Scaling laws for mixed-heated stagnant-lid convection and application to Europa

Frederic Deschamps¹ and Kenny vilella²

¹Academia Sinica ²Hokkaido University

November 23, 2022

Abstract

Because rocks and ices viscosities strongly depend on temperature, planetary mantles and ice shells are often thought to be animated by stagnant-lid convection. Their dynamics is further impacted by the release of internal heat either through radioactive isotopes decay or tidal dissipation. Here, we quantify the impact of internal heating on stagnant-lid convection. We performed numerical simulations of convection combining strongly temperature-dependent viscosity and mixed (basal and internal) heating in 3D-Cartesian and spherical geometries, and used these simulations to build scaling laws relating surface heat flux, Φ_{surf} , interior temperature, T_{m} , and stagnant lid thickness, d_{lid} , to the system Rayleigh number, heating rate, H, and top-to-bottom viscosity ratio, $\Delta\eta$. These relationships show that T_{m} increases with H but decreases with $\Delta\eta$, while Φ_{surf} increases with H and $\Delta\eta$. Importantly, they also describe heterogeneously heated systems well, provided that the maximum dissipation occurs in hottest regions. For H larger than a critical value H_{crit} , the bottom heat flux turns negative and the system cools down both at its top and bottom. Two additional interesting results are that 1) while the rigid lid stiffens with increasing H, it also thins; and 2) H_{crit} increases with increasing $\Delta\eta$. We then use our scaling laws to assess the impact of tidal heating on Europa's ice shell properties and evolution. Our calculations suggest a shell thickness in the range 20-80 km, depending on viscosity and dissipated power, and show that internal heating has a stronger influence than the presence of impurities in the sub-surface ocean.

Scaling laws for mixed-heated stagnant-lid convection and application to Europa

Frédéric Deschamps¹ and Kenny Vilella²

¹ Institute of Earth Sciences, Academia Sinica, 128 Academia Road Sec. 2, Nangang, Taipei 11529, Taiwan.

² JSPS International Research Fellow, Hokkaido University, Sapporo, Japan.

Corresponding author: Frédéric Deschamps; email: frederic@earth.sinica.edu.tw

Submitted to *Journal of Geophysical Research Planets* May 24th 2021, 42 pages, including 3 tables and 8 figures

Abstract. Because rocks and ices viscosities strongly depend on temperature, planetary mantles and ice shells are often thought to be animated by stagnant-lid convection. Their dynamics is further impacted by the release of internal heat either through radioactive isotopes decay or tidal dissipation. Here, we quantify the impact of internal heating on stagnant-lid convection. We performed numerical simulations of convection combining strongly temperature-dependent viscosity and mixed (basal and internal) heating in 3D-Cartesian and spherical geometries, and used these simulations to build scaling laws relating surface heat flux, Φ_{surf} , interior temperature, T_{m} , and stagnant lid thickness, d_{lid} , to the system Rayleigh number, heating rate, H, and top-to-bottom viscosity ratio, $\Delta \eta$. These relationships show that $T_{\rm m}$ increases with H but decreases with $\Delta\eta$, while Φ_{surf} increases with H and $\Delta\eta$. Importantly, they also describe heterogeneously heated systems well, provided that the maximum dissipation occurs in hottest regions. For H larger than a critical value H_{crit} , the bottom heat flux turns negative and the system cools down both at its top and bottom. Two additional interesting results are that 1) while the rigid lid stiffens with increasing H, it also thins; and 2) H_{crit} increases with increasing $\Delta \eta$. We then use our scaling laws to assess the impact of tidal heating on Europa's ice shell properties and evolution. Our calculations suggest a shell thickness in the range 20-80 km, depending on viscosity and dissipated power, and show that internal heating has a stronger influence than the presence of impurities in the sub-surface ocean.

45 **Plain language summary.** Convection is a mode of heat transfer that is thought to play or 46 have played a key role in the cooling of planetary mantles and ice shells of icy bodies. The 47 convection vigor, efficiency and ability to transport heat are all controlled by the properties of 48 the systems in which it settles. In planetary mantles and ice shells, two important parameters 49 are the variations of viscosity triggered by changes in temperature, which lead to the formation 50 of a rigid lid at the top of the system, and the production of heat within the system, which 51 weakens hot plumes rising from its base. In this article, we assess the combined effects of these 52 two parameters. For this, we perform numerical simulations of convection, from which we 53 deduce quantitative relationships between input and output parameters, the later including 54 internal temperature and surface heat flux. We show that both heat flux and temperature 55 increase with increasing internal heat production, while increasing the thermal viscosity 56 contrast increases heat flux, but reduces temperature. We then apply our relationships to the 57 case of Europa, a moon of Jupiter, and show that the thickness of its ice shell should be in the 58 range 20-80 km.

59

60

61 Key points.

62

• We run simulations of stagnant-lid mixed-heated convection and build temperature and heat
 flux scaling laws from them

- Stagnant lid stiffens and thins with increasing rate of internal heating
- The critical rate of internal heating at which bottom heat flux turns negative increases with
- 67 increasing viscosity ratio
- 68

69 1. Introduction

70 Heat transfer through planetary mantles and ice shells of large icy bodies is controlled by the 71 properties of these systems. Due to the strong temperature-dependence of silicate rocks and ices 72 viscosities, convection within these systems is likely to operate in the so-called stagnant-lid 73 regime (e.g., Christensen, 1984; Moresi and Solomatov, 1995), unless, as in the case of the 74 Earth, specific conditions allow the development of plate tectonics. In stagnant-lid convection, 75 a rigid layer forms at the top of the system as an extension of the top thermal boundary layer 76 (TBL). Because this layer is not mobile and transports heat by conduction, its presence strongly 77 alters heat transfer through the system. Another process altering the ability of convection to 78 transfer heat towards the surface is the production of heat within the system. In systems heated 79 both from their bases and their interiors, hot plumes rising from the bottom TBL get weaker 80 with increasing rate of internal heating, and may not reach the surface if heat production is too 81 high (e.g., Travis and Olson, 1994; Deschamps et al., 2010a). As a result, the amount of heat 82 that can be extracted from regions located beneath the system is reduced. Ultimately, for 83 internal heating rate larger than a critical value, the bottom heat flux turns negative, meaning that the system cools down both from its top and its base. In rocky planets, a source of internal 84 heating is the decay of long-lived radio elements (²³⁵U, ²³⁸U, ²³²Th, and ⁴⁰K). Short-lived 85 elements, mainly ²⁶Al, may have further played a role in the evolution of planetesimals, the 86 87 parent bodies of rocky planets and asteroids. In the case of icy moons, tidal dissipation provides 88 a source of heat within or at the bottom of the ice shell. The amount of heat released, and thus 89 the evolution of the body, depends on its orbital properties and may vary with time (e.g., Tobie 90 et al., 2003, 2005; Roberts and Nimmo, 2008), with internal heating being null or negligible if 91 the body is tidally locked or if it moves on a quasi-circular orbit. Quantifying the influence of 92 internal heating on the ability of rocky mantles and ice shells to transport heat towards the

93 surface is therefore essential to model accurately the long term evolution of icy bodies and94 rocky planets.

95 A convenient way to quantify these effects is to build relationships (or scaling laws) 96 between the key parameters describing thermal evolution (mainly interior temperature and 97 surface heat flux) and the system properties, for instance its rheology, Rayleigh number (which 98 measures the vigor of convection and depends itself on the system physical and thermal 99 properties), and rate of internal heating. Scaling laws may be built from series of numerical 100 simulations of convection, in which one or more parameters are systematically varied. Here, 101 we conduct such a study in the case of mixed-heated systems animated by stagnant-lid 102 convection. In addition to building scaling laws, we parameterize the value of internal heating 103 at which the bottom heat flux turns negative. Finally, we use our results to model the properties 104 and evolution of Europa's outer ice shell.

105

106 2. Numerical model and simulations

We performed numerical simulations of thermal convection for an incompressible, infinite
Prandtl number fluid using StagYY (Tackley, 2008). The fluid is heated both from the bottom
and from within, and the internal heating is controlled by the heat production per unit of mass, *H*. The conservation equations of momentum, mass, and energy are then

111
$$\nabla \bar{\sigma} - \nabla P = -\alpha \rho g T \boldsymbol{e}_{\boldsymbol{z}}$$
(1)

112
$$\nabla \cdot \boldsymbol{v} = 0 \tag{2}$$

113 and
$$\rho C_P \frac{\partial T}{\partial t} = k \nabla \cdot (\nabla T) - \rho C_P \nu \cdot \nabla T + \rho H$$
, (3)

114 where the elements of the deviatoric stress tensor, $\overline{\sigma}$, are $\sigma_{ij} = \eta (\partial v_i / x_j + \partial v_j / x_i)$, *P* is the 115 non-hydrostatic pressure, *v* the velocity, *T* the temperature, e_z the radial unit vector, α , ρ , and 116 C_P , and *k* the fluid thermal expansion, density, heat capacity and thermal conductivity (all 117 assumed constant throughout the system), *g* the gravity acceleration, and η the fluid viscosity, 118 which here varies with temperature. Numerical methods used to solve Eqs. (1) to (3) are detailed 119 in Tackley (2008).

120 The geometry is either 3D-Cartesian or 3D-spherical. In this later case, the spherical shell 121 is modelled with a set of Yin and Yang stripes (Kageyama and Sato, 2004), the shell curvature 122 being controlled by the ratio between its inner and outer radii, $f = r_c/R$. Depending on the 123 curvature and on the effective Rayleigh number, Raeff (defined below), the resolution of each 124 Yin or Yang stripe varies between 192×576 and 512×1536 (corresponding to spherical grids of 125 384×768 to 1024×2048 points), and the radial resolution of the shell varies between 96 and 192 126 points. 3D-Cartesian simulations are performed in boxes with a horizontal to vertical aspect ratio equal to 4 in both x and y directions, and a grid resolution of $128 \times 128 \times 64$ points for Ra_{eff} 127 $< 10^{6}, 256 \times 256 \times 128$ points for $10^{6} \le Ra_{eff} < 10^{8}$, and $384 \times 384 \times 192$ points for $Ra_{eff} \ge 10^{8}$. In 128 129 addition, for both 3D-Cartesian and 3D-spherical cases, the grid is vertically refined at the top 130 and at the bottom of the domain. Overall, this provides a good sampling of plumes and thermal 131 boundary layers, when they exist. The top and bottom boundaries are free slip and isothermal, 132 and reflective boundary conditions are imposed on sidewalls of 3D-Cartesian simulations. In 133 all cases, initial temperature distributions are built from random perturbations superposed on a 134 1D radial adiabatic profile with thin TBLs at top and bottom.

135 Conservation equations are non-dimensionalized with the characteristic properties of the 136 system. Hereafter, non-dimensional quantities are distinguished from their dimensional forms 137 by adding a tilde, ~. We used the thickness of the fluid layer, *D*, as characteristic length, and 138 the super-adiabatic temperature jump across this layer, ΔT , as characteristic temperature. The 139 non-dimensional temperature and internal heating rate are then given by $\tilde{T} = (T - T_{surf})/\Delta T$, 140 where T_{surf} is the surface temperature, and

141
$$\widetilde{H} = \frac{\rho H D^2}{k \Delta T}.$$
 (4)

142 Non-dimentionalization further implies to replace the source term of momentum equation, 143 $\alpha \rho g T$, by the Rayleigh number,

144
$$Ra = \frac{\alpha \rho g \Delta T D^3}{\eta \kappa}, \qquad (5)$$

145 where $\kappa = k/\rho C_P$ is the thermal diffusivity. This number measures the ratio between buoyancy 146 and viscous forces, and is an input parameter of our simulations.

147 The viscosity of ice strongly depends on temperature. Here, we modelled this dependency148 using the Frank-Kamenetskii (FK) approximation,

149
$$\eta = \eta_0 \exp\left[-a_\eta \frac{(T-T_0)}{\Delta T}\right],\tag{6}$$

150 where η_0 and T_0 are the reference viscosity and temperature, and a_{η} a parameter that controls the amplitude of viscosity variations. This approximation overestimates the surface heat flux 151 152 by up to 30 % (e.g., Reese et al., 1999), and it does not account for dependencies of viscosity 153 on strain rate and grain size. Nevertheless, it facilitates the calculations and allows capturing 154 the role of one specific parameter (here, internal heating), since a large number of FK simulations are available in the literature and can be used for comparisons. In the FK 155 approximation, the non-dimensional viscosity, $\tilde{\eta} = \eta/\eta_0$, is given as a function of the non-156 dimensional temperature, \tilde{T} , by 157

158
$$\tilde{\eta} = \exp(-a_{\eta}\tilde{T}).$$
 (7)

The top-to-bottom viscosity ratio, $\Delta \eta = \exp(a_{\eta})$, is an input parameter of our simulations. For viscosity ratios larger than 10⁴, convection generally operates in the so-called stagnant-lid regime (*e.g.*, Christensen, 1984; Davaille and Jaupart, 1993; Moresi and Solomatov, 1995), in which a highly viscous (stagnant) lid develops at the top of the fluid. In this layer, heat is transported by conduction, thus reducing the heat transfer. Experimental rheological laws for ice Ih (Durham et al., 2010) imply that the top-to-bottom viscosity ratios through the outer ice 165 shells of icy bodies are much larger than 10^4 . Convection within these shells, if occuring, should 166 then operate in the stagnant-lid regime.

In most cases, we assumed homogeneous heating, *i.e.*, *H* is constant throughout the system. Tidal dissipation within icy bodies may however depends on viscosity (Tobie et al., 2005), which, in our simulations, varies with temperature. We therefore calculated a few cases with viscosity-dependent internal heating. Following Roberts and Nimmo (2008), we assumed that internal heating is given by

172
$$H = H_0 \left[\frac{\omega \eta / \mu}{1 + (\omega \eta / \mu)^2} \right] / \left[\frac{\omega \eta_{ref} / \mu}{1 + (\omega \eta_{ref} / \mu)^2} \right]$$
(8)

173 where η_{ref} and μ are the reference viscosity and rigidity of ice, H_0 a constant, and ω the orbital 174 frequency. Note that the reference viscosity in Eq. (8) may be different from that defined in Eq. 175 (6), provided that in calculations a correction is applied for consistency. Here, because we 176 assumed that the strongest dissipation occurs close to the melting point of ice, η_{ref} is defined at 177 the bottom of the ice shell (*i.e.*, for $\tilde{T} = 1$). In Eq. (6), by contrast, the reference viscosity η_0 is 178 the surface viscosity (for $\tilde{T} = 0$), which implies $\eta_{ref} = \eta_0 \exp(-a_\eta)$. The non-dimensional 179 internal heating rate may then be written

180
$$\widetilde{H} = \widetilde{H}_0 \left[\frac{\zeta_{ref} \widetilde{\eta} \exp(a_\eta)}{1 + \left(\zeta_{ref} \widetilde{\eta} \exp(a_\eta)\right)^2} \right] / \left[\frac{\zeta_{ref}}{1 + \zeta_{ref}^2} \right]$$
(9)

181 where $\zeta_{ref} = \omega \eta_{ref} / \mu$ and $\tilde{\eta}$ is given by Eq. (7). The viscosity at which dissipation is maximal 182 depends on the exact value of ζ_{ref} . With μ around 4.0×10^9 Pa, orbital period of a few hours to a 183 few days (corresponding to ω in the range $3.0 \times 10^{-5} \cdot 3.0 \times 10^{-6}$ s⁻¹), and $5.0 \times 10^{12} \le \eta_{ref} \le 5.0 \times 10^{14}$ 184 Pas, ζ_{ref} may be chosen in the range $4.0 \times 10^{-3} \cdot 4.0$. Here, we fixed ζ_{ref} to 1, so that the maximum 185 dissipation occurs exactly at η_{ref} . This further implies that dissipation is strongest in hottest 186 regions, including plumes heads, as done in Tobie et al. (2003). 187 Because in our simulations viscosity varies throughout the system, the definition of the 188 Rayleigh number, Ra (Eq. 5), is ambiguous. The input Ra can however be defined at a specific 189 viscosity (or equivalently, a specific temperature), such that it does not vary during the 190 simulations. Here, we prescribed the surface Rayleigh number, Ra_{surf} , defined from the surface 191 viscosity and temperature. In stagnant-lid convection, a better description of the vigor of 192 convection beneath the lid is given by the effective Rayleigh number, Ra_{eff} , calculated with the 193 viscosity at the temperature of the well-mixed interior (or interior temperature), \tilde{T}_m , which is 194 defined as the volume averaged temperature within the adiabatic region. Following Eqs. (6) and 195 (8), $Ra_{\rm eff}$ is given by

196
$$Ra_{eff} = Ra_{surf} \exp(a_{\eta} \tilde{T}_m).$$
(10)

197 Note that \tilde{T}_m , and thus $Ra_{\rm eff}$, are outputs of the simulations.

198 A key output observable is the amount of heat transported to the surface, measured with 199 the heat flux. In mixed-heated systems, the conservation of energy implies that its top and 200 bottom values, Φ_{top} and Φ_{bot} , satisfy

201
$$\Phi_{top} = f^2 \Phi_{bot} + \frac{(1+f+f^2)}{3} H, \qquad (11)$$

202 where f is the ratio between the inner and outer shell radii, equal to 1 in Cartesian geometry. 203 The characteristic heat flux is defined as the conductive heat flux for pure basal heating in 204 Cartesian geometry, $\Phi_{carac} = k \Delta T/D$, such that the non-dimensional form of Eq. (11) is simply obtained by replacing each variable by its non-dimensional equivalent, $\tilde{\Phi}_{top}$, $\tilde{\Phi}_{bot}$ and 205 206 \tilde{H} . Equation (11) indicates that, for a given Φ_{top} , the production of heat within the system lowers 207 the amount of heat that can be extracted from regions located below (for instance, planetary 208 cores). If internal heating is too large, the system cannot extract heat from the bottom but cools 209 down both from its top and its bottom (e.g., Moore 2008; Vilella and Deschamps, 2018), 210 meaning that Φ_{bot} is negative. It is useful to introduce the Urey ratio, measuring the ratio 211 between the internal heat production and the surface heat flux.

212
$$Ur = \frac{(1+f+f^2)}{3} \frac{H}{\Phi_{top}}.$$
 (12)

Eqs. (11) and (12) imply that Ur > 1 if Φ_{bot} is negative, and $0 \le Ur \le 1$ otherwise.

214 Convection operates only if the convective heat flux is larger than the conductive heat 215 flux Φ_{cond} , which, for a mixed heated system, depends on depth (Table S1). Its surface 216 expression is given by

217
$$\Phi_{cond,top} = f \frac{k\Delta T}{D} + (f+2) \frac{\rho HD}{6},$$
 (13)

218 whose non-dimensional form (with respect to the characteristic heat flux) writes

219
$$\widetilde{\Phi}_{cond,top} = f + \frac{(f+2)}{6}\widetilde{H}.$$
(14)

The efficiency of heat transfer is measured with the Nusselt number, Nu, defined as the ratio between the convective and conductive heat flux. Convection operates if Nu > 1. As an example, in Cartesian geometry (f = 1), Nu > 1 requires that the surface non-dimensional convective heat flux, $\tilde{\Phi}_{top}$, is larger than $(1 + \tilde{H}/2)$.

224 Using this setup, we performed 63 simulations in 3D-Cartesian geometry (including 9 225 cases with heterogeneous heating) and 25 in 3D-spherical geometry (Table 1). For comparison, 226 we also listed 5 cases with pure bottom heating taken from Deschamps and Lin (2014). Surface 227 Rayleigh number, top-to-bottom viscosity ratio, and non-dimensional heating rate are taken in the ranges $1 \le Ra_{surf} \le 180$, $10^4 \le \Delta \eta \le 10^8$, and $0.5 \le \widetilde{H} \le 10$ respectively, leading to effective 228 Rayleigh numbers between 2.0×10^5 and 2.0×10^8 . In 3D-spherical cases, the inner-to-outer radii 229 230 ratio is chosen between 0.6 and 0.85. For these ranges of values, the flow is time-dependent and reaches a quasi-stationary state (meaning that output properties, including \tilde{T}_m and $\tilde{\Phi}_{top}$, 231 232 oscillate around constant values) after some time. Output properties are estimated after the 233 quasi-stationary phase has been reached, by time-averaging of each property over several 234 oscillations.

236 3. Flow pattern and thermal structure

3.1 Flow pattern

238 Stagnant-lid convection appears for top-to-bottom viscosity ratios larger than 10⁴ (Moresi and 239 Solomatov, 1995), but its occurrence requires larger viscosity contrasts as the Rayleigh number 240 (Deschamps and Sotin, 2000) or shell curvature (Yao et al., 2014; Guerrero et al., 2018) 241 increases. Stein et al. (2013) proposed two criteria to assess the presence of a stagnant lid. First, 242 a non-dimensional surface velocity, \tilde{v}_{surf} , lower than 1; and second a mobility, M, defined as the ratio between \tilde{v}_{surf} and the root mean square velocity of the whole system, smaller than 243 244 0.01. All our simulations satisfy these criteria (Table 1), and should thus belong to the stagnant-245 lid regime.

246 Figures 1 to 3 show snapshots of temperature fields and associated horizontally averaged 247 profiles for 3D-Cartesian cases with same surface Rayleigh number ($Ra_{surf} = 25$) and viscosity 248 ratio ($\Delta \eta = 10^6$), but different rates of internal heating, and for 3D-spherical cases with f = 0.6, $Ra_{surf} = 16$, $\Delta \eta = 10^6$ and, again, different values of \widetilde{H} . A stagnant lid is clearly visible in all 249 250 cases. A closer examination (section 3.2) indicates that the lid is thinning with increasing \tilde{H} . 251 Internal heating has a strong impact on the flow structure beneath the lid. With increasing \tilde{H} , 252 we observe a trend similar to that reported for isoviscous fluids (e.g., Travis and Olson, 1994; 253 Deschamps et al., 2010a). Plumes are getting thinner, more diffuse and may not reach the 254 bottom of the stagnant lid, indicating that the growth of hot instabilities in the base thermal 255 boundary layer (TBL) is more difficult. The flow is progressively controlled by downwellings and return flow. Importantly, if \tilde{H} is large enough (Figs. 1g-h, and 2c-d), the bottom TBL 256 257 disappears and the heat flux turns negative (Figure 3d and 3f). The system then cools down 258 both at its top and its bottom, and the Urey ratio (Eq. 12) is larger than 1.

260 3.2 Properties of the stagnant lid

We measured the (non-dimensional) thickness of the stagnant lid, \tilde{d}_{lid} , using the method developed by Davaille and Jaupart (1993), in which the base of the stagnant lid is defined by the intersection between the tangent at the point of inflexion of the horizontally averaged profile of vertically advected heat, $\tilde{v}_z \tilde{T}$, with the origin axis ($\tilde{v}_z \tilde{T} = 0$; left plots in Figure 3). The values of \tilde{d}_{lid} we obtained are reported in Table 1. All other parameters being equal, \tilde{d}_{lid} decreases with increasing rate of internal heating, while both \tilde{v}_{surf} and M are decreasing. Increasing internal heating thus results in thinner but stronger stagnant lids.

Because heat is transported by conduction in the stagnant lid, it is possible to derive analytical expressions for the horizontally averaged temperature in this region by solving the conduction heat equation. Assuming that internal heating rate and density are constant and that the surface temperature and heat flux (T_{surf} and Φ_{surf}) are known, the (dimensional) temperature profile is given either by Eq. (S7) in Cartesian geometry, or Eq. (S8) in spherical geometry of (Supporting Information, SI). Note that these expressions are independent of the lid thickness. Their non-dimensional forms are

275
$$\langle \tilde{T} \rangle = \tilde{z} \tilde{\Phi}_{top} - \frac{\tilde{H}}{2} \tilde{z}^2$$
 (15)

276 where \tilde{z} is the non-dimensional depth, and

277
$$< \tilde{T} >= -\frac{\tilde{\Phi}_{top}}{(1-f)} \left[1 - \frac{\tilde{R}}{\tilde{r}} \right] + \frac{\tilde{H}}{6(1-f)^2} \left[2 \left(1 - \frac{\tilde{R}}{\tilde{r}} \right) + \left(1 - \frac{\tilde{r}^2}{\tilde{R}^2} \right) \right],$$
(16)

where $\tilde{r} = (1 - f)^{-1} - \tilde{z}$ and $\tilde{R} = (1 - f)^{-1}$ are the non-dimensional radius and total radius, respectively. Solving heat equation for viscosity-dependent internal heating is more complex in the general case. In our case, however, imposing the maximum dissipation at lowest viscosity implies that dissipation in the lid is close to zero. A good description of the temperature profile within the lid is then obtained by setting $\tilde{H} = 0$ in Eqs. (15) and (16). The horizontally averaged heat flux within the stagnant lid is given by Eqs. (S11) and (S12) of SI, whose non-dimensional versions are

285
$$\widetilde{\Phi}(\widetilde{z}) = \frac{\widetilde{T}_{lid}}{\widetilde{d}_{lid}} + \frac{\widetilde{H}}{2} \left(\widetilde{d}_{lid} - 2\widetilde{z} \right)$$
(17)

286 and

287
$$\widetilde{\Phi}(\widetilde{z}) = \frac{\widetilde{T}_{lid}}{\widetilde{d}_{lid}} f_{lid} \frac{\widetilde{R}^2}{\widetilde{r}^2} - \frac{\widetilde{H}}{6(1-f)} \Big[f_{lid} (1+f_{lid}) \frac{\widetilde{R}^2}{\widetilde{r}^2} - 2\frac{\widetilde{r}}{\widetilde{R}} \Big],$$
(18)

where \tilde{d}_{lid} and \tilde{T}_{lid} are the non-dimensional stagnant lid thickness and basal temperature, respectively, and $f_{lid} = (R - d_{lid})/R = 1 - (1 - f) d_{lid}/D$ is the ratio between the radius of its base and the total radius. To obtain Eq. (18), it is useful to recall that $\tilde{R} = (1 - f)^{-1}$. Equations (17) and (18) can be used to estimate the temperature at the bottom of the lid as a function of the surface heat flux and stagnant lid thickness. Setting $\tilde{z} = 0$ in Eq. (17) and $\tilde{r} =$ \tilde{R} in Eq. (18), and re-arranging the terms leads to

294
$$\tilde{T}_{lid} = \tilde{d}_{lid} \left(\tilde{\Phi}_{top} - \frac{1}{2} \tilde{H} \tilde{d}_{lid} \right)$$
(19)

in Cartesian geometry, and

296
$$\tilde{T}_{lid} = \frac{\tilde{d}_{lid}}{f_{lid}} \left[\tilde{\Phi}_{top} - \frac{1}{6} \tilde{H} \frac{\left(2 - f_{lid} - f_{lid}^2\right)}{(1 - f)} \right]$$
(20)

in spherical geometry. Values of \tilde{T}_{lid} deduced either from Eq. (19) or Eq. (20) are reported in Table 1.

To check the validity of our approach, we inserted the values of \tilde{d}_{lid} we measured (Table 1) and the values of \tilde{T}_{lid} calculated by Eqs. (19) and (20) in Eqs. (15) and (16), respectively. This provides an excellent description of the top part of the horizontally averaged temperature profiles, corresponding to the stagnant lid (dashed dark red curves in Fig. 3). Note that the values of \tilde{T}_{lid} obtained with Eq. (19) or Eq. (20) are slightly larger than that measured on the horizontally averaged profiles of temperature.

306 4. Scaling laws

307 Reconstructing potential thermal evolutions of planets and satellites with parameterized 308 convection methods requires the knowledge of appropriate relationships between input 309 parameters (Rayleigh number, viscosity ratio, and rate of internal heating) and observables 310 (interior temperature, surface heat flux, stagnant lid thickness), or scaling laws for short. Results 311 from our numerical simulations allow us to build such scaling laws. These are detailed below 312 and summarized in Table 2.

313

314 4.1 Temperature of the well-mixed interior

Numerical simulations indicate that the interior temperature of an isoviscous, mixed-heated fluid is well described by a relationship combining the interior temperature for pure bottom and pure internal heating (Sotin and Labrosse, 1999; Deschamps et al., 2010a). Here, we followed a similar approach and built a scaling that combines the interior temperature for a bottom-heated fluid animated by stagnant-lid convection (Deschamps and Lin, 2014; Yao et al., 2014), and for an internally-heated fluid, leading to

321
$$\tilde{T}_m = 1 - \frac{a_1}{f^{a_2}\gamma} + (c_1 + c_2 f) \left[\frac{(1+f+f^2)}{3}\tilde{H}\right]^{c_4} \frac{1}{Ra_{eff}^{c_3}},$$
 (21)

where parameters a_1 , a_2 and c_1 to c_4 can be obtained by inversion of the \tilde{T}_m predicted by simulations (Table 1), and $\gamma = \Delta T / \Delta T_v$ is the non-dimensional inverse of the viscous temperature scale, ΔT_v , defined as

325
$$\Delta T_{\nu} = \left(-\frac{1}{\eta} \frac{d\eta}{dT}\Big|_{T=T_{m}}\right)^{-1}.$$
 (22)

In the case of Frank-Kamenetskii approximation (Eq. 6), $\gamma = a_{\eta} = \ln(\Delta \eta)$. For consistency with scaling laws obtained for pure bottom heating, we fixed a_1 and a_2 to the values obtained by Yao et al. (2014), $a_1 = 1.23$ and $a_2 = 1.5$. We then performed two separate inversions, for Ur < 1 and 329 Ur > 1, in which we excluded simulations with heterogeneous heating. The inversion method 330 follows the generalized inversion method of Tarantola and Valette (1982), and we assumed relative uncertainties of 0.5 % on observed \tilde{T}_m , accounting for the time-variations of this 331 observable during the steady-state phase. For Ur < 1, the best fitting values are $c_1 = 4.3$, $c_2 = -$ 332 2.8, $c_3 = 0.26$ and $c_4 = 0.96$, with a chi-square of 20 (the total number of experiments used for 333 334 this inversion being 46). The value of c_3 is fairly close to the theoretical value of the Rayleigh 335 number exponent for a purely internally heated fluid, 0.25 (Parmentier and Sotin, 2000). We 336 therefore did an additional inversion in which we fixed c_3 to 0.25, and (for simplicity) c_4 to 1.0, 337 and found $c_1 = 3.5 \pm 0.12$ and $c_2 = -2.3 \pm 0.11$, still with a good chi-square, around 30. We 338 followed a similar procedure for Ur > 1 (28 simulations). In that case, the best fit is obtained 339 for $c_1 = 4.5$, $c_2 = -3.1$, $c_3 = 0.34$ and $c_4 = 1.75$. Fixing, for simplicity, c_3 to 1/3, we obtained c_1 340 $= 4.4 \pm 0.22$, $c_2 = -3.0 \pm 0.17$, and $c_4 = 1.72 \pm 0.02$, with a chi-square of 39. Figure 4a compares 341 modelled and observed values of \tilde{T}_m . Note that the calculations with heterogeneous heating, 342 which were all conducted with Ur < 1 but were not included in the inversion process, are well 343 described by the scaling law for Ur < 1.

Because the effective Rayleigh number, Ra_{eff} , depends on \tilde{T}_m , solving Eq. (21) for \tilde{T}_m 344 345 requires the use of a zero-search method. As a consequence, identifying trends in the variations of \tilde{T}_m with the input model parameters (surface Rayleigh number, rate of internal heating, 346 347 thermal viscosity ratio, and curvature) is not straightforward. However, a close examination of Table 1 indicates that, other parameters being fixed, \tilde{T}_m increases with \tilde{H} and f, but decreases 348 with Ra_{surf} . Changes of \tilde{T}_m with $\Delta \eta$ are more complex (Figure S2a). For \tilde{H} around 0.5-1.0 and 349 higher, \tilde{T}_m first decreases with increasing $\Delta \eta$, reaches a minimum for a value of $\Delta \eta$ that 350 increases with \tilde{H} , and then starts increasing again. For $\tilde{H} < 1$, \tilde{T}_m increases monotonically with 351 352 $\Delta\eta$, as observed for purely bottom heated convection. Figures S1 and S2, built from Eq. (21) further illustrate these trends. Interestingly, for the range of γ expected in ice layers, around 15-20 (section 5.1), and $\tilde{H} > 1$ one expects \tilde{T}_m to decrease with increasing viscosity ratio.

355

356 4.2 Surface heat flux

357 Heat flux through thermal boundary layers (TBL) scales as a power law of the Rayleigh number 358 and of the temperature jump across the TBL (e.g., Moore and Weiss, 1973), implying that in 359 stagnant-lid convection it also scales as the temperature viscous scale. The horizontally 360 averaged non-dimensional surface heat flux may then be written as a function of the Rayleigh 361 number and of the parameter γ (section 4.1), which is, again, equal to $\ln(\Delta \eta)$ in the case of the Frank-Kamenetskii approximation. Figure 4b shows that regardless of \tilde{H} , the surface heat flux 362 observed in our simulations with Ur < 1 is very well described by the scaling obtained by 363 364 Deschamps and Lin (2014) and may thus be written

365
$$\widetilde{\Phi}_{top} = a \frac{Ra_{eff}{}^{b}}{\gamma^{c}}, \qquad (23)$$

where Ra_{eff} is the effective Rayleigh number (Eq. 10), and the constants a, b, and c are equal to 366 1.46, 0.27, and 1.21, respectively. Spherical cases for Ur < 1 also fit well along this 367 368 parameterisation, and do not require small correction for f, as suggested by Yao et al. (2014). 369 A reappraisal of Yao et al. (2014) calculations further shows that for f > 0.6 such a correction is not needed. Note that $\tilde{\Phi}_{top}$ implicitly depends on f through Ra_{eff} , which increases with 370 interior temperature \tilde{T}_m . Because \tilde{T}_m decreases with f, $\tilde{\Phi}_{top}$ also decreases with increasing 371 372 curvature. Interestingly, heat fluxes observed in cases with heterogeneous heating are slightly 373 lower than those predicted by our scaling, but still fit very well along it, suggesting that the 374 distribution of heat within the system does not substantially affect the surface heat flux. For Ur > 1, our calculations indicate that $\tilde{\Phi}_{top}$ also fits well along Eq. (23) with a = 1.57 and values of 375

376 *b* and *c* similar to those for Ur < 1 (Figure 4b). Finally, the bottom heat flux, $\tilde{\Phi}_{bot}$, can easily 377 be calculated by inserting Eq. (23) in the non-dimensional version of Eq. (11).

While increasing \tilde{H} results, of course, in larger $\tilde{\Phi}_{top}$ and smaller $\tilde{\Phi}_{bot}$, the influence of 378 the thermal viscosity ratio, $\Delta \eta$, on $\tilde{\Phi}_{top}$ is less intuitive. The $1/\gamma^c$ term and, if γ is not too high, 379 the decrease of \tilde{T}_m , both lower $\tilde{\Phi}_{bot}$ as $\Delta \eta$ increases. However, the exponential term in the 380 definition of Ra_{eff} (Eq. 10) remains dominant, such that for given values of Ra_{surf} and \tilde{H} , $\tilde{\Phi}_{top}$ 381 increases with increasing $\Delta\eta$ (Figure S2). An interesting consequence is that the Urey ratio (Eq. 382 383 12) decreases with increasing thermal viscosity ratio, as also shown in Table 1. In other words, 384 given the properties (thickness, density, thermal expansion and diffusivity, super-adiabatic 385 temperature jump, gravity acceleration, and rate of internal heating) of a mixed-heated shell 386 animated by stagnant-lid convection, increasing viscosity ratio allows the system to extract 387 more heat from the underlying layer (i.e., the bottom heat flux increases). This somewhat 388 counter-intuitive feature results from the strong increase in $Ra_{\rm eff}$ with increasing $\Delta \eta$, implying 389 that convection in the well-mixed interior gets more vigorous.

390

4.3 Transition between positive and negative bottom heat flux

If internal heating is too large, convection cannot evacuate all the heat produced towards the surface. A fraction of this heat is released at the base of the system, resulting in a negative bottom heat flux, $\tilde{\Phi}_{bot}$. Setting $\tilde{\Phi}_{bot} = 0$ in Eq. (12) provides a criterion for the maximum amount of internal heat that can be transported to the surface as a function of the system properties (Rayleigh number, curvature, and viscosity ratio),

397
$$\widetilde{H}_{crit} = \frac{3a}{(1+f+f^2)\gamma^c} Ra_{eff}^{\ b}$$
(24)

398 Again, because Ra_{eff} depends implicitly (through \tilde{T}_m) on \tilde{H} , Eq. (24) does not have analytical 399 solutions, but can be solved with a zero search method. An additional difficulty in estimating 400 \tilde{H}_{crit} is that, while the scalings obtained for Ur < 1 and Ur > 1 overlap at $\tilde{\Phi}_{bot} = 0$ within the 401 error bars on scaling parameters values, they are not continuous when using the average values 402 of these parameters (Table 2). A simple solution to this problem is to first calculate threshold 403 values of \tilde{H} with both Ur < 1 and Ur > 1 scalings, \tilde{H}_{crit}^{-} and \tilde{H}_{crit}^{+} , respectively, and second to 404 define the value of \tilde{H}_{crit} as the average of these two bounds.

405 We then solved Eq. (24) for Ra_{surf} in the range 0.3-300, $\Delta\eta$ in the range 10^4 - 10^8 , and f406 between 1 (Cartesian geometry) and 0.6, and found that \tilde{H}_{crit} is well described by

407
$$\widetilde{H}_{crit} = \frac{3}{(1+f+f^2)} a_H \exp(c_H \gamma) R a_{surf}^{b_H}, \qquad (25)$$

408 where $a_{\rm H} = 0.184$, $b_{\rm H} = 0.31$, and $c_{\rm H} = 0.19$. Equation (25) provides a convenient way to estimate \tilde{H}_{crit} and is in good agreement with our numerical simulations (Figure 5). It shows that 409 \tilde{H}_{crit} increases with Ra_{surf} , $\Delta\eta$, and curvature (decreasing f). Note that rescaling Eq. (25) implies 410 to multiply each of its member by $k\Delta T/\rho D^2$ (Eq. 4). Because Ra_{surf} is proportional to D^3 , one 411 412 expects the dimensional critical heating rate, H_{crit} , to decrease approximately as 1/D. Thus, the 413 transition to a negative heat flux is reached for lower heating rates in thick layers than in thin 414 layers, unless the thermal viscosity ratio and/or the super-adiabatic temperature jump increase 415 dramatically with D.

Finally, an interesting result is that, because $\tilde{\Phi}_{bot}$ increases with the thermal viscosity ratio $\Delta\eta$ (section 4.2), \tilde{H}_{crit} also increases with $\Delta\eta$. Therefore, given the properties of a mixedheated shell animated by stagnant-lid convection, increasing $\Delta\eta$ allows the system to extract heat from the underlying core up to higher rate of internal heating.

420

421 **4.4** Thickness of the stagnant lid

Following Eqs. (15) and (16), the temperature profile within the lid is not a linear function ofdepth. However, Figure 3 suggests that these profiles are, at first order, well described by a

424 linear function. This, in turn, implies that the thickness of this lid should approximately scales425 as the inverse of the heat flux, leading to

426
$$\tilde{d}_{lid} = a_{lid} \frac{\gamma^c}{Ra_{eff}^b},$$
(26)

where the values of parameters *b* and *c* are identical to those for surface heat flux (b = 0.27 and c = 1.21), and a_{lid} is a constant. Figure 4c shows that Eq. (26) provides a good description of the stagnant lid thickness, with best fit to the measured stagnant lid thicknesses obtained for a value of $a_{\text{lid}} = 0.633 \pm 0.03$ for Ur < 1, and $a_{\text{lid}} = 0.667 \pm 0.01$ for Ur > 1.

431

432 5. Application to Europa

We now use the results obtained in section 4 to estimate the properties and thermal evolution of Europa outer ice shell. Our purpose is not to provide a detailed description of Europa's evolution, since we do not consider time-dependent internal heating based on Europa's orbital evolution, but instead to assess quantitatively the role played by tidal heating within the ice layer. This approach can easily be extended to other bodies, including Pluto, which is today tidally locked but may have experienced tidal heating early in its history.

439 A feature common to many (if not all) large icy bodies of the outer solar System is the 440 persistence of a sub-surface ocean beneath an outer ice Ih shell (e.g., Hussmann et al., 2007). 441 In the case of Europa the presence of a sub-surface ocean is supported by anomalies in its 442 external magnetic field, attributed to an internal magnetic field induced within a sub-surface 443 liquid layer (Khurana et al., 1998). Europa's average density suggests that its rocky core is 444 large, ~ 70 % in volume, corresponding to a radius of ~ 1400 km. Given Europa's gravity acceleration, 1.31 m/s^2 , the pressure at the surface of the core is not large enough to allow the 445 presence of high pressure ices. Europa's radial structure therefore likely consists of a large 446 447 rocky core, surrounded by a liquid layer composed of water and impurities, and an outer ice 448 layer. The exact nature of impurities is still debated. Present species may include salts, in 449 particular magnesium sulfate (MgSO₄) (Vance et al. 2018), and volatile compounds such as 450 ammonia (NH₃), methanol (CH₃OH), and methane (CH₄), which are all predicted to condensate 451 in giant planets environments with amounts up to a few per cent (e.g., Mousis et al., 2009; Deschamps et al., 2010b). The presence of impurities acts as an anti-freeze, opposing or 452 453 delaying the crystallization of the sub-surface ocean. Interestingly, while the exact nature of 454 impurities may affect the sub-surface ocean physical properties, including its density, it does 455 not qualitatively impact the crystallization process, *i.e.*, different species present in different 456 amounts would lead to similar evolution. For instance, Vilella et al. (2020) pointed out that the 457 impact of 30 % MgSO₄ on the liquidus is equivalent to that of 3.5 % NH₃.

458 Our modelling approach is detailed in SI. It is mostly similar to the one used in 459 Deschamps (2021a), except for the treatments of the interior temperature, $T_{\rm m}$, and of the 460 stagnant lid thickness, d_{lid} . Another important difference is that two sets of parameters are used to calculate $T_{\rm m}$ and the surface heat flux, $\Phi_{\rm surf}$, depending on whether the bottom heat flux, $\Phi_{\rm bot}$, 461 462 is positive (Ur < 1) or negative (Ur > 1) (Table 2). Note that instead of solving Eq. (25) to decide which set of parameters to use, we apply a simpler procedure, which accounts for the 463 464 fact that temperature and heat flux scalings are not continuous at Ur = 1. First, we calculate T_m and Φ_{surf} assuming parameter values for Ur < 1. If the corresponding Φ_{bot} calculated with Eq. 465 (11) is negative, we calculate $T_{\rm m}$ and $\Phi_{\rm surf}$ again, but with parameter values for Ur > 1. If the 466 resulting Φ_{bot} turns back to positive, we set arbitrarily its value to zero and recalculate Φ_{surf} and 467 468 $T_{\rm m}$ accordingly.

Physical properties of Europa and ice Ih used for calculations are listed in Table 3, and we considered two possible initial compositions for the subsurface ocean, pure water and a mix of water and ammonia. In this later case, we fixed the initial amount of ammonia, x_{NH3}^{init} , to 3.0 vol%, corresponding to about 2.2 wt%. This value may be considered as an upper (possibly 473 exaggerated) bound, and we chose it to obtain a conservative estimate of the impact of 474 impurities on the ice shell properties and evolution. Concentration in ammonia then increases 475 as the ice layer thickens, since only water crystalizes, while impurities are left in the sub-surface ocean. The reference viscosity, $\eta_{\text{ref}},$ is taken as a free parameter and varied between 10^{12} and 476 10¹⁵ Pas, a range extended from Montagnat and Duval (2000) estimates of polar ice sheet flow. 477 478 Results are presented either for a given rate of heating per mass unit, H, or a given total power 479 dissipated in the ice shell, P_{tide} . For an ice shell thickness D_{ice} , H and P_{tide} are related by (see 480 also Figure S3)

481
$$H = \frac{3P_{tide}}{4\pi R^3 \left[1 - \left(1 - \frac{D_{ice}}{R}\right)^3\right]},$$
 (27)

482 where *R* is the total radius of Europa.

483

484 5.1 Ice shell properties

485 As heat dissipation in the ice shell increases, two transitions may occur. First, at heating rate $H_{\rm crit}$ the heat flux at the bottom of the shell may turn negative, heating up the underlying sub-486 487 surface ocean and delaying its crystallization. Convection can still operate within the shell, but 488 would be driven by downwellings and described with scaling laws for Ur > 1 (section 4). 489 Second, at heating rate H_{melt} the bottom temperature exceeds the water liquidus, triggering 490 melting at the bottom of the shell. This implies that the ice shell cannot be thicker than a critical 491 value, D_{melt} . Local pockets of partial melt may further appear in hottest regions (plumes head), 492 introducing additional complexities that are not accounted for by our modelling (see Vilella et 493 al., 2020 for a discussion on this topic). Here, we estimate H_{melt} by comparing the liquidus of 494 pure water with the ice shell horizontally averaged temperature, which underestimates the 495 presence of local pockets of melt. However, because the inverse of the non-dimensional viscous temperature scale γ , which is here equal to $E\Delta T/RT_m^2$ (SI), is somewhat high, this bias should 496

497 be limited (Vilella et al., 2020). Figure 6 shows that both H_{crit} and H_{melt} decrease with increasing 498 ice layer thickness, D_{ice} . The decrease in H_{crit} is mostly related to the thickening of the ice layer 499 (section 4.3). The decrease in H_{melt} is a consequence of the water liquidus, which is itself 500 decreasing with depth, thus favoring partial melting at lower heating rates. In other words, D_{melt} decreases with increasing H. Taking $H = 10^{-9}$ W/kg and a reference viscosity $\eta_{ref} = 10^{14}$ Pas, 501 502 for instance, D_{melt} is around 45 km, corresponding to a total power of ~ 1.2 TW. Figure S4 503 further indicates that all other parameters being the same, D_{melt} decreases with increasing η_{ref} . 504 As one would expect, in the case of a pure water ocean H_{crit} is very close to H_{melt} . It is also 505 worth noting that the addition of ammonia in the sub-surface ocean moderates the effects of H. 506 allowing slightly thicker ice shells at a given H.

507 Figure 7 plots the surface heat flux, interior temperature, and stagnant lid thickness as a 508 function of the dissipated power, P_{tide} , and for different shell thicknesses. For the two ocean 509 compositions we tested, and independently of the ice shell thickness, both $T_{\rm m}$ and $\Phi_{\rm surf}$ increase 510 with increasing P_{tide} , while the stagnant lid thins. At a given P_{tide} , thicker shells are cooler and 511 transfer less heat, but these changes attenuate as P_{tide} increases. Interestingly, for values of P_{tide} 512 estimated by Hussmann and Spohn (2004), in the range 0.6-1.0 TW, and despite the fact that 513 the bottom heat flux may turn negative (in particular for cases with NH₃ in the ocean), the ice 514 shell may be as thick as 160 km (see also Fig. 6). For slightly larger values, however, D_{melt} 515 sharply decreases with increasing P_{tide} . In the case of a pure water ocean, for instance, it is equal 516 to 120 and 40 km at P_{tide} of 1.1 and 1.3 TW, respectively. Finally, given D_{ice} and P_{tide} , Φ_{surf} 517 decreases with increasing η_{ref} , while T_m increases and the stagnant lid thickens (Figure S5).

518

519 5.2 Thermal evolution

We model the ice shell thermal evolution following the approach of Grasset and Sotin (1996),solving the conservation equation of energy at the boundary between this shell and the sub-

surface ocean (SI). Again, a detailed reconstruction of this evolution would require to couple
Europa's internal and orbital evolutions (Hussmann and Spohn, 2004), implying that the tidal
power dissipated within the shell is time-dependent. Instead, we assumed that the dissipated
heat does not vary with time.

Examples of evolutions for $\eta_{ref} = 10^{14}$ Pas are shown in Figure S6. The ice shell first 526 527 thickens up to a maximum value, and then starts to thin again after a time that depends on input 528 parameters. Note that values of P_{tide} around or larger than 1.5 TW prevents the ocean 529 crystallization. The shell remains thinner than 10 km, and is not animated by convection. Figures 8 and S7 plot the shell properties at time t = 4.55 Gyr as a function of P_{tide} and η_{ref} , 530 531 respectively. As one could expect, increasing P_{tide} and/or η_{ref} reduces the final shell thickness, 532 $D_{\rm ice}$, and increases its internal temperature, $T_{\rm m}$. In addition, the stagnant lid thickness, $d_{\rm lid}$, decreases, and convection shuts off at lower η_{ref} . Dissipated powers around or lower than 0.1 533 534 TW have no or small impact on D_{ice} and T_m , but still influences d_{lid} substantially. If η_{ref} and/or P_{tide} are too small, the ocean crystallizes completely and remains frozen up to 4.55 Gyr. These 535 conclusions hold for both a pure water ocean and for an ocean with $x_{NH3}^{init} = 3.0$ vol%. In this 536 later case, however, full crystallization cannot be completed even at low η_{ref} and/or P_{tide} . 537 538 Furthermore, the effects of impurities are reduced as *H* increases, such that the shell properties 539 get close to those for a pure water ocean. Internal heating therefore appears as a stronger 540 controlling parameter than the presence of impurities. Finally, it is worth noting that for P_{tide} in the range 0.6-1.0 TW, relevant to Europa (Hussmann et al., 2004), and $\eta_{ref} = 10^{14}$ Pas, Europa's 541 ice shell should be thin, around 20-40 km at a maximum (see also Fig. S6). Lower η_{ref} allows 542 thicker shells, for instance, with $\eta_{ref} = 3.0 \times 10^{13}$ Pas, up to 120 km (pure water ocean) or 80 km 543 $(x_{NH3}^{init} = 3.0 \text{ vol}\%).$ 544

546 6. Conclusions and perspectives

The numerical simulations we performed allowed us to quantify the influence of the rate of internal heating, *H*, on stagnant-lid convection, through the determination of scaling laws for interior temperature, surface heat flux, and stagnant lid thickness (Table 2). We observed two different regimes depending on the sign of the bottom heat flux, Φ_{bot} , or equivalently, whether the Urey ratio is smaller or larger than 1. Interestingly, our simulations show that the value of *H* at which Φ_{bot} turns negative increases with increasing thermal viscosity ratio, $\Delta\eta$. Another interesting finding is that, while the stagnant lid stiffens with increasing *H*, it also thins.

554 Our simulations include a few simplifications. The rheology of ices is certainly more 555 complex than the Frank-Kamenetskii law we used. Compared to an Arrhenius-type of law, this 556 approach overestimates heat flux by up to 30 % (e.g., Reese et al., 1999). In addition, different 557 mechanisms may control ice Ih deformation depending on the strain rate and/or the grain size, 558 but are not accounted for in our modelling. A full description of ice viscosity may instead 559 require the definition of a composite viscosity law, as proposed by Harel et al. (2020). Our 560 approach further neglects the possible presence of pockets of partially melted ice. Such pockets may occur in plumes heads, right beneath the stagnant lid, in which case they could trigger the 561 562 formation of chaos and lenticulae regions (Tobie et al., 2003). Melt may also influence the 563 physical properties of ice, in particular its viscosity and density. This would in turn affect the 564 buoyancy of plumes and reduce tidal dissipation, leading to alternate phases of melting and 565 crystallization (Tobie et al., 2003). Vilella et al. (2020) further studied the impact of melt on 566 heat budget, and showed that for internal heating larger than a critical value, heat flux reaches 567 a plateau, as most of the heating is used to generate more melt. While these limitations may 568 quantitatively alter the scaling laws we build, the main trends indicated by our simulations and 569 the conclusions drawn from them should remain unchanged.

570 A full description of Europa's ice shell evolution requires coupling its orbital and thermal 571 evolutions to capture time-variations in tidal heating. By contrast, calculations coupling Io and 572 Europa evolutions suggest that tidal dissipation within Europa's ice shell may have remained 573 fairly constant around 0.6-1.0 TW during the past 4.5 Gyr (Hussmann and Spohn, 2004). If 574 true, our evolution model should provide first order, but relevant estimates of today Europa's ice shell properties. Taking a reference viscosity in the range 3.0×10^{13} - 3.0×10^{14} Pas and 575 576 assuming the presence of impurities, the thickness of this shell should be in the range 20-75 577 km. This is larger than estimates from mechanical studies based on surface geology 578 observations (e.g., Billings and Katternhorn, 2005; Damptz and Dombard, 2011; Silber and 579 Johnson, 2017), but consistent with estimates from thermal evolution models (e.g., Tobie et al., 580 2003; Hussmann and Spohn, 2004; Allu Peddenti and McNamara, 2019; Green et al., 2021) 581 and estimates of the thickness needed to generate melts needed for cryovolcanism (Vilella et 582 al., 2020).

583 In addition to the evolution of icy bodies, our findings may have some implications for 584 the evolution of planetesimals that formed in early in solar System history. These bodies are 585 thought to have reached a few hundreds of kilometers in size and to have differentiated in a 586 core and a mantle. The decay of ²⁶Al may have released huge amounts of heat in their mantles, 587 which may, in turn, have delayed the cooling of their cores. The scaling laws we obtained can 588 be inserted in thermal evolution models of planetesimals as built, for instance, by Kaminski et 589 al. (2020). Of particular importance is the fact that, all other parameters being the same, Φ_{bot} 590 increases with increasing $\Delta \eta$ and turns negative for values of H that increase with $\Delta \eta$. This 591 suggests that, if stagnant-lid convection, triggered by large top-to-bottom temperature jump, 592 operated within the mantles of planetesimals, large amounts of heat released by the decay of 593 ²⁶Al may have helped, rather than prevented, the cooling of planetesimals cores, and possibly 594 the generation of magnetic fields within these cores.

596 Acknowledgments

597 The research presented in this article was supported by the National Science Council of Taiwan 598 (MoST) under grants 108-2116-M-001-017 and 107-2116-M-001-029 (FD), and JSPS 599 KAKENHI Grant JP19F19023 (KV). The data used for generating the figures displayed in this 600 article are available for academic purposes on Academia Sinica institutional repository 601 (Deschamps, 2021b). The code used in this work is not publicly available but was thoroughly 602 described in Tackley (2008).

603

604 References

- Allu Peddinti, D., & McNamara, A. K. (2019). Dynamical investigation of a thickening iceshell: Implications for the icy moon Europa. *Icarus*, *329*, 251–269.
- 607 Billings, S. E., & Kattenhorn, S. A. (2005). The great thickness debate: Ice shell thickness
- models for Europa and comparisons with estimates based on flexure at ridges. *Icarus*, 177,
 397–412.
- 610 Christensen, U.R. (1984). Heat transport by variable viscosity convection and implications for
 611 the Earth's thermal evolution. *Phys. Earth Planet. Inter.*, *35*, 264-282.
- 612 Croft, S.K., Lunine, J.I, & Kargel, J. (1988). Equation of state of ammonia-water liquid:
 613 .derivation and planetological applications. *Icarus*, *73*, 279-293.
- Damptz, A.L., & Dombard, A.J. (2001). Time-dependent flexures of the lithospheres on the
 icy satellites of Jupiter and Saturn. *Icarus*, *216*, 86-88.
- 616 Davaille, A., & Jaupart, C. (1993). Transient high-Rayleigh-number thermal convection with
- 617 large viscosity variations. J. Fluid Mech., 253, 141-166.

- 618 Deschamps, F. (2021a). Stagnant lid convection with temperature-dependent thermal
 619 conductivity and the thermal evolution of icy worlds. *Geophys. J. Int.*, 224, 1870-1890.
- 620 Deschamps, F. (2021b). Scaling laws for mixed-heated stagnant-lid convection and application
- 621 to Europa. Myspace, Academia Sinica.
- 622 https://myspace.sinica.edu.tw/public.php?service=files&t=C3tVH2NGcIc1AQSYmc9GAJUCe
- 623 <u>bt5XxUPwhU0Tu1pJeQN8J3cySepAxRwvuUk-W6Q</u>
- Deschamps, F., & Sotin, C. (2000). Inversion of two-dimensional numerical convection
 experiments for a fluid with a strongly temperature-dependent viscosity. *Geophys. J. Int.*, *143*, 204-218.
- 627 Deschamps, F., Tackley, P.J., & Nakagawa, T. (2010a). Temperature and heat flux scalings for
- 628 isoviscous thermal convection in spherical geometry. *Geophys. J. Int.*, 182, 137-154.
- Deschamps, F., Mousis, O., Sanchez-Valle, C. & Lunine, J.I. (2010b). The role of methanol on
 the crystallization of Titan's primordial ocean. *Astrophys. J.*, 724, 887-894.
- 631 Deschamps, F., & Lin, J.-R. (2014). Stagnant lid convection in 3D-Cartesian geometry: Scaling
- laws and applications to icy moons and dwarf planets. *Phys. Earth Planet. Inter.*, 229, 40-54.
- Grasset, O. & Sotin, C. (1996). The cooling rate of a liquid shell in Titan's interior. *Icarus*, *123*,
 101-112.
- Green, A.P., Montesi, L.G.J., & Cooper, C.M. (2021). The growth of Europa's icy shell:
 convection and crystallization. J. Geophys. Res. Planets, 126, e2020JE006677, doi:
 10.1029/2020JE006677.
- 638 Guerrero, J., Lowman, J.P., Deschamps, F., & Tackley, P.J. (2018). The influence of curvature
- 639 on convection in a temperature-dependent viscosity fluid: implications for the 2D and 3D
- 640 modeling of moons. J. Geophys. Res. Planets, 123, 1863-1880, doi:10.1029/2017JE005497.

- Harel, L., Dumoulin, C., Choblet, G., Tobie, G., & Besserer, J. (2020). Scaling of heat transfer
- 642 in stagnant lid convection for the outer ice shells of icy moons: influence of rheology. *Icarus*,
- 643 *338*, 113448, doi: 10.106/j.icarus.2019.113448.
- Hussmann H. & Spohn, T. (2004). Thermal-orbital evolution of Io and Europa. *Icarus*, 171,
 391-410.
- Hussmann, H., Sotin, C. & Lunine, J.I. (2007). Interiors and evolution of icy satellites, in
 planets and moons, *Treatise on Geophysics*, vol. 10, Elsevier, 509–539.
- Kageyama, A., & Sato, T. (2004). "Yin-Yang grid": an overset grid in spherical geometry. *Geochem. Geophys. Geosys.*, 5, doi: 10.1029/2004GC000734.
- 650 Kaminski, E., Limare, A., Kenda, B., & Chaussidon, M. (2020). Early accretion of
- planetesimals unravelled by the thermal evolution of the parent bodies of magmatic iron
 meteorites. *Earth Planet. Sci. Lett.*, *548*, 116469, doi: 10.1016/j.epsl.2020.116469.
- 653 Khurana, K.K., Kivelson, M.G., Stevenson, D.J., Schubert, G., Russell, C.T., Walker, R.J. &
- Polanskey, C. (1998). Induced magnetic field as evidence for subsurface ocean in Europa and
 Callisto. *Nature*, *395*, 777-780.
- 656 Montagnat, M. & Duval, P. (2000). Rate controlling processes in the creep of polar ice,
- 657 influence of grain boundary migration associated with recrystallization. *Earth Planet. Sci.*658 *Lett.*, 183, 179-186.
- Moore, W.B. (2008). Heat transport in a convecting layer heated from within and below. J. *geophys. Res.*, 113, doi:10.1029/2006JB004778.
- Moresi, L.-N. & Solomatov, V.S. (1995). Numerical investigation of 2D convection with
 extremely large viscosity variations. *Phys. Fluids*, 7, 2154-2162.
- Mousis, O., Lunine, J. I., Thomas, C., Pasek, M., Marboeuf, U., Alibert, Y., Ballenegger, V.,
- 664 Cordier, D., Ellinger, Y., Pauzat, F., & Picaud, S. (2009). Clathration of volatiles in the Solar

- nebula and implications for the origin of Titan's atmosphere. *The Astrophysical Journal*, 691,
 1780-1786.
- Roberts, J. H., & Nimmo, F. (2008). Tidal heating and the long-term stability of a subsurface
 ocean on Enceladus. *Icarus*, *194*, 675–689.
- Silber, E. A., & Johnson, B. C. (2017). Impact crater morphology and the structure of Europa's
 ice shell. *Journal of Geophysical Research Planets*, *122*, 2685–2701.
- Stein, C., Lowman, J.P., & Hansen, U. (2013). The influence of mantle internal heating on
 lithospheric mobility: Implications for super-Earths. *Earth and Planetary Science Letters*, *361*, 448–459.
- 674 Sotin, C., & Labrosse, S. (1999). Three-dimensional thermal convection in an iso-viscous,
- 675 infinite Prandtl number fluid heated from within and from below: applications to the transfer
 676 of heat through planetary mantles. *Phys. Earth Planet. Inter.*, *112*, 171-190.
- Tackley, P.J. (2008). Modelling compressible mantle convection with large viscosity contrasts
- in a three-dimensional spherical shell using the yin-yang grid. *Phys. Earth Planet. Inter.*, 171,
- 679 7-18.
- Tarantola, A. & Valette, B. (1982). Generalized nonlinear inverse problems solved using the
 least square criterion. *Rev. Geophys. Space Phys.*, 20, 219-232.
- Tobie, G., Choblet, G., & Sotin, C. (2003). Tidally heated convection: constraints on Europa's
 ice shell thickness. *J. Geophys. Res.*, *108*, doi: 10.1029/2003JE002099.
- Tobie, G., Mocquet, A., & Sotin, C. (2005). Tidal dissipation within large icy satellites:
 Applications to Europa and Titan. *Icarus*, *177*, 534-549.
- Travis, B. & Olson, P. (1994). Convection with internal sources and turbulence in the Earth's
 mantle. *Geophys. J. Int.*, *118*, 1-19.
- 688 Vance, S.D., Panning, M.P., Stähler, S. & al., (2018). Geophysical investigations of habitability
- 689 in ice-covered ocean worlds. J. Geophys. Res. Planets, 123, 180-205.

- 690 Vilella, K., & Deschamps, F. (2018). Temperature and heat flux scaling laws for isoviscous
 691 infinite Prandtl number mixed heating convection. *Geophys. J. Int.*, 214, 265-281.
- 692 Vilella, K., Choblet, G., Tsao, W.E., & Deschamps, F. (2020). Tidally heated convection and
- 693 the occurrence of melting in icy satellites: application to Europa. J. Geophys. Res. Planets,
- 694 *125*, e2019JE006248, doi: 10.1029/2019JE006248.
- 695 Yao, C., Deschamps, F., Lowman, J.P., Sanchez-Valle, C., & Tackley, P.J. (2014). Stagnant-
- 696 lid convection in bottom-heated thin 3-D spherical shells: influence of curvature and
- 697 implications for dwarf planets and icy moons. J. Geophys. Res. Planets, 119, 1895-1913.

699 Tables

702	Table 1. Simulations of stagnant-lid convection with mixed heating	ng.
	U	\sim

Ra_{surf}	f	Δη	\widetilde{H}	\widetilde{H}_0	Grid size	\tilde{T}_m	$\widetilde{\Phi}_{top}$	$\widetilde{\Phi}_{bot}$	Ur	$rms(\tilde{v})$	\tilde{v}_{surf}	$Ra_{\rm eff}$	\tilde{d}_{lid}	\tilde{T}_{lid}
3D-Cartesian														
16.0	-	10^{4}	4.0		128×128×64	1.075	3.458	-0.543	1.16	26.9	1.2×10 ⁻¹	3.19×10 ⁵	0.316	0.892
32.0	-	10^{4}	2.0		128×128×64	0.969	2.836	0.837	0.71	40.6	3.8×10 ⁻¹	2.40×10^{5}	0.324	0.814
32.0	-	10^{4}	3.0		128×128×64	1.016	3.222	0.223	0.93	39.9	2.9×10 ⁻¹	3.71×10^{5}	0.302	0.835
32.0	-	10^{4}	4.0		128×128×64	1.051	3.678	-0.323	1.09	41.3	2.2×10 ⁻¹	5.12×10^{5}	0.280	0.872
75.0	-	10^{4}	1.5		128×128×64	0.937	3.202	1.702	0.47	76.9	1.01	4.21×10^{5}	0.268	0.804
75.0	-	10^{4}	3.0		128×128×64	0.998	3.668	0.670	0.82	69.7	5.4×10 ⁻¹	7.36×10^{5}	0.249	0.820
75.0	-	10^{4}	5.0		256×256×128	1.059	4.577	-0.422	1.09	73.4	2.5×10^{-1}	1.05×10^{6}	0.221	0.889
17.9	-	3.2×10^{4}	2.0		128×128×64	0.977	2.887	0.887	0.69	52.1	1.3×10^{-1}	4.51×10^{5}	0.323	0.828
17.9	-	3.2×10^{4}	4.0		128×128×64	1.042	3.740	-0.260	1.07	53.5	8.2×10 ⁻²	8.85×10^{5}	0.276	0.880
55.9	-	3.2×10^{4}	0.0		128×128×64	0.874	3.000	3.001	0.00	112.5	7.1×10^{-1}	4.84×10^{5}	0.254	0.762
55.9	-	3.2×10^{4}	1.0		128×128×64	0.922	3.374	2.376	0.30	120.7	5.6×10 ⁻¹	7.92×10^{5}	0.252	0.818
55.9	-	3.2×10^{4}	2.0		256×256×128	0.962	3.649	1.649	0.55	117.4	2.6×10 ⁻¹	1.21×10^{6}	0.249	0.847
55.9	-	3.2×10^{4}	3.0		256×256×128	0.990	3.959	0.959	0.76	104.8	2.0×10^{-1}	1.62×10^{6}	0.233	0.840
55.9	-	3.2×10^{4}	6.0		256×256×128	1.069	5.352	-0.648	1.12	116.2	1.3×10^{-2}	3.65×10^{6}	0.192	0.917
178.9	-	3.2×10^{4}	4.0		256×256×128	0.985	5.344	1.343	0.75	198.9	5.4×10^{-1}	4.92×10^{6}	0.168	0.843
10.0	-	10^{5}	2.0		128×128×64	0.975	2.976	0.977	0.67	69.7	5.7×10 ⁻¹	7.62×10^{5}	0.319	0.849
10.0	-	10^{5}	4.0		256×256×128	1.034	3.818	-0.181	1.05	69.1	4.1×10 ⁻²	1.38×10^{6}	0.273	0.894
10.0	-	10^{5}	6.0		256×256×128	1.110	5.007	-0.994	1.20	96.4	2.4×10^{-1}	3.60×10^{6}	0.224	0.971
31.6	-	10^{5}	0.0		256×256×128	0.891	3.143	3.144	0.00	148.1	3.0×10 ⁻¹	9.06×10^{5}	0.257	0.809
31.6	-	10^{5}	0.492	1.0	256×256×128	0.915	3.276	2.785	0.15	156.4	2.4×10^{-1}	1.18×10^{6}	0.257	0.842
31.6	-	10^{5}	2.0		256×256×128	0.964	3.772	1.774	0.53	148.2	1.2×10^{-1}	2.08×10^{6}	0.244	0.860
31.6	-	10^{5}	2.096	3.0	256×256×128	0.982	3.493	1.397	0.60	132.2	7.2×10^{-2}	2.57×10^{6}	0.246	0.858
31.6	-	10^{5}	4.0		256×256×128	1.006	4.471	0.471	0.89	132.6	7.3×10 ⁻²	3.38×10^{6}	0.214	0.866
31.6	-	10^{5}	5.0		256×256×128	1.028	4.898	-0.101	1.02	133.7	7.2×10 ⁻²	4.35×10^{6}	0.202	0.889
31.6	-	10^{5}	6.0		256×256×128	1.054	5.447	-0.553	1.10	145.7	6.6×10 ⁻²	5.88×10^{6}	0.187	0.915
50.6	-	10^{5}	2.0		256×256×128	0.957	4.236	2.236	0.47	195.8	1.5×10^{-1}	3.08×10^{6}	0.214	0.861
50.6	-	10^{5}	3.0		256×256×128	0.979	4.501	1.502	0.67	169.1	1.3×10^{-1}	3.97×10^{6}	0.201	0.843
50.6	-	10^{5}	3.022	4.0	256×256×128	0.991	4.182	1.159	0.72	170.4	8.9×10 ⁻²	4.58×10^{6}	0.206	0.860
50.6	-	10 ⁵	6.0		256×256×128	1.035	5.710	-0.290	1.05	179.8	1.1×10 ⁻¹	7.60×10^{6}	0.174	0.900

705 Table I (continued)

<i>Ra</i> surf	f	Δη	Ĥ	\widetilde{H}_0	Grid size	\tilde{T}_m	$\widetilde{\Phi}_{top}$	$\widetilde{\Phi}_{bot}$	Ur	$rms(\tilde{v})$	\tilde{v}_{surf}	$Ra_{\rm eff}$	$ ilde{d}_{lid}$	\tilde{T}_{lid}
50.6	_	10 ⁵	8.0		256×256×128	1.089	6.970	-1.029	1.15	226.5	7.9×10 ⁻²	1.41×10^{7}	0.149	0.952
50.6	-	10^{5}	10.0		256×256×128	1.145	8.427	-1.573	1.19	313.9	6.7×10 ⁻²	2.70×10^{7}	0.129	1.007
56.6	-	3.2×10^{5}	4.0		256×256×128	0.980	5.685	1.685	0.70	326.6	9.4×10 ⁻²	1.41×10^{7}	0.161	0.864
56.6	-	3.2×10^{5}	8.0		256×256×128	1.040	7.461	-0.539	1.07	354.8	7.4×10^{-2}	2.99×10^{7}	0.132	0.914
113.1	-	3.2×10^{5}	4.0		256×256×128	0.970	6.600	2.602	0.61	486.3	2.3×10 ⁻¹	2.48×10^{7}	0.137	0.867
10.0	-	10^{6}	1.0		256×256×128	0.940	3.913	2.914	0.31	266.4	4.0×10^{-2}	4.38×10^{6}	0.232	0.880
10.0	-	10^{6}	2.0		256×256×128	0.963	4.162	2.164	0.48	240.1	2.7×10 ⁻²	5.97×10^{6}	0.227	0.894
10.0	-	10^{6}	3.0		256×256×128	0.981	4.455	1.456	0.67	213.7	2.1×10^{-2}	7.74×10^{6}	0.209	0.865
10.0	-	10^{6}	4.0		256×256×128	0.995	4.755	0.754	0.84	216.69	1.5×10^{-2}	9.33×10 ⁶	0.203	0.882
10.0	-	10^{6}	4.251	5.5	256×256×128	1.012	4.442	0.192	0.96	216.81	1.7×10^{-2}	1.18×10^{7}	0.205	0.912
10.0	-	10^{6}	5.0		256×256×128	1.010	5.179	0.178	0.97	217.0	1.8×10^{-1}	1.15×10^{7}	0.190	0.895
10.0	-	10^{6}	6.0		256×256×128	1.030	5.689	-0.312	1.05	227.3	1.5×10^{-1}	1.51×10^{7}	0.177	0.915
10.0	-	10^{6}	8.0		256×256×128	1.082	7.007	-0.994	1.14	308.3	1.4×10^{-1}	3.11×10^{7}	0.150	0.963
25.0	-	10^{6}	0.0		256×256×128	0.912	4.415	4.416	0.00	456.1	2.8×10^{-1}	7.38×10^{6}	0.193	0.850
25.0	-	10^{6}	2.0		256×256×128	0.952	5.234	3.235	0.38	371.5	3.8×10 ⁻²	1.29×10^{7}	0.171	0.868
25.0	-	10^{6}	2.059	3.0	256×256×128	0.959	4.984	2.926	0.41	370.8	4.5×10^{-2}	1.42×10^{7}	0.179	0.892
25.0	-	10^{6}	3.0		256×256×128	0.969	5.426	2.428	0.55	367.6	2.2×10 ⁻²	1.63×10^{7}	0.169	0.873
25.0	-	10^{6}	3.041	4.0	256×256×128	0.977	5.059	2.016	0.60	355.2	3.3×10 ⁻²	1.82×10^{7}	0.176	0.892
25.0	-	10^{6}	4.0		256×256×128	0.981	5.639	1.637	0.71	361.3	4.3×10 ⁻²	1.92×10^{7}	0.165	0.876
25.0	-	10^{6}	4.929	6.0	256×256×128	1.001	5.564	0.635	0.89	361.2	4.6×10 ⁻²	2.54×10^{7}	0.163	0.905
25.0	-	10^{6}	6.0		256×256×128	1.006	6.392	0.394	0.94	366.8	3.6×10 ⁻²	2.71×10^{7}	0.150	0.889
25.0	-	10^{6}	8.0		256×256×128	1.037	7.450	-0.550	1.07	403.1	3.3×10 ⁻²	4.14×10^{7}	0.133	0.922
45.0	-	10^{6}	4.0		256×256×128	0.973	6.377	2.380	0.63	515.9	8.5×10 ⁻²	3.10×10^{7}	0.144	0.875
5.6	-	3.2×10^{6}	4.0		256×256×128	0.992	4.957	0.956	0.81	281.6	1.1×10^{-1}	1.57×10^{7}	0.194	0.886
5.6	-	3.2×10^{6}	8.0		256×256×128	1.065	7.130	-0.870	1.12	377.2	7.4×10 ⁻²	4.71×10^{7}	0.146	0.958
41.9	-	3.2×10^{6}	4.0		256×256×128	0.967	7.495	3.496	0.53	937.6	5.4×10 ⁻²	8.18×10^{7}	0.123	0.892
10.0	-	107	0.0		256×256×128	0.923	5.271	5.278	0.00	964.4	8.9×10 ⁻²	2.94×10^{7}	0.165	0.869
10.0	-	107	2.948	4.0	256×256×128	0.969	6.089	3.142	0.48	753.1	1.9×10^{-2}	6.14×10^{7}	0.152	0.926
10.0	-	107	4.0		256×256×128	0.975	6.568	2.571	0.61	732.8	1.5×10^{-2}	6.76×10^{7}	0.149	0.935
10.0	-	107	8.0		384×384×192	0.940	8.115	0.114	0.99	734.9	1.1×10^{-2}	1.15×10^{8}	0.121	0.920
10.0	-	10^{7}	10.0		384×384×192	1.035	9.325	-0.676	1.07	851.5	1.0×10 ⁻²	1.79×10^{8}	0.108	0.948
3.2	-	10^{8}	0.0		384×384×192	0.934	6.000	5.999	0.00	1228.2	2.0×10 ⁻²	9.46×10^{7}	0.147	0.885
3.2	-	10^{8}	2.871	4.0	384×384×192	0.967	6.857	3.986	0.42	1353.7	4.8×10 ⁻³	1.74×10^{8}	0.137	0.941
3.2	-	10^{8}	4.0		384×384×192	0.971	7.403	3.401	0.54	1328.9	4.0×10 ⁻³	1.88×10^{8}	0.127	0.909

<i>Ra</i> _{surf}	f	Δη	\widetilde{H}	\widetilde{H}_0	Grid size	\tilde{T}_m	$\widetilde{\Phi}_{top}$	$\widetilde{\Phi}_{bot}$	Ur	$rms(\tilde{v})$	\tilde{v}_{surf}	$Ra_{\rm eff}$	$ ilde{d}_{lid}$	\tilde{T}_{lid}
	Spherical													
16.0	0.60	106	4.0		192×576×96×2	0.932	3.970	3.767	0.67	242.7	8.0×10 ⁻²	6.25×10 ⁶	0.221	0.859
16.0	0.60	106	10.0		192×576×128×2	1.034	6.313	-0.607	1.03	307.6	2.2×10 ⁻²	2.55×10^{7}	0.155	0.918
5.1	0.60	10^{7}	4.0		192×576×128×2	0.929	4.389	4.928	0.60	396.0	2.1×10^{-2}	1.64×10^{7}	0.204	0.887
10.0	0.70	10^{6}	8.0		192×576×96×2	0.964	3.964	2.129	0.74	193.4	3.5×10 ⁻²	6.12×10^{6}	0.228	0.861
10.0	0.70	10^{6}	8.0		256×768×128×2	1.035	5.596	-0.498	1.04	234.5	1.7×10^{-2}	1.61×10^{7}	0.177	0.918
3.2	0.70	10^{7}	8.0		256×768×128×2	1.011	5.833	-0.011	1.00	366.1	4.4×10^{-2}	3.78×10^{7}	0.167	0.908
10.0	0.70	10^{7}	2.0		192×576×128×2	0.907	4.914	7.046	0.30	622.9	6.1×10 ⁻²	2.23×10^{7}	0.171	0.854
15.8	0.70	10^{7}	3.0		192×576×128×2	0.917	5.845	7.461	0.37	797.0	5.8×10 ⁻²	4.15×10^{7}	0.149	0.879
3.2	0.75	10^{7}	4.0		256×768×128×2	0.964	4.530	2.568	0.68	346.0	9.2×10 ⁻²	1.78×10^{7}	0.206	0.894
3.2	0.75	107	8.0		256×768×128×2	1.017	6.049	-0.208	1.02	378.8	4.2×10 ⁻²	4.19×10^{7}	0.162	0.915
3.2	0.75	107	10.0		256×768×128×2	1.054	7.188	-0.923	1.07	496.3	3.5×10 ⁻²	7.49×10^{7}	0.141	0.950
10.0	0.75	10^{7}	4.0		256×768×128×2	0.945	5.834	4.883	0.53	660.7	2.0×10^{-2}	4.19×10^{7}	0.155	0.892
10.0	0.75	107	10.0		384×1152×192×2	1.005	7.817	0.192	0.99	716.3	9.1×10 ⁻³	1.10×10^{8}	0.123	0.912
10.0	0.80	10^{6}	2.0		256×768×96×2	0.938	3.617	3.107	0.45	223.4	7.4×10 ⁻²	4.25×10^{6}	0.242	0.859
10.0	0.80	10^{6}	4.0		256×768×96×2	0.977	4.267	1.580	0.76	192.0	3.0×10 ⁻²	7.18×10^{6}	0.213	0.857
10.0	0.80	10^{6}	8.0		256×768×128×2	1.051	6.039	-0.729	1.08	258.1	1.6×10 ⁻²	2.02×10^{7}	0.168	0.931
10.0	0.80	106	10.0		256×768×128×2	1.098	7.289	-1.315	1.12	366.4	1.3×10 ⁻²	3.89×10^{7}	0.144	0.978
32.0	0.80	10^{6}	4.0		256×768×128×2	0.952	5.412	3.368	0.60	404.7	1.7×10^{-1}	1.65×10^{7}	0.167	0.879
3.2	0.80	10^{7}	4.0		256×768×128×2	0.972	4.666	2.205	0.70	359.2	1.1×10^{-2}	2.02×10^{7}	0.201	0.893
3.2	0.80	10^{7}	8.0		256×768×128×2	1.025	6.253	-0.395	1.04	401.8	3.8×10 ⁻³	4.69×10^{7}	0.159	0.922
3.2	0.80	10^{7}	10.0		256×768×128×2	1.058	7.503	-0.979	1.08	557.2	4.6×10 ⁻³	7.99×10^{7}	0.134	0.951
1.0	0.80	10^{8}	4.0		512×1536×192×2	0.966	5.266	3.157	0.62	626.5	2.5×10 ⁻³	5.32×10^{7}	0.181	0.920
3.2	0.80	10^{8}	3.0		512×1536×192×2	0.940	6.482	6.387	0.38	1250.1	1.1×10 ⁻²	1.11×10^{8}	0.144	0.930
10.0	0.85	10^{6}	4.0		256×768×128×2	0.983	4.403	1.345	0.78	201.5	2.6×10 ⁻²	7.86×10^{6}	0.215	0.883
10.0	0.85	10^{6}	8.0		256×768×128×2	1.058	6.253	-0.838	1.10	283.1	1.4×10 ⁻²	2.23×10^{7}	0.164	0.940

Listed parameters are the surface Rayleigh number, Ra_{surf} , the inner-to-outer radii ratio (for spherical cases), *f*, the top-to-bottom thermal viscosity ratio, $\Delta\eta$, the non-dimensional rate of internal heating, \tilde{H} , the constant \tilde{H}_0 (for heterogeneous internal heating cases, Eq. 9), the grid size, the average non-dimensional temperature of the well-mixed interior, \tilde{T}_m , the top and bottom non-dimensional heat fluxes, $\tilde{\Phi}_{top}$ and $\tilde{\Phi}_{bot}$, the Urey ratio, Ur (Eq. 12), the root mean square velocity of the whole system, $rms(\tilde{v})$, the average surface velocity, \tilde{v}_{surf} , the effective Rayleigh number, Ra_{eff} (Eq. 10), the non-dimensional thickness of the stagnant lid, \tilde{d}_{lid} , calculated following the method of Davaille and Jaupart (1993), and the temperature at the base of this lid, \tilde{T}_{lid} , deduced from Eq. (19) or Eq. (20) with observed values of $\tilde{\Phi}_{top}$ and \tilde{d}_{lid} . Calculations with pure bottom heating ($\tilde{H} = 0$) are taken from Deschamps and Lin (2014).

714

Table 1 (continued).

715716 Table 2. Summary of scaling laws

Quantity	Expression	Parameters					
		Symbol	<i>Ur</i> < 1	<i>Ur</i> > 1			
		a_1	1.23	1.23			
		a_2	1.5	1.5			
Interior temperature	$\tilde{T}_m = 1 - a_1 / f^{a_2} \gamma + (c_1 + c_2 f) \left[\tilde{H} (1 + f + f^2) / 3 \right]^{c_4} / R a_{eff}^{c_3}$	<i>C</i> 1	3.5	4.4			
		<i>C</i> 2	-2.3	-3.0			
		<i>C</i> 3	0.25	1/3			
		<i>C</i> 4	1.0	1.72			
		а	1.46	1.57			
Surface heat flux	$\widetilde{\Phi}_{top} = a R a_{eff}^{\ b} / \gamma^c$	b	0.27	0.27			
		С	1.21	1.21			
		$a_{ m lid}$	0.633	0.667			
Stagnant lid thickness	$\tilde{d}_{lid} = a_{lid} \gamma^c / R a_{eff}^{\ b}$	b	0.27	0.27			
		С	1.21	1.21			
		a_{H}	0.	184			
Threshold internal heating	$\widetilde{H}_{crit} = 3a_H \exp(c_H \gamma) Ra_{surf}^{b_H} / (1 + f + f^2)$	$b_{ m H}$	0	.31			
-		$c_{ m H}$	0	.19			

Listed expressions are scaling laws for non-dimensional interior temperature, \tilde{T}_m , surface heat flux, $\tilde{\Phi}_{top}$, stagnant lid thickness, \tilde{d}_{lid} , and internal heating at the transition between positive (Ur < 1) and negative (Ur > 1) bottom heat flux, \tilde{H}_{crit} . In these expressions, \tilde{H} is the internal heating, fthe ratio between inner and outer radii (equal to 1 for Cartesian geometry), Ra_{surf} the surface Rayleigh number, and Ra_{eff} the effective Rayleigh number calculated at $\tilde{T} = \tilde{T}_m$, given by Eq. (10). The parameter γ , controlling the amplitude of viscosity changes with temperature, is given by $\gamma = \Delta T / \Delta T_v$, where ΔT_v is the viscous temperature scale (Eq. 22). Parameter values are inferred by best fitting these expressions to the results of numerical simulations listed in Table 1.

725 **Table 3.** Europa and materials properties

Parameter	Symbol	Unit	Value/Expression	Europa
Ice Ih properties				
Density	ρι	kg/m ³	920	
Thermal expansion	αι	1/K	1.56×10 ⁻⁴	
Thermal conductivity	k_{I}	W/m/K	566.8/T	
Heat capacity	$C_{ m p}$	J/kg/K	7.037T + 185	
Thermal diffusivity	KI	m ² /s	$k/ ho_{ m I}C_{ m p}$	
Latent heat of fusion	L_{I}	kJ/kg	284	
Reference bulk viscosity	η_{ref}	Pa s	$10^{12} - 10^{15}$	
Activation energy	E	kJ/mol	60	
Liquid water/ammonia properties				
Density (water)	$\rho_{\rm W}$	kg/m ³	1000	
Density (ammonia)	PNH3	kg/m ³	734	
Thermal expansion (water)	$\alpha_{\rm w}$	1/K	3.0×10 ⁻⁴	
Heat capacity (water)	$C_{ m w}$	J/kg/K	4180	
Silicate core properties				
Density	ρ _c	kg/m ³	3300	
Thermal diffusivity	ĸ	m^2/s	10-6	
Europa properties				
Total radius	R	km		1561
Core radius	$r_{\rm c}$	km		1400
Gravity acceleration	g	m/s^2		1.31
Surface temperature	$T_{\rm surf}$	Κ		100
Surface thermal conductivity	$k_{ m surf}$	W/m/K		5.7

All data for ice Ih and liquid water properties are similar to that used by Kirk and Stevenson (1987) (see references therein), except liquid ammonia density, which is from Croft et al. (1988), bulk viscosity, which is a free parameter with possible range of values extended from Montagnat and Duval (2000) estimates, and the activation energy, which is taken from the intermediate regime of Durham et al. (2010).

731

732 Figures

733





Figure 1. Snapshots of the temperature field (left) and vertical slices of the residual temperature relative to the temperature at the bottom of the stagnant lid \tilde{T}_{lid} (right) for cases with surface Rayleigh number $Ra_{surf} = 25$, thermal viscosity ratio $\Delta \eta = 10^6$ and different values of the nondimensional rate of internal heating, \tilde{H} . (a-b) $\tilde{H} = 0$ (pure bottom heating), (c-d) $\tilde{H} = 2$, (e- $\tilde{H} = 4$, and (g-h) $\tilde{H} = 8$. Isosurface values are (a) $\tilde{T} = 0.95$, (c) $\tilde{T} = 0.97$, (e) $\tilde{T} = 0.95$, and (g) $\tilde{T} = 1.015$. In the case with $\tilde{H} = 8$ (plots g-h) the bottom heat flux is negative, *i.e.*, the system cools down both at its top and its bottom. Value of \tilde{T}_{lid} are indicated on each panel.



Figure 2. Isosurface of the temperature (left) and polar slices of the residual temperature relative to the temperature at the bottom of the stagnant lid \tilde{T}_{lid} (right) for snapshots of two cases in 3D-spherical geometry with f = 0.6, surface Rayleigh number $Ra_{surf} = 16$, thermal viscosity ratio $\Delta \eta = 10^6$ and two values of the non-dimensional rate of internal heating, $\tilde{H} = 4$, (a-b) and $\tilde{H} = 10$ (c-d). Isosurface values are $\tilde{T} = 0.95$ in plot (a) and $\tilde{T} = 1.03$ in plot (c). In the case with $\tilde{H} = 10$ (plots c-d), the bottom heat flux is negative, *i.e.*, the system cools down both at its top and its bottom. Value of \tilde{T}_{lid} are indicated on each panel.





756 Figure 3. Horizontally averaged profiles of temperature (right plot in each panel) and vertically 757 advected heat flow (left plot) for four cases in 3D-Cartesian geometry (plots a-d) and two cases 758 in 3D-spherical geometry with inner-to-outer radii ratio f = 0.6 (plots e-f). Surface Rayleigh 759 number, Rasurf, is equal to 25 for 3D-Cartesian cases and 16 for spherical cases, and the top-tobottom viscosity ratio is $\Delta \eta = 10^6$ in all cases. The non-dimensional heating rate is (a) $\tilde{H} = 0$, 760 (b) $\tilde{H} = 2$, (c) $\tilde{H} = 4$, (d) $\tilde{H} = 8$, (e) $\tilde{H} = 4$, (f) $\tilde{H} = 10$. The grey areas denote the vertical 761 extension of the stagnant lid. The dashed lines in the plots of advected heat flow show the 762 763 tangent to the point of inflexion, whose intersection with the origin axis defines the bottom of 764 the lid. The dashed dark-red curves in the plots of temperature are determined assuming a 765 conductive temperature profile in the stagnant lid, and are calculated following either $\tilde{T}(\tilde{z}) =$ $\tilde{T}_{lid} z/\tilde{d}_{lid}$ (panel a), Eq. (15) (panels b-d) or Eq. (16) (panel e-f) with values of \tilde{d}_{lid} listed in 766 Table 1, and values of \tilde{T}_{lid} estimated from Eq. (19) or Eq. (20). 767 768



Figure 4. Comparison between observed and modelled output properties. (a) Temperature of the well-mixed interior, \tilde{T}_m . Observed values are listed in Table 1, and modelled values are given by Eq. (21) with parameter values discussed in section 4.1. (b) Surface heat flux, $\tilde{\Phi}_{top}$. Observed values are listed in Table 1, and modelled values are calculated by Eq. (23) with parameter values discussed in section 4.2. (c) Thickness of the stagnant lid, \tilde{d}_{lid} . Observed values are listed in Table 1, and modelled values are calculated by Eq. (26) with parameter values discussed in section 4.4.





Figure 5. Reduced non-dimensional heating rate, $\tilde{H}_{red} = \exp(0.19\gamma) (1 + f + f^2)/3$, as a function of surface Rayleigh number, Ra_{surf} . Blue and red symbols plot our numerical simulations (Table 1) with positive and negative bottom heat flux, respectively, and the dashed curve shows the (reduced) critical rate of internal heating for which the bottom heat flux turns negative, \tilde{H}_{crit} , calculated with Eq. (25).







793 Figure 6. Critical values of internal heating for the transition between a positive and negative 794 bottom heat flux, H_{crit} , and for partial melting of the ice shell, H_{melt} , as a function of the ice shell thickness. Calculations are made with the properties of Europa (Table 3), $\eta_{ref} = 10^{14}$ Pa s, and 795 796 for two possible compositions of the sub-surface ocean, pure water and an initial mix (*i.e.*, for 797 a shell thickness equal to 0) of water and 3.0 vol% ammonia. Dashed parts of the curves indicate 798 that the system is not animated by convection, based on the observation that the convective heat 799 flux is smaller than the corresponding conductive heat flux. The grey dashed curves represent 800 the heating rate for three values of the total power dissipated within the ice shell (values in TW 801 indicated on each curve). 802



804

805 Figure 7. Properties of Europa's outer ice shell as a function of the power dissipated within this 806 shell, and for three selected shell thicknesses (color code). (a) and (d) Surface heat flux. (b) and (e) Interior temperature. (c) and (f) Stagnant lid thickness. Physical properties used for 807 calculations are listed in Table 3, the reference viscosity η_{ref} is equal to 10^{14} Pa s, and two initial 808 809 compositions of the ocean are considered, pure water (left column), and an initial mix of water 810 and 3.0 vol% ammonia (right column). Curves interruptions indicate that the average interior 811 temperature is larger than the liquidus of pure water at this depth. For the cases with ammonia, 812 two regimes occur depending on whether the Urey ratio (Ur, Eq. 12) is smaller or larger than 813 1, leading to discontinuities at $Ur \sim 1$. The grey shaded bands represent the possible range of 814 dissipated power according to Hussman and Spohn (2004).





818 Figure 8. Properties of Europa's outer ice shell at t = 4.55 Gyr as a function of the power dissipated within the shell and for three values of the reference viscosity, η_{ref} (color code). (a) 819 820 and (d) Thickness. (b) and (e) Interior temperature. (c) and (f) Stagnant lid thickness. Physical 821 properties used for calculations are listed in Table 3, and two initial compositions of the ocean are considered, pure water (left column), and an initial mix of water and 3.0 vol% ammonia 822 823 (right column). Dashed parts of the curves indicate that the system is not animated by 824 convection. The grey shaded bands represent the possible range of dissipated power according 825 to Hussman and Spohn (2004).

Supporting information for "Scaling laws for mixed-heated stagnant lid convection and application to Europa"

Frédéric Deschamps¹ and Kenny Vilella² 4 5 6 ¹ Institute of Earth Sciences, Academia Sinica, 128 Academia Road Sec. 2, Nangang, Taipei 7 11529. Taiwan.² JSPS International Research Fellow, Hokkaido University, Sapporo, Japan. 8 9 10 This supporting information provides details on the calculation of radial conductive profiles of 11 temperature and heat flux for a mixed-heated system (section S1 and Table S1), and on the 12 trends predicted by scaling laws for interior temperature and surface heat flux (section S2 and 13 Figures S1 and S2). It further describes the methods used to calculate the ice shell properties (heat flux, interior temperature, and stagnant lid thickness; section S3 and Figures S3 to S5) 14 15 and the thermal evolution of this shell (section S4 and Figures S6 and S7). Our modelling is 16 mostly similar to that used in Deschamps (2021). Major differences are the treatments of the 17 interior temperature and stagnant lid thickness.

18

S1. Temperature and heat flux profiles for stagnant lids inmixed-heated systems

21 S1.1 Temperature and heat flux profiles in conductive mixed-heated systems

22 Radial profiles of temperature and heat flux for a purely conductive system with internal heat

23 production may be obtained by integrating the heat equation, which writes

24
$$\frac{\partial}{\partial z} \left(k \frac{\partial T}{\partial z} \right) + \bar{\rho} H = 0$$
 (S1)

25 in Cartesian geometry, and

26
$$\frac{1}{r^2}\frac{\partial}{\partial r}\left(r^2k\frac{\partial T}{\partial r}\right) + \bar{\rho}H = 0$$
 (S2)

in spherical geometry, where *T* is the temperature, *z* (in Eq. S1) the depth, *r* (in Eq. S2) the radius, *k* the thermal conductivity, $\bar{\rho}$ the density and *H* the heating rate per unit of mass. Considering that *k*, $\bar{\rho}$ and *H* are constant throughout the system, and taking surface and bottom 30 temperatures, T_{surf} and T_{bot} , as boundary conditions, integrations of Eqs. (S1) and (S2) lead to 31 the expressions listed in Table S1 for the temperature and heat flux profiles. Note that in 32 Cartesian geometry, D is the thickness of the domain, and in spherical geometry, R and r_c are the total and core radii, $f = r_c/R$ the ratio between these radii, and $D = (R - r_c)$, again, the 33 34 thickness of the conductive layer. Expressions for radial profiles of heat flux (also listed in 35 Table S1) are obtained by derivating the radial profiles for temperature with respect to either z36 in Cartesian geometry, or r in spherical geometry. In this later case, one may recall that the heat 37 flux is defined as the opposite of the temperature derivative with respect to radius.

In the case of the outer shells of icy bodies, the bottom temperature is known from the liquidus at the bottom of the ice shell. Instead of using T_{bot} as boundary condition, one may use the surface heat flux, Φ_{surf} . This surface heat flux is given by

41
$$\Phi_{surf} = k \frac{\Delta T}{D} + \frac{\bar{\rho} H D}{2}$$
(S3)

42 in Cartesian geometry (z = 0), and, noting that $(2 - f - f^2) = (1 - f)(2 + f)$ and R = 43 D/(1 - f), by

44
$$\Phi_{surf} = k \frac{\Delta T}{D} f + \frac{\overline{\rho} H D}{6} (2+f) , \qquad (S4)$$

45 in spherical geometry (r = R). Temperature profiles then write

46
$$T(z) = T_{surf} + z \frac{\Phi_{surf}}{k} - \frac{\bar{\rho}Hz^2}{2k},$$
 (S5)

47 in Cartesian geometry, and

48
$$T(r) = T_{surf} - \frac{\Phi_{surf}}{k} R\left(1 - \frac{R}{r}\right) + \frac{\bar{\rho}HR^2}{6k} \left[2\left(1 - \frac{R}{r}\right) + \left(1 - \frac{r^2}{R^2}\right)\right]$$
(S6)

49 in spherical geometry.

50

51 S1.2 Application to stagnant lids

52 Depending on whether the bottom temperature, T_{bot} , or the surface heat flux, Φ_{surf} , is known or 53 easier to access, either expressions in Table S1 or Eqs. (S5) and (S6) may be used to describe 54 temperature profiles within conductive systems or conductive layers. These equations may, in 55 particular be used to infer the thermal profile within the rigid lid that forms at the top of a system 56 animated with stagnant-lid convection (section 3.2), which writes

57
$$\langle T \rangle = T_{surf} + z \frac{\Phi_{surf}}{k} - \frac{\bar{\rho}Hz^2}{2k}$$
 (S7)

58 in Cartesian geometry, and

59
$$< T >= T_{surf} - \frac{\Phi_{surf}}{k} R\left(1 - \frac{R}{r}\right) + \frac{\bar{\rho}HR^2}{6k} \left[2\left(1 - \frac{R}{r}\right) + \left(1 - \frac{r^2}{R^2}\right)\right]$$
 (S8)

60 in spherical geometry. If Φ_{surf} is known, Eqs. (S7) and (S8) can be directly used to determine 61 the temperature profiles within the stagnant lid.

62 If the thickness of the stagnant lid, d_{lid} , and the temperature at its bottom, T_{lid} , are specified 63 instead of the surface heat flux, expressions given in Table S1 lead to

$$64 \qquad < T >= T_{surf} + \Delta T_{lid} \frac{z}{d_{lid}} + \frac{\rho Hz}{2k} (d_{lid} - z) \tag{S9}$$

65 in Cartesian geometry, and

$$66 \qquad < T >= T_{surf} - \Delta T_{lid} \frac{R}{d_{lid}} f_{lid} \left(1 - \frac{R}{r}\right) + \frac{\rho H R^2}{6k} \left[f_{lid} \left(1 + f_{lid}\right) \left(1 - \frac{R}{r}\right) + \left(1 - \frac{r^2}{R^2}\right) \right]$$
(S10)

in spherical geometry, where $\Delta T_{lid} = (T_{lid} - T_{surf})$ is the temperature jump across the stagnant lid, and $f_{lid} = (R - d_{lid})/R = 1 - (1 - f) d_{lid}/D$ the ratio between the radius of its base and the total radius. Numerical simulations of stagnant lid convection give easily access to the surface heat flux, while the average temperature at the bottom of the stagnant lid, T_{lid} , is more difficult to estimate. To calculate the temperature profiles within stagnant lids Eqs. (S7) and (S8) are thus handier than Eqs. (S9) and (S10).

Heat flux equations in Table S1 may further be used to estimate the temperature at the
bottom of stagnant lids given the surface heat flux and the lid thickness. In this case, heat flux
writes

76
$$\Phi(z) = k \frac{\Delta T_{lid}}{d_{lid}} + \frac{\bar{\rho}H}{2} (d_{lid} - 2z)$$
(S11)

in Cartesian geometry, and

78
$$\Phi(r) = k \frac{\Delta T_{lid}}{d_{lid}} f_{lid} \left(\frac{R}{r}\right)^2 + \frac{\bar{\rho}Hr}{3} \left[1 - \frac{f_{lid}(1+f_{lid})}{2} \frac{R^3}{r^3}\right]$$
(S12)

in spherical geometry. Taking Eqs. (S11) and (S12) at the surface (z = 0 or r = R), and rearranging the terms, one gets the temperature at the bottom of the lid, $T_{lid} = T_{surf} + \Delta T_{lid}$, as a function of the surface heat flux and stagnant lid thickness, following

82
$$T_{lid} = T_{surf} + \frac{d_{lid}}{k} \left(\Phi_{surf} - \frac{\bar{\rho}Hd_{lid}}{2} \right)$$
(S13)

83 in Cartesian geometry, and

84
$$T_{lid} = T_{surf} + \frac{d_{lid}}{kf_{lid}} \Big[\Phi_{surf} - \frac{\bar{\rho}HR}{6} (2 - f_{lid} - f_{lid}^2) \Big]$$
(S14)

85 in spherical geometry.

86

87 S2. Trends in scaling laws for temperature and heat flux

Supplementary Figures S1 and S2 plot the non-dimensional interior temperature, \tilde{T}_m , and 88 surface heat flux, $\tilde{\Phi}_{top}$, as a function of the input parameters of numerical simulations and 89 90 following scaling laws inferred in sections 4.1 and 4.2 of the main article (Eqs. 21 and 23). Input parameters are the surface Rayleigh number, Ra_{surf} , the ratio between the inner and outer 91 radii of the shell, f (with f = 1 for Cartesian geometry), the non-dimensional rate of internal 92 heating, \tilde{H} , and the non-dimensional inverse of the viscous temperature scale, γ , controlling the 93 94 amplitude of viscosity variations with temperature. The viscosity law follows the Frank-95 Kamenetskii approximation, implying that $\gamma = \ln(\Delta \eta)$, where $\Delta \eta$ is the top-to-bottom viscosity 96 ratio. As discussed in sections 4.1 and 4.2, two sets of parameters are needed to explain the results of the simulations, depending on whether the Urey number, Ur, defined by Eq. (12) of 97 the main text, is smaller or larger than 1. This leads to discontinuities for cases where $Ur \sim 1$. 98

Figure S1 shows that \tilde{T}_m increases with \tilde{H} , as one would expect, but decreases with 99 increasing Ra_{surf} , while $\tilde{\Phi}_{top}$ increases monotically with both \tilde{H} and Ra_{surf} . Interior temperature 100 further decreases as curvature gets larger (f decreases). The amplitude of variations in \tilde{T}_m with 101 f are rather limited compared to variations of \tilde{T}_m with \tilde{H} , but comparable to those induced by 102 changes in $Ra_{surf.}$ Note that $\tilde{\Phi}_{top}$ does not depend explicitly on f (Eq. 23 of main text), but is 103 nevertheless sensitive to this parameter because the effective Rayleigh number, Ra_{eff} (Eq. 10 of 104 main article) depends on temperature. As a consequence, $\tilde{\Phi}_{top}$ decreases with increasing 105 106 curvature, but these variations are relatively limited compared to those induced by changes in Ra_{surf} or \tilde{H} . 107

The influence of γ on \tilde{T}_m is more complex and depends in particular on the value of \tilde{H} 108 (plots a and b in Figure S2). For $\tilde{H} < 1$, \tilde{T}_m monotically increases with γ (and thus with $\Delta \eta$), as 109 observed for stagnant-lid convection with a bottom heated-fluid, i.e., $\tilde{H} = 0$ (e.g., Moresi and 110 Solomatov, 1995; Deschamps and Sotin, 2000). By contrast, for \tilde{H} around 1 and higher, \tilde{T}_m 111 112 first decreases with increasing γ , reaches a minimum value for a value of γ that increases with \widetilde{H} , and starts increasing again. It is also interesting to note that the influence of \widetilde{H} becomes 113 smaller as γ increases, *i.e.*, for high values of γ (typically, larger than 25-30), \tilde{T}_m is mostly 114 115 controlled by γ (and thus by the thermal viscosity contrast) regardless of \tilde{H} . As a consequence, $\tilde{T}_m < 1$ (and thus Ur < 1) for such values of γ , and \tilde{T}_m tends asymptotically to 1 as γ goes to 116 infinity. Finally, plots c and d in Figure S2 indicate that $\tilde{\Phi}_{top}$ increases monotically with γ . As 117 discussed in section 4.2, γ acts on $\tilde{\Phi}_{top}$ directly, through $1/\gamma^c$ and the exponential term defining 118 $Ra_{\rm eff}$, and indirectly through \tilde{T}_m . Both the $1/\gamma^c$ term in Eq. (23) and, if γ is not too large, the 119 decrease in \tilde{T}_m (and thus in Ra_{eff}) lead to a decrease in $\tilde{\Phi}_{top}$ as $\Delta \eta$ gets larger. However, the 120 121 exponential term in the definition of Ra_{eff} is dominant, such that for given values of Ra_{surf} and $\widetilde{H}, \widetilde{\Phi}_{top}$ increases with $\Delta \eta$. Again, it is worth noting that the influence of \widetilde{H} diminishes as γ 122

123 gets larger, and that for high viscosity ratios the value of $\tilde{\Phi}_{top}$ is mostly controlled by the 124 amplitude of these variations.

125

126 S3. Modelling of ice shell properties

127 For applications to Europa, we assumed that the viscosity of ice Ih is described by

128
$$\eta(T) = \eta_{ref} exp\left[\frac{E}{RT_{ref}} \left(\frac{T_{ref}}{T} - 1\right)\right]$$
(S15)

129 where E is the activation energy, R the ideal gas constant, and η_{ref} the reference viscosity at 130 temperature T_{ref} . The reference viscosity is not well constrained. Close to the melting point, *i.e.* 131 for $T_{\rm ref}$ equal to the liquidus temperature of pure water at the bottom of the ice shell, $T_{\rm H2O, bot}$, a range of values based on polar ice sheet creep is 10¹³-10¹⁵ Pa s (Montagnat and Duval, 2000). 132 Here, we considered this parameter as a free parameter and varied it in the range 10¹²-10¹⁵ Pa 133 134 s, extending the range of possible values estimated by Montagnat and Duval (2000). Activation 135 energy is better constrained, with values in the range 49-60 kJ/mol depending on the creep 136 regime (Durham et al., 2010), and around 60 kJ/mol for atomic diffusion (Weertman, 1983). 137 Here, we used E = 60 kJ/mol in all calculations. Under icy moons conditions, ice Ih rheology 138 is likely more complex than the diffusion creep mechanism assumed in Eq. (S15), but it is 139 reasonable to think that the impact of internal heating on ice shell dynamics follows a similar 140 trend for different rheologies.

141

1 Following Eq. (22) and the viscosity law (Eq. S15), the viscous temperature scale is

142
$$\Delta T_{v} = \frac{RT_{m}^{2}}{E}, \qquad (S16)$$

such that the inverse of the non-dimensional viscous temperature scale, $\gamma = \Delta T / \Delta T_v$, which controls the thermal viscosity contrast, is given by

145
$$\gamma = \frac{E\Delta T}{RT_m^2},$$
 (S17)

146 where $\Delta T = (T_{bot} - T_{surf})$ is the top to bottom temperature jump. Still following Eq. (S16), 147 rescaling Eq. (21) of main text gives the interior temperature

148
$$T_m = T_{bot} - \frac{a_1 R}{E f^{a_2}} T_m^2 + (a_1 + a_2 f) \left[\frac{(1 + f + f^2)}{3} \frac{\rho_I H D^2}{k_I \Delta T} \right]^{c_3} \frac{\Delta T}{R a_{eff}^{c_4}},$$
(S18)

where T_{bot} is the bottom temperature defined as the liquidus of the water + impurities system, *H* the internal heating rate per mass unit, ρ_{I} and k_{I} the density and thermal conductivity of the ice Ih, respectively, *D* the thickness of the ice layer, and Ra_{eff} the Rayleigh number calculated with the viscosity temperature T_{m} ,

153
$$Ra_{eff} = \frac{\alpha_I \rho_I g \Delta T D^3}{\eta(T_m) \kappa_I},$$
 (S19)

154 In Eq. (S19), α_I and κ_I are the thermal expansion and thermal diffusivity of ice Ih, and $\eta(T_m)$ is 155 calculated with Eq. (S15). The values of the parameters a_1 , a_2 , and c_1 to c_4 are given in section 156 4.1. Note that parameters c_1 to c_4 have different values depending on whether the Urey ratio 157 (Ur, Eq. 12 of main text) is smaller or larger than 1. It is also worth noting that if the sub-surface 158 ocean is composed of pure water, the bottom temperature T_{bot} is equal to the reference 159 temperature defined in the viscosity law (Eq. S15), but is lower than this reference temperature 160 if impurities (e.g., ammonia) are also present (see next paragraph). Equation (S18) does not 161 have analytical solution, and we solved it following a Newton-Raphson zero-search method.

162 Impurities act as an anti-freeze and may include ammonia (NH₃), methanol (CH₃OH), 163 and salts (e.g., magnesium sulfate, MgSO₄). Here, we more specifically considered ammonia, 164 which is predicted to condensate in giant planets environments with amounts up to a few per 165 cent (Mousis et al., 2009; Deschamps et al., 2010). In the case of Europa, magnesium sulfate 166 may further be an important compound of the ocean (Vance et al. 2018). Qualitatively, however, 167 the evolution of the icy bodies is not significantly impacted by the nature of the impurities, but 168 only by their amount. For instance, Vilella et al. (2020) pointed out that the impact of 30 % 169 MgSO₄ on the liquidus is equivalent to that of 3.5 % NH₃. On another hand, it should be noted

170 that different compositions may impact physical properties of the ocean, in particular its density. Adding 30 % MgSO₄ would increase density by about 150 kg/m³, while 3.5 % NH₃ 171 172 would reduce it. Details on the calculation of the water-ammonia system liquidus can be found 173 in Deschamps and Sotin (2001). Practically, we prescribed the initial fraction of ammonia, 174 corresponding to the concentration of ammonia in the initial ocean. The concentration in 175 ammonia then increases as the ocean starts to freeze, since up to the eutectic composition (equal 176 to 32.2 wt% in the case of NH₃), only water ice crystalizes, while impurities are left in the 177 subsurface ocean, whose volume decreases due to the thickening of the outer ice layer. Note 178 that in phase diagrams, concentrations in impurities are usually measured in wt%. For practical 179 reasons, we perform calculations with the volume fraction, which we correct to weight fraction 180 when determining the liquidus, following (in the case of ammonia)

181
$$x_{NH3}^{wt} = \frac{x_{NH3}^{vol}\rho_{NH3}}{x_{NH3}^{vol}\rho_{NH3} + (1 - x_{NH3}^{vol})\rho_w},$$
(S20)

182 where ρ_w and ρ_{NH3} are the densities of liquid water and ammonia, respectively.

183 The surface heat flux is obtained by rescaling the heat flux scaling law (Eq. 23 of main text) with the characteristic heat flux, $\Phi_{carac} = k_{ref} \Delta T/D$, where k_{ref} is the characteristic 184 185 thermal conductivity. Most reconstruction of icy bodies thermal evolutions used values of k_{ref} in the range 2.0-3.0 W/m/K, corresponding to the conductivity at the temperature of the well 186 187 mixed interior or at the bottom of the shell (e.g., Grasset et al., 1996; Tobie et al., 2003; 188 Běhounková et al., 2010). Here, we fixed k_{ref} to 2.6 W/m/K (Grasset and Sotin, 1996). 189 Interestingly, in the case of Europa, this value leads to ice shell properties and thermal evolution 190 very close to those obtained with temperature-dependent thermal conductivity (Deschamps, 191 2021). Accounting for the shell's curvature, measured with the ratio between the inner and outer 192 radii, f, the basal and surface heat fluxes write

193
$$\Phi_{surf} = \Phi_{carac} \widetilde{\Phi}_{top}$$
(S21)

194 and
$$\Phi_{bot} = \Phi_{carac} \tilde{\Phi}_{top} / f^2$$
. (S22)

195 Note that this formulation is slightly different from that used in Deschamps (2021), where the non-dimensional convective heat flux ($\tilde{\Phi}_{conv}$) was inferred from 3D-Cartesian calculations and 196 a correction for spherical geometry was assumed, leading to $\Phi_{surf} = f \Phi_{carac} \tilde{\Phi}_{conv}$ and 197 $\Phi_{bot} = \Phi_{carac} \tilde{\Phi}_{conv} / f$. Because the curvature of outer ice layers of large icy bodies remains 198 large (typically, f > 0.7), this difference only triggers small to moderate effects on the 199 200 calculations of ice shell properties and thermal evolution. Note that if the surface heat flux is 201 lower than the conductive characteristic heat flux, Φ_{carac} , the system is not animated by 202 convection and transfers heat by conduction. This occurs, for instance, if the ice shell is too thin 203 or, in the case of a sub-surface ocean containing impurities, too thick. In this later case, the 204 temperature at the bottom of the shell is much lower than in the case of a pure water ocean. As 205 a result, reference and interior viscosities are higher, decreasing the vigor of convection or even 206 shutting off convection (Deschamps and Sotin, 2001).

207 As discussed in main text, two sets of parameters for Eq. (23) may be used, depending on 208 whether the bottom heat flux, Φ_{bot} , is positive (Ur < 1) or negative (Ur > 1). The threshold (non-209 dimensional) internal heating is given by Eq. (25) of main text, and may be used as a criteria to 210 decide which set of parameters to use. Here, instead, we used a simpler procedure, which 211 accounts for the fact that temperature and heat flux scalings are not continuous at Ur = 1. First, 212 we calculate the internal temperature $T_{\rm m}$ (Eq. S18) and the surface heat flux, $\Phi_{\rm surf}$, assuming 213 parameter values for Ur < 1. If the corresponding Φ_{bot} (calculated with Eq. (11) of main text) 214 is negative, we re-evaluate $T_{\rm m}$ and $\Phi_{\rm surf}$, but with parameter values for Ur > 1. If the resulting Φ_{bot} is positive again, we set arbitrarily its value to zero, and recalculate Φ_{surf} and T_m 215 216 accordingly.

To calculate the thickness of the stagnant lid, Deschamps (2021) assumed that the temperature at the bottom of the lid is well described by $T_{\text{lid}} = 2T_{\text{m}} - T_{\text{bot}}$, and then deduced d_{lid} from the expression of the conductive temperature profile within the lid. However, the relationship between T_{lid} and T_{m} assumes that temperature jump in the bottom and top thermal boundary layers (excluding the stagnant lid) are equal, which is not valid for mixed-heating convection. Here, instead, we estimated the thickness of the stagnant lid by rescaling Eq. (26) of the main article, leading to

224
$$d_{lid} = \frac{a_{lid}\gamma^c}{Ra_{eff}{}^b}D, \qquad (S23)$$

where γ and Ra_{eff} are given by Eqs. (S17) and (S19), respectively, the constant a_{lid} is equal to 0.633 for Ur < 1 and 0.667 for Ur > 1, b = 0.27, and c = 1.21. The temperature at the bottom of the stagnant lid can then be calculated using Eq. (S14).

228

S4. Thermal evolution

230 The present day radial structure of icy bodies may be estimated from appropriate thermal 231 evolution modelling. Here, we followed the approach of Grasset and Sotin (1996), which 232 calculates the evolution of ice layers thicknesses based on an energy balance accounting for the 233 production of heat in the silicate core, the cooling of the ocean, and the crystallization of ice 234 shells. Europa is not large enough to host high pressure ices, such that the inner radius of the 235 outer ice Ih shell, r_{bot} , can be calculated by solving the energy conservation equation at the 236 boundary between this shell and the sub-surface ocean. Energy conservation at this boundary 237 then writes

238
$$\frac{dr_{bot}}{dt} \left[\rho_w C_w \left(-\frac{\partial T_{ad}}{\partial r} + \frac{\partial T_{bot}}{\partial r} \right) \frac{(r_{bot}^3 - r_c^3)}{3} - \rho_I L_I r_{bot}^2 \right] = r_{bot}^2 \Phi_{bot} - r_c^2 \Phi_c \quad (S24)$$

where *t* is time, T_{bot} and Φ_{bot} are the temperature and heat flux at the bottom of the ice layer, given by the liquidus of the ocean and by Eq. (S22), respectively, r_c is the core radius, Φ_c the heat flux at the top of the core, ρ_w and C_w the liquid water density and heat capacity, ρ_I and L_I the density and latent heat of fusion of ice Ih, respectively, and T_{ad} , the adiabatic temperature in the ocean, given by

244
$$T_{ad}(r) = T_{bot}(r_{bot}) \left[1 - \frac{\alpha_w}{\rho_w c_w} \rho_I g(r - r_{bot}) \right],$$
(S25)

with α_w being the thermal expansion of liquid water. Within the silicate core, heat is assumed to be produced by the decay of 4 radiogenic elements, 40 K, 232 Th, 235 U, and 238 U. The heat flux at the top of the core is then calculated following Kirk and Stevenson (1987) by

248
$$\Phi_c = 2\sqrt{\frac{\kappa_c t}{\pi}} \rho_c \sum_{i=1}^4 C_{0,i} H_i \frac{[1 - \exp(-\lambda_i t)]}{\lambda_i t}, \qquad (S26)$$

249 where κ_c and ρ_c are the thermal diffusivity and density of the silicate core, and the subscript *i* 250 refers to the 4 radiogenic elements, whose properties are listed in Table S2. We solved Eq. 251 (S24) up to t = 4.55 Gyr using an adaptative stepsize control Runge-Kutta method (Press et al., 252 1992), and assuming an initial ice Ih thickness equal to 10 km together with the material and 253 physical properties listed in Table 3 of the main text. Again, because the reference viscosity η_{ref} is a sensitive parameter but is poorly constrained, we performed calculations for values of η_{ref} 254 in the range 10^{12} - 10^{15} Pa s, corresponding to an extended range of the values estimated by 255 256 Montagnat and Duval (2000).

257

258 **References**

- 259 Běhounková, M., Tobie, G., Choblet, G., & Čadek, O. (2010). Coupling mantle convection and
- tidal dissipation : Applications to Enceladus and terrestrial planets, J. Geophys. Res. Planets,

261 *115*, doi: 10.1029/2009JE003564.

- Davaille, A., & Jaupart, C. (1993). Transient high-Rayleigh-number thermal convection with
 large viscosity variations. *J. Fluid Mech.*, 253, 141-166.
- Deschamps, F. (2021). Stagnant lid convection with temperature-dependent thermal
 conductivity and the thermal evolution of icy worlds, *Geophys. J. Int.*, 224, 1870-1890.

- Deschamps, F., & Sotin, C. (2000). Inversion of two-dimensional numerical convection
 experiments for a fluid with a strongly temperature-dependent viscosity. *Geophys. J. Int.*, *143*, 204-218.
- Deschamps, F. & Sotin, C. (2001). Thermal convection in the outer shell of large icy satellites, *J. Geophys. Res.*, *106*, 5107-5121.
- Deschamps, F., Mousis, O., Sanchez-Valle, C. & Lunine, J.I. (2010). The role of methanol on
 the crystallization of Titan's primordial ocean. *Astrophys. J.*, 724, 887-894.
- Durham, W.B., Prieto-Ballesteros, O., Goldsby, D.L. & Kargel, J.S. (2010). Rheological and
 thermal properties of icy minerals, *Space Sci. Rev.*, *153*, 273-298.
- Grasset, O. & Sotin, C. (1996). The cooling rate of a liquid shell in Titan's interior, *Icarus*, *123*,
 101-112.
- Kirk, R.L. & Stevenson, D.J. (1987). Thermal evolution of a differentiated Ganymede and
 implications for surface features, *Icarus*, *69*, 91-134.
- Montagnat, M. & Duval, P. (2000). Rate controlling processes in the creep of polar ice,
 influence of grain boundary migration associated with recrystallization, *Earth Planet. Sci. Lett.*, 183, 179-186.
- Moresi, L.-N. & Solomatov, V.S. (1995). Numerical investigation of 2D convection with extremely large viscosity variations, *Phys. Fluids*, **7**, 2154-2162.
- 284 Mousis, O., Lunine, J. I., Thomas, C., Pasek, M., Marboeuf, U., Alibert, Y., Ballenegger, V.,
- 285 Cordier, D., Ellinger, Y., Pauzat, F., & Picaud, S. (2009). Clathration of volatiles in the Solar
- nebula and implications for the origin of Titan's atmosphere. *The Astrophysical Journal*, 691,
 1780-1786.
- 288 Press, W.H., Flannery, B.P. Teukolsky, S.A. & Vetterling, W.T. (1992). Numerical Recipes,
- 289 2nd editon, Cambridge University Press, pp. 701-725.

- 290 Tobie, G., Choblet, G. & Sotin, C. (2003). Tidally heated convection: constraints on Europa's
- 291 ice shell thickness, J. Geophys. Res., 108, doi: 10.1029/2003JE002099.
- 292 Vilella, K., Choblet, G., Tsao, W.E. & Deschamps, F. (2020). Tidally heated convection and
- the occurrence of melting in icy satellites: application to Europa, J. Geophys. Res. Planets,
- 294 *125*, e2019JE006248, doi: 10.1029/2019JE006248.
- 295 Weertman, J. (1983). Creep deformation of ice, Ann. Rev. Earth Planet. Sci., 11, 215-240.
- 296

Table S1. Relationships for radial profiles of temperature and heat flux for a conductivemixed-heated system.

Quantity	Geometry	Expression
Temperature	Cartesian	$T_{surf} + \Delta T \frac{z}{D} + \frac{\bar{\rho}Hz}{2k}(D-z)$
-	Spherical	$T_{surf} - \Delta T \frac{R}{D} f\left(1 - \frac{R}{r}\right) + \frac{\bar{\rho}HR^2}{6k} \left[f(1+f)\left(1 - \frac{R}{r}\right) + \left(1 - \frac{r^2}{R^2}\right)\right]$
Heat flux	Cartesian	$k\frac{\Delta T}{D} + \frac{\bar{\rho}H}{2}(D-2z)$
-	Spherical	$k\frac{\Delta T}{D}f\left(\frac{R}{r}\right)^2 + \frac{\bar{\rho}Hr}{3}\left[1 - \frac{f(1+f)R^3}{2r^3}\right]$

 $\Delta T = (T_{bot} - T_{surf})$ is the bottom-to-top temperature jump, where T_{surf} and T_{bot} are the surface 302 and bottom temperature and *D* is the thickness of the shell. In Cartesian geometry, *z* is depth, 303 and in spherical geometry, *r* is radius, *R* the total radius, and $f = r_{bot}/R$ the ratio between the 304 inner and outer radii of the shell. *k* is the thermal conductivity, *H* the rate of internal heating, 305 and $\bar{\rho}$ the average density, which are here all considered as being constant.

Table S2. Properties of long-lived radioactive isotopes.

Element	Decay constant, λ (1/yr)	Heat release, <i>H</i> (W/kg)	Initial abundance, C ₀ (ppb)
⁴⁰ K	5.4279×10^{-10}	2.917×10 ⁻⁵	738.0
²³² Th	4.9405×10^{-11}	2.638×10 ⁻⁵	38.7
²³⁵ U	9.8485×10^{-10}	5.687×10 ⁻⁴	5.4
²³⁸ U	1.5514×10^{-10}	9.465×10 ⁻⁵	19.9





Figure S1. Non-dimensional interior temperature \tilde{T}_m deduced from Eq. (21) (top row) and 317 surface heat flux $\tilde{\Phi}_{top}$ calculated from Eq. (23) (bottom row) as a function of the surface 318 Rayleigh number Ra_{surf} (left column) and non-dimensional rate of internal heating \tilde{H} (right 319 320 column), and for several values of the ratio between the inner and outer shell radii f (color code; 321 f = 1 indicates Cartesian geometry). Two sets of parameters for Eqs. (21) and (23) are used, 322 depending on whether the Urey ratio (Ur, Eq. 12) is smaller or larger than 1 (see main article), 323 leading to discontinuities at $Ur \sim 1$. For calculations as a function of Ra_{surf} (left column), \tilde{H} is set to 4, and for calculations as a function of \tilde{H} (right column), Ra_{surf} is equal to 10. In all cases, 324 325 the surface top-to-bottom viscosity ratio is fixed to 10^6 .







Figure S2. Non-dimensional interior temperature \tilde{T}_m deduced from Eq. (21) (top row) and 329 surface heat flux $\tilde{\Phi}_{top}$ calculated from Eq. (23) (bottom row) as a function inverse of the non-330 dimensional viscous temperature scale, $\gamma = \Delta T / \Delta T_{v}$ (see main text), and for several values of 331 332 the non-dimensional rate of internal heating (color code). The viscosity is described by a Frank-333 Kamenetskii law (Eq. 7), such that γ is equal to the logarithm of the top-to-bottom viscosity 334 ratio. Two sets of parameters for Eqs. (21) and (23) are used, depending on whether the Urey 335 ratio (Ur, Eq. 12) is smaller or larger than 1 (see main article) and leading to discontinuities at 336 $Ur \sim 1$. In all cases, the surface Rayleigh number is equal to 10, and geometry is Cartesian (f =337 1).





Figure S3. Rate of internal heating per mass unit as a function of the ice shell thickness and for several values of the total power dissipated in the ice layer (color code). The density of the ice shell is $\rho_I = 920 \text{ kg/m}^3$.





Figure S4. Critical values of internal heating for partial melting of the ice shell, H_{melt} , as a function of the ice shell thickness and for different values of the reference viscosity, η_{ref} . Calculations are made with the properties of Europa (Table 3) and assuming a sub-surface ocean composed of pure water. Dashed parts of the curves indicate that the system is not animated by convection, based on the observation that the convective heat flux is smaller than the conductive heat flux. The grey dashed curves represent the heating rate for three values of the total power dissipated within the ice shell (values in TW indicated on each curve).



356

357 **Figure S5.** Properties of a 40 km thick ice Ih shell as a function of the reference viscosity, η_{ref} , 358 and for several values of the total power dissipated in the ice layer (color code). (a) and (d) 359 Surface heat flux. (b) and (e) Interior temperature. (c) and (f) Stagnant lid thickness. Physical 360 properties used for calculations are listed in Table 3, and two initial compositions of the ocean 361 are considered, pure water (left column), and an initial mix of water and 3.0 vol% ammonia 362 (right column). Curves interruptions indicate that the average interior temperature is larger than 363 the liquidus of pure water at that depth. Two regimes occur depending on whether the Urey ratio (Ur, Eq. 12) is smaller or larger than 1, leading to discontinuities at $Ur \sim 1$. 364



Figure S6. Evolution of the ice shell thickness as a function of time for reference viscosity η_{ref} 369 = 10¹⁴ Pa s and several values of the total power dissipated in the ice layer (color code). The 370 composition of the ocean is (a) pure water, or (b) an initial mix of water and 3.0 vol% ammonia. 371 Note the logarithmic scale for the time axis.





374

Figure S7. Properties of Europa's outer ice shell at t = 4.55 Gyr as a function of the reference viscosity, η_{ref} , and for several values of the total power dissipated in the ice layer (color code). (a) and (d) Ice shell thickness. (b) and (e) Interior temperature. (c) and (f) Stagnant lid thickness. Physical properties used for calculations are listed in Table 3, and two initial compositions of the ocean are considered, pure water (left column), and an initial mix of water and 3.0 vol% ammonia (right column). Dashed parts of the curves indicate that the system is not animated by convection.