# GOLD Observations of Longitudinal Variations in the Nighttime Equatorial Ionization Anomaly (EIA) Crests Latitudes

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March 16, 2023

- 1 GOLD Observations of Longitudinal Variations in the Nighttime Equatorial Ionization Anomaly
- 2 (EIA) Crests' Latitudes

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# 12 Key Points:

- The average nighttime EIA crests' latitudes observed by GOLD during 2020 equinoxes
   and December solstice have a longitudinal dependence
- Crests' latitude dependence on longitude changes near 47°W geographic, where the magnetic and geographic equators cross
- Crests' latitudes over 75°W-10°E geographic longitude during equinoxes and winter
   solstice depend on the subsolar point magnetic latitude

#### 20 Abstract

21 Each day the GOLD (Global-scale Observations of the Limb and Disk) imager observes 22 the Equatorial Ionization Anomaly (EIA) near sunset from  $\sim 10^{\circ}$  E to  $\sim 80^{\circ}$  W geographic longitude. 23 Most images cover ~45° of longitude (~3 hours), and most longitudes are observed multiple times. 24 Monthly averages of EIA crests' latitude (EIA lats) versus longitude during March, September and 25 December 2020 have been analyzed. The EIA lats reflect the combined influence of winds, solar 26 radiation, and fields (electric and magnetic) in the equatorial region. Winter solstice differs 27 significantly from the equinoxes, which are similar, but there are notable similarities between all 28 three. The similarities in the EIA lats during the seasons examined indicates that the magnetic 29 equator to subsolar point separation influences them in all three seasons and that it has a more 30 distinct, possibly more significant, influence than winds on the average latitudes.

31

### 32 Plain Language Summary

33 Each day the GOLD (Global-scale Observations of the Limb and Disk) imager observes 34 the nighttime Equatorial Ionization Anomaly (EIA) near sunset as the terminator progresses 35 westward from Africa to across South America. Most images cover ~45° of longitude (~3 hours 36 of local time), and most longitudes are observed multiple times. In 2020 seasonal averages of the crests' latitude versus longitude during the equinoxes (March and September) and winter solstice 37 38 (December) show significant, important similarities. The observed latitude versus longitude 39 dependence in all three show a dependence on the distance between the magnetic equator and the 40 subsolar point. Greater knowledge of the seasonal-longitudinal dependence of the EIA crests' 41 latitudes contributes significantly to understanding the crests' response to fields (electric and 42 magnetic) and winds in the equatorial region.

#### 43 **1 Introduction**

44 Longitudinal differences in the EIA are a source of information about spatial differences 45 in the processes that influence their creation and decay. Changes in their radiance have often been linked to differences in the equatorial electrojet (EEJ) strength which are attributable to 46 47 longitudinal differences in solar tidal forcing (England et al., 2006; Fang et al., 2009; Jin et al., 48 2008; Lühr et al., 2008; Pedatella et al., 2012; Lühr & Manoj, 2013; Venkatesh et al., 2015; 49 Yamazaki and Maute, 2016; Zhou et al., 2016; Mo and Zhang, 2020). Variations in the optical 50 davglow radiances at the northern crest's latitudes over Indian longitudes show a similar 51 relationship with the EEJ strength (Karan et al., 2016). The small scale (3°-7°) longitudinal 52 differences in the dayglow radiances have been attributed to variations of the equatorial electric 53 fields (Karan et al., 2017). Differences in the EEJ and and tides might also alter the latitudes. While 54 there have been multiple studies of the EIA radiances, the longitudinal dependence of the EIA 55 crests' latitudes (EIA lats), where the radiances peak, is less studied. Observing such effects is 56 complicated by longitudinal variation in both the location of the magnetic equator and the magnetic 57 declination at the equator as well as neutral winds which are additional influences on the EIA.

58 Significant short term variability is seen in the EIA lats. Recently, Rodríguez-Zuluaga, et 59 al. (2021) found that the EIA lats observed by GOLD sometimes exhibit a mesoscale structure (a 60 few thousand km scale size), indicating that they may be influenced by waves from the lower 61 atmosphere. Other periodic, longitudinal variations in the EIA crests' brightnesses are also seen 62 (e.g., Eastes et al., 2019, Fig 3; Gan et al 2020, Fig. 2), as are occasional decreases in their 63 separation (e.g., Basu et al., 2009). Such decreases have sometimes been sufficient to reduce the

two crests typically observed to a single crest located near the magnetic equator (e.g., Aa et al.2022).

66 The EIA around the South Atlantic Anomaly (SAA), where largest low-latitude deviations in Earth's magnetic field occur, is one of the most interesting longitude regions. The magnetic 67 68 equator is north of the geographic equator near Africa and south of it over South America, 69 Unfortunately, concurrent observations of the EIA across the SAA region have been insufficient 70 for a comprehensive study. Both ground and space based observations have limitations - there are 71 few ground based observations over Africa and radiation in the South Atlantic Anomaly (SAA) 72 can interfere with space-based observations by low Earth orbiting satellites (e.g., Kil et al., 2006). 73 These same longitudes, in and near the SAA, are observed during GOLD's nighttime disk imaging. 74 Each night GOLD, whose orbit places it safely above radiation in the SAA, routinely images the 75 EIA crests throughout this region. These images provide an opportunity to study the EIA and, 76 possibly, to better understand it at all longitudes.

77 Every evening the GOLD mission observes the same region of the EIA near sunset with 78 sufficient signal-to-noise to locate the EIA crests over South America, the Atlantic Ocean, and 79 West Africa. This paper presents our analysis of EIA lats versus longitude during both equinoxes 80 and December solstice in 2020, a solar minimum year with relatively quiet geomagnetic 81 conditions. Monthly averages are used to simplify the identification of systematic, global-scale 82 dependences. Such averages suppress the effects of short term variability in the EEJ and the winds. 83 Any remaining, distinct patterns should be indicative of underlying processes that influence the 84 EIA. The results indicate that the EIA lats have an often neglected dependence on the latitude of 85 the subsolar point.

## 86 **2 Data**

The observations used in the analysis are nighttime, partial disk scans by the GOLD imager, which is hosted on the SES 14 satellite in geostationary orbit at 47.5° W. The instrument and its observations have been described by McClintock et al. (2020a, b) and Eastes et al. (2017, 2020). Each of the two, independent channels contains a scan mirror and interchangeable slits, enabling each channel to observe either the northern or southern hemisphere.

92 During 2020 GOLD conducted nightside imaging observations (partial disk scans, NI1 93 mode) daily, beginning at 1700 satellite local time (SLT) (2010 UT) with a 30-min cadence, using 94 only one channel (designated as channel B), alternating between the northern and southern 95 hemispheres. At 2000 SLT (2310 UT), when both channels were used for nightside scans, both 96 hemispheres were imaged simultaneously at a 15-min cadence. This imaging continued until 2130 97 SLT (0040 UT) when observations were suspended to avoid accidentally observing the Sun. Each 98 night 24 images (each covering half the latitudes) were obtained. The medium-resolution slit used 99 during the partial-disk scans provides ~0.3 nm spectral resolution and a spatial resolution of ~93 km (96 km  $\times$  80 km) at the sub-satellite location (0° geographic latitude, 47.5° W longitude). 100 101 Observations were made at a constant angular resolution.

102 Most individual scans cover  $\sim 45^{\circ}$  in longitude,  $\sim 3$  hours in local time, just east of the sunset 103 terminator. Scans shift westward throughout the evening, observing the low latitudes from  $\sim 8^{\circ}$  E 104 to  $\sim 80^{\circ}$  W longitude. Each scan images EIA locations from  $\sim 1$  hour after sunset to later local times. 105 The observation sequence is described in detail by Karan, et al. (2020). At night the EIA typically 106 appears as bands of enhanced OI 135.6 nm emission, one on each side of the magnetic equator. An example of the EIA images (from March 30, 2020) can be seen in Figure 1(a). The OI 135.6 nm emission altitude is assumed to be 300 km, a typical altitude for the F region peak, when geolocating the pixels. An example of the EIA lats obtained from the nightly sequence of images is shown in Figure 1(b) where each data point represents a peak in a crest's brightness (the method is described in Section 3). The geographic longitude corresponding to the magnetic longitude (used during the analysis) is also shown in this figure.

113 The OI 135.6 nm emission from the EIA is produced by recombination of atomic oxygen 114 ions and electrons in the Earth's ionosphere. In the F region oxygen ions are the dominant ion 115 species. Since the emission rate varies approximately as the square of the electron density, as  $n_e \times$ 

 $n_{0+}$ , the observed radiance is dominated by emissions from near the peak in electron density and

117 is indicative of the peak plasma density,  $N_{\text{max}}$ .



**Figure 1**. (1a) Composite image constructed from portions of four nightside (NI1) images taken on March 30, 2020. The geographic latitude variation of the EIA crests with geographic longitude seen in these images can be clearly seen in the crests' peak latitudes, as shown in (1b) after mapping into magnetic coordinates. (1b) Both crests' latitudes identified in nightside images from March 30, 2020 (DOY 90) at latitudes within 21° of the magnetic equator are included in the analysis. At most longitudes there are multiple images of the crests, hence multiple latitudes are shown in 1b.

118 This paper examines the nightside, NI1 observations during March, September and 119 December 2020. The first two months are near equinox and the latter is near winter solstice in the 120 northern hemisphere, periods when bright emissions are observed from the EIA. Bright emissions 121 allow a more accurate and consistent determination of the EIA lats; brightnesses during June 122 solstice were too low to obtain good statistics or sufficient longitudinal coverage. In June crest 123 latitudes were retrieved on only 9 days and only over Africa. Geomagnetic conditions, which can 124 be a significant influence on the EIA lats, were relatively quiet during most of these three months. 125 A notable exception was a late September CIR storm (Sept. 22-30; Karan et al., 2021), which was 126 excluded from the analysis. After excluding that storm, any changes associated with the remaining 127 Dst and ap fluctuations during March, September and December (Figure 2) have little effect in 128 monthly averages.

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**Figure 2**. Geophysical data for September, December and March 2020. The Dst and ap indexes are plotted in black and red, respectively. The F10.7 index is plotted in blue. Crest latitudes during the CIR storm in late September (22-30) were excluded from the monthly average.

# 133 3 Analysis Method

134 While the GOLD data are mapped into geographic coordinates during processing, the EIA 135 crests are more readily analyzed in their natural coordinates, magnetic. Therefore, all images are 136 remapped using the quasi-dipole (QD) magnetic coordinates corresponding to the observed 137 geographic coordinates (Laundal and Richmond, 2017; Thébault et al., 2021). A one day sequence 138 of such images is shown in supplementary Figure S1. From the remapped images, the OI 135.6 nm 139 radiance as a function of magnetic latitude (mlat) is obtained for each 1° in magnetic longitude 140 (mlon). These radiances are smoothed over 3 points (i.e., 1.5° in mlat), and the maxima correspond 141 to the EIA lats used in subsequent analysis. Following this procedure, EIA lats are obtained for 142 each night. An example, from March 30, is shown in Figure 1(b).

143 While longitudes near the western edge of South America are measured only once each 144 night (see the data points near 80° W in Figure 1b), most longitudes are sampled multiple times, 145 typically 4-5 times. Consequently, multiple crest latitudes are identified at most longitudes on a 146 given night (see Figure S1). While there is some variation in the EIA lats, due to the signal-to-147 noise limitations of the data, the latitudinal changes with longitude shown in Figure 1b are 148 consistent with the images shown in Figure 1a and the other images (Figure S1) from March 30. 149 A similar correspondence is typically seen throughout March, September and December 2020. 150 Therefore, the observed latitude versus longitude dependence is primarily attributable to the EIA development prior to the earliest NI1 observations, which was ~1 hour after sunset. While the EIA 151 152 brightness typically decreases with time due to recombination, the latitude-longitude variation appears to persist from the earliest to the latest observations. During GOLD's observations (1700-153 154 2130 SLT; 2010-0040 UTC, observations end on the following day) the observed latitudes are 155 more indicative of the crests' development before 1700-2130 SLT.

156 Before further processing, the crests' latitudes in each 1° mlon bin for a day (e.g., as in 157 Figure 1b) were filtered to remove outliers (crest latitudes likely attributable to noise). Filtering 158 removed latitudes (a) deviating by more than half the mean value (when  $\geq 3^{\circ}$  latitudes were 159 available in the bin), (b) attributable to known artifacts in the data, or (c) at magnetic latitudes greater than 21°, 22°, and 30° from the magnetic equator in the March, September and December 160 161 observations respectively. From the EIA lats satisfying these conditions, a daily mean was calculated for each hemisphere, north and south, using 5° magnetic longitude bins. The monthly 162 163 average of the crests' magnetic latitudes as a function of mlon for September is shown in Figure 164 3a. Magnetic latitudes of the northern and southern (absolute value) crests are plotted as red o's 165 and blue  $\times$ 's, respectively. The latitudes shown have been smoothed over three, 5° longitude bins. The average latitudes of the crests (north and south) are plotted as green \*'s. The error bars are 166 167 representative of the  $\pm 1\sigma$  standard deviations at selected longitude bins. However, the latitudes are 168 not completely independent, substantial correlations (>0.6 in Sept.) occur between the latitude 169 displacements from the means for widely separated longitudes. Consequently, differences 170 persisting across multiple longitudes can be meaningful even when they are smaller than the error 171 bars for individual longitude bins. Persistent deviations that are of similar magnitude to the uncertainties shown but persist for numerous longitude bins are plausibly significant. Such 172 173 behavior is also consistent with the influence of physical processes in the thermosphere and 174 ionosphere having spatial scales larger than the longitudinal grid used.



Average Magnetic Latitude versus Magnetic Longitude for Months Studied

**Figure 3**. Monthly average of the absolute value of the magnetic latitude of the EIA crests as a function of magnetic longitude (green \*) for (a) September (equinox), (b) December (solstice), and (c) March 2020 (equinox). Also shown are the latitudes of the northern crest (red  $\circ$ ), and absolute value of southern crest (blue  $\times$ ). (See text for a complete explanation of the averaging process). The yellow vertical line at 26.5° mlon indicates the longitude where the magnetic and geographic equators cross. The magnetic latitude shifts are consistent with displacing both crests in the same geographic direction. The changes in separation and location near where the equators cross indicate there are significant seasonal-longitudinal differences in the EIA latitudes.

#### 176 4 Results and Discussion

177 As shown in Figure 3a and 3c, during equinox the monthly averaged EIA lats have a minimum near 26.5° magnetic longitude (vertical, yellow line, ~50° W geographic longitude). 178 179 Each of the crests, north and south, exhibits a slightly different longitude dependence. The crests 180 in both hemispheres show similar trends during Fall and Spring equinox. Relative to the north-181 south averages, the southern crests in the west are at slightly lower latitudes ( $\sim 2^{\circ}$ ) than the northern, 182 and the displacements reverse at the most eastern longitudes. During December solstice even larger 183 southward shifts, ~4°, are seen at the eastern longitudes. These shifts are consistent with a 184 southward shift, geographically, of both crests which, at the eastern longitudes, increases the 185 magnetic latitude of the southern crest and decreases that of the northern. These relatively 186 consistent patterns in the monthly averages indicate that identifiable processes are probably 187 responsible, that the shifts are not random.

As discussed in the introduction the EIA lats and their symmetry with respect to the magnetic equator depend on several influences, including the ion production and loss rates (primarily at local times prior to sunset), neutral winds, tides, electric fields (e.g., Fang et al., 2018), and magnetic field configuration. Next, these influences on the nighttime crests' latitudes are discussed.

Neutral winds, tides and electric fields: Since the effects of neutral winds, tides and the resulting electric fields are intertwined, their effect will be discussed jointly. Each is a possible source of latitudinal shifts. Shifts due to meridional winds have been reported (e.g., Khadka et al., 2018 and references therein). Both longitudinal (e.g. England et al., 2006) and seasonal (e.g. Hedin et al., 1991 or Liu et al., 2006) differences in the winds are widely recognized. These differences may introduce seasonal (equinox versus solstice) and the longitudinal (east versus west) signatures in the EIA lats. Similarly, longitudinal differences in the winds have been associated with tides.

One of the most prominent differences in the winds due to tides is associated with the landmass distribution (e.g., England et al., 2006). Since most of the eastern longitude region that GOLD observes is over the Atlantic while the western longitudes are over South America, stronger DE3 tides and associated tidal intereactions (Oberheide et al., 2011) would be expected at the western longitudes. Consequently, similar longitudinal effects would be expected at solstice and equinox, but there are significant seasonal differences. Therefore, other effects or combination of effects appear to have a larger influence on the EIA lats.

207 Seasonal differences in the meridional winds are associated with differences in the solar 208 energy received, north versus south. Models and the available observations both indicate that the 209 winds are similar for all three seasons.

210 Winds from HWM 14 (Drob et al., 2015) near the start of the observations, approximately 211 one hour after sunset, are similar at all the longitudes and months that GOLD observed. For 212 example, in September, at both 75° W (near the maximum separation) and 45° W (near the 213 minimum) northward winds reach ~10 m/s and zonal winds ~100 m/s at latitudes near the EIA. 214 Winds could shift the crests' latitudes northward, increasing northern and decreasing southern 215 latitudes. Winds also alter the neutral composition slightly (Liu et al., 2009), which would slightly 216 alter ion production and loss (discussed later) rates, but the effects are expected to be small. At the 217 western longitudes, where the largest displacements are seen, the northward latitude shifts shown in Figure 3a for September are consistent with those from HWM 14. However, the HWM 14 winds 218 219 in December are inconsistent with the observations because the crests at the eastern longitudes are 220 shifted opposite the meridional winds. The data used in HWM 14 could be a factor. HWM 14 is 221 expected to perform best in the American sector where it was updated from a previous version to 222 include ground-based observations from the American sector and the South Pole. The model 223 incorporates fewer data from the African sector where Fisher et al., (2015) reported Fabry-Perot 224 Interferometer measurements from the Ukaimeden Observatory in Morocco (31.2° N, 7.9°W; 225 19.7° magnetic) from Nov 2013 to Dec 2014 that differ from the model.

226 However, longitudinally averaged dayside observations by the Ionospheric CONnections 227 (ICON) mission also show northward winds, consistent with HWM 14, during all three months. 228 ICON observed significantly (2-5 times) stronger winds at low latitudes during solstice (Figure 4). 229 Immel et al., (2021) found, in 23-31 March 2020 ICON observations, meridional ion velocities at 230 the equator that varied with longitude (e.g., their Figure 3) but were smaller (< 10 m/s) than the 231 mean, equatorial neutral winds in Figure 4. The ion velocities observed in March are even less 232 significant relative to the neutral winds in September and December. Without the strong northward 233 winds during solstice, even larger southward displacements would be expected.

234 Nightside winds during the GOLD observations are probably northward also. Since the 235 available observations are consistent HWM 14 on the dayside, it is plausible to infer nightside 236 winds from HWM 14. Since those are also northward, significant longitudinal differences in the 237 average neutral winds would be necessary to explain the observed latitudes. This seems unlikely 238 for all three months. During both equinoxes the EIA crests are displaced northward over South 239 America but southward over Africa. Similarly, during December only the eastern longitudes are 240 significantly displaced relative to the magnetic equator. These results are consistent with dayside 241 processes playing a key role in the post-sunset structure of the EIA crests. Furthermore, this 242 indicates that the observed displacements probably have other or additional causes than winds 243 alone.



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Figure 4. Mean meridional winds at 250 km observed by ICON as function of latitude and day of 2020. Tides have been removed.

247 Winds can also influence the vertical drifts, which have been linked to the strength of the 248 EIA, but few studies have investigated the longitude dependence. Fejer et al. (2008) have shown 249 that during equinox upward drifts increase from  $\sim 70^{\circ}$  W to  $0^{\circ}$  in the evening (see their Figure 8). 250 Therefore, their observations would not explain the longitude dependence shown in Figure 3. Kil et al. (2008) used ROCSAT-1 data to extract vertical plasma drifts near the equator and found a 251 252 wave-4 pattern that was used to reproduce a similar pattern in electron density (Fang et al., 2009). 253 Models have shown similar results (e.g., Pedatella et al., 2011). Since the wave-4 pattern is a tidal 254 effect further discussion is presented in the next section.

255 Magnetic declination, PRE and PSSR: In addition to the electric fields associated with 256 the neutral winds, there are longitudinally localized field effects in the pre-reversal enhancement 257 (PRE) and post sunset rise (PSSR), both of which have longitudinal-seasonal variations related to 258 the magnetic declination. These are a recognized cause of longitudinal differences in the EIA, and 259 the largest declinations at low-latitude are over South America. GOLD nighttime observations 260 encompass these longitudes which also have the largest offsets between Earth's magnetic and 261 geographic latitudes. At the westernmost and easternmost longitudes, the magnetic declination  $\delta$ is near zero. Near the region where magnetic and geographic equators cross,  $\delta$  is ~ -20° as shown 262 263 in Figure 5. Consequently, the interaction between constant winds and the Earth's magnetic field 264 can vary substantially with longitude, and similar effects are expected during Spring and Fall 265 equinoxes. If we denote the geographic zonal and meridional wind components as  $U_G$  and  $V_G$ , then the magnetic meridional component can be expressed as  $V_M = V_G \cos(\delta) \mp U_G \sin(\delta)$  (equation 3) 266 of Luan and Solomon, 2008). Note that, the equatorward meridional wind components are positive 267 - i.e., southward directed in the northern hemisphere and northward in the southern hemisphere -268 269 and the upper (lower) sign applies to the northern (southern) hemisphere. Consequently,  $V_M$  and 270  $V_G$  can differ significantly when  $\delta$  is nonzero and a strong zonal wind is present. Thus, longitudinal 271 asymmetries in the crests can occur, depending on the local magnetic field configuration. Eastward 272 of the equators crossing, approximately where changes in the crests' latitudes occur,  $V_M$  decreases;

therefore, one should expect the minimum displacements and symmetric locations if windmagnetic field effects dominante. Since the observed latitudes during some months are clearly inconsistent with that expectation, that effect alone is clearly insufficient.

276 The magnetic declination angle also affects the PRE which is stronger when the sunset 277 terminator and magnetic field lines are aligned (Abdu et al., 1981; Tsunoda, 1985). These effects 278 are most prominent over eastern South America, resulting in stronger postsunset vertical ion drifts 279 during December solstice, when the sunset terminator and magnetic field lines are more aligned 280 than during June solstice, as observed in AE-E and ROCSAT-1 measurements (Fejer et al., 1995, 281 2008). Fejer et al. (2008, Figure 8) found that the seasonal, longitudinal variation of PRE during 282 December solstice had a maximum at ~45° W. That does not match the observed variations the 283 EIA lats in December (Figure 3b). During eqinox, the PRE trends higher rom 75°W to 0° longitude, 284 which does not match the observed EIA latitudinal variation (Figures 3a, 3c). Therefore, 285 longitudinal differences in PRE (alone) would not explain the observed EIA latitudes.

286 Even though the PRE is weak the PSSR of the F layer, to which PRE contributes, can be 287 strong (Tsunoda et al., 2018). The similarity between the longitudinal behavior of the EIA crests 288 latitudes in September (Figure 3a) and the magnetic declination versus longitude plot, shown in 289 Figure 5, prompted us to examine the solar terminator to magnetic field alignment. Because, when 290 the terminator and magnetic field are aligned, sunset is near simultaneous at magnetically 291 conjugate points, as discussed earlier. Since this can enhance the PSSR, producing conditions 292 favorable for ionospheric depletions and scintillation, it might also influence the EIA latitudes. 293 Figure 6 shows how the average magnetic latitude changes with the magnetic declination to solar 294 terminator (D-T) angle for March (blue), September (red ), December (green), at locations west 295  $(+, \times \text{ or } \star)$  and east  $(\diamond, \star \text{ or } \circ)$  of the minimum declination. For equinox the terminator is 296 assumed to be parallel to the latitude meridians; for December solstice the average inclination to 297 the meridians  $(21.2^{\circ})$  is assumed.

298 As seen from Figure 6, there is no apparent dependence on the declination-terminator angle 299 in March and December. If the D-T alignment alone were responsible for the observed longitudinal 300 changes in magnetic latitude, a similar dependence should occur during both equinoxes. It should 301 also be expected during December solstice because the D-T angles over Africa reach similar values 302 to those during equinox over South America. In December the D-T angle varies significantly 303 across the longitudes observed, and as discussed earlier, east of  $\sim 40^{\circ}$  W longitude both crests are 304 offset south (magnetically) relative to the magnetic equator. Clearly the D-T alignment alone is 305 not the dominant factor in determining the EIA lats.

306 The large latitude shifts observed during December solstice probably eliminate the PSSR 307 electric fields from a primary role. During December previous analyses at solar minimum indicate 308 that the average PSSR would be weak over the longitudes with large shifts. Consequently, a  $\sim 4^{\circ}$ 309 shift in latitude seems unlikely during the brief ~1 hour period typical for the PSSR. Also, the 310 observed latitude of the EIA at each longitude is observed from ~1 hour after sunset to later local times and changes little during the evening hours. Similarly, comparisons with TEC data from 311 312 earlier local times appear to show a similar latitude versus longitude structure pre and post sunset 313 (e.g., Cai et al., 2020). To summarize, since latitude shifts of ~4° are observed and these shifts do 314 not show a distinct dependence on the D-T angle, additional mechanisms must be significant and 315 declination effects alone cannot explain the observed crests' latitudes.



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**Figure 5**. At equinox the angle between the solar terminator and the magnetic field is approximately the declination angle. Declination angles from https://www.ngdc.noaa.gov/geomag/calculators/magcalc.shtml.



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Figure 6. The dependence of the crests' average latitude on the magnetic declination to terminator (D-T) angle in December and March are similar. Both differ from September. This indicates that other factors must influence the latitudes. Averaged magnitudes of the EIA lats for March, September and December are shown in blue, red and green respectively. For each month points west (+,  $\times$  or \*) and east ( $\diamond$ ,  $\star$  or  $\circ$ ) of the minimum declination are shown.

325 **Ionization and radiative recombination**: Ion production at night is insignificant, 326 essentially all the nighttime ionosphere is residual ionization produced during the daytime. While

327 the daytime production provides the initial ion densities which the other processes modify, 328 resulting in the peak densities GOLD observes. Radiative recombination  $(O^+ + e^-)$  can play a more 329 important role. It influences the peak electon densites and TEC in the nighttime ionosphere. Since 330 winds preferentially move ions along the magnetic field, the peak altitude of ionosphere in the 331 equatorial region also changes, increasing or decreasing as they move along the magnetic field. As 332 the altitude increases (decreases) the volume number densities decrease (increase) which decreases 333 (increases) the recombination rates. Consequently, asymmetries in the peak densities are likely to 334 increase with time. Radiative recombination effects should not significantly affect the average 335 locations shown in Figure 3 because all longitudes are observed near sunset, at similar local times. 336 This expectation is consistent with a small set of GOLD images of the Eastern longitudes were an 337 additional images were made at later local times. Those images were approximately 1 hour later 338 than the first images near Africa. Changes in the EIA latitude were small even with approximately 339 twice the time since sunset. Given the seasonal dependence of the neutral winds, etc. radiative 340 recombination is expected to influence the observed latitudes of the EIA, but it is expected to be a 341 small effect based on the observations. That expectation is consistent with modeling (Cai et al., 342 2022) for times when distinct changes in the EIA latitude are observed on an hour time scale. 343 Furthermore, observations of such rapid changes in latitude are atypical in the GOLD data.

344 Subsolar point: There is an additional seasonal effect to consider beyond those associated 345 with the winds, tides, and magnetic field geometry discussed above: the subsolar point location. 346 During both equinoxes the subsolar point at most longitudes over South America is north of the 347 magnetic equator and south of it over the Atlantic Ocean and Africa. During December solstice 348 the subsolar point to magnetic equator separation reaches  $\sim 30^{\circ}$  at longitudes over the Atlantic 349 Ocean and Africa. Shown in Figure 7 is the average separation for each month analyzed. Average 350 values for the two equinox months, September and March, differ slightly because days during the 351 CIR storm in late September were omitted. Some dependence on the subsolar point is not 352 surprising since ion production rates on the dayside have a solar zenith angle dependence and the 353 nightside ionosphere in the EIA within a few hours of sunset (where GOLD observes) consists 354 primarily of ions produced on the dayside.

The equinox observations exhibit a dependence on the subsolar point location at longitudes where the separation between the equators is greatest within GOLD's field-of-regard (FOR). The EIA lats in March and September, Figure 3a and 3c, show northward displacements (north crests to higher and south crests to lower magnetic latitudes) of both crests at most longitudes west of  $\sim$ -40° W. Shifts in the opposite direction are seen at the most eastern longitudes but are less prominent, although the subsolar-magnetic equator separation is similar in both the east and west.

A dependence on the subsolar point location is also consistent with the December observations (Figure 3b) and the stronger, northward winds during solstice. At the eastern longitudes the crests are displaced as much as 6° southward, but in the west displacements are negligible. If the subsolar point location were the only influence, as seen from Figure 7 one would expect similar latitude changes at the western longitudes for both seasons since subsolar-equator separations are similar.

A similar dependence on subsolar point to magnetic equator separation has been reported for dayside observations. Balan et al. (2013) found that the north-south asymmetry of TEC in the EIA at equinox during magnetically quiet conditions seemed more dependent on the displacement of the equators than on the declination angle. They examined Constellation Observing System for Meteorology Ionosphere and Climate (COSMIC) data and found weaker asymmetries at smaller magnetic equator displacements  $(15^{\circ} \text{ W-75}^{\circ} \text{ W} \text{ geographic})$  than at larger  $(60^{\circ} \text{ E-120}^{\circ} \text{ E} \text{ and} 60^{\circ} \text{ W-120}^{\circ} \text{ W})$ . As shown in their Figure 1, the displacements within GOLD's FOR have similar maximum absolute values in the Eastern (~10° W geographic) and Western (~70° W) observations, but GOLD shows more prominent displacements at the Western longitudes. Although their results indicated that the declination angle is a significant factor in the TEC asymmetry, that is not apparent in GOLD's peak density observations. As in GOLD's observations, their TEC data indicated that the displacement of the equators is a significant factor.

379 Likewise, a study of daytime GNSS TEC asymmetries in the 110° E geographic longitude 380 sector during the 2000 to 2011 solar cycle by Huang et al. (2013) also found that the subsolar point 381 location helped explain changes in the EIA. Their results indicated that the magnetic declination 382 angle was also a factor but did not quantify the relative contribution. The observations from GOLD 383 indicate that the subsolar point is a significant factor at night also, that its average influence may 384 be more significant than that of neutral winds, tides or magnetic declination. Consequently, the 385 subsolar point should be considered when examining the EIA lats (e.g., Oyedokun et al., 2020) at 386 longitudes where the equators have significant separation.

387



Figure 7. Each month studied has the same separation change with longitude but the mean subsolar-magnetic equator distances differ. March and September differ because some geomagnetically disturbed days during the latter were excluded. The geographic latitudes of the magnetic equator shown were used to calculate the mean separation from the subsolar point.

388 Since ion production rates are highest at the subsolar point a depdence on it during the 389 daytime is expected. Some dependence at night, especially near sunset, should be expected since 390 changes in the distribution are a result of time dependent processes that occur after the ions are 391 produced. Depending on the rate of these processes and the time of the observations, the 392 ionospheric densities observed at night can and should have a significant dependence on the local 393 ion production rates on the dayside. Effects from the other processes discussed earlier (winds, 394 tides, PSSR, PRE, etc.) will influence the electron distrution, but all begin with the electron 395 production rates on the dayside. Electron production maximizes at the minimum solar zenith angle, 396 corresponding to the subsolar point. GOLD's observations indicate that at solar minimum the

average nightime EIA latitude for times near sunset is not completely dependent on the processes

398 occurring after the ions are produced.

# 399 Conclusions

400 Monthly averaged latitudes of the EIA crests from GOLD nightside disk (NI1) 401 observations show statistically significant differences during geomagnetically quiet conditions. 402 Observations from two equinox (March and September) and one solstice (December) months of 403 2020 were analyzed. The magnetic latitudes of the north and south crests during all three months 404 (Figure 3) have a different dependence on longitude.

405 Several possible causes for the observed changes were considered. The influence of winds, 406 tides, and magnetic fields (declination) were not consistently distinguishable over seasons and 407 longitudes. Since data from GOLD cover the same range of EIA longitudes (~80° W to 10° E 408 geographic) each day, the observed variations with longitude were examined for consistency with 409 the longitudal dependence of each possible cause of latitude changes with longitude. While winds, 410 tides and magnetic declination will contribute to the observed changes, all of them were 411 inconsistent with the observed behavior. For example in December, the crests over Africa are 412 displaced southward while those over South America were not displaced, although the neutral 413 winds should be northward and similar in both regions.

414 One influence that does appear to be consistent with the observed latitude changes in all three months is the subsolar point to magnetic equator separation. As the distance to the magnetic 415 416 equator increases at either the east or west longitudes, both the northern and southern crests exhibit 417 shifts toward the subsolar point. For example, during December solstice both crests are shifted 418 southward by as much as  $\sim 6^{\circ}$  at longitudes over Africa where the subsolar point is south of the 419 magnetic equator by as much as  $\sim 30^{\circ}$  but are near the equator at longitudes over South America 420 where the subsolar point is near the magnetic equator. In the data examined there is a 421 correspondence between the latitudinal displacement of the crests and distance between the 422 subsolar point and magnetic equator.

423 While some previous studies of TEC in the daytime EIA found both the subsolar point and 424 the magnetic declination have identifiable effects, GOLD's observations of the peak densites in 425 the nighttime EIA indicate that the subsolar point dependence is the most distinct in seasonal 426 averages. Since previous studies spanning a complete solar cycle found similar effects in daytime 427 TEC during solar minimum and maximum, the effects observed by GOLD during solar minimum 428 are also expected during solar maximum. Further studies, including simulations and modeling, that 429 consider all the influences on the location of the EIA crests will be needed to understand the effects 430 of each in GOLD's observations.

These results suggest that GOLD's observations of the nightside EIA structure could provide a daily picture of the integrated effects of ionospheric processes, with information about the longitudinal differences in the low latitude processes from western South America to Western Africa. Since LEO satellite observations at low latitudes over South America are frequently problematic due to the South Atlantic Anomaly (SAA), GEO observations of that region provide a unique opportunity to expand our understanding of the thermosphere-ionosphere system.

# 437 Acknowledgments and Data

438 This research was supported by NASA contract 80GSFC18C0061to the University of Colorado.

### 439 **Open Reserch**

440 The geomagnetic indices and solar flux data are obtained from https://omniweb.gsfc.nasa.gov/.

- 441 The GOLD data used in this study are available from the GOLD Science Data Center 442 (http://gold.cs.ucf.edu/search/) and NASA's Space Physics Data Facility
- 443 (https://spdf.gsfc.nasa.gov/pub/data/gold/).

#### 444 **References**

Aa, E., Zhang, S.-R., Wang, W., Erickson, P. J., Qian, L., Eastes, R., et al. (2022). Pronounced
suppression and X-pattern merging of equatorial ionization anomalies after the 2022 Tonga
volcano eruption. Journal of Geophysical Research: Space Physics, 127, e2022JA030527.
https://doi.org/10.1029/2022JA030527

- 449 Abdu, M. A., J. A. Bittencourt, and I. S. Batista (1981). Magnetic declination control of the
- 450 equatorial F-region dynamo electric field development and spread-F, J. Geophys. Res. Space
- 451 *Physics*, 86, 11,443–11,446, doi:10.1029/JA086iA13p11443
- Alken, P., Thébault, E., Beggan, C.D. et al. (2021). International Geomagnetic Reference Field:
  the thirteenth generation. Earth Planets Space 73, 49. doi:10.1186/s40623-020-01288-x
- Balan, N., Rajesh, P.K., Sripathi, S., Tulasiram, S., Liu, J.Y., and Bailey, G.J. (2013). Modeling
  and observations of the north–south ionospheric asymmetry at low latitudes at long deep solar
  minimum; *Advances in Space Research*, **52** 375–382, doi:10.1016/j.asr.2013.04.003.
- 457 Basu, Su., Basu, S., Huba, J., Krall, J., McDonald, S. E., Makela, J. J., Miller, E. S., Ray, S., and

458 Groves, K. (2009). Day-to-day variability of the equatorial ionization anomaly and scintillations

459 at dusk observed by GUVI and modeling by SAMI3, J. Geophys. Res. Space Physics, 114,

- 460 A04302, doi:10.1029/2008JA013899
- 461 Cai, X., Burns, A. G., Wang, W., Coster, A., Qian, L., Liu, J., ... & McClintock, W. E. (2020).
- 462 Comparison of GOLD nighttime measurements with total electron content: Preliminary results. J.
  463 *Geophys. Res. Space Physics*, 125(9), e2019JA027767. doi:10.1029/2019JA027767
- 464 Cai, X., Qian, L., Wang, W., McInerney, J. M., Liu, H.-L., & Eastes, R. W. (2022).
  465 Hemispherically asymmetric evolution of nighttime ionospheric equatorial ionization anomaly in
  466 the American longitude sector. Journal of Geophysical Research: Space Physics, 127,
  467 e2022JA030706. doi:10.1029/2022JA030706
- Drob, D. P., Emmert, J. T., Meriwether, J. W., Makela, J. J., Doornbos, E., Conde, M., et al. (2015).
  An update to the Horizontal Wind Model (HWM): The quiet time thermosphere, *Earth and Space Science*, *2*, 301–319, doi:10.1002/2014EA000089
- 471 Eastes, R. W., McClintock, W. E., Burns, A. G., Anderson, D. N., Andersson, L., Codrescu, M.,
- et al. (2017). The Global-scale Observations of the Limb and Disk (GOLD) mission. *Space Sci Rev*, 212, 383, doi:10.1007/s11214-017-0392-2

Eastes, R. W., Solomon, S. C., Daniell, R. E., Anderson, D. N., Burns, A. G., England, S. L.,
Martinis, C. R., & McClintock, W. E. (2019). Global-scale observations of the equatorial
ionization anomaly. *Geophysical Research Letters*, *46*, *9318-9326*, doi:10.1029/2019GL084199h

477 Eastes, R. W., McClintock, W. E., Burns, A. G., Anderson, D. N., Andersson, L., Aryal, S., et al.
478 (2020). Initial observations by the GOLD mission. J. Geophys. Res. Space Physics, 125,
479 e2020JA027823, doi:10.1029/2020JA027823

England, S. L., Maus, S., Immel, T. J., & Mende, S. B. (2006). Longitudinal variation of the *E*region electric fields caused by atmospheric tides. *Geophysical Research Letters*, 33, L21105,
doi:10.1029/2006GL027465

483 Fang, T.-W., Kil, H., Millward, G., Richmond, A. D., Liu, J.-Y., & Oh, S.-J. (2009). Causal link

484 of the wave-4 structures in plasma density and vertical plasma drift in the low-latitude ionosphere. 485  $L = \frac{1}{2} \frac{1$ 

485 *J. Geophys. Res. Space Physics*, *114*, A10315, doi:10.1029/2009JA014460

Fang, T.-W., Fuller-Rowell, T., Yudin, V., Matsuo, T., & Viereck, R. (2018). Quantifying the
sources of ionosphereday-to-day variability. J. Geophys. Res. Space Physics, 123,9682–9696.
doi:10.1029/2018JA025525

Fejer, B. G., E. R. de Paula, R. A. Heelis, and W. B. Hanson (1995). Global equatorial ionospheric
vertical plasma drifts measured by the AE-E satellite, *J. Geophys. Res. Space Physics*, 100(A4),
5769–5776, doi:10.1029/94JA03240

492 Fejer, B. G., J. W. Jensen, and S.-Y. Su (2008). Quiet time equatorial F region vertical plasma drift

model derived from ROCSAT-1 observations, J. Geophys. Res. Space Physics, 113, A05304,
 doi:10.1029/2007JA012801

Fisher, D. J., J. J. Makela, J. W. Meriwether, R. A. Buriti, Z. Benkhaldoun, M. Kaab, and A.
Lagheryeb (2015). Climatolo- gies of nighttime thermospheric winds and temperatures from
Fabry-Perot interferometer mea- surements: From solar minimum to solar maximum, *J. Geophys. Res. Space Physics*, *120*, 6679–6693, doi:10.1002/2015JA021170

Forbes, J. M., Maute, A., Zhang, X., & Hagan, M. E. (2018). Oscillation of the ionosphere at
planetary-wave periods. J. Geophys. Res. Space Physics, 123, 7634–7649
doi:10.1029/2018JA025720

Gan, Q., W. Wang, J. Yue, H. Liu, L. C. Chang, S. Zhang, A. Burns, and J. Du (2016). Numerical
simulation of the 6 day wave effects on the ionosphere: Dynamo modulation, *J. Geophys. Res. Space Physics*, 121, 10,103–10,116, doi:10.1002/2016JA022907

Gan, Q., Eastes, R. W., Burns, A. G., Wang, W., Qian, L., Solomon, S. C., et al. (2020). New
observations of large - scale waves coupling with the ionosphere made by the GOLD Mission:
Quasi - 16 - day wave signatures in the F - region OI 135.6 - nm nightglow during sudden
stratospheric warmings. J. Geophys. Res. Space Physics, 125, e2020JA027880,
doi:10.1029/2020JA027880

- 510 Hedin, A. E., Biondi, M. A., Burnside, R. G., Hernandez, G., Johnson, R. M., Killeen, T. L., ... &
- 511 Virdi, T. S. (1991). Revised global model of thermosphere winds using satellite and ground-based
- 512 observations. *Journal of Geophysical Research: Space Physics*, *96*(A5), 7657-7688.
- 513 Huang, L., Huang, J., Wang, J., Jiang, Y., Deng, B., Zhao, K., & Lin, G. (2013). Analysis of the
- 514 north-south asymmetry of the equatorial ionization anomaly around 110 E longitude. Journal of
- 515 Atmospheric and Solar-Terrestrial Physics, 102, 354-361. doi:10.1016/j.jastp.2013.06.010
- 516 Jin, H., Miyoshi, Y., Fujiwara, H., & Shinagawa, H. (2008). Electrodynamics of the formation of
- 517 ionospheric wave number 4 longitudinal structure. J. Geophys. Res. Space Physics, 113,
- 518 A09307.doi:10.1029/2008JA013301
- 519 Immel, T. J., Harding, B. J., Heelis, R. A., Maute, A., Forbes, J. M., England, S. L., ... & Makela,
- 520 J. J. (2021). Regulation of ionospheric plasma velocities by thermospheric winds. Nature
- 521 geoscience, 14(12). 893-898., doi:10.1038/s41561-021-00848-4
- 522 Karan, D. K., D. Pallamraju, K. A. Phadke, T. Vijayalakshmi, T. K. Pant, and S. Mukherjee (2016).
- 523 Electrodynamic influence on the diurnal behavior of neutral daytime airglow emissions. Ann.
- 524 Geophys., 34, 1019-1030, doi:10.5194/angeo-34-1019-2016
- Karan, D. K., and D. Pallamraju, (2017). Small-scale longitudinal variations in the daytime
   equatorial thermospheric wave dynamics as inferred from oxygen dayglow emissions. *J. Geophys. Res. Space Physics*, 122, 6528-6542, doi:10.1002/2017JA023891
- Karan, D. K., Daniell, R. E., England, S. L., Martinis, C. R., Eastes, R. W., Burns, A. G., &
  McClintock, W. E. (2020). First zonal drift velocity measurement of equatorial plasma bubbles
  (EPBs) from a geostationary orbit using GOLD data. J. Geophys. Res. Space Physics, 125,
  e2020JA028173, doi:10.1029/2020JA028173
- Karan, D. K., Eastes, R. W., Daniell, R. E., Martinis, C. R., Burns, A. G., & McClintock, W. E.
  (2021). Effects of 2020 September Moderate Geomagnetic Storms in the Nighttime Equatorial
  Ionization Anomaly (EIA) as Observed by GOLD Mission. AGU Fall Meeting 2021, New
  Orleans, LA, 13-17 December 2021, id. SA25D-1993, <u>2021AGUFMSA25D1993K</u>.
- Khadka, S. M., Valladares, C. E., Sheehan, R., & Gerrard, A. J. (2018). Effects of electric field
  and neutral wind on the asymmetry of equatorial ionization anomaly. *Radio Science*, 53, 683–697,
  doi:10.1029/2017RS006428
- Kil, H., DeMajistre, R., Paxton, L. J., & Zhang, Y. (2006). Nighttime F-region morphology in the
  low and middle latitudes seen from DMSP F15 and TIMED/GUVI. *Journal of atmospheric and solar-terrestrial physics*, 68(14), 1672-1681, doi:10.1016/j.jastp.2006.05.024
- 542 Kil, H., E. R. Talaat, S.-J. Oh, L. J. Paxton, S. L. England, and S.-J. Su (2008). Wave structures of
- 543 the plasma density and vertical  $E \times B$  drift in low-latitude F region, J. Geophys. Res. Space
- 544 *Physics*, 113, A09312, doi:10.1029/2008JA013106

- 545 Laundal, K. M., and Richmond, A. D. (2017). Magnetic Coordinate Systems, Space Sci Rev, 206, 546 27-59, doi:10.1007/s11214-016-0275-y
- 547 Liu, H., Lühr, H., Watanabe, S., Köhler, W., Henize, V., & Visser, P. (2006). Zonal winds in the 548 equatorial upper thermosphere: Decomposing the solar flux, geomagnetic activity, and seasonal 549 dependencies. Journal of Geophysical Research: Space Physics, 111(A7).
- 550 Liu, H., Yamamoto, M. & Lühr, H. (2009). Wave-4 pattern of the equatorial mass density anomaly:
- 551 A thermospheric signature of tropical deep convection, Geophysical Research Letters, 36, L18104,
- 552 doi:10.1029/2009GL039865
- 553 Luan, X., and S. C. Solomon (2008). Meridional winds derived from COSMIC radio occultation 554 measurements, J. Geophys. Res. Space Physics., 113, A08302, doi:10.1029/2008JA013089.
- 555 Lühr, H., M. Rother, K. Häusler, P. Alken, and S. Maus (2008). The influence of nonmigrating
- 556 tides on the longitudinal variation of the equatorial electrojet, J. Geophys. Res. Space Physics, 113,
- 557 A08313, doi:10.1029/2008JA013064
- 558 Lühr, H., & Manoj, C. (2013). The complete spectrum of the equatorial electrojet related to solar
- 559 tides: CHAMP observations. Annales Geophysicae, 31, 1315-1331, doi:10.5194/angeo-31-1315-
- 560 2013
- 561 McClintock, W. E., Eastes, R. W., Beland, S., Bryant, K. B., Burns, A. G., Correira, J., et al. 562 (2020a). Global-scale observations of the limb and disk mission implementation: 2. Observations, data pipeline, and level 1 data products. J. Geophys. Res. Space Physics, 125, e2020JA027809, 563 564 doi:10.1029/2020JA027809
- 565 McClintock, W. E., Eastes, R. W., Hoskins, A. C., Siegmund, O. H. W., McPhate, J. B., Krywonos, A., et al. (2020b). Global-scale observations of the limb and disk Mission implementation: 1. 566 567 Instrument design and early flight performance. J. Geophys. Res. Space Physics, 125, 568 e2020JA027797, doi:10.1029/2020JA027797
- 569 Mo, X. H., & Zhang, D. H. (2020). Six-day periodic variation in equatorial ionization anomaly 570 region. J. Geophys. Res. Space Physics, 125, e2020JA028225. doi:10.1029/2020JA028225
- 571 Oberheide, J., J. M. Forbes, X. Zhang, and S. L. Bruinsma (2011). Wave-driven variability in the 572 ionospherethermosphere-mesosphere system from TIMED observations: What contributes to the "wave 4"?, J. Geophys. Res., 116, A01306, doi:10.1029/2010JA015911. 573
- 574 Oyedokun, O. J., Akala, A. O., & Oyeyemi, E. O. (2020). Characterization of African equatorial
- 575 ionization anomaly during the maximum phase of Solar Cycle 24. Journal of Geophysical
- 576 Research: Space Physics, 125, e2019JA027066. doi:10.1029/2019JA027066
- 577 Pedatella, N. M., J. M. Forbes, A. Maute, A. D. Richmond, T.-W. Fang, K. M. Larson, and G.
- 578 Millward (2011), Longitudinal variations in the F region ionosphere and the topside ionosphere-

579 plasmasphere: Observations and model simulations, J. Geophys. Res. Space Physics., 116, 580 A12309, doi:10.1029/2011JA016600

Pedatella, N. M., Hagan, M. E., & Maute, A. (2012). The comparative importance of DE3, SE2,
and SPW4 on the generation of wavenumber-4 longitude structures in the low-latitude ionosphere
during September equinox. *Geophysical Research Letters*, 39, L19108,
doi:10.1029/2012GL053643

- Rodríguez-Zuluaga, J., Stolle, C., Yamazaki, Y., Xiong, C., & England, S. L. (2021). A synopticscale wavelike structure in the nighttime equatorial ionization anomaly. Earth and Space Science,
  8, e2020EA001529, doi:10.1029/2020EA001529
- Thébault, E., Finlay, C.C., Beggan, C.D., Alken, P., Aubert, J., Barroi, O. *et al.* (2015). International
  Geomagnetic Reference Field: the 12th generation, *Earth Planet Sp* 67, 79, doi:10.1186/s40623-
- 590 015-0228-9
- 591 Tsunoda, R. T. (1985). Control of the seasonal and longitudinal occurrence of equatorial 592 scintillations by the longitudinal gradient in integrated E region Pedersen conductivity. *Journal of*
- 593 *Geophysical Research: Space Physics*, 90(A1), 447-456. doi:10.1029/JA090iA01p00447
- 594 Tsunoda, R. T., Saito, S., and Nguyen, T. T. (2018). Post-sunset rise of equatorial F layer—or 595 upwelling growth? *Progress in Earth and Planetary Science*, 5(1), 1-28. doi:10.1186/s40645-
- 596 018-0179-4
- 597 Venkatesh, K., P. R. Fagundes, D. S. V. V. D. Prasad, C. M. Denardini, A. J. de Abreu, R. de Jesus,
- and M. Gende (2015). Day-to-day variability of equatorial electrojet and its role on the day-to-day

599 characteristics of the equatorial ionization anomaly over the Indian and Brazilian sectors, J.

- 600 Geophys. Res. Space Physics, 120,9117–9131, doi:10.1002/2015JA021307
- Yamazaki, Y. and Maute, A. (2016). Sq and EEJ—A Review on the Daily Variation of the
  Geomagnetic Field Caused by Ionospheric Dynamo Currents. *Space Sci Rev* 206, 299–405,
  doi:10.1007/s11214-016-0282-z
- 604 Zhou, Yun-Liang & Lühr, Hermann & Alken, Patrick & Xiong, Chao. (2016). New perspectives
- 605 on equatorial electrojet tidal characteristics derived from the Swarm constellation: Tidal features 606 on equatorial electrojet. *J. Geophys. Res. Space Physics*, 10.1002/2016JA022713