Sea surface height anomalies of the Arctic Ocean from ICESat-2: a first examination and comparisons with CryoSat-2

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November 22, 2022

Abstract

Accurately resolving spatio-temporal variations in sea surface height across the polar oceans is key to improving our understanding of ocean circulation variability and change. Here, we examine the first two years (2018-2020) of Arctic Ocean sea surface height anomalies (SSHA) from the photon-counting laser altimeter onboard NASA's ICE, Cloud, and Land Elevation Satellite-2 (ICESat-2). ICESat-2 SSHA estimates are compared to independent estimates from the CryoSat-2 mission, including available semi-synchronous along-track measurements from the recent CRYO2ICE orbit alignment campaign. There are documented residual centimeter-scale range biases between the ICESat-2 beams (in the current data release, r003) and we opted for a single-beam approach in our comparisons. We find good agreements in the along-track estimates (correlations > 0.8 and differences < 0.03 m) as well as in the gridded monthly SSHA estimates (correlation 0.76 and mean difference 0.01 m) from the two altimeters, suggesting ICESat-2 adds to the SSHA estimates from CryoSat-2.

Sea surface height anomalies of the Arctic Ocean from ICESat-2: a first examination and comparisons with CryoSat-2

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Key Points:

- We present the first (multi-year) examination of Arctic Ocean sea surface height anomalies (SSHA) from the ICESat-2 laser altimeter.
- ICESat-2 SSHA estimates compare well with near-coincident (*CRYO2ICE*) radar altimetry-derived SSHA estimates from CryoSat-2.
- ICESat-2 and CryoSat-2 show good agreement in the seasonal variability in SSHA suggesting ICESat-2 adds to the time-series of Arctic SSHA.

Abstract.

Accurately resolving spatio-temporal variations in sea surface height across the polar oceans is key to improving our understanding of ocean circulation variability and change. Here, we examine the first two years (2018-2020) of Arctic Ocean sea surface height anomalies (SSHA) from the photon-counting laser altimeter onboard NASA's ICE, Cloud, and Land Elevation Satellite-2 (ICESat-2). ICESat-2 SSHA estimates are compared to independent estimates from the CryoSat-2 mission, including available semi-synchronous along-track measurements from the recent *CRYO2ICE* orbit alignment campaign. There are documented residual centimeter-scale range biases between the ICESat-2 beams (in the current data release, r003) and we opted for a single-beam approach in our comparisons. We find good agreements in the along-track estimates (correlations > 0.8 and differences < 0.03 m) as well as in the gridded monthly SSHA estimates (correlation 0.76 and mean difference 0.01 m) from the two altimeters, suggesting ICESat-2 adds to the SSHA estimates from CryoSat-2.

Plain Language Summary

The polar oceans, with warming and dramatic declines in sea ice coverage, are experiencing some of the most rapid environmental changes on Earth. These changes have direct impacts on ocean circulation and freshwater distribution, with observable changes in sea surface height. Measuring and monitoring basin-scale variability of sea level of the ice-covered oceans has proven challenging because the surface of these oceans is only exposed within narrow openings in the sea ice, requiring high spatial resolution and bespoke measurement techniques. This study takes a first look at new high-resolution laser altimetry measurements of sea level over the Arctic Ocean collected by NASA's ICESat-2 satellite since its launch in 2018. We compare the results with those obtained using independent data from the CryoSat-2 satellite radar altimeter. By looking at near-synchronous data from when the orbit of the two satellites coincide over the Arctic Ocean, and by comparing sea surface height maps from both sensors during the two years of overlap (2018-2020) between the two missions, we find good agreement between the sea surface height estimates, providing additional confidence that ICESat-2 can be used to infer regional and seasonal polar sea surface height variability.

1 Introduction

Satellite observations of the Arctic Ocean have shown significant changes in ocean circulation, fresh water storage and energy balance since at least the 1980s, (Armitage et al., 2020; Morison et al., 2012, 2021; Polyakov et al., 2017; Proshutinsky et al., 2019; Timmermans & Marshall, 2020). Routine and accurate profiling of the sea surface height (SSH) in the Arctic is needed to continue these crucial time-series and provide more detailed insights into these changes. While we can reliably monitor the sea surface height of the open oceans at low-to-mid latitudes using satellite altimetry data (IPCC, 2019), continuous and widespread measurements at high-latitude ice-covered seas have remained limited. The main challenges are the reduced coverage due to the low inclination orbit of most satellite altimeters, sea surface sampling limited to narrow openings in the sea ice cover, and the need to accurately discriminate between sea ice and ocean surface altimetry returns.

Measurements of Arctic SSH from satellite altimetry started with low resolution radar data collected by the European Space Agency's (ESA) ERS and Envisat radar missions (1995–2010; Giles et al., 2012; Peacock & Laxon, 2004). However, the orbit inclination of these satellites limited measurements to 81.5° latitude. NASA's ICESat satellite, which operated between 2003 and 2009 (Zwally et al., 2002), offered higher resolution lidar data that improved lead classification and SSH estimates (Kwok & Morison, 2011) while its orbit inclination resulted in more extensive coverage of the Arctic Ocean. Since 2010, ESA's CryoSat-2 satellite has been acquiring unfocussed synthetic aperture radar (SAR) altimetry data over the polar regions. CryoSat-2's high orbit inclination and continuous data collection have enabled basin-scale mapping of seasonal and interannual SSH variability up to 88° latitude (Wingham et al., 2006). The SSH data from CryoSat-2 have been compared with Arctic tide gauge measurements and ocean mass variations (e.g., GRACE) and basin-scale, monthly, estimates of dynamic ocean topography (DOT; Armitage et al., 2016, 2018; Kwok & Morison, 2016) have been produced.

In September 2018, NASA launched the Ice, Cloud, and Land Elevation Satellite-2 (ICESat-2) laser altimetry mission, which has since been providing year-round profiling of the Earth's surface up to 88° latitude (Neumann et al., 2019). The novel photon-counting Advanced Topographic Laser Altimeter (ATLAS) on ICESat-2 provides high-resolution surface height measurements across its six-beam configuration. For the polar oceans, the data collected by ICESat-2 are currently being used to produce routine estimates of sea ice height, type (e.g., lead/ice), and freeboard (Kwok et al., 2020). The ICESat-2 processing algorithms utilize specular returns to discriminate open-water leads from sea ice, and the laser's spatial resolution (~11 m diameter footprint; Magruder et al., 2020) is significantly higher than that of CryoSat-2 (380 m along-track and 1650 m across-track pulse limited footprint; Scagliola, 2013). Also, contamination by off-nadir specular returns from up to 15 km across-track can potentially bias CryoSat-2 surface height retrievals (Armitage & Davidson, 2014). On the other hand, laser altimetry measurements are often hindered by the presence of clouds, which are otherwise penetrated by radar. Measurements of sea ice height and freeboard by ICESat-2 have been validated against coincident laser profiles collected during targeted underflights by NASA's

Operation IceBridge (OIB) airborne mission (Kwok et al., 2019) and the sea ice classification algorithm has been shown to agree well with coincident imagery (R. Kwok et al., 2021; Petty et al., 2021). At the time of writing, sea surface height measurements have yet to be compared against independent height data.

As of August 2020, the orbit of CryoSat-2 has been modified as part of the *CRYO2ICE* campaign, such that every 19 orbits (20 orbits for ICESat-2) the two satellites are aligned for hundreds of kilometers over the Arctic Ocean, acquiring data along near-coincident ground tracks with a minimum time difference of approximately three hours. In this study, we present a first comparison of semi-synchronous along-track SSHA retrievals from ICESat-2 and CryoSat-2 from several *CRYO2ICE* profiles. We examine SSHA from individual ICESat-2 beams and assess inter-beam range biases. We produce gridded SSHA composite maps of the Arctic Ocean and examine the relative agreement of the monthly, seasonal, and multi-year SSHA from the two altimeters. Daily/monthly gridded SSHA measurements over both polar oceans are planned to be released as an official ICESat-2 data product (ATL21) in 2021, and this study offers an examination of this type of composite SSHA data over the Arctic.

2 Data and Methods

2.1 ICESat-2 data

The ICESat-2 photon-counting laser altimeter transmits laser pulses split into a six-beam configuration of three beam pairs (each having a strong and a weak beam), where beam numbers 1, 3, and 5 identify the strong beams, and 2, 4, and 6 the weak beams (Neumann et al., 2019). The 10 kHz pulse repetition rate leads to a 0.7 m along-track separation between subsequent laser pulses of the ~11 m lidar footprint (Magruder et al., 2020). Among the ICESat-2 data products, the Level 3A sea ice products ATL07 (sea ice height and type, https://nsidc.org/data/ATL07) and ATL10 (freeboard, https://nsidc.org/data/ATL10) provide along-track measurements for six individual ground tracks (targeted at reference ground tracks, RGTs), and up to 16 satellite passes per day over both the Arctic and the Southern Ocean. The along-track surface heights are generated by aggregating 150 geolocated signal photon heights from the primary science Level 2A ATL03 data product (Neumann et al., 2019). ATL10 data coverage is limited to areas that have an ice concentration > 50% (15% for ATL07), as inferred from passive microwave satellite measurements, and up to 25 km distance from land. A full description of the ATL07/10 products can be found in the Algorithm Theoretical Basis Document (ATBD, Kwok et al., 2020) and recent changes to the algorithm are further discussed in (Kwok et al., 2021). In this study we use release 003 (r003) ATL10 data.

In ATL10, the SSHA represents the measured sea surface elevation relative to a multi-year mean sea surface (MSS, see Section 2.3) after various geophysical and atmospheric corrections have been applied (see Table S1). Note that we adjust the solid earth tide correction included in each ICESat-2 segment's SSHA from r003 ATL10 data to correct a discrepancy in the permanent tide system. The adjustment is described in the supporting information (Text S1). The SSHA is provided for each beam at three different length-scales: (1) the *height_segment_height* variable where *ssh_flag* = 1 or 2 (height segments classified as sea surface after radiometric classification as specular returns and height filtering), provides SSHA measurements calculated from the Gaussian fit to the height distribution of 150 photons within a segment (~7 m mean along-track SSH segment length for strong beams); (2) *lead_height*, expresses the weighted mean height from consecutive segments forming an individual lead; (3) *beam_refsurf_height*, represents the SSHA for a ~10-km along-track section, calculated as the weighted mean of all leads within a given section for each beam, or linearly interpolated from two adjacent sections, and smoothed using a 3-point point smoother. In subsequent analyses we use (1) but note that ATL21 data products will be formed using (3) to be consistent with the reference sea surface heights used to calculate freeboards (ATL10 and ATL20). This choice does not introduce significant differences in the gridded SSHA estimates (not shown) but allows us to take advantage of higher spatial resolution and of non-interpolated data when comparing results with CryoSat-2.

2.2 CryoSat-2 data

We use data acquired by the SIRAL K_u band SAR altimeter in the SAR mode, one of CryoSat-2's three modes of operation. We use intermediate Level 2 (L2) ice products processed at Baseline-D (Meloni et al., 2020) and available from ESA's CryoSat-2 Science Server (https://science-pds.cryosat.esa.int/). L2 data provide geolocated heigh measurements above the reference ellipsoid (WGS84) computed from each echo at intervals of approximately 300 meters. The data are already corrected for instrument effects, propagation delays, measurement geometry, and other geophysical effects (e.g., atmospheric delays and tides, see Table S1). Waveform retracking is also already applied in L2 data and determined using a model-fitting method to specular lead waveforms described by Giles et al. (2007). Further details and information can be found in the CryoSat-2 Baseline D Product Handbook (ESA, 2019) and in Meloni et al. (2020). Data coverage is controlled by the operational geographical mode mask for SAR data (https://earth.esa.int/web/guest/-/geographical-mode-mask-7107) and updated weekly to account for changes in sea-ice extent.

2.3 Mean sea surface (MSS)

To consistently compute the SSHA for CryoSat-2 we remove a mean sea surface height from each ellipsoidal elevation from L2 data (*height_sea_ice_lead_20_ku*, which includes all instrumental and geophysical corrections) by bilinearly interpolating MSS values from a 2.5 km grid (Kwok et al., 2020 – https://zenodo.org/record/4294048) to the interval centroids. The MSS grid and the interpolation approach are the same as those used in the ICESat-2 sea ice data products. The MSS includes the geoid component and is in the mean-tide system.

2.4 SSHA data binning and gridding

In along-track comparisons for the *CRYO2ICE* campaign (Figure 1, Section 3.1), we first identify measurement overlaps by selecting ICESat-2 SSHA segments from a given beam that fall within the theoretical CryoSat-2 pulse-limited across-track footprint ($\pm$825 m across-track from the centroid of each footprint; Scagliola, 2013). We then bin individual SSHA segments for ICESat-2 and SSHA intervals for CryoSat-2 in coincident 10-km sections and calculate the simple mean value from all measurements within each bin (shown as stars in Figure 1). For each profile we calculate the mean (μ) and standard deviation (*SD*) of the differences from all bins, and the correlation coefficient (*R)* between the two datasets.

To generate composite maps of the Arctic Ocean SSHA, along-track data from ICESat-2 and CryoSat-2 are first reprojected from the WGS 84 (EPSG:4326) to the NSIDC Sea Ice Polar Stereographic North coordinate system (EPSG:3411). The SSHA data are then gridded to the 25-km SSM/I polar stereographic grid by calculating the mean value within each grid cell for all data acquired within a given time period. Finally, we apply to both datasets a mask based on the NSIDC Arctic regional mask, in order to limit our assessment to the Beaufort, Chukchi, East

Siberian, Laptev, Kara, Barents, and Greenland seas, and the Central Arctic (see Figure S1 and black dashed outline in maps shown in Figure 2-4).

3 Results and discussion

3.1 Along-track CRYO2ICE SSHA comparison

There have been 77 nominal orbit overlaps between ICESat-2 and CryoSat-2 since the start of the *CRYO2ICE* campaign on 4 August 2020 (ICESat-2 RGT 606) and 11 November 2020 (ICESat-2 RGT 739), the date of the last ICESat-2 r003 ATL10 dataset available at NSIDC. For some overlaps the data products are not available and for many other overlaps, data are missing/invalid (e.g., because of cloud cover in ICESat-2). From the subset of available data, we find 4 overlaps that extend for at least 400 km with >1000 valid sea surface height segments/intervals. Most overlaps over the Arctic Ocean, including those in our subset, are with ICESat-2's beam 1 (*gt1l*). Note that the first three overlaps in our subset (Figure 1a-c) occur within summer, and while there are possible benefits from a higher lead fraction and increased number of SSH segments/intervals, we recognize that the presence of melt ponds due to snow melt on sea ice may interfere with the sea surface type retrieval algorithms, especially in the mid-August data when melt ponds are thought to be more prevalent (Kwok et al., 2020; Tilling et al., 2020).

Figure 1 shows the along-track SSHA estimates for the four selected *CRYO2ICE* overlaps (12 August to 22 September 2020). Of the four examples, three (14, 15 August and 22 September, Figure 1b-d) show mean differences of 0.01 m and one (12 August, Figure 1a) of –0.03 m. The standard deviations are 0.02–0.03 m and the correlation coefficients (R) vary between 0.83 and 0.90. The relative differences between 10-km SSHA sections are shown in

Figure S2 together with differences between geophysical corrections (i.e., tides and inverted barometer). Note that applying the geophysical corrections is key when doing these comparisons, as the lack of time-coincidence can cause significant (up to 20 cm) differences (Figure S2).

The larger (> 0.20 m) SSHA excursion seen in Figure 1b and smaller but still significant short-scale variability in the other profiles may be localized geoid features (e.g., associated to deep ocean ridges) that are not represented properly in the current MSS, and unlikely to be ocean circulation features.

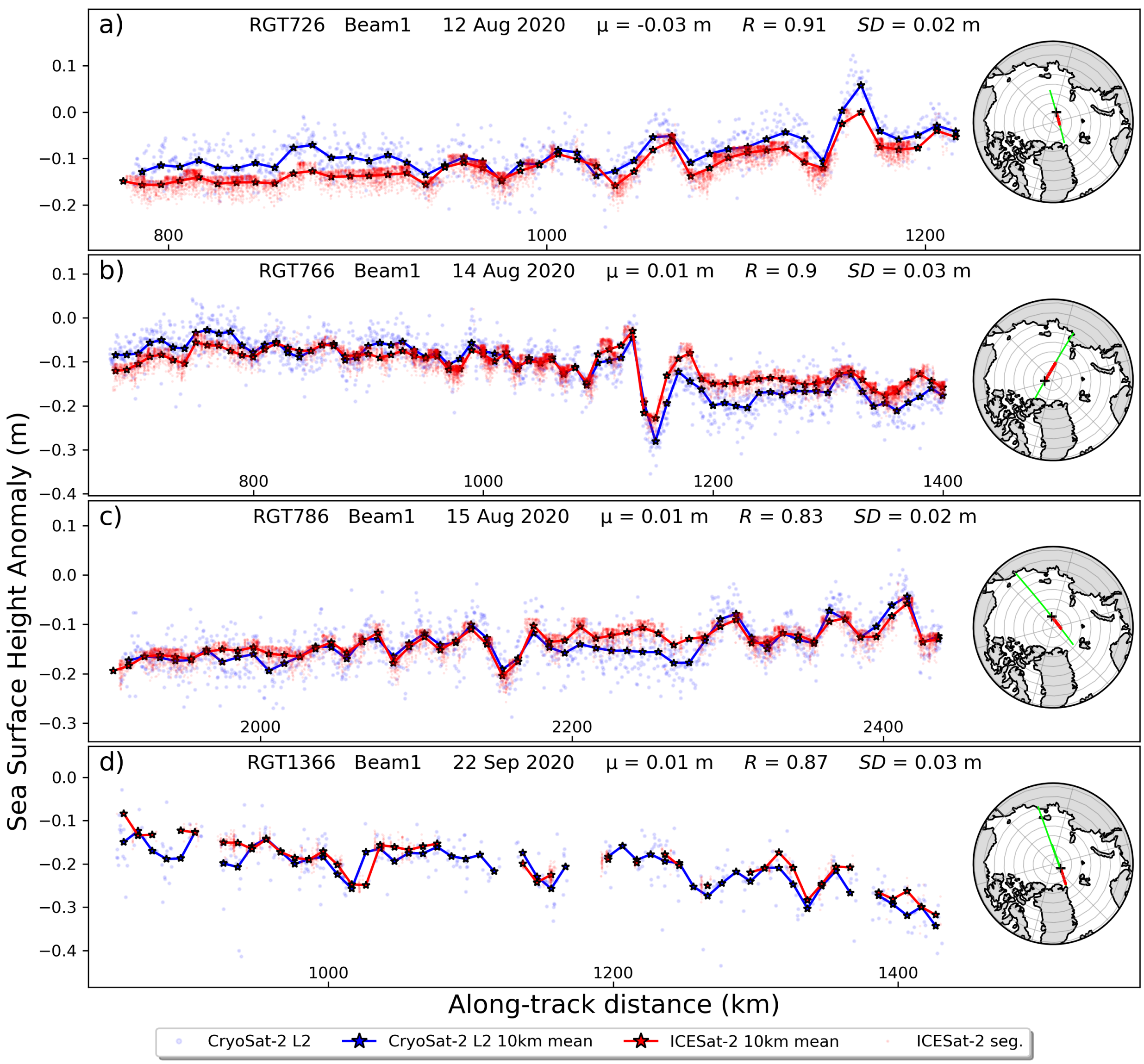

Figure 1: *CRYO2ICE* along-track SSHA comparisons. Red dots represent ICESat-2 sea surface segments, red stars show the mean value for 10-km sections. Blue dots represent CryoSat-2 sea surface intervals and blue stars the mean value for the same 10-km sections as for ICESat-2. The RGT number identifies the ICESat-2 reference ground track number. The date of acquisition of both datasets (separated by ~3 hours) is shown for each panel. For each overlap we report the mean difference (μ), the standard deviation (*SD*) of differences between the two datasets, and the correlation coefficient *R* from the least-squares regression. Map insets show the CryoSat-2 ground track in green and the extent of the overlap with ICESat-2 in red. The black + symbol marks the beginning of the profile (left side in main panels).

3.2 ICESat-2 beam comparison

Preliminary analyses by the ICESat-2 Project Science Office (PSO) have suggested that the ATLAS beams have different range biases and that these can vary through time – i.e. the height profiles from the 6 beams are not yet fully calibrated/reconciled and centimeter-level differences between beams remain. To understand the inter-beam range variability from SSHA estimates we calculate the monthly mean SSHA value over the Arctic since the start of the mission for the three strong beams independently (Figure 2a). The monthly SSHA estimate from beam 1 presents the largest differences with respect to the two other strong beams (up to ~ 0.07 m in July 2019) while differences between beam 3 and beam 5 are consistently ≤ 0.02 m. Correlation coefficients are 0.76, 0.66, and 0.93 for beam 1 – beam 3, beam 1 – beam 5, and beam 3 – beam 5, respectively. In Figure 2b-d we show the spatial distribution of the beam-to-beam differences for a given month (January 2019, gray bar in Figure 2a), which show that differences exhibit no obvious spatial correlation. This remains valid for all months since the start of the mission. The same beam-to-beam differences are also shown as histograms in Figure 2e-g, further demonstrating the clear inter-beam bias associated with beam 1 (mean of -0.03 m when compared to beam 3 and 5) and that differences between beam 3 and 5 are normally distributed around a mean of 0.00 m with a standard deviation of 0.05 m. The significant larger differences with beam 1 are also consistent with the findings of Brunt et al. (2021) estimated over the interior ice sheets of Antarctica (beam 1-3: 0.039 m; 1-5: 0.036 m; 3-5: 0.003 m), suggesting that these are sensor- or pointing solution-related.

For all of our subsequent analyses (Section 3.3 and 3.4), and until range differences between beams are fully characterized, we opt to use just a single strong beam when estimating Arctic SSHA. Based on the results presented above we select the middle strong beam (beam 3) since, despite its lower transmitted energy level (~80% of beam 1 and 5), the steeper incidence angle results in a stronger backscatter in the presence of highly reflective surfaces (e.g., leads) consistently increasing the number of specular lead returns compared to other strong beams (Kwok et al., 2021). This is currently our recommended strategy for the initial production and release of ICESat-2 ATL21 data.

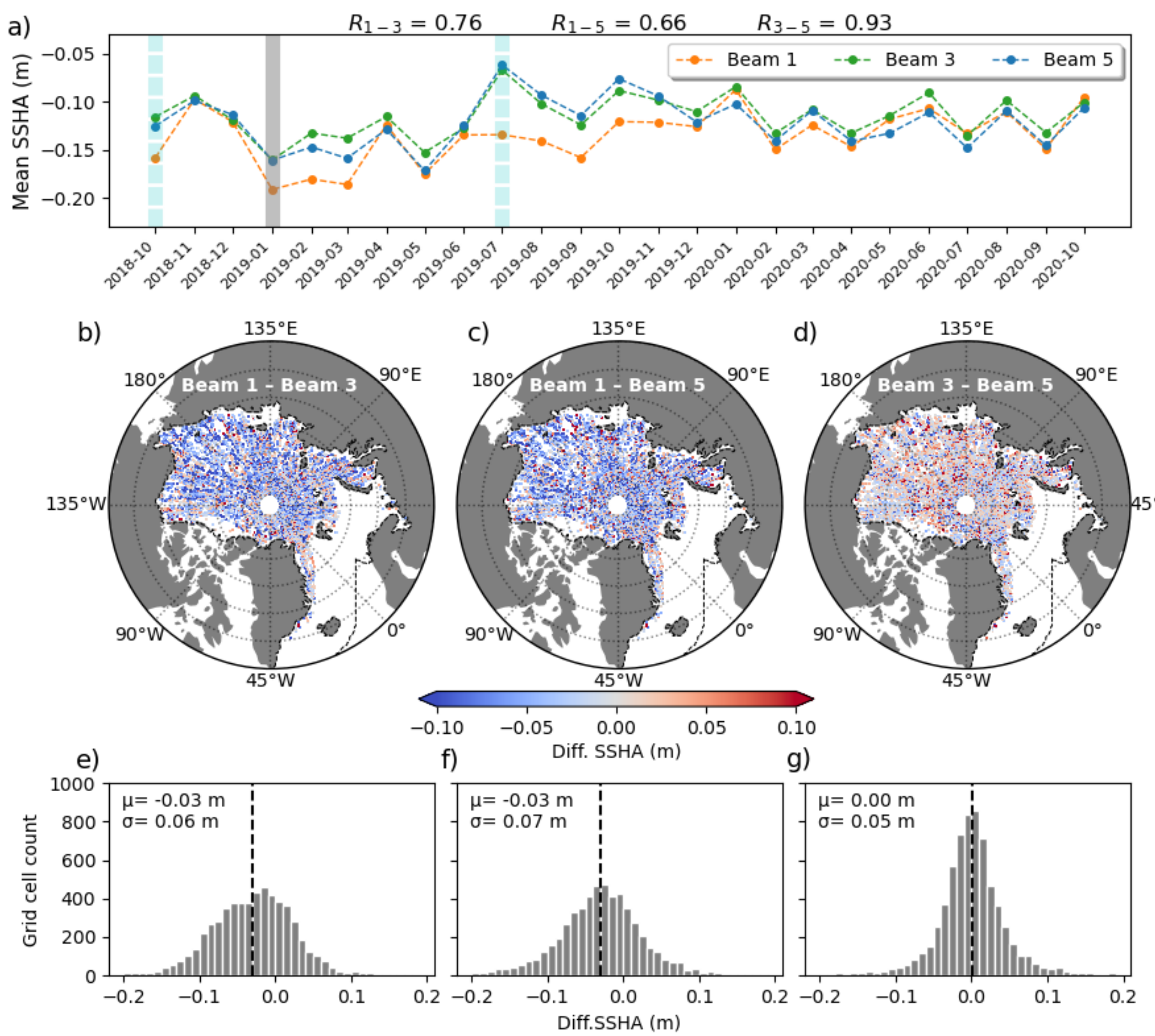

Figure 2: ICESat-2 beam comparison. **a)** Monthly mean for the Arctic Ocean calculated using data from each beam. The cyan dashed bars mark months for which data do not cover the entire month, October 2018–beginning of science data acquisition on 14 October– and July 2019–data between 1 and 8 July are not available due to satellite safe mode operations. The gray bar marks the month for which data are shown in panel b-g. Correlation coefficients (R) between beams are shown at the top. **b-d)** Maps showing the differences between beams for the month of January 2019. The black dashed line marks the extent of the area of interest. **e-g)** histograms showing the distribution of the differences presented in panels b-d. The black dash lines mark the mean (μ) and σ is the standard deviation.

3.3 Monthly and multi-year SSHA comparison

In Figure 3a we compare monthly SSHA means calculated using ICESat-2 beam 3 to those calculated using CryoSat-2 Level-2 data. We limit this comparison to the Central Arctic, the area outlined by the green dashed line in Figure 3b, where we expect consistent year-round ice cover and to exclude effects introduced by season-dependent changes in sea-ice extent and different data coverage near the coastal regions. Further details for each monthly comparison (mean, number of valid grid cells, number of data points) are provided in Table S2. A decrease in mean SSHA is shown by both sensors during fall-winter months and is followed by a mean SSHA increase during spring-summer months. Differences across all months between the two sensors have a mean of 0.01 m (SD = 0.02 m), and the correlation coefficient from a least-squares regression (*R)* is 0.76 (slope = 0.95, intercept = – 0.02 m). We find that up to 0.03 m of the observed monthly SSHA differences, especially during fall/winter, are caused by differences in the inverted barometer correction applied to each dataset. Our comparisons between heights from ICESat-2 with those from CryoSat-2 show a better agreement than has been shown by Brunt et al. (2021), who compared absolute ice height over the flat interiors of the Antarctic ice sheet and found differences > 0.3 m. This larger discrepancy, however, is likely due to the much greater penetration depth of the Ku band radar in firn compared to sea water.

We then compare the Arctic SSHA calculated from data spanning the two-year mission overlap, from November 2018 through October 2020. The ICESat-2 mean 2018-2020 SSHA in shown in Figure 3b and that from CryoSat-2 is presented in Figure 3c. Both maps show a positive SSHA in the southern Beaufort Sea, a strong negative anomaly in the Chukchi/Siberian seas and a weaker negative SSHA in Central Western Arctic, a spatial pattern consistent with recent positive phase in the Arctic Oscillation (Armitage et al., 2018; Morison et al., 2021). In Figure 3d we show a histogram of the differences between ICESat-2 and CryoSat-2 SSHA, while a map of the SSHA differences is presented in Figure 3e, which shows the ICESat-2 SSHA to be generally higher in the more marginal seas (Barents, Kara, East Siberian, and Chukchi) and slightly lower in the Central Arctic. The marginal seas are areas of large SSH variability where the different acquisition times between the two satellites can capture different parts of these cycles (see Figure S3 for the standard deviation of each dataset, showing higher values in the marginal seas) and can therefore explain much of these differences. Increased data acquisition from both missions will enable a more reliable comparison of the mean SSHA from ICESat-2 and CryoSat-2.

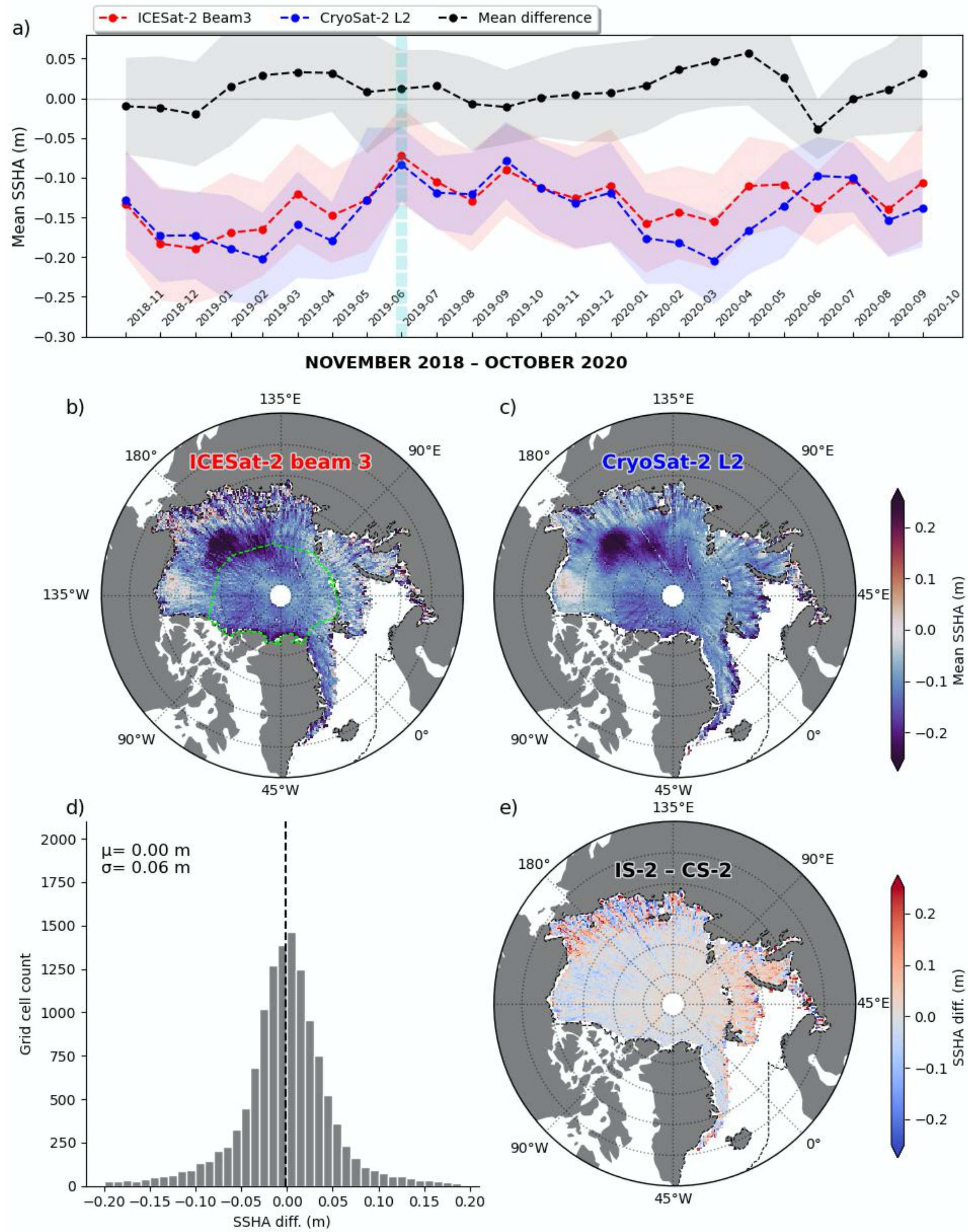

Figure 3: Comparison between ICESat-2 and CryoSat-2 Arctic SSHA. **a)** time series of monthly mean SSHA for the Central Arctic (area outlined by green dashed-line in panel b) from ICESat-2 (red) and CryoSat-2 (blue), with shaded areas representing one standard deviation from the

mean. **b-c)** Multi-year mean SSHA estimated using data acquired between November 2018 and October 2020. **d)** Histogram showing the distribution of the differences between ICESat-2 and CryoSat-2, also shown in map view in panel **e)**. In d) the black dash line marks the mean (μ) and σ is the standard deviation. The black dashed line in e) marks the extent of the area of interest (data outside this line are masked out).

3.4 Seasonal SSHA variations from ICESat-2

In Figure 4 we present seasonal maps of Arctic SSHA for three-month periods starting in October 2018 and ending in September 2020. The top row (Figure 4a-d) can be directly compared to the bottom row (Figure 4e-h) to assess year-to-year differences, while from left to right we track the temporal progression during two entire freezing-melting seasons (2018–2019 and 2019–2020). Note that variations in spatial coverage are dictated by variations in sea ice extent since ICESat-2 ATL10 data are only provided for areas that have an ice concentration > 50%. Comparisons to CryoSat-2 for each three-month period are presented in Figure S4, and confirm similar SSHA spatio-temporal variations providing some confidence in the capability of ICESat-2 to produce consistent estimates of Arctic SSHA.

A positive SSHA centered on the Beaufort Sea (a strengthened Beaufort Gyre) is clearly visible during winter months but less apparent in 2020 (see Figure 4 c-d compared to Figure 4 g-h). Large variability in the Siberian and Chukchi seas also corresponds to areas characterized by high short-term SSH variability.

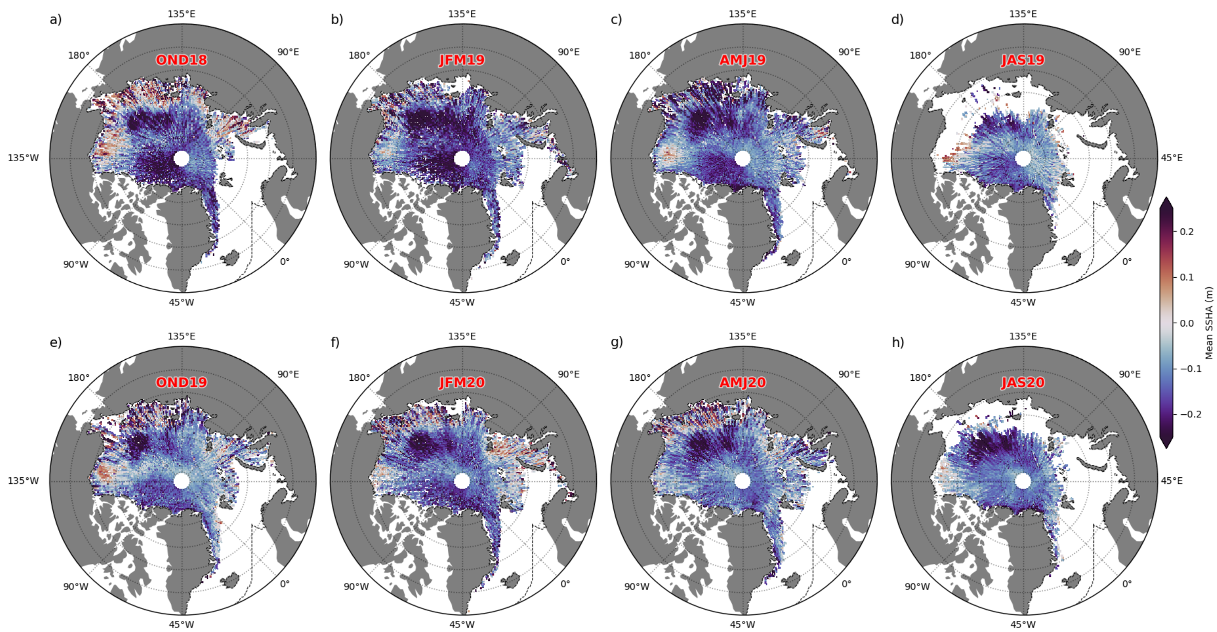

Figure 4: Seasonal mean SSHA maps from ICESat-2. OND = October, November, December; JFM = January, February, March; AMJ = April, May, June; JAS = July, August, September. The black dashed line marks the extent of the area of interest (data outside this line are masked out).

5 Summary and conclusions

Here we have presented a first examination of Arctic sea surface height anomalies (SSHA) from NASA's ICESat-2 laser altimeter during the first two years of the mission (2018-2020). We analyzed beam-to-beam differences and provided an independent assessment of inter-beam range biases for the ATLAS altimeter. We compared the ICESat-2 SSHA estimates with L2 sea ice data obtained from ESA's CryoSat-2 radar altimeter. We provided a brief description of the necessary steps to reconcile the SSHA data from the two altimetry missions by imposing the same permanent tide system, MSS, and geophysical corrections. A careful reconciliation of the data is needed in future efforts to blend data from ICESat-2 with those from CryoSat-2 (and potentially other airborne and space-borne altimetry missions).

The strong agreement between both the semi-synchronous along-track estimates from the *CRYO2ICE* overlaps and basin-scale gridded SSHA estimates between the two sensors suggests that the higher resolution ICESat-2 data can be used to estimate monthly/seasonal SSHA and perhaps resolve 10 km-scale spatial variability in SSHA. The multi-year record of overlap also opens up the potential to produce a new, high-resolution, blended, estimate of the mean sea surface of the Arctic Ocean (and indeed Southern Ocean) which could better resolve what we believe to be anomalously large SSHA spatial deviations shown in the *CRYO2ICE* overlaps. Finally, our results provide a first evaluation of the approach used for the production of ICESat-2 SSHA gridded data products for the polar oceans (ATL21). Future work will extend this analysis to the Southern Ocean, pending *CRYO2ICE* orbit maneuvers for the Southern Hemisphere.

Acknowledgments, Samples, and Data

ICESat-2 ATL10 data products were obtained from NSIDC and are available at http://nsidc.org/data/atl10. CryoSat-2 Level-2 data (SIR_SAR_L2) were obtained from ESA at https://science-pds.cryosat.esa.int/#. The mean sea surface grid is available at https://zenodo.org/record/4294048.

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Geophysical Research Letters (GRL)

Supporting Information for

Sea Surface Height Anomalies of the Arctic Ocean From ICESat-2: A First Examination and Comparisons with CryoSat-2

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Contents of this file

Introduction

This supplement contains text, figures and tables in support of the main document.

Text S1 describes an adjustment needed to correct a discrepancy in the permanent tide system between geophysical corrections applied to ICESat-2 release 003 ATL10 data.

Figure S1 shows a map of the Arctic Ocean with the names of the main seas and the extent of the NSIDC Sea Ice Index Arctic regional mask used to define the area of interest.

Figure S2 is complementary to Figure 1 and shows the differences between SSHA estimates from ICESat-2 and CryoSat-2 for four *CRYO2ICE* overlaps. This figure also highlights the importance of geophysical corrections when comparing non-synchronous

SSHA measurements by showing the differences introduced by the ~3-hour time lag between the passes of the two satellites.

Figure S3 shows the variability (standard deviation) and number of data points (segment/interval count) associated with each dataset used to produce the multi-year (2018-2020) SSHA composite maps shown in Figure 3b – ICESat-2 – and Figure 3c – CryoSat-2.

Figure S4 shows SSHA differences between ICESat-2 and CryoSat-2 in three-month composites for the duration of the mission overlap. The figure is complementary to Figure 4 and used to assess the ability of ICESat-2 in tracking seasonal and inter-annual changes in SSHA.

Table S1 lists all the geophysical corrections that are applied to obtain SSHA estimates from ICESat-2 and CryoSat-2 and provides information on the tidal and atmospheric models used for each sensor.

Table S2 provides the statistics for monthly SSHA estimates for ICESat-2 and CryoSat-2 used to generate the time-series plot presented in Figure 3a. The table also provides the count of the total number of segments/intervals used to generate each monthly estimate, and the number of valid grid cell in each monthly SSHA composite.

Text S1: Permanent tide adjustment for ICESat-2

The ICESat-2 release 003 (r003) ATL03 photon heights used for the production of r003 ATL07/10 data products are corrected for the solid earth tide (SET) and include both the time-dependent (periodic) and the time-independent (permanent) components, so that ellipsoidal heights are in a tide-free system. However, the mean sea surface (MSS) grid used to estimate sea surface height anomalies (SSHA) in ATL07/10 (see Section 2.3) is in the mean-tide system (i.e., it includes the distortion of the geoid due to the permanent tide). While this inconsistency will be rectified in future ICESat-2 data releases (starting from r004 all data will be in a tide-free system), in this study we adjust the SET correction for ICESat-2 by reintroducing the time-independent component as follows:

$$SSHA_{mean-tide} = SSHA_{r003} + 0.060292 - 0.180873\, sin^2\varphi$$

where the degree-2 Love number, *h2*=0.609 is implicit in the equation (IERS2010 Conventions) and φ is the latitude in radians.

The approach described above makes sure that SSHA estimates from both ICESat-2 and CryoSat-2 are in the same permanent tide system (mean tide system), enabling SSHA comparative analyses.

Figure S1. Map of the Arctic Ocean showing the extent of the area of interest in blue.

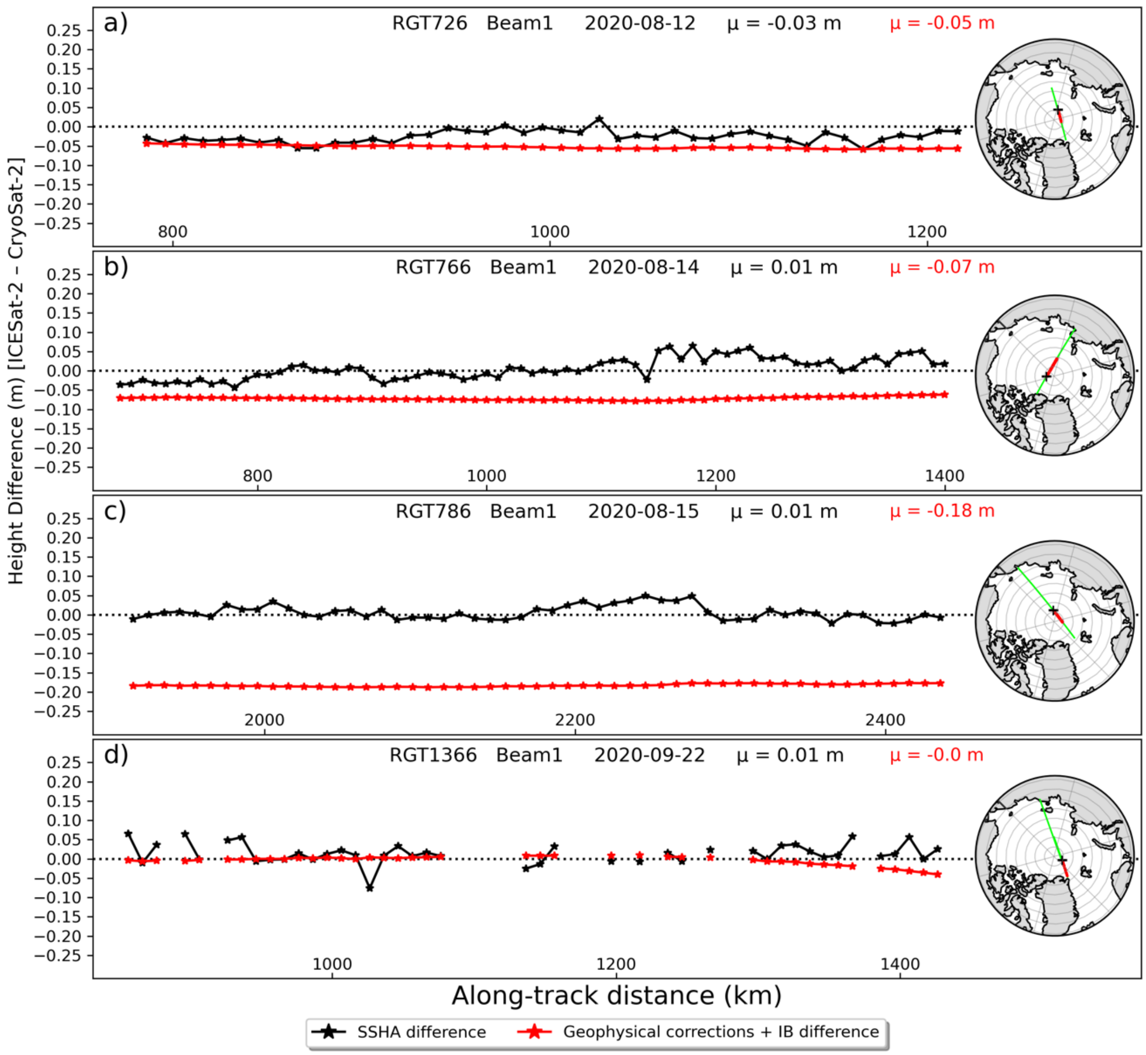

Figure S2: *CRYO2ICE* differences between ICESat-2 and CryoSat-2. Black stars show the SSHA difference (ICESat-2 – CryoSat-2) calculated for 10-km sections presented in Figure 1. Red stars show differences in geophysical corrections (ocean tide, ocean long period tide, pole tide, solid earth tide, ocean loading, and inverted barometer) for the same 10-km segments. These differences are mainly due to the time difference between acquisitions (~3 hours), and to a minor extent to the different tidal/IB models, and do not include other sensor-specific corrections for instrumental and atmospheric effects (e.g., wet and dry troposphere) that are applied to SSHA estimates. The RGT number identifies the ICESat-2 reference ground track number. For each overlap we report the SSHA (in black) and geophysical corrections (in red) mean difference ($\boldsymbol{\mu}$). Map insets show the CryoSat-2 ground track in green and the extent of the overlap with ICESat-2 in red. The black + symbol marks the beginning of the overlap (left side of main panel).

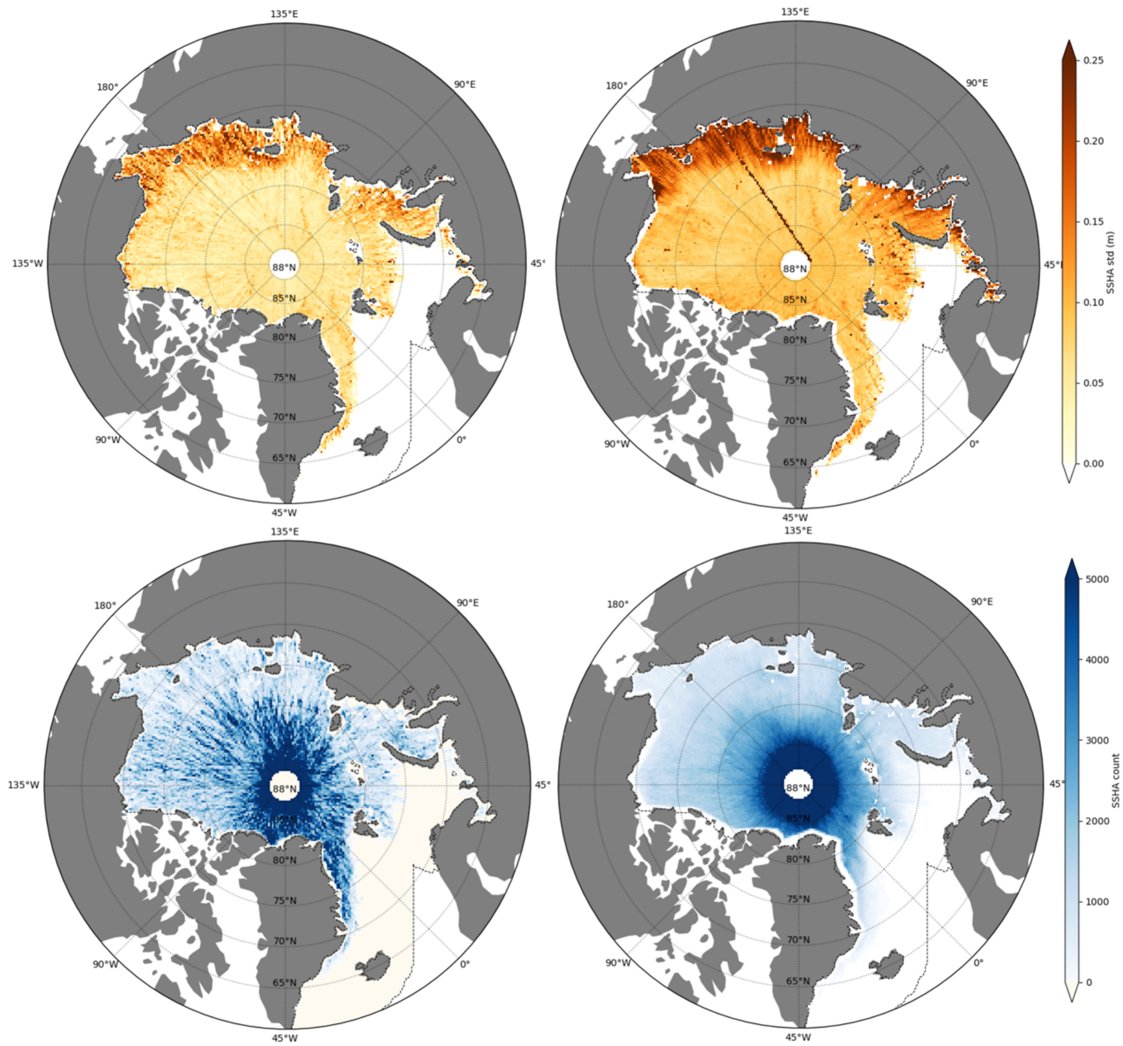

Figure S3. Maps showing (top) the standard deviation of SSHA measurements and (bottom) the number of data segments/intervals in each grid cell for the entire mission overlap (November 2018 – October 2020). (left) ICESat-2 and (right) CryoSat-2.

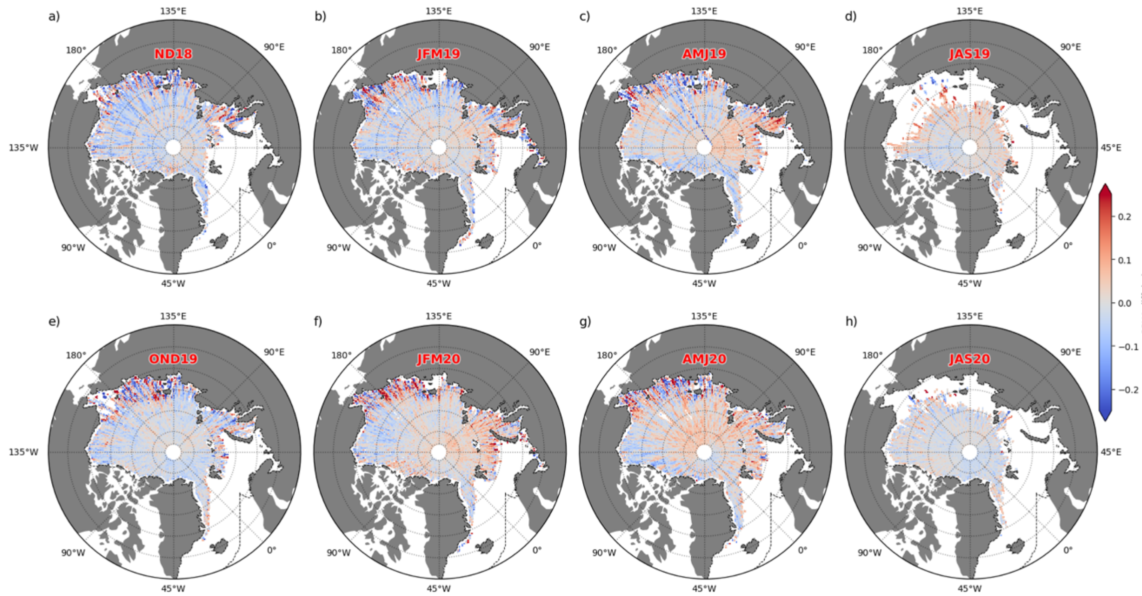

Figure S4: Seasonal mean SSHA difference maps (ICESat-2 – CryoSat-2). OND = October, November, December; JFM = January, February, March; AMJ = April, May, June; JAS = July, August, September. The black dashed line marks the extent of the area of interest (data outside this line are masked out).

Correction	ICESat-2		CryoSat-2	
Solid Earth Tide	IERS 2010 conventions	Y	Cartwright model	Y
Ocean Loading	GOT 4.8 ocean tide model	Y	FES 2004 model	Y
Ocean Tides	GOT 4.8 ocean tide model	Y	FES 2004 model	Y
Long Period Equilib. Tide	GOT 4.8 ocean tide model	Y	FES 2004 model	Y
Solid Earth Pole Tide	IERS 2010 conventions	Y	SSALTO	Y
Ocean Pole Tide	N/A	N		?
Inverted Barometer	From sea level pressure (ATL09)	Y	CNES SSALTO (ECMWF)	Y
Total column atm. delay	Luthcke & Petrov, ATBD ATL03a	Y	N/A	N
Dry troposphere	N/A	N	CNES SSALTO (ECMWF)	Y
Wet troposphere	N/A	N	CNES SSALTO (ECMWF)	Y
Ionosphere	N/A	N	GIM/Bent model	Y

Table S1. Geophysical and atmospheric corrections applied to each dataset. Y= yes, N=no, ?=unknown.

DATE	IS-2 SSHA MEAN	CS-2 SSHA MEAN	IS-2 SSHA GRID COUNT	CS-2 SSHA GRID COUNT	ICESAT-2 N. SEGMENTS	CRYOSAT-2 N. INTERVALS	MEAN (SD) OF **DIFFERENCES IS-2 – CS-2**
201811	-0.133	-0.128	3,084	4,291	651,527	870,650	-0.010 (0.061)
201812	-0.183	-0.173	3,218	4,356	484,112	883,514	-0.012 (0.065)
201901	-0.189	-0.173	3,170	4,402	546,014	808,559	-0.020 (0.066)
201902	-0.169	-0.190	2,756	4,408	520,729	686,661	0.015 (0.075)
201903	-0.165	-0.202	2,581	4,370	678,093	672,523	0.029 (0.054)
201904	-0.120	-0.159	3,131	4,312	964,540	707,937	0.033 (0.061)
201905	-0.148	-0.180	3,891	4,433	819,505	1,111,238	0.032 (0.049)
201906	-0.128	-0.128	3,748	4,421	1,476,440	2,352,078	0.008 (0.077)
201907	-0.072	-0.083	3,226	4,404	2,079,310	3,131,588	0.012 (0.049)
201908	-0.105	-0.119	3,859	4,403	2,830,539	2,607,746	0.016 (0.045)
201909	-0.129	-0.121	3,932	4,290	1,180,395	1,280,231	-0.007 (0.059)
201910	-0.090	-0.078	3,767	4,397	859,163	797,832	-0.011 (0.047)
201911	-0.113	-0.113	3,600	4,415	733,714	944,903	0.001 (0.050)
201912	-0.126	-0.132	3,798	4,416	646,454	959,979	0.005 (0.059)
202001	-0.110	-0.119	3,374	4,395	648,188	819,654	0.007 (0.062)
202002	-0.158	-0.176	2,998	4,377	573,857	680,526	0.016 (0.058)
202003	-0.144	-0.182	3,192	4,397	620,692	748,099	0.036 (0.056)
202004	-0.156	-0.204	3,583	4,388	982,488	897,103	0.047 (0.058)
202005	-0.110	-0.167	3,726	4,399	955,037	1,363,667	0.057 (0.066)
202006	-0.109	-0.135	4,071	4,443	2,200,551	2,602,271	0.026 (0.070)
202007	-0.138	-0.098	4,224	4,415	7,609,865	3,117,697	-0.039 (0.038)
202008	-0.102	-0.100	3,718	4,191	4,466,378	2,652,584	-0.001 (0.047)
202009	-0.140	-0.153	3,002	3,640	898,732	1,193,840	0.011 (0.056)
202010	-0.106	-0.137	3,411	4,130	563,092	1,940,856	0.031 (0.072)

Table S2. Monthly mean SSHA for the Central Arctic (see Figure 3a) from ICESat-2 and CryoSat-2. YYYYMM = year and month. Mean values are in meters. Grid cell counts are calculated as the total number of 25-km grid cells within the area of interest containing at least one data point.