FA ₂	243	38	L ₂
308	67	S ₂	126	2	FA ₂	41	18	L ₂
292	65	S ₂	248	7	FA ₂	40	11	L ₂
292	64	S ₂	11	24	FA ₂	235	28	L ₂
279	65	S ₂	90	71	FA ₃	234	35	L ₂
317	74	S ₂	129	74	FA ₃	45	8	L ₂
348	84	S ₂	87	70	FA ₃	90	47	L ₂
311	63	S ₂	86	85	FA ₃	84	55	L ₂
319	58	S ₂	120	75	FA ₃	82	40	L ₂
333	75	S ₂	93	78	FA ₃	251	16	L ₂
330	68	S ₂				88	65	L ₂
164	90	S ₂				67	31	L ₂
338	77	S ₂				68	29	L ₂
138	62	S ₂				267	5	L ₂

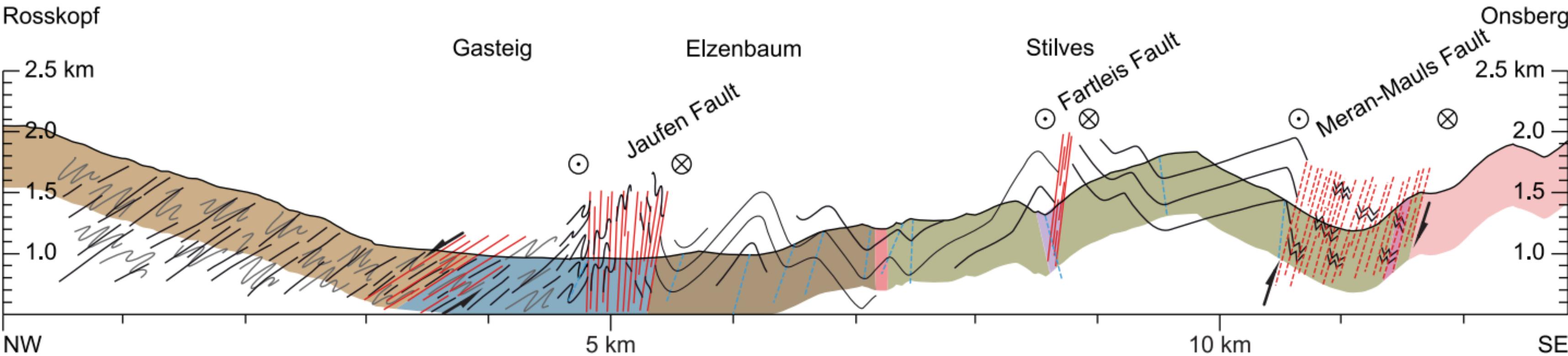


AhSZ: Ahrental Shear Zone, **ASZ:** Ahorn Shear Zone, **BF:** Brenner Fault, **DAV:** Deferegggen-Antholz-Vals Fault, **GÖF:** Görtzschitztal Fault, **GSZ:** Greiner Shear Zone, **HoF:** Hochstuhl Fault, **IF:** Inntal Fault, **IS:** Iseltaal Fault, **JF:** Jaufen Fault, **KLT:** Königssee-Lammertal-Traunsee Fault, **KF:** Katschberg Fault, **LF:** Lavanttal Fault, **MMB:** Meran-Mauls Basement, **MMF:** Meran-Mauls Fault, **MMZF:** Mur-Mürz Fault; **MöF:** Mölltal Fault, **NGF:** Northern Giudicarie Fault, **OSZ:** Olperer Shear Zone, **PaF:** Passeier Fault, **PeF:** Pejo Fault, **PLF:** Palten-Liesing-Fault, **PÖF:** Pöls Fault, **PGF:** Pustertal-Gailtal Fault, **SEMP:** Salzach-Ennstal-Mariazell-Puchberg Fault, **SGF:** Southern Giudicarie Fault, **TF:** Tonale Fault, **WeF:** Wyer Fault, **ZWF:** Zwischenberge-Wöllentratten Fault, **AA:** Ahorn Antiform, **GS:** Greiner Synform, **HD:** Hochalm Dome, **HS:** Habach Synform, **MS:** Mallnitz Synform, **SD:** Sonnblick Dome, **TA:** Tuxer Antiform, **ZA:** Zillertaler Antiform



AhSZ: Ahrntal Shear Zone, ASZ: Ahorn Shear Zone, BM: Brenner Mylonites, DAV: Deferggen-Antholz-Vals Fault, E: Elzenbaum, G: Gasteig, GSZ: Greiner Shear Zone, IF: Inntal Fault, IsF: Iseltaal Fault, JF: Jaufen Fault, MMF: Meran-Maules Fault, OSZ: Olperer Shear Zone, PaF: Passeier Fault, PGF: Pustertal Fault, RSZ: Rinderkar Shear Zone, SEMP: Salzach-Ennstal-Puchberg-Mariazell Fault, SF: Silltal Fault, SMF: Sprechenstein-Maules Fault, SpSz: Speikboden Shear Zone, TNBF: Tauern Northern Boundary Fault, TSZ: Tuxer shear zones, VJ: Valler Jöchl

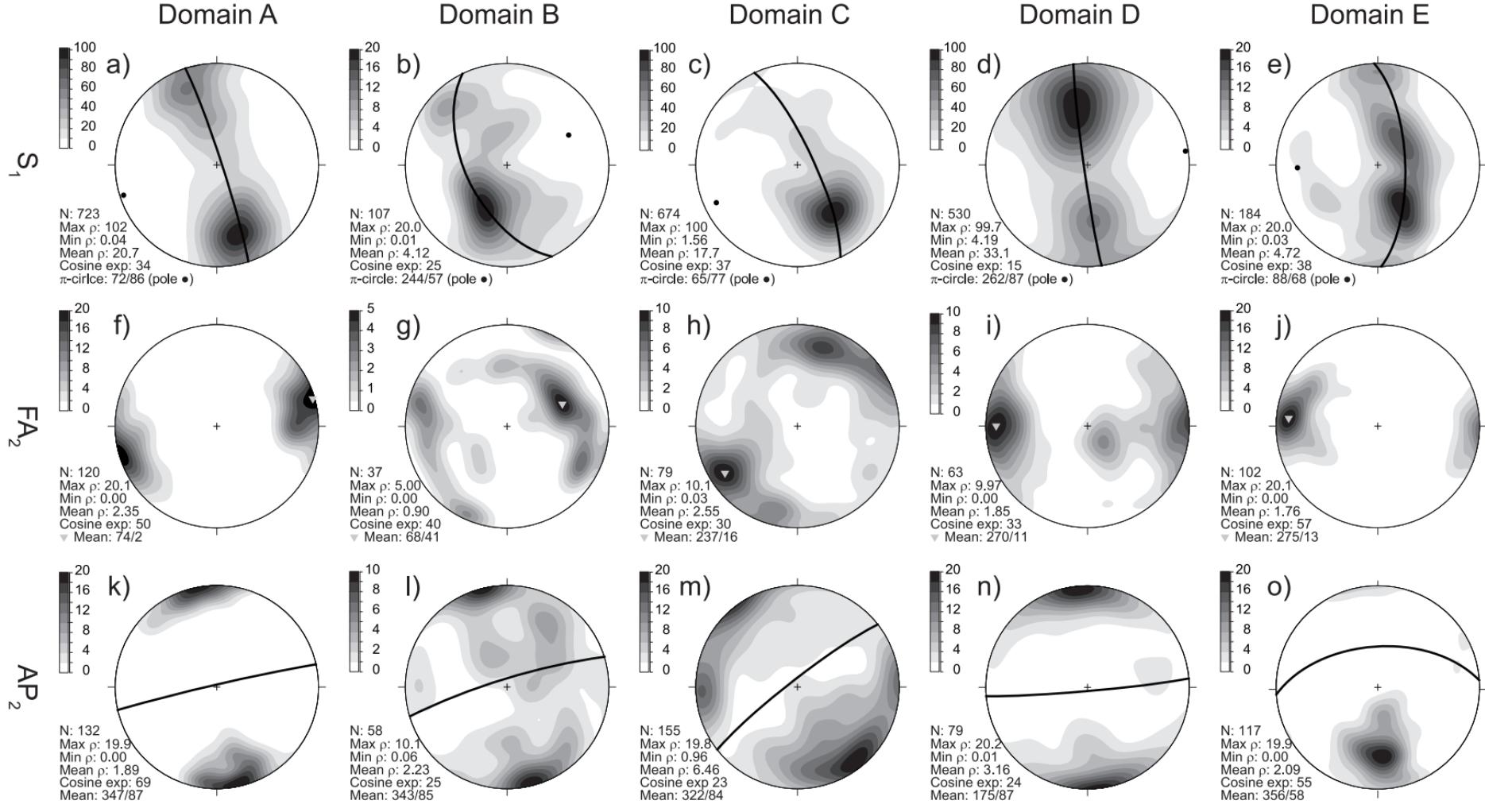




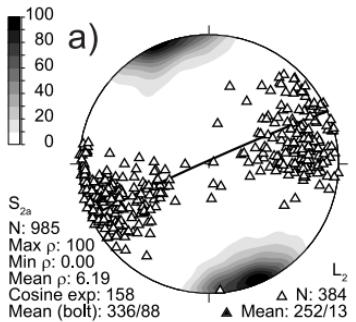
Brixen Granite	(SA)
Granodiorite - Tonalite	(PI)
Gnt-Bt-bearing Micaschist	(UA)
Granitegneiss	(UA)

Permotriassic sediment	(SA)
Bt-Mus-Paragneiss	(UA)
Phyllites, Micaschist	(UA)
Calcareous-micaschist	(GN)

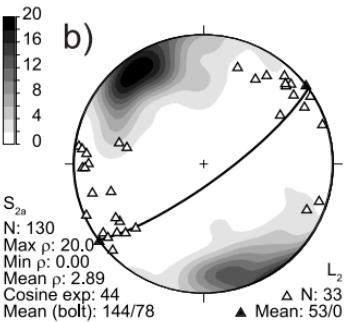
S_0 , isoclinally folded schistosity
 S_1 , axial plane foliation, pre- D_2
 S_2 , axial plane foliation, D_2
 C_{sin}^{2a} and C_{nor} , ductile and semiductile mylonites, D_2



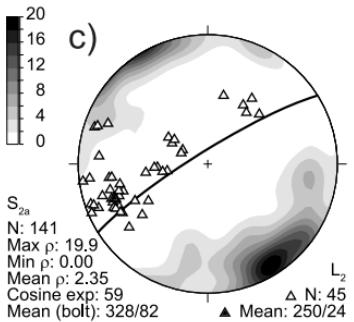
Domain A



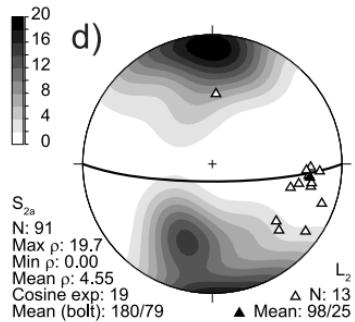
Domain B



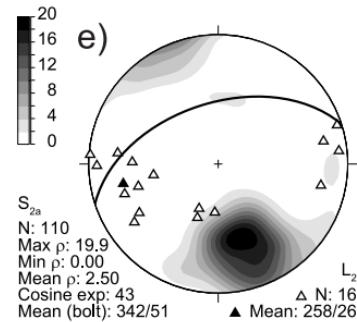
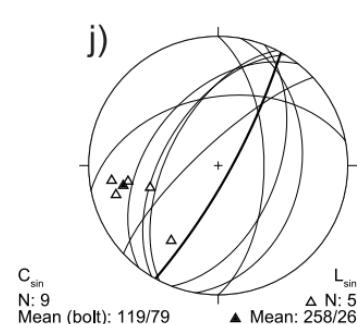
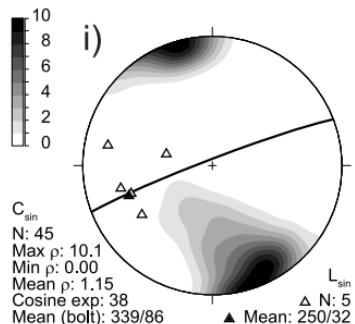
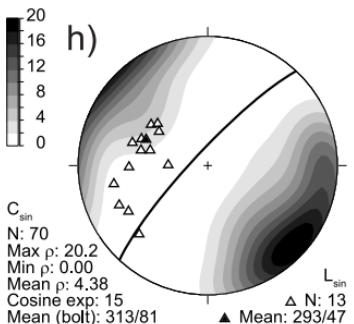
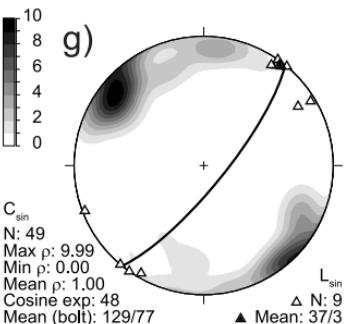
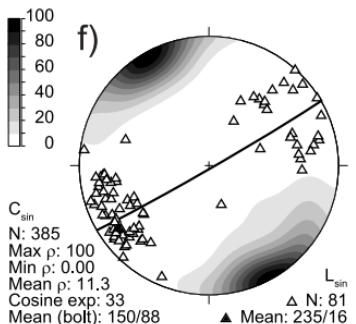
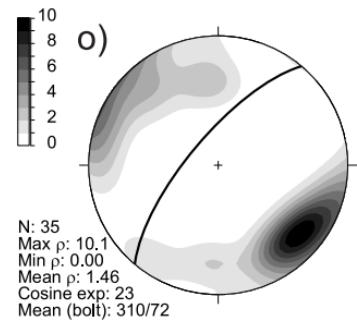
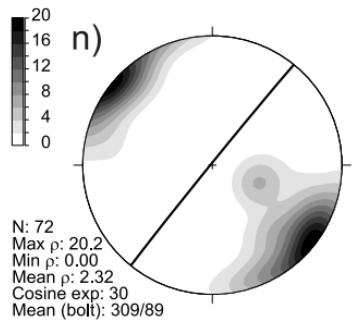
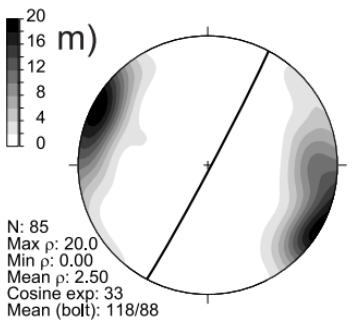
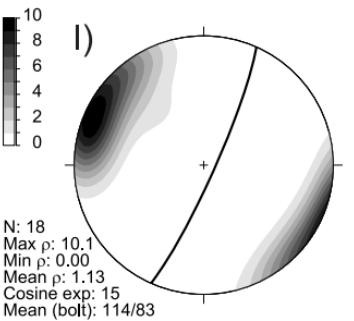
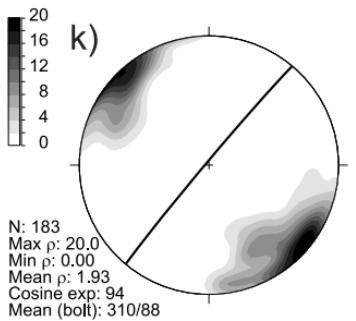
Domain C



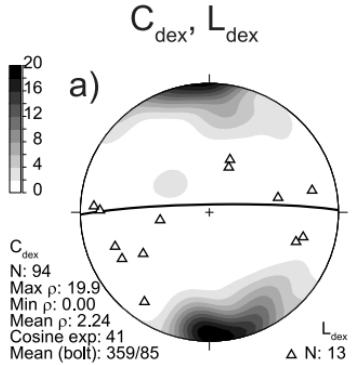
Domain D



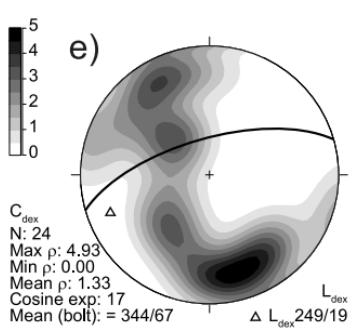
Domain E

 S_{2a}, L_2  C_{\sin}, L_{\sin}  C'_{\sin}

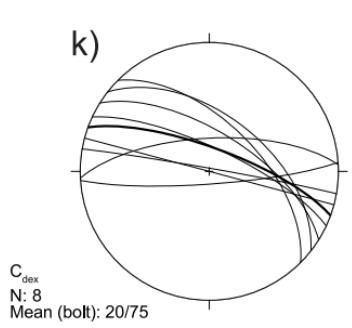
Domain A

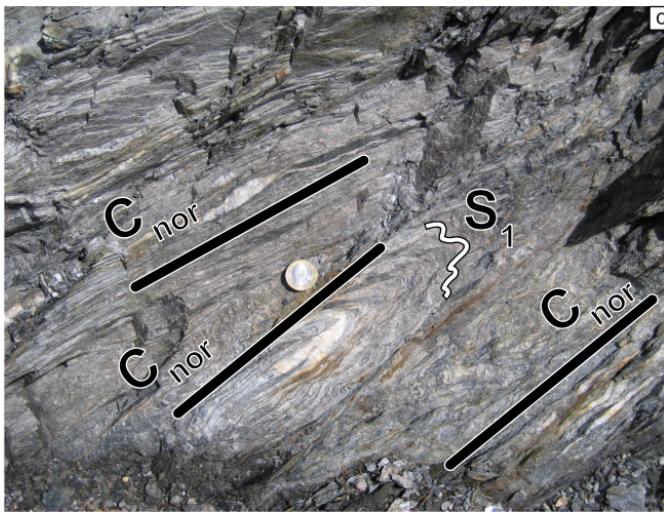
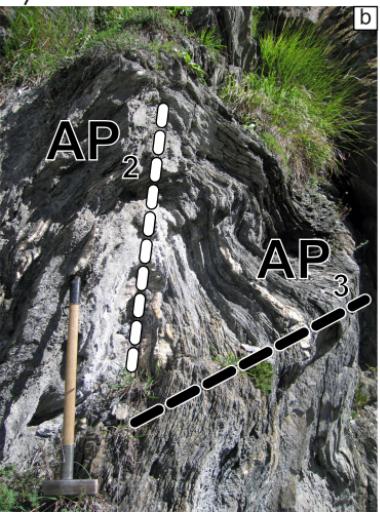
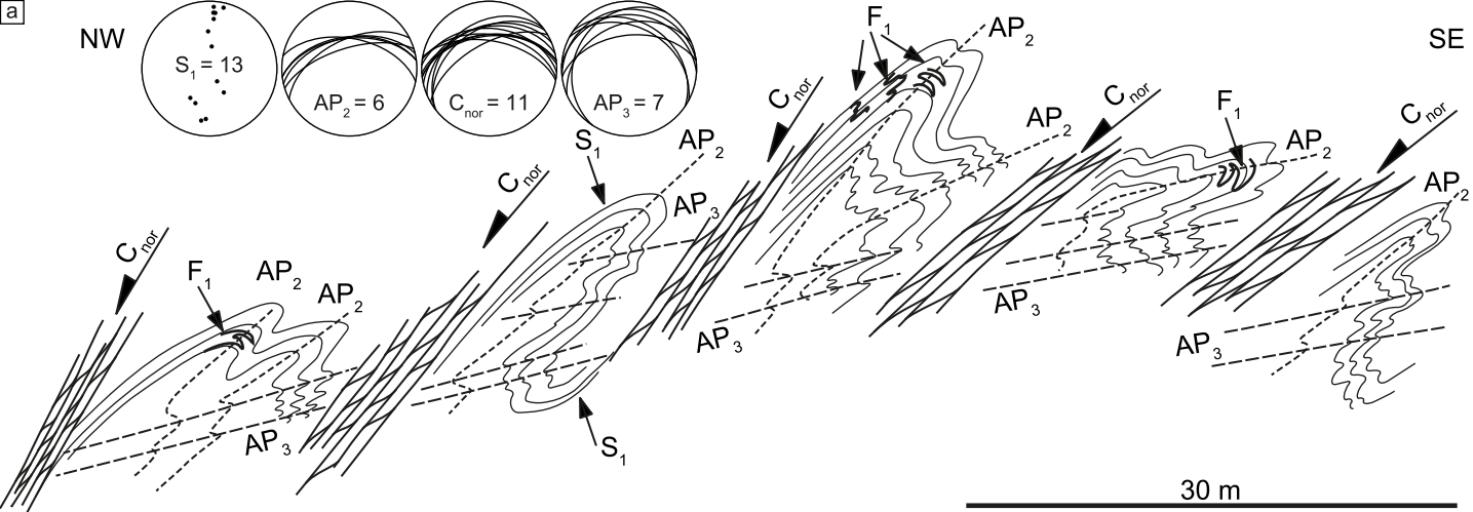


Domain C

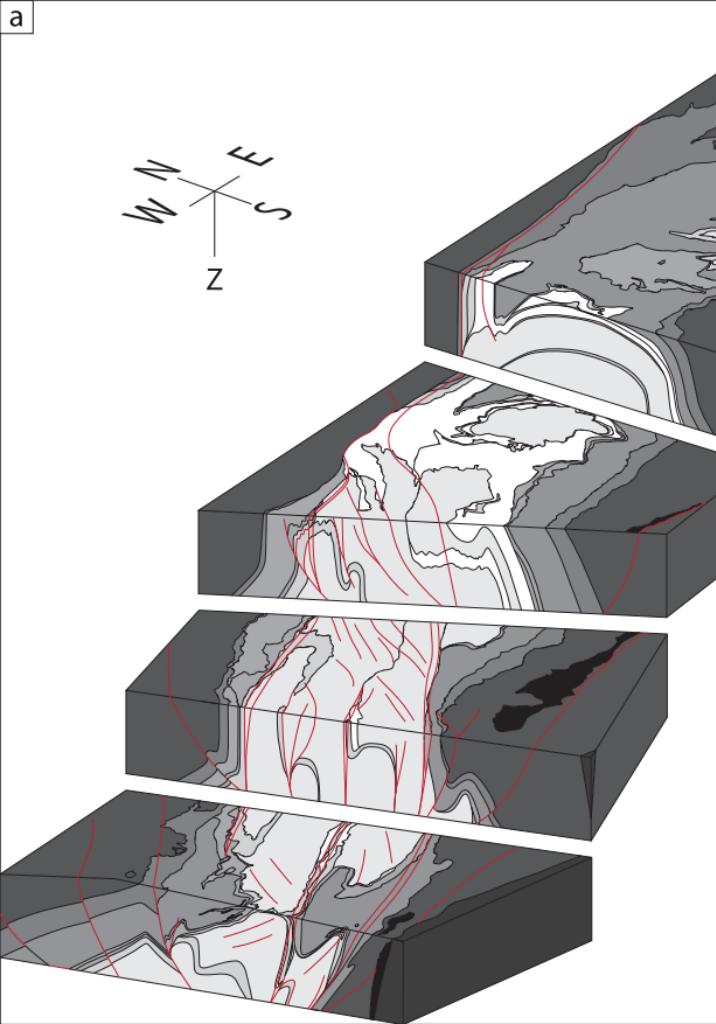


Domain E

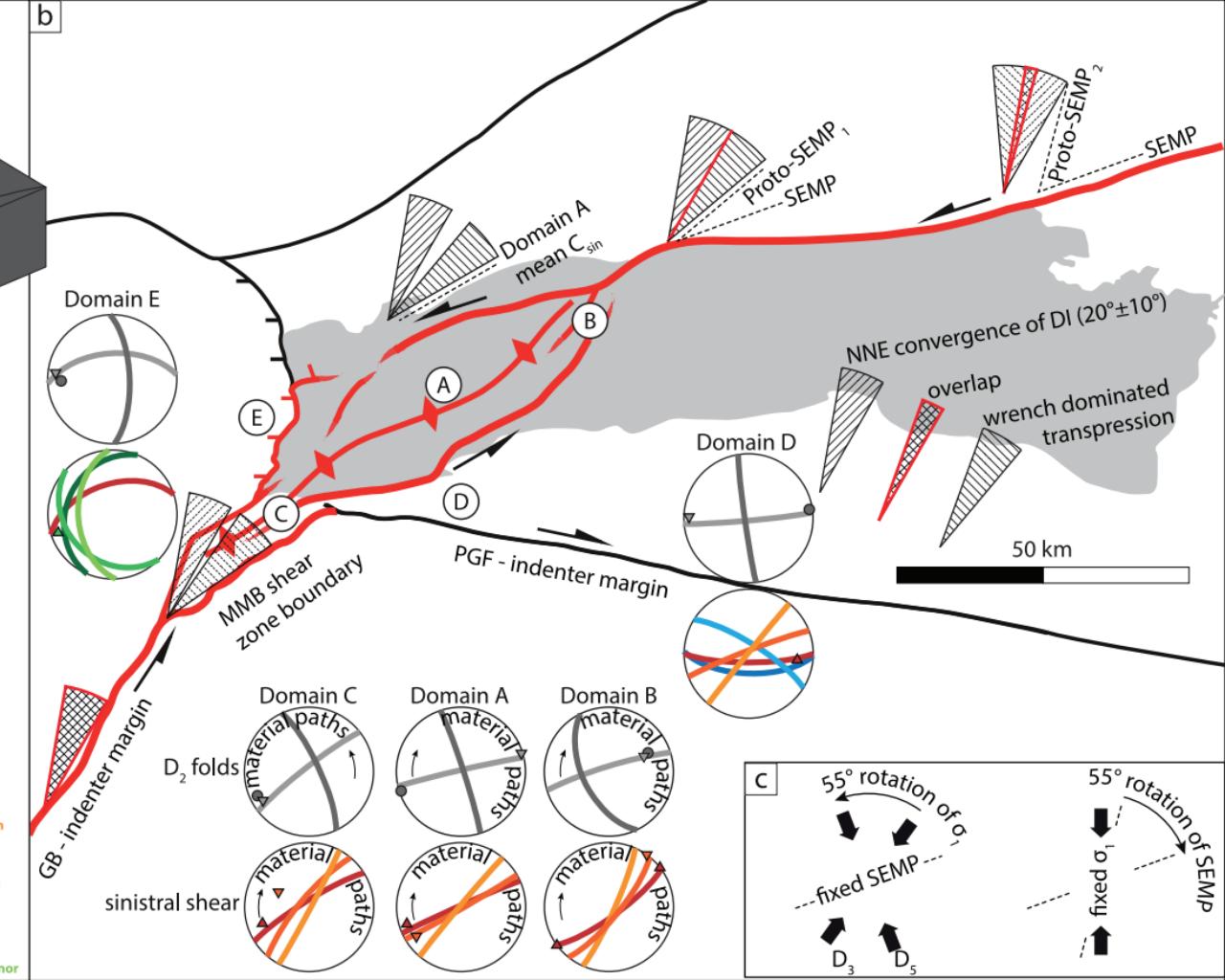


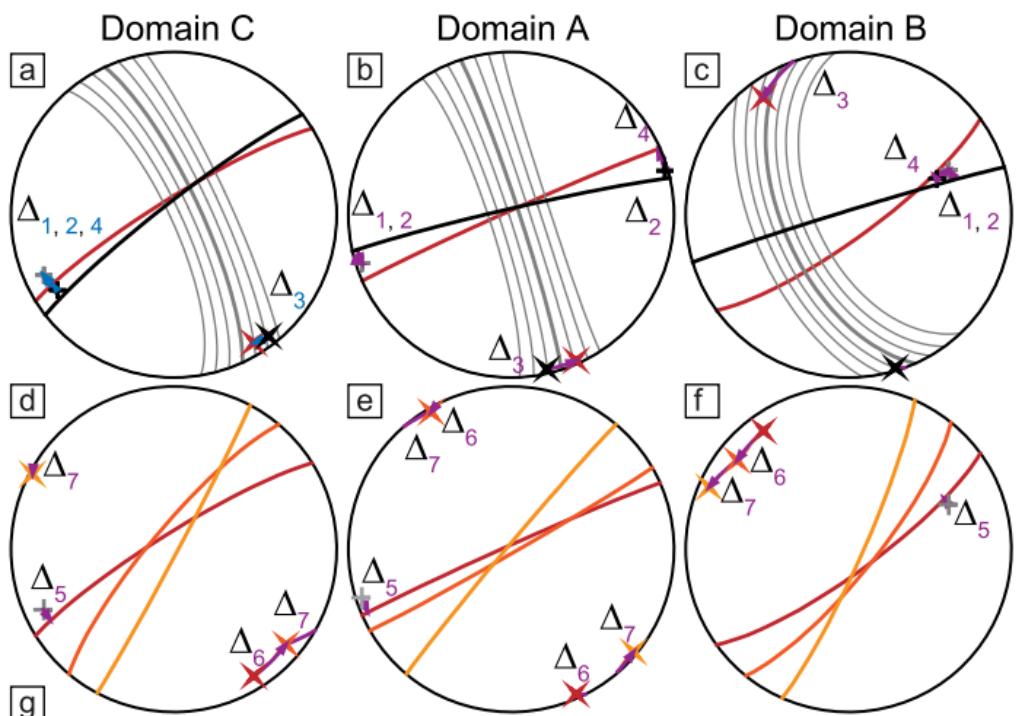


a

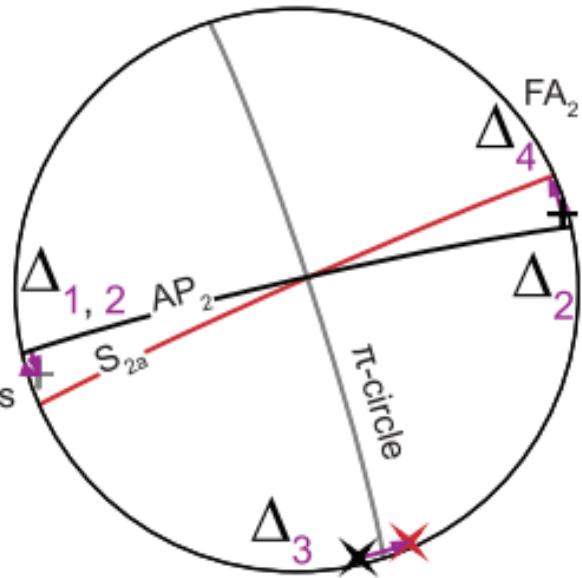
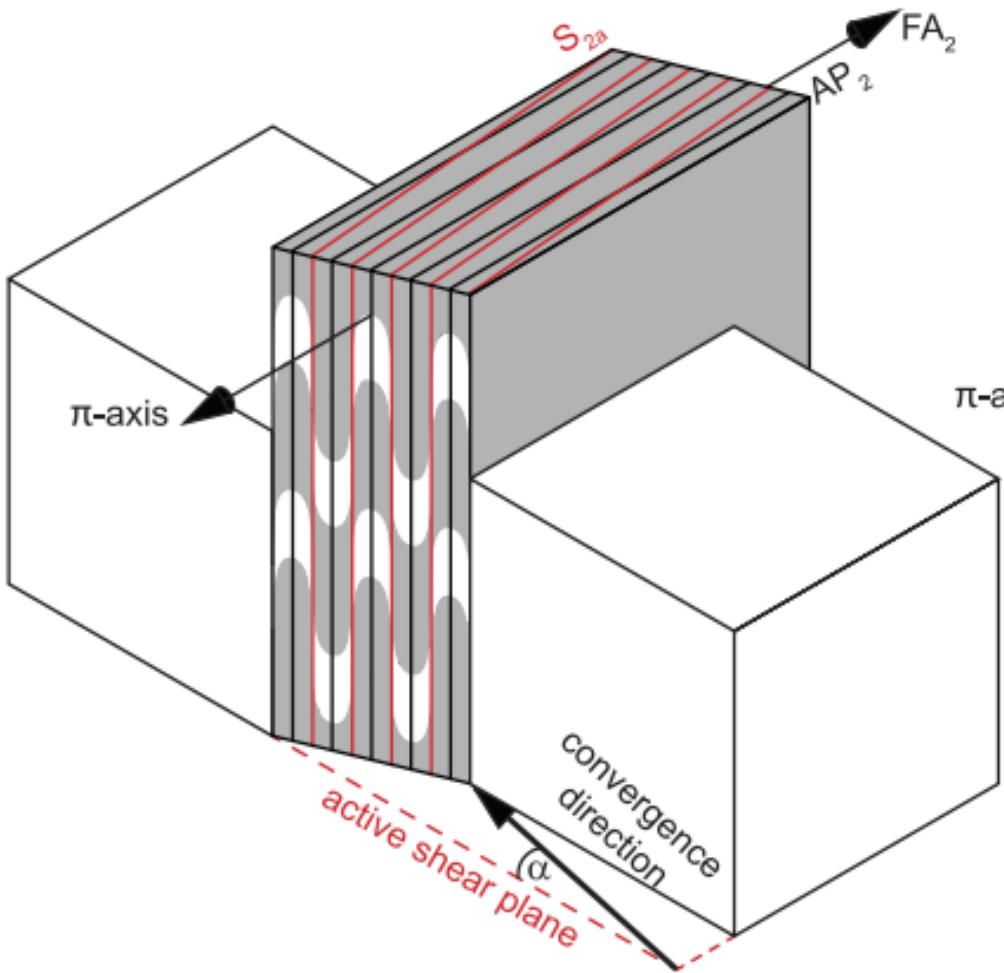


b



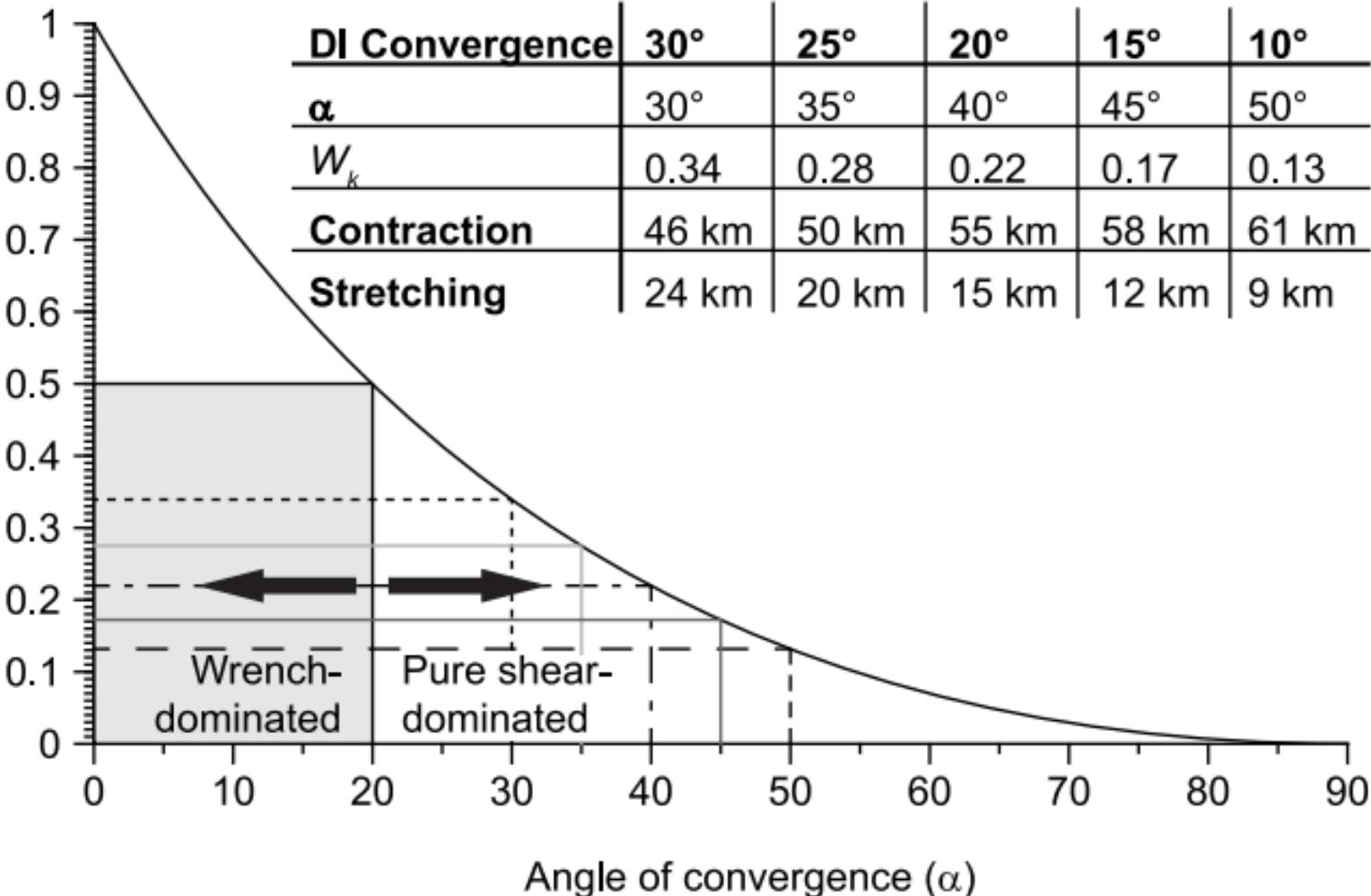


Dihedral angles (in degree °)	Domain C	Domain A	Domain B
$\Delta_1 = (\text{AP}_2 - \pi\text{-axis})$ (/, +)	11	-5	-5
$\Delta_2 = (\text{FA}_2 - \pi\text{-axis})$ (+, +)	8	-6	-8
$\Delta_3 = (\text{AP}_2 - \text{S}_{2a})$ (x, x)	6	-11	-25
$\Delta_4 = (\text{FA}_2 - \text{S}_{2a})$ (+, /)	3	-9	-3
$\Delta_5 = (\pi\text{-axis} - \text{S}_{2a})$ (+, /)	-5	-6	-4
$\Delta_6 = (\text{S}_{2a} - \text{C}_{\sin})$ (x, x)	-18	-7	-15
$\Delta_7 = (\text{C}_{\sin} - \text{C}'_{\sin})$ (x, x)	-18	-20	-16



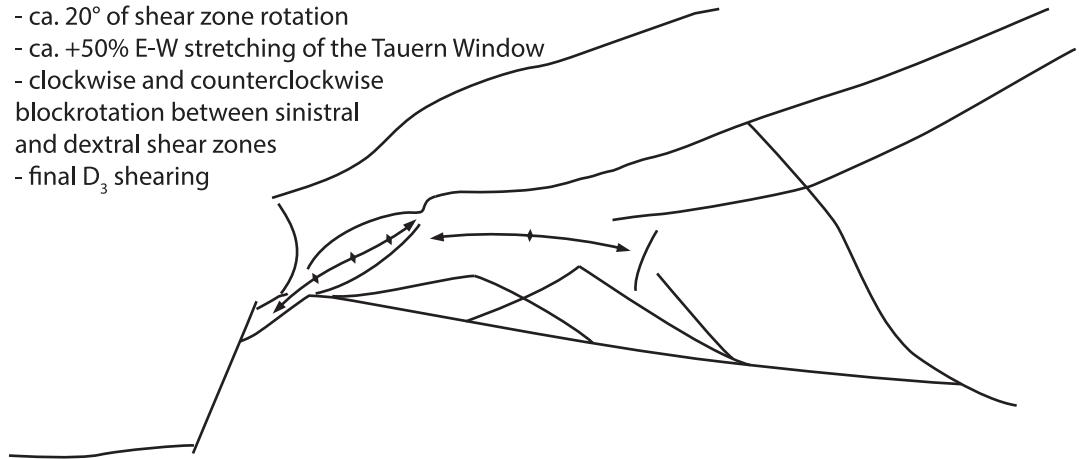
Dihedral angles (D_2 , folding)	
$\Delta_1 = (AP_2 - \pi\text{-axis})$	(↙, +)
$\Delta_2 = (FA_2 - \pi\text{-axis})$	(↗, +)
$\Delta_3 = (AP_2 - S_{2a})$	(✗, ✗)
$\Delta_4 = (FA_2 - S_{2a})$	(↗, ↘)

Extension rate / Contraction rate (W_k)



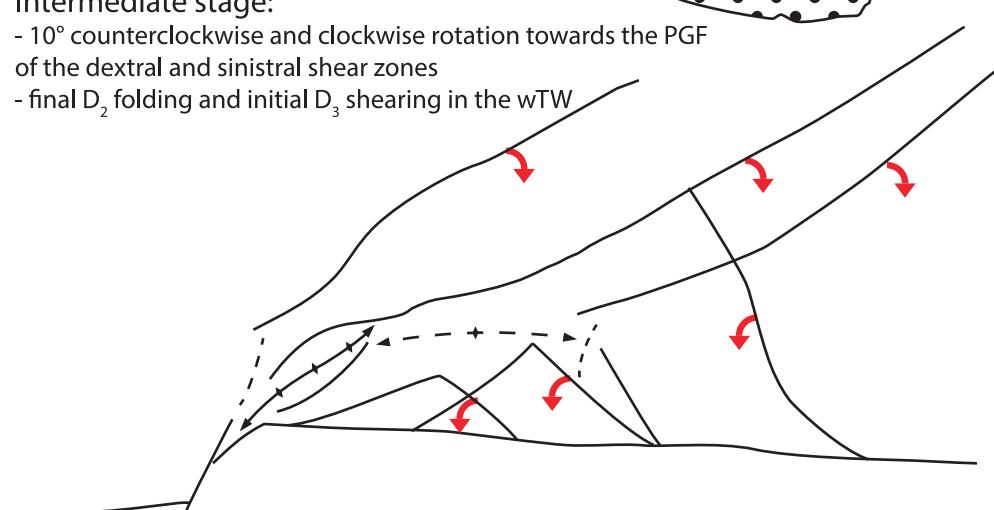
Final stage:

- ca. 20° of shear zone rotation
- ca. +50% E-W stretching of the Tauern Window
- clockwise and counterclockwise blockrotation between sinistral and dextral shear zones
- final D_3 shearing



Intermediate stage:

- 10° counterclockwise and clockwise rotation towards the PGF of the dextral and sinistral shear zones
- final D_2 folding and initial D_3 shearing in the wTW



Initial stage:

- N-S to NNE-SSW directed shortening
- angles between the conjugated shear zones are larger in the west than in the east
- initial D_2 folding in the western TW

