

# Characterization of Transient-Large-Amplitude Geomagnetic Perturbation Events

Brett A. McCuen<sup>1</sup>, Mark B. Moldwin<sup>2</sup>, and Mark J. Engebretson<sup>3</sup>

<sup>1</sup>University of Michigan

<sup>2</sup>University of Michigan-Ann Arbor

<sup>3</sup>Department of Physics, Augsburg University

November 24, 2022

## Abstract

We present a characterization of transient-large-amplitude (TLA) geomagnetic disturbances that occurred at six stations of the Magnetometer Array for Cusp and Cleft Studies throughout 2015. TLA events are defined as one or more short-timescale ( $< 60$  seconds) dB/dt signature with magnitude  $> 6$  nT/s. A semi-automated dB/dt search algorithm was developed to identify TLA events in ground magnetometer data and used to identify 40 TLA dB/dt events. We demonstrate the existence of large-amplitude dB/dt with timescale less than 10 seconds in nine of the events. The association of these events to sudden commencements is relatively weak, rather the events are more likely to occur in relation to substorm onsets. However, 15% of TLA events show no direct association to geomagnetic storms, substorms or nighttime magnetic impulse events.

**Characterization of Transient-Large-Amplitude Geomagnetic Perturbation Events**

Brett A. McCuen<sup>1</sup>, Mark B. Moldwin<sup>1</sup>, Mark Engebretson<sup>2</sup>

<sup>1</sup>University of Michigan, Ann Arbor, Michigan, USA, <sup>2</sup>Augsburg University, Minneapolis, Minnesota, USA

**Contents of this file**

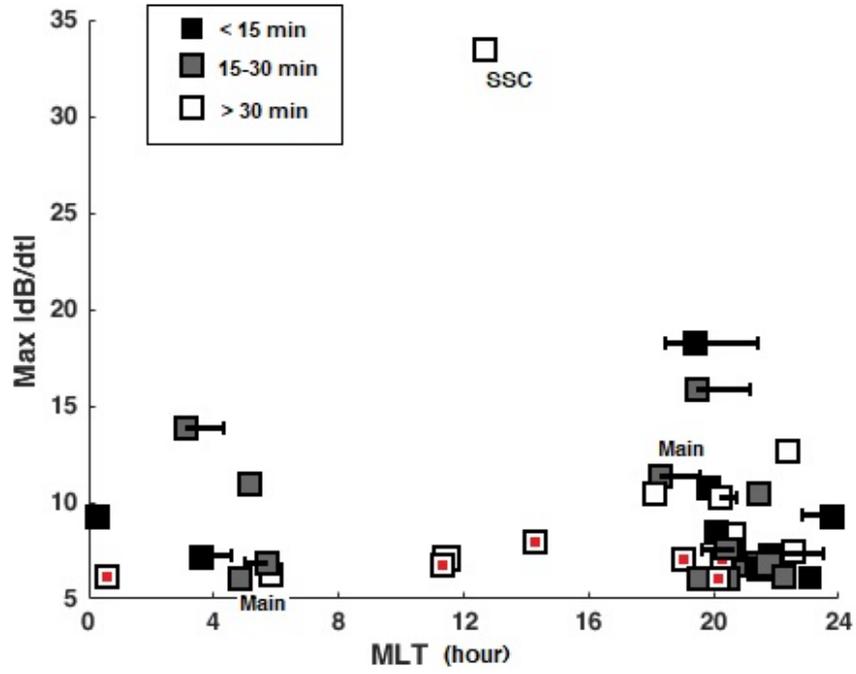
Table S1

**Introduction**

This supporting information provides a table with the geographic latitude and longitude, and the corrected geomagnetic latitude and longitude of the six MACCS stations used in this study. These corrected geomagnetic coordinates were calculated for the year of 2015 with the IGRF transformation tool of the World Data Center (WDC) for Geomagnetism, Kyoto. These stations can be found on the map of Figure 1 of the main article.

Station	Geographic Latitude	Geographic Longitude	Corrected Geomagnetic Latitude	Corrected Geomagnetic Longitude
IGL	69.30	278.2	78.63	343.3
GJO	68.63	264.2	76.86	320.5
RBY	66.52	273.8	75.62	22.33
PGG	66.1	294.2	75.53	11.16
CDR	64.2	283.4	73.70	353.8
NAN	56.4	298.3	65.67	14.80

**Table S1.** Locations of MACCS tables used in this study.



1                   **Characterization of Transient-Large-Amplitude**  
2                   **Geomagnetic Perturbation Events**

3                   **Brett A. McCuen<sup>1</sup>, Mark B. Moldwin<sup>1</sup>, Mark Engebretson<sup>2</sup>**

4                                   <sup>1</sup>University of Michigan, Ann Arbor, Michigan  
5                                   <sup>2</sup>Augsburg University, Minneapolis, Minnesota

6                   **Key Points:**

- 7                   • Short-timescale (< 60 s) geomagnetic perturbation events found at 6 high-latitude  
8                   MACCS stations throughout 2015 are characterized.  
9                   • The existence of large-amplitude  $dB/dt$  at Earth's surface with timescale 1-10 sec-  
10                   onds is demonstrated.  
11                   • The exact physical mechanisms driving TLA events are still unclear

**Abstract**

We present a characterization of transient-large-amplitude (TLA) geomagnetic disturbances that occurred at six stations of the Magnetometer Array for Cusp and Cleft Studies throughout 2015. TLA events are defined as one or more short-timescale ( $< 60$  seconds)  $dB/dt$  signature with magnitude  $\geq 6$  nT/s. A semi-automated  $dB/dt$  search algorithm was developed to identify TLA events in ground magnetometer data and used to identify 40 TLA  $dB/dt$  events. We demonstrate the existence of large-amplitude  $dB/dt$  with timescale less than 10 seconds in nine of the events. The association of these events to sudden commencements is relatively weak, rather the events are more likely to occur in relation to substorm onsets. However, 15% of TLA events show no direct association to geomagnetic storms, substorms or nighttime magnetic impulse events.

**Plain Language Summary**

Severe space weather events like geomagnetic storms and substorms cause geomagnetically induced currents (GIC) in electrically conducting material on Earth that are capable of damaging transformers and causing large-scale power grid failure. Models have been developed to forecast GIC that rely on estimation of the surface magnetic field fluctuations,  $dB/dt$ , but require knowledge of the geomagnetic field behavior on short-timescales ( $< 60$  seconds) to more accurately predict GICs. Further, there is some evidence to suggest that extreme second-timescale geomagnetic perturbations may play a role in GIC production that has been previously overlooked. In this study, we investigate transient-large-amplitude (TLA) surface magnetic field disturbances in an effort to better understand the second-timescale nature of the geomagnetic field.

**1 Introduction**

Space weather events cause disturbances of the magnetosphere-ionosphere (M-I) system that result in fluctuations in the surface geomagnetic field. These large-amplitude surface magnetic perturbations, or  $dB/dt$ , generate geomagnetically induced currents (GIC) in electrical systems on Earth. GICs can be large enough to cause damage to transformers resulting in major power outages and costly equipment damage (Pulkkinen et al., 2017). The time derivative of the surface magnetic field,  $dB/dt$ , is often used to study GICs as it is proportional to the spatially-varying geoelectric field via Faraday's Law. In an effort to mitigate potential hazards and safeguard power systems, several models have been developed (e.g., Toth et al., 2005; Ngwira et al., 2014) to predict the geomagnetic behavior following severe space weather events. These models have been evaluated by their success in predicting whether the local horizontal  $dB/dt$  exceeds a threshold within a 20-minute time interval and have been validated by the geospace community (Pulkkinen et al., 2013).

However, there are still many challenges in accurately predicting  $dB/dt$ . Currently available models cannot predict minute or second-scale variation of the local magnetic field. An investigation of a method to improve the efficiency of the Space Weather Modeling Framework (SWMF) predictions by Toth et al, (2019) found that it is reasonably successful at predicting whether  $dB/dt$  will exceed some threshold in the next 20-minutes but it is unlikely that it would be successful in making this prediction in the next 5-minutes, ultimately concluding that 1-minute observed data are insufficient to accurately estimate  $dB/dt$ . While technological advancements in recent years have enabled surface magnetic field measurements with 1 s temporal resolution, geomagnetic perturbations in the  $\sim 1$  Hz frequency band have not been well studied due to their similarity to lightning signals whose impacts on technology are often well mitigated (e.g., Rivera et al., 2016; Gombosi et al., 2017), but analysis of transient-large-amplitude (TLA) geomagnetic disturbances could greatly improve space weather forecasting models.

61 In relation to GICs, second-scale  $dB/dt$  are generally attributed only to sudden com-  
 62 mencements (SC) as an M-I driver (Kataoka & Ngwira, 2016). However, there is evidence  
 63 to suggest that SCs are not the only driver for large-amplitude transient  $dB/dt$  at the  
 64 surface. Several studies suggest that there are more complex, small-scale and localized  
 65 processes involved in generating some extreme GICs (e.g., Engebretson et al., 2021; Ng-  
 66 wira et al., 2015, 2018). Further, a study by Simpson (2011) used a finite-difference time  
 67 domain model to conclude that rapid ionospheric current fluctuations of order 1-second  
 68 can induce substantial currents in power transmission lines following a severe coronal mass-  
 69 ejection (CME).

70 Understanding the transient behavior of the surface geomagnetic field will help to  
 71 improve GIC forecasting models to predict hazardous GICs more quickly and accurately,  
 72 as well as enable a future investigation into whether large-amplitude, second-scale  $dB/dts$   
 73 play a role in producing GIC at the surface. In this study, we surveyed transient (<60  
 74 s), large-amplitude (> 6 nT/s) surface geomagnetic perturbation events that occurred  
 75 at six stations of the Magnetometer Array for Cusp and Cleft Studies (MACCS) through-  
 76 out 2015. We characterize these TLA signatures by their frequency of occurrence, tempo-  
 77 ral dependence and relation (or lack thereof) to various space weather events.

## 78 2 Data Set and Identification Technique

79 The data used in this study are from six ground magnetometer stations of the Mag-  
 80 netometer Array for Cusp and Cleft Studies (MACCS). The stations are located in north-  
 81 east Nunavut, Canada, shown on the map in Figure 1 in corrected geomagnetic (CGM)  
 82 coordinates (geographic and CGM coordinates are listed in Supporting Information Ta-  
 83 ble S1). The CGM coordinates were calculated for the year of 2015 with the IGRF trans-  
 84 formation tool of the World Data Center (WDC) for Geomagnetism, Kyoto. The MACCS  
 85 magnetometers collect 8 samples per second in three axes, then averages and records the  
 86 data at two samples per second (Hughes and Engebretson, 1997). The half-second sam-  
 87 pling rate and high sensitivity (0.01 nT resolution) of the MACCS magnetometers is suf-  
 88 ficient to detect shorter period Pc 1 and 2 pulsations. The geomagnetic variations mea-  
 89 sured by the magnetometers are in local geomagnetic coordinates:  $X$  (north-south),  $Y$   
 90 (east-west) and  $Z$  (vertical).

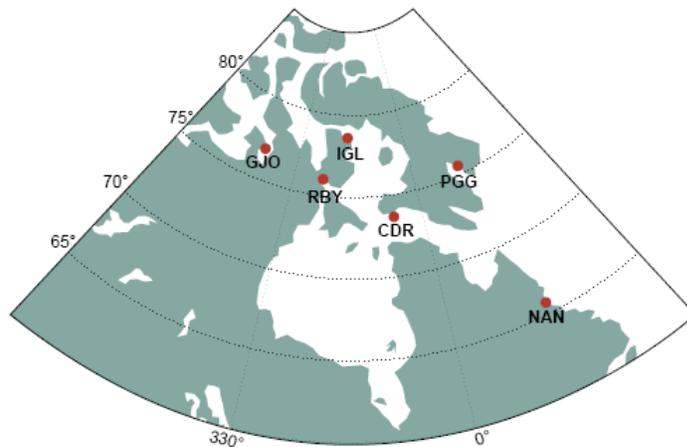


Figure 1: Map of the six MACCS stations used in this study with grid lines in corrected geomagnetic coordinates. Geographic and CGM coordinates for each station are listed in Table S1 in Supporting Information.

91 A semi-automated algorithm was developed to identify  $dB/dt$  signatures in mag-  
92 netometer data with user-specified duration and magnitude. After initial data process-  
93 ing to remove instrument artifacts and smooth the data with a sliding average (if desired  
94 and with user-specified window length), the algorithm is essentially a series of filters. First  
95 the algorithm calculates the slope between each and every data point and determines  
96 the sign of the slope (assigns a 1 if positive slope, 0 if negative slope). If the sign of the  
97 slope changes for at least 1-second (two data points), the data point at which this change  
98 occurs (i.e. local minima or maxima) is flagged. Then the last filter recalculates the new  
99  $dB/dt$  between each local maxima and minima and returns the information of the sig-  
100 nature if it meets the conditions of the defined thresholds for  $dB/dt$  and  $\Delta t$ . The final  
101 product returned from the algorithm is a seven column matrix, each row represents an  
102 individual event and provides the start and end time of the event, start and end B value,  
103 the time elapsed of the event:  $dt$ , the change in magnetic field amplitude:  $dB$ , and finally  
104 the total perturbation:  $dB/dt$ .

105 We used this algorithm to identify  $dB/dt$  signatures with amplitude 6 nT/s or higher  
106 and duration less than 60 seconds. The  $dB/dt$  threshold is comparable to the surface mag-  
107 netic field perturbations (approximately  $\pm 8$  nT/s) that caused the HydroQuebec power  
108 grid to fail during the geomagnetic storm of March 1989 (Kappenman, 2006). We char-  
109 acterize a transient-large-amplitude (TLA)  $dB/dt$  event as an occurrence of one or more  
110 of these signatures if they occur within 1-hour of another (regardless of the axis mea-  
111 sured in and the station measured at). Because of the timescale and magnitude of the  
112  $dB/dts$  sought, many of these signatures are similar in nature to magnetometer noise caused  
113 either by instrumental artifacts or magnetic deviation due to by interference by ferro-  
114 magnetic materials in the vicinity of the magnetometer (Nguyen et al., 2020). Thus, each  
115 event returned from the routine was visually inspected to confirm that it appeared to  
116 be of physical nature or remove it if it was a result of noise. In our manual inspection  
117 process, we found that the events resulting from magnetometer noise have several char-  
118 acteristics that make them possible to automatically detect. Our future work will incor-  
119 porate a machine learning noise identification method that will help to fully automate  
120 the  $dB/dt$  search algorithm.

121 After the filtering process, a total of 181 transient-large-amplitude  $dB/dt$  signa-  
122 tures were identified. The majority ( $\sim 63\%$ ) of these signatures were measured in the x-  
123 component, 29.5% in the y-component and 7.5% in the z-component. Finally, grouping  
124 the  $dB/dts$  if they occurred within 1 hour of another signature resulted in a total of 40  
125 TLA  $dB/dt$  events. While the primary temporal periods of interest in this study are 1-  
126 60 seconds, we also ran the algorithm with the upper limit for the duration of events ex-  
127 tended to 5 minutes in order to compare to the 5-10 minute lasting magnetic impulse  
128 events (MIE) studied in Engebretson et al. (2019). Note that we used cleaned, full res-  
129 olution half-second magnetic field data in this study and GIC measurement often involves  
130 averaging magnetometer data over 1 minute (e.g., Pulkkinen et al., 2006; Ngwira et al.,  
131 2008). Because our identification method relies on changes of the magnetic field lasting  
132 at least 1 second, some larger and more extended  $dB/dts$  are undetected by our algo-  
133 rithm due to more rapid changes of the slope within.

134 Our analysis of TLA event dependence on space weather events relies on several  
135 databases. The SuperMAG Ring Current (SMR) index was used to determine geomag-  
136 netic storm activity and the SuperMAG Electrojet indices (SME) were used to exam-  
137 ine auroral substorm activity during the events ([supermag.jhuapl.edu/indices/](http://supermag.jhuapl.edu/indices/)). The as-  
138 sociation of TLA events with SCs was determined with the International Service of Ge-  
139 omagnetic Indices Sudden Commencement event list ([isgi.unistra.fr/events\\_sc.php](http://isgi.unistra.fr/events_sc.php)).

140

### 3 Occurrence of Transient-Large-Amplitude (TLA) $dB/dt$ Events

141

142

143

144

145

146

147

We identified 40 TLA events consisting of one or more  $dB/dt$  signatures with magnitude 6 nT/s or higher and duration less than 60 seconds. Figure 2 shows three panels with examples of distinct TLA events identified at the MACCS stations in 2015. The hollow circles in all three panels of Figure 2 mark the start of each  $dB/dt$  within the TLA event and the solid dots mark the end. Note that axes in all plots of Figure 2 have been adjusted by subtracting the mean  $B_{x,y,z}$  value from the interval, so the magnitude of the rate of change of the magnetic field is still to scale.

148

149

150

151

152

153

154

155

156

157

158

159

160

We expected to find many events occurring due to SCs as they have been considered the primary driver for the most rapid GICs (Kataoka & Ngwira, 2016). However, we found only one SC-related event, shown in Figure 2a. This is the only SC-related event despite five recorded SCs in 2015 that occurred when the MACCS stations were located on the dayside; the other four SCs caused  $dB/dt$ s at the MACCS stations that all lasted less than 60 seconds but did not exceed the 6 nT/s threshold. This TLA event started on 22 June 2015 at 18:33:22 UT (12:41:22 MLT, at RBY), just seconds after a large CME reached Earth causing an SSC at 18:33 UT. The largest  $dB/dt$  signature of the entire data set occurred in this event at RBY in the y-component, lasting 9.5 seconds with a magnitude of -33.49 nT/s. The  $dB/dt$ s measured in the y- and z-components at PGG and CDR all last 10.5 seconds or less, with the shortest event in the y-component at CDR with a magnitude of 13.3 nT/s and lasting just 5 seconds. All four stations were on the dayside during the time of the event.

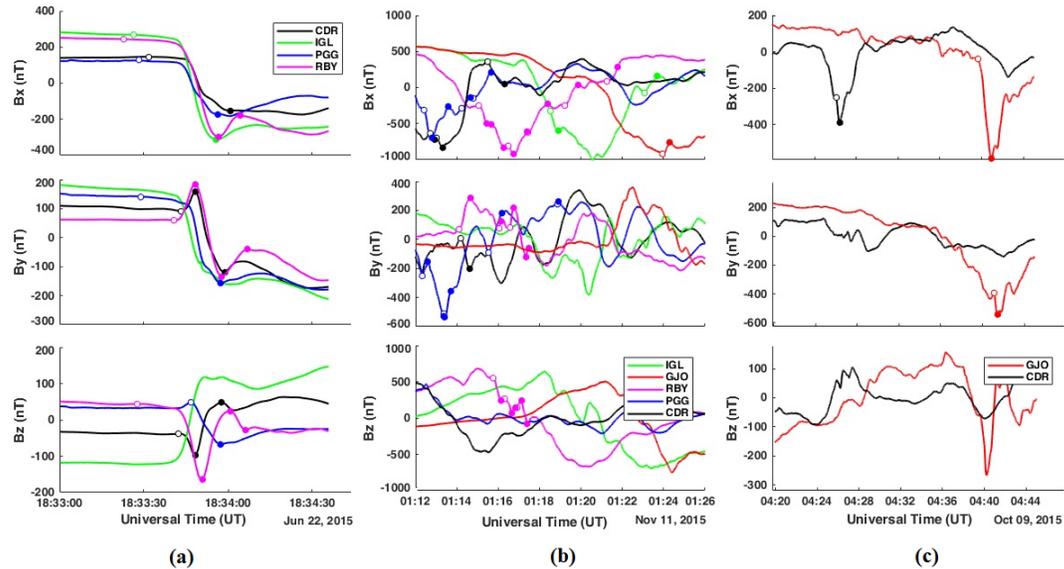


Figure 2: (a): A TLA event that occurred on 22 June 2015. (b): An event that occurred on 11 November 2015. (c) An event that occurred on 9 October 2015. All three panels show the x, y and z components of the surface magnetic field from top to bottom, respectively. Hollow circles mark the start of a  $dB/dt$  signature and the dots mark the end.

161

162

163

164

165

Shown in Figure 2b is an event that occurred on 11 November 2015 beginning at 01:12:20 UT (21:22:36 MLT of 10 November 2015). This event consists of 34  $dB/dt$ s measured at all but the NAN station. Of these 34  $dB/dt$ s, six have magnitude greater than 10 nT/s and five have duration < 10 seconds. One of the largest  $dB/dt$ s (16.2 nT/s) was measured at PGG at 1:13:21 UT in the y-component and lasted only 1 second. The over-

all event lasts about 10 minutes and occurs within a larger, longer ( $\sim 1$  hour) magnetic impulse event (MIE) that is investigated by Engebretson et al. (2019). The MIE and the TLA event are not associated with a geomagnetic storm, although a substorm onset occurred at 01:07 UT, about 5 minutes prior to the start of the event. The MIE was preceded by a steady magnetic field for at least an hour prior to the start of the disturbance around 00:40 UT.

Finally, Figure 2c shows a TLA event on 9 October 2015 starting at 04:26:06 UT at the CDR station (23:31:06 MLT of 8 October 2015) where  $B_x$  decreases by 135.9 nT in 21 seconds ( $dB_x/dt = -6.46$  nT/s). Then about 14 minutes later, two similar signatures occurred at GJO: a  $dB_x/dt$  of  $-6.87$  nT/s at 04:49:37 UT and a  $dB_y/dt$  of  $-6.52$  nT/s at 04:41:05 UT. Note, however, that the  $dB_x/dt$  at GJO actually lasted 80 seconds, this is one of the signatures identified when extending the upper threshold for the duration of the sample in the search algorithm to 5 minutes rather than 60 seconds. This TLA event occurred on the second day of recovery from a moderate geomagnetic storm (the SuperMAG Ring Current (SMR) index reached  $-123$  nT in hour 23 of 7 October but recovered to around  $-34$  nT during the hour of the event on 9 October) and there were marked substorm onsets occurring at 04:13 UT and 4:34 UT. Further, a nighttime MIE was identified at RBY at 04:37 UT but was not identified at CDR (note that GJO, the other station that measured this TLA event was not one of the stations used in the statistical study of Engebretson et al., 2019). There did occur a nighttime MIE measured at CDR later on at 22:00 UT of 9 October, and while no TLA signatures were identified at CDR during this time, a TLA event ( $dB_x/dt = -10.43$  nT/s) was identified at PGG at 21:56:02 UT, preceding that MIE by several seconds.

We demonstrate the existence of significant magnetic disturbances with timescale  $\leq 10$  seconds in nine of the 40 TLA events identified. In five of these events, the shortest-timescale signatures exhibit the largest amplitude disturbances of the entire set of events ( $|dB/dt| \geq 10$  nT/s). Further, there are seven cases in which these signatures precede a larger, longer timescale ( $< 60$  seconds)  $dB/dt$ . Examples of these signatures can be seen in Figure 2a ( $B_y$  at RBY:  $dB_y/dt = -33.49$  nT/s), and in Figure 2b (the decrease in  $B_y$  at CDR at 18:33:43 UT lasts for 5 seconds and has rate of change of  $13.23$  nT/s; the two signatures in the  $z$ -component at CDR last 6 and 9.5 seconds with magnitudes of  $-9.85$  and  $15.28$  nT/s, respectively).

#### 4 Spatial and Temporal Characteristics and Space Weather Dependence

Of the 40 identified events, 27.5% consist of at least one  $dB/dt$  signature with magnitude exceeding 10 nT/s and half of these occurred within an event that has at least one other  $|dB/dt| \geq 10$  nT/s. These ten largest events were measured primarily between  $73^\circ$  and  $76^\circ$  CGM latitude at the PGG and CDR stations: PGG and CDR not only recorded the majority of the largest events but a substantial fraction (50% and 43%, respectively) of events in general. The GJO ( $76.86^\circ$ ) station recorded 9 events and RBY ( $75.62^\circ$ ) and IGL ( $78.63^\circ$ ) recorded 3 and 4 events, respectively. The southern-most station, NAN ( $65.67^\circ$ ), recorded just two events that were not recorded at any other station. In fact, 75% of the events were measured locally at only one station (the average, absolute distance from one station to the nearest station is  $\sim 580$  km. Note this average excludes NAN as it is the lowest latitude station with only two locally recorded events). Of the other 25% of events measured at more than one station, 4 were recorded relatively simultaneously (as shown in Figures 2a and 2b) while 6 other events had  $dB/dts$  at more than one station delayed by at least 2 minutes (and at most 14 minutes, shown in Figure 2c).

TLA events occurred substantially more often in the Fall-Winter months with exactly 60% of events occurring in October through December. To illustrate the occurrence of TLA events as a function of magnetic local time as well as the association to geomag-

217 netic storms and substorms, Figure 3 shows the maximum  $dB/dt$  of each TLA event through-  
 218 out 2015 as a function of MLT. The events that occurred between 18-6 MLT are plot-  
 219 ted as squares with opacity according to temporal proximity of prior substorm onset: the  
 220 black squares signify that the event started within 15 minutes after the nearest substorm  
 221 onset and during nighttime hours of 18-6 MLT, the grey squares are events that occurred  
 222 15-30 minutes after substorm onset and the white squares occurred more than 30 min-  
 223 utes after the nearest substorm onset (daytime events were automatically marked as white  
 224 squares). These onset delays were determined with the SuperMAG Newell and Gjerloev  
 225 (2011) Substorm Event List ([supermag.jhuapl.edu/substorms/](http://supermag.jhuapl.edu/substorms/)). The bars extending from  
 226 some of the squares in Figure 3 signify the full duration of the event if it consisted of mul-  
 227 tiple  $dB/dts$ , showing at what point throughout the event that the maximum  $dB/dt$  oc-  
 228 curred. Only three events occurred in the commencement or main phase of a geomag-  
 229 netic storm, these are labeled in Figure 3. There are also five events that occurred on  
 230 the first day of recovery from a geomagnetic storm and four events that occurred on the  
 231 second day of recovery.

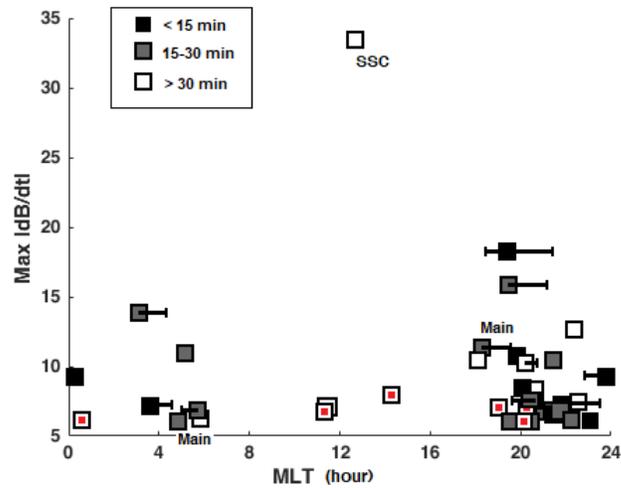


Figure 3: Maximum  $dB/dt$  as a function of magnetic local time (MLT) of each TLA event found in 2015. The bars extended from some squares signifies the duration of an event with multiple  $dB/dts$ . The opacity of squares is based on the temporal proximity after the nearest substorm onset.

232 Figure 3 shows that a vast majority (90%) of events occur at nighttime between  
 233 18-6 MLT with peak number of events (70%) in the pre-midnight sector from 18-24 MLT.  
 234 A large number of the events (65%) occurred within 30 minutes of substorm onset, but  
 235 it is clear from Figure 3 that not all of the nighttime events show this association to sub-  
 236 storm onsets (see white squares occurring at nighttime). While there is a strong asso-  
 237 ciation of TLA events to substorm onsets, 30% of events occurred more than 30 minutes  
 238 after a substorm onset, with a small subset of events (6) that occurred more than 2 hours  
 239 after substorm onset. Figure 3 also shows that the eleven *largest* TLA events ( $\geq 10$  nT/s)  
 240 are more likely to occur between 18-24 MLT, but these are not necessarily more likely  
 241 to occur within 30 minutes of substorm onset as about half of the set of largest events  
 242 occurred more than 30 minutes after. As previously stated, five of these eleven largest  
 243 events also have signatures lasting 10 seconds or less, with magnitude exceeding 10 nT/s.  
 244 Comparison to the nighttime MIE events of Engebretson et al., (2019) found that 70%  
 245 are related: either preceding the MIE within 30 minutes or occurring within the longer-  
 246 timescale perturbation. Eight of the largest amplitude events were associated to a night-

247 time MIE. While the set of events exhibit a clear association to substorm activity and  
 248 nighttime MIEs, there exists a subset of TLA events (15%) that occur more than 30 min-  
 249 utes prior to a nighttime MIE, more than 30 minutes after a substorm onset, and dur-  
 250 ing relatively quiet geomagnetic conditions (i.e. not during any phase of a geomagnetic  
 251 storm, nor occurring within two days of recovery), we classify these as unrelated events.  
 252 These six events are expressed in Figure 3 as squares with red dots in the center. None  
 253 of these unrelated events are in the set of largest disturbances, but they do show more  
 254 temporal spread than the majority of events as two of these unrelated events are within  
 255 the only four events that occurred during the daytime.

## 256 5 Discussion and Conclusions

257 While we identified a fairly small number of TLA events, the set exhibits several  
 258 cases of large-amplitude ( $\geq 10$  nT/s) and very short-timescale ( $\leq 10$  s) disturbances. We  
 259 found that SCs were not the main driver for these transient magnetic disturbances, al-  
 260 though the large SSC that occurred on 22 June did cause the largest amplitude pertur-  
 261 bation, it was the only TLA event associated to an SC despite many occurring over the  
 262 course of the year. Rather, TLA  $dB/dts$  occurred most often during local magnetic night-  
 263 time, with the highest frequency of events in the pre-midnight sector from 18-24 MLT.  
 264 There is a clear association of these events to the onset of substorms as well as associ-  
 265 ation to nighttime MIEs (about two-thirds occurring at nighttime within 30 minutes of  
 266 substorm onset and about two-thirds related to MIEs), but there is not a perfect cor-  
 267 relation between nighttime events and substorm-related events (i.e. not all nighttime events  
 268 are substorm-related). Further, the relationship with substorm onsets appears to be a  
 269 complicated one, as several events occurred multiple hours after the nearest substorm  
 270 onset; this association will be investigated further in a future study extending the search  
 271 for TLA events to many other stations and for a longer period of time.

272 In addition to a clear association to substorm onsets, we found that a majority of  
 273 our events either preceded or occurred within a nighttime MIE (Engebretson et al., 2019).  
 274 These nighttime MIEs are large-amplitude magnetic disturbances with 5-10 minute timescale  
 275 occurring in this region of north-east Canada, the study surveyed MIEs from 2014-2017.  
 276 Like MIEs, the TLA events identified were often but not always associated with substorms  
 277 on a similar two-thirds basis. Using the spherical elementary current systems (SECS)  
 278 method (Amm & Viljanen, 1999) and the implementation of this technique by Weygand  
 279 et al. (2011), a superposed epoch analysis was conducted to investigate the average equiv-  
 280 alent ionospheric currents (EIC) and inferred field-aligned currents (FAC) during 21 night-  
 281 time MIEs that occurred at CDR from mid-2014 to 2016. Engebretson et al. (2019a) found  
 282 that the largest of these MIEs were associated to intense westward ionospheric currents  
 283 100 km above CDR, coinciding with a region of shear between upward and downward  
 284 FAC. They also found that the largest horizontal  $dB/dts$  occurred slightly south of CDR  
 285 in a localized region of  $\sim 275$  km. Our TLA events show some similarities to these MIEs:  
 286 1) Of all six stations, the PGG and CDR stations measured the greatest number of events  
 287 as well as the largest-amplitude events ( $|dB/dt| \geq 10$  nT/s) and 2) we found only nine  
 288 events that were measured by more than one station, so the majority of our events ( $\sim 75\%$ )  
 289 were measured locally at just one station. The localized nature of many TLA disturbances  
 290 implies that the source currents are localized in the ionosphere (Boteler & Beek, 1999).  
 291 More recent research has found extreme local enhancements of the geoelectric field with  
 292 spatial scale  $\sim 250$ -1600 km (Ngwira et al., 2015); these localized peak geoelectric fields  
 293 occur during geomagnetic storms but the exact physical mechanism responsible for gen-  
 294 erating them is yet unknown. The TLA events studied in this paper are consistent with  
 295 that of Ngwira et al. (2015), but also occur independent of geomagnetic storms (as well  
 296 as auroral substorms and MIEs). Our future work will expand the data set to include  
 297 more stations over an extended period of time and will include a superposed epoch anal-  
 298 ysis to investigate the ionospheric activity during TLA events.

299 In order to better understand our events in the context of these MIEs, we extended  
300 the upper threshold of the search algorithm to identify disturbances lasting up to 5 min-  
301 utes with magnitude of 6 nT/s or greater. We found 25 additional  $dB/dt$ s that were all  
302 related to TLA events that we had already identified. Interestingly, only one signature  
303 lasted slightly longer than 2 minutes. We hypothesized that the absence of magnetic per-  
304 turbations in the 2-5 minute timescale range could be due to algorithm bias. Because  
305 the method of the routine searches for changes in the direction of the slope ( $dB/dt$ ) with  
306 the condition that the change last for at least 1 second and we used raw magnetic field  
307 data without any smoothing method, it was possible that the algorithm could be miss-  
308 ing collections of  $dB/dt$  signatures lasting 2-5 minutes because there are shorter timescale  
309 variations occurring within them that did not meet the threshold of 6 nT/s. To test this  
310 theory, we applied a 10-point sliding average filter on the magnetic field data so that any  
311 of these shorter variations would be smoothed over, then ran the search algorithm for  
312 disturbances lasting up to 5 minutes again. Engebretson et al. (2019) also used a 10-point  
313 sliding average smoothing on the data. We found when the data were smoothed around  
314 10-points, the algorithm identified all the same events as the raw data and identified 17  
315 new events. All the events with signatures lasting  $> 60$  seconds were the same apart from  
316 one case where the smoothed data marked the magnetic field response to the SSC at RBY  
317 as a disturbance lasting 60.5 seconds rather than 34 seconds. This occurred in many cases  
318 where the smoothed data identified the same signatures as longer e vents; because the  
319 algorithm searches for changes in the direction of the  $dB/dt$ , the 10-point smoothing was  
320 altering the exact moment that the slope changed sign and the signature started or ended.  
321 While the smoothing method resulted in many signatures marked as having longer du-  
322 ration, there was still only a small number of  $dB/dt$ s with  $> 1$  minute timescale (32 as  
323 opposed to 25 with raw data) and the longest signature lasted 147 seconds. By compar-  
324 ing our results with smoothed data, we verified the methodology of the algorithm and  
325 determined that the absence of large-amplitude ( $\geq 6$  nT/s) magnetic disturbances with  
326 timescale  $\sim 2.5$ -5 minutes is not due to algorithm bias. This finding suggests that all longer-  
327 timescale magnetic perturbations at these stations consist of more rapid variations last-  
328 ing less than 2.5 minutes, with a vast majority  $< 60$  seconds.

329 While TLA events show a clear association with substorm activity as well as many  
330 shared characteristics with nighttime MIEs, TLA  $dB/dt$  events are not consistently re-  
331 lated to these space weather events. We found a small subset of TLA events that are un-  
332 related to geomagnetic storms, auroral substorms and nighttime MIEs. TLA events show  
333 a similar localized behavior with a weak association to geomagnetic storms, suggesting  
334 that there are other physical mechanisms, even beyond substorms, for localized peak en-  
335 hancements in the geoelectric field (roughly proportional to the  $dB/dt$ ). What we learned  
336 from the error analysis of this study is that a common smoothing method on the data  
337 altered the timing and amplitude of the events, suggesting that the short-timescale na-  
338 ture of the geomagnetic field could often be removed with common data processing meth-  
339 ods or missed altogether with 1-minute or even 10-second averaged magnetic field data.  
340 Finally, we show that these signatures can have amplitude of the same order as longer-  
341 timescale events that are relevant to GICs. Our future work will include a statistical anal-  
342 ysis on an expanded set of TLA events as well as an investigation of the geoelectric fields  
343 resulting from TLA  $dB/dt$  events in order to assess the potential threat they pose on tech-  
344 nological infrastructure on Earth.

## 345 Acknowledgments

346 The tables of TLA events and space weather association information, as well as the al-  
347 gorithm developed for this research are available on the University of Michigan Deep Blue  
348 data repository ([doi.org/10.7302/9t46-0092](https://doi.org/10.7302/9t46-0092)).

349 The authors thank the MACCS team for data (space.augsburg.edu/maccs/). MACCS  
 350 is operated by the University of Michigan and Augsburg University and funded by the  
 351 U.S. National Science Foundation via grants AGS-2013433 and AGS-2013648.

352 We gratefully acknowledge the SuperMAG collaborators (supermag.jhuapl.edu/info  
 353 /?page=acknowledgement)

354 The results presented in this paper rely on the SC Event List calculated and made  
 355 available by Observatori de L'Ebre, Spain from data collected at magnetic observato-  
 356 ries. We thank the involved national institutes, the INTERMAGNET network and ISGI  
 357 (isgi.unistra.fr).

## 358 References

- 359 Amm, O., & Viljanen, A. (1999). Ionospheric disturbance magnetic field contin-  
 360 uation from the ground to the ionosphere using spherical elementary current  
 361 systems. *Earth, Planets and Space*, *51*(6), 431–440. doi: 10.1186/BF03352247
- 362 Boteler, D. H., & Beek, G. J. V. (1999). August 4, 1972 revisited: A new look at the  
 363 geomagnetic disturbance that caused the L4 cable system outage. *Geophysical*  
 364 *Research Letters*, *26*(5), 577–580.
- 365 Boteler, D. H., Pirjola, R. J., & Nevanlinna, H. (1998). The effects of geomagnetic  
 366 disturbances on electrical systems at the Earth's surface. *Advances in Space*  
 367 *Research*, *22*(1), 17–27. doi: 10.1016/S0273-1177(97)01096-X
- 368 Engebretson, M. J., Hughes, W. J., Alford, J. L., Zesta, E., Cahill, L. J., Arnoldy,  
 369 R. L., & Reeves, G. D. (1995). Magnetometer array for cusp and cleft studies  
 370 observations of the spatial extent of broadband ULF magnetic pulsations at  
 371 cusp/cleft latitudes. *Journal of Geophysical Research*, *100*(A10), 19371. doi:  
 372 10.1029/95ja00768
- 373 Engebretson, M. J., Pilipenko, V. A., Ahmed, L. Y., Posch, J. L., Steinmetz, E. S.,  
 374 Moldwin, M. B., ... Vorobev, A. V. (2019). Nighttime Magnetic Perturba-  
 375 tion Events Observed in Arctic Canada: 1. Survey and Statistical Analysis.  
 376 *Journal of Geophysical Research: Space Physics*, *124*(9), 7442–7458. doi:  
 377 10.1029/2019JA026794
- 378 Engebretson, M. J., Pilipenko, V. A., Steinmetz, E. S., & Moldwin, M. B. (2021).  
 379 Nighttime magnetic perturbation events observed in Arctic Canada : 3 . Oc-  
 380 currence and amplitude as functions of magnetic latitude , local time , and  
 381 magnetic disturbance indices.
- 382 Engebretson, M. J., Steinmetz, E. S., Posch, J. L., Pilipenko, V. A., Moldwin, M. B.,  
 383 Connors, M. G., ... Kistler, L. M. (2019). Nighttime Magnetic Perturba-  
 384 tion Events Observed in Arctic Canada: 2. Multiple-Instrument Observations.  
 385 *Journal of Geophysical Research: Space Physics*, *124*(9), 7459–7476. doi:  
 386 10.1029/2019JA026797
- 387 Gjerloev, J. W. (2012). The SuperMAG data processing technique. *Journal of Geo-*  
 388 *physical Research: Space Physics*, *117*(9), 1–19. doi: 10.1029/2012JA017683
- 389 Gombosi, T. I., Baker, D. N., Balogh, A., Erickson, P. J., Huba, J. D., & Lanzerotti,  
 390 L. J. (2017). Anthropogenic Space Weather. *Space Science Reviews*, *212*(3-4),  
 391 985–1039. Retrieved from <http://dx.doi.org/10.1007/s11214-017-0357-5>  
 392 doi: 10.1007/s11214-017-0357-5
- 393 Hughes, W., & Engebretson, M. (1997). MACCS: Magnetometer array for cusp and  
 394 cleft studies. (1).
- 395 Kappenman, J. G. (2006). Great geomagnetic storms and extreme impulsive geo-  
 396 magnetic field disturbance events - An analysis of observational evidence  
 397 including the great storm of May 1921. *Advances in Space Research*, *38*(2),  
 398 188–199. doi: 10.1016/j.asr.2005.08.055
- 399 Kataoka, R., & Ngwira, C. (2016). Extreme geomagnetically induced currents. ,  
 400 *3*(1). Retrieved from <http://dx.doi.org/10.1186/s40645-016-0101-x> doi:

- 10.1186/s40645-016-0101-x
- 401 Khomutov, S. Y., Mandrikova, O. V., Budilova, E. A., Arora, K., & Manjula,  
402 L. (2017). Noise in raw data from magnetic observatories. *Geoscientific Instrumentation, Methods and Data Systems*, 6(2), 329–343. doi:  
403 10.5194/gi-6-329-2017
- 404  
405
- 406 Newell, P. T., & Gjerloev, J. W. (2011). Substorm and magnetosphere characteristic  
407 scales inferred from the SuperMAG auroral electrojet indices. *Journal of Geo-*  
408 *physical Research: Space Physics*, 116(12), 1–15. doi: 10.1029/2011JA016936
- 409 Newell, P. T., & Gjerloev, J. W. (2012). SuperMAG-based partial ring current in-  
410 dices. *Journal of Geophysical Research: Space Physics*, 117(5), 1–15. doi: 10  
411 .1029/2012JA017586
- 412 Nguyen, N., Muller, P., & Collin, J. (2020). The Statistical Analysis of Noise  
413 in Triaxial Magnetometers and Calibration Procedure. *2019 16th Work-*  
414 *shop on Positioning, Navigation and Communications (WPNC)*, 1–6. doi:  
415 10.1109/wpnc47567.2019.8970255
- 416 Ngwira, C. M., Pulkkinen, A., Kuznetsova, M. M., & Glocer, A. (2014). Modeling  
417 extreme "carrington-type" space weather events using three-dimensional global  
418 MHD simulations. *Journal of Geophysical Research: Space Physics*, 119(6),  
419 4456–4474. doi: 10.1002/2013JA019661
- 420 Ngwira, C. M., Pulkkinen, A., McKinnell, L. A., & Cilliers, P. J. (2008). Improved  
421 modeling of geomagnetically induced currents in the South African power  
422 network. *Space Weather*, 6(11). doi: 10.1029/2008SW000408
- 423 Ngwira, C. M., Pulkkinen, A. A., Bernabeu, E., Eichner, J., Viljanen, A., & Crow-  
424 ley, G. (2015). Characteristics of extreme geoelectric fields and their possible  
425 causes: Localized peak enhancements. *Geophysical Research Letters*, 42(17),  
426 6916–6921. doi: 10.1002/2015GL065061
- 427 Pulkkinen, A., Bernabeu, E., Thomson, A., Viljanen, A., Pirjola, R., Boteler, D.,  
428 ... MacAlester, M. (2017). Geomagnetically induced currents: Science, en-  
429 gineering, and applications readiness. *Space Weather*, 15(7), 828–856. doi:  
430 10.1002/2016SW001501
- 431 Pulkkinen, A., Rastätter, L., Kuznetsova, M., Singer, H., Balch, C., Weimer, D., ...  
432 Weigel, R. (2013). Community-wide validation of geospace model ground mag-  
433 netic field perturbation predictions to support model transition to operations.  
434 *Space Weather*, 11(6), 369–385. doi: 10.1002/swe.20056
- 435 Pulkkinen, A., Viljanen, A., & Pirjola, R. (2006). Estimation of geomagnetically in-  
436 duced current levels from different input data. *Space Weather*, 4(8). doi: 10  
437 .1029/2006SW000229
- 438 Rivera, M. K., Backhaus, S. N., Woodroffe, J. R., Henderson, M. G., Bos, R. J.,  
439 Nelson, E. M., & Kelic, A. (2016). *Emp/gmd phase 0 report, a review*  
440 *of emp hazard environments and impacts* (Tech. Rep.). Retrieved from  
441 [http://permalink.lanl.gov/object/view?what=info:lanl-repo/](http://permalink.lanl.gov/object/view?what=info:lanl-repo/lareport/LA-UR-16-28380)  
442 [lareport/LA-UR-16-28380](http://permalink.lanl.gov/object/view?what=info:lanl-repo/lareport/LA-UR-16-28380) (Dataset: RASSTI)
- 443 Simpson, J. J. (2011). On the possibility of high-level transient coronal mass  
444 ejection-induced ionospheric current coupling to electric power grids. *Jour-*  
445 *nal of Geophysical Research: Space Physics*, 116(11), 1–12. doi: 10.1029/  
446 2011JA016830
- 447 Tóth, G., Sokolov, I. V., Gombosi, T. I., Chesney, D. R., Clauer, C. R., De Zeeuw,  
448 D. L., ... Kóta, J. (2005). Space weather modeling framework: A new tool for  
449 the space science community. *Journal of Geophysical Research: Space Physics*,  
450 110(A12), 1–21. doi: 10.1029/2005JA011126
- 451 Weygand, J. M., Amm, O., Viljanen, A., Angelopoulos, V., Murr, D., Engebret-  
452 son, M. J., ... Mann, I. (2011). Application and validation of the spherical  
453 elementary currents systems technique for deriving ionospheric equivalent  
454 currents with the North American and Greenland ground magnetometer ar-  
455 rays. *Journal of Geophysical Research: Space Physics*, 116(3), 1–8. doi:



Figure 1.

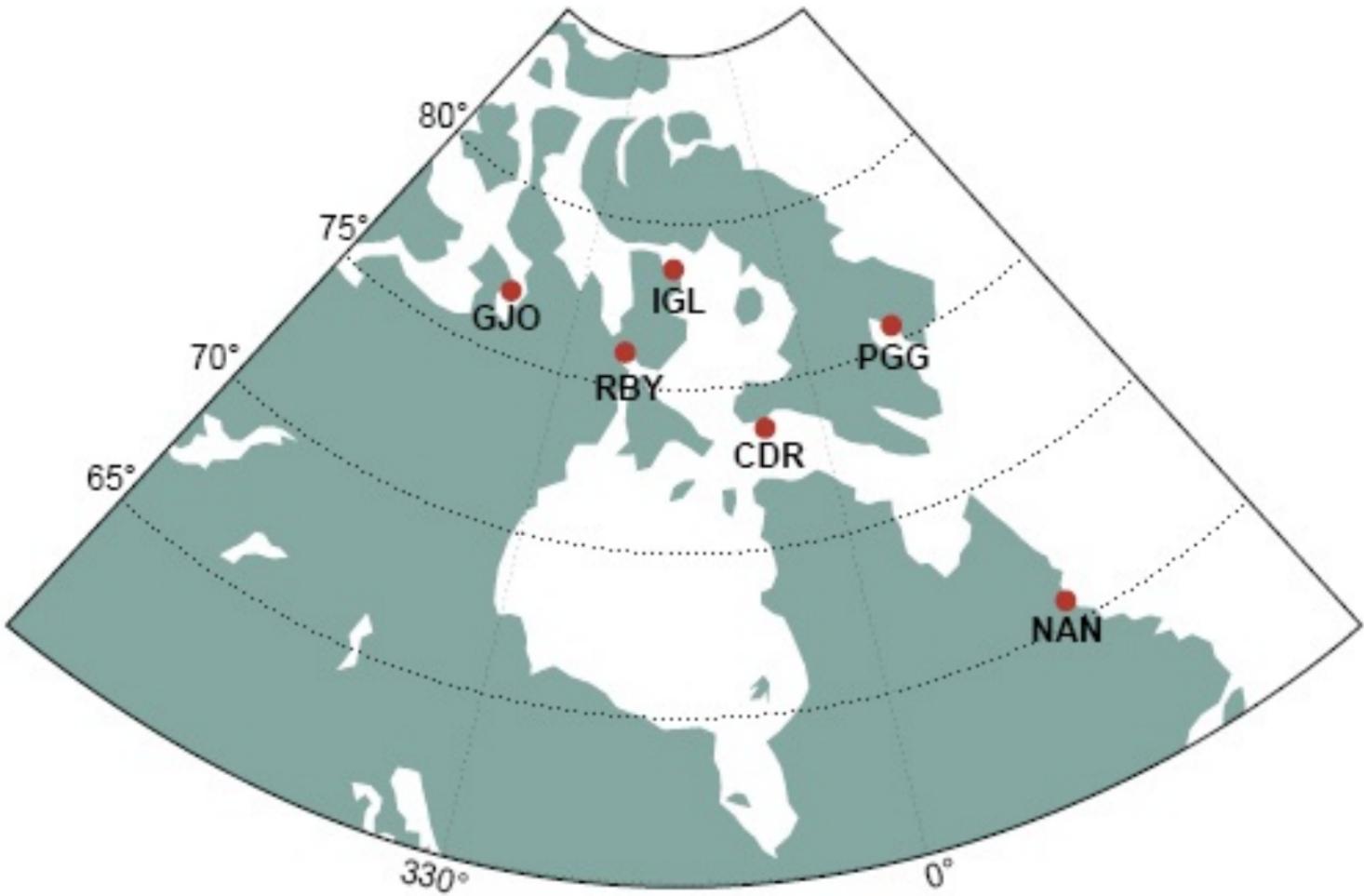
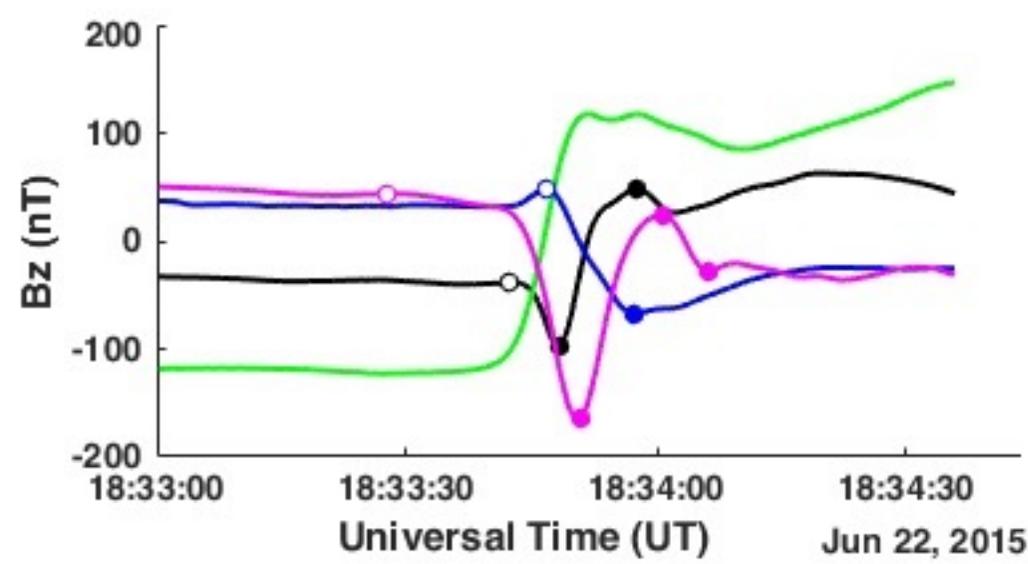
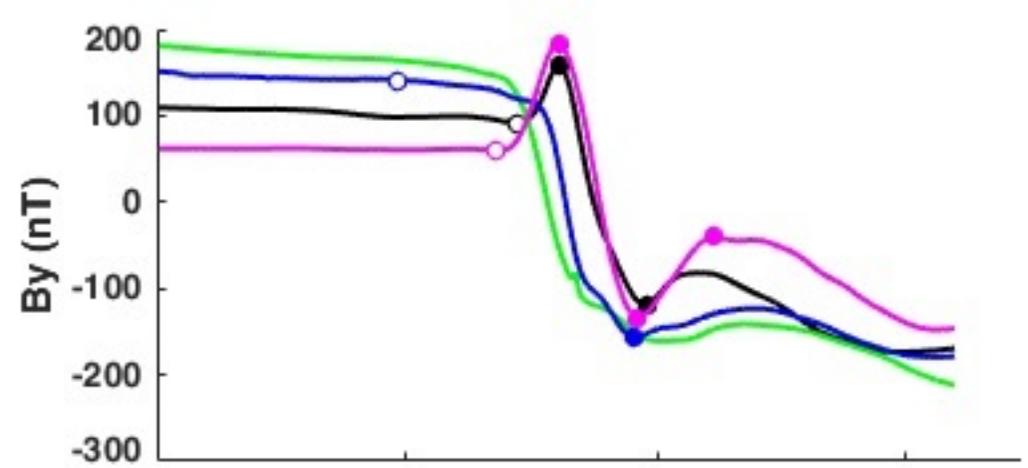
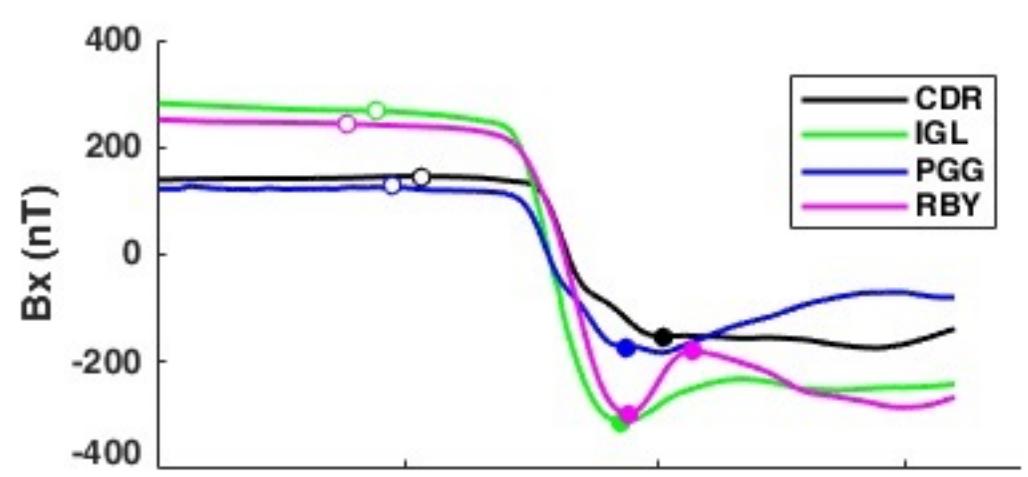
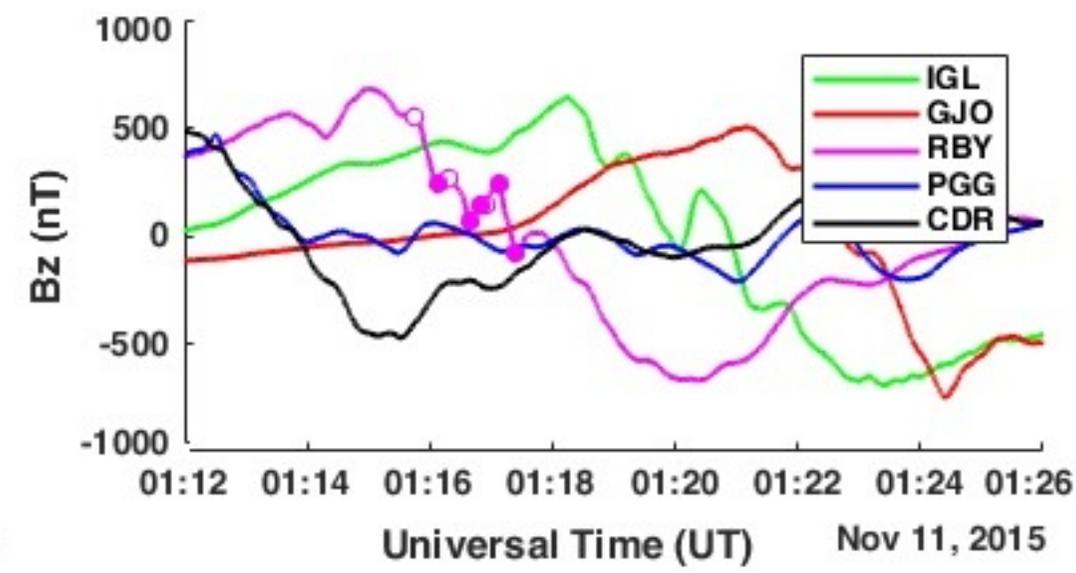
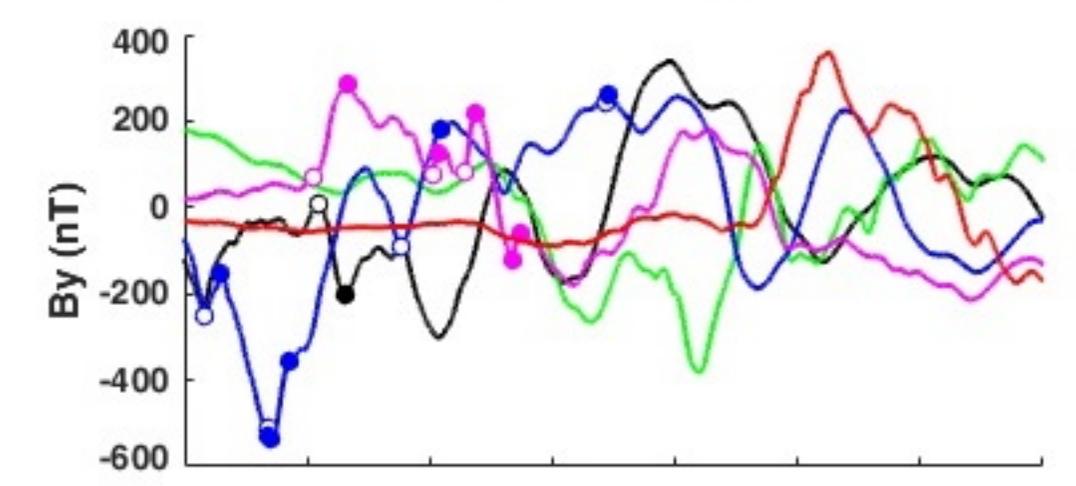
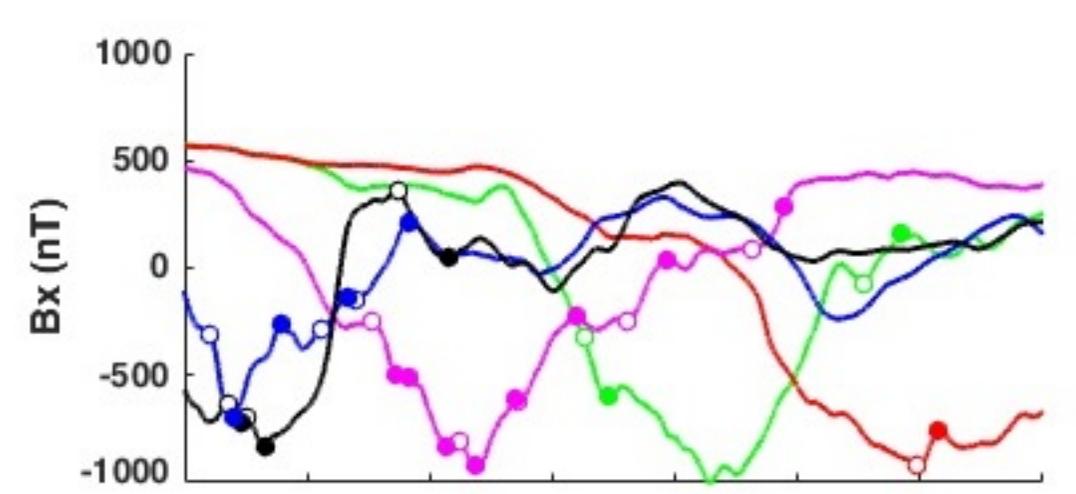


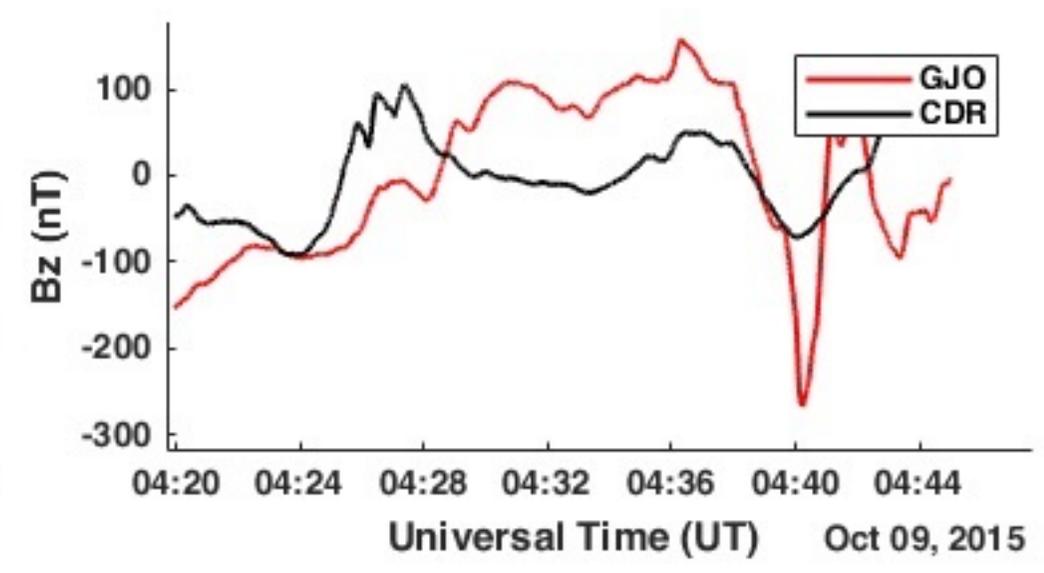
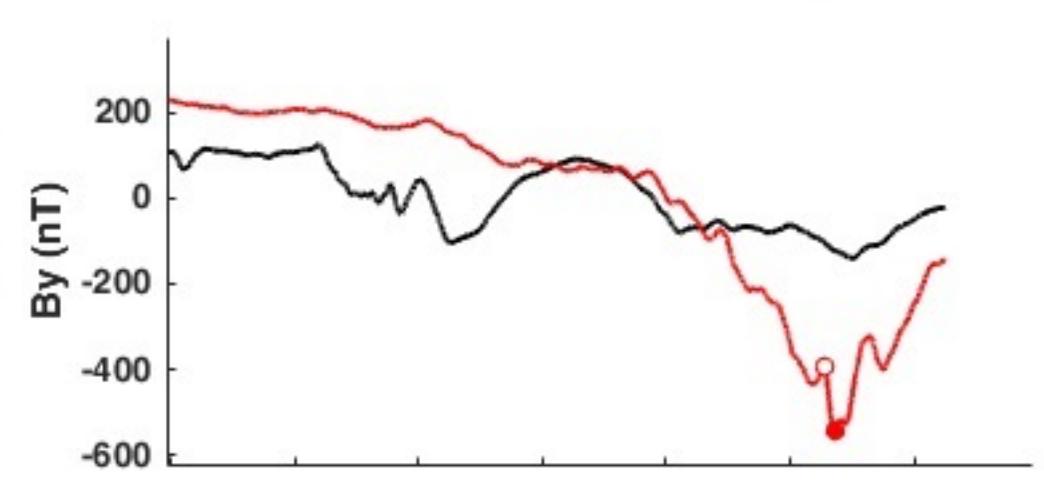
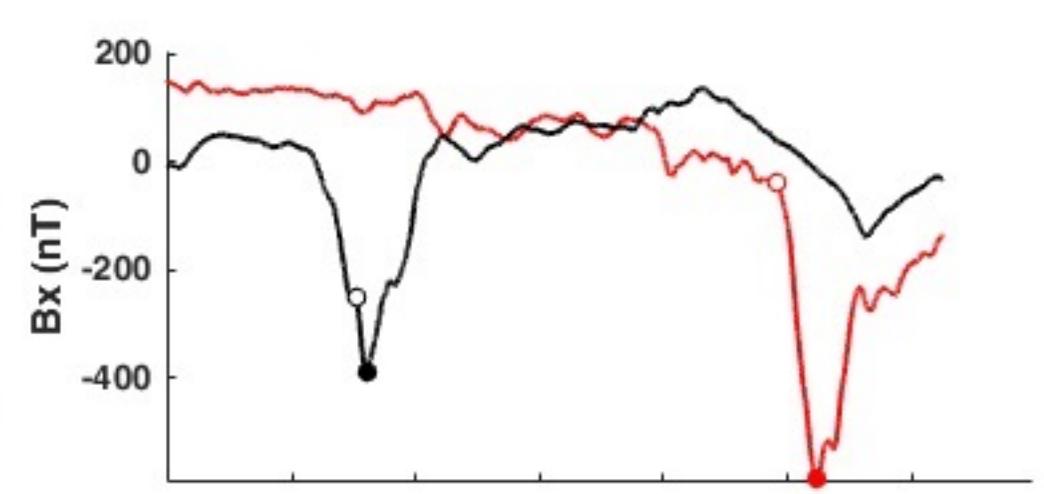
Figure 2.



(a)



(b)



(c)

Figure 3.

