Bias Corrected Estimation of Paleointensity (BiCEP): An improved methodology for obtaining paleointensity estimates

Brendan J Cych^{1,1}, Matthias Morzfeld^{2,2}, and Lisa Tauxe^{3,3}

¹Scripps Institution of Oceanography ²UCSD ³University of California, San Diego

November 30, 2022

Abstract

The assumptions of paleointensity experiments are violated in many natural and archaeological materials, leading to Arai plots which do not appear linear and yield inaccurate paleointensity estimates, leading to bias in the result. Recently, paleomagnetists have adopted sets of "selection criteria" that exclude specimens with non linear Arai plots from the analysis, but there is little consensus in the paleomagnetic community on which set to use. In this paper, we present a statistical method we call Bias Corrected Estimation of Paleointensity (BiCEP), which assumes that the paleointensity recorded by each specimen is biased away from a true answer by an amount that is dependent a single metric of nonlinearity (the curvature parameter \sqrt{k}) on the Arai plot. We can use this empirical relationship to estimate the recorded paleointensity for a specimen where $\sqrt{k}=0$, i.e., a perfectly straight line. We apply the BiCEP method to a collection of 30 sites for which the true value of the original field is well constrained. Our method returns accurate estimates of paleointensity, with a higher level of accuracy and precision than the strict CCRIT selection criteria, and with higher accuracy and similar precision to the modified PICRIT03 criteria. The BiCEP method has a significant advantage over using these selection criteria because it achieves these accurate results without excluding large numbers of specimens from the analysis.

Bias Corrected Estimation of Paleointensity (BiCEP): An improved methodology for obtaining paleointensity estimates

Brendan Cych¹, Matthias Morzfeld¹, Lisa Tauxe¹

¹University of California, San Diego

Key Points:

4

5

6

7	•	Empirical evidence suggests that paleointensity estimates for non-ideal specimens
8		are biased.
9	•	BiCEP is a method for estimating paleointensity for ensembles of specimens, cor-
10		recting for bias
11	•	BiCEP produces accurate results when applied to data where the true field strength
12		is known.

Corresponding author: Brendan Cych, bcych@ucsd.edu

13 Abstract

The assumptions of paleointensity experiments are violated in many natural and archae-14 ological materials, leading to Arai plots which do not appear linear and yield inaccurate 15 paleointensity estimates, leading to bias in the result. Recently, paleomagnetists have 16 adopted sets of "selection criteria" that exclude specimens with non linear Arai plots from 17 the analysis, but there is little consensus in the paleomagnetic community on which set 18 to use. In this paper, we present a statistical method we call Bias Corrected Estimation 19 of Paleointensity (BiCEP), which assumes that the paleointensity recorded by each spec-20 imen is biased away from a true answer by an amount that is dependent a single met-21 ric of nonlinearity (the curvature parameter \vec{k}) on the Arai plot. We can use this em-22 pirical relationship to estimate the recorded paleointensity for a specimen where \vec{k} 23 0, i.e., a perfectly straight line. We apply the BiCEP method to a collection of 30 sites 24 for which the true value of the original field is well constrained. Our method returns ac-25 curate estimates of paleointensity, with similar levels of accuracy and precision to restric-26 tive sets of paleointensity criteria, but accepting as many sites as permissive criteria. The 27 28 BiCEP method has a significant advantage over using these selection criteria because it achieves these accurate results without excluding large numbers of specimens from the 29 analysis. It yields accurate, albeit imprecise estimates from sites whose specimens all fail 30 traditional criteria. BiCEP combines the accuracy of the strictest selection criteria with 31 the low failure rates of the less reliable 'loose' criteria. 32

³³ Plain Language Summary

Paleomagnetists perform experiments on rocks and pottery sherds (among other 34 things) to estimate the strength of the ancient Earth's magnetic field (the paleointen-35 sity) through time. These make assumptions that are frequently violated, leading to bias. 36 Quantitative metrics (selection criteria) attempt to screen out 'bad' data. If a partic-37 ular experiment fails the criteria, the results are ignored. However, there is a lack of agree-38 ment as to which set of criteria are the most important and what is considered a fail-39 ure. One of these criteria quantifies the deviation from the fundamental assumption of 40 linearity between the ancient and laboratory magnetizations. We present a new Bayesian 41 method called Bias Corrected Estimation of Paleointensity (BiCEP), in which we assume 42 that the estimated paleointensity depends on this deviation. We can then use this de-43 pendency to correct the paleointensity made on an ensemble of specimens with differ-44 ing deviations from ideal behavior. BiCEP allows us to calculate accurate estimates of 45 the ancient magnetic field, without ignoring results from non-ideal specimens. We test 46 BiCEP on paleomagnetic data for which the original field strength is well constrained. 47 BiCEP recovers the field strength with similar accuracy to stricter sets of criteria, but 48 gets results for a greater number of sites. 49

50 1 Introduction

Estimates of the strength of the ancient Earth's magnetic field are currently made 51 by performing experiments that compare the natural remanent magnetization (NRM) 52 acquired by a specimen while cooling in the Earth's field, to a remanence known as ther-53 mal remanent magnetization (TRM) acquired by the specimen while cooling in a known 54 laboratory field. Such experiments include the Königsberger-Thellier-Thellier (KTT) fam-55 ily of experiments (Königsberger, 1938; Thellier & Thellier, 1959), the Shaw family of 56 experiments (Shaw, 1974), and the multi-specimen family of experiments (Hoffman et 57 al., 1989), among others. All of these experimental families make assumptions about the 58 relationship between the magnetic field and the remanent magnetization which may or 59 may not be applicable (see the review by Tauxe & Yamazaki, 2015). In this paper, we 60 will focus on the KTT family of experiments. 61

KTT type experiments involve a double heating protocol in which a specimen is 62 heated two or more times to a series of temperatures up to the Curie Temperature. At 63 each temperature, the specimen is cooled in two different fields. This has the effect of 64 replacing the NRM with a TRM acquired in a known laboratory field. Data from KTT-65 type experiments are normally represented by the Arai diagram (Nagata et al., 1963), 66 which plots the NRM magnetization remaining at each temperature step against the mag-67 netization imparted in the laboratory (often referred to as partial TRM or pTRM). The 68 ratio of these two magnetizations, as represented by the slope of the best fitting line to 69 the Arai plot data, is generally taken to be the ratio of the two magnetizing fields (an-70 cient, B_{anc} and laboratory, B_{lab}). 71

KTT-type experiments rely on several assumptions which are frequently violated
in paleointensity experiments. These include thermochemical alteration of specimens which
may lead to the production of new magnetic minerals, and an assumption known as reciprocity, which requires that the blocking temperature (the temperature below which grains
retain their magnetization after an external field is removed) is the same as the unblocking temperature (the temperature above which grains equilibrate with the external field).

The reciprocity assumption of Thellier and Thellier (1959) is fundamental to Néel's 78 theory for uniaxial single domain grains (Néel, 1949). Néel theory assumes that the elec-79 tronic spins within magnetic grains are fully aligned, and that the alignment is in one 80 of two directions along an energetically favorable 'easy' axis. In zero field, there is no pref-81 erence for either direction, but in the presence of a field there is a slight preference for 82 the direction along the easy axis with the smallest angle to the applied field. If the reci-83 procity assumption is met, then the energy required for the magnetization to change di-84 rections along the easy axis is always the same regardless of whether the specimen is cooled 85 from higher temperature (blocking) or heated from room temperature (unblocking) and 86 the two temperatures are identical. 87

By assuming that electronic spins within magnetic grains are fully aligned, Néel 88 theory fails to take into account a term in the magnetic energy of grains which causes 89 deviations from full alignment, resulting in structures such as the vortex state of, e.g., 90 Williams and Dunlop (1989). Although this effect is present in nearly all magnetic grains, 91 it is insignificant over short length scales (10s of nm) and so uniaxial single domain the-92 ory may be a reasonable approximation for smaller, elongate grains. Specimens in pa-03 leointensity experiments contain mixtures of grains with different sizes and shapes and a specimen used for paleointensity is likely to include grains for which the applicability 95 of single domain theory does not hold. 96

Failure of reciprocity and other fundamental assumptions embedded in the KTT 97 family of experiments (laid out by e.g., Thellier & Thellier, 1959) provides a challenge for those analyzing paleointensity data. Paleomagnetists generally use a set of selection 99 criteria which reject an intensity result if the NRM and pTRM data behave in a way which 100 deviate from single domain theory (linear on the Arai plot, see Figure 1a) by more than 101 some arbitrarily chosen threshold value. This is because data that contain a large pro-102 portion of non single domain-like grains or which otherwise violate the assumptions of 103 the experiment are likely to give biased results (Tauxe et al., 2021). Selection criteria 104 generally operate in a binary way, with specimens either being 'accepted' or 'rejected' 105 from the estimation of the site mean, where 'site' is the collection of specimens assumed 106 to have cooled in identical external magnetic fields (say, a lava flow or ceramic fragment). 107

Figure 1 gives a demonstration of biased results in specimens from prepared magnetite powders of increasing grain size that were magnetized in a 60 μ T field (Krása et al., 2003). If all assumptions of Thellier and Thellier (1959) were obeyed, we would expect the best fitting lines to data on Arai plots to give a range of values distributed closely about a mean of 60 μ T. As the grain size of the powder increases, the Arai plot becomes more curved and the best fitting line to the Arai plot yields a progressively lower intensity estimate. As all the paleointensities estimated from the curved plots are below the
expected value, the estimate for the ensemble can be biased, with the high temperature
segment having an even lower mean value, and the low temperature segment having a
high mean value. The data of Tauxe et al. (2021) also demonstrate downward curved
Arai plots in natural samples are biased so this problem may effect many of the results
compiled in paleointensity databases like the MagIC database (Tauxe et al., 2016) or PINT
(Biggin, 2010).

The curvature of an Arai plot can be quantified using the curvature criterion (\vec{k}) of Paterson (2011) (see also Paterson et al., 2014). Curvature is calculated using the reciprocal of the radius of a circle fit to scaled Arai plot data (see Section 2.2.1). While there is no theoretical basis for a circular fit (as opposed to the linear fit, which is firmly rooted in Néel theory), it is a useful approximation that we will exploit in this paper.



Figure 1. Arai plots from prepared magnetite powders given a TRM in a 60 μ T field (Krása et al., 2003). The curvature criterion, \vec{k} (Paterson, 2011) and specimen level paleointensity estimate B_m estimated from fitting a line to the entire Arai plot are plotted on the figure as text. The grain size of the magnetite powders increases from left to right. The coarser grains have non ideal domain state, leading to curved Arai plots and estimates of paleointensity which are biased to lower values than the expected 60 μ T. a) Nominal grain size of 23 nm. b) Mean grain size of 70 nm. c) Mean grain size of 12.1 μ m.

The practice of using binary (pass/fail) selection criteria is problematic for many 126 reasons. Paleomagnetic specimens generally contain magnetic carriers which span a range 127 of grain sizes and may or may not conform to the assumptions of the method. In addi-128 tion, micromagnetic simulations (e.g., Williams & Dunlop, 1989; Nagy et al., 2017) demon-129 strate that the change in magnetic domain state with grain size is a continuum, and so 130 one individual grain's behaviour may be more or less ideal than any other's. With bi-131 nary pass/fail criteria, the distinction between 'good' and 'bad' data must be assessed 132 with an arbitrary threshold value, which does not reflect the range of behaviors within 133 both groups. Consequently there are a large number of selection criteria in common use 134 (over 40 in Paterson et al., 2014), most of which have some empirical rationale, but there 135 is little agreement on which set to use or their threshold values. 136

In this paper, we describe a new approach for paleointensity estimation that treats the quality of paleointensity data as a continuum as opposed to the binary 'in' or 'out' approach using selection criteria. We assume that paleointensities become more biased as specimens' magnetic behaviors become more non-ideal and their Arai plots become less linear. By allowing the data interpretation for specimens to be based on the shape of their Arai plots, we are able to obtain unbiased estimates of paleointensity without the need for many specimen level (binary) selection criteria. We call this method the 'Bias

Corrected Estimation of Paleointensity' or BiCEP. In the next section, we develop a Bayesian 144 approach to obtain accurate paleointensity estimates with realistic uncertainties, using 145 k as a metric of bias, and show how to combine data at the site level. In Section 3 we 146 compare results from the BiCEP method to those of more traditional selection criteria 147 based approaches. We discuss the results in Section 4 and summarize our conclusions 148 in Section 5. Accompanying this paper, we release a Graphical User Interface (GUI) which 149 can apply the BiCEP method to MagIC formatted data. Links and instructions on how 150 to access the code can be found in Appendix 6.3. 151

152 2 Methods

153

186

187

188

189

2.1 Accounting for bias in paleointensity experiments

Paleomagnetists determine the paleointensity for a site by performing a Thelliertype double heating experiment on multiple specimens from that site. According to the theory for single domain grains (assuming no alteration of the specimen during heating), the ratio of NRM lost to pTRM gained is the ratio of the ancient field to the laboratory field. If the specimen conforms to theory, the Arai plot data will fall along a line the slope of which is equal to the ratio of ancient to the laboratory field (see Figure 1a).

We expect that the field strength predicted by the slope of the line on the Arai plot 160 for each specimen (here called B_m) will be distributed about the true (expected) ancient 161 field (B_{exp}) at the site with a Gaussian distribution. However, rarely do a set of spec-162 imens from a site all produce linear Arai plots that are easily interpretable. For exam-163 ple, interpretation of data from specimens with magnetic grains exhibiting non single do-164 main magnetic domain states produce non-linear Arai plots which violate the assump-165 tions of the method (e.g., Dunlop & Özdemir, 2001). Fitting lines to the data on such 166 Arai plots often produces estimates of paleointensity which are biased (see Figure 1c, Krása 167 et al., 2003), which in turn would bias site level estimates. 168

Paleomagnetists generally approach non-ideal data by using certain quantitative 169 criteria chosen to eliminate results suffering from one or more pathologies (Paterson et 170 al., 2014). If a particular criterion calculated for a specimen fails to meet some thresh-171 old value, then the specimen is excluded from the analysis. In this paper, we present an 172 alternative approach in which we allow for specimens to behave in a non-ideal (non-linear) 173 fashion when considering how specimen intensity estimates are distributed about a site 174 mean and weight the contribution of individual specimen estimates according to linear-175 ity. Under such a scheme, we start by predicting a bias for each specimen, and the spec-176 imens with the smallest predicted bias most strongly determine the paleointensity at that 177 site. In this way, biased specimens do not strongly affect our site intensity estimate, as 178 they are down-weighted, yet provide useful constraints on the uncertainty. 179

To predict the amount of bias a specimen is likely to have, we require a proxy for bias in paleointensity experiments. For this we use the curvature criterion \vec{k} of Paterson (2011) (see Section 2.2.1). There are several reasons that make this criterion a useful proxy for bias in paleointensity experiments:

- Specimens that are highly linear have, by definition, low values for $|\vec{k}|$ and will generally give unbiased paleointensity estimates (e.g., Cromwell et al., 2015).
 - By contrast, specimens with higher $|\vec{k}|$ yield biased paleointensities, with the magnitude of the bias generally increasing with the magnitude of $|\vec{k}|$ (e.g., Tauxe et al., 2021).
 - $|\vec{k}|$ has an empirical correlation with magnetic grain size (Paterson, 2011).

	MSH FreshTRM	Paterson et al. (2010) Santos and Tauxe (2010)	Lithic Clasts remagnetized /svnthetic	46.2 -12 N/A N	2.2 1980 / A N / A	55.6 19 70 0 24	
Table 1. Table of sites	used for analysis in this st	udy, including original study locs	ations, latitude, longitude	and year	of magnet	ization (wh	ere applicable),
expected field at that loca	tion (B_{exp}) , number of sp	ecimens used for analysis at that	t site M . Lat.: site latitude	(°N). Lc	ng. site l	ongitude (°	E. N/A: Not Ap-
plicable (Synthetic). B_{exp}	is either a known laborat	ory field, from the International 6	Geomagnetic Reference Fie	eld (IGRF	, Thébau	lt et al., 20	15 or in two cases
(hw226,hw108) using the .	Arch3k.1 model of Korte	et al., 2009					

To predict bias, we can use a method by which we minimize the misfit to a model assuming that B_m is linearly related to \vec{k} for all specimens. In other words, we say that:

$$B_m = B_{exp} + c\vec{k}_m + \epsilon \tag{1}$$

where m is an index reflecting the specimen number, ϵ is an error term and B_{exp} is the 192 true value of B. Effectively, our model just becomes a linear fit between the specimen 193 estimate B_m and \vec{k} , the y-intercept of which is the true value of the field B_{exp} and c is 194 a slope constant. While there is no theoretical justification (yet) for why B_m would be 195 related to k_m , although it has been observed empirically (by Paterson, 2011 using the 196 data in Figure 1, and more recently by Tauxe et al., 2021), a linear model is the simplest 197 one to relate the two. We demonstrate in Section 3.3 that more complex models with 198 a quadratic and cubic fit relating B_m to \vec{k}_m perform worse than the linear model when 199 predicting the paleointensity for sites for which the paleointensity is well constrained (his-200 torical lava flows or laboratory remanences). 201

Arai plot curvature is not the sole cause of bias in paleointensity experiments. In 202 some cases, specimens with Arai plots which do not have high |k| but are still non lin-203 ear (e.g., 'zig-zagged' as in, e.g., Yu et al., 2004), may still cause bias in paleointensity 204 experiments. To counteract this, we use a Bayesian method of calculating \vec{k}_m and B_m 205 which provides an uncertainty for both of these parameters. The benefit of this approach 206 is that specimens whose Arai plots are not well fit by a line or an elliptical arc have less 207 influence on the linear fit. Therefore, the specimens with the lowest uncertainty in k are 208 generally the most linear, and will have the most influence on the linear fit. Yet, for each 209 specimen, there is a trade off between minimizing the circle fit at a specimen level and 210 the linear fit between B_m and \vec{k} for specimens from the same site, an issue we will deal 211 with in Section 2.2.3. 212

Figure 2 shows results from our method (detailed in Section 2.2) applied to sev-213 eral sites for which the true value of B_{anc} (here, B_{exp}) is either calculated from the In-214 ternational Geomagnetic Reference Field (IGRF, Thébault et al., 2015) or Arch3k.1 (Korte 215 et al., 2009) for historical flows, or known as the NRM is a laboratory TRM imparted 216 to the specimens. Following Equation 1, the uncertainty in the intercept value of these 217 linear fits gives us the uncertainty for our site value of B_{anc} . In this way, we can obtain 218 an unbiased estimate of B_{anc} without relying on arbitrary binary (accept/reject) crite-219 ria to exclude specimen results. 220

In the following, we detail how the specimen level circle fit \vec{k} and site level paleointensity for unknown values for B (here called B_{anc}) can be calculated. We then compare the efficacy of several different versions of our model to classical selection criteria. We do this using a data compilation from 30 sites updated from Paterson et al. (2014) and Tauxe et al. (2016) for which B_{exp} is well constrained (see Table 1 for details concerning the original publications of the data).

227

228

2.2 Statistical Methodology

2.2.1 Estimating curvature

Paterson (2011) proposed a least squares fit of circles in Arai plot data. The parameter \vec{k} of Paterson (2011) is defined as the reciprocal of the radius of a best-fitting circle through the data. It is positive if the circle center is to the upper right of the Arai plot data (concave up, Figure 3a) and negative if the circle center is below and to the left of the Arai plot data (concave down, Figure 3b).

Before fitting to the Arai plot data, Paterson (2011) scales the pTRMs by the maximum pTRM to ensure that the paleointensity data are independent of the laboratory field. For estimating \vec{k} , we also subtract the minimum remaining NRM (NRM_{min}) for specimens for which full demagnetization has not been completed and we subtract the



Figure 2. Example of results from the BiCEP method for several sites used as examples in this study. Lines (in blue) are fit to the values of B_m and \vec{k} for each specimen (blue dots, with uncertainties as black lines). The values of linear fits at $\vec{k} = 0$ (blue histograms) provide an unbiased estimate of the expected paleointensity value at the site from the known field (red lines). a,d) hw126. b,e) hw201. c,f) BBQ. See Table 1 for sampling and citation details and Section 3 for comparison with the expected field values, B_{exp} .

minimum pTRM (pTRM_{min}) for specimens for which the low temperature steps were excluded from the analysis (e.g., because of viscous remanent magnetization).

For the BiCEP method, we define two sets of data vectors x and y:

240

$$x_n = \frac{\text{pTRM}_n - \text{pTRM}_{min}}{\text{pTRM}_{max}}, \quad y_n = \frac{\text{NRM}_n - \text{NRM}_{min}}{\text{NRM}_0}, \quad (2)$$

where *n* is the index of the data point. Because scaling should be by the total (original) TRM (the NRM), we also exclude specimens whose NRM_{min} is more than 25% of the initial NRM. This is justified by the assumption that the experimenter did not carry out demagnetization to fully replace the NRM. Then, to fit a circle with center x_c, y_c and radius *R* to the data, we try to minimize the squared perpendicular distance d_n^2 (Figure 3a) of all the *n* data points to the circle edge:

$$\sum_{n=1}^{N} d_n^2 \quad \text{where} \quad d_n^2 = \left(\sqrt{(x_n - x_c)^2 + (y_n - y_c)^2} - R\right)^2. \tag{3}$$

In a total least squares fit, Equation 3 would be our objective function that we would minimize. To fit circles to the Arai plot using a Bayesian method, we use Bayes' formula (Equation 4). This formula allows us to assign a probability distribution to the values of different parameters (in this case, \vec{k}_m and B_m), rather than just finding the 'best' value of the parameters. In a Bayesian context, we can simply assume that the data have some Gaussian noise distribution with some unknown standard deviation σ and apply Bayes' formula (e.g., Gelman et al., 2004):

$$P(\text{Parameters}|\text{Data}) = \frac{P(\text{Data}|\text{Parameters})P(\text{Parameters})}{P(\text{Data})},$$
(4)

where the left hand side is the probability of the parameters given the data and the right hand side is the probability of the data given the parameters times the probability of the



Figure 3. Example circles with different values for parameters \vec{k} and D with the same ϕ , showing how these parameters define a circle. a) Positive \vec{k} . Red dots are example data, and the green star is the intersection of D, ϕ with the circle edge (see text for definitions). d is the distance of an individual data point from the best-fit curve (blue). b) Negative \vec{k} . Note that in this case, ϕ could take any value as the circle center is at the origin, making the definition of ϕ meaningless in this case. c) Example showing how two sets of the parameters \vec{k}, ϕ, D can describe the same circle.

parameters, normalized by the probability of the data. In our case, the parameters are x_c, y_c, R and σ and our data are x and y so we rewrite Equation 4 as:

$$P(x_{c}, y_{c}, R, \sigma | x, y) = \frac{P(x, y | x_{c}, y_{c}, R, \sigma) P(x_{c}, y_{c}, R, \sigma)}{P(x, y)}.$$
(5)

The term $P(x, y|x_c, y_c, R, \sigma)$ is known as the "likelihood" and is based on the prob-258 ability of generating the observed data from a given set of parameters using the assumed 259 Gaussian distribution. The term $P(x_c, y_c, R, \sigma)$ is known as the "prior" and is a prob-260 ability distribution for values of x_c, y_c, R and σ we consider to be reasonable before we 261 see any data. We consider the priors on these parameters to be independent of one an-262 other, so we could rewrite this as $P(x_c)P(y_c)P(R)P(\sigma)$. The term P(x,y) is known as 263 the "evidence", and is simply a normalizing constant that makes the "posterior" prob-264 ability distribution, $P(x_c, y_c, R, \sigma | x, y)$, integrate to 1. In our application, we can sim-265 plify the relationship by ignoring the normalization. Furthermore, we can say from the 266 definition of the Gaussian distribution that: 267

$$P(x, y|x_c, y_c, R, \sigma) = \left(\frac{1}{2\pi\sigma^2}\right)^{N/2} \exp\left(\sum_{n=1}^N -\frac{d_n^2}{\sigma^2}\right).$$
(6)

268 Now we have an expression for our posterior probability distribution:

$$P(x_c, y_c, R, \sigma | x, y) \propto \left(\frac{1}{2\pi\sigma^2}\right)^{N/2} \exp\left(\sum_{n=1}^N -\frac{d_n^2}{\sigma^2}\right) P(x_c, y_c, R) P(\sigma).$$
(7)

Because the actual noise distribution of the Arai plot data is quite complicated (Paterson et al., 2012), we do not know the value of σ , so we use the uninformative prior $P(\sigma) \propto \frac{1}{\sigma}$; in other words, the smaller σ , the more likely the result. We can then substitute this prior into Equation 7 and integrate out σ to obtain:

$$P(x_c, y_c, R|x, y) \propto \left(\sum_{n=1}^N d_n^2\right)^{-N/2} P(R, x_c, y_c) \tag{8}$$

where N is the total number of measurements considered.

The set of parameters x_c, y_c and R is not easy to solve for, because Equation 3 has 274 multiple local minima (see Chernov and Lesort (2005) for a more detailed discussion). 275 Consider the simple case of a specimen with a linear Arai plot; in even this simplest case, 276 there are four minima, as both R and x_c, y_c will be either positive or negative and very 277 large. To avoid this complexity, we can use instead a change of parameters similar to that 278 of Chernov and Lesort (2005) which Paterson (2011) used as a basis for the circle fit-279 ting protocol. Based on this, we define a set of three new parameters which avoid the 280 281 problem of multiple minima.

Firstly, we require a point on the Arai plot which can be related to a unimodal distribution. We know that linear data will plot along the edge of a circle (the tangent), so if we draw a line from the origin toward the center (x_c, y_c) (not shown), this will touch the edge of the circle at some distance D (green star in Figure 3a). The angle to the horizontal of this line we call ϕ and we can directly estimate the \vec{k} parameter of Paterson (2011) using Equations 9,10,11. We can then establish equations for transforming between these two sets of parameters (see Appendix 6.1 for a more detailed derivation):

$$x_c = \left(D + \frac{1}{\vec{k}}\right)\cos(\phi),\tag{9}$$

$$y_c = \left(D + \frac{1}{\vec{k}}\right)\sin(\phi),\tag{10}$$

289

$$R = \frac{1}{|\vec{k}|}.$$
(11)

²⁹¹ Despite this transformation, the circle fitting equation can still have multiple min-²⁹² ima, even with \vec{k}, D, ϕ as our parameters, as the line connecting the origin to the hor-²⁹³ izontal touches the circle edge in two locations (see Figure 3c). However, we can use prior ²⁹⁴ distributions to avoid this.

²⁹⁵ Chernov and Lesort (2005) define a function of the data d_{max} to define the region ²⁹⁶ of possible values for \vec{k} :

$$d_{max} = \max_{i,j} \sqrt{(x_i - x_j)^2 + (y_i - y_j)^2}$$
(12)

Additionally, we define distance from the origin to the centroid of the data, d_{cent} :

$$d_{cent} = \sqrt{\bar{x}^2 + \bar{y}^2} \tag{13}$$

Using this function, we can assume that $D < 2d_{cent}$ and $|\vec{k}| < N/d_{max}$ and can define priors for our parameters:

$$P(D) \sim \text{Uniform}(0, 2d_{cent}),$$
 (14)

$$P(\phi) \sim \text{Uniform}(0,\pi),$$
 (15)

301 and

300

$$P(\vec{k}) \sim \text{Uniform}(-N/d_{max}, N/d_{max}).$$
 (16)

Using these priors gives us a posterior with a single maximum in most cases, which makes the problem much easier to solve computationally.

We can now apply a Bayesian approach to estimate \vec{k} for all temperature steps for a given specimen *m*. It is frequently useful to choose a subset of the temperature steps



Figure 4. Examples of circle fits to Arai plots (left column) and approximate probability densities of \vec{k} (right column). Dashed lines in the left hand plots are the tangents to circles with the median values for ϕ and D. We use tangents to the circle to get an estimate for B_m as outlined in Section 2.2.2. Triangles in a), c), e) are repeated lower temperature steps (pTRM checks) that indicate alteration of magnetic minerals during the experiment when offset from the original measurements (red dots). a) Specimen hw126a1. A fit to a straight line yields a precise \vec{k} distribution with a maximum close to zero (b). c) Specimen hw126a7. A curved Arai plot with a high amount of scatter/zigzagging (left) results in a higher uncertainty in the value of \vec{k} (d). e) Specimen hw126a6. Arai plot for a specimen that underwent thermochemical alteration at high temperature. A circle fit to just the low temperature steps results in a high uncertainty in the value of \vec{k} (f). Note that we do not exclude any measurements due to thermal alteration in our results section, and that this is only done here for illustrative purposes.

(e.g., if there is evidence for multiple components of the NRM or heating related alteration, as detected by repeated lower temperature pTRM steps). When using a subset of steps, we scale by the maximum pTRM for all temperature steps and the NRM at room temperature; in this way we can predict the curvature for the part of the Arai plot that is missing. This means that interpretations based on a small fraction of the Arai plot will have large uncertainties in the value of \vec{k} . Therefore, our circle fit can prioritize interpretations using the largest fraction of the NRM.

Figures 4a,c,e show circle fits sampled from the posterior distributions for speci-313 mens from site hw126 (site level results shown in Figure 2a). The probability densities 314 of all the k values for each specimen are plotted in Figures 4b,d,f. The plot demonstrates 315 how a straight Arai plot (Figure 4a) produces a narrow posterior about k = 0 (Figure 4b), 316 while a curved one (Figure 4c) produces a posterior which does not contain k = 0 (Fig-317 ure 4d). In the example with failed pTRM checks at higher temperatures (offset trian-318 gles in Figure 4e), we exclude the data points represented by open circles and use a lin-319 ear segment with only a portion of the results, the posterior distribution of k has a larger 320 uncertainty on the value, translating to a larger uncertainty in the bias for that spec-321 imen. We do not advocate for any particular method of checking for alteration, and do 322 not exclude any measurement steps in our results section. However, our circle fitting rou-323 tine allows for measurement steps to be excluded and accounts for the increased uncer-324 tainty in doing so. 325

326

2.2.2 Obtaining a specimen level paleointensity estimate

Analogous to the case in which paleointensity estimates are made using the slope 327 of a fitted line to the Arai plot data, we can obtain a similar "slope" value for a circu-328 lar arc fit to the data. Consider the case in which the edge of the circle forms an exact 329 line (k=0, see Figure 4a). In this case, the slope of the line can be given by the tangent 330 to the circle at the point where it intersects a line drawn from the origin (0,0) to the cir-331 cle center (Figure 3a). In other words, the "slope" of the Arai plot can be estimated as 332 $\cot \phi$, which gives the tangent to the circle. We can then turn this into an intensity es-333 timate B_m using the formula: 334

$$B_m = \frac{B_{lab}\cot(\phi)}{\text{pTRM}_{max}},\tag{17}$$

where B_{lab} is the laboratory field used to impart a pTRM to the specimen. If a specimen is corrected for anisotropy, cooling rate, or non-linear acquisition of TRM, we apply this correction to Equation 17.

We now have a way of obtaining estimates for B_m and \vec{k}_m for each specimen. We 338 use the methodology laid out in Sections 2.2.1 and 2.2.2 to plot the median value of the 339 posterior for these parameters (with error bars) in Figure 5a, and examples of circle fits 340 in Figures 5c, e, g. For specimens with values of \vec{k} that are approximately 0 (Figure 5g). 341 the B_m values are quite accurate. There appears to be a bias for specimens with large 342 \vec{k} , with the amount of bias increasing as \vec{k} increases. In this example, large positive val-343 ues of \vec{k} lead to a large underestimates of B_m while negative values of \vec{k} lead to overes-344 timates of B_m (although small in this example). 345

346

2.2.3 Obtaining a site level paleointensity estimate

The main problem with the method presented thus far is that we still do not have 347 a way of obtaining an estimate for B_{anc} , the unknown value at the site level. However, 348 in Figure 5a there appears to be a dependence between k_m and B_m as suggested ear-349 lier, with most of the specimens showing a quasi-linear relationship (the only exception 350 being the point labeled e) whose Arai plot is shown in Figure 5e) and suggests there is 351 a great deal of uncertainty in the value of \vec{k} itself. Because of this, we can modify our 352 model slightly by imposing the extra restriction that B_m must be linearly dependent on 353 \vec{k}_m (with noise) using Equation 1 (substituting B_{anc} for the unknown value of B_{exp}). 354

Previous papers have assumed that B_{anc} for selected specimens follows a Gaussian distribution and we can also make this assumption here. In the following, we will show how this modification can shift results from specimens that are offset from the linear relationship toward the line (as in the point labeled 'f' in Figure 5b) and produce mod-

els (shown as blue lines) that estimate all of our B_m . We can then use the resulting mod-

 $_{360}$ els to estimate the probability distribution for B_m as:



Figure 5. Examples demonstrating how the predicted \vec{k} and B_m for each specimen are modified for a site by using a hierarchical model (Equation 14). The left column shows draws from the posterior for an "unpooled" model where we estimate B_m and \vec{k}_m independently. The right column shows draws from the posterior for the BiCEP method where we assume a linear relationship between B_m and \vec{k}_m . a) Red horizontal line is B_{exp} (hw126, see Table 1). 95% credible intervals for $\vec{k_m}$ and B_m are plotted using black error bars, with the medians as green points. b) Representative draws from the posterior distribution are plotted as blue lines assuming that the individual specimen values B_m follow the relationship stated in Equation 14. Note that the higher curvature specimens with large uncertainty in \vec{k} follow a linear trend away from B_{exp} . c),e),g): [Symbols same as in Figure 4.] Arai plots of particular specimens are shown with circle fits sampled from the posterior of the unpooled model shown in a) and plotted in green. In d), f), h), same specimens as in c), e), g) but using the posterior of the BiCEP model in b). Note that there is little change in the specimen in d) for which a close fit to the data is possible, but in f) and h) the curvature (and intensity) of the specimen are modified to fit the line better.

$$P(B_m|k_m, B_{anc}, \sigma_{site}, c) = \frac{1}{\sqrt{2\pi\sigma_{site}^2}} \exp\left(-\frac{(B_{anc} + c\vec{k}_m - B_m)^2}{2\sigma_{site}^2}\right).$$
 (18)

Now we can combine our expressions for B_m and \vec{k}_m (Equations 17, Sections 2.2.1 and 2.2.2) with the new constraint of a linear relationship between B_m and \vec{k}_m (Equation 18). This allows us to obtain an expression for the site level intensity estimate B_{anc} :

$$P(B_{anc}, \sigma_{site}, c, B_m, k_m, D_m | x_m, y_m) \propto$$

$$P(x_m, y_m | k_m, D_m, B_m) P(B_m | k_m, B_{anc}, \sigma_{site}, c) P(B_{anc}, \sigma_{site}, c) P(D_m, k_m). \tag{19}$$

Equation 19 may look complicated, but we defined each of the terms already. The ben-365 efit of this treatment is that we can obtain $P(x_m, y_m | k_m, D_m, B_m)$ from our circle fit-366 ting in Equation 8 (see also Appendix 6.1). We defined $P(B_m|k_m, B_{anc}, \sigma_{site}, c)$ in Equa-367 tion 18. The values of \vec{k} and B_m for each specimen are needed to fit both of these terms. 368 This means that specimens with large scatter in their Arai plots (those which have Arai 369 plots that are not fit well by a line or a circle) are more strongly affected by the site level 370 fit B_{anc} , and therefore by the specimens with more linear (or circular) Arai plots. Con-371 versely, those specimens with a small uncertainty in \vec{k} or B_m are tightly constrained by 372 the Arai plot fit and so have more control over the fit at the site level. 373

The other two terms on the right side of Equation 19 $(P(B_{anc}, \sigma_{site}, c)P(D_m, k_m))$, 374 are priors. $P(D_m, k_m)$ were defined in Equations 14 and 16 respectively. Now, we need 375 to define priors for $P(B_{anc}, \sigma_{site}, c)$. For this purpose, we use a poorly constrained prior 376 for the slope, c, where $P(c) \propto 1$. Although this is not a probability distribution, the 377 resulting posterior distribution for B_{anc} is always a real probability distribution if the 378 number of specimens is greater than one. We use a uniform prior between 0 and 250 μ T 379 for $P(B_{anc})$ as intensity values can never be negative and in databases such as the MagIC 380 database (Tauxe et al., 2016) or the PINT database of Biggin (2010) rarely (if ever) ex-381 ceed 250 μ T. For $P(\sigma_{site})$ we use a normal distribution with zero mean and standard 382 deviation of 5 μ T, truncated to always be positive. 383

Figure 5b shows our median estimates for B_m and \vec{k}_m after applying the linear re-384 striction. Here, there is a tradeoff between fitting the Arai plot data with the circle, and 385 fitting the linear trend at a site level. The effect of the linear fitting is apparent when 386 compared to estimating \vec{k}_m and B_m for each specimen in isolation, which is shown in 387 Figure 5a. With the linear restriction, the k and B_m of specimens are "pulled" closer 388 to a linear trend by modifying the Arai plot fits; specimens with more uncertain k_m are 389 more strongly affected (e.g., specimen labeled e) and f) in Figure 5a and b). The spec-390 imens with highly linear Arai plots (for which we have small uncertainty in \vec{k}_m), the cir-391 cle fits (see g and h) are mostly unchanged. Despite this modification of the circle fits 392 to the Arai plots by the linear model, the circle fits to those specimens do not look un-393 reasonable. 394

395 **2.3** Metrics of success

364

In order to 'ground-truth' the method, we rely on a compilation of paleointensity 396 data updated from that of Paterson et al. (2014) and Tauxe et al. (2016). This compi-397 lation has data from 30 sites for which B_{anc} is well constrained (hence we use B_{exp}), ei-398 ther through the IGRF, or because the specimens were given TRMs in a known lab field 399 before the Thellier experiment. One exception to this is for hw226 and hw108, lava flows 400 erupted in Hawaii in 1843 and 1859, prior to the range included in the IGRF. For these 401 sites, we used the Arch3k.1 model of Korte et al. (2009). A list of sites used here is given 402 in Table 1. Instead of choosing a range of temperatures for each site, we simply use ev-403 ery temperature on the Arai plot for all specimens. 404



Figure 6. Examples of accuracy and precision metrics used in this study with simulated Gaussian distributions of B_{anc} for illustration. a) An accurate and precise estimate, b) An accurate but imprecise estimate, c) An inaccurate and imprecise estimate. d) A slightly inaccurate and highly precise estimate. Accuracy check used for n_{acc} checks whether the black line intersects the expected value (B_{exp}) . f_{prob} is the area of the blue histogram that lies within the red shaded area. Δ_{median} is the length of the green line.

Because we have to estimate multiple parameters for each specimen, our method involves a high dimensional optimization problem. Therefore, we generate the estimates for B_{anc} for a given site using a Markov chain Monte Carlo (MCMC) method which approximates the posterior distribution by generating pseudosamples from it (see Appendix 6.2 for details). MCMC techniques are frequently used to solve high dimensional problems of this kind.

For each site, we quantify the effectiveness of the BiCEP method using several metrics, f_{prob} , Δ_{median} (see Figure 6 for graphical representation), \bar{f}_{prob} , and n_{acc} :

- 4131. f_{prob} : We report the median value of our posterior distribution and the 2.5th and
97.5th percentile of the Monte Carlo sample (95% credible interval) as error bars.41497.5th percentile of the Monte Carlo sample (95% credible interval) as error bars.415To quantify the effectiveness of our method, we look at the proportion of the pos-
terior distribution that lies within 3 μ T of the expected value of $B(B_{exp})$ and call417this proportion f_{prob} .4182. \bar{f}_{prob} : the mean value of f_{prob} over all sites included in the study. A value of 1 is
 - 2. \bar{f}_{prob} : the mean value of f_{prob} over all sites included in the study. A value of 1 is the best possible value and means all our results are accurate and precise to better than 3 μ T.

419 420

- 421 3. Δ_{median} : the difference (in μ T) between the median value of the MCMC sample 422 (see Section 6.2 for explanation) and B_{exp} . The median value of Δ_{median} is $\tilde{\Delta}_{median}$. 423 Values of $\tilde{\Delta}_{median}$ close to zero are best.
- 424 4. n_{acc} : the number of sites for which B_{exp} lies within our 95% credible interval. A 425 related parameter, f_{acc} is the fraction of results that are accurate (n_{acc}/n_{sites}) ,

where n_{sites} is the total number of sites analyzed. We expect this number to be 0.95 in ideal circumstances.

We use these metrics to compare the BiCEP results to those obtained by several differ-428 ent sets of selection criteria: CCRIT (Cromwell et al., 2015), Paterson's modified PICRIT03 429 (here called PICRITMOD) and SELCRIT Criteria (here called SELCRITMOD, Paterson 430 et al., 2014). For this exercise, we also calculated these two criteria with the addition 431 of the curvature criterion of $|\vec{k}| < 0.270$, which we refer to as PICRITMODk and SEL-432 CRITMODk. We apply these criteria using the standard deviation optimization method 433 in Thellier GUI. Most sets of commonly used selection criteria rely on an assumption of 434 a Gaussian probability distribution for the site level estimate B_{anc} , which allows us to 435 calculate these same metrics. 436

For our analyses of our success metrics, we exclude sites that contain fewer than 437 three specimens. For fair comparison, we do not exclude sites from our analyses with tra-438 ditional selection criteria which have high standard deviation, as we do not do this for 439 BiCEP. If a site fails to produce an estimate of B_{anc} for any reason (for example, selec-440 tion criteria passed less than two specimens), we assume the prior distribution of a uni-441 form distribution between 0 and 250 μT . This allows us to compare methods directly, 442 with a penalty applied for excluding sites. An excluded site will have $f_{prob}=0.012$, whereas 443 a site with a highly inaccurate result can have f_{prob} of 0, so exclusion is considered only 444 slightly better than an inaccurate result in this scheme. We discuss the results of this 445 comparison in Section 3.1. 446

447

426

427

2.4 Width of prior and order of fit

Here we consider several alternative contingent models in order to explore our choices for $P(\sigma_{site})$ and assumptions about the relationship of B_m and \vec{k} . In addition to using a standard deviation of 5 μ T for $P(\sigma_{site})$, we use standard deviations of 10 μ T and 20 μ T. The effect of this is hard to conceptualize, but wider priors will prioritize fitting circles to the individual specimens over fitting the linear relationship between B_m and \vec{k}_m at a site level. The practical effect of this is wider posteriors for sites where the number of specimens is small.

⁴⁵⁵ So far, we have assumed a priori that B_m is linearly dependent on \vec{k}_m . Because there is no theoretical reason why this should be the case, we test models for which the relationship between B_m and \vec{k}_m is described by a quadratic polynomial and a cubic polynomial. We would expect a higher order model to more closely fit the individual \vec{k}_m and B_m values, but with a loss of precision due to the more complicated model.

Results for our method, as well as for two sets of selection criteria, are given in Table 2. For each model, we calculate \bar{f}_{prob} , $\tilde{\Delta}_{median}$ and f_{acc} for comparison. In this table, our models are named for the value of the standard deviation of $P(\sigma_{site})$ as well as the order of the fit. Our preferred model is referred to as "Linear 5 μ T", and this is the model used in this paper where otherwise unspecified.

465

2.5 MCMC sampler diagnostics

466 MCMC samplers are only ever an approximation of the posterior distribution, and 467 the number of Monte Carlo samples needed to make an accurate approximation is not 468 the same for every site, or every run of the sampler. To determine whether we are ac-469 curately sampling the posterior distribution, we look at three diagnostics which are also 470 described in Appendix 6.2:

- 1. \dot{R} : (Gelman & Rubin, 1992) quantifies convergence between chains in the MCMC method. This parameter is required to be between 1.1 and 0.9 for the sampler to converge.
- 2. n_{eff} : the effective MCMC sample size. We are using 30,000 Monte Carlo samples and n_{eff} should be large (> 1000) to have a good representation of our parameters.
- 3. f_{div} : the proportion of divergent transitions f_{div} in the MCMC sample. This should ideally be zero, but it does not appear to cause large problems for the estimate of B_{anc} if it is non zero (see Section 6.2).

The diagnostics n_{eff} and R are produced for each of our parameters (each of our $B_{m}, \vec{k}_{m}, D_{m}$ and B_{anc}, σ_{site}). When reporting these values, we look at the worst value of \hat{R} (furthest from unity) and the value of n_{eff} for B_{anc} . If $\hat{R} > 1.1$, we replace the distribution on B_{anc} with a uniform distribution between 0 and 250 μ T (the prior). The results of the MCMC sampler are presented in Section 3.4.

485 **3 Results**

471

472

473

474

475 476

477

478

479

486

3.1 Comparison of BiCEP to Selection Criteria

In this section, we compare the BiCEP to several sets of selection criteria (see Section 2.3). The full set of results for all sites can be seen in Figure 7, and are summarized in Supplementary Data Set S1.

Model Name	\bar{f}_{prob}	$\tilde{\Delta}_{median} \ (\mu T)$	f_{acc}	Number of Sites
Linear, 5 μ T (BiCEP)	0.63	1.7	0.85	25
Linear, 10 μT	0.62	1.7	0.85	25
Linear, 20 μT	0.61	1.7	0.85	25
Quadratic, 5 μ T	0.56	1.7	0.81	25
Quadratic, 10 μT	0.55	1.6	0.85	25
Quadratic, 20 μT	0.55	1.8	0.85	25
Cubic, 5 μ T	0.45	2.6	0.85	22
Cubic, 10 μ T	0.45	2.3	0.85	24
Cubic 20 μT	0.44	2.5	0.85	24
CCRIT	0.47	1.9	0.88	22
PICRIT (Modified)	0.56	2.2	0.77	23
PICRIT (Modified with \vec{k})	0.61	1.9	0.69	21
SELCRIT (Modified)	0.53	2.9	0.58	25
SELCRIT (Modified with \vec{k})	0.58	2.3	0.58	23

Table 2. Results comparing the models used in this study to results using CCRIT (Cromwell et al., 2015) as well as PICRITMOD and SELCRITMOD (Paterson et al., 2014), both with and without the \vec{k} criterion. See details in text and Figure 6 for explanations of the different parameters presented here. Results are sorted by the number of specimens in the site used to make the estimate using our method.

Figure 7 shows the 95% credible intervals for each method, normalized by the expected value at the site. The median values of our results are generally similar to those found by our selection criteria. BiCEP yields the largest number of accurate and precise results, with CCRIT being generally less precise and slightly less accurate. PICRIT-MOD and SELCRITMOD are generally less accurate and precise than BiCEP, however





introducing the curvature criterion for PICRITMODk and SELCRITMODk improve the
 accuracy and precision significantly. Both PICRITMODk and SELCRITMODk boast
 highly precise estimates for passing sites, with similar levels of accuracy to BiCEP. How ever, this improved accuracy and precision is achieved by excluding more sites, which
 penalizes these methods using our success metrics.

Sites in Figure 7 are sorted by the number of specimens used by BiCEP for the anal-500 ysis. Unique to our method, sites with low numbers of specimens (M) have wide cred-501 ible intervals and sites with high M have narrow credible intervals, so the estimate of 502 B_{anc} becomes more precise as more specimens are measured. This is because calculat-503 ing the credible interval for a B_{anc} is more similar to calculating the standard error of 504 the mean than the site level standard deviation, which is done for our traditional selec-505 tion criteria. The increasing precision on B_{anc} leads to some sites with high M having 506 estimates of B_{anc} which are seemingly too precise. These estimates are still generally only 507 a few μT away from the expected value, however, and we discuss potential reasons for 508 this in Section 4.4. 509

Our results in Table 2 indicate that BiCEP is the method that yields the largest 510 number of accurate and precise results, having a higher \bar{f}_{prob} and lower Δ_{median} than 511 all of our sets of selection criteria. For selection criteria which include a curvature cri-512 terion, much of this improvement comes from BiCEP's inclusion of accurate results for 513 two sites, remag-rs78 and Synthetic60. If we look exclusively at the sites which passed 514 each criterion, PICRITMODk and SELCRITMODk achieve higher levels of precision for 515 those sites (higher f_{prob} than BiCEP if only passing sites considered), with PICRITMODk 516 achieving similar levels of accuracy to BiCEP (similar $\overline{\Delta}_{median}$ for passing sites). This 517 higher level of precision is likely an outcome of using the standard deviation optimiza-518 tion procedure, and is probably not reflective of the true uncertainty judging by the low 519 f_{acc} for both PICRITMODk and SELCRITMODk. CCRIT still achieves lower f_{prob} and 520 higher Δ_{median} than BiCEP even if only passing sites are considered, indicating a slightly 521 lower accuracy and precision overall. Our two selection criteria which do not include a 522 curvature criterion (PICRITMOD and SELCRITMOD) have a larger number of pass-523 ing sites, including remag-rs78 and Synthetic60, but still have reduced \bar{f}_{prob} and Δ_{median} . 524 Ultimately it seems that BiCEP offers the best of both worlds, passing at least as many 525 sites as the more permissive criteria, and achieving higher accuracy and more realistic 526 precision than the more restrictive criteria. 527

3.2 Width of the prior

528

To investigate the role of the prior distribution $(P(\sigma_{site}))$, we apply the BiCEP method on the data compilation using a variety of values for its standard deviation (see Table 2). The main effect of varying σ_{site} is that for smaller values, the estimates of B_m and \vec{k}_m for specimens are "pulled" closer to the line being fitted at a site level (see Figure 5a,b). For our estimate of B_{anc} , this means that sites with fewer specimens will be more precise, as it is unlikely that specimen B_m will deviate strongly from the mean. For sites with many specimens, there is little effect as σ_{site} is well constrained by the data.

From Table 2, we see that changes to $P(\sigma_{site})$ seem to have little influence on the effectiveness of the model, as all our f_{acc} values are the same for our linear model regardless of the prior distribution used. We can also see graphically in Figure 7 that our precision is low for sites with small number of specimens (M). Because of this, we favor the version of the model with a 5 μ T standard deviation on $P(\sigma_{site})$, as models with higher standard deviations reduce precision without capturing any more sites within their 95% credible intervals.

⁵⁴³ 3.3 Order of polynomial fit

The results for our test sites (Table 2) demonstrate that increasing the order of the 544 polynomial fit decreases the precision of the estimate as demonstrated by reduced val-545 ues of f_{prob} . This is expected as there are more parameters to be estimated with the same 546 number of data. The level of accuracy is not significantly improved by increasing the model 547 order. The best quadratic model produced a $\tilde{\Delta}_{median}$ of 1.6 μ T, which is not a signif-548 icant improvement over the value of 1.7 μ T for the best linear model to account for the 549 reduction in precision. The number of passing sites is reduced for the cubic model, in-550 551 dicating that the sampler is struggling to fit this model. Consequently, the cubic model produces more inaccurate and less precise results. For this reason, we assume a linear 552 relationship between B_m and \vec{k}_m . 553

⁵⁵⁴ 3.4 Sampler Diagnostics

Site Name	Worst \hat{R}	n_{eff}	f_{div}
1991-1992 Eruption Site	1.00	59741	0.00
hw108	1.00	77959	0.00
hw123	1.01	11687	0.00
hw126	1.00	36130	0.01
hw128	1.00	78978	0.00
hw201	1.00	10641	0.01
hw226	1.00	7139	0.05
hw241	1.00	66565	0.00
BR06	1.01	451	0.00
Р	1.00	62252	0.00
VM	1.05	1447	0.00
BBQ	1.00	63082	0.00
rs25	1.00	5614	0.00
rs26	1.00	11866	0.00
rs27	1.00	22211	0.00
remag-rs61	1.00	26746	0.00
remag-rs62	1.00	16916	0.00
remag-rs63	1.00	3788	0.00
remag-rs78	1.00	12388	0.00
kf	1.02	2712	0.00
Hawaii 1960 Flow	1.00	60184	0.00
SW	1.00	36390	0.00
TS	1.00	56518	0.00
ET1	1.01	995	0.00
ET2	6.93	6	0.03
ET3	1.01	424	0.00
Synthetic60	1.00	36572	0.01
LV	1.02	5931	0.08
MSH	2.78	24	0.45
FreshTRM	1.00	81007	0.00

Table 3. Sampler diagnostics (see Section 2.5 for an explanation of each diagnostic) for each site using the BiCEP method.

The sampler diagnostics for each site are given in Table 3. Indicators of poor MCMC sampler performance (worst $\hat{R} > 1.1$, low n_{eff} , high f_{div}) tend to occur at sites with four or fewer specimens, or for specimens where the Arai plots are extremely scattered and the sampler struggles to fit them. In the latter case, it may be possible to exclude these specimens by looking at which specimen level parameters have high \hat{R} , as this indicates that fitting a circle to these specimens is inappropriate. We did not exclude specimens on this basis in our analysis, however, we include an option to do this in the BiCEP GUI software (see Appendix 6.3).

The prevalence of high \hat{R} for sites with low numbers of specimens indicates that to get a strongly reproducible answer from this method, paleomagnetists ought to measure five or more specimens per site. In practice, most studies already do this in order to have enough specimens that pass the chosen selection criteria, yet many specimens may be excluded from analysis. Here, we can use all of the specimens measured so there may be no additional burden.

3.5 Summary of Results

After testing all of our contingent models, we prefer the model which assumes the 570 relationship between B_m and k_m is linear, and which uses a 5 μ T standard deviation on 571 $P(\sigma_{site})$. This model performs better than classical sets of selection criteria, either pass-572 ing a greater number of sites (than CCRIT, PICRITMODk, SELCRITMODk) or hav-573 ing significantly higher accuracy and precision (than PICRIT, SELCRIT). Our preci-574 sion increases for sites for which the number of specimens is large, similar to calculat-575 ing the standard error of the mean when using selection criteria. Unlike selection crite-576 ria, the BiCEP method does not require exclusion of large numbers of specimens to ob-577 tain an accurate result, which leads us to prefer it over those methods. 578

579 4 Discussion

580

569

4.1 Advantages of BiCEP compared to selection criteria

BiCEP has significant advantages over the classical selection criteria approach. Firstly, 581 we obtain paleointensity estimates for all sites with at least three specimens, including 582 some which do not contain any specimens that would pass classical selection criteria (see 583 Figure 7). In most cases, our estimates have similar or higher accuracy than the selec-584 tion criteria approach (evidenced by Δ_{median} and Figure 7), and this is accomplished 585 while only excluding specimens from the analysis which were not fully demagnetized. In 586 some cases, our method yields results even if none of the selection criteria accept any spec-587 imens or are inaccurate. For example, for sites remag-rs78 and Synthetic60, our strict 588 criteria (CCRIT, PICRITMODk, SELCRITMODk) produce no results, and our more 589 permissive criteria (PICRITMOD, SELCRIT) produce less accurate (and in the case of 590 Synthetic60, much less precise) results than BiCEP. 591

Secondly, the increasing precision of our paleointensity estimate as the number of 592 specimens increases allows for an improved workflow when compared to classical pale-593 ointensity criteria. Instead of needing a minimum number of specimens to pass our se-594 lection criteria, we can keep measuring specimens until we reach a desired level of pre-595 cision. We discuss this workflow in more detail in Section 4.3. The property of increas-596 ing precision with number of specimens is inherent to Bayesian models and can also be 597 found in the method of Kosareva et al. (2020), although their method does not include 598 the bias correction found in our method. 599

Thirdly, the BiCEP method propagates the uncertainties from a specimen to the site level. Specimens with more scattered (or non linear, or non circular) Arai plots will have less influence over the specimen mean than those with highly linear Arai plots. In addition to this, the BiCEP method foregoes the need for criteria which are concerned with the length of the line on the Arai plot used to make an interpretation, like the NRM Fraction (e.g., FRAC of Shaar & Tauxe, 2013). Using a set of temperatures with small FRAC will cause an increase in the uncertainty in \vec{k} (see Figure 4e, f), which will cause this specimen to have less effect on the estimate of B_{anc} , without excluding it from the analysis entirely. We discuss this further in Section 4.5.



4.2 Predictive ability of the method

609

Figure 8. Example of the BiCEP method applied to three subsets of 6 specimens from site hw108 ($B_{exp} = 39.3 \ \mu\text{T}$). The left column in each subplot shows histograms of the BiCEP results, and the right column shows plots of specimen \vec{k} vs B_{anc} with the BiCEP line fits. Light green shaded regions with dashed edges represent the 2σ interval of the PICRITMOD estimate for these subsamples. In a) there is a small range of \vec{k} and B_{anc} values which leads to an imprecise estimate of c, but an accurate and precise estimate of B_{anc} . In b) there is a large range of values on \vec{k} , but all specimens have high \vec{k} . This leads to an estimate with a relatively precise estimate of c, and an accurate but imprecise estimate of B_{anc} . In c) there is a reasonably small range of values on B_{anc} , and the relationship between B_{anc} and \vec{k} is not linear, but BiCEP attempts to find a linear model. This leads to an imprecise and inaccurate estimate of both c and B_{anc} .

Although our results are promising, it is worth noting that traditional selection cri-610 teria also perform well for the majority of our sites. To see if the BiCEP method offers 611 accurate results with poorer quality data, we subsampled results from site hw108, which 612 had a range of good and poor quality specimens. Figure 8 shows the results of BiCEP 613 applied to three different subsets of six specimens taken from this site, along with the 614 results of the PICRITMOD criteria applied to this site (in green). It is worth noting that 615 only the specimens in Figure 8a would pass the CCRIT criteria which gave a highly ac-616 curate result (within 1 μ T), or any of our more restrictive criteria. 617

We identify three behaviours for which BiCEP results deviate from a linear model with high precision on the slope and intercept. Figure 8a shows a subset of specimens for which the range of \vec{k} values of the specimens is very small, and so the uncertainty

of the slope of the linear relationship between \vec{k}_m and $B_m(c)$ is high. In this case, how-621 ever, because these specimens all have k close to zero, the estimate of B_{anc} is accurate 622 and precise. Figure 8b shows a different subset of specimens for which the range of k val-623 ues is large, but there are no k values close to zero. This results in an estimate of B_{anc} 624 which is still accurate, but imprecise due to the uncertainty in extrapolating the linear 625 relationship between k_m and B_m back to zero. The PICRITMOD result for this subset 626 returns an average value which underestimates B_{exp} by around ~13 μ T or ~30%, and 627 criteria using the curvature criterion return no values, as all specimens have curvature 628 values higher than the threshold. The high uncertainty in B_{anc} might still be considered 629 a problem, but this result indicates that measuring more specimens would likely yield 630 a more precise result. 631

Figure 8c shows a set of specimens where the range of \vec{k} is low, so the \vec{k}_m versus 632 B_m relationship is not particularly linear. BiCEP attempts to find a linear trend with 633 these data, and extrapolates back to a B_{anc} which is both highly inaccurate and impre-634 cise. This might be considered a problem for BiCEP, but it is possible to detect such be-635 havior as the uncertainty on both B_{anc} and the slope relating the B_m versus k_m (c) are 636 large. This indicates to us that we can use a metric of the uncertainty in both the slope 637 and intercept of the linear fit in BiCEP to decide whether a site level result is accurate 638 or not. This leads us to a laboratory workflow which uses BiCEP results to decide if a 639 site is acceptable, might be acceptable with further work or is unlikely to give a reason-640 able result. 641

642

4.3 Workflow with BiCEP

Figure 9 plots the 95% credible interval on B_{anc} as a percentage against the 95% 643 credible interval on c as a proportion of the median \tilde{B}_{anc} for all sites where $\hat{R} > 1.1$. 644 The sizes of the points on the plot represents the number of specimens per site (M), with 645 squares representing sites with M < 5. The colors show the percentage deviation from 646 B_{exp} using BiCEP, with redder colors for more inaccurate results. With the exception 647 of two sites (VM and hw123), as the number of specimens increases, sites trend towards 648 the bottom left region of this plot, indicating an increase in precision. This has dimin-649 ishing returns as the number of specimens increases above five. The increase in preci-650 sion is also accompanied with an increase in accuracy. Almost all sites with an estimated 651 precision on B_{anc} better than 40% have median values within 20% of B_{exp} . Our outlier 652 sites VM and hw123, which are imprecise despite having large numbers of specimens (M=12653 and 18 respectively), are also inaccurate. This indicates that the width of the 95% cred-654 ible intervals is a useful statistic for diagnosing inaccuracy in the BiCEP method. 655

We have divided Figure 9 into four regions (labeled A-D). Sites in region A have 656 high precision on both B_{anc} and c and are representative of the results for the major-657 ity of sites in this study; sites in this region are highly accurate. Sites in region B have 658 high precision on B_{anc} (better than 40%, which for a Gaussian distribution would be equiv-659 alent to a standard deviation of $\pm 10\%$) but low precision on c (95% credible interval on 660 $c/B_{anc} > 1$). These sites are usually analogous to the example shown in Figure 8a, with 661 low Arai plot curvature and similar intensities for all specimens. Sites in region C have 662 high precision on B_{anc} but a low precision on c. These sites may have a large number 663 of curved specimens which follow a linear trend that can be extrapolated back to the cor-664 rect B_{exp} , and are analogous to our example in Figure 8b. Region D is representative 665 of the worst constrained estimates, with low precision on B_{anc} and c. Sites in this re-666 gion may have highly inaccurate estimates of B_{anc} , often with low M. If these sites have 667 high M, they may be similar to our example in Figure 8c in which a linear relationship 668 between B_{anc} and c is not well determined, and the average $|\vec{k}|$ is large, leading to an 669 inaccurate estimate of B_{anc} . 670



Figure 9. Plot of the 95% credible interval on B_{anc} against the 95% confidence interval on c (slope between intensity estimate and \vec{k}), normalized by the median B_{anc} for all sites with $\hat{R} < 1.1$. Circles indicate sites for which the number of specimens $M \geq 5$, and squares indicate sites where M < 5. Colors indicate the deviation of the median value of the estimate from the expected site value (B_{exp}) as a percentage. The size of markers is used to represent M. The horizontal dashed line indicates a value of 40% for the full width of the 95% credible interval on B_{anc} , which for a Gaussian distribution would correspond to a standard deviation of $\pm 10\%$. The vertical dashed line represents a value of 1 for the 95% confidence interval of c/\tilde{B}_{anc} . Suggested workflow for sites in regions: A) or B), accept the site or continue measuring if improve precision is desired. C) Continue measuring specimens, as improved precision is likely. D) If $M \geq 5$ stop measuring the site as further effort is likely to be futile. Otherwise continue measuring specimens until M = 5.

Considering the region in which a particular site plots leads to a workflow based 671 on the likelihood of success. In general, sites with very low numbers of specimens, (M =672 2 or 3), will begin in region D, and migrate to regions C, B or A as M increases to around 673 five. If a site has migrated to region A or B after five specimens have been measured, 674 then we likely have an accurate and precise estimate of B_{anc} , and we can finish measur-675 ing specimens (or continue to measure if a higher level of precision is desired). If a site 676 has migrated to region C, it is likely that our estimate is accurate and that our uncer-677 tainty in B_{anc} can be reduced by increasing the number of specimens. If our site remains 678 in region D after five specimens have been measured, the site level estimate may be in-679 accurate, and measuring more specimens would be unlikely to reduce the site level un-680 certainty. 681

Because the regions in Figure 9 define a workflow based on measuring five spec-682 imens, we wanted to test whether our methodology could identify sites which will remain 683 in region D after a large number of specimens were measured. We randomly subsampled 684 100 sets of 5 specimens from sites hw108, VM and hw123 and calculated B_{anc} and c us-685 ing BiCEP. Site hw108 was chosen because contains specimens which exhibit a wide range 686 of behaviours, with a large number of specimens having high k, but yields an accurate 687 result. Sites VM and hw123 were chosen because these are our sites which remain in re-688 gion D after measuring a large number of specimens. Our results for these three sites 689 are given in Supplemental Figure S2. Our subsampled hw108 obtained results in region 690

⁶⁹¹ D 5 times out of 100, whereas our subsampled VM and hw123 obtained results in region ⁶⁹² D 94 and 90 times respectively. This indicates that sites which remain in region D af-⁶⁹³ ter measuring 5 specimens are likely to remain there after measuring many more, and ⁶⁹⁴ so measuring more specimens is usually a futile effort.

695

4.4 Overly precise estimates of B_{anc}

⁶⁹⁶ The BiCEP method has a lower f_{acc} than CCRIT, despite having a similar degree ⁶⁹⁷ of accuracy when using a metric like $\tilde{\Delta}_{median}$. The reason for this is that the increas-⁶⁹⁸ ing precision on the BiCEP estimate leads to estimates which are highly precise when ⁶⁹⁹ M is large. This is the case shown in Figure 6d.

Labeling sites with extremely high precision in the estimate as inaccurate may be 700 misleading, as we have not taken into account uncertainties in the value of the expected 701 fields at the sites in this study. For example, using differences between the observed di-702 rections and the IGRF, Yamamoto and Hoshi (2008) quoted the expected value at the 703 site "SW" as $46.0\pm2.6 \ \mu\text{T}$, which is just consistent with the 95% credible interval for our 704 specimen (48.2-49.7 μ T). Because of this, we prefer to use f_{prob} as a metric of how well 705 a model performs as it allows for a few μT of uncertainty in the expected field value. Ad-706 ditionally, Yamamoto and Yamaoka (2018) suggested that the IZZI-Thellier results for 707 sites SW and TS may be biased slightly high due to acquisition of a thermo-chemical re-708 manent magnetization (TCRM), which is not detectable by our method. Yamamoto et 709 al. (2003) also invoke a TCRM mechanism to explain the paleointensity overestimate for 710 the Hawaii 1960 Flow, which is another of their sites for which we overestimate the ex-711 pected intensity (see Figure 7 and Supplementary Data Set S1). We note that Cromwell 712 et al. (2015) also sampled the 1960 flow (hw241 which targeted the fine grained flow top) 713 and all selection criteria resulted in accurate results, with BiCEP producing the tight-714 est confidence interval. 715

716

4.5 Exclusion of measurement level data

It is frequently possible to improve the accuracy and precision of results by find-717 ing the 'best' set of temperature steps to use in the intensity interpretation. Two situ-718 ations frequently occur for which this might be justified. The first is the case in which 719 thermochemical alteration occurs at high temperature (e.g., Figure 4e). For such spec-720 imens, the low temperature measurements can be used to make a paleointensity estimate 721 (colored dots in the figure). Figures 4e and f show how our method can be used on a re-722 duced range of temperature steps on the Arai plot at the cost of precision. The plot of 723 circle fits (green lines in Figure 4e) demonstrates that the Arai plot interpretations are 724 poorly constrained and can continue in any direction after the last temperature step cho-725 sen. This results in a higher uncertainty in the curvature associated with this (Figure 4f). 726 The second case in which a portion of the data could be excluded from the calculation, 727 would be when the magnetization has multiple components (Figure 10a). In such a case, 728 a paleointensity estimate can only be made using the small range of temperature steps 729 that correspond to the characteristic component. We currently do not have an objec-730 tive method to choose which set of temperature steps on the Arai plot to use. We sug-731 gest that decisions about which data points to include should not be made based on the 732 original in-field or zero field Arai plot measurements (dots in the Arai plots), but rather 733 exclusively on deviating pTRM checks (triangles in, e.g., Figure 4e) or other indicators 734 of alteration for the first case and on the directions of the magnetization vector (it must 735 trend to the origin and be well defined) in the second case, e.g., Figure 10a. 736

Caution should be used when excluding a particular temperature steps for reasons
 other than this. If the set of temperature steps chosen does not represent the character istic component of magnetization, this can alter the outcome of the BiCEP method, especially if a large part of the Arai plot is excluded. Additionally, excluding more points

on the Arai plot tends to increase the chance that a specimen will cause \hat{R} failure. As 741 such, we recommend using as many points on the Arai plot as possible unless done for 742 one of the reasons stated above.

743

744



Figure 10. a) Example of vector endpoint diagram for specimen FB2-B1 from Lisé-Pronovost et al. (2020). The magnetization is rotated so that the principal component of the TRM direction for all steps lies along the x axis. Green line fit to the low temperature component and cyan line fit to the high temperature component. b) Arai plot and c) "corrected" Arai plot for a specimen from the data shown in b). NRM values for the low temperature component (filled circles) are usually calculated by taking the magnitude of the vector endpoint (blue dashed lines in the vector endpoint diagram in a). In c), these NRM values are calculated by vector subtracting the high temperature component (cyan line), taking the magnitude of our new NRM vectors (distance along green line), and adding the magnitude of the low temperature component (length of cyan line). Both b) and c) are scaled by the total NRM distance along both components (total distance along both green and cyan lines).

4.6 Application to multi-component magnetizations

We test an application of the BiCEP method on data with multi-component di-745 rections as shown in Figure 10a using the data of Lisé-Pronovost et al. (2020). The data 746 are from Scottish firebricks which were used in a foundry in Australia. The date and lo-747 cation of firing are both well constrained, hence we have a reasonably well constrained 748 value for B_{exp} . The bricks all contained a low temperature component associated with 749 the Australian field. Some also displayed a high temperature component associated with 750 the original firing in Scotland as shown in Figure 10a. Lisé-Pronovost et al. (2020) al-751 ready have interpretations which separate these components in the original study. To 752 account for the change in direction of the NRM, we subtract the high temperature com-753 ponent from the low temperature component, and then add the magnitude of these val-754 ues to the magnitude of the low temperature component (see Figure 10 for a graphical 755 explanation). The vector subtraction is necessary for the low-temperature component 756 as we need a total TRM $(pTRM_{max})$ to scale by in order to penalize the result for shorter 757 components. We then proceed to use the BiCEP method as previously described, using 758 the original interpretations for the different components. For the sake of simplicity, we 759 do not perform the magnetomineralogical change (MMC) correction (Valet et al., 1996). 760 We also do not apply the corrections for anisotropy of TRM or cooling rate with these 761 data, as they appeared to be negligible. Of course these could be applied in the usual 762 fashion if necessary. 763

We display the results from multi-component remanences in Figure 11. We find that 764 for the low temperature, Australian field, component (Figure 11a), our estimates for all 765 firebricks contain the expected answer (61.17 μ T) within the 95% credible interval. Our 766



Figure 11. Expected and predicted intensities on the data of Lisé-Pronovost et al. (2020) using BiCEP (blue circles) and the method used in the original study (black diamonds). a) Results for the low temperature component (Australia, expected field value 61.17 μ T) for each firebrick. b) Results for the high temperature component (Scotland, expected field value 48.3 μ T), where this component was present. The dashed blue line indicates that the MCMC sampler failed to converge for site FB1.

interpretation for site FBG is slightly less accurate than the original analysis but with
 much higher precision. This difference is likely caused by not applying the MMC cor rection, as the specimens at this site were mostly of good quality, with none being ex cluded from the original analysis.

For the high temperature component (Figure 11b) our results behave differently. 771 The sampler does not converge for site FB1, indicating too few specimens in the anal-772 ysis. For site FB2, we have a result that is less accurate, but more precise than in the 773 original study. The lack of MMC correction may contribute to the decreased accuracy 774 in this example, whereas the reduced precision is likely caused by the smaller length of 775 the interpretation on the Arai plot, leading to a higher uncertainty in the curvature for 776 that specimen. Our results for this study demonstrate that BiCEP will obtain precise 777 estimates for components which represent most of the magnetization, and be imprecise 778 for components which have small NRM fraction. 779

4.7 Implications for bias in curved Arai plots

780

The success of our method demonstrates that Arai plot "curvature" or sagging does 781 lead to a progressive bias in paleointensity estimation which increases as the amount of 782 curvature increases as described by Tauxe et al. (2021) and strongly suggested by the 783 data of Krása et al. (2003) (see Figure 1). Our estimates are made by using the tangent 784 to a circle fit rather than fitting a line to part of the data, so one might expect them to 785 be biased. However, it has been demonstrated by e.g. the data of (Krása et al., 2003) 786 that fitting lines to the high temperature or low temperature slope of Arai plots yields 787 even more biased results than using the total TRM, which is more similar to the tan-788 gent. The scaling used by our method incorporates the added uncertainty in the line slope 789 and k associated with choosing one of these slopes, which allows for more consistent anal-790 ysis between specimens with interpretations of varying quality. The bias seen generally 791 underestimates paleointensity with higher (positive) curvature, but this is not the case 792 for all sites, some of which exhibit the opposite trend. 793

The assumption of a quasi-linear dependence between the specimen level paleointensities and the curvature of the Arai plot does not have any theoretical basis. This does imply that the curvature is linearly related to the change in TRM susceptibility (or decay of the original magnetization) between the original and lab coolings, a relationship which should be further investigated. We stress that this relationship only needs to be loosely followed for our method to work. In cases where there does not appear to be a strong linear relationship between B_m and \vec{k}_m (e.g. in Figure 8a), an accurate paleointensity estimate is still possible if there are enough specimens with low $|\vec{k}|$, as the intercept of the linear fit is still well constrained even if the slope is not. Conversely, if there are few specimens with low $|\vec{k}|$ and there is a poor linear relationship, then both the slope and intercept are poorly constrained, resulting in a huge uncertainty in B_{anc} , as is seen in Figure 8c.

5 Conclusions

807

808

- We present a new Bayesian method (BiCEP) which accounts for bias in paleointensity estimates in specimens.
- Instead of excluding specimens from the paleointensity analysis in the traditional (binary) selection criteria based approach, our method predicts an amount of bias for each specimen, using the curvature of the Arai plot as a metric of non-linearity and a predictor of bias. In this way, the BiCEP method is quite different from the recently published Bayesian approach of Kosareva et al. (2020).
- When tested on a compilation of sites for which an approximate paleointensity is known *a priori*, our method yields levels of accuracy and precision similar to, or better than restrictive paleointensity criteria, whilst accepting as many results as permissive criteria.
- Our method generates some slightly inaccurate paleointensity estimates with high levels of precision, but these can generally be explained with inaccuracies in the expected field (see Section 4.4).
- The BiCEP method handles uncertainties in a different way than using classical selection criteria, as the uncertainty in site level estimates decreases as the number of specimens increases, but this uncertainty remains high when the number of specimens is low due to inclusion of prior information. The Bayesian uncertainties are in this way more similar to the 'extended error bars' in the Thellier_GUI auto-interpreter of Shaar and Tauxe (2013).
- We propose a workflow in which sites are accepted and measurement of specimens can cease once a desired level of confidence in the site level estimate has been reached. Sites which do not reach this level of confidence after measuring several (> 5) specimens likely do not contain useful information and can be discarded.

⁸³¹ Data Availability Statement

Data used in this paper may be found in the MagIC database at: https://earthref .org/MagIC/17104/0326fdaa-4bcf-44f3-989d-0116b9a2fb75 for review and will be available to the public at https://earthref.org/MagIC/17104 on publication.

6 Appendix

836

6.1 Change of variables

In Section 2.2.1 we mention that we need to use a change of variables to get from our original circle fitting parameters R, x_c, y_c to our new set of parameters \vec{k}, D, ϕ . We can use the Jacobian of the parameter change to get the new formula for the posterior probability under our new parameters:

$$P(D,\phi,\vec{k}|x,y) = P(x_c,y_c,R|x,y) \left| \frac{\partial(x_c,y_c,R)}{\partial(D,\phi,\vec{k})} \right|.$$
(20)

⁸⁴¹ We can evaluate this Jacobian as:

$$\left|\frac{\partial(x_c, y_c, R)}{\partial(D, \phi, \vec{k})}\right| = \left|\frac{\vec{k}}{|\vec{k}^3|} \left(D + \frac{1}{\vec{k}}\right)(\cos\phi + \sin\phi)\right|.$$
(21)

So our posterior looks like:

$$P(D,\phi,\vec{k}|x,y) \propto \left(\sum_{n=1}^{N} \sqrt{((D+\frac{1}{\vec{k}}\cos\theta) - x_n)^2 + ((D+\frac{1}{\vec{k}}\sin\theta) - y_n)^2} - \frac{1}{|\vec{k}|})^2}\right)^{-N/2} \\ \left|\frac{\vec{k}}{|\vec{k}^3|} \left(D + \frac{1}{\vec{k}}\right)(\cos\phi + \sin\phi)\right| P(\vec{k},\phi,D).$$
(22)

843

844

870

877

842

6.2 Markov chain Monte Carlo sampling

The Markov chain Monte Carlo (MCMC) sampling method generates a set of samples from the posterior probability distribution of B_{anc} which allows us to approximate it. We use the python bindings for the Stan software package (http://mc-stan.org) to generate these samples which provides diagnostic information and runs relatively quickly. For each site we run four Markov chains and generate 30,000 samples of B_{anc} in each chain. We discard the first half of the chain as 'burn in' for a total of 60,000 samples.

Stan provides several diagnostics that tell us whether we have successfully sampled 851 the posterior distribution. These include the R score (Gelman & Rubin, 1992) which tells 852 us about the convergence between chains, and is required to be between 1.1 and 0.9 which 853 is necessary for convergence, the effective sample size, n_{eff} which should be large (> 1000) 854 for a good sample and the number of divergent transitions (f_{div}) which should be zero 855 in ideal cases. In most cases our results display high degrees of convergence with R close 856 to 1 and high effective sample sizes. Some sites included divergent transitions in small 857 numbers. These seem to occur at a specimen level for specimens where the posterior dis-858 tribution of one of the circle parameters is long-tailed. In theory this can mean the pos-859 terior was inefficiently sampled, but because these specimens generally have large uncertainties on their k parameter, the final results do not change, even under a change of 861 parameters. The sampler struggled to converge, with $\hat{R} > 1.1$ for several sites with very 862 few specimens, where once again the distributions are extremely long tailed. The sam-863 pler also did not converge for site MSH, where the Arai plots were so non linear, with 864 few points, that BiCEP struggled to fit circles to them. We consider these sites to have 865 "failed" using our method (grade of 'D' in Figure 9) and use the prior distribution on 866 B_{anc} (uniform between 0 and 250 μ T) as an estimate of their intensity. We calculate the 867 R furthest from unity, the n_{eff} for B_{anc} and the proportion of divergent samples f_{div} 868 for our model. 869

6.3 Code and GUI

We present a simple GUI that can perform the BiCEP method on data in the MagIC format. The code uses Jupyter notebooks and can be found at (http://github.com/ bcych/BiCEP_GUI) and contains a readme file detailing how to use the notebook. The GUI can also be accessed at the Earthref JupyterHub site (http://jupyterhub.earthref .org). To access the GUI this way:

- Sign up to Earthref at (http://earthref.org)
 - Navigate to the Earthref JupyterHub site at (http://jupyterhub.earthref.org)
- Open and run all the cells in the "BiCEP GUI Setup.ipynb" notebook.

- Upload MagIC formatted "sites", "samples", "specimens" and "measurements" files to the BiCEP_GUI directory in JupyterHub. These can be formatted using pmag_gui (Tauxe et al., 2016).
 - Open the BiCEP GUI notebook and press the "App Mode" button.
- 883

882

For more detailed instructions, read the included readme file at the github site.

Acknowledgments

We are deeply grateful to Lennart de Groot and Greig Paterson for their very helpful reviews and for the advice and guidance given by Andrew Roberts, David Heslop and Joseph Wilson. This research was supported in part by NSF Grant EAR1827263 to LT. We are also grateful to Agnes Lisé-Pronovost for sharing her measurement level data for use in section 4.6.

890 References

- ⁸⁹¹ Biggin, A. J. (2010). Paleointensity database updated and upgraded. *EOS*, *91*, 15.
- Biggin, A. J., Perrin, M., & Dekkers, M. J. (2007). A reliable absolute palaeointensity determination obtained from a non-ideal recorder. *Earth and Planetary Science Letters*, 257(3), 545-563. doi: https://doi.org/10.1016/j.epsl.2007.03
 .017
- Bowles, J., Gee, J. S., Kent, D. V., Perfit, M. R., Soule, S. A., & Fornari, D. J.
 (2006). Paleointensity applications to timing and extent of eruptive activity,
 9°-10°n east pacific rise. *Geochemistry, Geophysics, Geosystems*, 7(6). doi: https://doi.org/10.1029/2005GC001141
- Chernov, N., & Lesort, C. (2005). Least squares fitting of circles. Journal of Mathematical Imaging and Vision, 23(3), 239–252. doi: 10.1007/s10851-005-0482-8
- Cromwell, G., Tauxe, L., Staudigel, H., & Ron, H. (2015). Paleointensity estimates from historic and modern hawaiian lava flows using glassy basalt
 as a primary source material. *Phys. Earth Planet. Int.*, 241, 44–56. doi: 10.1016/j.pepi.2014.12.007
- 906Donadini, F., Kovacheva, M., Kostadinova, M., Casas, L., & Pesonen, L. (2007).907New archaeointensity results from scandinavia and bulgaria: Rock-magnetic908studies inference and geophysical application.909Planetary Interiors, 165(3), 229-247.910j.pepi.2007.10.002
- Dunlop, D., & Ozdemir, O. (2001). Beyond Néel's theories: thermal demagnetization
 of narrow-band partial thermoremanent magnetization. *Phys. Earth Planet. Int.*, 126, 43-57.
- Gelman, A., Carlin, J. B., Stern, H. S., & Rubin, D. B. (2004). Bayesian data analy sis (Second ed.). Chapman & Hall/CRC, Boca Raton, FL.
- Gelman, A., & Rubin, D. B. (1992, 11). Inference from iterative simulation using
 multiple sequences. *Statist. Sci.*, 7(4), 457–472. doi: 10.1214/ss/1177011136
- Hoffman, K. A., Constantine, V. L., & Morse, D. L. (1989). Determinaton of absolute palaeointensity using a multi-specimen procedure. *Nature*, 339, 295-297.
- Königsberger, J. G. (1938). Natural residual magnetism of eruptive rocks. Ter restrial Magnetism and Atmospheric Electricity, 43(3), 299-320. doi: 10.1029/
 TE043i003p00299
- Korte, M., Donadini, F., & Constable, C. G. (2009). Geomagnetic field for 0–3 ka: 2.
 a new series of time-varying global models. *Geochemistry, Geophysics, Geosystems*, 10(6). doi: https://doi.org/10.1029/2008GC002297
- Kosareva, L. R., Kuzina, D. M., Nurgaliev, D. K., Sitdikov, A. G., Luneva, O. V.,
 Khasanov, D. I., ... Spassov, S. (2020). Archaeomagnetic investiga-

928	tions in Bolgar (Tatarstan). Stud. Geophys. Geod., $64(2)$, 255–292. doi: 10.1007/s11200.019.0403.3
929	Vráce D. Heynemann C. Leenhardt D. & Deterson N. (2002) Experimental
930	Krasa, D., Heunemann, C., Leonnardt, K., & Petersen, N. (2005). Experimental
931	procedure to detect multidomain remainence during themer-themer experi-
932	ments. Phys. Chem Edith $(A/B/C)$, 28(10), 081 - 087. (Paleo, Rock and Environmental Magnetic 2002) drive 10.1016 (C1474.7065 (02)00122.0)
933	Environmental Magnetism 2002) doi: $10.1016/514/4-7065(03)00122-0$
934	Lise-Pronovost, A., Mallett, T., & Herries, A. I. R. (2020). Archaeointensity of
935	nineteenth-century scottish firebricks from a foundry in melbourne, australia:
936	comparisons with field models and magnetic observatory data. Geological Soci-
937	ety, London, Special Publications, 497(1), 27–45. doi: 10.1144/SP497-2019-72
938	Muxworthy, A. R., Heslop, D., Paterson, G. A., & Michalk, D. (2011). A preisach
939	method for estimating absolute paleoneld intensity under the constraint of
940	using only isothermal measurements: 2. experimental testing. Journal of
941 942	<i>Geophysical Research: Solid Earth</i> , 116 (B4). doi: https://doi.org/10.1029/ 2010JB007844
943	Nagata, T., Arai, Y., & Momose, K. (1963). Secular variation of the geomagnetic to-
944	tal force during the last 5000 years. J. Geophys. Res., 68(18), 5277-5281. doi:
945	0.1029/j.2156-2202.1963.tb00005.x
946	Nagy, L., Williams, W., Muxworthy, A. R., Fabian, K., Almeida, T. P., Conbhuí,
947	P. Ó., & Shcherbakov, V. P. (2017). Stability of equidimensional pseudo-
948	single-domain magnetite over billion-year timescales. Proc. Natl. Acad. Sci.
949	U.S.A., 114(39), 10356-10360.doi: 10.1073/pnas.1708344114
950	Néel, L. (1949). Théorie du traînage magnétique des ferromagnétiques en grains fins
951	avec applications aux terres cuites. Ann. géophys., 5, 99–136.
952	Paterson, G. A. (2011). A simple test for the presence of multidomain be-
953	havior during paleointensity experiments. J. Geophys. Res., 116. doi:
954	$10.1029/2011 \mathrm{JB008369}$
955	Paterson, G. A., Biggin, A. J., Yamamoto, Y., & Pan, Y. (2012). Towards the ro-
956	bust selection of Thellier-type paleointensity data: The influence of experimen-
957	tal noise. Geochem. Geophys. Geosyst., 13(5). doi: 10.1029/2012GC004046
958	Paterson, G. A., Muxworthy, A. R., Roberts, A. P., & Mac Niocaill, C. (2010).
959	Assessment of the usefulness of lithic clasts from pyroclastic deposits for pa-
960	leointensity determination. Journal of Geophysical Research: Solid Earth,
961	115(B3). doi: https://doi.org/10.1029/2009JB006475
962	Paterson, G. A., Tauxe, L., Biggin, A. J., Shaar, R., & Jonestrask, L. C. (2014). On
963	improving the selection of thellier-type paleointensity data. Geochem. Geophys.
964	<i>Geosyst.</i> , 15(4), 1180–1192. doi: 10.1002/2013GC005135
965	Pick, T., & Tauxe, L. (1993). Holocene paleointensities: Thellier experiments on
966	submarine basaltic glass from the east pacific rise. Journal of Geophysical
967	Research: Solid Earth, 98 (B10), 17949-17964. doi: https://doi.org/10.1029/
968	93JB01160
969	Santos, C. N., & Tauxe, L. (2019). Investigating the accuracy, precision, and cooling
970	rate dependence of laboratory-acquired thermal remanences during paleoin-
971	tensity experiments. Geochem., Geophys., Geosyst., 20(1), 383–397. doi:
972	Shaan D. Dan H. Tauwa I. Kassal D. & Aman A. (2011). Deleannemetic field
973	intensity derived from non ad. Testing the thellion iggi technique on md der
974	and a new bootstrap procedure Farth and Planatary Science Letters 210(2)
975 976	213-224. doi: https://doi.org/10.1016/j.epsl.2011.08.024
977	Shaar, R., Ron, H., Tauxe, L., Kessel, R., Agnon, A., Ben-Yosef, E., & Feinberg,
978	J. M. (2010). Testing the accuracy of absolute intensity estimates of the
979	ancient geomagnetic field using copper slag material. Earth and Planetary Sci-
980	ence Letters, $290(1)$, 201-213. doi: https://doi.org/10.1016/j.epsl.2009.12.022
981	Shaar, R., & Tauxe, L. (2013). Thellier gui: An integrated tool for analyzing pa-
982	leointensity data from theilier-type experiments. Geochem. Geophys. Geosys.,

983	14, 677–692. doi: doi:10.1002/ggge.20062
984	Shaw, J. (1974). A new method of determining the magnitude of the paleomanetic
985	field application to 5 historic lavas and five archeological samples. <i>Geophys. J.</i>
986	<i>R. astr. Soc.</i> , 39, 133-141.
987	Tanaka, H., Hashimoto, Y., & Morita, N. (2012, 05). Palaeointensity determina-
988	tions from historical and Holocene basalt lavas in Iceland. Geophysical Journal
989	International, 189(2), 833-845. doi: 10.1111/j.1365-246X.2012.05412.x
990	Tauxe, L., Santos, C., Cych, B., Zhao, X., Roberts, A., Nagy, L., & Williams, W.
991	(2021). Understanding non-ideal paleointensity recording in igneous rocks:
992	Insights from aging experiments on lava samples and the causes and conse-
993	quenes of 'fragile' curvature in arai plots. <i>Geochem. Geophys. Geosyst.</i> , 22,
994	e2020GC009423. doi: 10.1029/2020GC009423
995	Tauxe, L., Shaar, R., Jonestrask, L., Swanson-Hysell, N. L., Minnett, R., Koppers,
996	A. a. P., Fairchild, L. (2016). PmagPy: Software package for paleomag-
997	netic data analysis and a bridge to the magnetics information consortium
998	(MagIC) database. Geochem., Geophys., Geosyst., $17(6)$, $2450-2463$. doi:
999	10.1002/2016GC006307
1000	Tauxe, L., & Yamazaki, T. (2015). Paleointensities. In M. Kono (Ed.), Geomag-
1001	netism (2nd Edition ed., Vol. 5, p. 461-509). Elsevier.
1002	Thébault, E., Finlay, C. C., Beggan, C. D., Alken, P., Aubert, J., Barrois, O.,
1003	Zvereva, T. (2015). International Geomagnetic Reference Field: the 12th
1004	generation. Earth Planets Space, 67(1), 79. doi: 10.1186/s40623-015-0228-9
1005	Thellier, E., & Thellier, O. (1959). Sur l'intensité du champ magnétique terrestre
1006	dans le passé historique et géologique. Ann. Geophys., 15, 285.
1007	Valet, JP., Brassart, J., Le Meur, I., Soler, V., Quidelleur, X., Tric, E., & Gillot,
1008	PY. (1996). Absolute paleointensity and magnetomineralogical changes.
1009	Journal of Geophysical Research: Solid Earth, 101 (B11), 25029-25044. doi:
1010	https://doi.org/10.1029/96JB02115
1011	Williams, W., & Dunlop, D. J. (1989). Three-dimensional micromagnetic modelling
1012	of ferromagnetic domain structure. Nature, 337, 634–637.
1013	Yamamoto, Y., & Hoshi, H. (2008). Paleomagnetic and rock magnetic studies of
1014	the sakurajima 1914 and 1946 and esitic lavas from japan: A comparison of
1015	the ltd-dht shaw and thellier paleointensity methods. <i>Phys. Earth and Planet.</i>
1016	Inter., 167, 118-143.
1017	Yamamoto, Y., Tsunakawa, H., & Shibuya, H. (2003). Palaeointensity study of
1018	the hawaiian 1960 lava: implications for possible causes of erroneously high
1019	intensities. Geophys J Int, $153(1)$, $263-276$.
1020	Yamamoto, Y., & Yamaoka, R. (2018). Paleointensity study on the Holocene surface
1021	lavas on the Island of Hawaii using the Tsunakawa-Shaw method. Front. Earth
1022	Sci., $6.$ doi: 10.3389/feart.2018.00048
1023	Yu, Y., Tauxe, L., & Genevey, A. (2004). Toward an optimal geomagnetic field in-
1024	tensity determination technique. Geochem., Geophys., Geosyst., 5(2). doi: 10

1025 .1029/2003GC000630

Supporting Information for "Bias Corrected Estimation of Paleointensity (BiCEP): An improved methodology for obtaining paleointensity estimates"

Brendan Cych¹, Matthias Morzfeld¹, Lisa Tauxe¹

 $^1 \mathrm{University}$ of California, San Diego

 19500 Gilman Drive, La Jolla, CA, 92093, USA

Contents of this file

1. Figure S1- plots of B_{exp} vs B_{anc} results using the BiCEP method and other sets of selection criteria used in this study.

2. Figure S2- plots of the credible intervals of BiCEP results for subsamples of 5 specimens from sites hw108, VM and hw123.

3. Figures S3 to S32- plots of unpooled B_{anc} and \vec{k} fits and the BiCEP method applied to all sites used in this study.

Additional Supporting Information (Files uploaded separately)

Caption for Data Set S1 - This dataset contains the full set of results shown in Figure
 7 and described in Section 3.

Data Set S1

This data set is a csv file containing the full set of results for each version of the BiCEP method (Linear, Quadratic and Cubic models with 5, 10 and 20 μ T standard deviations

on the prior for σ_{site} , vs classic selection criteria. Columns are named by the model class (Linear, Quadratic, Cubic, or Selection Criteria, followed by the criteria name/prior standard deviation in μ T) and then the parameter in question, either the percentile (2.5%, 50%, 97.5%) or f_prob for the f_{prob} value. The first three columns of the csv file contain the site name, B_{exp} and number of specimens (M).





Figure S1. Plot of B_{anc} vs B_{exp} for the passing sites for BiCEP and each of our criteria. a) BiCEP results, b) PICCRITMOD results, c) PICRITMODk results, d) SELCRIT results, e) SELCRITMODk results, f) CCRIT results. Red shaded area represents $B_{exp} \pm 3 \mu$ T. Note that for BiCEP, most of the severely underestimated and highly inaccurate results are for sites with low numbers of specimens (M < 5) which did not pass the other criteria.



Figure S2. Figure in the style of Figure 9 from the main manuscript. A set of 100 random subsamples of 5 specimens were taken from a) site hw108, b) site VM and c) site hw123. We applied the BiCEP method to these subsamples and plot the 95% credible interval on B_{anc} as a percentage against the 95% credible interval on c/\tilde{B}_{anc} . Dashed lines represent the boundaries between regions A, B, C and D, defined in Section 4.3, and point colors represent the percentage deviation of the median from B_{exp} . It is apparent that for our sites where we obtain an inaccurate answer for a large number of specimens (VM and hw123), our subsamples fall in region D the majority of the time (94/100 and 90/100 respectively). Conversely, our site hw108, which returns an accurate result with higher numbers of specimens only has 5 results in region D. This indicates that we should either continue measuring specimens (region C, 24/100 results) or that we have a relatively accurate and precise result already (regions A and B, 66/100 results). Although few sites/subsamples were used for this analysis, this test indicates that our labelling scheme has a predictive accuracy of 90-95%



Figure S3. Unpooled and BiCEP models applied to site 1991-1992 Eruption Site a). Independent estimates of B_{anc} vs \vec{k} are plotted as green circles. These estimates are made without assuming a linear relationship between B_{anc} and \vec{k} , similar to the analysis in Figure 9a) Black error bars represent 95% credible intervals. b) The BiCEP method applied to this site. Green circles and error bars represent the 95% credible interval, blue lines represent draws from the posterior distribution. The $\vec{k}=0$ axis is plotted as a black vertical line, and the value of B_{exp} is plotted as a red horizontal line. For an accurate estimate, the blue lines should cross the point where these two lines intersect.



Figure S4. The same methodology described in the caption for Figure S3 applied to site BBQ

X - 5



:

Figure S5. The same methodology described in the caption for Figure S3 applied to site BR06



Figure S6. The same methodology described in the caption for Figure S3 applied to site ET1



Figure S7. The same methodology described in the caption for Figure S3 applied to site ET2



Figure S8. The same methodology described in the caption for Figure S3 applied to site ET3



Figure S9. The same methodology described in the caption for Figure S3 applied to site FreshTRM



Figure S10. The same methodology described in the caption for Figure S3 applied to site Hawaii 1960 Flow

:



Figure S11. The same methodology described in the caption for Figure S3 applied to site hw108



Figure S12. The same methodology described in the caption for Figure S3 applied to site hw123



Figure S13. The same methodology described in the caption for Figure S3 applied to site hw126



Figure S14. The same methodology described in the caption for Figure S3 applied to site hw128



Figure S15. The same methodology described in the caption for Figure S3 applied to site hw201



Figure S16. The same methodology described in the caption for Figure S3 applied to site hw226



Figure S17. The same methodology described in the caption for Figure S3 applied to site hw241



Figure S18. The same methodology described in the caption for Figure S3 applied to site kf



Figure S19. The same methodology described in the caption for Figure S3 applied to site LV



Figure S20. The same methodology described in the caption for Figure S3 applied to site MSH



Figure S21. The same methodology described in the caption for Figure S3 applied to site P



Figure S22. The same methodology described in the caption for Figure S3 applied to site remag-rs61



Figure S23. The same methodology described in the caption for Figure S3 applied to site remag-rs62



Figure S24. The same methodology described in the caption for Figure S3 applied to site remag-rs63



Figure S25. The same methodology described in the caption for Figure S3 applied to site remag-rs78



Figure S26. The same methodology described in the caption for Figure S3 applied to site rs25



Figure S27. The same methodology described in the caption for Figure S3 applied to site rs26



Figure S28. The same methodology described in the caption for Figure S3 applied to site rs27



Figure S29. The same methodology described in the caption for Figure S3 applied to site SW



Figure S30. The same methodology described in the caption for Figure S3 applied to site Synthetic60



Figure S31. The same methodology described in the caption for Figure S3 applied to site TS



:

Figure S32. The same methodology described in the caption for Figure S3 applied to site VM