# Sources, Occurrence and Characteristics of Fluorescent Biological Aerosol Particles Measured over the Pristine Southern Ocean

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#### Abstract

In this study we investigate the occurrence of primary biological aerosol particles (PBAP) over all sectors of the Southern Ocean (SO) based on a 90-day dataset collected during the Antarctic Circumnavigation Expedition (ACE) in austral summer 2016-2017. Super-micrometer PBAP (1 to 16  $\mu$ m diameter) were measured by a wide band integrated bioaerosol sensor (WIBS-4). Low (3 $\sigma$ ) and high (9 $\sigma$ ) fluorescence thresholds are used to obtain statistics on fluorescent and hyper-fluorescent PBAP, respectively. Our focus is on data obtained over the pristine ocean, i.e. more than 200 km away from land. The results indicate that (hyper-)fluorescent PBAP are correlated to atmospheric variables associated with sea spray aerosol (SSA) particles (wind speed, total super-micrometer aerosol number concentration, chloride and sodium concentrations). This suggests that a main source of PBAP over the SO is SSA. The median fraction of fluorescent and hyper-fluorescent PBAP to super-micrometer particles positively correlates with concentrations of bacteria and several taxa of phytoplankton measured in seawater, indicating that marine biota concentrations modulate the PBAP source flux. We investigate the fluorescent properties of (hyper-)fluorescent PBAP for several events that occurred near land masses. We find that the fluorescence signal characteristics of particles near land is much more variable than over the pristine ocean. We conclude that the source and concentration of fluorescent PBAP over the open ocean is similar across all sectors of the SO.

#### Sources, Occurrence and Characteristics of Fluorescent Biological Aerosol Particles 1

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| 23          |   |  |
| 24          | Key Points:   |  |
| 25<br>26    | • Fluorescent primary bioaerosol particles (PBAP) were measured over all sectors of the Southern Ocean  |  |
| 27<br>28    | • Moderate to good correlations were observed between PBAP and sea spray aerosol (SSA) proxies  |  |
| 29<br>30    | • PBAP fractions in SSA were positively correlated to concentrations of certain marine biota  |  |
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## 37 Abstract

In this study we investigate the occurrence of primary biological aerosol particles (PBAP) over 38 all sectors of the Southern Ocean (SO) based on a 90-day dataset collected during the Antarctic 39 40 Circumnavigation Expedition (ACE) in austral summer 2016-2017. Super-micrometer PBAP (1 to 16 µm diameter) were measured by a wide band integrated bioaerosol sensor (WIBS-4). Low 41  $(3\sigma)$  and high  $(9\sigma)$  fluorescence thresholds are used to obtain statistics on fluorescent and hyper-42 fluorescent PBAP, respectively. Our focus is on data obtained over the pristine ocean, i.e. more 43 44 than 200 km away from land. The results indicate that (hyper-)fluorescent PBAP are correlated to atmospheric variables associated with sea spray aerosol (SSA) particles (wind speed, total 45 super-micrometer aerosol number concentration, chloride and sodium concentrations). This 46 suggests that a main source of PBAP over the SO is SSA. The median fraction of fluorescent and 47 hyper-fluorescent PBAP to super-micrometer SSA is 1.6% and 0.13%, respectively. We 48 demonstrate that the fraction of (hyper-)fluorescent PBAP to total super-micrometer particles 49 positively correlates with concentrations of bacteria and several taxa of phytoplankton measured 50 in seawater, indicating that marine biota concentrations modulate the PBAP source flux. We 51 investigate the fluorescent properties of (hyper-)fluorescent PBAP for several events that 52 occurred near land masses. We find that the fluorescence signal characteristics of particles near 53 54 land is much more variable than over the pristine ocean. We conclude that the source and concentration of fluorescent PBAP over the open ocean is similar across all sectors of the SO. 55 56 57

### 67 **1 Introduction**

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69 Primary biological aerosol particles (PBAP) are ubiquitous atmospheric particles emitted from the biosphere, which encompass intact microorganisms (e.g., bacteria, viruses, pollen, 70 fungal spores etc.), or fragments of such microorganisms (Després et al., 2012; Fröhlich-71 72 Nowoisky et al., 2016). PBAP have major impacts on public health, as certain types of PBAP are known to act as allergens or spread disease (Després et al., 2012; Fröhlich-Nowoisky et al., 73 2016; Taylor et al., 2004). Furthermore, long-range transport of PBAP, such as bacteria, could 74 influence the ecosystem and biome diversity of the environments to which they are transported 75 (Burrows et al., 2009; Hervàs et al., 2009; Kellogg & Griffin, 2006). Moreover, PBAP have the 76 potential to affect cloud formation, for example by acting as giant cloud condensation nuclei 77 (Pope, 2010) at low supersaturations. A number of studies have demonstrated that PBAP are 78 effective ice nucleating particles (INP) (Després et al., 2012; Tobo et al., 2013), thereby 79 facilitating glaciation of super-cooled liquid clouds via heterogeneous ice nucleation (Kanji et 80 al., 2017). Such aerosol-cloud interactions can modify cloud optical properties and precipitation 81 82 patterns with important atmospheric impacts on regional and global scales (Kanji et al., 2017).

PBAP originate from both the terrestrial and marine biosphere (Després et al., 2012). In 83 the oceanic environment, primary aerosol particles, known as sea spray aerosol (SSA) particles, 84 are produced through a combination of processes, which includes breaking of waves, generation 85 86 of bubbles in the oceanic water, rising of bubbles to the ocean surface and the subsequent bubble bursting and aerosol ejection (de Leeuw et al., 2011; Lewis & Schwartz, 2004). Additionally, 87 larger sea spray droplets known as spume droplets can be torn directly from wave crests during 88 strong wind conditions (Monahan et al., 1986). In addition to inorganic sea salt, SSA consists of 89 90 complex arrays of organic compounds (Brooks & Thornton, 2018; Hawkins & Russell, 2010; O'Dowd & de Leeuw, 2007; Prather et al., 2013). SSA organic compounds have their origin in 91 seawater dissolved organic matter (DOM) (Hawkins & Russell, 2010), particulate organic matter 92 (POM) such as polysaccharides and proteinaceous gel-like particles (Aller et al., 2017), and 93 microorganisms such as bacteria, viruses and phytoplankton (Quinn et al., 2015). Studies on the 94 chemical composition of laboratory-generated SSA indicate that seawater bioactivity influences 95 the fraction of organic matter in SSA by altering the abundance of microorganisms in water 96 97 (Ault et al., 2013; Lee et al., 2020; Wang et al., 2015). In addition to laboratory-based studies,

analysis of aerosol samples collected in different global oceanic regions have demonstrated that

99 marine microorganisms and associated organic components are incorporated into SSA (e.g.

100 Ceburnis et al., 2016; Mayol et al., 2017; Orellana et al., 2011; Russell et al., 2010). More

recently, sequencing analysis of aerosol samples from the Southern Ocean (SO) also

102 demonstrated that bacteria were present in the SSA (Uetake et al., 2020). These studies indicate

103 that PBAP contribute to SSA-associated primary organic matter.

Previous studies indicate that some SSA particles possess ice nucleating properties (Bigg, 104 105 1973; Schnell & Vali, 1976), and it was suggested that this could be related to marine biological 106 activity. More recent studies have demonstrated that SSA containing both dissolved and/or particulate organic matter are capable of nucleating ice crystals efficiently at temperatures in the 107 range -20 to -35°C (DeMott et al., 2016; McCluskey et al., 2018; Wang et al., 2015; Wilbourn et 108 al., 2020; Wilson et al., 2015). Such SSA particles tend to nucleate ice at lower temperatures 109 110 than their terrestrial counter-parts, i.e. they are less effective INP (DeMott et al., 2010), which necessitates the segregation of terrestrial and marine INP parametrizations in global atmospheric 111 112 models (Vergara-Temprado et al., 2017). Overlooking such a distinction in INP parametrizations can increase the uncertainty in global atmospheric models. 113

114 The Southern Ocean (SO) is a pristine environment (e.g. Hamilton et al., 2014; Schmale, Baccarini, et al., 2019) as well as the roughest ocean on Earth in terms of surface winds and 115 waves (Young, 1999). This makes the SO an extremely promising location to study SSAs and 116 their associated PBAP. However, our knowledge regarding the regional distribution and 117 118 composition of SO SSA and PBAP is still very limited (Middlebrook et al., 1998; Murphy et al., 119 1998; Uetake et al., 2020). In addition, studies have indicated considerable uncertainties in calculated radiative forcing over the SO (Flato et al., 2013). These uncertainties are partly 120 attributed to misrepresentation of SO aerosol and associated processes, e.g. excessive 121 heterogeneous ice crystal formation and subsequent precipitation in global atmospheric models 122 (Vergara-Temprado et al., 2017, 2018). Considering the unique properties of marine PBAP and 123 their potential effects on cloud microphysics, identification, quantification and source 124 apportionment of these particles is an important step towards improving the representation of SO 125 aerosols in global climate models. 126

Identification and quantification of atmospheric PBAP of oceanic origin is prone to several challenges. Conventional methods rely on atmospheric sample extraction and offline analysis (Després et al., 2012; Fröhlich-Nowoisky et al., 2016). Although the analysis of offline samples can provide detailed morphological, chemical and biological information on PBAP, it remains time consuming. It limits the obtainable sample sizes through offline analysis, making it difficult to gain quantitative insights. In addition, the relatively poor time resolution of offline samples complicates source identification.

134 More recently, online PBAP detection methods based on aerosol auto-fluorescent 135 properties have become available (e.g. Fennelly et al., 2018). Online PBAP detection methods typically rely on ultra-violet light induced fluorescence (UV-LIF). These methods employ UV 136 excitation of single particles, followed by spectrally resolved or waveband integrated detection 137 of the resulting fluorescent light. The wavelength detection ranges are chosen to match regions 138 139 of fluorescence for biological compounds that are found ubiquitously in PBAP, such as tryptophan and Nicotinamide Adenine Dinucleotide (NADH) (Fennelly et al., 2018; Kaye et al., 140 141 2005). To date, online PBAP measurements have been employed in both laboratory studies (Hernandez et al., 2016; Savage et al., 2017) and field measurements (Crawford et al., 2016, 142 2017; Healy et al., 2014; Perring et al., 2015; Pöhlker et al., 2012; Toprak & Schnaiter, 2012; 143 Ziemba et al., 2016). In the context of field measurements, the key advantage of online UV-LIF 144 techniques is that they facilitate size-resolved quantitative measurements of PBAP 145 concentrations at high time resolution. This makes it possible to compare them to other highly 146 variable environmental parameters, thereby facilitating identification of PBAP sources. To the 147 best of our knowledge, only two studies have used online UV-LIF methods to investigate PBAP 148 in the Antarctic and SO regions (Crawford et al., 2017; McFarquhar et al., 2020). Crawford et al. 149 (2017) identified fluorescent particles measured in the Halley VI station along the Antarctic 150 coast as dust and/or pollen particles transported from the Antarctic and South American 151 continents, or as PBAP transported from biologically active coastal marginal ice zones. 152 However, it is not clear if these results are representative of other SO regions, particularly remote 153 oceanic regions far from continental influence. McFarquhar et al. (2020) report median PBAP 154 concentrations measured during the Measurements of Aerosols, Radiation, and Clouds over the 155 Southern Ocean research cruise (MARCUS, October 2017 – April 2018) with no additional 156

analyses of possible sources and sinks. Therefore, further online, fluorescence-based
 measurements are required to gain better insights into PBAP over the SO.

In the current study, we strive to explore the occurrence and origin of marine PBAP in 159 the pristine SO region with an extensive database of new measurements. Co-located marine and 160 atmospheric measurements were performed during the research cruise Antarctic 161 Circumnavigation Expedition (ACE) between December 2016 and March 2017 (Schmale, 162 Baccarini, et al., 2019), including online auto-fluorescence measurements of PBAP made with a 163 164 wideband integrated bioaerosol sensor (WIBS-4). This unique dataset represents one of the 165 largest sets of aerosol measurements ever collected over all sectors of the SO. Section 2 describes the details of the dataset, instrumentation and data analysis assumptions. We 166 investigate the link between PBAP and SSA in sections 3.1 and 3.2. Additionally, a 167 comprehensive set of measurements of seawater chemical composition and biological activity 168 169 were conducted during ACE. In section 3.3, we compare the variability of the seawater measurements to that of the fluorescent aerosols in order to explore the ocean-originating source 170 171 of the PBAP. Finally, we present the spatial concentration distribution, and microphysical and fluorescent properties of PBAP in Sections 3.4 to 3.8. Overall, this study provides 172 comprehensive insights into the distribution of SSA-related PBAP over the SO, and sheds light 173 on the marine biological components responsible for the observed PBAP. 174

- 175 **2 Materials and Methods**
- 176 2.1 Campaign Description
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We acquired the results presented in this study during the Antarctic Circumnavigation 178 179 Expedition (ACE) conducted from December 2016 to March 2017 (Schmale, Baccarini, et al., 2019). A detailed overview of the ACE campaign can be found in the ACE cruise report (Walton 180 & Thomas, 2018). In this campaign, we performed co-located marine and atmospheric 181 measurements aboard the research vessel Akademik Tryoshnikov. ACE covered an extensive 182 range of geographical locations (Figure 1) starting from Cape Town, South Africa, and 183 circumnavigating the SO before returning back to Cape Town. To simplify the geographical 184 extent for analyses in this work, we divided the route into three segments. Segment 1 represents 185 samples collected from January 6<sup>th</sup> 2017 to January 31<sup>st</sup> 2017, which covers the route from 186

187 Kerguelen Islands to the Mertz Glacier in Antarctica. Segment 2 represents samples collected

188 from January 31<sup>st</sup> 2017 to Feb 22<sup>nd</sup> 2017, which covers the track from the Mertz Glacier to Punta

Arenas in Chile. Segment 3 represents samples collected from Feb 22<sup>nd</sup> 2017 to March 19<sup>th</sup> 2017,

which covers the area between Punta Arenas and Cape Town. Although the campaign started in

191 Cape Town, we used only those data acquired after the Kerguelen Islands because the internal

192 pump of the instrument did not function properly at the beginning of the cruise. This pump was

replaced by an external pump during the station at Kerguelen Islands. It should be noted that in

this study the campaign route is divided differently than in previously published ACE studies

195 (Schmale, Baccarini, et al., 2019), in which divided segments are referred to as *legs*.

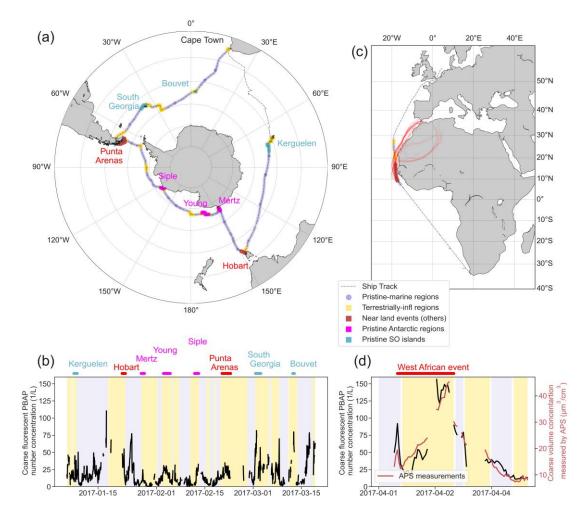
Additionally, Figure 1c shows the ship return path from Cape Town to Bremerhaven, Germany,

along the west coast of Africa, where a Saharan dust plume was likely intercepted by the ship

based on high aerosol loads and modelled air mass back trajectories. We compared the results

199 from this period against the SO measurements in Section 3.3.

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202 203 Figure 1 (a) Map of the investigated segments (segments 1, 2 and 3) of the ACE cruise. The map shows regions defined as pristine-marine in blue (further than 200 km from land masses) and terrestrially-influenced in yellow (closer than 200 km to land masses). Several 204 terrestrially-influenced regions where relatively strong fluorescent particle events occurred are also shown in the figure, and these are 205 further classified into pristine SO islands (cyan), pristine Antarctic (magenta), and near populated continenetal regions events (red). (b) 206 207 Time series of the fluorescent PBAP number concentration measured during ACE campaign (3o threshold). The time periods of pristinemarine (blue shade), terrestrially-influenced (yellow shade) and other selected events are highlighted in the time series. (c) Map of the 208 return path along West Africa where a dust plume was intercepted by the ship (red region). Five-day air mass back trajectories 209 calculated with the Lagrangian analysis tool LAGRANTO (Sprenger & Wernli, 2015; Thurnherr et al., 2020) are shown as red lines. The 210 back trajectories during the event (red lines) indicate that air travelled south along West Africa before reaching the ship. (d) Time series 211 of fluorescent PBAP number concentrations and total coarse aerosol volume concentrations (measured by the APS) for the West African 212 section shown in (c). The blue shades correspond to pristine-marine time periods, the yellow shades are the terrestrially-influenced and 213 the red line is the presumed dust event.

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2.2 Fluorescent Aerosol Measurements

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We used a wideband integrated bioaerosol sensor (WIBS-4, University of Hertfordshire, Hatfield, UK) to measure fluorescent aerosol particles on a single-particle basis. Ambient air was drawn from a standard Global Atmospheric Watch air inlet mounted onto a laboratory container, where the WIBS was located along with other aerosol instruments (more details can be found in

220 Schmale, Baccarini, et al., 2019). The WIBS inlet flow rate was 2.5 l/min, of which 0.23 l/min is

the sample flow and the remaining 2.27 l/min are filtered and used as sheath flow. The WIBS 221 measures the aerosol optical diameter based on elastic light scattering by exposing incoming 222 aerosol particles to a continuous 635 nm diode laser. The light scattered from individual particles 223 is measured in the forward direction by a quadrant photo multiplier tube (PMT) detector, and at a 224 90° angle relative to the laser beam by a second PMT. The aerosol optical diameter in the size 225 range from 0.5 to 14 µm is inferred from the 90° side scattering measurements. The forward 226 scattered signals measured by the quadrant detector are used with Eq. S1 in the supporting 227 information to derive the aerosol asymmetry factor (AF), which is a measure of aerosol 228 morphology. Toprak and Schnaiter (2012) demonstrated that an AF value of 8 represents 229 spherical particles, while higher AF values are associated with non-spherical particles. 230

Upon detection of an aerosol particle through the scattering signal, two xenon flash lamps 231 provide UV excitation at wavelengths of 280 and 370 nm sequentially. The fluorescent signals 232 233 from individual particles are measured in two different channels with bands of 310-400 nm and 420-650 nm. The frequency of the xenon flash lamps and hence the single particle detection 234 frequency is 125 Hz, which implies that a portion of fluorescent particles will not be detected if 235 the aerosol number concentration is above 25'000 L<sup>-1</sup>. On the other hand, based on the elastic 236 scattering measurements, the WIBS provides the number of missed particle counts between 237 sequential UV source activations. Analysis of the data for different segments revealed that the 238 median of the missing particle fraction to total aerosol number concentration measured by the 239 WIBS ranged between 5 to 8 % (Figure S1). Due to the small portion of missed particles, we did 240 not consider their contribution in this study. 241

The combination of two excitation wavelengths (ExWL) and two emission wavebands 242 (EmWB) provides three different valid fluorescent signal configurations, while one combination 243 is invalid due to interference from the excitation laser. The configuration of the fluorescent 244 channels are: 245

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Channel 1 (or A): ExWL of 280 nm and EmWB of 310-400 nm Channel 2 (or B): ExWL of 280 nm and EmWB of 420-650 nm • Channel 3 (or C): ExWL of 370 nm and EmWB of 420-650 nm •

It should be noted that the ExWL of 280 nm is selected to excite tryptophan while the ExWL of 370 nm is targeted toward excitation of NADH. Moreover, the peaks in the fluorescent signals for tryprophan and NADH occur at EmWB of 310-400 nm and 420-650 nm, respectively.

An exited aerosol particle is considered to be fluorescent if its emitted fluorescent signal 252 detected by any of the fluorescent channels is above certain thresholds. The fluorescent 253 thresholds are defined based on the fluorescent signals of the instrument background which are 254 measured periodically through the so called "forced triggering" process. Each threshold is 255 defined based on a certain increment above the mean value of the raw signal during forced 256 trigger mode. It is common practice to choose a fixed multiple of the raw signal standard 257 deviation ( $\sigma$ ) as increments in order to account for random noise. In this study, we applied and 258 compared increments of  $3\sigma$  and  $9\sigma$  as two alternative threshold settings, as previously applied by 259 Savage et al. (2017). We distinguish the results obtained with these two different threshold 260 settings by referring to them as the fluorescent particle  $(3\sigma)$  and hyper-fluorescent particle  $(9\sigma)$ 261 results. It is important to note that the hyper-fluorescent particles are the subset of fluorescence 262 263 particles displaying the strongest fluorescent signals.

Fluorescent aerosols can be classified into different groups based on combinations of the emitted signals detected in the different fluorescent channels. We use the classification scheme introduced by Perring et al. (2015). In this method, the fluorescent particles are divided into 7 different classes (A, B, C, AB, AC, BC, and ABC) based on the logical combination of emitted signals in the 3 fluorescent channels. Table 1 provides the description for all seven fluorescent particle types defined based on the Perring et al. (2015) method.

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Table 1. Description of different fluorescence classes following the classification scheme presented by Perring et al. (2015).
 The AND and NOT in this table correspond to logical 'and' and 'not', respectively.

| Fluorescence | Definition of fluorescence class                                     |
|--------------|--|
| class        |  |
| А            | Fluorescent aerosol detected in channel 1 but NOT in channel 2 and 3 |
| В            | Fluorescent aerosol detected in channel 2 but NOT in channel 1 and 3 |
| С            | Fluorescent aerosol detected in channel 3 but NOT in channel 1 and 2 |

| AB  | Fluorescent aerosol detected in channel 1 AND 2 but NOT in channel 3 |
|-----|--|
| AC  | Fluorescent aerosol detected in channel 1 AND 3 but NOT in channel 2 |
| BC  | Fluorescent aerosol detected in channel 2 AND 3 but NOT in channel 1 |
| ABC | Fluorescent aerosol detected in channel 1 AND 2 AND 3                |

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It should be noted that other methods for classifying fluorescent particles are also available. Toprak and Schnaiter (2012) used a slightly different classification method. Their study indicated fluorescent particles detected simultaneously in WIBS channels 1 and 3 could be defined as a robust indicator class for fluorescent bioaeorsol particles (FBAP) with low crosssensitivity to non-biogenic aerosol. The FBAP class defined in Toprak and Schnaiter is equivalent to particles identified as AC or ABC based on the classification scheme used in this study.

One of the major challenges in processing WIBS measurements is to consider the 281 interference of fluorescent aerosols of non-biological origin, e.g., fluorescent particulate matter 282 in the ship exhaust such as polycyclic aromatic hydrocarbons (PAHs). For the atmospheric 283 samples measured during ACE, we used an empirical masking technique to remove samples that 284 were suspected to be contaminated by ship exhaust (Schmale, Baccarini, et al., 2019). In short, 285 aerosol number concentrations (CN, measured by a condensation particle counter with a time 286 287 resolution of 10 seconds) and ambient CO<sub>2</sub> concentrations obtained by a PICARRO (measured at 1 Hz) were used. Then binomial smoothing over 60 data points was applied to both time series. 288 Periods were classified as polluted when the ratio of the 1 minute CN over the smoothed time 289 series was greater than 1.24 or smaller than 0.51, or when the ratio of the 1 minute CO<sub>2</sub> signal 290 over the smoothed CO<sub>2</sub> time series deviated by 20 %, or when the absolute change between CN 291 292 at time t and t+1 was larger than 50. In addition to this mask, we used a second filter based on wind direction to further minimize the risk of including ship-exhaust-influenced measurements 293 294 in the analyzed dataset. Specifically, periods when the wind was blowing from between 90 and 270 ° relative to the ship's main axis sample (with 0 ° referring to the ships bow being pointed 295 296 into the wind and 90° referring to wind coming from starboard). Approximately 44% of the measurements acquired during the campaign were discarded by these two filters. 297

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Mineral dust particles can also generate measurable fluorescence signals in the WIBS instrument (Savage et al., 2017). However, the results that we present later in Section 3.3 suggest that long-range transported dust aerosols did not contribute substantially to the remote oceanic measurements. Therefore, we assume that all measured particles remaining after application of the ship exhaust filters are PBAP, and we refer to these hereafter as '(hyper-) fluorescent PBAP'.

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2.3 Auxiliary atmospheric measurements used as proxies for SSA concentrations

We use a range of auxiliary atmospheric measurements in this study as proxy variables for the concentration of SSA in the air. It is necessary to use proxies for the concentration of SSA since it is difficult to measure this parameter directly, due to the fact that it is difficult to isolate SSA from other aerosol types found in the marine atmosphere like non-sea-salt sulfates (e.g. Modini et al., 2015).

Wind speed is often used as an indicator for SSA since SSA source strength and number 311 concentration depend strongly on wind speed through wave breaking (Lewis & Schwartz, 2004). 312 Although wind speed is a useful indicator of SSA production, one must always keep in mind 313 potential differences between wind speeds at the point of SSA production and wind speeds at the 314 point of measurement (in this case the research vessel), which complicates SSA-concentration-315 wind-speed relationships. Here we report wind speeds as 10-meter neutral wind speeds, which 316 were derived from the on board measurements (including a correction for air-flow distortion) as 317 described in Landwehr, Thurnherr, et al. (2020). 318

319 The dominant inorganic chemical component of SSA is NaCl (e.g. Bates et al., 2008). Therefore, the concentrations of sodium and chloride are useful markers for SSA (e.g. Modini et 320 al., 2015; Quinn et al., 2017). Sodium ion concentrations were measured for sub-10 µm aerosols 321 using ion chromatography, which was performed offline on filter samples that had been collected 322 323 over 24 hours (Tatzelt et al., 2020). Inorganic chloride concentrations (Chen et al., 2019) were 324 measured by a time-of-flight aerosol chemical speciation monitor (ToF-ACSM, Aerodyne Research, Inc.; Fröhlich et al., 2013). The ACSM is only sensitive to the non-refractory, 325 submicrometer fraction of the total aerosol (i.e., the fraction that undergoes flash vaporization at 326

600 °C). Therefore, the ACSM is only able to detect a very small fraction of the total chloride in SSA. This signal can be easily overwhelmed by anthropogenic sources of non-refractory chloride (e.g. ammonium chloride), which prevents the use of ACSM chloride as a marker for SSA in environments with strong continental or anthropogenic influences. In the remote SO such influences are largely absent, and we assume that ACSM chloride represents SSA chloride qualitatively well.

For the same reason of geographical remoteness, we also assume that the number 333 334 concentration of particles with diameters larger than 1 µm is a good proxy variable for SSA 335 concentrations. That is, we assume that super-micrometer particles with optical diameter larger than 1 µm, hereafter referred to as coarse mode, are composed predominantly of SSA particles. 336 This is a reasonable assumption to make in remote marine locations since there are no major 337 sources of coarse mode particles other than SSA production (on a number basis). The number 338 339 size distributions of total aerosol particles (i.e. both fluorescent and non-fluorescent particles) was obtained from the elastic scattering measurements performed with the WIBS, and these were 340 341 integrated over diameters greater than 1 µm to calculate super-micrometer number concentrations. Coarse aerosol number size distributions (Schmale, Henning, et al., 2019) were 342 also measured by an Aerodynamic Particle Sizer (APS, TSI Inc., Model 3321). Integrated super-343 micrometer number concentrations from the WIBS and APS correlated well during segments 1-344 3, lending confidence to the measurements from both instruments (Figure S2). The integrated 345 number concentrations also correlated well for the subset of measurements acquired after the 346 ACE campaign (i.e., during the passage from Cape Town back to Europe), but the absolute ratio 347 between these two parameters was higher compared to the value measured during segments 1-3. 348 This suggests a drift in one or both of these instruments. Therefore, we consider the WIBS data 349 measured during the return passage from Cape Town to Europe to be more uncertain than the 350 351 WIBS data measured during segments 1-3.

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We used additional measurements from other ACE projects, No. 1 and 8 (Walton &

355 2.4 Oceanic measurements

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Thomas, 2018) to investigate links between airborne fluorescent PBAP and seawater 358 composition (including dissolved compounds and microbial characteristics). 359 Seawater from approximately 5 meter depth was sampled from an underway seawater 360 supply and preserved for later analysis or measured on-board. In this study we use the 361 measurements of microbial composition (phytoplankton taxa relative pigment biomass 362 contributions) (Antoine et al., 2019), biomass (particulate organic carbon (Thomalla et al., 2020), 363 total chlorophyll-a concentration, and absorption by coloured dissolved oranic matter), microbial 364 cell abundance (e.g. bacterial cell number concentration), and concentrations of transparent 365 exopolymeric particles (TEPs) and coomasie stainable particles (CSPs) (measured as in 366 Zamanillo et al., 2019). 367 A complete description of all ocean measurements is available in supplementary Section 368 369 S.3 and Tables S.1 to S.3, while the ACE Cruise Report (Walton & Thomas, 2018) provides further information on the objectives and sampling methods. 370 371 2.5 Data analysis considerations and segregation of the measurements 372 373 374 The main objective of this study is to investigate ocean-derived fluorescent PBAP, (i.e. those primary biological particles that are thought to be emitted with SSA). To isolate such 375 376 particles we segregated our measurements into two main categories: pristine-marine and *terrestrially-influenced* samples. This segregation was performed based on proximity to land. 377 Measurements that were performed within 200 km distance from any coastline (continental land 378 mass or island) were classified as terrestrially-influenced, while all other measurements were 379 380 identified as pristine-marine. The 200 km threshold was chosen by examining the coefficients of correlation between fluorescent particle number concentrations and three of the proxy variables 381 382 for SSA concentrations (wind speed, total number of coarse mode particles, chloride concentration) as a function of the proximity to land. This analysis is shown in Figure S3 for the 383

fluorescent particle category and Figure S4 for the hyper-fluorescent particle category (as defined in Section 2.2). For both categories, the coefficients of correlation reach a plateau at a distance greater than approximately 200 km. Therefore, we chose this distance as the threshold to segregate pristine-marine and terrestrially-influenced samples.

It should be noted that other methods for segregating land-influenced and oceanic 388 samples are also possible. For example, air mass back trajectories could be used to perhaps 389 obtain a clearer separation of the terrestrially-influenced measurements. We did not apply this 390 391 method in this study because it carries a greater risk that some terrestrially-influenced samples 392 are classified as pristine-marine samples due to uncertainties in the calculated air mass back trajectories. Since our goal was to focus specifically on ocean-derived particles, we instead opted 393 for the simple but conservative threshold value of 200 km from any land mass. The corollary of 394 this approach is that our terrestrially-influenced category likely also contains a sizeable fraction 395 396 of pristine-marine measurements, which we deemed to be an acceptable consequence since mixed marine-terrestrial aerosols are not the focus of our study. At the same time, this approach 397 398 provides a good estimate of the radius of influence of terrestrial PBAP sources.

399 We segregated the aerosol fluorescence measurements by optical particle diameter as 400 measured by the WIBS-4. In particular, we categorized the measurements into *fine* (optical diameter  $< 1 \mu m$ ) and *coarse* (optical diameter  $> 1 \mu m$ ) aerosol categories. Our main focus is on 401 the coarse particles since: 1) larger particles are less likely to be long-range transported and can 402 therefore be more confidently attributed to local, oceanic sources; 2) any contamination particles 403 404 such as soot remaining after application of the ship exhaust post-processing filters described in Section 2.2 are more likely to reside in the fine category than the coarse category; and 3) the 405 WIBS counting efficiency deteriorates for particles with diameters less than 0.7 µm (Healy et al., 406 2012). The consequence of our decision to focus on coarse particles is that we possibly exclude 407 certain types of PBAP, e.g. bacteria with sizes below 1 µm (Fröhlich-Nowoisky et al., 2016). 408

409 **3 Results and discussion** 

410

3.1 Time series of fluorescent PBAP number concentrations over the campaign

411

Figure 1 presents the time series of coarse fluorescent PBAP number concentrations 412 measured over the entire ACE campaign. During pristine-marine conditions, fluorescent PBAP 413 number concentrations varied considerably and ranged between 0.17 and 120.1 L<sup>-1</sup>. The median 414 number concentration was 11.4  $L^{-1}$  with interquartile range (IOR) ranging between 5.6  $L^{-1}$  to 415 21 L<sup>-1</sup>. The median number concentration of coarse hyper-fluorescent PBAP was 0.87 L<sup>-1</sup> with 416 IOR ranging between  $0.37 L^{-1}$  to  $1.95 L^{-1}$ . The corresponding concentrations in the terrestrially-417 influenced regions were higher than those in the pristine-marine regions. The median number 418 concentration of fluorescent particles in the terrestrially-influenced regions was 17.3 L<sup>-1</sup> and the 419 IOR ranging between 6.5  $L^{-1}$  to 27.8  $L^{-1}$ . For hyper-fluorescent particles under terrestrial 420 influence, the median number concentration was 1.52 L<sup>-1</sup> with IQR ranging between 0.58 L<sup>-1</sup> and 421 2.9 L<sup>-1</sup>. 422

It is important to note the diversity of the terrestrial areas that contributed to the land-423 424 influenced measurements. As shown by the cruise map and time series displayed in Figure 1 the terrestrially-influenced samples comprised measurements that were performed near the continent 425 426 of Antarctica, near pristine and unpopulated islands in the SO, and near the Australian (Hobart) and South-American (Punta Arenas) continents. When the ship passed through terrestrially-427 influenced regions close to uninhabited islands and coastal regions, as well as more populated 428 continental areas, high peaks in concentrations of fluorescent PBAP, reaching up to 90  $L^{-1}$ , were 429 occasionally observed. We visually identified nine of these high-conentration events, as 430 indicated in Figure 1: three occurred in the vicinity of pristine SO islands (Kerguelen, South 431 Georgia, and Bouvet), three near continental Antarctica (Mertz Glacier, Young and Siple 432 Islands), and three near populated continental regions (Hobart, Punta Arenas, and West Africa on 433 the return route). The highest fluorescent particle concentrations were measured during the West 434 African event, when hourly-averaged concentrations reached up to  $160 L^{-1}$ . The back trajectories, 435 which are included in Figure 1c, indicate that some air masses passed over the Saharan desert. In 436 addition, Figure 1d shows the integrated aerosol volume concentration of coarse particles 437 obtained from APS measurements for the West African event, indicating an increase in 438 integrated volume concentration of aerosol particles during this period. Therefore, we identify 439 the fluorescent particles measured during the West African event as Saharan dust particles. 440 Although our main focus in this work is on pristine-marine PBAP, the various different near-land 441 measurements provide an insightful contrast for the remote ocean measurements. Difference 442

between near land events and pristine-marine samples are further investigated through themeasured fluorescence classes in section 3.5.

- 3.2 Demonstration of the link between fluorescent PBAP and SSA particles in the
   pristine-marine atmosphere
- 447

Based on previous studies of SSA composition it is hypothesized that SSA production is 448 449 the dominant source of PBAP in the remote oceanic regions far from land where the contribution of long-range transported aerosol particles is less likely (see Section 1). To investigate this 450 hypothesis, we assessed the level of correlation between measured fluorescent particle number 451 concentrations and four proxy variables for SSA concentrations (Section 2.3; wind speed, total 452 coarse aerosol number concentrations, and aerosol chloride and sodium mass concentrations). 453 We used the combined results from segments 1-3 of the research cruise for this correlation 454 analysis. 455

Figure 2 presents scatter plots of hourly averaged hyper-fluorescent PBAP number 456 concentrations against the four variables (note Figure 2d presents 24h averaged measurements to 457 match the filter sample collection periods). The results are split into pristine-marine (blue points) 458 and terrestrially-influenced samples (red points) as described in Section 2.5. For the pristine-459 marine samples, moderate correlation is observed between the hyper-fluorescent number 460 concentrations and all four proxies for SSA concentrations (Pearson's R values ranging from 461 0.37 to 0.61). These results suggest that the same underlying process drives the variability in all 462 463 of these measured quantities, which supports the hypothesis that sea spray is an important source of fluorescent PBAP in the pristine-marine atmosphere. 464

Further supporting evidence for the hypothesis is provided by the terrestrially-influenced 465 results. In all four correlations shown in Figure 2, the correlation coefficients are lower for the 466 terrestrially-influenced samples than the corresponding pristine-marine samples. In contrast, the 467 absolute concentrations of hyper-fluorescent PBAP of pristine-marine and terrestrially-468 influenced samples are similar (as indicated in Sec 3.1 and depicted in Figure 2), with respecting 469 IQRs spanning  $0.37-1.95 \text{ L}^{-1}$  and  $0.58-2.9 \text{ L}^{-1}$ . The median value is slightly higher for the 470 terrestrially-influenced  $(1.52 \text{ L}^{-1})$  than pristine-marine  $(0.87 \text{ L}^{-1})$  samples. Altogether, this 471 indicates that the lower correlation values for the terrestrially-influenced samples are primarily 472

the result of few observations of much higher PBAP concentrations, which we attribute to additional PBAP sources near coastlines. A similar picture emerges when including the PBAP with weaker fluorescence ( $3\sigma$  threshold) as shown in Figure S5. The conservative, 200 km distance-from-land threshold we applied to segregate the measurements (Section 2.5) explains why the terrestrially-influenced samples remain similar to the pristine-marine subset, while the loss of correlation demonstrates the necessity of properly segregating the dataset to exclusively isolate those fluorescent particles that are related to SSA production.



481

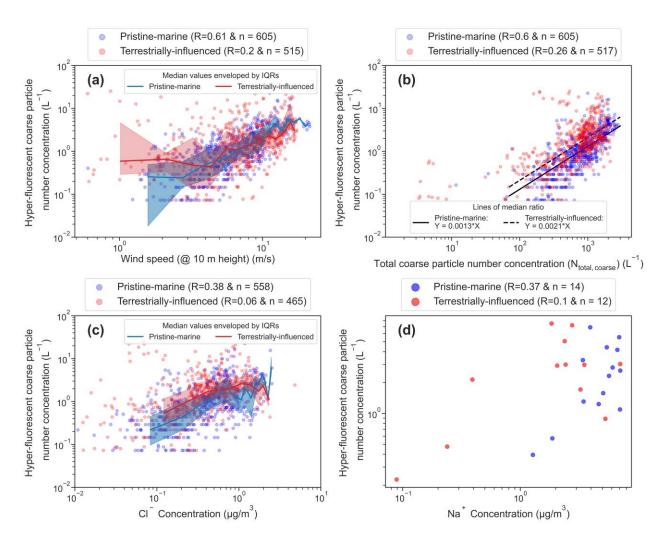


Figure 2. Scatter plots of pristine-marine (blue) and terrestrially-influenced (red) samples of hyper-fluorescent particles
 vs four proxy variables for SSA concentrations: a) wind speed, b) total coarse particle concentrations, c) chloride (Cl-)
 concentrations as measured by the ACSM, and d) sodium (Na+) concentrations measured offline from filter samples.
 Measurements from all segments are shown. The red and blue solid lines and shaded areas correspond to the medians
 and IQRs of the measurements, which were calculated by separating the dataset into ten equidistant logarithmic bins.
 Median lines and IQRs are not shown for the sodium ion measurements due to the small sample size. The number of
 tested samples for each condition (n) is included in the subplots.

489 3.3 Quantification of the contributions of fluorescent PBAP to coarse SSA concentrations
 490 in the pristine marine atmosphere

491

The moderate correlation observed between (hyper-)fluorescent PBAP number 492 concentrations and total coarse particle concentrations (Figures 2b and S4b, respectively) for the 493 pristine-marine samples suggests that the former quantities can be estimated from measurements 494 or calculations of the latter. Histograms of the ratios of hyper-fluorescent and fluorescent PBAP 495 concentrations to total particle concentrations are shown in the Figures S6 and S7. The median 496 values of these ratios are plotted as straight lines in Figures 2b and S5b. These results indicate 497 that for pristine-marine samples the median contributions of hyper-fluorescent and fluorescent 498 PBAP to total super-micrometer SSA concentrations were 0.13 and 1.6 %, respectively. For the 499 terrestrially-influenced samples, the median contributions of hyper-fluorescent and fluorescent 500 PBAP to total fluorescent were 0.21 and 2.2 %. Although it remains to be seen if similar 501 fractions are obtained in other oceanic regions and during different seasons, these estimates 502 503 provide a means for estimating super-micrometer fluorescent PBAP number concentrations from measured or modelled SSA concentrations. 504

505

# 3.4 Modulation of fluorescent PBAP number fractions in SSA by marine biologicalactivity

508

We have demonstrated a clear link between fluorescent PBAP and SSA concentrations in 509 the pristine-marine atmosphere. According to the previous studies discussed in the Introduction, 510 this link is likely formed by marine microorganisms and DOM that are co-emitted with sea salt 511 during the SSA production process. Therefore, fluctuations in the abundance of marine biota 512 could potentially modulate the fraction of observed fluorescent aerosols. It is important to note 513 that fixed relationships should not necessarily be expected, given the complex, intermediate 514 515 aerosol generation and loss processes that link seawater composition with atmospheric aerosol 516 properties. Nevertheless in this section, we qualitatively assess any potential links by examining correlations between seawater composition measurements and the fluorescent aerosol 517 measurements. 518

#### manuscript submitted to Journal of Geophysical Research

Twenty four different types of marine biological and chemical measurements were considered in this analysis. A description of these marine variables is provided in the SI (Section S.3). In short, the marine variables consisted of three distinct classes: 1) number concentrations of different microorganisms obtained from flow cytometry measurements, 2) mass concentrations of different phytoplankton taxa inferred from phytoplankton pigment measurements, and 3) organic matter (OM) measurements which corresponds to DOM (CDOM) and gel-like POM (TEP and CSP) measurements.

We performed correlation analysis separately for the pristine-marine and terrestrially-526 influenced groups of measurements in order to isolate the SSA-related fluorescent PBAP. The 527 number fractions of fluorescent PBAP were considered rather than absolute number 528 concentrations to minimise the risk of falsely identifying associations between the oceanic and 529 atmospheric measurements due to cross-correlation (e.g. to wind speed, which is an important 530 531 driver of SSA and marine PBAP production, as shown in Figure 2a, and which might also influence some of the marine variables). In addition, absolute aerosol concentrations are affected 532 533 by variable atmospheric loss processes, which complicates their use in such a correlation analysis. It is reasonable to assume that similar loss processes occur for similarly sized 534 fluorescent PBAP and non-fluorescent aerosol particles, and therefore that fluorescent PBAP 535 fractions are much less sensitive to variations in these loss processes. 536

537 Number fractions of fluorescent PBAP were calculated by normalizing the coarse 538 fluorescent PBAP number concentrations by the total coarse particle number concentrations 539 simultaneously measured by the WIBS. Marine point samples were extracted from oceanic water 540 with sampling frequencies which varied from 1 to 6 hours for different marine variables. To 541 perform the correlation analysis, the results of each marine point sample were simply paired with 542 the overlapping 1h average of fluorescent PBAP concentration data, justified by limited variation 543 of the latter during 1 hour intervals.

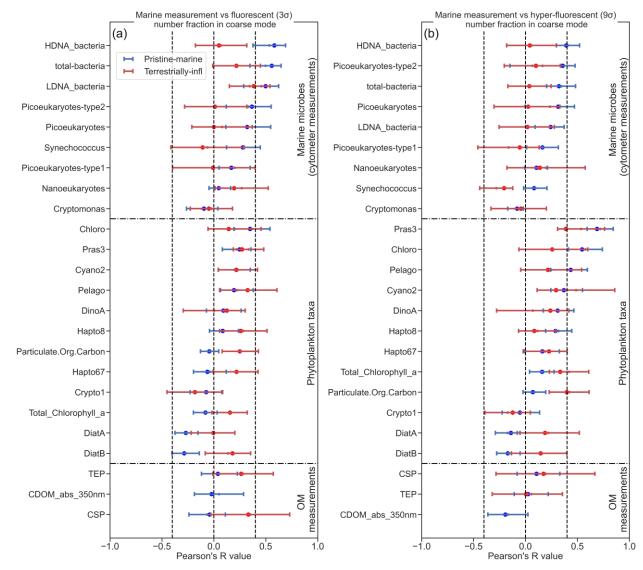




Figure 3. Pearson's correlation coefficients (Pearson's R) from the correlation analysis of marine variables against
 number fractions of (a) fluorescent PBAP (3σ threshold) and (b) hyper-fluorescent PBAP (9σ threshold), both relative to
 total coarse aerosol particle number concentration. The horizontal dot lines separate different types of marine
 measurements namely, cytometer marine microbe number concentration measurement, phytoplankton taxa mass
 concentration results and OM measurements. The error bars are obtained from the bootstrap analysis.

The Pearson coefficients of correlation between coarse fluorescent and hyper-fluorescent 550 number fractions and the different marine variables are displayed in Figure 3. Corresponding p-551 values calculated with a permutation test are shown in Figure S18. The p-values indicate that the 552 correlation results are statistically significant at the 90% level (i.e., p-values less than 0.1), with 553 the exception of the Chloro, Cyano2, DinoA, Hapto and Crypto1 results. The results are grouped 554 according to the three marine variable categories (microorganism number concentrations, 555 phytoplankton mass concentrations, and OM measurements). To obtain a measure of 556 uncertainties on the correlation coefficient values a bootstrap analysis was performed for each 557

pair of the analysed variables (the correlation coefficient calculation was repeated 100 times with

random selections containing 60% of all the available data points for each pair of variables).

560 Only those pairs of variables with more than 25 simultaneous data points were considered in this

561 correlation analysis. The corresponding scatter plots for all of the tested variables are displayed

562 in Figures S8 to S17.

For both pristine-marine and terrestrially-influenced samples the absolute Pearson's R 563 values associated with the majority of the marine variables were low (-0.4 < R < 0.4). Only a few 564 variables demonstrated more pronounced correlation with the fluorescent (as opposed to hyper-565 566 fluorescent) fraction of coarse particles, with Pearson's R values beyond the -0.4 to 0.4 range. For pristine-marine results, the variables displaying R > 0.4 were the concentrations of bacteria 567 with high DNA content (HDNA), bacteria with low DNA content (LDNA), total bacteria (sum of 568 the former two types) and picoeukaryotes (small-sized eukaryotic phytoplankton, typically 1-569 570 3 µm). The other types of marine measurements (phytoplankton taxa mass concentrations and DOM related measurements) correlated only weakly with fluorescent PBAP number fractions (R 571 572 < 0.4). This correlation analysis, hence, suggests that the variance of fluorescent particles over the pristine ocean were largely influenced by surface-ocean bacteria and, to a lesser extent, small 573 phytoplankton. 574

For the hyper-fluorescent PBAP number fractions the correlation results are distinct and 575 their rankings are different from those of the fluorescent PBAP number fractions. For the 576 pristine-marine samples, the prominent correlating marine variables were from the 577 578 phytoplankton taxa mass concentration results: prasinophytes (Pras3; R≈0.69), chlorophytes 579 (Chloro;  $R\approx 0.55$ ), and pelagophytes (Pelago;  $R\approx 0.44$ ). In addition to the phytoplankton taxa, concentrations of HDNA bacteria showed moderate correlation ( $R\approx0.4$ ), while LDNA bacteria 580 had an R value of 0.27. Similarly to the fluorescent PBAP number fraction results, the DOM 581 related variables only weakly correlate (R values < 0.4) with the hyper-fluorescent number 582 fractions. The larger contribution of phytoplankton over bacteria to the variance of hyper-583 fluorescent PBAP can be expected since phytoplankton cells are larger than bacteria and 584 therefore likely contain more fluorescent components. 585

The lack of correlation between the OM measurements and the fluorescent and hyperfluorescent PBAP fractions does not imply that DOM or gel-like POM do not contribute to the

biologically-derived organic matter in SSA. Indeed, transparent expolymeric particles (TEPs) 588 and coomasie stainable particles (CSPs) were abundant in seawater throughout the entire ACE 589 cruise and therefore, these organic matter components were likely incorporated into SSA 590 particles. The low correlations observed for the OM category in Figure 3 could be due to weak 591 fluorescent emission of the organic compounds comprising DOM and gel-like POM within the 592 WIBS detection range. Additionally, DOM is expected to be distributed more homogeneously 593 across individual SSA particles compared to insoluble POM. Hence, DOM may be less likely to 594 produce single particles with sufficiently strong fluorescence for detection by the WIBS. 595

596 The terrestrially-influenced results indicate systematically lower R values for those marine variables that display the highest correlation coefficients with the pristine-marine 597 samples. Such systematic deterioration of correlation is consistent with the correlation analysis 598 performed in Section 3.2 with the proxy variables for SSA concentrations (Figure 2), which 599 600 further strengthens the point that the presence of terrestrial aerosols weakens correlations between atmospheric aerosols and marine variables. It is likely that marine biological activity 601 602 could be enhanced near some of the land masses due to nutrient abundance (Gove et al., 2016), and a few marine variables show correlation coefficients of  $\sim+0.4$  or greater for the air masses in 603 proximity to land. However, such results could be due to cross-correlations with changes in 604 marine biota near land. Therefore, no attempt is made to further interpret this subset of data. 605 Additionally, it should be noted that terretrially-influenced samples typically possess smaller 606 sample size (the average terrestrially-influenced marine samples were ~ 25% of the total marine 607 samples) and are statistically less significant than the oceanic samples, as seen from the error 608 bars. 609

Overall, two main points can be drawn from these correlation results. Firstly, 610 (hyper-)fluorescent aerosol number fractions in the coarse mode correlate best with variables 611 related to marine microorganisms (bacteria and phytoplankton types). This suggests that marine 612 microorganisms are likely incorporated into SSA, and that variations of their concentrations in 613 the ocean modulates the fluorescent fraction of SSA. This strengthens the hypothesis that the 614 observed fluorescent particles are indeed PBAP. Secondly, the results suggest that those aerosol 615 particles possessing the strongest auto-fluorescent properties (hyper-fluorescent particles) 616 correlate to different marine variables than the regularly fluorescing particles. Specifically, the 617

hyper-fluorescent PBAP fraction correlates more strongly with phytoplankton than bacteria,presumably because phytoplankton are larger and contain more fluorescent material.

Further elaboration is required regarding the different sizes of the marine microbes 620 measured in this study relative to the size detection limits of the WIBS (i.e., aerosol particle 621 diameters from 0.5 to 14 µm). For example, prasinophytes (Pras3) – the mass concentrations of 622 which correlated most strongly with hyper-fluorescent PBAP number fractions - are amongst the 623 smallest-sized microalgae. Bacteria, which were among the highest correlating variables with 624 respect to the fluorescent PBAP fractions, are even smaller, with typical sizes in the range of 0.5 625 626 to 1 µm. Conversely, the number concentrations of cryptomonas correlated very weakly with (hyper-)fluorescent PBAP number fractions. Cryptomonas particles have typical sizes of  $\sim 40$ 627 µm, which is generally larger than the other microbes measured in this study, and which may 628 have rendered them undetectable by the WIBS even if they were injected into the atmosphere in 629 SSA. However, such large airborne microbes would display relatively high settling rates and 630 short atmospheric lifetimes, meaning they are less likely to be transported far from their source 631 632 regions. Therefore, regardless of the limitations of the WIBS measurements, baceteria and small phytoplankton are anyway more likely to contribute substantially to pristine marine PBAP than 633 much larger airborne microbes like cryptomonas. 634

In conclusion, this correlation analysis suggests that certain types of marine microbes have the potential to modulate the fractions of fluorescent particles in SSA, which is generally consistent with previous studies (e.g. Mayol et al., 2017; Uetake et al., 2020). Further dedicated and targeted measurements are required to confirm if the most highly correlating marine variables observed in this study (concentrations of bacteria and certain phytoplankton types) also have an impact on fluorescent PBAP away from immediate source areas in other oceanic regions and during other seasons.

642

3.5 Classification of different fluorescent particle types

643

In this section the fluorescent aerosols are discussed according to the ABC classification scheme of Perring et al. (2015) (Section 2.2 and Table 1). We present classification results for both the pristine-marine samples and the nine near-land events identified in Section 3.1, in order to compare and contrast the fluorescent properties of particles originating from sea spray versusthose from the various different terrestrial sources.

Figure 4 shows the number fractions of each ABC fluorescence class for the three
pristine-marine cruise segments and the nine near-land events. Results are displayed for both the
fluorescent and hyper-fluorescent particles. Since it only made negligible contributions, type AC
particles are excluded from Figure 4 for visual clarity.

The most prominent fluorescence classes in the pristine-marine samples are A, B, AB and 653 ABC. Class C is a prominent class for the particles at the  $3\sigma$  fluorescence threshold (Figure 4a), 654 but its fractional contribution is substantially reduced for the hyper-fluorescent (9 $\sigma$  threshold) 655 particles (Figure 4b). The fluorescent particle results indicate that the relative proportions of 656 these classes (mean, median and IQRs) are very similar throughout the cruise segments 1 to 3 657 (top three rows in each panel). For the hyper-fluorescent particles, the mean fractions of each 658 class (red triangle markers) are very consistent across segments 1 to 3, while the IQR results for 659 segment 2 are less consistent with the other segments. This could be due to the fact that segment 660 2 samples were collected further south compared to the other segments, where the presence of 661 sea ice may have resulted in different types of marine microorganisms contributing to the 662 pristine-marine hyper-fluorescent PBAP. 663

The fluorescence class fractions varied more substantially between the nine near-land events than they did between the different pristine-marine cruise segments. For example, the median fractions of type A and B fluorescent particles ( $3\sigma$  threshold) ranged between ~20 to 65% and 5 to 50%, respectively, for the near-land events, while the corresponding ranges for the pristine-marine samples were only 25 to 30% and 30 to 40%, respectively. Such large variability for the near land events can be expected since the composition of fluorescent aerosols and their respective sources might vary substantially between different types of geographical locations.

The fluorescence class fractions for the pristine SO island events (e.g. Kerguelen and South Georgia), and to a lesser extent for the Mertz glacier, are similar to the fluorescence class fractions of the pristine-marine samples. This might suggest that these near land events were mainly influenced by pristine-marine aerosol sources. Interestingly, Hobart and Punta Arenas events, which are not regarded as pristine, show fluorescence class fraction compositions which are not significantly distinct from the other pristine near-land samples. The only noticeable difference is the higher relative prominence of type ABC and AB particles during the Hobartevent.

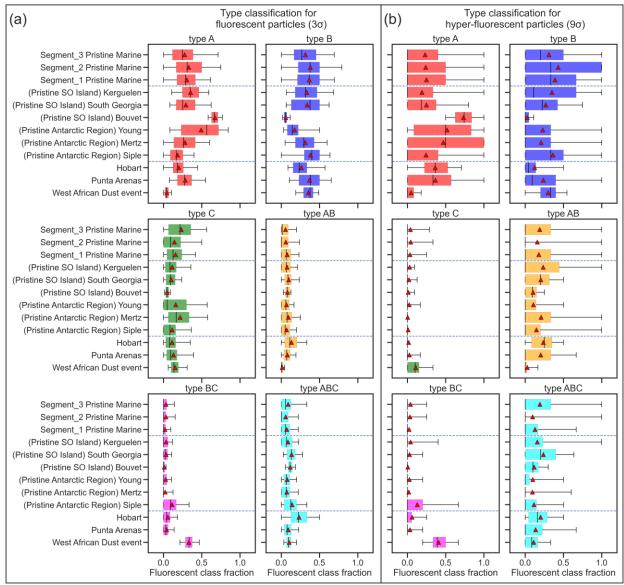
We also show the fluorescence class fraction West African dust event in Figure 4 (this 679 event was discussed and identified in Section 3.1). The fluorescence class make-up of the 680 particles measured during this event was distinctly different to those measured during both the 681 pristine-marine segments and during the other near-land events. During the West African event 682 type BC particles were very prominent for the fluorescent samples, while type BC and type C 683 684 particles were prominent for the hyper-fluorescent samples. This is an indication that the 685 particles observed during this event possessed distinctly different fluorescent properties compared to the particles that were measured in the SO region. As discussed in Section 3.1, we 686 interpret the fluorescent particles measured during the West Africa event as fluorescing dust 687 aerosols. Therefore, this comparison suggests that long-range transported dust particles – at least 688 those originating in the Saharan desert and/or those having a similar fluorescence class make-up 689 as Saharan dust particles – did not contribute substantially to the particles measured over the 690 691 remote SO during the ACE cruise. This result is in contrast to the study conducted by Crawford et al. (2017) at the Halley VI Research Station in Antarctica in austral summer 2015. They 692 concluded that long-range transported dust particles, perhaps transported from the southern tip of 693 South America, contributed substantially to the fluorescent particles observed at that Antarctic 694 695 site.

The observed differences and larger variability in relative fractions of fluorescent particle 696 697 types for the near-land events compared to pristine-marine samples may also be partly due to the fact that the sample durations of the individual near-land events (which only lasted from  $\sim 12$  to 698 48 hours) are much shorter than the averages over entire segments for the pristine-marine 699 samples. To investigate this further a bootstrap analysis was performed separately for each 700 pristine-marine cruise segment based on 288 randomly-selected pristine-marine data points 701 702 (which is equivalent to 12 hour periods of 5 min averaged data points). These results are presented in Figures S19 to S21. They indicate that the subsamples are consistent with the 703 704 overall results for each segment, which demonstrates that 12 hours of data are sufficient to provide statistically robust medians and IQRs for the relative fractions of different fluorescence 705 706 classes.

707 A second type of subsampling bootstrap analysis, presented in Figures S22 to S24, was based on fixed 24 hours' time windows that contained at least 12 hours of pristine-marine data 708 and that were randomly positioned in a segment. They demonstrate some degree of variability in 709 the median and IQR values for subsamples relative to the entire segments, in particular for 710 segment 3. These deviations might reflect inhomogeneity in the types of local marine 711 microorganisms that contribute to the fluorescent particle populations, as well as variations in 712 atmospheric conditions that affect the aerosol sources and sinks on time scales of 24 hours, such 713 as passing storms. The variability in fluorescent particle type fractions of terrestrially-influenced 714 samples (Figure 4) is larger than that of pristine-marine subsamples (Figures S22-S24) of 715 comparable duration indicating additional or different fluorescent particle sources. 716 Size-resolved ABC classification for pristine-marine conditions is shown in Figure S25. 717

These results suggest that single type classes (A, B, C) are more dominant in the smaller 2  $\mu$ m size range, while the fraction of multi-type classes (AB, BC, ABC) strongly increases for sizes above 2  $\mu$ m. The increasing contribution of multi-type classes could be explained by the fact that greater particle volumes are more likely to accommodate sufficient fluorophores of multiple

types to exceed the signal thresholds of the corresponding channels.



723 724

Figure 4. Box and whisker plots of fraction of fluorescent type for (a) coarse fluorescent ( $3\sigma$ ) and (b) coarse hyper-725 fluorescent ( $9\sigma$ ) particles. The v axis reperent different name of different test cases for pristine-marine samples from 726 segement 1 to 3 and near land events. The selected near land events occurred in Hobart-Tazmenia, Mertz Glacier, Young 727 Island, Siple Island, King Edward Point-South Georgia Island, and Bouvet Island. In these plots the black whiskers correspond to 5<sup>th</sup> and 95<sup>th</sup> percentile and the boxes are the interquartile ranges. The red dots in the plots represent the 728 729 mean values.

- 3.6 Further investigation of the fluorescent particle types approximate humification 730 index results 731
- 732

In addition to the ABC WIBS classification scheme, other metrics have been devised to 733 interpret and classify the types of fluorescing compounds and particles that have been observed 734 in various environments. For example, the so-called humification index has been applied 735 extensively to excitation-emission spectroscopic measurements of organic matters found in 736

seawater, freshwater, and soils (e.g. Chen et al., 2016; Fu et al., 2015; Zsolnay et al., 1999). In 737 these contexts, the humification index is typically defined as the ratio of emission intensity in the 738 wavelength range from  $\sim 400 - 480$  nm to emission intensity in the wavelength range from  $\sim 300$ 739 - 350 nm, given an excitation wavelength of 255 nm. The rationale behind this metric is that at 740 this excitation wavelength, protein-like organic matters tends to display sharper emission profiles 741 at shorter wavelengths, while humic-like organic matters display broader emission profiles that 742 are shifted to larger wavelength ranges. Therefore, large humification index values (i.e.,  $> \sim 10$ ) 743 correspond to samples with strong contributions of humified and aromatic organics, while lower 744 humification index values correspond to samples that are either dominated by or contain large 745 contributions from microbially-derived protein-like organic molecules (e.g. Fu et al., 2015). 746

Unlike emission-excitation spectroscopy measurements which are typically performed at high spectral resolutions, the WIBS only excites particles at two, discrete excitation wavelengths and then detects the resulting fluorescence signals within two broad emission wavebands. Nevertheless, we can still define an approximate humification index for application to the WIBS measurements, which we denote as the  $R_{B2A}$  ratio to highlight that it is not directly comparable to other humification index results reported in the literature, although they are strongly related. We define this ratio as:

$$R_{B2A} = \frac{FL_B}{FL_A} \tag{1}$$

With  $FL_A$  and  $FL_B$  being the fluorescence signal amplitude in channel A and B, 754 respectively. That is, the  $R_{B2A}$  ratio is defined as the ratio of fluorescent signal intensity in the 755 wavelength range from 420 - 650 nm to the fluorescent signal intensity in the range from 310 - 650756 400 nm, given an excitation wavelength of 280 nm. One key difference between the  $R_{B2A}$ 757 parameter and the ABC scheme presented in Section 3.4 is that the  $R_{B2A}$  parameter is a 758 continuous variable, while the ABC approach is a binary classification method (i.e., a given 759 signal is either above or below a given channel's threshold). Thus, we calculated  $R_{B2A}$  values for 760 all types of particles, regardless of whether they displayed fluorescent signals above or below the 761 relevant thresholds in channels A and B. However, to prevent measurement noise at low signal 762 levels influencing the results, measured intensities below the 3 (or 9) standard deviation 763

detection thresholds in channels A and B were simply set equal to the mean value of the forcedtriggering signal for the calculation.

Given that the  $R_{B2A}$  parameter depends on absolute signal intensities, drifts in either or 766 both of the detector channels A and B could contribute to its variability. However, no evidence 767 of substantial drift was observed in the forced trigger data for these two channels over segments 768 1 to 3, suggesting detector drift didn't contribute substantially to the observed variations in  $R_{B2A}$ . 769 Furthermore, given that the absolute signal intensities measured in each channel were not 770 771 routinely calibrated during the campaign (as is standard operating practice for the instrument, 772 routine calibration is not typically required), we focus here only on the relative comparison between the measurements in this study, and refrain from comparing our  $R_{B2A}$  measurements 773 with humification index results reported in other studies. 774

Figure 5 shows the box and whisker plots of  $R_{B2A}$  for both fluorescent and hyper-775 fluorescent cases. These results show that the IQRs for the pristine-marine segments 1 to 3 are 776 very similar, while the IQRs for the near-land events were more diverse and considerably 777 different from the pristine-marine results. These results are consistent with the ABC 778 779 classification results presented in Section 3.4. Both approaches indicate high similarity between 780 pristine-marine air masses throughout all campaign segments and more variability between individual near-land events. This further corroborates sea spray as dominant and quite 781 homogeneous source of fluorescent PBAP in pristine-marine conditions in the SO. 782

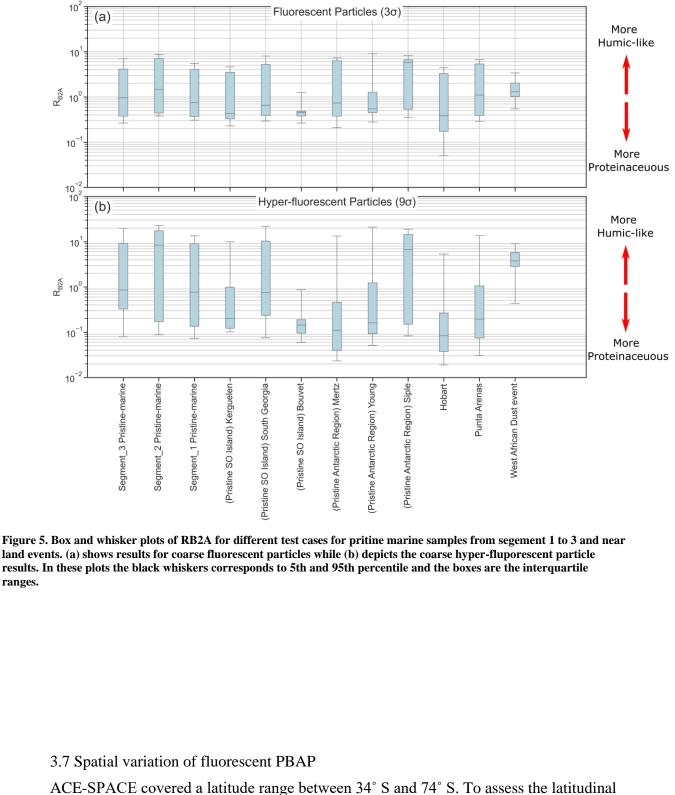
As evidenced by the broad IQRs displayed in Figure 5, considerable variability in  $R_{B2A}$ was observed for all the events except the Bouvet and West African dust events. This suggests that a broad range of different fluorophore types (both protein- and humic-like) contributed to the fluorescent particles observed during most of the events, whereas specific types of fluorescing matter likely dominated the Bouvet and West African dust events.

The highest median value for  $R_{B2A}$  for the fluorescent aerosol condition is 5.8 which corresponds to the event in Siple Island, indicating that the fluorescent particles in Siple are potentially more humic-like than the particles observed during the other events. The Siple event was characterized by highly microbially active waters, as well as land-based penguin colonies and areas of bare soil. Thus, the humic-like signals may have been caused by high levels of humified and aromatic organics, which may have been produced by increased heterotropy (e.g. as occurs during the decay phase of a phytoplankton bloom), or from water outflows off the Siple coast. The median values for other events are considerably lower and range between 0.4 to 1.5, with the Kerguelen, Bouvet, and Hobart events having the lowest  $R_{B2A}$  (median below 0.5) suggesting that fluorescent particles measured during these events are more protein-like on

average.

799 The IQRs and median  $R_{B2A}$  results for the hyper-fluorescent particles differed noticeably from those for the fluorescent particles. In particular, the median  $R_{B2A}$  value for the pristine-800 801 marine segment 2 was substantially higher than the median values during pristine-marine 802 segments 1 and 3, a difference which was not observed for the fluorescent particles. Indeed, under the hyper-fluorescent condition, the  $R_{B2A}$  values for pristine-marine segment 2 are very 803 similar to those measured during the Siple event: median values of 8.4 and 6.8, respectively, the 804 highest median values out of all the events. This indicates higher contributions of humic-like 805 matter to the most strongly fluorescent particles observed during these two events. In contrast, 806 six of the events (Kerguelen, Bouvet, Mertz, Young, Hobart, and Punta Arenas) displayed 807 808 median  $R_{B2A}$  values below 0.5 under the hyper-fluorescent condition (compared to only three events under the fluorescent condition; i.e., Kerguelen, Bouvet, and Hobart). This indicates that 809 for these events, the most strongly fluorescent particles contained greater contributions of 810 protein-like organic matters than the weakly fluorescent particles. 811

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variability of fluorescent PBAP in pristine-marine air masses, we grouped the WIBS data in 

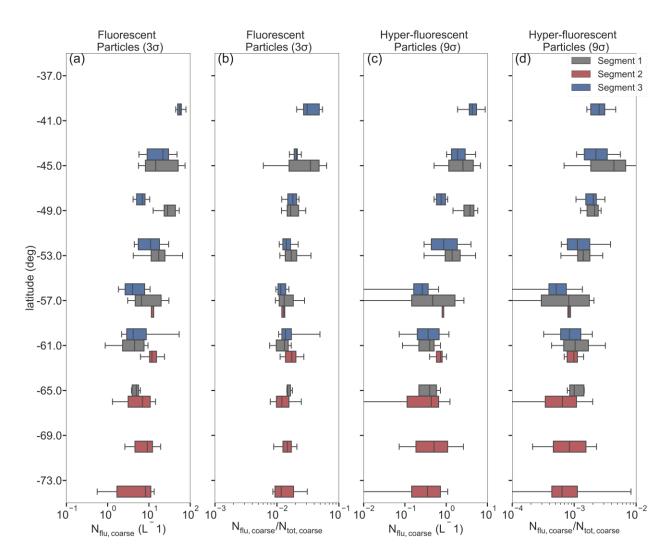
intervals of 4° latitude for each of the three pristine-marine segments of the cruise. Figure 6a and 

825 6b presents box and whisker plots of fluorescent coarse particle number concentrations and

fluorescent fractions (relative to total coarse mode number concentrations), and alike for hyper-826

fluorescent particles in Figure 6c and 6d. 827

828





830 Figure 6. (a) Variation coarse fluorescent number concentration and (b) fraction of coarse fluorescent number 831 concentration to total coarse aerosol number concentration for pristine-marine samples from different segments of the 832 campaign. (c) Variation of coarse hyper-fluorescent number concentration and (d) fraction of coarse hyper-fluorescent 833 number concentration to total coarse aerosol number concentration for pristine-marine samples. The boxes represent 834 IQR and the error bars are the 5<sup>th</sup> and 95<sup>th</sup> percentiles.

835



The median number concentrations ranged from 0.26 to 4.3  $L^{-1}$  and 4 to 56.6  $L^{-1}$  for 836 hyper-fluorescent (9 $\sigma$ ) and fluorescent particles (3 $\sigma$ ), respectively. Ovearall the particle number 837 concentrations decrease from North to South over the study area. At the same latitude the median 838 839 values for segment 3 are consistently smaller than for segment 1 (except for hyper-fluorescent

particles near 61 °S). This could be interpreted as a seasonal signal (since the segment 1
measurements were performed in January and the segment 3 measurements in March), or a
geographical signal, since the segment 1 measurements were performed in the Indian Ocean and
the segment 3 measurements were performed in the Atlantic Ocean.

For segment 2 a clear latitudinal trend could not be observed due to the small latitudinal range covered by the segment and the relatively broad IQRs. However, it can be noted that the ranges of median values for the hyper-fluorescent and fluorescent PBAP (0.35 to 0.8 L<sup>-1</sup> and 7 to 11.8 L<sup>-1</sup>, respectively) are not significantly smaller than the corresponding ranges for the most southern parts of the other cruise segments.

The median fluorescent particle fractions in the coarse size range for hyper-fluorescent and fluorescent particles ranged from 0.05 % to 0.43 % and 1.1 % to 3.4 %, respectively. The latitudinally-binned results shown in Figure 6 indicate that the median of (hyper-)fluorescent number fractions exhibit a weaker latitudinal trend compared to the absolute (hyper-)fluorescent particle number concentrations. The greater variability in the absolute concentrations is likely to simply reflect the different sea state and meteorological conditions affecting the SSA source flux, given the correlations observed in Figure 2.

856 Comparing to previous measurements of fluorescent particle number concentrations at 857 high southern latitudes, Crawford et al (2017) reported average fluorescent number 858 concentrations (based on a  $3\sigma$  threshold) of  $1.9 \pm 2.6$  L<sup>-1</sup> at the Halley VI Research Station in 859 Antarctica in austral summer 2015. This corresponded to average fluorescent particle number 860 fractions of 1.9 %. These values are comparable with the corresponding values reported in the 861 present study. However, it should be noted that sampling locations are quite different hindering 862 further detailed interpretations.

863

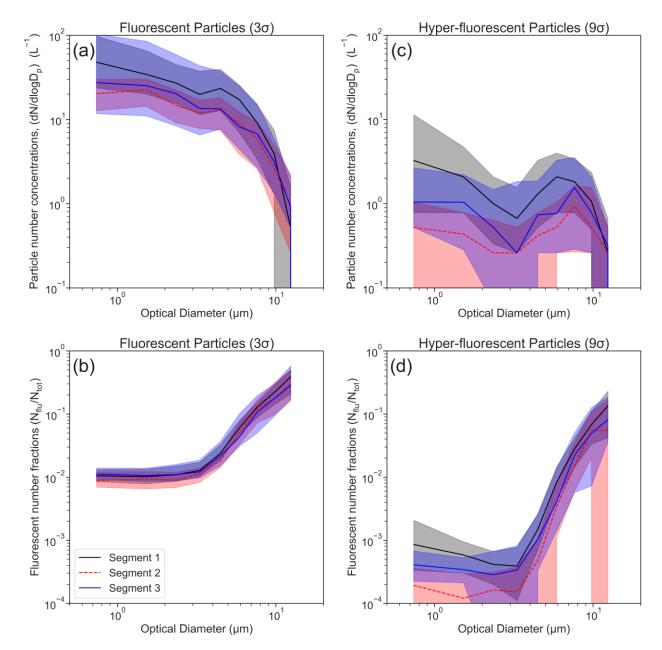
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3.8 Size and Asymmetry Factor (AF) distributions of Fluorescent Particles

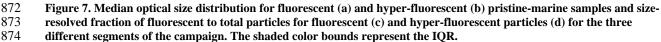
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Figure 7 shows the median (hyper-)fluorescent particle size distributions (PSD) along with the corresponding size-resolved (hyper-)fluorescent particle fractions for each of the three pristine-marine segments of the cruise.





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- 875

The size distribution trends are consistent for segments 1 to 3. Most notably, there is a 876 clear difference between the size-distributions of the fluorescent and hyper-fluorescent particles. 877 For the fluorescent particles the size-resolved concentration decreases continuously as the optical 878 879 diameter increases. The trend in the hyper-fluorescent number concentration indicates an initial

decrease leading to a minimum at  $\sim 3 \mu m$ , followed by a peak number concentration in the range 880 from 5 to 8 µm. This difference in the size distribution shapes suggests that particles larger than 881 ~3 µm particles emit stronger fluorescence signals compared to smaller particles. The correlation 882 analysis presented in Section 3.4 suggests that phytoplankton are main contributors to hyper-883 fluorescent PBAP, whereas bacteria are main contributors to fluorescent PBAP (both assessed 884 for coarse particles with optical diameter > 1  $\mu$ m). These correlation results are consistent with 885 the size distribution measurements: they suggest that relatively large phytoplankton – a dominant 886 contributor to hyper-fluorescent PBAP – constitute the mode in the hyper-fluorescent particles 887 size distributions observed between 5 and 8 µm, while the small phytoplankton, i.e. 888 prasinophytes, might be responsible for the signal  $< 3 \mu m$ . Secondly, bacteria, which have 889 generally smaller sizes, have a higher contribution in the fluorescent particle fraction, resulting in 890 891 higher absolute sub-micrometer than super-micrometer fluorescent particle concentrations. 892 The general trends of the size-resolved fluorescent particle fractions are similar for both fluorescent and hyper-fluorescent PBAP (Figure 7b & 7d). The contribution of fluorescent 893 894 particles in the size range between 0.5 and 3 µm is between 1 to 2%. For particle sizes above  $3 \,\mu\text{m}$ , the fractions increasing continuously reaching values of approximately 30 to 40 % for the 895 largest size bin (14 µm). In the case of the hyper-fluorescent particles, the fractions for the size 896 range from 0.5 to 3 µm are between 0.01 to 0.1%, followed by a significant increase to fractions 897 from ~ 5 to 13 % in the largest size bin. These results indicate that, over the SO, the relative 898 contribution of fluorescent particles to total particle number increases substantially with particle 899

size. This does not necessarily imply an increasing fraction of PBAP with increasing size of SSA
 particles, as the size dependence of the detected fluorescent particle fraction could be due to

<sup>902</sup> larger particles carrying greater quantities of fluorescent compounds.

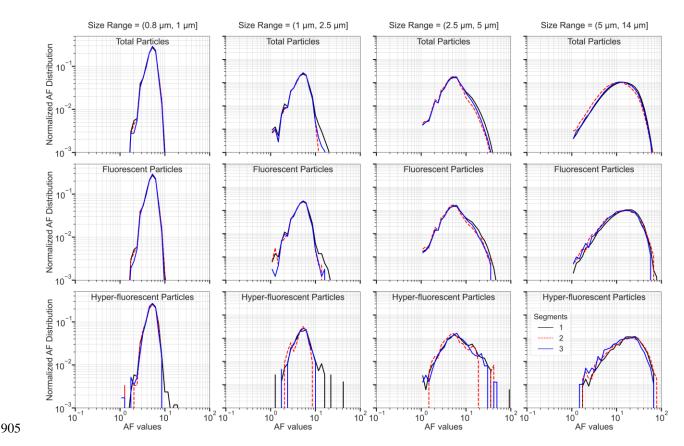


Figure 8. Probability density functions (PDFs) of asymmetry factor (AF) measurements for the pristine-marine samples
 from segment 1 (black line), segment 2 (red line) and segment 3 (blue line). Different columns represent different size
 ranges as indicated in the column titles, while the different rows represent distributions for total aerosol particles (top
 row), fluorescent particles (3σ) (middle row), and hyper-fluorescent particles (9σ) (bottom row).

In addition to the size distribution, WIBS also provides information about the shapes of 910 particles through the asymmetry factor (AF) measurements. Toprak & Schnaiter (2012) 911 previously showed that probability density functions (PDF) of spherical particles with different 912 sizes ranging from 1 to 2 µm peak at AF values between ~8 to 10. Figure 8 shows the normalized 913 914 PDFs of single particle AF values segregated by size and for total aerosol, fluorescent  $(3\sigma)$  and hyper-fluorescent (9 $\sigma$ ) particles. For particles smaller than 2.5  $\mu$ m the AF distributions are 915 unimodal with a peak at AF values of ~ 5. The AF distribution results are consistent across 916 (hyper-)fluorescent and total aerosol particles. This indicates that total aerosols and fluorescent 917 918 particles are essentially spherical in the sub 2.5 µm size range within the limits of the AF resolving power. Given that bacteria likely make an important contribution to the sub-2.5 µm 919 (hyper-)fluorescent particle fractions, we speculate that the apparent sphericity of these particles 920 could be due to either bacteria that possess sphere-like morphologies, or internal mixing of non-921

spherical bacteria with DOM components and/or sea salts within individual particles such that
the overall particles possess spherical shapes.

The AF PDFs for particles larger than 2.5 µm are also unimodal but with broader 924 distributions compared to the AF PDFs of the sub-2.5 µm particles. Similarly to the sub-2.5 µm 925 results, the shapes and widths of the AF PDFs are similar for all particles types (total, 926 fluorescent, and hyper-fluorescent) in the particle size ranges above 2.5 µm. The modes of the 927 AF PDFs are  $\sim$  5 for particles with diameters in the size range from 2.5 to 5 µm, while larger 928 929 mode values between 10 and 20 are observed for those particles with diameters in the size range 930 from 5 to 14 µm. The larger modes and increased widths for particles with diameters greater than 2.5 µm suggest that these particles are less spherical than the smaller particles. Considering the 931 large (hyper-)fluorescent fractions for super-3 µm particles (Figure 7b and d), one hypothesis is 932 that the constitutive marine microorganisms of larger PBAP particles are less spherical than the 933 934 microbes in smaller PBAP particles. On the other hand, the lack of difference between the AF PDFs for the total aerosol particles and for the (hyper-)fluorescent particles indicates that the 935 936 morphologies of (hyper-)fluorescent PBAPs are quite similar to the morphologies of the total aerosol particles, which are dominated by cubic-shaped SSA. That is, the AF results indicate that 937 the biological compounds embedded in PBAP (whether POM or DOM) do not have a major 938 influence on the shapes of SSA particles. 939

In regards to this discussion on particle shapes it should be noted that the AF 940 measurement applied in the WIBS is not a comprehensive nor sensitive method for investigating 941 942 particle morphology. For example, a previous laboratory study observed that WIBS-measured AF increases roughly linearly with increasing particle size for a range of different fluorescing 943 particle types (Savage et al., 2017). These authors were not able to determine if this trend was 944 real or an artefact of the WIBS AF measurement. Therefore, more robust methods (e.g. electron 945 microscopy) should be performed on marine PBAP to further investigate the trends observed in 946 the present study as well as other morphological properties of these aerosols. 947

### 948 4 Conclusions

In this study we presented a comprehensive dataset of fluorescent aerosol particle measurements over vast regions of the Southern Ocean (SO). In our analysis we focused on coarse particles (optical diameter > 1  $\mu$ m) and separated the data into two categories: samples acquired further than 200 km from any land mass (pristine-marine samples), and samples collected within 200 km from any land mass (terrestrially-influenced samples). Furthermore, we used two different instrument fluorescent thresholds ( $3\sigma$  and  $9\sigma$ ) to identify both fluorescent and hyper-fluorescent particles. The median fluorescent particle number concentrations for the pristine-marine and terrestrially-influenced influenced samples were 11 L<sup>-1</sup> and 16.6 L<sup>-1</sup>, respectively, while the median hyper-fluorescent particle number concentration for pristinemarine and terrestrially-influenced samples were 0.87 L<sup>-1</sup> and 1.47 L<sup>-1</sup>, respectively.

959 To investigate the relationship between (hyper-)fluorescent PBAP and SSA a correlation 960 analysis was conducted with four different proxy variables for SSA concentrations (wind speed, total coarse mode particle concentration, Cl- and Na+ concentrations). Moderately high 961 correlations were observed between pristine-marine (hyper-)fluorescent PBAP number 962 concentrations and the SSA proxy variables (e.g. Pearson's R values of 0.76 and 0.61 were 963 964 obtained between total coarse particle number concentrations and fluorescent and hyperfluorescent particle number concentrations, respectively). For all four SSA proxy variables, 965 966 lower correlation values were obtained for the terrestrially-influenced samples relative to the pristine-marine samples due to existence of outlying measurements that we attribute to potential 967 terrestrial PBAP sources. These results support the hypothesis that SSA is the main source of 968 fluorescent PBAP in pristine marine environments, while also demonstrating the importance of 969 970 fully isolating pristine-marine from terrestrially-influenced PBAP measurements in order to 971 study them.

Given the high correlation between total and fluorescent particle number concentrations for the pristine-marine samples, we calculated that fluorescent PBAP represent 1.6% (median value) of the total number of coarse aerosol particles over the pristine SO, while hyperfluorescent PBAP represent 0.13% (median value) of the same total. Assuming that in the pristine SO atmosphere SSA is the only significant source of coarse aerosols (on a number basis), these fractions provide a useful means for estimating PBAP number concentrations using measured or modelled SSA number concentrations.

To identify the potential marine sources that modulate fluorescent PBAP concentrations we conducted further correlation analysis with the (hyper-)fluorescent particle fractions and thirty different marine variables measured in seawater. The results indicated that for pristine-marine

samples, fluorescent particles correlated best with the number concentrations of marine bacteria 982 (Pearson's R = 0.4 - 0.5), while hyper-fluorescent particles correlated best with mass 983 concentrations of several different phytoplankton taxa (Pearson's R = 0.4-0.7). In this correlation 984 analysis the terrestrially-influenced samples also had systematically lower correlation 985 coefficients compared to the pristine-marine samples, confirming that the terrestrially-influenced 986 samples are likely influenced by non-marine sources. Overall, the two correlation analyses 987 indicate that the PBAP source flux in the pristine SO is primarily driven by the SSA source flux, 988 with further modulation by seawater concentrations of marine biota such as bacteria and 989 phytoplankton. 990

To gain insight into the fluorescence characteristics of the measured PBAP, we classified 991 the WIBS measurements using the ABC fluorescence classification scheme. The fluorescence 992 class compositions for the three pristine-marine segments of the cruise were relatively consistent, 993 994 which suggests that the sources of pristine marine PBAP were relatively homogenous across all 995 sectors of the SO. In contrast, much more variability was observed between the fluorescence 996 class compositions of nine near-land events, which indicates greater diversity in the terrestrial sources of PBAP that contributed to these events. This is not surprising since these events 997 occurred in a wide variety of different environments, including near the Antarctic coast, pristine 998 SO islands, populated continental regions, and even the Saharan desert (the latter occurring 999 1000 during the ship's return voyage back to Europe). The fluorescence class composition of the 1001 Saharan dust event was particularly unique (prominent contribution of type BC particles), which suggests that the long-range transport of dust particles with fluorescence signatures like those of 1002 Saharan dust did not contribute substantially to the SO measurements performed during the ACE 1003 1004 campaign.

In addition to the ABC classification scheme, we investigated a complementary approach for characterizing aerosol fluorescence properties based on the ratio of fluorescent intensities in channels B and A of the WIBS instrument (termed the  $R_{B2A}$  parameter). The  $R_{B2A}$  results were generally consistent with the ABC classification results: the  $R_{B2A}$  distributions for the three pristine-marine segments of the cruise were similar while the distributions for the nine near-land events were much more variable. The highest median  $R_{B2A}$  value was observed for the Siple Island event, which suggests a greater contributions of humic-like fluorescing matters to the 1012 particles comprising this event. The lowest median  $R_{B2A}$  value was observed during the Hobart 1013 event, suggesting greater contributions from protein-like organic matters during this event.

1014 Finally, we summarized the latitudinal variations in (hyper-)fluorescent particle 1015 concentrations and fractions, as well as the (hyper-)fluorescent particle size and shape parameter (asymmetry factor) distributions. These summaries aim to provide a useful point of comparison 1016 for future studies of marine PBAP over the SO as well as other oceanic regions. Of particular 1017 interest are the size distribution results, which indicates that while the concentrations of 1018 1019 fluorescent particles decreased monotonically from small to large particle diameters, the hyper-1020 fluorescent particle number size distributions contained a mode between 5 and 7  $\mu$ m. We suggest that this size distribution mode is associated with the phytoplankton taxa that were observed to 1021 correlate highly with the fractions of hyper-fluorescent PBAP. 1022

1023

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#### 1038 Data availability statement

- 1039 The dataset used in this study are available in (1) Antoine et al. (2019) available at
- 1040 <u>https://doi.org/10.5281/zenodo.3406983;</u> (2) Chen et al. (2019) available at
- 1041 <u>https://doi.org/10.5281/zenodo.3559982</u>; (3) Landwehr, Thomas, et al. (2020) available at
- 1042 <u>https://doi.org/10.5281/zenodo.3836439</u>; (4) Landwehr, Thurnherr, et al. (2020) available at
- 1043 <u>https://doi.org/10.5194/amt-13-3487-2020;</u> (5) Schmale, Henning, et al. (2019) available at
- 1044 <u>https://doi.org/10.5281/zenodo.2636709</u>; (6) Tatzelt et al. (2020) available at
- 1045 <u>https://doi.org/10.5281/zenodo.3922147;</u> (7) Thomalla et al., (2020) available at
- 1046 <u>https://doi.org/10.5281/zenodo.3859515;</u> (8) Thurnherr et al. (2020) available at
- 1047 https://doi.org/10.5281/zenodo.4031705; The archiving of the fluorescent aerosol and gel-like
- 1048 POM measurments are ongoing. Currently these data have been uploaded as supporting
- 1049 information. They will be uploaded to a Zenodo repository if the paper is accepted for
- 1050 publication.

#### 1051 Author contributions

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### manuscript submitted to Journal of Geophysical Research

| 1062         | References   |
|--------------|--|
| 1063         | Aller, J. Y., Radway, J. C., Kilthau, W. P., Bothe, D. W., Wilson, T. W., Vaillancourt, R. D., et al. (2017). Size-        |
| 1064         | resolved characterization of the polysaccharidic and proteinaceous components of sea spray aerosol.                        |
| 1065         | Atmospheric Environment, 154, 331-347. https://doi.org/10.1016/j.atmosenv.2017.01.053                                      |
| 1066         | Antoine, D., Thomalla, S., Berliner, D., Little, H., Moutier, W., Olivier-Morgan, A., et al. (2019). Phytoplankton         |
| 1067         | pigment concentrations of seawater sampled during the Antarctic Circumnavigation Expedition (ACE)                          |
| 1068         | during the Austral Summer of 2016/2017. (Version 1.0) [Data set]. Zenodo.  |
| 1069         | https://doi.org/10.5281/zenodo.3406983   |
| 1070<br>1071 | Ault, A. P., Moffet, R. C., Baltrusaitis, J., Collins, D. B., Ruppel, M. J., Cuadra-Rodriguez, L. A., et al. (2013). Size- |
| 1072         | Dependent Changes in Sea Spray Aerosol Composition and Properties with Different Seawater Conditions.                      |
| 1073         | Environmental Science & Technology, 47(11), 5603-5612. https://doi.org/10.1021/es400416g                                   |
| 1074         | Bates, T. S., Quinn, P. K., Coffman, D., Schulz, K., Covert, D. S., Johnson, J. E., et al. (2008). Boundary layer          |
| 1075         | aerosol chemistry during TexAQS/GoMACCS 2006: Insights into aerosol sources and transformation                             |
| 1076         | processes. Journal of Geophysical Research: Atmospheres, 113(D7).  |
| 1077         | https://doi.org/10.1029/2008JD010023   |
| 1078         | Bigg, E. K. (1973). Ice Nucleus Concentrations in Remote Areas. Journal of the Atmospheric Sciences, 30(6), 1153-          |
| 1079         | 1157. https://doi.org/10.1175/1520-0469(1973)030<1153:INCIRA>2.0.CO;2  |
| 1080         | Brooks, S. D., & Thornton, D. C. O. (2018). Marine Aerosols and Clouds. Annual Review of Marine Science, 10(1),            |
| 1081         | 289-313. https://doi.org/10.1146/annurev-marine-121916-063148  |
| 1082         | Burrows, S. M., Elbert, W., Lawrence, M. G., & Poschl, U. (2009). Bacteria in the global atmosphere - Part 1:              |
| 1083         | Review and synthesis of literature data for different ecosystems. Atmos. Chem. Phys., 18.                                  |
| 1084         | Ceburnis, D., Masalaite, A., Ovadnevaite, J., Garbaras, A., Remeikis, V., Maenhaut, W., et al. (2016). Stable              |
| 1085         | isotopes measurements reveal dual carbon pools contributing to organic matter enrichment in marine                         |
| 1086         | aerosol. Scientific Reports, 6(1), 36675. https://doi.org/10.1038/srep36675  |
| 1087         | Chen, G., Schmale, J., & Landwehr, S. (2019). Non-refractory particulate sulfate and chloride data from a time of          |
| 1088         | flight aerosol chemical speciation monitor around the Southern Ocean in the austral summer of 2016/17,                     |
| 1089         | during the Antarctic Circumnavigation Expedition (ACE). (Version 1.0) [Data set]. Zenodo.                                  |
| 1090         | https://doi.org/10.5281/zenodo.3559982   |
|              |  |

- 1091 Chen, Q., Miyazaki, Y., Kawamura, K., Matsumoto, K., Coburn, S., Volkamer, R., et al. (2016). Characterization of
- 1092 Chromophoric Water-Soluble Organic Matter in Urban, Forest, and Marine Aerosols by HR-ToF-AMS
- 1093 Analysis and Excitation–Emission Matrix Spectroscopy. *Environmental Science & Technology*, 50(19),
- 1094 10351–10360. https://doi.org/10.1021/acs.est.6b01643
- 1095 Crawford, I., Lloyd, G., Herrmann, E., Hoyle, C. R., Bower, K. N., Connolly, P. J., et al. (2016). Observations of
- 1096 fluorescent aerosol–cloud interactions in the free troposphere at the High-Altitude Research 1097 Station Jungfraujoch. *Atmospheric Chemistry and Physics*, *16*(4), 2273–2284. https://doi.org/10.5194/acp-
- 1098 16-2273-2016
- Crawford, I., Gallagher, M. W., Bower, K. N., Choularton, T. W., Flynn, M. J., Ruske, S., et al. (2017). Real-time
  detection of airborne fluorescent bioparticles in Antarctica. *Atmospheric Chemistry and Physics*, *17*(23),
  14291–14307. https://doi.org/10.5194/acp-17-14291-2017
- DeMott, P. J., Prenni, A. J., Liu, X., Kreidenweis, S. M., Petters, M. D., Twohy, C. H., et al. (2010). Predicting
  global atmospheric ice nuclei distributions and their impacts on climate. *Proceedings of the National Academy of Sciences*, *107*(25), 11217–11222. https://doi.org/10.1073/pnas.0910818107
- DeMott, P. J., Hill, T. C. J., McCluskey, C. S., Prather, K. A., Collins, D. B., Sullivan, R. C., et al. (2016). Sea spray
  aerosol as a unique source of ice nucleating particles. *Proceedings of the National Academy of Sciences*,
- 1107 *113*(21), 5797–5803. https://doi.org/10.1073/pnas.1514034112
- 1108 Després, V. R., Huffman, J. A., Burrows, S. M., Hoose, C., Safatov, A. S., Buryak, G., et al. (2012). Primary
- 1109 biological aerosol particles in the atmosphere: a review. *Tellus B: Chemical and Physical Meteorology*,
- 1110 64(1), 15598. https://doi.org/10.3402/tellusb.v64i0.15598
- Fennelly, M. J., Sewell, G., Prentice, M. B., O'Connor, D. J., & Sodeau, J. R. (2018). Review: The Use of RealTime Fluorescence Instrumentation to Monitor Ambient Primary Biological Aerosol Particles (PBAP). *Atmosphere*, 9(1), 1. https://doi.org/10.3390/atmos9010001
- 1114 Flato, G. M., Marotzke, J., Abiodun, B., Braconnot, P., Chou, S. C., Collins, W., et al. (2013). Evaluation of climate
- 1115 models. Climate Change 2013 The Physical Science Basis: Working Group I Contribution to the Fifth
- 1116 Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press.
- 1117 Retrieved from https://hal.archives-ouvertes.fr/hal-01644494

- 1118 Fröhlich-Nowoisky, J., Kampf, C. J., Weber, B., Huffman, J. A., Pöhlker, C., Andreae, M. O., et al. (2016).
- Bioaerosols in the Earth system: Climate, health, and ecosystem interactions. *Atmospheric Research*, *182*,
  346–376. https://doi.org/10.1016/j.atmosres.2016.07.018
- Fu, P., Kawamura, K., Chen, J., Qin, M., Ren, L., Sun, Y., et al. (2015). Fluorescent water-soluble organic aerosols
  in the High Arctic atmosphere. *Scientific Reports*, 5(1). https://doi.org/10.1038/srep09845
- 1123 Gove, J. M., McManus, M. A., Neuheimer, A. B., Polovina, J. J., Drazen, J. C., Smith, C. R., et al. (2016). Near-

1124 island biological hotspots in barren ocean basins. *Nature Communications*, 7(1), 10581.
1125 https://doi.org/10.1038/ncomms10581

- 1126 Hamilton, D. S., Lee, L. A., Pringle, K. J., Reddington, C. L., Spracklen, D. V., & Carslaw, K. S. (2014).
- 1127Occurrence of pristine aerosol environments on a polluted planet. Proceedings of the National Academy of1128Sciences. https://doi.org/10.1073/pnas.1415440111
- Hawkins, L. N., & Russell, L. M. (2010). Polysaccharides, Proteins, and Phytoplankton Fragments: Four Chemically
   Distinct Types of Marine Primary Organic Aerosol Classified by Single Particle Spectromicroscopy.
- 1131 Advances in Meteorology, 2010, 1–14. https://doi.org/10.1155/2010/612132
- 1132 Healy, D. A., O'Connor, D. J., & Sodeau, J. R. (2012). Measurement of the particle counting efficiency of the
- 1133 "Waveband Integrated Bioaerosol Sensor" model number 4 (WIBS-4). Journal of Aerosol Science, 47, 94–
- 1134 99. https://doi.org/10.1016/j.jaerosci.2012.01.003
- 1135 Healy, D. A., Huffman, J. A., O'Connor, D. J., Pöhlker, C., Pöschl, U., & Sodeau, J. R. (2014). Ambient
- 1136 measurements of biological aerosol particles near Killarney, Ireland: a comparison between real-time
- fluorescence and microscopy techniques. *Atmospheric Chemistry and Physics*, 14(15), 8055–8069.
- 1138 https://doi.org/10.5194/acp-14-8055-2014
- 1139 Hernandez, M., Perring, A. E., McCabe, K., Kok, G., Granger, G., & Baumgardner, D. (2016). Chamber catalogues
- of optical and fluorescent signatures distinguishbioaerosol classes. *Atmospheric Measurement Techniques*,
  9(7), 3283–3292. https://doi.org/10.5194/amt-9-3283-2016
- 1142 Hervàs, A., Camarero, L., Reche, I., & Casamayor, E. O. (2009). Viability and potential for immigration of airborne
- bacteria from Africa that reach high mountain lakes in Europe. *Environmental Microbiology*, 11(6), 1612–
- 1144 1623. https://doi.org/10.1111/j.1462-2920.2009.01926.x

- 1145 Kanji, Z. A., Ladino, L. A., Wex, H., Boose, Y., Burkert-Kohn, M., Cziczo, D. J., & Krämer, M. (2017). Overview
- 1146 of Ice Nucleating Particles. *Meteorological Monographs*, 58, 1.1-1.33.
- 1147 https://doi.org/10.1175/AMSMONOGRAPHS-D-16-0006.1
- 1148 Kaye, P. H., Stanley, W. R., Hirst, E., Foot, E. V., Baxter, K. L., & Barrington, S. J. (2005). Single particle
- 1149 multichannel bio-aerosol fluorescence sensor. *Optics Express*, *13*(10), 3583.
- 1150 https://doi.org/10.1364/OPEX.13.003583
- Kellogg, C. A., & Griffin, D. W. (2006). Aerobiology and the global transport of desert dust. *Trends in Ecology & Evolution*, 21(11), 638–644. https://doi.org/10.1016/j.tree.2006.07.004
- 1153 Landwehr, S., Thomas, J., & Schmale, J. (2020). Five-minute average horizontal wind velocity data combined from
- 1154 both sensors (which has been corrected for air-flow distortion) from the Antarctic Circumnavigation
- 1155 Expedition (ACE) 2016/2017 legs 0 to 4. [Data set]. (Version 1.0) Zenodo.
- 1156 https://doi.org/10.5281/zenodo.3836439
- Landwehr, S., Thurnherr, I., Cassar, N., Gysel-Beer, M., & Schmale, J. (2020). Using global reanalysis data to
   quantify and correct airflow distortion bias in shipborne wind speed measurements. Atmospheric

1159 Measurement Techniques, 13(6), 3487–3506. https://doi.org/10.5194/amt-13-3487-2020

- 1160 Lee, H. D., Morris, H. S., Laskina, O., Sultana, C. M., Lee, C., Jayarathne, T., et al. (2020). Organic Enrichment,
- 1161 Physical Phase State, and Surface Tension Depression of Nascent Core–Shell Sea Spray Aerosols during
- 1162 Two Phytoplankton Blooms. *ACS Earth and Space Chemistry*, 4(4), 650–660.
- 1163 https://doi.org/10.1021/acsearthspacechem.0c00032
- de Leeuw, G., Andreas, E. L., Anguelova, M. D., Fairall, C. W., Lewis, E. R., O'Dowd, C., et al. (2011). Production
  flux of sea spray aerosol. *Reviews of Geophysics*, 49(2). https://doi.org/10.1029/2010RG000349
- 1166 Lewis, E. R., & Schwartz, S. E. (2004). Sea Salt Aerosol Production Fluxes: Estimates and Critical Analysis. In Sea
- 1167 Salt Aerosol Production: Mechanisms, Methods, Measurements and Models (pp. 299–344). American
- 1168 Geophysical Union (AGU). https://doi.org/10.1002/9781118666050.ch5
- 1169 Mayol, E., Arrieta, J. M., Jiménez, M. A., Martínez-Asensio, A., Garcias-Bonet, N., Dachs, J., et al. (2017). Long-
- 1170 range transport of airborne microbes over the global tropical and subtropical ocean. *Nature*
- 1171 *Communications*, 8(1), 1–9. https://doi.org/10.1038/s41467-017-00110-9

- 1172 McCluskey, C. S., Hill, T. C. J., Humphries, R. S., Rauker, A. M., Moreau, S., Strutton, P. G., et al. (2018).
- 1173 Observations of Ice Nucleating Particles Over Southern Ocean Waters. *Geophysical Research Letters*,
- 1174 *45*(21), 11,989-11,997. https://doi.org/10.1029/2018GL079981
- 1175 McFarquhar, G. M., Bretherton, C., Marchand, R., Protat, A., DeMott, P. J., Alexander, S. P., et al. (2020).
- 1176 Observations of clouds, aerosols, precipitation, and surface radiation over the Southern Ocean: An
- 1177 overview of CAPRICORN, MARCUS, MICRE and SOCRATES. Bulletin of the American Meteorological
- 1178 Society, 1(aop), 1–92. https://doi.org/10.1175/BAMS-D-20-0132.1
- Middlebrook, A. M., Murphy, D. M., & Thomson, D. S. (1998). Observations of organic material in individual
  marine particles at Cape Grim during the First Aerosol Characterization Experiment (ACE 1). *Journal of*
- 1181 *Geophysical Research: Atmospheres*, *103*(D13), 16475–16483. https://doi.org/10.1029/97JD03719
- 1182 Modini, R. L., Frossard, A. A., Ahlm, L., Russell, L. M., Corrigan, C. E., Roberts, G. C., et al. (2015). Primary
- 1183 marine aerosol-cloud interactions off the coast of California. *Journal of Geophysical Research:*
- 1184 Atmospheres, 120(9), 4282–4303. https://doi.org/10.1002/2014JD022963
- 1185 Monahan, E. C., Spiel, D. E., & Davidson, K. L. (1986). A Model of Marine Aerosol Generation Via Whitecaps and
- 1186 Wave Disruption. In Edward C. Monahan & G. M. Niocaill (Eds.), Oceanic Whitecaps: And Their Role in
- 1187 Air-Sea Exchange Processes (pp. 167–174). Dordrecht: Springer Netherlands. https://doi.org/10.1007/9781188 94-009-4668-2\_16
- 1189 Murphy, D. M., Thomson, D. S., Middlebrook, A. M., & Schein, M. E. (1998). In situ single-particle
- 1190 characterization at Cape Grim. *Journal of Geophysical Research: Atmospheres*, *103*(D13), 16485–16491.
- 1191 https://doi.org/10.1029/97JD03281
- 1192 O'Dowd, C. D., & de Leeuw, G. (2007). Marine aerosol production: a review of the current knowledge.
- 1193 Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences,
- 1194 365(1856), 1753–1774. https://doi.org/10.1098/rsta.2007.2043
- 1195 Orellana, M. V., Matrai, P. A., Leck, C., Rauschenberg, C. D., Lee, A. M., & Coz, E. (2011). Marine microgels as a
- 1196 source of cloud condensation nuclei in the high Arctic. *Proceedings of the National Academy of Sciences*,
- 1197 *108*(33), 13612–13617. https://doi.org/10.1073/pnas.1102457108

- 1198 Perring, A. E., Schwarz, J. P., Baumgardner, D., Hernandez, M. T., Spracklen, D. V., Heald, C. L., et al. (2015).
- 1199Airborne observations of regional variation in fluorescent aerosol across the United States. Journal of1200Geophysical Research: Atmospheres, 120(3), 1153–1170. https://doi.org/10.1002/2014JD022495
- Pöhlker, C., Huffman, J. A., & Pöschl, U. (2012). Autofluorescence of atmospheric bioaerosols fluorescent
  biomolecules and potential interferences. *Atmospheric Measurement Techniques*, 5(1), 37–71.
- 1203 https://doi.org/10.5194/amt-5-37-2012
- Pope, F. D. (2010). Pollen grains are efficient cloud condensation nuclei. *Environmental Research Letters*, 5(4),
  044015. https://doi.org/10.1088/1748-9326/5/4/044015
- Prather, K. A., Bertram, T. H., Grassian, V. H., Deane, G. B., Stokes, M. D., DeMott, P. J., et al. (2013). Bringing
   the ocean into the laboratory to probe the chemical complexity of sea spray aerosol. *Proceedings of the National Academy of Sciences*, *110*(19), 7550–7555. https://doi.org/10.1073/pnas.1300262110
- Quinn, P. K., Coffman, D. J., Johnson, J. E., Upchurch, L. M., & Bates, T. S. (2017). Small fraction of marine cloud
  condensation nuclei made up of sea spray aerosol. *Nature Geoscience*, *10*(9), 674–679.
- 1211 https://doi.org/10.1038/ngeo3003
- 1212 Quinn, P. K., Collins, D. B., Grassian, V. H., Prather, K. A., & Bates, T. S. (2015). Chemistry and Related

1213 Properties of Freshly Emitted Sea Spray Aerosol. *Chemical Reviews*, 115(10), 4383–4399.

- 1214 https://doi.org/10.1021/cr500713g
- Russell, L. M., Hawkins, L. N., Frossard, A. A., Quinn, P. K., & Bates, T. S. (2010). Carbohydrate-like composition
  of submicron atmospheric particles and their production from ocean bubble bursting. *Proceedings of the*
- 1217 National Academy of Sciences, 107(15), 6652–6657. https://doi.org/10.1073/pnas.0908905107
- 1218 Savage, N. J., Krentz, C. E., Könemann, T., Han, T. T., Mainelis, G., Pöhlker, C., & Huffman, J. A. (2017).
- 1219 Systematic characterization and fluorescence threshold strategies for the wideband integrated bioaerosol
- 1220 sensor (WIBS) using size-resolved biological and interfering particles. Atmospheric Measurement
- 1221 Techniques, 10(11), 4279–4302. https://doi.org/10.5194/amt-10-4279-2017
- Schmale, J., Henning, S., Tummon, F., Hartmann, M., Baccarini, A., Welti, A., et al. (2019). Coarse mode aerosol
   particle size distribution collected in the Southern Ocean in the austral summer of 2016/2017, during the
   Antarctic Circumnavigation Expedition. [Data set]. Zenodo. https://doi.org/10.5281/zenodo.2636709

- 1225 Schmale, J., Baccarini, A., Thurnherr, I., Henning, S., Efraim, A., Regayre, L., et al. (2019). Overview of the
- 1226 Antarctic Circumnavigation Expedition: Study of Preindustrial-like Aerosols and Their Climate Effects
- 1227 (ACE-SPACE). Bulletin of the American Meteorological Society, 100(11), 2260–2283.
- 1228 https://doi.org/10.1175/BAMS-D-18-0187.1
- 1229 Schnell, R. C., & Vali, G. (1976). Biogenic Ice Nuclei: Part I. Terrestrial and Marine Sources. *Journal of the*
- 1230 Atmospheric Sciences, 33(8), 1554–1564. https://doi.org/10.1175/1520-
- 1231 0469(1976)033<1554:BINPIT>2.0.CO;2
- Sprenger, M., & Wernli, H. (2015). The LAGRANTO Lagrangian analysis tool version 2.0. Geoscientific Model
   Development, 8(8), 2569–2586. https://doi.org/10.5194/gmd-8-2569-2015
- 1234 Tatzelt, C., Henning, S., Tummon, F., Hartmann, M., Baccarini, A., Welti, A., et al. (2020). Ionic composition of
- 1235 particulate matter (PM10) from high-volume sampling over the Southern Ocean during the austral summer
- 1236 of 2016/2017 on board the Antarctic Circumnavigation Expedition (ACE). [Data set]. Zenodo.
- 1237 https://doi.org/10.5281/zenodo.3922147
- Taylor, P. E., Flagan, R. C., Miguel, A. G., Valenta, R., & Glovsky, M. M. (2004). Birch pollen rupture and the
   release of aerosols of respirable allergens. *Clinical & Experimental Allergy*, *34*(10), 1591–1596.
- 1240 https://doi.org/10.1111/j.1365-2222.2004.02078.x
- 1241 Thomalla, S., Antoine, D., Berliner, D., Little, H., Moutier, W., Olivier-Morgan, A., et al. (2020). Particulate
- 1242 organic carbon and particulate organic nitrogen concentrations and stable isotope composition of seawater
- 1243 sampled during the Antarctic Circumnavigation Expedition (ACE) during the Austral Summer of
- 1244 2016/2017. (Version 1.0) [Data set]. Zenodo. https://doi.org/10.5281/zenodo.3859515
- 1245 Thurnherr, I., Wernli, H., & Aemisegger, F. (2020). 10-day backward trajectories from ECMWF analysis data along
- 1246 the ship track of the Antarctic Circumnavigation Expedition in austral summer 2016/2017. (Version 1.0)
- 1247 [Data set]. Zenodo. https://doi.org/10.5281/zenodo.4031705Tobo, Y., Prenni, A. J., DeMott, P. J.,
- 1248 Huffman, J. A., McCluskey, C. S., Tian, G., et al. (2013). Biological aerosol particles as a key determinant
- 1249 of ice nuclei populations in a forest ecosystem. Journal of Geophysical Research: Atmospheres, 118(17),
- 1250 10,100-10,110. https://doi.org/10.1002/jgrd.50801

- Toprak, E., & Schnaiter, M. (2012). Fluorescent biological aerosol particles (FBAPs) measured with the Waveband
   Integrated Bioaerosol Sensor WIBS-4: laboratory tests combined with a one year field study. *Atmospheric*
- 1253 *Chemistry and Physics Discussions*, *12*(7), 17607–17656. https://doi.org/10.5194/acpd-12-17607-2012
- 1254 Uetake, J., Hill, T. C. J., Moore, K. A., DeMott, P. J., Protat, A., & Kreidenweis, S. M. (2020). Airborne bacteria
- 1255 confirm the pristine nature of the Southern Ocean boundary layer. *Proceedings of the National Academy of* 1256 *Sciences*, *117*(24), 13275–13282. https://doi.org/10.1073/pnas.2000134117

Vergara-Temprado, J., Murray, B. J., Wilson, T. W., O'Sullivan, D., Browse, J., Pringle, K. J., et al. (2017).

- 1258 Contribution of feldspar and marine organic aerosols to global ice nucleating particle concentrations.
- 1259 Atmospheric Chemistry and Physics, 17(5), 3637–3658. https://doi.org/10.5194/acp-17-3637-2017
- 1260 Vergara-Temprado, J., Miltenberger, A. K., Furtado, K., Grosvenor, D. P., Shipway, B. J., Hill, A. A., et al. (2018).
- Strong control of Southern Ocean cloud reflectivity by ice-nucleating particles. *Proceedings of the National Academy of Sciences*, *115*(11), 2687–2692. https://doi.org/10.1073/pnas.1721627115
- Walton, D. W. H., & Thomas, J. (2018). Cruise Report Antarctic Circumnavigation Expedition (ACE) 20th
   December 2016 19th March 2017. Zenodo. https://doi.org/10.5281/zenodo.1443511
- Wang, X., Sultana, C. M., Trueblood, J., Hill, T. C. J., Malfatti, F., Lee, C., et al. (2015). Microbial Control of Sea
  Spray Aerosol Composition: A Tale of Two Blooms. *ACS Central Science*, 1(3), 124–131.
- 1267 https://doi.org/10.1021/acscentsci.5b00148
- Wilbourn, E. K., Thornton, D. C. O., Ott, C., Graff, J., Quinn, P. K., Bates, T. S., et al. (2020). Ice Nucleation by
  Marine Aerosols Over the North Atlantic Ocean in Late Spring. *Journal of Geophysical Research:*
- 1270 Atmospheres, 125(4), e2019JD030913. https://doi.org/10.1029/2019JD030913
- 1271 Wilson, T. W., Ladino, L. A., Alpert, P. A., Breckels, M. N., Brooks, I. M., Browse, J., et al. (2015). A marine
- 1272 biogenic source of atmospheric ice-nucleating particles. *Nature*, *525*(7568), 234–238.
- 1273 https://doi.org/10.1038/nature14986
- 1274 Young, I. R. (1999). Seasonal variability of the global ocean wind and wave climate. *International Journal of*
- 1275 Climatology, 19(9), 931–950. https://doi.org/10.1002/(SICI)1097-0088(199907)19:9<931::AID-
- 1276 JOC412>3.0.CO;2-O

- 1277 Zamanillo, M., Ortega-Retuerta, E., Nunes, S., Estrada, M., Sala, M. M., Royer, S.-J., et al. (2019). Distribution of
- transparent exopolymer particles (TEP) in distinct regions of the Southern Ocean. Science of The Total
   *Environment*, 691, 736–748. https://doi.org/10.1016/j.scitotenv.2019.06.524
- 1280 Ziemba, L. D., Beyersdorf, A. J., Chen, G., Corr, C. A., Crumeyrolle, S. N., Diskin, G., et al. (2016). Airborne
- 1281 observations of bioaerosol over the Southeast United States using a Wideband Integrated Bioaerosol
- 1282 Sensor: Airborne Observations of Bioaerosol. Journal of Geophysical Research: Atmospheres, 121(14),
- 1283 8506–8524. https://doi.org/10.1002/2015JD024669
- 1284 Zsolnay, A., Baigar, E., Jimenez, M., Steinweg, B., & Saccomandi, F. (1999). Differentiating with fluorescence
- spectroscopy the sources of dissolved organic matter in soils subjected to drying. *Chemosphere*, 38(1), 45–
- 1286 50. https://doi.org/10.1016/S0045-6535(98)00166-0

#### 1287 References From the Supporting Information

- 1288 Antoine, D., Thomalla, S., Berliner, D., Little, H., Moutier, W., Olivier-Morgan, A., et al. (2019). Phytoplankton
- 1289 pigment concentrations of seawater sampled during the Antarctic Circumnavigation Expedition (ACE)

1290 during the Austral Summer of 2016/2017. (Version 1.0) [Data set]. Zenodo.

- 1291 https://doi.org/10.5281/zenodo.3406983
- Cassar, N., Wright, S. W., Thomson, P. G., Trull, T. W., Westwood, K. J., Salas, M. de, et al. (2015). The relation of
   mixed-layer net community production to phytoplankton community composition in the Southern Ocean.
- 1294 Global Biogeochemical Cycles, 29(4), 446–462. https://doi.org/10.1002/2014GB004936
- 1295 Clementson, L. A., Parslow, J. S., Turnbull, A. R., & Bonham, P. I. (2004). Properties of light absorption in a highly
- 1296 coloured estuarine system in south-east Australia which is prone to blooms of the toxic dinoflagellate
- 1297 Gymnodinium catenatum. Estuarine, Coastal and Shelf Science, 60(1), 101–112.
- 1298 https://doi.org/10.1016/j.ecss.2003.11.022
- 1299 Cook, S. S., Whittock, L., Wright, S. W., & Hallegraeff, G. M. (2011). Photosynthetic Pigment and Genetic
- 1300 Differences Between Two Southern Ocean Morphotypes of Emiliania Huxleyi (haptophyta)1. Journal of
- 1301 Phycology, 47(3), 615–626. https://doi.org/10.1111/j.1529-8817.2011.00992.x
- 1302 Gabey, A. M., Gallagher, M. W., Whitehead, J., Dorsey, J. R., Kaye, P. H., & Stanley, W. R. (2010). Measurements

1303 and comparison of primary biological aerosol above and below a tropical forest canopy using a dual

manuscript submitted to Journal of Geophysical Research

- 1304 channel fluorescence spectrometer. Atmospheric Chemistry and Physics, 10(10), 4453–4466.
- 1305 https://doi.org/10.5194/acp-10-4453-2010
- 1306 Mackey, M. D., Mackey, D. J., Higgins, H. W., & Wright, S. W. (1996). CHEMTAX a program for estimating
- class abundances from chemical markers: application to HPLC measurements of phytoplankton. Marine
   Ecology Progress Series, 144, 265–283. https://doi.org/10.3354/meps144265
- 1309 Nunes, S., Latasa, M., Delgado, M., Emelianov, M., Simó, R., & Estrada, M. (2019). Phytoplankton
- 1310 community structure in contrasting ecosystems of the Southern Ocean: South Georgia, South
- 1311
   Orkneys and western Antarctic Peninsula. Deep Sea Research Part I: Oceanographic Research
- 1312 Papers, 151, 103059. https://doi.org/10.1016/j.dsr.2019.06.005
- 1313 Rodriguez, F., Varela, M., & Zapata, M. (2002). Phytoplankton assemblages in the Gerlache and Bransfield Straits
- 1314 (Antarctic Peninsula) determined by light microscopy and CHEMTAX analysis of HPLC pigment data.
- 1315 Deep Sea Research Part II: Topical Studies in Oceanography, 49(4), 723–747.
- 1316 https://doi.org/10.1016/S0967-0645(01)00121-7
- 1317 Roesler, C. S., & Barnard, A. H. (2013). Optical proxy for phytoplankton biomass in the absence of
- 1318 photophysiology: Rethinking the absorption line height. Methods in Oceanography, 7, 79–94.
- 1319 <u>https://doi.org/10.1016/j.mio.2013.12.003</u>
- Roy, S., Llewellyn, C. A., Egeland, E. S., & Johnsen, G. (2011). Phytoplankton Pigments: Characterization,
  Chemotaxonomy and Applications in Oceanography. Cambridge University Press.
- 1322 Thomalla, S., Antoine, D., Berliner, D., Little, H., Moutier, W., Olivier-Morgan, A., et al. (2020). Particulate
- 1323 organic carbon and particulate organic nitrogen concentrations and stable isotope composition of seawater
- 1324 sampled during the Antarctic Circumnavigation Expedition (ACE) during the Austral Summer of
- 1325 2016/2017. (Version 1.0) [Data set]. Zenodo. https://doi.org/10.5281/zenodo.3859515
- Walton, D. W. H., & Thomas, J. (2018). Cruise Report Antarctic Circumnavigation Expedition (ACE) 20th
   December 2016 19th March 2017. Zenodo. https://doi.org/10.5281/zenodo.1443511
- 1328 Zapata, M., Rodríguez, F., & Garrido, J. L. (2000). Separation of chlorophylls and carotenoids from marine
- 1329 phytoplankton: a new HPLC method using a reversed phase C8 column and pyridine-containing mobile
- 1330 phases. Marine Ecology Progress Series, 195, 29–45. https://doi.org/10.3354/meps195029

manuscript submitted to Journal of Geophysical Research

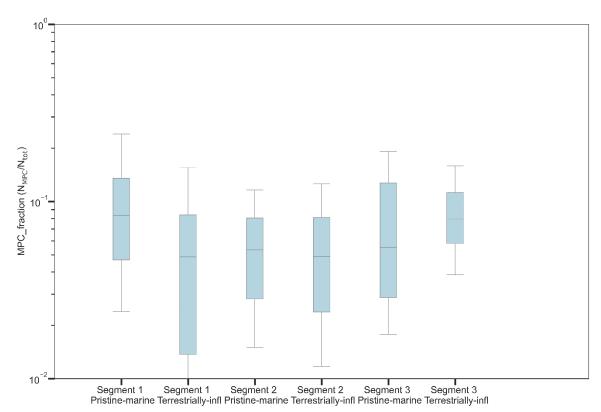
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| 2           | Journal of Geophysical Research Atmospheres  |
| 3           | Supporting Information for   |
| 4<br>5      | Sources, Occurrence and Characteristics of Fluorescent Biological Aerosol Particles<br>Measured over the Pristine Southern Ocean   |
| 6<br>7<br>8 | Alireza Moallemi <sup>1</sup> , Sebastian Landwehr <sup>1,2</sup> , Charlotte Robinson <sup>3</sup> , Rafel Simó <sup>4</sup> , Marina Zamanillo <sup>4</sup> ,<br>Gang Chen <sup>1</sup> , Andrea Baccarini <sup>1,2</sup> , Martin Schnaiter <sup>5,6</sup> , Silvia Henning <sup>7</sup> , Robin L. Modini <sup>1</sup> ,*, Martin<br>Gysel-Beer <sup>1</sup> , and Julia Schmale <sup>1,2</sup> ,* |
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### 31 Introduction

- 32 This supporting document contains information on the analysis of wide band integrated
- 33 bioaerosol senor data, complementary results related to fluorescent and hyper-fluorescent
- 34 aerosol number concentration, description of marine biological and chemical measurements,
- and further details regarding ABC fluorescent classification of aerosol particles. Moreover, the
- 36 document contains scatter plots of (hyper-)fluorescent particle fractions against marine
- biological and chemical variables. The results demonstrated here were acquired during the
- 38 Antarctic Circumnavigation Expedition (ACE) in austral summer 2016-2017.
- 39

### 40 Text S1: Missing Particle Count Fraction

- 41 To show the contribution of the missing particle counts by the WIBS, the fraction of missed
- 42 particle counts (MPC) to the total detected particles particle number (Ntot) are is presented in
- 43 Figure S1. These results indicate that the median of the fraction of missing particle counts
- 44 ranges from ~ 5 to 8 %.
- 45



47 **Figure S1.** Variation of the fraction of missing particle count to total particle number

- 48 concentration measured by the WIBS.
- 49

46

### 51 Text S2: APS vs WIBS coarse mode aerosol measurements

Figure S2 shows the scatter plot of hourly averaged integrated number concentrations
 of total aerosol particles measured by APS and WIBS for particles within the size range of 1 μm
 to 20 μm, for samples collected in segments 1-3 of ACE and samples collected after segment 3
 during the return route from Cape Town to Europe.

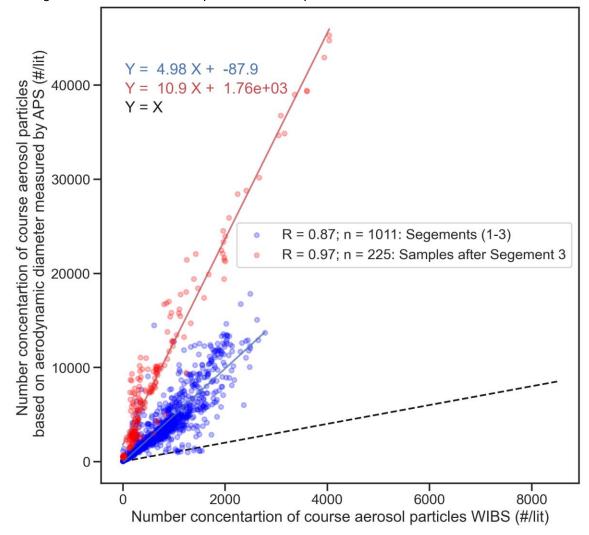


Figure S2. Comparison of particle number concentrations for the diameter range 1 – 20 μm
 obtained with the WIBS and APS, for samples collected during segments 1-3 and samples
 collected after segment 3. The correlation coefficient (R) included in the plot corresponds to
 Pearson's correlation coefficient.

### 67 **Text S3: Marine Measurement Description**

68 In this section, marine variables and their analysis methods are presented.

# 69 **Table S1.** Description of marine microbe measurement used in the correlation study against70 fluorescent aerosol particles.

| Variable                    | Units                  | Description                                 | Methods     |
|-----------------------------|------------------------|---|-------------|
| HDNA_bacteria-sea-p8        | Cells mL <sup>-1</sup> | Concentration of high DNA containing        | See section |
|                             |                        | bacteria                                    | S 4.1       |
| LDNA_bacteria-sea-p8        | Cells mL <sup>-1</sup> | Concentration of low DNA containing         | See section |
|                             |                        | bacteria                                    | S 4.1       |
| Total-bacteria-sea          | Cells mL <sup>-1</sup> | Concentration of total bacteria (high & low | See section |
|                             |                        | DNA containing) bacteria                    | S 4.1       |
| Synechococcus-sea-p8        | Cells mL <sup>-1</sup> | Concentration of Synechoccocus sp. Cells    | See section |
|                             |                        | _   | S 4.1       |
| Picoeukaryotes-type1-sea-p8 | Cells mL <sup>-1</sup> | Concentration of picoeukaryote type 1 cells | See section |
|                             |                        |   | S 4.1       |
| Picoeukaryotes-type2-sea-p8 | Cells mL <sup>-1</sup> | Concentration of picoeukaryote type 2 cells | See section |
|                             |                        |   | S 4.1       |
| Nanoeukaryotes-sea-p8       | Cells mL <sup>-1</sup> | Concentration of nanoeukaryote cells        | See section |
|                             |                        |   | S 4.1       |
| Cryptomonas-sea-p8          | Cells mL <sup>-1</sup> | Concentration of cryptomonas cells          | See section |
| •                           |                        |   | S 4.1       |
| Picoeukaryotes-sea-p8       | Cells mL <sup>-1</sup> | Concentration of picoeukaryote (type 1 &    | See section |
| · -                         |                        | type 2) cells                               | S 4.1       |

# **Table S2.** Description of marine phytoplankton taxa measurements used in the correlation study against fluorescent aerosol particles

| Variable                      | Units              | Description  | Methods              |
|-------------------------------|--------------------|--|----------------------|
| Particulate.Org.Carbon-p1     | $\mu M$            | Particulate organic carbon concentration                 | See section<br>S 4.2 |
| Total_Chlorophyll_a_merged-p1 | $\mu g L^{-1}$     | Total chlorophyll-a concentration                        | See section<br>S 4.3 |
| Chloro                        | $\mu g L^{-1}$     | chlorophyte contribution to chlorophyll biomass          | See section<br>S 4.4 |
| Crypto1                       | $\mu g L^{-1}$     | Cryptophyte contribution to chlorophyll biomass          | See section<br>S 4.4 |
| Cyano2                        | $\mu g L^{-1}$     | Cyanobacteria type 2 contribution to chlorophyll biomas  | See section<br>S 4.4 |
| DiatA                         | $\mu g L^{-1}$     | Diatom type contribution to chlorophyll biomas           | See section<br>S 4.4 |
| DiatB                         | μg L <sup>-1</sup> | Diatom type 2 contribution to chlorophyll biomas         | See section<br>S 4.4 |
| DinoA                         | μg L <sup>-1</sup> | Dinoflagellate type 1 contribution to chlorophyll biomas | See section<br>S 4.4 |
| Hapto8                        | μg L <sup>-1</sup> | Haptophyte type 8 contribution to chlorophyll biomas     | See section<br>S 4.4 |
| Haptophyte67                  | $\mu g L^{-1}$     | Haptophyte type 6&7 contribution to chlorophyll biomas   | See section<br>S 4.4 |
| Pras3                         | $\mu g L^{-1}$     | Prasinophyte type 3 contribution to chlorophyll biomas   | See section<br>S 4.4 |
| Pelago                        | μg L <sup>-1</sup> | Pelagophyte contribution to chlorophyll biomas           | See section<br>S 4.4 |

# 73 **Table S3.** Description of other marine organic measurements. Dissolved Compounds Variable Units Description

| Dissolved Compounds |  |  |  |  |  |  |
|---------------------|--|--|--|--|--|--|
| Units               | Description  | Methods  |  |  |  |  |
| m <sup>-1</sup>     | Colored dissolved organic material                       | See section  |  |  |  |  |
|                     | (CDOM) absorption at 350 nm                              | S 4.5  |  |  |  |  |
| µg XG eq            | Transparent Exopolymeric Particles                       | See section  |  |  |  |  |
| L-1                 |  | S 4.6  |  |  |  |  |
| µg BSA              | Coomasie Stainable Particles                             | See section  |  |  |  |  |
| eq L <sup>-1</sup>  |  | S 4.6  |  |  |  |  |
|                     | m <sup>-1</sup><br>μg XG eq<br>L <sup>-1</sup><br>μg BSA | m <sup>-1</sup> Colored dissolved organic material<br>(CDOM) absorption at 350 nm       μg XG eq<br>L <sup>-1</sup> Transparent Exopolymeric Particles       μg BSA     Coomasie Stainable Particles |  |  |  |  |

### 74 Text S4: Description of methods used for marine measurement

### 75 **S4.1** Marine microbe number concentration measurements

- 76 Number concentration of bacteria and pico-, nano- and microalgae in sea water were
- 77 measured through cytometry. After extraction, sea water samples were aliquoted in
- 78 cryovials. For each samples 4.5 ml duplicates and 1.8 ml replicate were collected. The
- <sup>79</sup> samples were treated by 1% paraformaldehyde plus 0.05% glutaraldehyde and kept at
- 80 80 °C until analysis on land. After thawing, samples were analysed with a PARTEC
- 81 Cube 8 flow cytometer equipped with a laser emitting at 488 nm. Heterotrophic
- 82 bacteria were counted by their signature in a plot of side scatter versus green
- fluorescence after being stained with 10 μM of SYBRGreen I. In separate runs of
   unstained samples, pico- and nano-phytoplankton and cryptomonas cells were
- identified and enumerated on the basis of the differences in autofluorescence and
- 86 light scattering characteristics.

## 87 **S4.2** Particulate organic carbon concentration measurements

- 88 Particulate organic carbon was measured by extracting 2000 ml of sea water samples
- and filtering them using 25 mm combusted 0.3 μm Glass Fibre filters (GF-75;
- 90 Sterlitech). After sample extraction, the filter papers were kept in combusted tinfoil
- and cooled down to -80 °C. The filters were analyzed in University of Cape Town using
- an elemental analyser-isotope ratio mass spectrometer (Walton and Thomas, 2018).
- 93 The particulate organic carbon data could be found in (Thomalla et al., 2020).
- 94

## 95 **S4.3 Merged total chlorophyll-a**

- 96 Absolute concentrations of total chlorophyll-a pigment concentration were derived
- via high performance liquid chromatography (HPLC, Antoine et al., 2019) at locations
- 98 roughly every 6-12 hours. Measurements of particulate absorption were collected at a
- 99 higher resolution, roughly every 3-6 hours. Using matched samples of HPLC derived
- 100 total chlorophyll-a and particulate absorption, the absorption line height method of
- 101 Roesler & Barnard (2013) for determining total chlorophyll-a concentration was
- 102 calibrated and applied to the whole particulate absorption dataset in order to increase
- 103 the resolution of the total chlorophyll-a concentration estimations

# 104 **S4.4 Phytoplankton CHEMTAX**

- 105 The data on phytoplankton taxonomy groups and their contributions were obtained
- 106 from the pigment concentration measurements (Antoine et al., 2019) and by using

107 CHEMTAX v1.95 chemical taxonomy software (Mackey et al., 1996). The quantified

108 taxonomy groups in this studies are: Chlorophytes type 1, cryptophytes type 2,

109 diatoms type 1, diatoms type 2, dinoflagellates type 1, haptophytes type 8,

- haptophytes types 6 + 7, prasinophytes, and pelagophytes (Higgins at al., 2011). 110
- 111 Before conducting CHEMTAX analysis, the data was pre-processed and clustered. The
- data was standardized was based on mean subtracted and divided by standard 112
- 113 deviation. Prior to clustering the data, a dissimilarity matrix was computed based
- 114 Manhattan's distances. Hierarchical clustering (Ward's method) was used for clustering
- 115 analysis and the Elbow, silouette and gap tests indicated the existence of 5 clusters.
- 116 The CHEMTAX analysis was conducted on the clustered data. Initially, to obtain the
- 117 matrices of optimized pigment rations, 60 analysis runs were performed on each 118
- individual clustered. This was followed by a final 20 analysis runs on the data to
- 119 calculate the taxonomic abundance. In this study the initial pigment ratios were 120 gathered from Rodriguez et al. (2002) (2002), Zapata et al. (2004), Cook et al. (2011)
- 121 and Higgins at al. (2011), Cassar et al. (2015), Nunes et al. (2019).
- 122

#### 123 S4.5 Coloured dissolved organic matter (CDOM) concentration measurements

124 Coloured dissolved organic matter is a component dissolved organic matter (DOM) in

- 125 seaweter which strongly absorbs light in the ultraviolet wavelengths. CDOM is
- 126 typically strongly correlated with DOM and could be used as a proxy for DOM.
- 127 The absorption spectra of the CDOM from the collected sample were measured
- 128 onboard with a UV-spectrometer, and the data included in this analysis corresponds to
- 129 the absorption of CDOM at wavelength of 350 nm. Further information can be found
- 130 in the cruise report (Walton and Thomas, 2018).

#### 131 S4.6 Transparent Exopolymeric Particles (TEPs) and Coomasie Stainable Particles (CSPs) 132 measurements

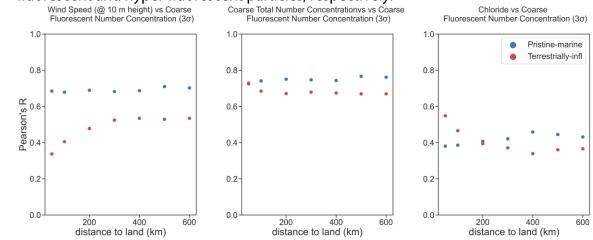
133

134 Transparent exopolymeric particles (TEP) and coomassie-blue stainable particles (CSP) 135 are gel-like compounds that are rich in polysaccharide and protein, respectively. 136 Seawater samples (150-300 ml) were filtered through 25 mm diameter 0.4 µm pore 137 size polycarbonate filters. For TEP analysis the filters were stained with 500 µL of Alcian 138 blue solution (0.02 %, pH 2.5) for 5 s, rinsed with Milli-Q water and stored frozen. For 139 CSP analysis, the filters were stained with 700 µL of a working Coomassie Brilliant Blue (CBB-G 250) solution (0.04 %, pH 7.4) for 30 s, rinsed with Milli-Q water and stored 140 141 frozen. For each batch of TEP and CSP samples duplicate blank filters which were not 142 stained were collected. Measurements of TEP and CSP were conducted in land 143 laboratories. For TEP all the samples and blank filters were treated in 5 ml of 80% 144 sulfuric acid and shaken intermittently for 3 h. The measurement was conducted by a 145 spectrophotometre at 787 nm (Varian Cary spectrophotometer). For CSP all the 146 samples and blank filters were treated in 4 mL of extraction solution (3 % SDS in 50 % 147 isopropyl alcohol) and sonicated in a water bath at 37° C for 2 hours. The CSP 148 measurement was conducted y a spectrophotometre at 615 nm (Shimadzu UV-Vis 149 UV120). The Alcian blue dye solution calibration was performed using a standard

- 150 solution of Xanthan Gum (XG). The CBB dye solution calibration was performed using
- 151 bovine serum albumin standard (BSA).

# 152 Text S5: Correlation analysis of SSA proxies vs fluorescent aerosols at different 153 land proximity values

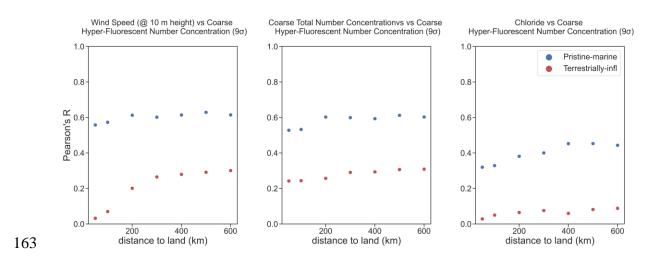
- 154 To find a reasonable proximity to land distance, the Pearson's R values of the different
- 155 proxy variables against fluorescent and hyper-fluorescent coarse particles were
- 156 obtained as a function of the distance to land. Figures S3 and S4 show the results for
- 157 fluorescent and hyper-fluorescent particles, respectively.



159 **Figure S3.** Pearson's R values for pristine-marine and terrestrially-influenced air masses of 160 fluorescent particles for different land proximity values.

161

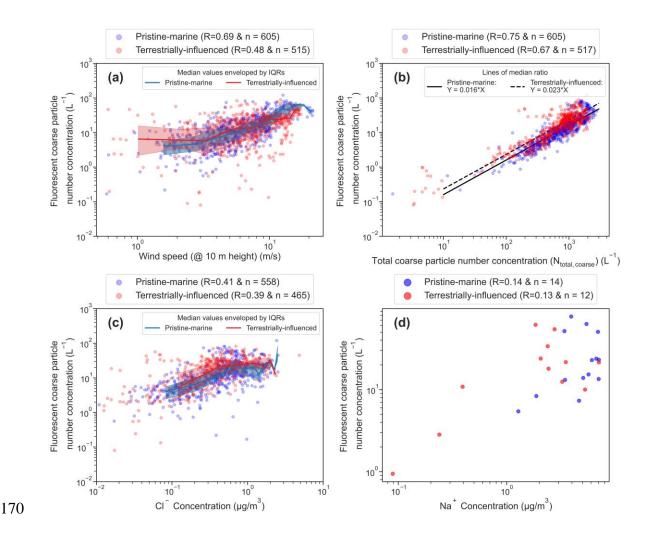
158



164 Figure S4. Pearson's R values for pristine-marine and terrestrially-influenced air masses of165 hyper-fluorescent particles for different land proximity values.

# Text S5: Scatterplots of fluorescent particle (3σ) concentrations against the four proxy variables for SSA concentrations

- 168 The scatter plots for fluorescent coarse particles vs SSA proxies are presented in Figure
- 169 S5.



171 **Figure S5.** Scatter plots of pristine-marine and terrestrially-influenced air masses of 172 fluorescent particles vs SSA proxies for the combined segment 1 to segment 3 result

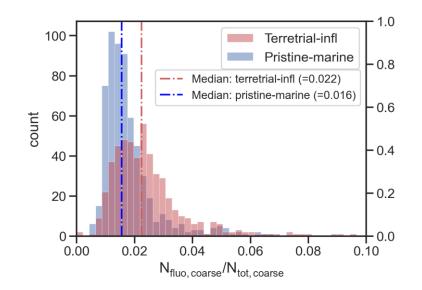
172 fluorescent particles vs SSA proxies for the combined segment 1 to segment 3 results. The red 173 and blue shades correspond to the interguartile ranges (IQR) of the measurements that were

174 calculated by binning the dataset into ten equidistant logarithmic bins.

- 175
- 176

# 177 Text S6: Distribution of the number concentration fraction of fluorescent PBAPs 178 to coarse SSA number concentrations

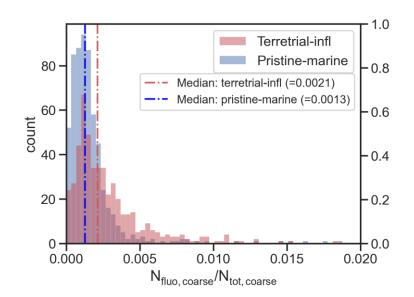
- 179 The histograms of the fraction of (hyper-)fluorescent number concentrations to total
- 180 coarse aerosol particle number concentrations based on hourly averaged data are
- 181 shown in Figures S6 and S7.





**Figure S6.** Distribution of number fraction of fluorescent PBAP to total coarse particle number

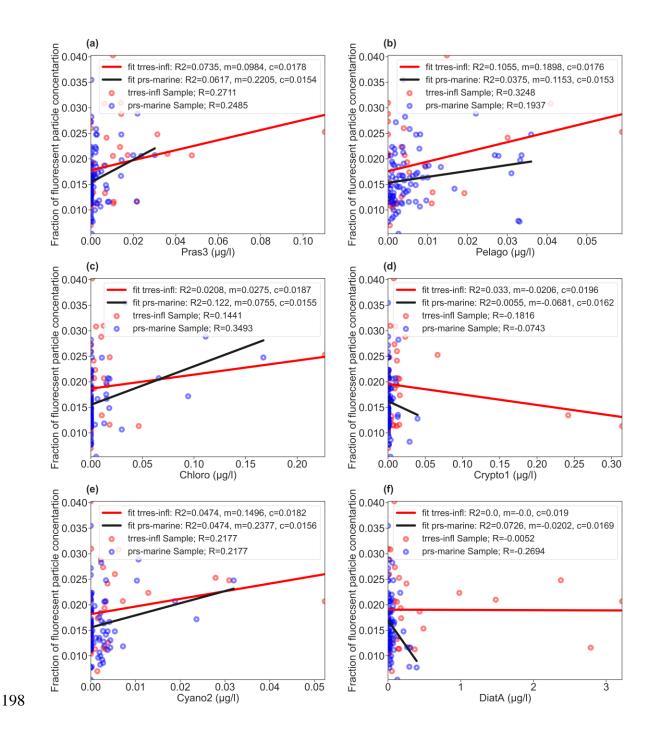
184 concentration.



- 187 Figure S7. Distribution of number fraction of hyper-fluorescent PBAP to total coarse particle188 number concentration.

# 190 Text S7: Scatter plots of different marine variables against normalized fluorescent 191 number concentration

- 192 **S7.1** Fluorescent particle number concentration fraction vs phytoplankton taxa
- 193 Figures S8 and S9 show the results of the fraction of coarse fluorescent particle
- 194 number concentrations to total coarse particles against marine measurements
- 195 associated with phytoplankton taxa. All the fit lines in the plots demonstrated in
- 196 section S7 correspond to linear regressions that were applied on the datasets. The
- 197 Pearson's R values are also included.



**Figure S8.** Scatter plot of fraction of coarse fluorescent particle number concentrations to

200 total coarse particles vs. different phytoplankton taxa measurements.

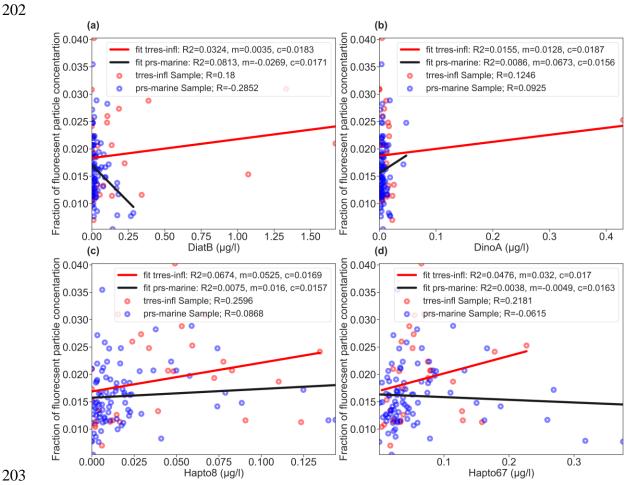
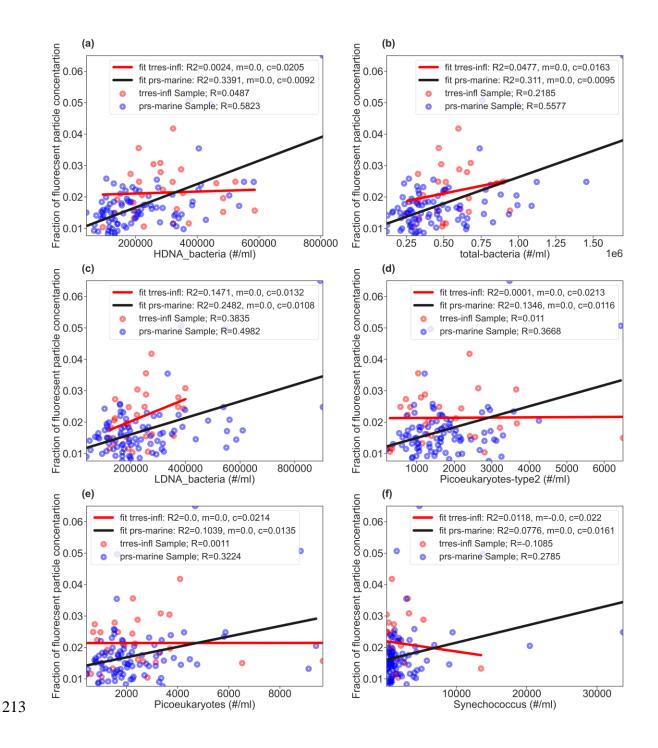


Figure S9. Scatter plot of fraction of coarse fluorescent particle number concentrations to total coarse particles vs. different phytoplankton taxa measurements

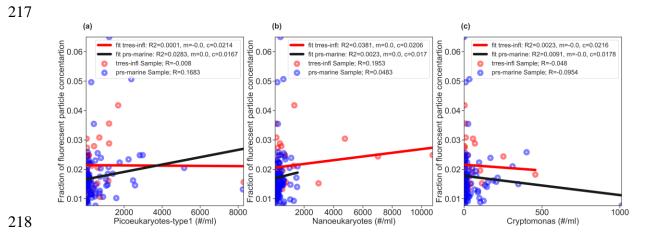


- Figures S10 and S11 show the results of the fraction of coarse fluorescent particle
- number concentrations to total coarse particles against marine measurements
- associated with marine microbe measurements. Fits are analogue to S 9.1.



**Figure S10.** Scatter plot of fraction of coarse fluorescent particle number concentrations to

215 total coarse particles vs. different marine microbe measurements



219 **Figure S11.** Scatter plot of fraction of coarse fluorescent particle number concentrations to

- 220 total coarse particles vs. different marine microbe measurements
- 221

# 57.3 Fluorescent particle number concentration fraction vs organic matter (OM) measurements

- 224 Figure S12 shows the results of fraction of coarse fluorescent particle number
- 225 concentrations to total coarse particles against OM measurements. Fits are analogue
- 226 to \$ 9.1.

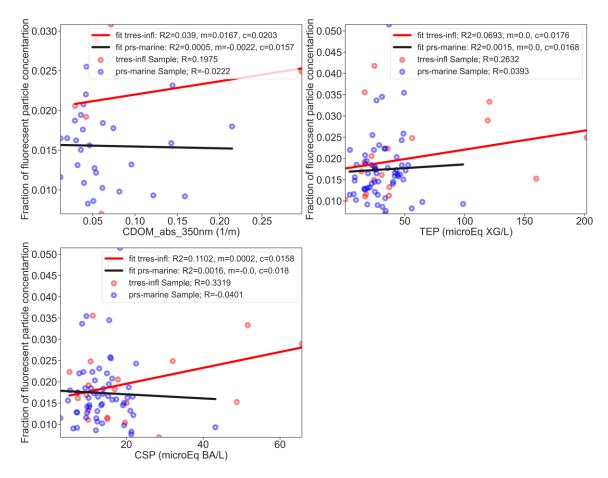


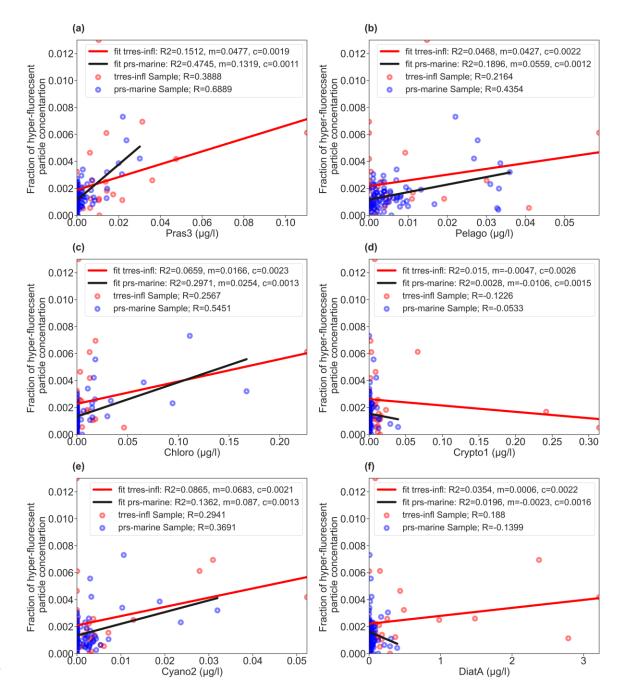
Figure S12. Scatter plot of fraction of coarse fluorescent particle number concentrations to
 total coarse particles vs. OM measurements.

227

231

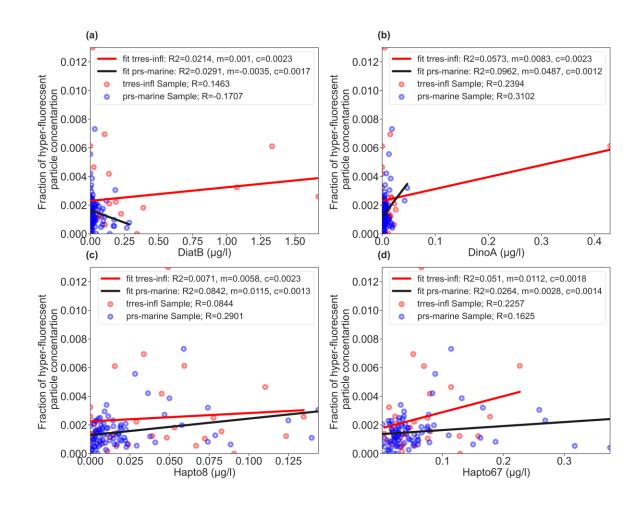
### 232 **S7.4** Hyper-fluorescent particle number concentration fraction vs phytoplankton taxa

- 233 Figures S13 and S14 show the scatter results of fraction of coarse hyper-fluorescent
- 234 particle number concentrations to total coarse particles against marine measurements
- associated with phytoplankton taxa.



**Figure S13.** Scatter plot of fraction of coarse hyper-fluorescent particle number

- 239 concentrations to total coarse particles vs. different phytoplankton taxa measurements



243

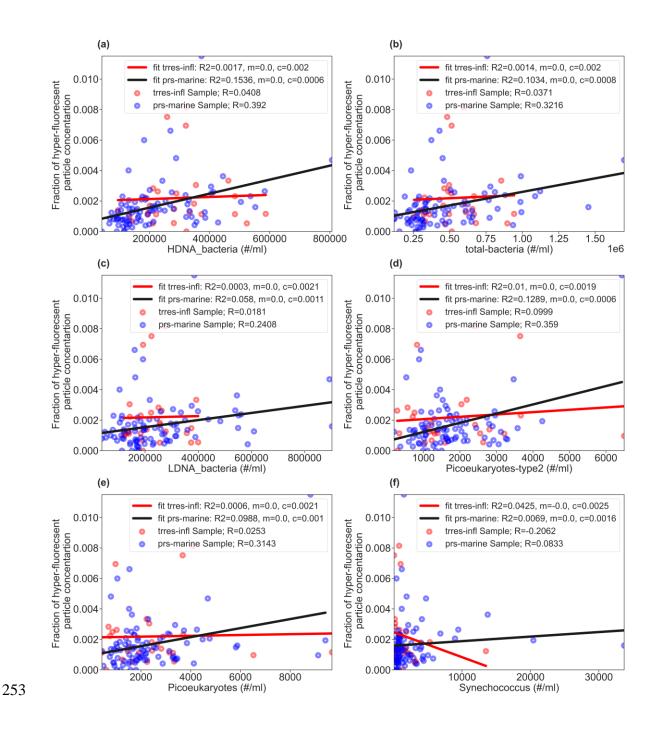
244 **Figure S14.** Scatter plot of fraction of coarse hyper-fluorescent particle number

245 concentrations to total coarse particles vs. different phytoplankton taxa measurements

246

# S7.5 Hyper-fluorescent particle number concentration fraction vs marine microbe measurements

- 249 Figures S15 and S16 show the scatter results of fraction of coarse hyper-fluorescent
- 250 particle number concentrations to total coarse particles against marine measurements
- associated with marine microbe measurements.



**Figure S15.** Scatter plot of fraction of coarse hyper-fluorescent particle number

255 concentrations to total coarse particles vs. different marine microbe measurements

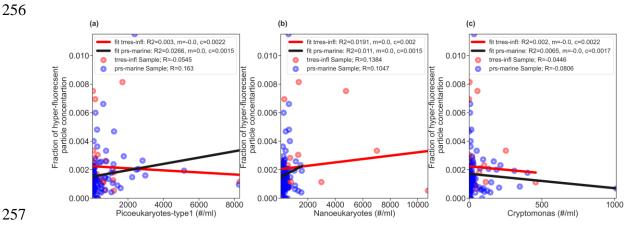
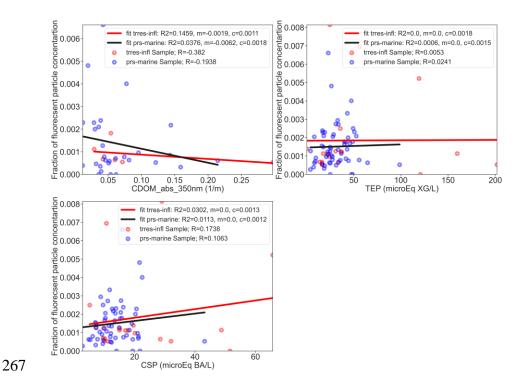


Figure S16. Scatter plot of fraction of coarse hyper-fluorescent particle number
 concentrations to total coarse particles vs. different marine microbe measurements

261 **S7.6** Hyper-fluorescent particle number concentration fraction vs OM measurements

- 262 Figure S17 shows the results of fraction of coarse fluorescent particle number
- 263 concentrations to total coarse particles against OM measurements. Fits are analogue
- 264 to \$ 9.1.ss

265



- 268 **Figure S17.** Scatter plot of fraction of coarse hyper-fluorescent particle number
- 269 concentrations to total coarse particles vs. OM measurements

## 270 Text S8: p value results of marine measurement

- 271 The p values for the marine varibales used in the correlation study against (hyper-
- 272 fluorescent particles) are demonstrated in Figure S18.

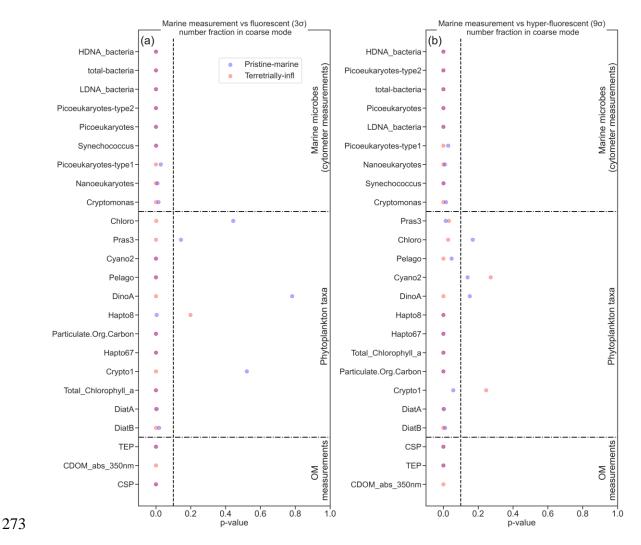


Figure S18. (a) p values of marine variables against fluorescent aerosols, and (b) p values of marine variables against hyper-fluorescent aerosols.

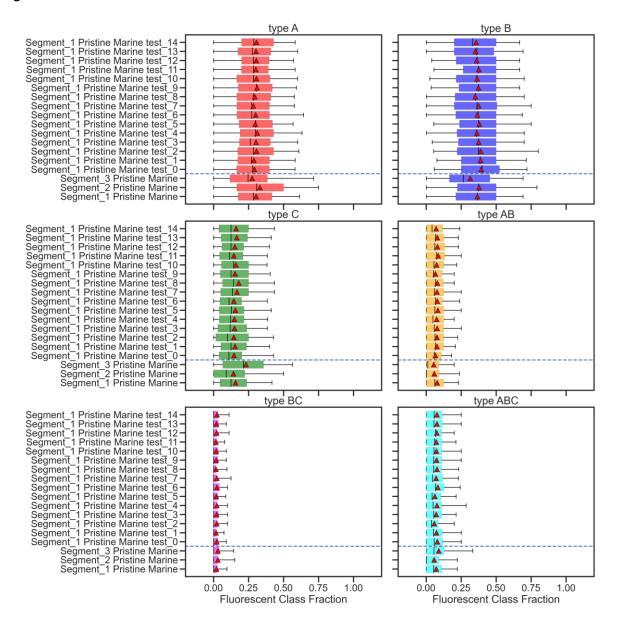
### 277 Text S8: Subsampling analysis of fluorescent type classification

# S.8.1 Variation of the fluorescent type fraction of pristine-marine segment samples based on 24 hour random data points subsampling

- 280 Figure S.19 to S.21 demonstrate the resampling results for segment 1 to 3. Random
- subsamples of 288 points (equivalent to 24 hours of data) from 5 min time average
- 282 datasets of fluorescent aerosol measurements from pristine-marine air masses of
- 283 different segments were drawn. The resampling process was repeated 15 times to

#### 284 provide a number of resample ensembles to compare their variability with results of

#### full segment data sets.



286

- 287 **Figure S19.** Fluorescent type fraction subsampling results for pristine-marine air masses from
- 288 segment 1 for coarse fluorescent particles (3σ)

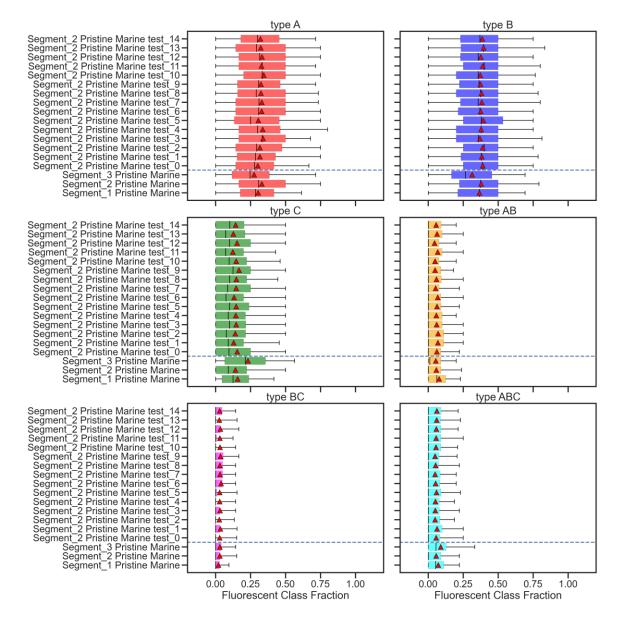




Figure S20. Fluorescent type fraction subsampling results for pristine-marine air masses from
 segment 2 for coarse fluorescent particles (3σ)

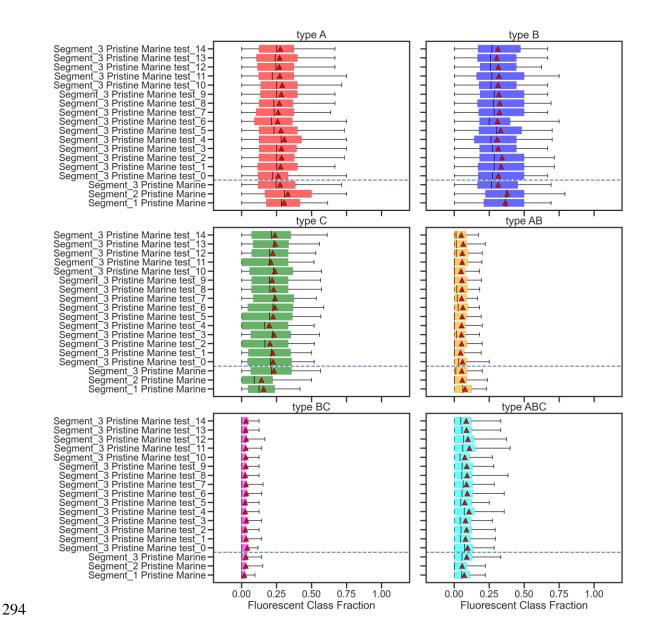


Figure S21. Fluorescent type fraction subsampling results for pristine-marine air masses from
 segment 3

#### 298 **S.8.2** Variation of fluorescent type fraction based on a constant time window of 24 hours

- 299 To investigate the variability of fluorescent type fraction of pristine-marine air masses
- 300 of each segment over different time periods, an additional subsampling analysis was
- 301 conducted by drawing subsamples from a fixed time interval of 24 hours. Figure S.22

to S.24 demonstrate the resampling results for segment 1 to 3. For this analysis, in
 each segment 15 different and randomly selected time intervals were used. Only time
 intervals containing a total number of data points equivalent to or longer than 12
 hours within the 24 hours were considered.

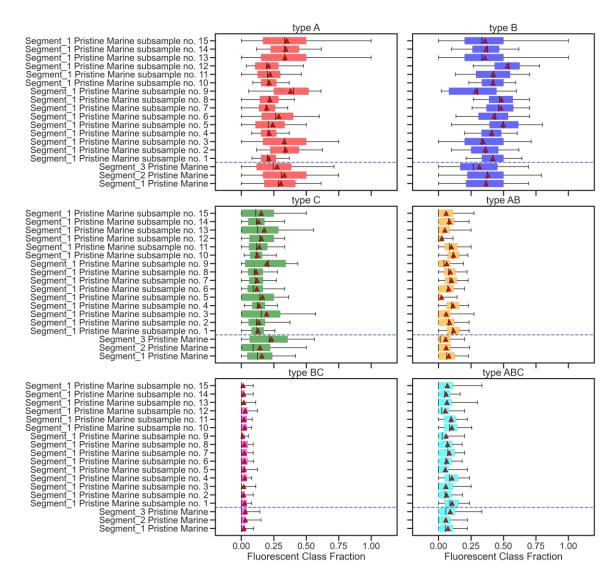
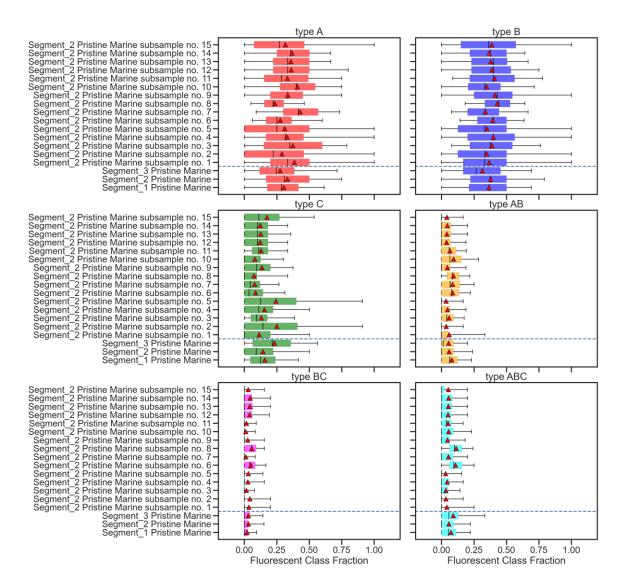
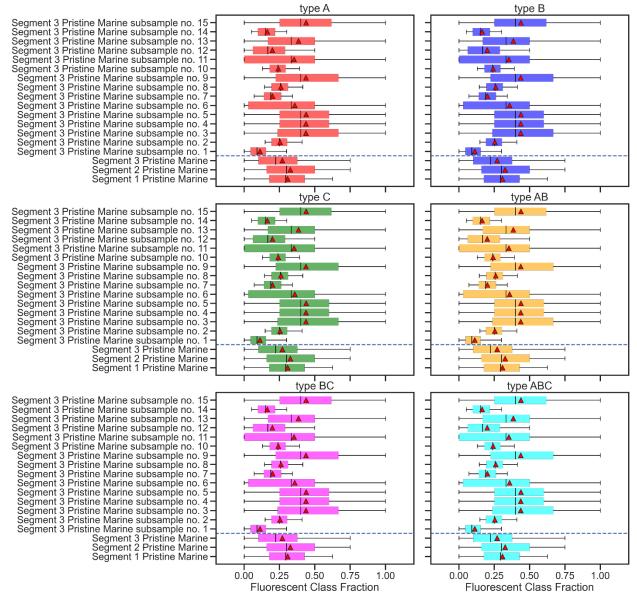


Figure S22. Results of fluorescent type fraction subsampling (based on fixed time windows)
 results for pristine-marine air masses from segment 1 for coarse fluorescent particles (3σ)

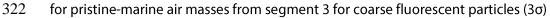


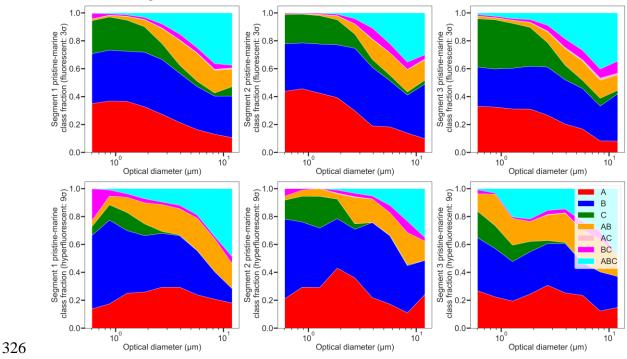
**Figure S23.** Results of fluorescent type fraction subsampling (based on fixed time windows)

- 317 results for pristine-marine air masses from segment 2 for coarse fluorescent particles (3σ)



320 Fluorescent Class Fraction Fluorescent Class Fraction Fluorescent Class Fraction Fluorescent Class Fraction Subsampling (based on fixed time windows)





325 Text S9: Average size distribution of aerosol fluorescent classes

Figure S25. Size distribution of fluorescent type fraction for fluorescent particles (3σ) (top
 row) and hyper-fluorescent particles (9σ) bottom row for pristine-marine air masses from
 segment 1 to 3

#### 331 Text S10: Asymmetry Factor (AF)

- 332 Once aerosols are illuminated by the continuous 635 nm laser beam of the WIBS, their
- 333 forward scattering is measured by a quadrant detector. The quadrant detector has
- four sensors, which measure a portion of the scattered light intensities. The
- asymmetry factor is obtained by combining these four measured light intensities
- through the following formula introduced by Gabey et al. (2010) and used in other
- 337 studies (Savage et al. 2017):

$$AF = \frac{k(\sum_{i=1}^{n} (E - E_{i})^{2})^{\frac{1}{2}}}{E}$$
 Equation S1

338 In Eq S1, *k* is an instrument constant, *E* is the mean forward scattering signal measured

- by all the detector sensors, and  $E_i$  the scattering signal detected by an individual
- 340 sensor and *n* is the number of sensors.