# Dynamical coupling between the low-latitude lower thermosphere and ionosphere via the non-migrating diurnal tide as revealed by concurrent satellite observations and numerical modeling

Federico Gasperini<sup>1,1,1</sup>, Irfan Azeem<sup>2,2,2</sup>, Geoff Crowley<sup>3,3,3</sup>, Michael Perdue<sup>4,4,4</sup>, Matthew Depew<sup>5,5,5</sup>, Thomas J. Immel<sup>6,6,6</sup>, Erik Stromberg<sup>2,2,2</sup>, Chad Fish<sup>7,7,7</sup>, Crystal Frazier<sup>1,1,1</sup>, Adam Reynolds<sup>8,8,8</sup>, Anthony Swenson<sup>1,1,1</sup>, Ted Tash<sup>1,1,1</sup>, Russell Gleason<sup>1,1,1</sup>, Ryan Blay<sup>2,2,2</sup>, Jordan Maxwell<sup>1,1,1</sup>, Keith Underwood<sup>1,1,1</sup>, Christian Frazier<sup>1,1,1</sup>, and Scott Jensen<sup>1,1,1</sup>

<sup>1</sup>ASTRA LLC

<sup>2</sup>ASTRA LLC.

<sup>3</sup>Atmospheric and Space Technology Research Associates, 12703 Spectrum Drive, Suite 101,San Antonio, TX 78249, USA. <sup>4</sup>Georgia Institute of Technology

<sup>5</sup>University of Texas at Dallas

<sup>6</sup>University of California, Berkeley

<sup>7</sup>Unknown

<sup>8</sup>ASTRA

November 30, 2022

### Abstract

The diurnal-eastward propagating tide with zonal wavenumber 3 (DE3) has gained significant attention due to its ability to preferentially propagate to the ionosphere and thermosphere (IT) from the tropical troposphere, thus effectively coupling these atmospheric regions. In this work, we demonstrate the existence of a pronounced zonal wavenumber 4 (WN4) structure in the low-latitude ionosphere during May 27 - June 5, 2020 using concurrent in-situ total ion number density measurements from the Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) and the Ionospheric Connection Explorer (ICON) satellites. Temperature observations from the Thermosphere Ionosphere Mesosphere Energetics Dynamics Sounding of the Atmosphere using Broadband Emission Radiometry (TIMED/SABER) instrument near 105 km and output from the Specified Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension (SD/WACCM-X) demonstrate that this global-scale ionospheric WN4 structure is due to DE3 propagating through the lower thermosphere.

# Dynamical coupling between the low-latitude lower thermosphere and ionosphere via the non-migrating diurnal tide as revealed by concurrent satellite observations and numerical modeling

1

2

3

4

20

5	${\bf Federico \ Gasperini^1, \ Irfan \ Azeem^1, \ Geoff \ Crowley^1, \ Michael \ Perdue^2,}$
6	${f Matthew \ Depew^2, \ Thomas \ Immel^3, \ Erik \ Stromberg^1, \ Chad \ Fish^1, \ Crystal}$
7	${\bf Frazier}^1, {\bf Adam \ Reynolds}^1, {\bf Anthony \ Swenson}^1, {\bf Ted \ Tash}^1, {\bf Russell \ Gleason}^1,$
8	${f Ryan}$ Blay <sup>1</sup> , Jordan Maxwell <sup>1</sup> , Keith Underwood <sup>1</sup> , Christian Frazier <sup>1</sup> , Scott
9	$\mathbf{Jensen}^1$
10	<sup>1</sup> Atmospheric and Space Technology Research Associates, 282 Contury Pl #1000, Louisville, CO, USA
10	<sup>2</sup> William B. Hanson Center for Space Sciences. Physics Department. University of Texas at Dallas.
12	Richardson, TX, USA
13	<sup>3</sup> Space Sciences Laboratory, University of California, Berkeley, CA, USA
14	Key Points:
15	• A large-amplitude zonal wavenumber 4 (WN4) structure in the low-latitude ionosphere-
16	thermosphere (IT) is seen in multiple satellite observations
17	• Numerical simulations demonstrate the non-migrating diurnal tide DE3 to be re-

Numerical simulations demonstrate the non-migrating diurnal tide DE3 to be re sponsible for this strong IT WN4 coupling

<sup>•</sup> SORTIE and ICON IVM observations provide useful insights into the influence

of tropical tropospheric deep convection on IT variability

Corresponding author: Federico Gasperini, fgasperini@astraspace.net

### 21 Abstract

The diurnal-eastward propagating tide with zonal wavenumber 3 (DE3) has gained sig-22 nificant attention due to its ability to preferentially propagate to the ionosphere and ther-23 mosphere (IT) from the tropical troposphere, thus effectively coupling these atmospheric 24 regions. In this work, we demonstrate the existence of a pronounced zonal wavenumber 25 4 (WN4) structure in the low-latitude ionosphere during May 27 - June 5, 2020 using 26 concurrent in-situ total ion number density measurements from the Scintillation Obser-27 vations and Response of The Ionosphere to Electrodynamics (SORTIE) and the Iono-28 spheric Connection Explorer (ICON) satellites. Temperature observations from the Ther-29 mosphere Ionosphere Mesosphere Energetics Dynamics Sounding of the Atmosphere us-30 ing Broadband Emission Radiometry (TIMED/SABER) instrument near 105 km and 31 output from the Specified Dynamics Whole Atmosphere Community Climate Model with 32 thermosphere and ionosphere eXtension (SD/WACCM-X) demonstrate that this global-33 scale ionospheric WN4 structure is due to DE3 propagating through the lower thermo-34 sphere. 35

### <sup>36</sup> Plain Language Summary

The extent to which terrestrial weather (below  $\sim 30$  km) can influence the ionosphere 37 and thermosphere (IT) is a fascinating discovery of the last two decades or so. The IT 38 is known to vary significantly from day to day, and this daytoday weather is largely driven 39 by processes originating in the lower atmosphere, especially during periods of quiet so-40 lar activity. Accurate forecasting of the IT variability thus depends on the ability to fore-41 cast the component that originates in the lower atmosphere. Ionospheric variability trans-42 lates to uncertainty in navigation and communications systems, while thermospheric vari-43 ability translates to uncertainty in orbital and reentry predictions. In this work, we present 44 evidence of a large amplitude structure with four longitudinally-distributed peaks in the 45 lower thermosphere and ionosphere associated with the well-known diurnal-eastward prop-46 agating tide with zonal wavenumber 3 (DE3) originating in the tropical troposphere. This 47 is accomplished by using concurrent SORTIE, ICON, and TIMED satellite observations 48 during May 27 - June 5, 2020 and a whole atmosphere model. SORTIE and ICON prove 49 to be excellent observational platforms for studying the influence of terrestrial weather 50 on IT variability. 51

### 52 1 Introduction

The lower atmosphere drives variability in the ionosphere-thermosphere (IT) sys-53 tem through the vertical propagation of waves, including tides, planetary waves, and Kelvin 54 waves. These waves are periodic in time and longitude due to the rotation of the Earth, 55 and interact with the lower IT region to modulate electric fields that map to higher al-56 titudes and redistribute plasma in the 200-1000 km region. Due to the geometry of mag-57 netic field lines near the equator, much of this variability occurs at low latitudes and is 58 driven by waves that are excited by deep convective processes in the tropical troposphere 59 and that propagate upwards into the IT system. Tropical troposphere variability is es-60 sentially mapped to the IT system through a variety of neutral-plasma coupling processes, 61 and over a range of spatial and temporal scales. One important class of global-scale at-62 mospheric waves characterized by periods that are harmonics of a solar day are thermal 63 tides. These atmospheric tides can be generated in different altitudinal regions due to 64 tropospheric latent heating, absorption of tropospheric infrared radiation by water va-65 por, absorption of solar ultraviolet radiation by stratospheric ozone, thermosphere molec-66 ular oxygen absorption of extreme ultraviolet radiation, and wave-wave interactions (e.g., 67 Chapman and Lindzen, 1970; Hagan and Forbes, 2002; Hagan et al., 2007; Liu, 2016). 68 The main pathways responsible for the modulation of the ionosphere by tides are direct 69 propagation of atmospheric tides into the ionosphere and thermosphere (e.g., Hagan et 70 al., 2007; Oberheide et al., 2009) and indirect coupling via the ionosphere E-region dy-71 namo (e.g., Jin et al., 2008; Ren et al., 2010; Wan et al., 2008, 2010, 2012). 72

Several studies investigated the importance of IT variability driven by lower-atmosphere 73 wave sources. Initial work (e.g., Hagan et al., 2007; Jin et al., 2008; Fang et al., 2009; 74 Wan et al., 2010) focused on verifying Sagawa et al.s (2005) and Immel et al.s (2006) sug-75 gestion that the wavenumber-4 (hereafter, WN4) structure seen in satellite-borne Sun-76 synchronous F-region ionospheric data was due to the modulation of dynamo electric fields 77 by non-migrating tides propagating from below, and in particular due to the diurnal eastward-78 propagating tide with zonal wavenumber 3 (s = -3, i.e., DE3). DE3 originates in the trop-79 ical troposphere by latent heat release in deep convective clouds (e.g., Hagan, 1996; Ha-80 gan and Forbes, 2002, Lieberman et al., 2007), and its first equatorially-symmetric Hough 81 mode is the largest component in the lower thermosphere (Oberheide et al., 2009; Truskowski 82 et al., 2014), capable of propagating well into the middle thermosphere (Oberheide et 83 al., 2011; Gasperini et al., 2015, 2017a, 2018) due to its longer vertical wavelength. When 84

-3-

#### manuscript submitted to Geophysical Research Letters

viewed at quasi-fixed local time (LT) from slowly precessing satellites, this feature manifests as a 4peak longitude structure. Subsequent studies examined the LT and seasonal
variations of related WN4 and WN3 structures (e.g., Lin et al., 2007; Liu and Watanabe, 2008; Ren et al., 2009), and investigated the underlying mechanisms in further depth
(e.g., Oberheide et al., 2011; He et al., 2011; Mukhtarov and Pancheva, 2011; Maute et
al., 2012; Chang et al., 2013; Lei et al., 2014; Cho et al., 2015; Onohara et al., 2018).

In this work, we present observational evidence of prominent WN4 coupling between 91 the lower thermosphere near 105 km and the ionospheric F-region at heights near 420 92 km and 590 km during May 27 - June 5, 2020. Our results take advantage of total ion 93 number density (hereafter, ion density) measurements from the Ion Velocity Meter (IVM) 0/ instruments onboard the Scintillation Observations and Response of The Ionosphere to 95 Electrodynamics (SORTIE) and Ionospheric Connection Explorer (ICON) spacecrafts 96 and kinetic temperature observations from the Sounding of the Atmosphere using Broad-07 band Emission Radiometry (SABER) instrument onboard the Thermosphere Ionosphere Mesosphere Energetics Dynamics (TIMED) satellite. We further connect this WN4 lon-99 gitude variability to DE3 and investigate its latitude-height amplitude and phase struc-100 ture using a Whole Atmosphere Community Climate Model with thermosphere and iono-101 sphere eXtension (SD/WACCM-X) simulation with realistic wave forcing imposed by 102 nudging Modern-Era Retrospective analysis for Research and Applications version-2 (MERRA-103 2) reanalysis data in the troposphere and stratosphere. After a brief description of the 104 data, models, and methods (Section 2), we present in detail the observational and mod-105 eling results regarding the WN4 variability and its connection to the DE3 tide (Section 106 3), and conclude with a brief summary (Section 4). 107

108

109

### 2 Data, Models, and Methods

2.1 SORTIE/IVM

SORTIE is a NASA Heliophysics System Observatory (HSO) 6U CubeSat mission to investigate the underlying causes behind the appearance of plasma structures in the F-region ionosphere, leading to equatorial plasma bubbles, and the evolution of these structures after their formation (Crowley et al., 2016). SORTIE was launched onboard Dragon CRS-19 to the International Space Station (ISS), from where it was deployed on February 19, 2020 in a nearly circular orbit near 420 km with  $\sim 51.6^{\circ}$  inclination. It carries

-4-

two science instruments, a miniature IVM (MIVM), and a micro-Planar Langmuir Probe
(Crowley et al., 2016). This study employs SORTIE's IVM Level 2 ion density data product.

119

### 2.2 ICON/IVM

ICON is a NASA HSO mission designed to study the fundamental connections be-120 tween the dynamics of the neutral atmosphere at altitudes between 100 km and 300 km 121 and the charged particle motions at low and middle latitudes from a nearly circular or-122 bit at an altitude near 590 km with  $\sim 27^{\circ}$  inclination (Immel et al., 2018). The ICON 123 payload includes an IVM instrument that provides in situ measurements of the ion drift 124 motions, density, temperature and major ion composition (Heelis et al., 2017). The IVM 125 is comprised of two instruments, the Retarding Potential Analyzer (RPA) and the Drift 126 Meter (DM). This study employs ICON's IVM-A Data Product 2.7 v4 determined from 127 RPA measurements. Preliminary validation work by the ICON IVM team reports ac-128 curacy below  $\pm 10^3$  cm<sup>-3</sup> for this data product (see the acknowledgments for further in-129 formation). 130

131

### 2.3 TIMED/SABER

The SABER instrument was launched onboard the TIMED satellite (also a NASA 132 HSO mission) on December 7, 2001. SABER provides measurements of kinetic temper-133 ature from  $\sim 20$  km to  $\sim 120$  km altitude (Mertens et al., 2001). SABER views the at-134 mospheric limb from an orbit of  $\sim 625$  km altitude and  $\sim 73^{\circ}$  inclination, so that the lat-135 itude coverage on a given day extends from about  $53^{\circ}$  in one hemisphere to about  $83^{\circ}$ 136 in the other. This viewing geometry alternates once every  $\sim 60$  days. Errors in the re-137 trieved temperatures in the 80-105 km region are estimated to be  $\pm 1.5$ -5 K (Garcia-Comas 138 et al., 2008). 139

140

### 2.4 WACCM-X

WACCM-X (Liu et al., 2018) is a configuration of the National Center for Atmospheric Research (NCAR) Community Earth System Model (CESM) (Hurrell et al., 2013) that extends the atmospheric component into the thermosphere, with a model top boundary between 500 and 700 km. The standard spatial resolution of WACCM-X is 1.9° by

-5-

 $2.5^{\circ}$  in latitude and longitude, respectively, and 1/4 scale height vertically above the up-145 per stratosphere. To simulate the lower and middle atmosphere, WACCM-X can be run 146 with specified dynamics (SD) that constrains the troposphere and stratosphere dynam-147 ics using MERRA-2 reanalysis data (Gelaro et al., 2017). This study employs output from 148 an extended SD/WACCM-X version v2.1 simulation that covers the May 27 - June 5, 149 2020 period. The empirical ion convection patterns are specified using the Heelis et al. 150 (1982) empirical model. The MERRA-2 forcing provides a realistic representation of the 151 wave forcing in the lower and middle atmosphere making this simulation an ideal physics-152 based framework for interpreting atmosphere-ionosphere coupling by tides. 153

154

### 2.5 Analysis of Satellite Data

When combined, the ascending and descending nodes of SORTIE and ICON cover 155 24-hour of LT in  $\sim$ 30 days (i.e., each node samples  $\sim$ 4 hours of LT in about 10 days). 156 Figure 1 shows the LT coverage of ICON (blue plus signs) and SORTIE (red squares) 157 IVM measurements at 25°N (panel b), at the equator (panel c), and at 25°S (panel d) 158 during May 27 - June 5, 2020. At the equator, SORTIE's (ICON's) orbit precesses from 159  ${\sim}2$  LT/{{\sim}14} LT ( ${\sim}7$  LT/{{\sim}19} LT) on May 27 to  ${{\sim}22}$  LT/{{\sim}10} LT ( ${{\sim}2}$  LT/{{\sim}14} LT) on 160 June 5 (Figure 1c). As shown in Figure 1a, this 10-day period is characterized by low 161 solar activity (average F10.7  $\sim$ 70) with two minor geomagnetic disturbances (ap $\sim$ 25 on 162 May 30, 2020 and ap $\sim$ 15 on June 1-2, 2020). For this study, ICON and SORTIE IVM 163 ion density are analyzed by both combining and separating ascending and descending 164 node measurements. Wave analysis is performed by least-squares fitting SORTIE IVM 165 ion density in 30-day sliding windows and 3-hourly SD/WACCM-X ion density and tem-166 perature during May 27 - June 5, 2020. Note that no least squares fitting of ICON data 167 is performed here as this analysis is left for future work. 168

### 169 **3 Results**

The global distribution of ion density observed by SORTIE (ICON) IVM near 420 km (590 km) altitude during May 27 - June 5, 2020 is presented in Figure 2. SORTIE daytime (~10-14 LT) and nighttime (~22-2 LT) 10-day averages are shown in panels *a* and *b*, respectively. Figures 2*a'* and 2*b'* show the corresponding ICON daytime (~14-18 LT) and nighttime (~2-6 LT) 10-day averages, respectively. Excellent low-latitude coverage is achieved by both SORTIE (about  $\pm 51.6^{\circ}$  latitude) and ICON (about  $\pm 27.1^{\circ}$  lat-

-6-

itude) during this 10-day period, as evidenced by the distribution of the black dots in 176 panels a-a' and b-b' showing the measurement locations. Ion density values up to  $\sim 4.4 \times 10^5$ 177  $\mathrm{cm}^{-3}$  (~2.2x10<sup>5</sup> cm<sup>-3</sup>) are observed during daytime and up to ~1.7x10<sup>5</sup> cm<sup>-3</sup> (~0.8x10<sup>5</sup>) 178  $\rm cm^{-3}$ ) during nighttime by SORTIE (ICON) IVM. Note that a scaling factor is applied 179 multiplying the ICON ion density by a factor of 2 given ICON's higher mean altitude 180 (i.e.,  $\sim 590$  km versus  $\sim 420$  km for SORTIE). Figures 2a''-2b'' display ascending and de-181 scending node differences for SORTIE and ICON IVM ion density, respectively, each scaled 182 (multiplied) by 0.5. Both SORTIE and ICON show enhanced ion density at low latitudes 183 around 30°S-30°N with a prominent WN4 structure and peaks near 180°E, 100°E, 45°E, 184 and  $100^{\circ}$ W. This type of feature in the low-latitude ionosphere was reported in previ-185 ous satellite-based measurements (e.g., Immel et al., 2006; Lin et al., 2007; Liu and Watan-186 abe, 2008, Liu et al., 2009), as discussed in Section 1. This notable WN4 structure ob-187 served by SORTIE and ICON points to a possible modulation of the background ion den-188 sity by the non-migrating DE3 tide as observed from their slowly precessing orbits. Note 189 that the equatorial ionization anomaly (EIA) is not evident in Figure 2 likely due SOR-190 TIE's and ICON's mean altitude ( $\sim$ 420 km and  $\sim$ 590 km, respectively) being signifi-191 cantly higher than the F-layer peak height ( $\sim 200-350$  km), in agreement with Figure 2 192 of Mukhtarov and Pancheva (2011). 193

It is important to realize that while the variability due to the semidiurnal eastward-194 propagating tide with s = -2 (SE2) and stationary planetary wave 4 (SPW4) would largely 195 be eliminated by the ascending/descending node differences shown in Figures 2a''-2b'', 196 aliasing from the diurnal westward-propagating tide with s = 5 (DW5) and from the ter-197 diurnal eastward-propagating tide with s = -1 (TE1) may also contribute to a WN4 lon-198 gitude structure seen at a constant LT (see Lieberman, 1991; Oberheide et al., 2000; Gasperini 199 et al., 2015, 2017b, 2018, 2020). Thus, to best investigate the origin of the observed WN4 200 structure we (1) combine 30 days of SORTIE data collected during May 27 - June 25, 201 2020 to acquire full 24-hour LT coverage, and (2) use output from a 3-hourly SD/WACCM-202 X simulation. Figure 3 shows the period versus zonal wavenumber amplitude spectra of 203 SORTIE IVM ion density obtained by combining ascending and descending node mea-204 surements during May 27 - June 25, 2020 (panels a-a'') and SD/WACCM-X ion density 205 during May 27 - June 5, 2020 derived using the full model output (panels b-b''). Around 206  $10^{\circ}$ S- $10^{\circ}$ N magnetic latitude (MLAT) both SORTIE (panels a') and WACCM-X (panel 207 b') ion density spectra reveal the existence of a pronounced DE3 component with am-208

plitudes of  $\sim 2 \times 10^4$  cm<sup>-3</sup>. DE3 is found to be nearly absent in both SORTIE and WACCM-209 X ion density at 40°N MLAT (panels a and b) and 40°S MLAT (panels a'' and b''), in 210 agreement with the well-known Kelvin wave behavior of the first-symmetric equatorially-211 trapped Hough mode of DE3 that efficiently propagates to the IT. Also present in the 212 spectra is the migrating DW1 tide, as shown in previous ionospheric observations (e.g., 213 Chang et al., 2013). SW2, DW5 and TE1 amplitudes are found to be negligible (TE1 214 results not shown here), providing confidence to our assertion that the WN4 structure 215 observed by SORTIE and ICON is in fact due to the non-migrating DE3 tide. 216

To examine the latitudinal structure associated with DE3 (and other diurnal tides), 217 Figure 4 shows the MLAT-wavenumber diurnal amplitude spectra of SORTIE IVM (pan-218 els a and a') and SD/WACCM-X (panels b and b') ion densities with and without fil-219 tering the migrating DW1 tide. Similar to the results presented in Figure 3, SORTIE 220 and WACCM-X show general agreement with large DW1 ( $\sim 7x10^4$  cm<sup>-3</sup>) and DE3 ( $\sim 3x10^4$ 221  $cm^{-3}$ ) amplitude variations primarily confined to low latitudes. Along with the pronounced 222 DW1 and DE3, SORTIE spectra also exhibit D0 and DW2 variations (and some DE2), 223 while WACCM-X spectra show DE1 and DW2 variations. Larger DE3 amplitudes (up 224 to  $3.06 \times 10^4$  cm<sup>-3</sup>) are found around  $25^{\circ}$ N-10°S MLAT in the SORTIE spectra (panel 225 4a') consistent with the results shown in Figure 2a'', while WACCM-X maxima are found 226 near 10°N-20°S MLAT. Some differences between the SORTIE and WACCM-X spec-227 tra are not unexpected and are likely due to (1) the different averaging window used (30 228 days for SORTIE and 10 day for WACCM-X; note that a 10-day window for WACCM-229 X is adopted for consistency with the SORTIE and ICON results shown in Figure 2), 230 (2) differences in the longitude-local time sampling between model and observations, and 231 (3) inherent tidal variability observed by SORTIE that is not reproduced by the model. 232 Even with the amplitude suppression associated with the monthly averaging, panel 4a'233 shows SORTIE IVM ion density DE3 amplitudes exceeding  $\sim 3 \times 10^4$  cm<sup>-3</sup>, i.e. over 15% 234 of the observed zonal mean. 235

Additional analysis of the DE3 tide in the ionosphere is performed using a 30-day sliding window of low-MLAT ( $\pm 20^{\circ}$ ) SORTIE IVM ion density during the 4-month interval from May 27, 2020 through September 15, 2020 to investigate its temporal variability. Figure 5 show daily DE3 amplitudes (panel *a*) and phases (panel *a'*) derived using a least squares method and the combined ascending/descending node data (i.e., covering ~24-hr LT). DE3 amplitudes are found to be larger (~1.9x10<sup>4</sup> cm<sup>-3</sup>) around May

-8-

27 - June 25, quickly decreasing to minima  $(10^3 - 10^4 \text{ cm}^{-3})$  near June - July, with largest 242 amplitudes ( $\sim 2.1 \times 10^4 \text{ cm}^{-3}$ ) observed during mid-August through early September 2020. 243 The latitude structure of DE3 presented in Figure 4 and its temporal variation shown 244 in Figure 5a, with largest amplitudes occurring around August-September and a second 245 period of enhanced activity around April-May, is consistent with previous modeling work 246 and observations (e.g., Mukhtarov and Pancheva, 2011; Truskowski et al., 2014; Gasperini 247 et al., 2015). To investigate the altitude-latitude structure of the DE3 tide in the IT re-248 gion, Figure 5b shows the SD/WACCM-X height ( $\sim$ 100-450 km) - MLAT (60°S-60°N) 249 structure of ion density (panels b and b') and temperature (panels c and c') amplitudes 250 and phases during May 27 - June 5, 2020. Ionospheric DE3 amplitudes are found to be 251 largest around the F-layer peak near 200-350 km, with a defined 2-peak structure and 252 maxima of  $\sim 8 \times 10^4$  cm<sup>-3</sup> around 15°N and 15°S associated with the EIA (Appleton, 1946; 253 Balan and Bailey, 1995). Previous studies (e.g., Sagawa et al., 2005; Immel et al., 2006; 254 Wan et al., 2008) showed the EIA to exhibit a WN4 longitudinal variation and it is now 255 well accepted that this structure in the ionospheric F-region can form as a result of the 256 combined effect of the E-region dynamo modulation by the lower thermospheric DE3 and 257 from the direct propagation of DE3 into the F-region. It would be beyond the scope of 258 this study to investigate the relative contribution of these two effects to the ionospheric 259 DE3 observed by SORTIE and ICON and this effort is left for follow-on work. The lower 260 thermospheric temperature DE3 exhibits largest amplitudes of  $\sim 12$  K around 110-120 261 km, in accord with previous modeling (e.g., Gasperini et al., 2015, 2017a) and observa-262 tional (e.g., Truskowski et al., 2014) results. The vertical wavelength of the modeled tem-263 perature DE3 inferred from its vertical phase progression (panel c') is  $\sim 49.8$  km, in line 264 with a predominant first symmetric Hough mode. Latitudinal asymmetries and broad-265 ening of the latitude structure with height in the thermospheric DE3 are likely due to 266 the combined effect of mean winds and dissipation (e.g., Forbes, 2000; Gasperini et al., 267 2015, 2017a). 268

Figure 6 displays maps of TIMED/SABER temperatures near 105 km observed concurrently (i.e., May 27 - June 5, 2020) with the SORTIE and ICON IVM ion density maps presented in Figure 2. Panels a and b show results for the descending (~3-5 LT) and ascending (~15-17 LT) nodes, respectively; while panel c contains results for half ascending/descending node differences. A prominent WN4 structure is found to dominate the low-latitude (±30°) lower thermosphere with observed amplitude maxima of ~30 K peak-

ing near 135°E, 75°E, 45°W, 135°W. Pearson correlation coefficients between the equa-275 torial WN4 structure in TIMED/SABER temperature and SORTIE (ICON) IVM ion 276 density observations are calculated to be r=0.71 (r=0.66). This level of correlation, along 277 with the modeling work discussed in Figure 5, provides strong evidence that the iono-278 spheric WN4 structure observed by SORTIE and ICON is in fact associated with the 279 lower thermospheric DE3. As noted earlier, a more comprehensive follow-on work (be-280 yond the purview of the current investigation) will address the extent to which this DE3 281 signature observed in the ionospheric F-region is a result of the effect of the E-region dy-282 namo modulation by the lower thermospheric DE3 versus the direct propagation of DE3 283 into the F-region. 284

#### 285

### 4 Summary and Conclusions

Results presented above provide a clear picture of a marked longitudinal WN4 vari-286 ation observed during May 27 - June 5, 2020 by SORTIE and ICON IVM in the low-287 latitude ionosphere near 420 km and 590 km, respectively. Taking advantage of output 288 from an SD/WACCM-X simulation nudged with MERRA-2 reanalysis data in the tro-289 posphere and stratosphere and monthly-averaged SORTIE IVM data, we demonstrated 290 that this prominent WN4 structure in the F-region ion density is due to the well-known 291 DE3 tide. This non-migrating tide has gained significant attention due to its ability to 292 preferentially propagate to the IT from the tropical troposphere, thus effectively coupling 293 these atmospheric regions. SORTIE IVM observations and SD/WACCM-X output showed 294 ion density DE3 amplitudes upward of  $3 \times 10^4$  cm<sup>-3</sup> (i.e., over 15% of the zonal mean) 295 at low-latitudes ( $\pm 20^{\circ}$  MLAT), with general agreement between model and observations. 296 Least-squares fitting of monthly-averaged SORTIE IVM ion density during May 27 - Septem-297 ber 15, 2020 showed larger ( $\sim 1.9 \times 10^4$  cm<sup>-3</sup>) DE3 amplitudes around May 27 - June 25, 298 rapidly decreasing to smaller values  $(10^3 - 10^4 \text{ cm}^{-3})$  around June - July, and becoming 299 largest ( $\sim 2.1 \times 10^4 \text{ cm}^{-3}$ ) around mid-August through early September 2020, in agree-300 ment with the well-known seasonal variation of DE3 amplitudes found in previous mod-301 eling and observational studies (e.g., Mukhtarov and Pancheva, 2011; Truskowski et al., 302 2014; Gasperini et al., 2015, 2017a). Concurrent TIMED/SABER temperature obser-303 vations near 105 km also showed a pronounced  $(\pm 30 \text{ K})$  WN4 structure in the low-latitude 304 lower thermosphere associated with DE3 that was found to be highly correlated with the 305 ionospheric WN4 (r=0.71 for SABER/SORTIE and r=0.66 for SABER/ICON). 306

-10-

307	The latitude-height structure of this prominent DE3 was further investigated us-
308	ing SD/WACCM-X. DE3 amplitudes were found to be largest around the F-layer peak
309	near 200-350 km, with two well-defined enhancements near $15^{\circ}\mathrm{N}$ and $15^{\circ}\mathrm{S}$ MLAT as-
310	sociated with the EIA and maxima of $\sim 8 \times 10^4$ cm <sup>-3</sup> . The modeled DE3 temperature am-
311	plitudes were found to exhibit largest amplitudes around 110-120 km, in accord with pre-
312	vious modeling and observational results (e.g., Truskowski et al., 2014; Gasperini et al.,
313	2015, 2017a). Departure from a purely equatorially symmetric latitude structure and broad-
314	ening of latitude structures with height of the thermospheric DE3 was explained in terms
315	of the combined effect of mean winds and dissipation. The vertical wavelength of the mod-
316	eled temperature DE3 was found to be ${\sim}49.8~\mathrm{km}$ indicative of a predominant first sym-
317	metric Hough mode.

This study presented first results of the DE3 tide as observed by the NASA-funded SORTIE CubeSat. The results herein provide evidence for the extent to which globalscale tidal activity related to the tropical troposphere can influence the IT system. This study also demonstrates SORTIE and ICON to be excellent observational platforms for studying the influence of terrestrial weather on IT variability. Complementary and concurrent measurements from SORTIE's and ICON's near identical IVM instruments at different altitudes are shown to be particularly valuable.

### 325 Acknowledgments

The SORTIE mission is supported from NASA HQ by Grant 80NSSC18K0094 to Atmospheric & Space Technology Research Associates (ASTRA) LLC. ICON is supported by NASAs Explorers Program through contracts NNG12FA45C and NNG12FA42I. AS-TRA is grateful for support from U.C. Berkeley and the ICON mission via Subcontract

No: 00008210. ICON IVM-A data products (Level 2.7, Version 4) are publicly available

at ftp: //icon - science.ssl.berkeley.edu/pub. Post-processed SORTIE IVM Level 2

ion density data can be accessed on Zenodo at doi:10.5281/zenodo.4589362. TIMED/SABER

temperatures (version v2.0) can be freely accessed at http: //saber.gats - inc.com/.

WACCM-X history files can be accessed at the NCAR/CDG (https://doi.org/10.26024/5b58-

nc53) and are archived on the NCAR/HAO Campaign Space. Further information on

the ICON IVM data used in this study can be found at: ftp://icon-science.ssl.berkeley.edu/.

337

### 338 **References**

339	Appleton, E. V. (1946), Two anomalies in the ionosphere, Nature, 157, 691.
340	Balan, N., and G. J. Bailey (1995), Equatorial plasma fountain and its effects:
341	Possibility of an additional layer, J. Geophys. Res., 100, 21421-21432,
342	doi:10.1029/95JA01555.
343	Chang, L. C., Lin, C.H., Liu, J.Y., Balan, N., Yue, J., and Lin, J.T. (2013), Seasonal
344	and local time variation of ionospheric migrating tides in $2007-2011$ FOR-
345	MOSAT3/COSMIC and TIEGCM total electron content, J. Geophys. Res.
346	Space Physics, 118, 2545-2564, doi:10.1002/jgra.50268.
347	Chapman, S., and R. S. Lindzen (1970), Atmospheric Tides, Springer, New York.
348	Cho, YM., and G. Shepherd (2015), Resolving daily wave 4 nonmigrating tidal
349	winds at equatorial and midlatitudes with WINDII: DE3 and SE2, J. Geophys.
350	Res. Space Physics, 120, 10,05310,068, doi:10.1002/ 2015JA021903.
351	Crowley, G., C. Fish, M. Pilinski, E. Stromberg et al. (2016), Scintillation Observa-
352	tions and Response of The Ionosphere to Electrodynamics (SORTIE), Proceed-
353	ings of the 30th Annual AIAA/USU SmallSat Conference, paper: SSC16-VI-3.
354	Fang, TW., H. Kil, G. Millward, A. D. Richmond, JY. Liu, and SJ. Oh
355	(2009), Causal link of the wave-4 structures in plasma density and vertical
356	plasma drift in the low-latitude ionosphere, J. Geophys. Res., 114, A10315,
357	doi:10.1029/2009JA014460.
358	Forbes, J. M. (2000), Wave coupling between the lower and upper atmosphere: Case
359	study of an ultra-fast Kelvin wave, J. Atmos. Terr. Phys, 62, 1603-1621.
360	Garcia-Comas, M., et al. (2008), Errors in SABER kinetic temperature
361	caused by non-LTE model parameters, J. Geophys. Res., 113, D24, doi:
362	10.1029/2008JD010105.
363	Gasperini, F., J. M. Forbes, E. N. Doornbos, and S. L. Bruinsma (2015), Wave cou-
364	pling between the lower and middle thermosphere as viewed from TIMED and
365	GOCE, J. Geophys. Res., 120, 5788-5804, doi:10.1002/2015JA021300.
366	Gasperini, F., Forbes, J. M., and Hagan, M. E. (2017a), Wave coupling from the
367	lower to the middle thermosphere: Effects of mean winds and dissipation, J.
368	Geophys. Res., 122, 7781-7797, doi:10.1002/2017JA024317.
369	Gasperini, F, M. E. Hagan, and Y. Zhao (2017b), Evidence of tropo-
370	spheric 90-day oscillations in the thermosphere, Geophys. Res. Lett.,

371	doi:10.1002/2017GL075445.
372	Gasperini, F., Forbes, J. M., Doornbos, E. N., and Bruinsma, S. L. (2018),
373	Kelvin wave coupling from TIMED and GOCE: Inter/intra-annual vari-
374	ability and solar activity effects, J. Atmos. SolTerr. Phys., 171, 176-187,
375	doi:10.1016/j.jastp.2017.08.034.
376	Gasperini, F., Liu, H., and McInerney, J. (2020), Preliminary evidence of Madden-
377	Julian Oscillation effects on ultrafast tropical waves in the thermosphere, J.
378	Geophys. Res., 125, e2019JA027649, doi:10.1029/2019JA027649.
379	Gelaro, R., McCarty, W., Suarez, M. J., Todling, R., Molod, A., Takacs, L., et al.
380	(2017), The modern-era retrospective analysis for research and applications,
381	version 2 (MERRA-2), J. Climate, 30, 5419-5454, doi:10.1175/JCLI-D-16-
382	0758.1.
383	Hagan, M. E. and J. M. Forbes (2002), Migrating and nonmigrating diurnal tides in
384	the middle and upper atmosphere excited by tropospheric latent heat release,
385	J. Geophys. Res., 107(D24), 4754, doi: 10.1029/2001JD001236.
386	Hagan, M.E., A. Maute, R.G. Roble, A.D. Richmond, T.J. Immel, and S.L. England
387	(2007), Connections between deep tropical clouds and the Earths ionosphere,
388	Geophys. Res. Lett., 34, L20109, doi:10.1029/2007GL030142.
389	Hagan, M. E. (1996), Comparative effects of migrating solar sources on tidal signa-
390	tures in the middle and upper atmosphere, J. Geophys. Res., 101, 21213-21222.
391	He, M., L. Liu, W. Wan, and Y. Wei (2011), Strong evidence for couplings between
392	the ionospheric wave4 structure and atmospheric tides, Geophys. Res. Lett.,
393	38, L14101, doi:10.1029/2011GL047855.
394	Heelis, R. A., Lowell, J. K., and Spiro, R. W. (1982), A model of the high-latitude
395	ionospheric convection pattern, J. Res. Lett., 87, 63390-6345.
396	Heelis, R.A., Stoneback, R.A., Perdue, M.D., Depew, M.D., Morgan, W.A., Mankey,
397	M.W., Lippincott, C.R., Harmon, L.L., and Holt, B.J., (2017), Ion Velocity
398	Measurements for the Ionospheric Connections Explorer, Sp. Sci. Rev., $212(1-$
399	2), 615-629, doi:10.1007/s11214-017-0383-3.
400	Hurrell, J. W., Holland, M. M., Gent, P. R., Ghan, S., Kay, J. E., Kushner, P. J., et
401	al. (2013), The community earth system model: A framework for collaborative
402	research, Bull. of the Am. Met. Soc., $94(9)$ , 1339-1360, doi:10.1175/BAMS-D-
403	12-00121.1.

404	Immel, T. J., E. Sagawa, S. L. England, S. B. Henderson, M. E. Hagan, S. B.
405	Mende, H. U. Frey, C. M. Swenson, and L. J. Paxton (2006), The control of
406	equatorial ionospheric morphology by atmospheric tides, Geophys. Res. Lett.,
407	33 (15), doi: $10.1029/2006$ GL026161.
408	Immel, T.J., England, S.L., Mende, S.B. et al, (2018), The Ionospheric Connec-
409	tion Explorer Mission: Mission Goals and Design, Sp. Sci. Rev. 214, 13,
410	doi:10.1007/s11214-017-0449-2.
411	Jin, H., Y. Miyoshi, H. Fujiwara, and H. Shinagawa (2008), Electrodynamics of the
412	formation of ionospheric wave number 4 longitudinal structure, J. Geophys.
413	Res., 113, A09307, doi:10.1029/2008JA013301.
414	Lei, J., J.P. Thayer, W. Wang, J. Yue, and X. Dou (2014), Nonmigrating tidal mod-
415	ulation of the equatorial thermosphere and ionosphere anomaly, J. Geophys.
416	Res., 119, 30363043, doi:10.1002/2013JA019749.
417	Lieberman, R. S., Riggin, D. M., Ortland, D. A., Nesbitt, S. W., and Vin-
418	cent, R. A. (2007), Variability of mesospheric diurnal tides and tropo-
419	spheric diurnal heating during 1997-1998, J. Geophys. Res., 112, D20110,
420	doi:10.1029/2007JD008578.
421	Lieberman, R. S. (1991), Nonmigrating diurnal tides in the equatorial middle atmo-
422	sphere, J. Atmos. Sci., 48, 1112 - 1123.
423	Lin, C.H., C.C. Hsiao, J.Y. Liu, and C.H. Liu (2007), Longitudinal structure of the
424	equatorial ionosphere: Time evolution of the four-peaked EIA structure, J.
425	Geophys. Res., 112, A12305, doi:10.1029/2007JA012455.
426	Liu, H., and S. Watanabe (2008), Seasonal variation of the longitudinal structure
427	of the equatorial ionosphere: Does it reflect tidal influences from below?, J.
428	Geophys. Res., 113, A08315, doi:10.1029/2008JA013027.
429	Liu, H., M. Yamamoto, and H. Luhr (2009), Wave-4 pattern of the equatorial mass
430	density anomaly: A thermosphere signature of tropical deep convection, Geo-
431	phys. Res. Lett., 36, L18104, doi:10.1029/2009GL039865.
432	Liu, HL. (2016), Variability and predictability of the space environment
433	as related to lower atmosphere forcing, Space Weather, 14, 634-658,
434	doi: $10.1002/2016$ SW001450.
435	Liu, J., Liu, H., Wang, W., Burns, A. G., Wu, Q., Gan, Q., et al. (2018),

436 First results from the ionospheric extension of WACCM-X during the

437	deep solar minimum year of 2008, J. Geophys. Res., 123, 1534-1553,
438	doi:10.1002/2017JA025010.
439	Maute, A., A. D. Richmond, and R. G. Roble, Sources of low-latitude ionospheric
440	E x B drifts and their variability (2012), J. Geophys. Res., 117, A06312,
441	doi:10.1029/2011JA017502.
442	Mertens Christopher J., M Martin G. Mlynczak, Manuel Lpez-Puertas, Peter P.
443	Wintersteiner, R. H. Picard, Jeremy R.Winick, Larry L.Gordley, and James
444	M. Russell III (2001), Retrieval of mesospheric and lower thermospheric kinetic
445	temperature from measurements of CO <sub>2</sub> 15 $\mu m$ Earth limb emission under
446	non-LTE conditions, Geophys. Res. Lett., 28, 1391-1394.
447	Mukhtarov, P., and D. Pancheva (2011), Global ionospheric response to nonmigrat-
448	ing DE3 and DE2 tides forced from below, J. Geophys. Res., 116, A05323,
449	doi:10.1029/2010JA016099.
450	Oberheide, J., J. M. Forbes, X. Zhang, and S. L. Bruinsma (2011), Wavedriven
451	variability in the ionosphere thermosphere mesosphere system from TIMED
452	observations: What contributes to the wave 4?, J. Geophys. Res., 116, A01306,
453	doi:10.1029/2010JA015911.
454	Oberheide, J., J. Forbes, K. Hausler, Q. Wu, and S. L. Bruinsma (2009), Tropo-
455	spheric tides from 80400 km: Propagation, inter-annual variabil- ity and solar
456	cycle effects, J. Geophys. Res., doi:10.1029/2009JD012388.
457	Oberheide, J., Hagan, M. E., Ward, W. E., Riese, M., and Offermann, D. (2000),
458	Modeling the diurnal tide for the Cryogenic Infrared Spectrometers and Tele-
459	scopes for the Atmosphere (CRISTA) 1 time period, J. Geophys. Res., $105($
460	A11), 24917-24929, doi:10.1029/2000JA000047.
461	Onohara, A. N., Batista, I. S., and Batista, P. P. (2018), Wavenumber4 structures
462	observed in the low latitude ionosphere during low and high solar activity peri-
463	ods using FORMOSAT/COSMIC observations, Ann. Geophy., $36(2)$ , $459-471$ ,
464	doi:10.5194/angeo-36-459-2018.
465	Ren, Z., W. Wan, L. Liu, and J. Xiong (2009), Intra-annual variation of wave num-
466	ber 4 structure of vertical E B drifts in the equatorial ionosphere seen from
467	ROCSAT-1, J. Geophys. Res., 114, A05308, doi:10.1029/2009JA014060.
468	Sagawa, E., T. J. Immel, H. U. Frey, and S. B. Mende (2005), Longitudinal struc-
469	ture of the equatorial anomaly in the nighttime ionosphere observed by IM-

470	AGE/FUV, J. Geophys. Res., 110, A11302, doi:10.1029/ 2004JA010848.
471	Truskowski, A.O., Forbes, J.M., Zhang, X., and S.E. Palo (2014), New perspec-
472	tives on thermosphere tides - 1. Lower thermosphere spectra and seasonal-
473	latitudinal structures, Earth, Planets and Space, $66-136$ , doi: $10.1186/s40623$ -
474	014-0136-4.
475	Wan, W., L. Liu, X. Pi, ML. Zhang, B. Ning, J. Xiong, and F. Ding (2008),
476	Wavenumber-4 patterns of the total electron content over the low latitude
477	ionosphere, Geophys. Res. Lett., 35, L12104, doi:10.1029/2008GL033755.
478	Wan, W., J. Xiong, Z. Ren, L. Liu, M.L. Zhang, F. Ding, B. Ning, B. Zhao,
479	and X. Yue (2010), Correlation between the ionospheric WN4 signature
480	and the upper atmospheric DE3 tide, J. Geophys. Res., 115, A11303,
481	doi:10.1029/2010JA015527.
482	Wan, W., Z. Ren, F. Ding, J. Xiong, L. Liu, B. Ning, B. Zhao, G. Li, ML. Zhang
483	(2012), A simulation study for the couplings between DE3 tide and longi-
484	tudinal WN4 structure in the thermosphere and ionosphere, J. of Atm. and
485	SolTerr. Phy.,, 90-91, 52-60, doi:10.1016/j.jastp.2012.04.011.

Figure 1. (a) Time series of the daily F10.7 solar flux (red curve) and 3-hourly ap (blue curve) during May 27 - June 5, 2020. (b)-(d) Time series of ICON (blue plus signs) and SORTIE (red squares) local solar time at  $25^{\circ}$ N (panel b), equator (panel c), and  $25^{\circ}$ S (panel d).

Figure 2. Global latitude (60°S-60°N) versus longitude (180°W-180°E) maps of SORTIE (upper panels) and ICON (lower panels) Ion Velocity Meter (IVM) ion density measured during May 27 - June 5, 2020. SORTIE (ICON) daytime, i.e., ~10-14 LT (~14-18 LT), averages are shown in panel a (panel b); while SORTIE (ICON) nighttime, i.e., ~22-2 LT (~2-6 LT) averages are shown in panel a' (b'). Panels a'' and b'' show maps of half ascending and descending node differences (i.e., daytime-nighttime differences). The black dots in panels a-a' and b-b' show the measurement locations. ICON ion density is scaled (multiplied) by a factor of 2 given ICON's higher mean altitude (~575-610 km) compared to SORTIE (~418-427 km). This scaling is performed to use the same color bar for both ICON and SORTIE.

Figure 3. Period versus zonal wavenumber amplitude spectra of SORTIE IVM ion density at MLAT 40°N (panel a), 10°S-10°N (panel a'), 40°S (panel a'') during May 27 - June 25, 2020, and of SD/WACCM-X ion density at MLAT 40°N (panel b), 10°S-10°N (panel b'), 40°S (panel b'') during May 27 - June 5, 2020. Note that this 10-day window for WACCM-X is adopted for consistency with the SORTIE and ICON results shown in Figure 2. Panels a' and b' show the same results as in panels a and b, but with the migrating tide set to zero. Eastward propagating waves are shown with negative values. The diurnal s=-3 DE3 is clearly seen in both the SORTIE and SD/WACCM-X low-latitude ionosphere spectra, along with a large DW1.

Figure 4. (a) MLAT versus zonal wavenumber diurnal amplitude spectra of SORTIE IVM ion density during May 27 - June 25, 2020. (b) Same as (a) but using SD/WACCM-X output during May 27 - June 5, 2020. The prominent ionospheric s=-3 DE3 exhibits amplitudes upward of  $3 \times 10^4$  cm<sup>-3</sup> in both observations (panel a') and the model (panel b').

Figure 5. Time series of SORTIE IVM ion density DE3 amplitudes (panel a) and phases (panel a') derived using least squares fitting within 30-day sliding windows during May 27 - September 15, 2020. The blue vertical boxes identify 1- $\sigma$  uncertainly estimates in the amplitudes and phases output by the fitting procedure. Panels b-b' show the SD/WACCM-X height (~100-450 km) versus MLAT (60°S-60°N) ion density DE3 amplitudes and phases, respectively. Panels c-c' show the same results as in panels b-b', but for neutral temperature DE3. The vertical wavelength of the modeled temperature DE3 inferred from its vertical phase progression (panel c') is ~49.8 km, consistent with a predominant first symmetric Hough mode.

Figure 6. Latitude ( $60^{\circ}S-60^{\circ}N$ ) versus longitude ( $180^{\circ}W-180^{\circ}E$ ) maps of TIMED/SABER kinetic temperatures near 105 km during May 27 - June 5, 2020 at the descending node (panel a) near 3-5 LT, ascending node (panel b) near 15-17 LT, and their half difference (panel c). A prominent longitudinal WN4 structure is evident in the differences (panel c) consistent with the concurrent ionospheric DE3 signature observed by SORTIE (ICON) IVM near 420 km (590 km). Figure 1.



Figure 2.

May 27 - June 5, 2020





Figure 3.



SORTIE IVM Ion Dens, May 27 - Jun 25, MLAT 10S-10N 4.5 a 2.0×10<sup>4</sup>



SORTIE IVM Ion Dens, May 27 - Jun 25, MLAT 40S

0

Zonal Wavenumber (s)

2

3

Eeriod (days) 0.5 (days) 0.5 2.5 0.5

1.5

1.0

0.5

-4

-3

-2





WACCM-X Ion Density, May 27 - Jun 5, MLAT 10S-10N



WACCM-X Ion Density, May 27 - Jun 5, MLAT 40S 2.0×10<sup>4</sup> 5.0 3 b'' 4.5 1.5×10⁴√ ⊑ 1.0×10⁴ ݠ 4.0 111 5.0×10<sup>3</sup> V 1.5 1.0 0.5 -3 -2 2 5 -4 3 -1 4 Zonal Wavenumber

Negative for eastward propagation

Figure 4.



Negative for eastward propagation

Figure 5.



Figure 6.

# TIMED/SABER Temperature V2.0, May 27 – June 5, 2020





# **Geophysical Research Letters**

### **RESEARCH LETTER**

10.1029/2021GL093277

### Key Points:

- First results from the Scintillation
   Observations and Response of The
   Ionosphere to Electrodynamics
   (SORTIE) CubeSat show a large low latitude ionospheric wavenumber-4
   (WN4) structure observed
   concurrently by Ionospheric
   Connection Explorer (ICON) and
   Sounding of the Atmosphere using
   Broad band Emission Radiometry
- Spectral analyses of SORTIE Ion Velocity Meter (IVM), ICON IVM, and model output demonstrate that the diurnal eastward-propagating tide with zonal wavenumber 3 tide is responsible for this strong ionosphere-thermosphere (IT) WN4 coupling
- SORTIE and ICON IVM provide insights into tropical tropospheric influences on the IT by measuring poorly sampled altitudes simultaneously

Correspondence to:

F. Gasperini, fgasperini@astraspace.net

### Citation:

Gasperini, F., Azeem, I., Crowley, G., Perdue, M., Depew, M., Immel, T., et al. (2021). Dynamical coupling between the low-latitude lower thermosphere and ionosphere via the nonmigrating diurnal tide as revealed by concurrent satellite observations and numerical modeling. *Geophysical Research Letters*, *48*, e2021GL093277. https://doi. org/10.1029/2021GL093277

Received 25 MAR 2021 Accepted 28 JUN 2021

© 2021. American Geophysical Union. All Rights Reserved.

## Dynamical Coupling Between the Low-Latitude Lower Thermosphere and Ionosphere via the Nonmigrating Diurnal Tide as Revealed by Concurrent Satellite Observations and Numerical Modeling

Federico Gasperini<sup>1</sup> <sup>(1)</sup>, Irfan Azeem<sup>1</sup> <sup>(1)</sup>, Geoff Crowley<sup>1</sup> <sup>(1)</sup>, Michael Perdue<sup>2</sup> <sup>(1)</sup>, Matthew Depew<sup>2</sup>, Thomas Immel<sup>3</sup>, Erik Stromberg<sup>1</sup>, Chad Fish<sup>1</sup>, Crystal Frazier<sup>1</sup>, Adam Reynolds<sup>1</sup> <sup>(1)</sup>, Anthony Swenson<sup>1</sup>, Ted Tash<sup>1</sup>, Russell Gleason<sup>1</sup>, Ryan Blay<sup>1</sup> <sup>(1)</sup>, Jordan Maxwell<sup>1</sup>, Keith Underwood<sup>1</sup>, Christian Frazier<sup>1</sup>, and Scott Jensen<sup>1</sup>

<sup>1</sup>Atmospheric and Space Technology Research Associates, Louisville, CO, USA, <sup>2</sup>William B. Hanson Center for Space Sciences, Physics Department, University of Texas at Dallas, Richardson, TX, USA, <sup>3</sup>Space Sciences Laboratory, University of California, Berkeley, CA, USA

**Abstract** The diurnal eastward-propagating tide with zonal wavenumber 3 (DE3) is an important tidal component due to its ability to effectively couple the ionosphere-thermosphere with the tropical troposphere. In this work, we present the first results of a prominent zonal wavenumber-4 (WN4) structure in the low-latitude ionosphere observed by the Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) CubeSat mission during May 27–June 5, 2020. Least squares analyses of concurrent in situ ion number density measurements from the SORTIE and the Ionospheric Connection Explorer satellites near 420 and 590 km show this pronounced WN4 to be driven by DE3. Thermosphere Ionosphere Mesosphere Energetics Dynamics Sounding of the Atmosphere using Broad band Emission Radiometry temperatures and Specified-Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension output demonstrate that the ionospheric WN4 structure is driven by DE3 propagating from the lower thermosphere.

**Plain Language Summary** The extent to which terrestrial weather (below ~30 km) can influence the ionosphere and thermosphere (IT) is a fascinating discovery of the last two decades or so. The IT is known to vary significantly from day to day, and this day-to-day weather is largely driven by processes originating in the lower atmosphere, especially during periods of quiet solar activity. Accurate forecasting of the IT variability thus depends on the ability to forecast the component that originates in the lower atmosphere. Ionospheric variability translates to uncertainty in navigation and communications systems, while thermospheric variability translates to uncertainty in orbital and reentry predictions. In this work, we present the first results from the Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) CubeSat showing a large amplitude structure with four longitudinal peaks in the ionosphere associated with the well-known diurnal eastward-propagating tide with zonal wavenumber 3 originating in the tropical troposphere. Our analyses presented in this paper use concurrent SORTIE, Ionospheric Connection Explorer (ICON), and Thermosphere Ionosphere Mesosphere Energetics Dynamics satellite observations during May 27-June 5, 2020 and a whole atmosphere model to interpret SORTIE measurements. Our results suggest SORTIE and ICON to be excellent and complementary observational platforms for studying the influence of terrestrial weather on IT variability.

### 1. Introduction

The lower atmosphere drives variability in the ionosphere-thermosphere (IT) system through the vertical propagation of waves, including tides, planetary waves, and Kelvin waves. These waves are periodic in time and longitude due to the rotation of the Earth, and interact with the lower IT region to modulate electric fields that map to higher altitudes and redistribute plasma in the 200–1,000 km region. Due to the geometry of magnetic field lines near the equator, much of this variability occurs at low latitudes and is driven by waves that are excited by deep convective processes in the tropical troposphere and that propagate upwards

into the IT system. Tropical troposphere variability is essentially mapped to the IT system through a variety of neutral-plasma coupling processes, and over a range of spatial and temporal scales. One important class of global-scale atmospheric waves characterized by periods that are harmonics of a solar day are thermal tides. These atmospheric tides can be generated in different altitudinal regions due to tropospheric latent heating, absorption of tropospheric infrared radiation by water vapor, absorption of solar ultraviolet radiation by stratospheric ozone, thermosphere molecular oxygen absorption of extreme ultraviolet radiation, and wave-wave interactions (e.g., Chapman & Lindzen, 1970; Hagan & Forbes, 2002; Hagan et al., 2007; H.-L. Liu, 2016). The main pathways responsible for the modulation of the ionosphere by tides are direct propagation of atmospheric tides into the ionosphere and thermosphere (e.g., Hagan et al., 2007; Oberheide et al., 2009) and indirect coupling via the ionosphere E-region dynamo (e.g., Jin et al., 2008; Ren et al., 2010; Wan et al., 2008, 2010, 2012).

Several studies investigated the importance of IT variability driven by lower-atmosphere wave sources. Initial work (e.g., Fang et al., 2009; Hagan et al., 2007; Jin et al., 2008; Wan et al., 2010) focused on verifying Sagawa et al.'s (2005) and Immel et al.'s (2006) suggestion that the wavenumber-4 (hereafter, WN4) structure seen in satellite-borne Sun-synchronous F-region ionospheric data was due to the modulation of dynamo electric fields by nonmigrating tides propagating from below, and in particular due to the diurnal eastward-propagating tide with zonal wavenumber 3 (s = -3, i.e., DE3). DE3 originates in the tropical troposphere by latent heat release in deep convective clouds (e.g., Hagan, 1996; Hagan & Forbes, 2002; Lieberman et al., 2007), and its first equatorially symmetric Hough mode is the largest component in the lower thermosphere (Oberheide et al., 2009; Truskowski et al., 2014), capable of propagating well into the middle thermosphere (Gasperini et al., 2015, 2018; Gasperini, Forbes, & Hagan, 2017; Oberheide et al., 2011) due to its longer vertical wavelength. When viewed at quasi-fixed local time (LT) from slowly precessing satellites, this feature manifests as a 4-peak longitude structure. Subsequent studies examined the LT and seasonal variations of related WN4 and WN3 structures (e.g., Lin et al., 2007; H. Liu & Watanabe, 2008; Ren et al., 2009), and investigated the underlying mechanisms in further depth (e.g., Chang et al., 2013; Cho & Shepherd, 2015; He et al., 2011; Lei et al., 2014; Maute et al., 2012; Mukhtarov & Pancheva, 2011; Oberheide et al., 2011; Onohara et al., 2018).

In this work, we present observational evidence of prominent WN4 coupling between the lower thermosphere near 105 km and the ionospheric F-region at heights near 420 and 590 km during May 27–June 5, 2020. Our results take advantage of ion number density (hereafter, ion density) measurements from the Ion Velocity Meter (IVM) instruments onboard the Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) and Ionospheric Connection Explorer (ICON) spacecraft and kinetic temperature observations from the Sounding of the Atmosphere using Broadband Emission Radiometry (SABER) instrument onboard the Thermosphere Ionosphere Mesosphere Energetics Dynamics (TIMED) satellite. We further connect this WN4 longitude variability to DE3 by spectrally analyzing ICON and SORTIE IVM data and investigate its latitude-height amplitude and phase structure using a Specified-Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension (SD/WACCM-X) simulation with realistic wave forcing imposed by nudging Modern-Era Retrospective analysis for Research and Applications version-2 (MERRA-2) reanalysis data in the troposphere and stratosphere. After a brief description of the data, models, and methods (Section 2), we present in detail the observational and modeling results regarding the WN4 variability and its connection to the DE3 tide (Section 3), and conclude with a brief summary (Section 4).

### 2. Data, Models, and Methods

### 2.1. SORTIE/IVM

SORTIE is a NASA Heliophysics System Observatory (HSO) 6U CubeSat mission to investigate the underlying causes behind the appearance of plasma structures in the F-region ionosphere, leading to equatorial plasma bubbles, and the evolution of these structures after their formation (Crowley et al., 2016). SORTIE was launched onboard Dragon CRS-19 to the International Space Station, from where it was deployed on February 19, 2020 in a nearly circular orbit near 420 km with  $\sim$ 51.6° inclination. It carries two science instruments, a miniature IVM, and a micro-Planar Langmuir Probe (Crowley et al., 2016). This study employs SORTIE's IVM Level 2 ion density data product with a 4-s temporal cadence.





**Figure 1.** (a) Time series of the daily F10.7 solar flux (red curve) and 3-hourly ap (blue curve) during May 27–June 5, 2020. (b–d) Time series of Ionospheric Connection Explorer (ICON) (blue plus signs) and Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) (red squares) local solar time at 25°N (panel b), equator (panel c), and 25°S (panel d).

### 2.2. ICON/IVM

ICON is a NASA HSO mission designed to study the fundamental connections between the dynamics of the neutral atmosphere at altitudes between 100 and 300 km and the charged particle motions at low and middle latitudes from a nearly circular orbit at an altitude near 590 km with  $\sim$ 27° inclination (Immel et al., 2018). The ICON payload includes an IVM instrument that provides in situ measurements of the ion drift motions, density, temperature and major ion composition (Heelis et al., 2017). The IVM comprises of two instruments, the Retarding Potential Analyzer (RPA) and the Drift Meter. This study employs ICON's IVM-A Data Product 2.7 v4 determined from RPA measurements, with a 1-s temporal cadence. Preliminary validation work by the ICON IVM team reports accuracy around ±10<sup>3</sup> cm<sup>-3</sup> for this data product (see the acknowledgments for further information).

### 2.3. TIMED/SABER

The SABER instrument was launched onboard the TIMED satellite (also a NASA HSO mission) on December 7, 2001. SABER provides measurements of kinetic temperature from ~20 km to ~120 km altitude (Mertens et al., 2001). SABER views the atmospheric limb from an orbit of ~625 km altitude and ~73° inclination, so that the latitude coverage on a given day extends from about 53° in one hemisphere to about 83° in the other. This viewing geometry alternates once every ~60 days. Errors in the retrieved temperatures in the 80–105 km region are estimated to be ±1.5–5 K (Garcia-Comas et al., 2008). This study takes advantage of the SABER Level 2a Data Product with a ~55-s temporal cadence.

### 2.4. WACCM-X

WACCM-X (J. Liu et al., 2018) is a configuration of the National Center for Atmospheric Research Community Earth System Model (Hurrell et al., 2013) that extends the atmospheric component into the thermosphere, with a model top boundary between 500 and 700 km. The standard spatial resolution of WACCM-X is 1.9° by 2.5° in latitude and longitude, respectively, and 1/4 scale height vertically above the upper stratosphere. To simulate the lower and middle atmosphere, WACCM-X can be run with SD that constrains the troposphere and stratosphere dynamics using MERRA-2 reanalysis data (Gelaro et al., 2017). This study employs output from an extended SD/WACCM-X version v2.1 simulation that covers the May 27–June 5, 2020 period. The empirical ion convection patterns are specified using the Heelis et al. (1982) empirical model. The MERRA-2 forcing provides a realistic representation of the wave forcing in the lower and middle atmosphere making this simulation an ideal physics-based framework for interpreting atmosphere-ionosphere coupling by tides.

### 2.5. Analysis of Satellite Data

When combined, the ascending and descending nodes of SORTIE and ICON cover 24-h of LT in ~30 days (i.e., each node samples ~4 h of LT in about 10 days). Figure 1 shows the LT coverage of ICON (blue plus signs) and SORTIE (red squares) IVM measurements at 25°N (panel b), at the equator (panel c), and at 25°S (panel d) during May 27–June 5, 2020. At the equator, SORTIE's (ICON's) orbit precesses from ~2 LT/~14 LT (~7 LT/~19 LT) on May 27 to ~22 LT/~10 LT (~2 LT/~14 LT) on June 5 (Figure 1c). As shown in Figure 1a, this 10-day period is characterized by low solar activity (average F10.7 ~70) with two minor





**Figure 2.** Latitude ( $60^{\circ}S-60^{\circ}N$ ) versus longitude ( $180^{\circ}W-180^{\circ}E$ ) maps of Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) total ion density (upper panels) and Ionospheric Connection Explorer (ICON) O<sup>+</sup> density (lower panels) Ion Velocity Meter (IVM) ion density measured during May 27–June 5, 2020. SORTIE (ICON) daytime, that is, ~10–14 local time (LT) (~14–18 LT), averages are shown in panel (a) (panel b); while SORTIE (ICON) nighttime, that is, ~22–2 LT (~2–6 LT) averages are shown in panel (a') (b'). Panels (a'' and b'') show maps of half ascending and descending node differences (i.e., daytime-nighttime differences). The black dots in panels (a–a') and (b–b') show the measurement locations. ICON O<sup>+</sup> ion density is scaled (multiplied) by a factor of 2 given ICON's higher mean altitude (~575–610 km) compared to SORTIE (~418–427 km). This scaling is performed to use the same color bar for both ICON and SORTIE. The correlation between the WN4 in SORTIE and ICON (panels a'' and b'') is *r* = 0.87 near the equator (10°S–10°N), as noted at the bottom of panel (a''). MIVM, miniature IVM.

geomagnetic disturbances (ap  $\sim$ 25 on May 30, 2020 and ap  $\sim$ 15 on June 1–2, 2020). For this study, ICON and SORTIE IVM ion density are analyzed by both adding and subtracting ascending and descending node measurements. Wave analysis is performed by applying least squares methods to SORTIE IVM and ICON IVM ion density, and to SD/WACCM-X ion density and neutral temperature.

### 3. Results

The global distribution of ion density observed by SORTIE (ICON) IVM near 420 km (590 km) altitude during May 27–June 5, 2020 is presented in Figure 2. SORTIE daytime (~10–14 LT) and nighttime (~22–2 LT) 10-day averages are shown in panels a and b, respectively. Figures 2a' and 2b' show the corresponding ICON daytime (~14-18 LT) and nighttime (~2-6 LT) 10-day averages, respectively. Excellent low-latitude coverage is achieved by both SORTIE (about  $\pm 51.6^{\circ}$  latitude) and ICON (about  $\pm 27.1^{\circ}$  latitude) during this 10day period, as evidenced by the distribution of the black dots in panels (a-a') and (b-b') showing the measurement locations. Ion density values up to  $\sim 4.4 \times 10^5$  cm<sup>-3</sup> ( $\sim 2.1 \times 10^5$  cm<sup>-3</sup>) are observed during daytime and up to  $\sim 1.7 \times 10^5$  cm<sup>-3</sup> ( $\sim 0.5 \times 10^5$  cm<sup>-3</sup>) during nighttime by SORTIE (ICON) IVM. Note that a scaling factor is applied multiplying the ICON ion density by a factor of 2 given ICON's higher mean altitude (i.e., ~590 km vs. ~420 km for SORTIE). Figures 2a''-2b'' display ascending and descending node differences for SORTIE and ICON IVM ion density, respectively, each scaled (multiplied) by 0.5. Both SORTIE and ICON show enhanced ion density at low latitudes around 30°S-30°N with a prominent WN4 structure and peaks near 180°E, 100°E, 45°E, and 100°W. This type of feature in the low-latitude ionosphere was reported in previous satellite-based measurements (e.g., Immel et al., 2006; Lin et al., 2007; H. Liu & Watanabe, 2008; H. Liu et al., 2009), as discussed in Section 1. This notable WN4 structure observed by SORTIE and ICON points to a possible modulation of the background ion density by the nonmigrating DE3 tide as observed from their slowly precessing orbits. Note that while SORTIE IVM observes an O<sup>+</sup>-dominated ionosphere, ICON IVM samples nonnegligible H<sup>+</sup> (especially during nighttime). Thus, to best interpret the concurrent





**Figure 3.** Period versus zonal wavenumber amplitude spectra of Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) Ion Velocity Meter (IVM) ion density at magnetic latitude (MLAT) 40°N (panel a), 10°S–10°N (panel a'), and 40°S (panel a'') during May 27–June 25, 2020. (b–b'') Same as (a–a''), but for Ionospheric Connection Explorer (ICON) IVM O<sup>+</sup> ion density at MLAT 25°N (panel b), 10°S–10°N (panel b'), and 25°S (panel a''). (c–c'') Same as (a–a''), but for Specified-Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension (SD/ WACCM-X) ion density during May 27–June 5, 2020 and sampled along SORTIE's orbit. This 10-day window for WACCM-X is adopted for consistency with the SORTIE and ICON results shown in Figure 2. Eastward-propagating waves are indicated with negative wavenumbers. A prominent diurnal *s* = –3 diurnal eastward-propagating tide with zonal wavenumber 3 signal is clearly seen in the SORTIE, ICON, and SD/WACCM-X low-latitude ionosphere spectra, along with large DW1.

ionospheric WN4 signatures observed by ICON and SORTIE, Figures 2b-2b'' show ICON's O<sup>+</sup> density. The Pearson correlation coefficient between the WN4 structure seen in the SORTIE total ion density and ICON O<sup>+</sup> density (see Figures 2a''-2b'') is calculated to be ~0.87. Also note that the equatorial ionization anomaly (EIA) is not evident in Figure 2 likely due SORTIE's and ICON's mean altitude (~420 km and ~590 km, respectively) being significantly higher than the F-layer peak height (~200–350 km), in agreement with Figure 2 of Mukhtarov and Pancheva (2011).

It is important to realize that while the variability due to the semidiurnal eastward-propagating tide with s = -2 (SE2) and stationary planetary wave 4 (SPW4) would largely be eliminated by the ascending/descending node differences shown in Figures 2a''-2b'' for latitudes less than about  $\pm 10^{\circ}$  ( $\pm 30^{\circ}$ ) for ICON IVM (SORTIE IVM), aliasing from the diurnal westward-propagating tide with s = 5 (DW5) and from the terdiurnal eastward-propagating tide with s = -1 (TE1) may also contribute to a WN4 longitude structure seen at a constant LT (see Gasperini et al., 2015, 2018, 2020; Gasperini, Hagan, & Zhao, 2017; Lieberman, 1991; Oberheide et al., 2000). Additionally, the LT difference between ICON's (SORTIE's) ascending and descending node IVM measurements is < 9 h for latitudes greater than around  $\pm 10^{\circ}$  ( $\pm 30^{\circ}$ ) (see Figure 1). Hence one can anticipate significant aliasing due to variability associated with SE2 and SPW4 at latitudes greater than around  $\pm 10^{\circ}$  for ICON and  $\pm 30^{\circ}$  for SORTIE due to this LT difference being <9 h. Thus, to avoid these aliasing issues in the analysis of the WN4 structure, we (a) combine 30 days of SORTIE and ICON data collected during May 27-June 25, 2020 to acquire full 24-h LT coverage, and (b) use output from a 3-hourly SD/ WACCM-X simulation. Figure 3 shows the period versus zonal wavenumber amplitude spectra of SORTIE IVM ion density and ICON IVM O<sup>+</sup> ion density obtained by combining ascending and descending node measurements during May 27-June 25, 2020, and of SD/WACCM-X ion density during May 27-June 5, 2020 derived using the full model output. Around 10°S-10°N magnetic latitude (MLAT) SORTIE (panel a'),





Negative for eastward propagation

**Figure 4.** (a) Magnetic latitude (MLAT) versus zonal wavenumber diurnal amplitude spectra of Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) Ion Velocity Meter (IVM) ion density during May 27–June 25, 2020. (b) Same as (a), but for Ionospheric Connection Explorer (ICON) IVM O<sup>+</sup> ion density. (c) Same as (a), but from Specified-Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension (SD/WACCM-X) output sampled along SORTIE's orbit during May 27–June 5, 2020. (a'-c') Same as (a-c), but with the migrating tide set to zero. The prominent ion density *s* = -3 diurnal eastward-propagating tide with zonal wavenumber 3 (DE3) variation exhibits amplitudes upward of  $3 \times 10^4$  cm<sup>-3</sup> in both the SORTIE observations (panel a') and the model (panel c'). ICON's DE3 ion density amplitude maxima are shown to be around  $2.1 \times 10^4$  cm<sup>-3</sup>, that is, 30% lower than those found in SORTIE. The monthly averaged SORTIE and ICON diurnal ion density spectra are in general agreement, with similar latitudinal structures observed and predominant DW1, DE3, D0, and DE2 tidal signals.

ICON (panel b'), and WACCM-X (panel c') ion density spectra reveal the existence of a pronounced DE3 component with amplitudes up to  $\sim 2 \times 10^4$  cm<sup>-3</sup>. DE3 is found to be nearly absent in both WACCM-X and the SORTIE (ICON) observations at 40°N MLAT (25°N MLAT) and 40°S MLAT (25°N MLAT), in agreement with the well-known Kelvin wave behavior of the first-symmetric equatorially trapped Hough mode of DE3 that efficiently propagates to the IT. The ion density DE3 observed by ICON is about 30% smaller than that observed by SORTIE. This amplitude difference between ~420 km and ~590 km is likely due to the combined influence of dissipation, zonal mean winds, wave-wave interactions, and inherent transience on the upward propagating thermospheric DE3. Also, present in the 10°S–10°N MLAT spectra is the migrating DW1 tide, as shown in previous ionospheric observations (e.g., Chang et al., 2013). SW2, DW5, and TE1 amplitudes are found to be negligible (TE1 results not shown here), providing confidence to our assertion that the WN4 structure observed by SORTIE and ICON is in fact due to the nonmigrating DE3 tide.

To examine the latitudinal structure associated with this prominent F-region ion density DE3 (and other diurnal tides), Figure 4 shows the MLAT-wavenumber diurnal amplitude spectra of SORTIE IVM (panels a and a'), ICON IVM O<sup>+</sup> (panels b and b') and SD/WACCM-X (panels c and c') ion densities with and without the migrating DW1 tide set to zero. Similar to the results presented in Figure 3, SORTIE, ICON, and WAC-CM-X show general agreement with large DW1 ( $\sim$ 7 × 10<sup>4</sup> cm<sup>-3</sup>) and DE3 ( $\sim$ 2.1–3 × 10<sup>4</sup> cm<sup>-3</sup>) amplitude variations primarily confined to low latitudes. Along with pronounced DW1 and DE3 signals, the SORTIE and ICON spectra also exhibit D0 and DW2 variations (and some DE2 for SORTIE), while the WACCM-X spectra show DE1 and DW2 variations. Note that a  $\sim 2 \times 10^4$  cm<sup>-3</sup> DE2 amplitude variation is found in the model around 5°N-20°N MLAT, but not in the SORTIE or ICON observations. Larger DE3 amplitudes (up to  $3.06 \times 10^4$  cm<sup>-3</sup>) are found in the SORTIE and ICON spectra around 25°N–10°S MLAT (panels 4a' and 4b', respectively) consistent with the results shown in Figures 3a' and 3b', while WACCM-X maxima are found near 10°N-20°S MLAT. Some differences between the observed and modeled diurnal spectra are not unexpected and are likely due to (a) the different averaging window used (30 days for SORTIE and ICON, and 10 day for WACCM-X; note that a 10-day window for WACCM-X is adopted for consistency with the SORTIE and ICON results shown in Figure 2), (b) differences in the longitude-LT sampling between model and observations, and (c) inherent tidal variability observed by SORTIE and ICON that is not reproduced by the model. Even with the amplitude suppression associated with the monthly averaging, panel 4a' shows SORTIE IVM ion density DE3 amplitudes exceeding  $\sim 3 \times 10^4$  cm<sup>-3</sup>, that is, over 15% of the observed zonal mean. As noted above in the context of Figure 3, the ICON observed ion density DE3 amplitudes are about

30% smaller than SORTIE. Although not addressed in this study, the noted amplitude difference may be ascribable to reduced thermospheric DE3 amplitudes near 590 km due to the combined influence of dissipation, zonal mean winds, and wave-wave interactions that may result in reduced in situ generated ion density DE3 at ICON heights. Follow on work will investigate these possible contributions to the observed amplitude differences (and day-to-day variability) using concurrent thermospheric observations by ICON. On the other hand, it is important to point out the remarkable degree of consistency between the SORTIE and ICON diurnal ion density spectra between ~420 km and ~590 km during May 27–June 25, 2020. The latitude structure of DW1, DE3, D0, and DE2 are similarly captured by both space-borne observational platforms. It will be interesting to investigate in detail the diurnal, semidiurnal, and terdiurnal tidal coupling between these two ionospheric regions and possible sources of latitude and temporal variability in the various tidal components. A few likely candidates responsible for the observed variability with height in the tidal amplitudes and latitude structures are noted above. A more systematic study that employs the entire ~5-month period of simultaneous SORTIE and ICON observations and leverages other ICON remotely sensed thermospheric observations is currently underway.

Additional analysis of the DE3 tide in the ionosphere is performed using a 30-day sliding window of low-MLAT (±20°) SORTIE and ICON IVM ion density during the 4-month interval from May 27, 2020 through September 15, 2020 to investigate its temporal variability. Figure 5 shows daily variations of the estimated DE3 amplitudes (panel a) and phases (panel a') derived using a least squares method and the combined ascending/descending node data (i.e., covering ~24-h LT). Both SORTIE and ICON DE3 amplitudes are found to be larger ( $\sim 1.9 \times 10^4$  cm<sup>-3</sup> for SORTIE and  $\sim 1 \times 10^4$  cm<sup>-3</sup> for ICON) around May 27–June 25, quickly decreasing to minima  $(10^3-10^4 \text{ cm}^{-3} \text{ for SORTIE} \text{ and } \sim 800-2 \times 10^3 \text{ cm}^{-3} \text{ for ICON})$  near June–July, with largest amplitudes ( $\sim 2.1 \times 10^4$  cm<sup>-3</sup> for SORTIE and  $\sim 1.3 \times 10^4$  cm<sup>-3</sup>) observed during mid-August through early September 2020. General agreement is found in the seasonal variation of the ion density DE3 amplitudes observed by SORTIE near 420 km and by ICON near 590 km, with some differences likely associated to additional complexities introduced by the influences of dissipation, zonal mean winds, and wavewave interactions. The latitude structure of DE3 presented in Figure 4 and its temporal variation shown in Figure 5a, with largest amplitudes occurring around August–September and a second period of enhanced activity around April-May, is consistent with previous modeling work and observations (e.g., Gasperini et al., 2015; Mukhtarov & Pancheva, 2011; Truskowski et al., 2014). To investigate the altitude-latitude structure of the DE3 tide in the IT region, Figure 5b shows the SD/WACCM-X height (~100-450 km)-MLAT (60°S-60°N) structure of ion density (panels b and b') and temperature (panels c and c') amplitudes and phases during May 27-June 5, 2020. Ionospheric DE3 amplitudes are found to be largest around the F-layer peak near 200–350 km, with a defined 2-peak structure and maxima of  $\sim 8 \times 10^4$  cm<sup>-3</sup> around 15°N and 15°S associated with the EIA (Appleton, 1946; Balan & Bailey, 1995). Previous studies (e.g., Immel et al., 2006; Sagawa et al., 2005; Wan et al., 2008) showed the EIA to exhibit a WN4 longitudinal variation and it is now well accepted that this structure in the ionospheric F-region can form as a result of the combined effect of the E-region dynamo modulation by the lower thermospheric DE3 and from the direct propagation of DE3 into the F-region. It is fairly well established that neutral density variations, changes in thermospheric atomic oxygen to nitrogen ratio, and meridional winds at F-region altitudes (e.g., England et al., 2010; H. Liu et al., 2009; Maute et al., 2012) can contribute to the coupling between the tides and the ionospheric plasma. Previous modeling and observational results (e.g., Oberheide et al., 2011) indicate small amplitudes (<1 m/s) meridional DE3 winds in the low latitude (±15-20°) middle and upper thermosphere, which suggests the thermospheric DE3 winds to be a minor contributor the observed ionospheric DE3 signature. Similarly, neutral density and thermospheric atomic oxygen to nitrogen ratio DE3 amplitudes are generally small at SORTIE and ICON altitudes. It would be beyond the scope of this study to investigate the relative contribution of these effects to the ionospheric DE3 observed by SORTIE and ICON and this effort is left for follow on work. The lower thermospheric temperature DE3 exhibits largest amplitudes of ~12 K around 110-120 km, in accord with previous modeling (e.g., Gasperini et al., 2015; Gasperini, Forbes, & Hagan, 2017) and observational (e.g., Truskowski et al., 2014) results. The vertical wavelength of the modeled temperature DE3 inferred from its vertical phase progression (panel c') is ~49.8 km, in line with a predominant first-symmetric Hough mode. Latitudinal asymmetries and broadening of the latitude structure with height in the thermospheric DE3 are likely due to the combined effect of mean winds and dissipation (e.g., Forbes, 2000; Gasperini et al., 2015; Gasperini, Forbes, & Hagan, 2017).





**Figure 5.** Time series of Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) (blue line) and Ionospheric Connection Explorer (ICON) (red line) Ion Velocity Meter (IVM) ion density diurnal eastward-propagating tide with zonal wavenumber 3 (DE3) amplitudes (panel a) and phases (panel a') derived using least squares fitting within 30-day sliding windows during May 27–September 15, 2020. The vertical boxes identify  $1-\sigma$  uncertainly estimates in the amplitudes and phases output by the fitting procedure. Panels (b–b') show the Specified-Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension (SD/WACCM-X) height (~100–450 km) versus magnetic latitude (MLAT) (60°S–60°N) ion density DE3 amplitudes and phases, respectively. Panels (c–c') show the same results as in panels (b–b'), but for neutral temperature DE3. The vertical wavelength of the modeled temperature DE3 inferred from its vertical phase progression (panel c') is ~49.8 km, consistent with a predominant first symmetric Hough mode.

Figure 6 displays maps of TIMED/SABER temperatures near 105 km observed concurrently (i.e., May 27– June 5, 2020) with the SORTIE and ICON IVM ion density maps presented in Figure 2. Panels a and b show results for the descending (~3–5 LT) and ascending (~15–17 LT) nodes, respectively; while panel c contains results for half ascending/descending node differences. A prominent WN4 structure is found to dominate the low-latitude (±30°) lower thermosphere with observed amplitude maxima of ~30 K peaking near 135°E, 75°E, 45°W, 135°W. Pearson correlation coefficients between the equatorial longitudinal WN4 structure in TIMED/SABER temperatures and the WN4 structure in SORTIE (ICON) IVM ion density are calculated to be r = 0.71 (r = 0.67). The correlation is computed on half ascending/descending node differences com-





### TIMED/SABER Temperature V2.0, May 27 – June 5, 2020

**Figure 6.** Latitude (60°S–60°N) versus longitude (180°W–180°E) maps of Thermosphere Ionosphere Mesosphere Energetics Dynamics Sounding of the Atmosphere using Broad band Emission Radiometry (TIMED/SABER) kinetic temperatures near 105 km during May 27–June 5, 2020 at the descending node (panel a) near 3–5 local time (LT), ascending node (panel b) near 15–17 LT, and their half difference (panel c). A prominent longitudinal wavenumber-4 structure is evident in the differences (panel c) consistent with the concurrent ionosphere to Electrodynamics (Ionospheric Connection Explorer) Ion Velocity Meter near 420 km (590 km).

bining data within 10°S–10°N MLAT. This level of correlation, along with the modeling work discussed in Figure 5, provides strong evidence that the ionospheric WN4 structure observed by SORTIE and ICON is in fact associated with the lower thermospheric DE3. This level of correlation also indicates that there is significant phase coherence between the lower atmospheric DE3 and the F-region ion density DE3 near 420 and 590 km. This result agrees with our hypothesis that the F-region ion density DE3 is primarily driven by the E-region dynamo. As previously discussed, modeling and observational results generally demonstrate small meridional wind, neutral density, and atomic oxygen to nitrogen ratio DE3 amplitudes in the low-latitude  $(\pm 15-20^\circ)$  middle and upper thermosphere. These considerations suggest that the in situ-driven component may be a minor component. An investigation on the relative contribution of the E-region dynamo modulation by the lower thermospheric DE3 versus the direct propagation of DE3 into the F-region is considered beyond the purview of the current investigation and will be the subject of a follow on work.

### 4. Summary and Conclusions

Results presented above provide a clear picture of a marked longitudinal WN4 variation observed during May 27–June 5, 2020 by SORTIE and ICON IVM in the low-latitude ionosphere near 420 and 590 km, respectively. Taking advantage of output from an SD/WACCM-X simulation nudged with MERRA-2 reanalysis data in the troposphere and stratosphere and monthly averaged SORTIE and ICON IVM data, this prominent WN4 structure in the F-region ion density is demonstrated to be due to the well-known DE3 tide. This nonmigrating tide has gained significant attention due to its ability to preferentially propagate to the IT from the tropical troposphere, thus effectively coupling these regions. SORTIE IVM (ICON IVM) observations and SD/WACCM-X output are shown to exhibit ion density DE3 amplitudes upward of  $3 \times 10^4$  cm<sup>-3</sup>( $2.1 \times 10^4$  cm<sup>-3</sup>), that is, over 15% of the zonal mean, around  $\pm 20^{\circ}$  MLAT, with general agreement between model and observations. The monthly averaged SORTIE and ICON diurnal ion density spectra are found to be in strong agreement, with similar latitudinal structures observed and predominant DW1, DE3, D0, and DE2 tidal signals. DE3 ion density amplitudes are found to be about 30% smaller in ICON observations near 590 km compared to SORTIE observations near 420 km, possibly due to the combined effect of dissipation, mean winds, and wave-wave interactions on the thermospheric DE3 propagating between 420 and 590 km (e.g., Gasperini, Forbes, & Hagan, 2017).

Least squares fitting of monthly averaged SORTIE and ICON IVM ion density during May 27-September 15, 2020 showed larger ( $\sim 1.9 \times 10^4$  cm<sup>-3</sup> for SORTIE and  $\sim 1 \times 10^4$  cm<sup>-3</sup> for ICON) DE3 amplitudes around May 27–June 25, rapidly decreasing to smaller values  $(10^3-10^4 \text{ cm}^{-3} \text{ for SORTIE and } \sim 800-2 \times 10^3 \text{ cm}^{-3} \text{ for})$ ICON) around June–July, and becoming largest ( $\sim 2.1 \times 10^4$  cm<sup>-3</sup> for SORTIE and  $\sim 1.3 \times 10^4$  cm<sup>-3</sup>) around mid-August through early September 2020, in agreement with the well-known seasonal variation of DE3 amplitudes highlighted in previous modeling and observational studies (e.g., Gasperini et al., 2015; Gasperini, Forbes, & Hagan, 2017; Mukhtarov & Pancheva, 2011; Truskowski et al., 2014). Concurrent TIMED/ SABER temperature observations near 105 km are shown to exhibit a pronounced WN4 structure ( $\pm$ 30 K) in the low-latitude lower thermosphere associated with DE3 that is found to be correlated with the ionospheric WN4 (r = 0.71 for SABER/SORTIE and r = 0.67 for SABER/ICON around 10°N-10°S MLAT). Strong correlation (r = 0.87) is found between the WN4 structure observed simultaneously by SORTIE and ICON IVM. This level of correlation and the similarities in the SORTIE and ICON ion density DE3 latitude structures demonstrate that this tidal component is effective at coupling these two different ionospheric regions near 420 and 590 km simultaneously. Note that this period is generally characterized by small thermospheric DE3 amplitudes (e.g., Forbes et al., 2014; Gasperini et al., 2015; Oberheide et al., 2011). The less than perfect agreement (i.e., r = 1.0) between the two heights is ascribable to additional complexities on the thermospheric DE3 introduced by wave dissipation, the presence of zonal mean winds, possible wave-wave interactions, and inherent transience.

The latitude-height structure of this prominent DE3 signal is further investigated using SD/WACCM-X. DE3 amplitudes were found to be largest around the F-layer peak near 200–350 km, with two well-defined enhancements near 15°N and 15°S MLAT associated with the EIA and maxima of  $\sim 8 \times 10^4$  cm<sup>-3</sup>. The modeled DE3 temperature amplitudes are found to exhibit largest amplitudes around 110–120 km, in accord with previous modeling and observational results (e.g., Gasperini et al., 2015; Gasperini, Forbes, & Hagan, 2017; Truskowski et al., 2014). Departure from a purely equatorially symmetric latitude structure and broadening of latitude structures with height of the thermospheric DE3 is explained in terms of the combined effect of mean winds and dissipation. The vertical wavelength of the modeled temperature DE3 is found to be  $\sim$ 49.8 km indicative of a predominant first symmetric Hough mode.

This study presented first, results of a prominent F-region ion density WN4 structure driven by the nonmigrating DE3 tide observed concurrently by the SORTIE CubeSat and ICON. The results herein contained provide evidence for the extent to which DE3 is capable of affecting simultaneously the global structure of the F-region ion density near 420 and 590 km. This study further demonstrates the degree to which global-scale tidal activity related to the tropical troposphere can influence the IT system. This study also demonstrates SORTIE and ICON to be excellent observational platforms for studying the influence of terrestrial weather on IT variability at low latitudes. Complementary and concurrent measurements from SORTIE's and ICON's near identical IVM instruments at different altitudes are shown to be particularly valuable.



Future work will take advantage of comprehensive sets of SORTIE and ICON observations to investigate in further detail the coupling of different IT regions by global-scale waves.

### Data Availability Statement

ICON IVM-A data products (Level 2.7, Version 4) are publicly available at https://spdf.gsfc.nasa.gov/pub/ data/icon/l2/. Post-processed SORTIE IVM Level 2 ion density data can be accessed on Zenodo at https:// doi.org/10.5281/zenodo.4589362. TIMED/SABER temperatures (version v2.0) can be freely accessed at http://saber.gats-inc.com/data.php/. WACCM-X history files can be accessed at the NCAR/CDG (https:// doi.org/10.26024/5b58-nc53) and are archived on the NCAR/HAO Campaign Space. For further information on the ICON IVM data used, see: https://spdf.gsfc.nasa.gov/pub/data/icon/documentation/. The post-processed data used for Figures 1–6 are available at https://doi.org/10.5281/zenodo.4746615.

### References

Appleton, E. V. (1946). Two anomalies in the ionosphere. Nature, 157, 691. https://doi.org/10.1038/157691a0

- Balan, N., & Bailey, G. J. (1995). Equatorial plasma fountain and its effects: Possibility of an additional layer. Journal of Geophysical Research, 100, 21421–21432. https://doi.org/10.1029/95JA01555
- Chang, L. C., Lin, C.-H., Liu, J.-Y., Balan, N., Yue, J., & Lin, J.-T. (2013). Seasonal and local time variation of ionospheric migrating tides in 2007-2011 FORMOSAT-3/COSMIC and TIE-GCM total electron content. *Journal of Geophysical Research: Space Physics*, 118, 2545– 2564. https://doi.org/10.1002/jgra.50268
- Chapman, S., & Lindzen, R. S. (1970). Atmospheric tides. Springer.
- Cho, Y.-M., & Shepherd, G. (2015). Resolving daily wave 4 nonmigrating tidal winds at equatorial and midlatitudes with WINDII: DE3 and SE2. Journal of Geophysical Research: Space Physics, 120(10), 10053–10068. https://doi.org/10.1002/2015JA021903

Crowley, G., Fish, C., Pilinski, M., Stromberg, E., Huang, C., Roddy, O., et al. (2016). Scintillation Observations and response of The Ionosphere to Electrodynamics (SORTIE). Proceedings of the 30th Annual AIAA/USU SmallSat Conference, Paper: SSC16-VI-3.

- England, S. L., Immel, T. J., Huba, J. D., Hagan, M. E., Maute, A., & DeMajistre, R. (2010). Modeling of multiple effects of atmospheric tides on the ionosphere: An examination of possible coupling mechanisms responsible for the longitudinal structure of the equatorial ionosphere. Journal of Geophysical Research, 115, A05308. https://doi.org/10.1029/2009JA014894
- Fang, T.-W., Kil, H., Millward, G., Richmond, A. D., Liu, J.-Y., & Oh, S.-J. (2009). Causal link of the wave-4 structures in plasma density and vertical plasma drift in the low-latitude ionosphere. *Journal of Geophysical Research*, 114, A10315. https://doi.org/10.1029/2009JA014460 Forbes, J. M. (2000). Wave coupling between the lower and upper atmosphere: Case study of an ultra-fast Kelvin wave. *Journal of Atmos*-
- orbes, J. M. (2000). wave coupling between the lower and upper atmosphere: Case study of an ultra-fast Kelvin wave. *Journal of Atmospheric and Terrestrial Physics*, 62, 1603–1621. https://doi.org/10.1016/s1364-6826(00)00115-2
- Forbes, J. M., Zhang, X., & Bruinsma, S. L. (2014). New perspectives on thermosphere tides: 2. Penetration to the upper thermosphere. *Earth, Planets and Space*, 66(1), 122. https://doi.org/10.1186/1880-5981-66-122
- Garcia-Comas, M., López-Puertas, M., Marshall, B. T., Wintersteiner, P. P., Funke, B., Bermejo-Pantaleón, D., et al. (2008). Errors in SABER kinetic temperature caused by non-LTE model parameters. *Journal of Geophysical Research*, *113*, D24. https://doi.org/10.1029/2008JD010105
- Gasperini, F., Forbes, J. M., Doornbos, E. N., & Bruinsma, S. L. (2015). Wave coupling between the lower and middle thermosphere as viewed from TIMED and GOCE. Journal of Geophysical Research: Space Physics, 120, 5788–5804. https://doi.org/10.1002/2015JA021300
- Gasperini, F., Forbes, J. M., Doornbos, E. N., & Bruinsma, S. L. (2018). Kelvin wave coupling from TIMED and GOCE: Inter/intra-annual variability and solar activity effects. *Journal of Atmospheric and Solar-Terrestrial Physics*, 171, 176–187. https://doi.org/10.1016/j. jastp.2017.08.034
- Gasperini, F., Forbes, J. M., & Hagan, M. E. (2017). Wave coupling from the lower to the middle thermosphere: Effects of mean winds and dissipation. Journal of Geophysical Research: Space Physics, 122, 7781–7797. https://doi.org/10.1002/2017JA024317
- Gasperini, F., Hagan, M. E., & Zhao, Y. (2017). Evidence of tropospheric 90-day oscillations in the thermosphere. *Geophysical Research Letters*, 44, 10125–10133. https://doi.org/10.1002/2017GL075445
- Gasperini, F., Liu, H., & McInerney, J. (2020). Preliminary evidence of Madden-Julian Oscillation effects on ultrafast tropical waves in the thermosphere. Journal of Geophysical Research: Space Physics, 125, e2019JA027649. https://doi.org/10.1029/2019JA027649
- Gelaro, R., McCarty, W., Suarez, M. J., Todling, R., Molod, A., Takacs, L., et al. (2017). The modern-era retrospective analysis for research and applications, version 2 (MERRA-2). *Journal of Climate*, 30, 5419–5454. https://doi.org/10.1175/JCLI-D-16-0758.1
- Hagan, M. E. (1996). Comparative effects of migrating solar sources on tidal signatures in the middle and upper atmosphere. Journal of Geophysical Research, 101, 21213–21222. https://doi.org/10.1029/96jd01374
- Hagan, M. E., & Forbes, J. M. (2002). Migrating and nonmigrating diurnal tides in the middle and upper atmosphere excited by tropospheric latent heat release. *Journal of Geophysical Research*, 107(D24), 4754. https://doi.org/10.1029/2001JD001236
- Hagan, M. E., Maute, A., Roble, R. G., Richmond, A. D., Immel, T. J., & England, S. L. (2007). Connections between deep tropical clouds and the Earth's ionosphere. *Geophysical Research Letters*, 34, L20109. https://doi.org/10.1029/2007GL030142
- He, M., Liu, L., Wan, W., & Wei, Y. (2011). Strong evidence for couplings between the ionospheric wave-4 structure and atmospheric tides. *Geophysical Research Letters*, 38, L14101. https://doi.org/10.1029/2011GL047855
- Heelis, R. A., Lowell, J. K., & Spiro, R. W. (1982). A model of the high-latitude ionospheric convection pattern. Journal of Geophysical Research, 87, 6339–6345. https://doi.org/10.1029/ja087ia08p06339
- Heelis, R. A., Stoneback, R. A., Perdue, M. D., Depew, M. D., Morgan, W. A., Mankey, M. W., et al. (2017). Ion velocity measurements for the Ionospheric Connections Explorer. Space Science Reviews, 212(1–2), 615–629. https://doi.org/10.1007/s11214-017-0383-3
- Hurrell, J. W., Holland, M. M., Gent, P. R., Ghan, S., Kay, J. E., Kushner, P. J., et al. (2013). The Community Earth System Model: A framework for collaborative research. *Bulletin of the American Meteorological Society*, 94(9), 1339–1360. https://doi.org/10.1175/BAMS-D-12-00121.1

#### Acknowledgments

The SORTIE mission is supported from NASA HQ by grant 80NSSC18K0094 to Atmospheric & Space Technology Research Associates (ASTRA) LLC. ICON is supported by NASA's Explorers Program through contracts NNG-12FA45C and NNG12FA42I. ASTRA is grateful for support from U. C. Berkeley and the ICON mission via Subcontract No: 00008210.

- Immel, T. J., England, S. L., Mende, S. B., Heelis, R. A., Englert, C. R., Edelstein, J., et al. (2018). The Ionospheric Connection Explorer mission: Mission goals and design. Space Science Reviews, 214, 13. https://doi.org/10.1007/s11214-017-0449-2
- Immel, T. J., Sagawa, E., England, S. L., Henderson, S. B., Hagan, M. E., Mende, S. B., et al. (2006). The control of equatorial ionospheric morphology by atmospheric tides. *Geophysical Research Letters*, 33(15). https://doi.org/10.1029/2006GL026161
- Jin, H., Miyoshi, Y., Fujiwara, H., & Shinagawa, H. (2008). Electrodynamics of the formation of ionospheric wave number 4 longitudinal structure. Journal of Geophysical Research, 113, A09307. https://doi.org/10.1029/2008JA013301
- Lei, J., Thayer, J. P., Wang, W., Yue, J., & Dou, X. (2014). Nonmigrating tidal modulation of the equatorial thermosphere and ionosphere anomaly. Journal of Geophysical Research: Space Physics, 119, 3036–3043. https://doi.org/10.1002/2013JA019749
- Lieberman, R. S. (1991). Nonmigrating diurnal tides in the equatorial middle atmosphere. Journal of the Atmospheric Sciences, 48, 1112–1123. https://doi.org/10.1175/1520-0469(1991)048<1112:ndtite>2.0.co;2
- Lieberman, R. S., Riggin, D. M., Ortland, D. A., Nesbitt, S. W., & Vincent, R. A. (2007). Variability of mesospheric diurnal tides and tropospheric diurnal heating during 1997-1998. Journal of Geophysical Research, 112, D20110. https://doi.org/10.1029/2007JD008578
- Lin, C. H., Hsiao, C. C., Liu, J. Y., & Liu, C. H. (2007). Longitudinal structure of the equatorial ionosphere: Time evolution of the fourpeaked EIA structure. *Journal of Geophysical Research*, 112, A12305. https://doi.org/10.1029/2007JA012455
- Liu, H., & Watanabe, S. (2008). Seasonal variation of the longitudinal structure of the equatorial ionosphere: Does it reflect tidal influences from below? Journal of Geophysical Research, 113, A08315. https://doi.org/10.1029/2008JA013027
- Liu, H., Yamamoto, M., & Luhr, H. (2009). Wave-4 pattern of the equatorial mass density anomaly: A thermosphere signature of tropical deep convection. *Geophysical Research Letters*, 36, L18104. https://doi.org/10.1029/2009GL039865
- Liu, H.-L. (2016). Variability and predictability of the space environment as related to lower atmosphere forcing. *Space Weather*, 14, 634–658. https://doi.org/10.1002/2016SW001450
- Liu, J., Liu, H., Wang, W., Burns, A. G., Wu, Q., Gan, Q., et al. (2018). First results from the ionospheric extension of WACCM-X during the deep solar minimum year of 2008. *Journal of Geophysical Research: Space Physics*, 123, 1534–1553. https://doi.org/10.1002/2017JA025010 Maute, A., Richmond, A. D., & Roble, R. G. (2012). Sources of low-latitude ionospheric E × B drifts and their variability. *Journal of Geo*theta and the state of the
- physical Research, 117, A06312. https://doi.org/10.1029/2011JA017502
  Mertens, C. J., Mlynczak, M. G., López-Puertas, M., Wintersteiner, P. P., Picard, R. H., Winick, J. R., et al. (2001). Retrieval of mesospheric and lower thermospheric kinetic temperature from measurements of CO<sub>2</sub> 15 µm Earth limb emission under non-LTE conditions. *Geo-physical Research Letters*, 28, 1391–1394. https://doi.org/10.1029/2000gl012189
- Mukhtarov, P., & Pancheva, D. (2011). Global ionospheric response to nonmigrating DE3 and DE2 tides forced from below. Journal of Geophysical Research, 116, A05323. https://doi.org/10.1029/2010JA016099
- Oberheide, J., Forbes, J., Hausler, K., Wu, Q., & Bruinsma, S. L. (2009). Tropospheric tides from 80–400 km: Propagation, inter-annual variability and solar cycle effects. *Journal of Geophysical Research*, 114. https://doi.org/10.1029/2009JD012388
- Oberheide, J., Forbes, J. M., Zhang, X., & Bruinsma, S. L. (2011). Wave-driven variability in the ionosphere- thermosphere-mesosphere system from TIMED observations: What contributes to the "wave 4"? Journal of Geophysical Research, 116, A01306. https://doi.org/10.1029/2010JA015911
- Oberheide, J., Hagan, M. E., Ward, W. E., Riese, M., & Offermann, D. (2000). Modeling the diurnal tide for the Cryogenic Infrared Spectrometers and Telescopes for the Atmosphere (CRISTA) 1 time period. *Journal of Geophysical Research*, 105(A11), 24917–24929. https://doi.org/10.1029/2000JA000047
- Onohara, A. N., Batista, I. S., & Batista, P. P. (2018). Wavenumber-4 structures observed in the low-latitude ionosphere during low and high solar activity periods using FORMOSAT/COSMIC observations. Annals of Geophysics, 36(2), 459–471. https://doi.org/10.5194/ angeo-36-459-2018
- Ren, Z., Wan, W., Liu, L., & Xiong, J. (2009). Intra-annual variation of wave number 4 structure of vertical E B drifts in the equatorial ionosphere seen from ROCSAT-1. Journal of Geophysical Research, 114, A05308. https://doi.org/10.1029/2009JA014060
- Ren, Z., Wan, W., Xiong, J., & Liu, L. (2010). Simulated wave number 4 structure in equatorial F-region vertical plasma drifts. *Journal of Geophysical Research*, 115, A05301. https://doi.org/10.1029/2009JA014746
- Sagawa, E., Immel, T. J., Frey, H. U., & Mende, S. B. (2005). Longitudinal structure of the equatorial anomaly in the nighttime ionosphere observed by IMAGE/FUV. Journal of Geophysical Research, 110, A11302. https://doi.org/10.1029/2004JA010848
- Truskowski, A. O., Forbes, J. M., Zhang, X., & Palo, S. E. (2014). New perspectives on thermosphere tides—1. Lower thermosphere spectra and seasonal-latitudinal structures. *Earth, Planets and Space*, 66, 66–136. https://doi.org/10.1186/s40623-014-0136-4
- Wan, W., Liu, L., Pi, X., Zhang, M.-L., Ning, B., Xiong, J., & Ding, F. (2008). Wavenumber-4 patterns of the total electron content over the low latitude ionosphere. *Geophysical Research Letters*, 35, L12104. https://doi.org/10.1029/2008GL033755
- Wan, W., Ren, Z., Ding, F., Xiong, J., Liu, L., Ning, B., et al. (2012). A simulation study for the couplings between DE3 tide and longitudinal WN4 structure in the thermosphere and ionosphere. *Journal of Atmospheric and Solar-Terrestrial Physics*, 90–91, 52–60. https://doi. org/10.1016/j.jastp.2012.04.011
- Wan, W., Xiong, J., Ren, Z., Liu, L., Zhang, M.-L., Ding, F., et al. (2010). Correlation between the ionospheric WN4 signature and the upper atmospheric DE3 tide. *Journal of Geophysical Research*, 115, A11303. https://doi.org/10.1029/2010JA015527