Toward Improved Physics-Based Simulations of the LEO Space Environment using GNSS-Enabled Small Satellites

Eric K. Sutton^{1,1}, Jeffrey P. Thayer^{2,2}, Marcin D. Pilinski^{3,3}, Shaylah M. Mutschler^{1,1}, Thomas E. Berger^{1,1}, Vu Nguyen^{4,5}, and Dallas Masters^{4,4}

¹University of Colorado at Boulder ²University of Colorado Boulder ³Laboratory for Atmospheric and Space Physics ⁴Spire Global, Inc ⁵Spire Global

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Abstract

Satellite-atmosphere interactions cause large uncertainties in low-Earth orbit determination and prediction. Thus, knowledge of and the ability to predict the space environment, most notably thermospheric mass density, are essential for operating satellites in this domain. Recent progress has been made toward supplanting the existing empirical, operational methods with physicsbased data-assimilative models by accounting for the complex relationship between external drivers such as solar irradiance, Joule, and particle heating, and their response in the upper atmosphere. Simultaneously, a new era of CubeSat constellations is set to provide data with which to calibrate our upper-atmosphere models at higher spatial resolution and temporal cadence. With this in mind, we provide an initial method for converting precision orbit determination (POD) solutions from global navigation satellite system (GNSS) enabled CubeSats into timeseries of thermospheric mass density. This information is then fused with a physics-based, data-assimilative technique to provide calibrated model densities.

1	Toward Accurate Physics-Based Specifications of						
2	Neutral Density using GNSS-Enabled Small Satellites						
3	Eric K. Sutton ¹ , Jeffrey P. Thayer ^{1,2} , Marcin D. Pilinski ³ , Shaylah M. Mutschler ² ,						
4	Thomas E. Berger ¹ , Vu Nguyen ⁴ , Dallas Masters ⁴						
5	¹ Space Weather Technology, Research and Education Center (SWx TREC), University of						
6	Colorado at Boulder.						
7	² Ann and H.J. Smead Aerospace Engineering Sciences, University of Colorado at Boulder.						
8	³ Laboratory for Atmospheric and Space Physics (LASP), University of Colorado at Boulder.						
9	⁴ Spire Global, Inc., Boulder, CO.						
10	Corresponding author: Eric Sutton (eric.sutton@colorado.edu)						
11	Key Points:						
12	• GNSS-enabled satellites are capable of monitoring the state of the thermosphere at much						
13	higher cadences than current operational datasets						
14	• We present an initial technique to infer neutral densities from orbit determination						
15	products of the Spire CubeSat constellation						
16	• Densities are used to drive a data-assimilative, physics-based model of the thermosphere						
17	and ionosphere during 23 Sept.–9 Dec. 2018						

18 Abstract

Satellite-atmosphere interactions cause large uncertainties in low-Earth orbit determination and 19 prediction. Thus, knowledge of and the ability to predict the space environment, most notably 20 21 thermospheric mass density, are essential for operating satellites in this domain. Recent progress has been made toward supplanting the existing empirical, operational methods with physics-22 based data-assimilative models by accounting for the complex relationship between external 23 drivers such as solar irradiance, Joule, and particle heating, and their response in the upper 24 atmosphere. Simultaneously, a new era of CubeSat constellations is set to provide data with 25 26 which to calibrate our upper-atmosphere models at higher spatial resolution and temporal cadence. With this in mind, we provide an initial method for converting precision orbit 27 determination (POD) solutions from global navigation satellite system (GNSS) enabled CubeSats 28 29 into timeseries of thermospheric mass density. This information is then fused with a physics-30 based, data-assimilative technique to provide calibrated model densities.

31 Plain Language Summary:

Satellites with heights below 1,000 kilometers (or about 600 miles) travel through the upper 32 atmosphere, which influences the path of their orbits. This influence has been monitored, in 33 some capacity, since the first man-made orbiting satellites were launched into space, but 34 predicting the effects is still quite difficult. Now commercial satellite "mega constellations" are 35 being launched into the region at a fast pace, which means that all satellite paths must be known 36 and projected into the future with great accuracy in order to avoid high-speed collisions. Using 37 Global Positioning System (GPS) signals, this work blends information from tracking the 38 39 position of the mega-constellation satellites themselves with a high-fidelity model of the upper 40 atmosphere, in an attempt to improve our knowledge of where satellites are and where they are going to be. 41

42 **1 Introduction**

Within low-Earth orbit (LEO), a region spanning roughly 100 to 1000 km in altitude for
the purposes of this paper, interactions between man-made satellites and the ambient atmosphere
cause large uncertainties in the orbit determination and prediction processes (Berger et al., 2020).
During episodic periods of moderate to severe space weather activity, such atmospheric drag

uncertainties can amplify by a factor of 2–5 in a matter of minutes to hours (Krauss et al., 2015; 47 Sutton et al., 2005). These uncertainties, when combined with the steadily growing launch rate of 48 small satellites and CubeSats and our advancing ability to track smaller and smaller objects, are 49 poised to overwhelm the U.S. Department of Defense infrastructure currently carrying out the 50 Detect-Track-Catalog mission. Products of this mission are pervasive across the Space 51 Situational Awareness (SSA) and Space Traffic Management (STM) enterprises and form a 52 critical infrastructure for nearly all space-based activities. Thus, knowledge and prediction of the 53 54 space environment, particularly the neutral mass density of the thermosphere and lower exosphere, are an essential part of satellite operations within LEO. 55

One of the major obstacles in predicting orbit trajectories hours to days in advance, and in 56 correlating consecutive or irregular object tracking data with a particular orbiting object, comes 57 from the legacy framework used to model the upper atmosphere's state and its interaction with 58 59 satellites and debris. The current model employed by the Combined Space Operations Center (CSpOC) and is the High Accuracy Satellite Drag Model (HASDM) (Storz et al., 2005), an 60 empirical model that self-calibrates by ingesting ground-based tracking data of a select set of 61 orbiting "calibration objects"-i.e., operational and defunct satellites passing through LEO with 62 reasonably stable ballistic coefficients. While this method provides an accurate global-average 63 snapshot of the upper atmosphere, its abilities to capture realistic spatial structure and forecast 64 65 into the future are limited, particularly ahead of geomagnetic storming that has the largest impact on LEO orbital tracking. Physics-based upper atmosphere simulation approaches offer a vast 66 potential improvement in this regard. Models in this category solve a set of Navier-Stokes fluid 67 equations that have been appropriately tailored for use in the upper atmosphere and are therefore 68 inherently better equipped for simulating a dynamic system response to impulsive energy input 69 from the solar wind and coronal mass ejections. For years the computational cost of these models 70 prohibited their use in an operational setting. However, present-day computing technology is 71 abundantly capable of running an ensemble of such models in near real time. Instead, the 72 primary reason that physics-based methods remain to be adopted by operational centers is the 73 lack of robust data assimilation schemes capable of self-calibrating at levels equal to or better 74 than those currently used in combination with empirical models. 75

Fortunately, significant strides have been made in recent years toward supplanting
 empirical methods with physics-based data assimilative models of the upper atmosphere. One

such advancement has been accomplished by accounting for the complex relationship between 78 external drivers—namely solar flux, Joule, and particle heating—and the response of the upper 79 atmosphere by employing a new least-squares filter called the Iterative Driver Estimation and 80 Assimilation (IDEA) technique (Sutton, 2018). The new filter operates similarly to an unscented 81 Kalman filter (UKF) with the addition of mechanisms to accommodate the lagged response of 82 the upper atmosphere to variations in the external drivers. Using this new technique, notable 83 improvements in neutral density specification—even during a geomagnetic storm—have already 84 been demonstrated (Sutton, 2018), which can help to lower the uncertainty of orbit determination 85 and prediction across the LEO catalog, thereby increasing the efficacy of STM activities, 86 including satellite conjunction assessment and collision avoidance. In addition, the emergence of 87 large constellations of commercial and academic CubeSats over the past 5 years brings with it an 88 89 excellent opportunity. Most newer SmallSats and CubeSats are equipped with Global Navigation Satellite System (GNSS) devices, making them valuable sources of Precision Orbit 90 Determination (POD) information. Many are also equipped with the ability to monitor their 91 attitude, allowing the construction of an accurate force model. This information can be combined 92 93 to initialize and constrain models of the upper atmosphere.

94 In order to track the state of the upper atmosphere with reasonable fidelity, the HASDM model ingests observations from ground-based radar tracks of known objects using a similar 95 technique to the one we present here. However, in order to make strides in specifying and 96 predicting the state of the thermosphere, new data sets with increased spatial resolution, temporal 97 cadence, and global coverage are needed (Bruinsma, Fedrizzi, et al., 2021). Satellite-based 98 GNSS observations are capable of describing the space environment at a much higher spatial 99 resolution and temporal cadence. Whereas the conventional radar-derived, satellite-drag data sets 100 operate on a multi-orbit to multi-day cadence, we will show that the GNSS-derived data sets are 101 capable of operating at a cadence of a single orbital period, i.e., on the order of hours rather than 102 days. Even higher cadences may also be possible but will require further development. The 103 remainder of the paper details our efforts to use the new set of information provided by CubeSats 104 to drive a physics-based, data-assimilative approach to simulating atmospheric densities in LEO. 105

106 2 Datasets

107 2.1 Spire CubeSats

Spire operates a constellation of over 100 CubeSats in LEO with altitudes ranging from
 400–650 km and inclinations spanning the globe, from equatorial to polar orbits. Figure 1 gives a
 snapshot of the distribution of altitude and orbit inclination of Spire CubeSats as of late January
 2021.



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Figure 1. Current coverage of altitude versus inclination for the Spire constellation of CubeSats (as of 26 January 2021). The error bars show the perigee-to-apogee range of altitudes. CubeSats are color coded by common launch dates with the total number of CubeSats in each launch group indicated in parentheses. The 21 May 2018 launch date is emphasized to indicate the launch date of the three satellites used in this study.

119 The data sets used in this study were provided by Spire Global as part of the NASA 120 Commercial SmallSat Data Pilot Program and cover the period of 23 Sept.–9 Dec. 2018. For the 121 purposes of our work, the following data products were utilized:

- POD solution ephemeris derived from GNSS tracking
- Satellite pointing in the form of attitude quaternions
- Satellite geometry model

POD solutions were typically available during the duty cycle of the GNSS/Radio

126 Occultation (RO) instrument. For the 2018 dataset, duty cycles were in the range of 30–40% of

the time, usually concentrated along 40- to 60-minute segments of an orbit (referred to hereafter as an orbit arc). This efficiency has increased with more recent CubeSat builds such that current duty cycles are beginning to approach 100%. For the current data set, ephemeris from each orbit arc were estimated using the RTOrb software (https://gps-

solutions.com/brochures/GPSS Brochure RTOrb Nov 2011.pdf). This software implements a 131 Kalman filter-based approach to estimate orbit ephemeris. As configured for the current dataset, 132 RTOrb considers Earth's gravity up to degree and order 120 from the EIGEN-2 model (Reigber 133 et al., 2003), Luni-Solar 3rd body perturbations, atmospheric drag assuming densities from the 134 Mass Spectrometer Incoherent Scatter extension (MSISe-90) model (Hedin, 1991), and solar 135 radiation pressure (SRP) with cylindrical Earth-shadowing effects. The latter two effects use a 136 cannonball approach in which coefficients of drag and reflectivity are estimated within each arc, 137 respectively, along with the orbit ephemeris. The treatment of drag and SRP in the POD process 138 is not to be confused with the force model described later in this section; instead, the parameters 139 140 estimated here have little bearing on our calculations of orbit energy.

The attitude of the Spire CubeSats is represented by a quaternion describing the 141 transformation from the body-fixed coordinate system (see Figure 4 below) to the vehicle 142 143 velocity/local horizontal (VVLH) orbit-based coordinate system at a given instance in time. These data enable the orientation of the satellite with respect to the final coordinate system 144 introduced in Section 3.2. In the initial phases of the NASA Data Pilot assessment, quaternions 145 were provided at an approximate cadence of 10 seconds during the duty cycle of the GNSS/RO 146 receiver, with nothing available outside of the duty cycle. However, it was realized early on in 147 the project that, due to frequent orientation maneuvers, the accuracy of the retrieved neutral 148 densities would be limited by any breaks in continuity of satellite attitude data (see Section 3.3 149 for further details). The attitude mode of the CubeSats frequently switched between an observing 150 mode aligning GNSS/RO antennas along track and a mode that maximizes the amount of solar 151 flux incident on the solar panels. Because these changes in orientation modify the integrated 152 effect that atmospheric drag has on the orbit parameters, the orientation must be monitored 153 154 constantly in order to convert orbital energy loss rates to an atmospheric density. Spire has since updated their processing chain for the entire fleet to ensure that a continuous stream of attitude 155

156 quaternions is available for any datasets originating after 2018. However, for the 2018 data set,

- 157 processing was limited to a small subset of three CubeSats from Spire Global's constellation for
- 158 which attitude data had been continuously downlinked and archived. These satellites, which will
- be used throughout the remainder of the paper, are referred to by Spire's internal satellite ID

numbers: 83, 84, and 85. These three CubeSats trace back to a common launch on 21 May 2018

- 161 into a 51.6° inclination orbit. During the time period of interest these satellites orbited between
- the altitudes of 467–492 km and remained within 800–2100 km (or 2–4.5 minutes) of one
- another along the orbit track. Additional properties and designations of these CubeSats can be
- 164 found in Table 1.
- Table 1. Properties of Spire CubeSats used in this study. Note: the last three columns apply to all
 three satellites

Spire ID	NORAD ID	COSPAR ID	Perigee/Apogee Altitude (km)	Inclination (degrees)	S/C Mass (g)
83	43560	2018-046G			
84	43559	2018-046F	467–492	51.6	4933 ± 4
85	43558	2018-046E			

Figure 2 shows the geometry for the three Spire CubeSats. The GNSS/POD antenna nominally points in the zenith direction while the front radio occultation (FRO) antenna generally points along the in-track or anti-in-track directions when the satellite is recording RO data. When the RO instrument is cycled off, the satellite reorients in such a way as to maximize illumination of the solar panels.



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174 2.2 Swarm Satellite Mission

As an independent data source, neutral densities from the Swarm satellite mission (Friis-175 176 Christensen et al., 2008) are used to compare with the assimilated model density output at locations that differ from the Spire dataset. Anomalies in the Swarm accelerometer data were 177 noticed early in the mission (Siemes et al., 2016), preventing their use for neutral density 178 179 determination using established methods (e.g., Bruinsma et al., 2004; Doornbos et al., 2010; Sutton et al., 2007). Instead, GNSS tracking data are used to produce POD solutions of neutral 180 density for the Swarm satellites at a temporal resolution of about 20 minutes, which is then used 181 to constrain the uncertainties in the accelerometer measuremens (van den IJssel et al., 2020). 182

The Swarm mission consists of three satellites: Swarm-A, -B, and -C. Swarm-A and -C 183 reside in essentially the same near-polar orbit, while the orbit of Swarm-B is higher in altitude 184 185 and slightly lower in inclination. Of the three satellites, accelerometer data is only currently available from Swarm-C. This data, referred to as Swarm-C ACC, spans the altitude range of 186 437–468 km during the 2018 period of interest. During this period, anomalies in the data can 187 cause the densities to attain non-physical values; these are removed from data prior to 188 189 performing any comparison. In addition, orbital averages of the Swarm-C densities are taken and compared with a corresponding orbital average of model densities in order to mitigate any 190 spurious errors in the accelerometer data. The orbital plane of Swarm-C precesses 12 hours with 191 respect to local time approximately every 133 days of the mission. Because the lower-inclination 192 193 Spire CubeSats precess much faster (i.e., 12 hours of local-time precession every 31.25 days),

Swarm-C data allows us to assess the accuracy of the assimilation model for local times and
locations far away from the ingested Spire data over the 2018 period of interest.

196 **3 Methods**

197 3.1 Orbital Energy Determination

To drive our data assimilative process, we use information from GNSS measurements 198 taken aboard CubeSats. There are several methods available to infer neutral densities from orbit 199 positioning information. For instance, this can be done by estimating a scaling correction for a 200 density model within a POD solution using two-line element (TLE, e.g., Brandt et al., 2020) sets 201 or GNSS tracking (e.g., van den IJssel & Visser, 2007). We choose instead to employ a model-202 203 agnostic energy tracking method that uses the existing POD solutions routinely obtained by Spire. The first step is to calculate the orbital energy at each available ephemeris data point and 204 track the change in this quantity between subsequent orbits. For an Earth-orbiting satellite, this 205 energy can be approximated in the following way: 206

207
$$\xi = \frac{v^2}{2} - \omega_{Earth}^2 \frac{x^2 + y^2}{2} - \frac{\mu}{r} - U_{nonSpherical}$$
(1)

where $r = \sqrt{x^2 + y^2 + z^2}$ and v are the satellite's respective position and velocity in an Earth-208 centered Earth-fixed (ECEF) coordinate frame, ω_{Earth} is the rotation rate of the Earth, μ is the 209 gravitational parameter for the Earth, and $U_{nonSpherical}$ is a potential function describing 210 deviations in Earth's gravitational field from the purely spherical (i.e., $-\mu/r$) term. $U_{nonSpherical}$ 211 is most commonly expressed as a spherical harmonic expansion of degree, n, and order, m. In 212 the absence of nonconservative forces (e.g., atmospheric drag or solar radiation pressure) or any 213 additional perturbing conservative forces not accounted for in Equation 1 (e.g., 3rd body 214 attraction, solid Earth tides, ocean tides, atmospheric tides, etc.), ξ is a conserved quantity along 215 the orbit of a satellite. 216

We have found that the choice of Earth-fixed coordinates becomes important when considering the non-spherical gravity terms in the energy equation (i.e., Equation 1), particularly any non-zonal terms (i.e., m > 0), which depend on longitude. In ECEF coordinates, $U_{nonSpherical}$ is clearly a function of position alone. The alternate formulation of the energy equation in an inertial coordinate frame, however, would require $U_{nonSpherical}$ to be a function

of both position and time, violating the assumptions underlying a potential function and its use in 222 the energy equation. As a result, the formulation of energy in an inertial coordinate frame does 223 not remain constant along an orbit when considering non-zonal terms-even in the absence of 224 nonconservative forces—and leads to twice-daily oscillations of approximately $\pm 130-140$ J/kg/s 225 or m^2/s^3 for the orbits analyzed in this paper, or equivalently, about $\pm 30-35$ m in the semi-major 226 axis. Much of this can be directly attributed to the n = m = 2 gravitational potential term, which 227 is the largest non-zonal term in $U_{nonSpherical}$. The 3rd body attraction from the sun and moon 228 depend on time in both Earth-fixed or inertial coordinates, although much less so in the latter. 229 While fairly minor, the work done by 3rd body attraction on a satellite's orbit can be taken into 230 account over time using the following equation: 231

232
$$\Delta\xi_{3B} = \int_{t_0}^{t_1} \vec{a}_{3B}(\vec{r}, t) \cdot \vec{v} \, dt \tag{2}$$

where $\Delta \xi_{3B}$ is the difference in orbital energy due to 3rd body acceleration between times t_0 and 233 t_1 ; \vec{r} and \vec{v} are the position and velocity vectors in ECEF coordinates; and \vec{a}_{3B} is the acceleration 234 vector of the satellite caused by the gravitational attraction from the sun and moon, also 235 expressed in the ECEF reference frame. In contrast to Equation 1, continuous knowledge of the 236 satellite ephemeris is required in order to carry out the integral calculation of Equation 2. While 237 this is not available directly from the GNSS measurements due to duty cycling, it can be 238 obtained at sufficient precision using Two-Line Element (TLE) sets along with the Simplified 239 General Perturbations (SGP4) satellite propagator, both available at https://space-track.org. The 240 continuous position of the sun and moon were obtained from JPL's planetary and lunar 241 ephemeris product (Park et al., 2021 and references therein). 242

If we describe the Earth's gravity field using the two-body approximation—ignoring for a 243 moment the non-spherical and 3rd body contributions—the energy dissipation due to atmospheric 244 drag remains obscured by the large variations in energy due to Earth's J₂ oblateness term (i.e., 245 n = 2, m = 0) and higher-order gravitational terms. The light blue data points in Figure 3 show 246 this simplified calculation of orbital energy for a single CubeSat from Spire Global's 247 constellation (satellite 83) during the period spanning 23 Sept.-9 Dec. 2018. However, when we 248 account for a spherical harmonic gravity field up to degree and order 36 and 3rd body effects, the 249 change in energy caused by atmospheric drag is more readily isolated from variations in the 250 251 gravity field as depicted by the dark blue curve.





Figure 3. Keplerian orbital energy (light blue curve, i.e., ignoring the $U_{nonSpherical}$ term from Equation 1) and total orbital energy (dark blue curve, i.e., including the $U_{nonSpherical}$ term from Equation 1) for Spire CubeSat 83 during the period of 23 Sept.–9 Dec. 2018.

Figure 4 depicts the orbital energy of all three CubeSats over the same time span as Figure 3 but zoomed in to reveal variations in the rate of decay. To conform with the POD solutions, we have used the non-spherical terms specified by the EIGEN-2 gravity model (Reigber et al., 2003). We found that, for our purposes, including terms of degree or order higher than 36 yielded diminishing returns.



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Figure 4. Orbital energy (i.e., including the $U_{nonSpherical}$ term from Equation 1) for Spire CubeSats 83, 84, and 85 during the period of 23 Sept.–9 Dec. 2018. The three timeseries lie approximately on top of one another, given that they reside in nearly the same orbit and therefore experience very similar accelerations.

During this period of time, the energy curves track one another quite well due, in part, to 266 the fact that all three CubeSats occupy essentially the same orbital plane, with separations along 267 the orbit track in the range of 800–2100 km (or 2–4.5 minutes). Changes in energy were on the 268 order of 5000 m²/s² over the entire period of analysis, or about 65 m²/s² per day. This is 269 equivalent to a change in the semi-major axis of 1.2 km total, or about 15 meters per day. These 270 magnitudes are specific to the size, shape and ballistic coefficients of the satellites, as well as the 271 altitude and prevailing geophysical conditions sampled during the time period of interest. After 272 applying a simple filter to reject erroneous arcs (note the obvious outliers on day 273, 320, and 273 342 in Figure 4), the noise level of these timeseries of orbital energy becomes low enough to 274 derive an effective energy dissipation rate between subsequent orbit arcs. 275

276 3.2 Satellite Force Model

To interpret the timeseries of energy from Figure 4 in terms of the behavior of the upper atmosphere, it is necessary to understand how the satellite drag interaction depends on atmospheric density. The rate at which energy is lost from a satellite's orbit to the atmosphere via the drag force, or the energy dissipation rate (EDR), can be related to atmospheric mass density through the following equation:

282
$$EDR \equiv -\frac{d\xi}{dt} = \frac{1}{2m} C_D A_{ref} \rho v^3$$
(3)

where C_D is the satellite's coefficient of drag, A_{ref} is the cross-sectional area of the satellite 283 projected in the direction of v, the velocity of the satellite in the ECEF coordinate frame, m is 284 the satellite mass, ρ is the mass density. Winds are neglected in this equation, however, the co-285 rotation of the atmosphere with the Earth is automatically considered through the use of ECEF 286 287 coordinates. The force model of Sutton (2009) is used to compute the coefficient of drag. From their Equation 7, we consider the transfer of momentum between incoming atmospheric particles 288 and the satellite surface assuming that particles are accommodated to the approximate surface 289 temperature of the satellite using an accommodation coefficient of $\alpha = 0.93$. While the 290 accommodation coefficient is kept constant, both C_D and A_{ref} can vary significantly over the 291 course of an orbit due to changes in the attitude of the satellite. 292

In order to compare two subsequent observations of orbital energy ξ_0 and ξ_1 calculated by Equations 1 and 2 at their respective epochs t_0 and t_1 , Equation 3 can be integrated to find the dependence on atmospheric density:

296
$$\xi_1 - \xi_0 = -\frac{1}{2m} \int_{t_0}^{t_1} C_D A_{ref} \rho v^3 dt = -\frac{1}{2m} \rho_{eff} \int_{t_0}^{t_1} C_D A_{ref} v^3 dt$$
(4)

297 Solving for ρ_{eff} , similar in theme to the work of Picone (2005), gives an effective mass 298 density between t_0 and t_1 along the orbit of the satellite.

299 Figure 5 shows the simulated change in orbital energy normalized by neutral density (EDR/ρ) as given by Equation 3 for one of Spire Global's CubeSats according to its orientation 300 over the course of a single day. This parameter, which we can refer to simply as the force model, 301 is the conversion factor between the observed energy dissipation rate and atmospheric density. 302 The periodic shift between pointing modes—one optimized for RO sensing and the other for 303 solar panel illumination—can be clearly seen in Figure 5. Accounting for the large variations in 304 the force model becomes crucial because a satellite can dwell in a given pointing mode for a 305 significant fraction of an orbit, and this dwell time is not necessarily consistent between orbits. If 306 neglected, these approximate factor-of-two variations in the force model have the potential of 307 causing errors of similar magnitude in the density retrievals. 308



Figure 5. Force model for Spire CubeSat 83 for a single day starting early on 7 Nov. The force model is the conversion factor between the observed energy dissipation rate and atmospheric density.

313 3.3 Data Assimilation

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After processing the GNSS measurements and applying the force model described above, 314 the final step in our process is to ingest these observations into a data assimilative framework to 315 correct the global upper atmospheric density. Here we briefly describe the Iterative Driver 316 Estimation and Assimilation (IDEA) technique, based on the method of Sutton (2018). This 317 method accounts for the complex relationship between external drivers-namely solar flux and 318 319 geomagnetic heating—and the resulting response of the upper atmosphere. In general, these drivers are poorly monitored and often rely on proxies that only very approximately represent the 320 321 physical mechanisms that heat and energize the upper atmosphere. To represent the absorption of 322 solar extreme and far ultraviolet (EUV/FUV) irradiance, the solar radio flux at 10.7 cm wavelength (F_{10.7}) is often used as a proxy. In terms of the solar wind-magnetosphere-323 ionosphere-thermosphere interaction, the geomagnetic Kp index is often used to characterize 324 heating and momentum exchange at high latitudes. Empirical formulas, such as the Heelis et al. 325 (1982) convection electric field model, are then used to help convert these proxies into 326 atmospheric heating, incurring further uncertainty into the overall modeling process. The 327 reliance on these proxies and their empirical coupling functions leads to large uncertainties when 328 driving a model of the thermosphere. 329

330 IDEA estimates corrections to the external forcing parameters and their empirical 331 coupling functions in order to bring a physics-based model into better agreement with direct 332 observations of the thermosphere. The discrepancies between model output and observations are 333 minimized by employing a least-squares filter similar in nature to an unscented Kalman filter 334 (UKF). Figure 6 compares the IDEA process (right) to that of a typical ensemble Kalman filter 335 (EnKF) configured for ionosphere/thermosphere modeling. IDEA runs several versions of the 336 thermosphere model, each experiencing slightly different external driving conditions.

In the current implementation of IDEA, the Thermosphere–Ionosphere–Electrodynamics 337 338 General Circulation Model (TIEGCM) (Qian et al., 2014; Richmond et al., 1992; Sutton et al., 2015) is used as the physics-based environment model. TIEGCM is a finite-difference solution to 339 the conservation equations of momentum, mass, and energy describing the upper atmosphere in 340 the presence of momentum and energy sources. TIEGCM accounts for the dominant features in 341 342 the upper atmosphere of molecular diffusion and circulation, solar heating in the EUV and FUV bands, and high-latitude auroral heating. TIEGCM also has the ability to simulate the ionosphere 343 344 and associated electrodynamic coupling between the neutral and plasma environment in a selfconsistent manner at middle and low latitudes. The model spans from 97 km at its lower 345 boundary to between 450 and 700 km at its upper boundary, mostly depending on the level of 346 solar flux. Migrating diurnal and semi-diurnal tides are specified at the lower boundary in a 347 348 climatological sense. Other dynamic features of the lower and middle atmosphere are ignored, which could lead to uncertainty when estimating corrections to the external forcing. 349

350 In terms of data assimilation, additional measures must be taken to deal with the lagged response of the upper atmosphere to variations in the external drivers. It is well known that the 351 response of the thermosphere can take on a large range of timescales depending on several 352 factors, height being among the largest contributors. In order for an estimated correction of the 353 external forcing parameters to have a timely effect on the model, the time-lagged response must 354 be accounted for. IDEA abandons the sequential filtering techniques typically used for 355 ionosphere/thermosphere applications (e.g., M. V. Codrescu et al., 2004, 2021; S. M. Codrescu 356 et al., 2018; Fuller-Rowell et al., 2004; Godinez et al., 2015; Matsuo et al., 2012, 2013; Minter et 357 al., 2004; Morozov et al., 2013; Murray et al., 2015). Instead, an iterative approach is adopted so 358 359 that estimated forcing parameters can be re-applied to a simulation over the course of a day so

that the model can respond to forcing (refer to the additional feedback loop on the right side of

361 Figure 6).



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Figure 6. Comparison of a typical Ensemble Kalman Filter as configured for use with a timedependent thermospheric model (left) with the IDEA technique (right; features in color differ from their counterparts in the EnKF flow chart on the left), where t_0 and t_1 are the respective start and end times of the model runs during a given data assimilation cycle (adapted from Sutton, 2018).

In Sutton (2018), satellite-borne accelerometer observations of thermosphere density were used to calibrate the external forcing parameters driving the TIEGCM. Here we use EDRs based on POD ephemeris derived from GNSS measurements from 3 satellites from Spire Global's constellation of CubeSats. A forward model, based on output from the TIEGCM, the satellite geometry model shown in Figure 2, and the force model of Sutton (2009), is used to synthesize orbital energy dissipation for each satellite according to Equation 4. Accelerometer instruments operate at high cadence (0.1–1 Hz) equating to a resolution of 7–70 km along the

satellite's orbit. The GNSS/POD data set yields a measurement of density more on the order of 375 once per orbit arc (possibly higher with additional development). This difference in information 376 content between data sets necessitates additional consideration when designing a thermospheric 377 estimation filter. In this case, we found that the observability of IDEA was limited to estimation 378 of the most recent daily F_{10.7} value and the most recent 6-hourly effective Kp value for each 379 assimilation cycle. The configuration used in this study iterates 3 times per assimilation cycle 380 and uses five 48-core nodes of a high-performance computer (HPC) to advance by 6 hours to the 381 next assimilation cycle in less than 3 minutes (i.e., >120x realtime). For comparison, Sutton 382 (2018) found it possible to estimate the most recent daily $F_{10,7}$ value and the three most recent 3-383 hourly Kp values for each assimilation cycle when using the high-resolution accelerometer-384 derived density data set. However, it is expected that improvements in observability will be 385 enabled through the use of more CubeSats in the estimation process. And considering the greater 386 coverage of CubeSats in altitude and local time, accuracy could very well exceed accelerometer-387 based density model corrections. 388

389 4 Results and Discussion

The period spanning 23 Sept.–9 Dec. 2018 (days 266–343) of our study was marked with 390 very low activity in terms of the magnitude and variation of solar EUV and FUV, as 391 approximated by measurements of the 10.7 cm solar radio flux ($F_{10.7}$; top panel of Figure 7). 392 Note that F_{10.7} has an approximate lower bound of 66 solar flux units (sfu) at solar minimum and 393 attains values above 200 during solar maximum. During the latter, 27-day solar rotational 394 modulation can also produce large swings in F_{10.7}, causing large signals in the thermospheric 395 density. Because the available data for this studyfalls firmly within solar minimum, the 396 variations seen here are quite small. In terms of geomagnetic activity, however, there were two 397 minor-to-moderate disturbances on 7 Oct. (day 280) and 4 Nov. (day 308) as shown by the 3-398 hourly Kp geomagnetic index (lower panel of Figure 7). 399



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Figure 7. Top: observed solar $F_{10.7}$ radio flux. The grey curve is the daily measured value from the Ottawa observatory normalized to 1 AU sun-earth distance; the black curve is an 81-day (~3 solar rotation) centered average. Bottom: the 3-hourly planetary magnetic index Kp. Both panels span the period of interest 23 Sept.–9 Dec. 2018.

Given observations of orbital variations and an appropriate force model as discussed in 406 the previous section, an effective atmospheric mass density can be inferred between consecutive 407 orbit arcs. Figure 8 shows such neutral mass densities derived from the three CubeSats (blue, red 408 and yellow curves) of Spire's constellation. The cadence of these densities is approximately one 409 measurement for each consecutive set of orbit arcs. For the time period studied, this equates to a 410 cadence of about 2–2.5 hours on average. This cadence depends on the instrument duty cycle, 411 which has steadily improved since 2018. A higher cadence may be possible in the future as duty 412 cycle improves, however, the exact allowable cadence will also depend on the altitude of the 413 satellite and the noise errors of the GNSS measurements. HASDM output is also shown with the 414 black curve for reference. This empirical model is calibrated by ground-based radar tracking 415 observations of approximately 70–90 orbiting objects. Because the individual tracking 416 417 observations are sparse-relative to those available from GNSS-densities derived from this technique have an effective cadence of several hours to several days (Storz et al., 2005). 418



419

Figure 8. Neutral mass densities derived from Spire CubeSats 83–85. Also shown is output from
HASDM as sampled on the orbit of satellite 84. The values plotted are the effective densities
(see the right-hand side of Equation 4) between subsequent orbit arcs.

The CubeSat-derived densities maintain good agreement with one another, given the 423 close proximity of all three spire satellites within 800–2100 km (or 2–4.5 minutes) along the 424 orbit track. Agreement with HASDM is also reasonable during this time period. As Figure 7 425 shows, there are several minor to moderate variations in Kp over the time interval. The 426 signatures of these disturbances are also seen in the neutral densities of Figure 8. There are 427 several deviations between data and model though, most notably around days 270, 290, and 300, 428 where CubeSat-derived densities are significantly lower than HASDM. We have not yet 429 concluded whether model or data are in error during these intervals, since very little ground-truth 430 data exists during this period for validation. Another period of discrepancy exists around the 431 geomagnetic disturbance on day 280, where CubeSat-derived densities experience a much larger 432 storm-time increase. While it is possible that the higher cadence GNSS densities are capturing 433 actual storm dynamics better than HASDM, we note that POD data were less frequent during this 434 particular event than during other times. Additionally, attitude data was unavailable for satellite 435 436 83 over much of the disturbance, particularly days 282–285. The discrepancy in amplitude during this event could also be a function of the higher cadence of the CubeSat POD data fit 437 spans (5–6 hours during this event) relative to that of the HASDM data fit spans (~1 day or 438 more), in which case, the CubeSat-derived densities would be expected to more accurately 439 440 resolve the storm-time disturbance.

In general, some error in the observations and modeled output is expected, of 441 instrumental, data sampling, and geophysical origins. Part of this error is caused by variations in 442 sampling location for a given data point. In other words, the data points presented in Figure 8 do 443 not represent the density averaged over a complete orbit; instead, each data point can be sampled 444 over a very different part of the globe than the previous. The resulting error can be seen in the 445 HASDM model, which if plotted as an average over full orbits, would appear much more 446 smooth. Another important source of error in the density timeseries comes directly from 447 uncertainties in the POD solutions themselves. Because the POD solutions were not designed 448 with a thermospheric application in mind, we expect that some of the estimation parameters may 449 have been overfit. And finally, there is certainly an amount of geophysical variability seen in the 450 observed density timeseries that is not captured by the HASDM model. While an in-depth error 451 452 analysis is beyond the scope of the present work, we will continue to investigate techniques to minimize these sources of error, including improving the underlying POD solutions and 453 454 combining timeseries from additional satellites.

A central goal of this work is to ingest multiple data sources into a physics-based, assimilative thermosphere model to combine information and mitigate the uncertainty of any one dataset. Figure 9 shows the baseline TIEGCM simulation without any assimilation (grey curve) driven externally by the observed geophysical indices (GPI) of Kp and $F_{10.7}$; the POD-based densities derived using the techniques described in the previous Section (blue, red, and yellow curves); and the IDEA output over the interval spanning 23 Sept.–9 Dec. 2018 (solid black curves).

The baseline TIEGCM-GPI simulation shows muted response to the Kp and $F_{10,7}$ inputs 462 during this solar-minimum interval, when compared with the IDEA output (or with the HASDM 463 output in Figure 8). CubeSat densities and IDEA output agree very well over the interval. There 464 are, however, several short periods when POD data from a single satellite becomes sparse, such 465 as the period around day 304–306 for satellite 85 (yellow curve), or when attitude data becomes 466 unavailable, such as the period around day 282-285 for satellite 83 (blue curve). There are also 467 several periods during which data from a single satellite becomes spurious, not agreeing with the 468 data from the other two satellites, such as the period around 335-340 for satellite 85 (yellow 469 470 curve). In these cases, the other two data sets tend to compensate for missing or spurious data 471 from the third satellite. This leads us to believe that adding data from additional satellites and



473 assimilation process.



474

Figure 9. Comparison of observations with model output. CubeSat-derived densities are given
by the colored curves for satellites 83 (top), 84 (middle), and 85 (bottom). Also shown is the
output from the baseline thermosphere model driven by measured geophysical indices
(TIEGCM-GPI, grey curves) F_{10.7} and Kp. The data assimilation IDEA output is given along
each of the CubeSat orbits by the black curves.

The performance of these models with respect to the CubeSat-derived densities are assessed using the metrics of Sutton (2018) and Bruinsma, Boniface, et al. (2021). These consist of the mean (μ), standard deviation (σ), and root mean square error (*RMSe*) of the ratio of model density to observed density, all computed in logarithmic space:

484
$$\mu(m/o) = \exp\left(\frac{1}{N}\sum_{i=1}^{N}\ln\frac{\rho_{m,i}}{\rho_{o,i}}\right)$$
(5)

485
$$\sigma(m/o) = \sqrt{\frac{1}{N} \sum_{i=1}^{N} \left(\ln \frac{\rho_{m,i}}{\rho_{o,i}} - \ln \mu(m/o) \right)^2}$$
(6)

486
$$RMSe(m/o) = \sqrt{\frac{1}{N} \sum_{i=1}^{N} \left(\ln \frac{\rho_{m,i}}{\rho_{o,i}} \right)^2}$$
(7)

As mentioned in Sutton (2018), these metrics have several properties that are desirable 487 when working with the ratio of two quantities that vary exponentially, such as neutral densities. 488 489 The RMSe(m/o) and $\sigma(m/o)$ quantities are best interpreted as a percentage in the following way: $\% = 100 \times (\exp \sigma(m/o) - 1)$. The RMSe(m/o) is a combination of $\mu(m/o)$ and 490 $\sigma(m/o)$, as can be see through the following relation: $RMSe(m/o)^2 = (\ln \mu(m/o))^2 +$ 491 $\sigma(m/o)^2$. The RMSe(m/o) is therefore a good indicator of total model errors. However, if the 492 intent is to drive a POD process using the density model, it may be more informative to use the 493 $\sigma(m/o)$ metric, since a ballistic coefficient is typically estimated per satellite. In practice, this 494 estimated ballistic coefficient will soak up errors not only in the assumed coefficient of drag, but 495 496 also in the mean bias of the density model. Table 2 shows the overall performance of the three models, TIEGCM-GPI, HASDM, and IDEA at re-creating the energy dissipation rates observed 497 by the Spire CubeSats' GNSS data. Table 2 also shows the performance of the three models in 498 synthesizing the independent data set of orbit-averaged neutral densities from the Swarm-C 499 satellite. 500

501 During the period of interest, the IDEA method outperforms HASDM in all three metrics with respect to the assimilated Spire data. This is true of both the prior and posterior IDEA 502 estimates of density. The 'posterior' IDEA estimate is the fully assimilated nowcast solution, 503 whereas the 'prior' IDEA estimate is a 6-hour forward simulation (i.e., forecast) by the TIEGCM 504 based on the initial conditions and estimated drivers from the previous posterior assimilation 505 506 cycle. The prior mean is expected to remain close to the posterior mean if no major changes in the actual geophysical conditions occur during this 6-hour span. Minor fluctuations in the actual 507 508 conditions could cause upward or downward trends over a given 6-hour span, resulting in density variations that tend to average out of $\mu(m/o)$ while slightly increasing the $\sigma(m/o)$ metric for 509 510 the prior IDEA estimates relative to the posterior. It should be noted, however, that IDEA has a clear advantage in this comparison over the other two models, since IDEA assimilates the very 511

data that it is now being compared against. This comparison confirms that the IDEA technique,
as an estimation filter, has the requisite control authority to sufficiently adjust the model to the
assimilated data set.

To go a step further, the Swarm-C ACC data is used as an independent data source to 515 assess the performance of IDEA in locations outside the vicinity of assimilated data. Due to the 516 517 differences in precession rates between the Spire CubeSats and the Swarm-C satellite, the local times of their orbital planes align approximately once every 41 days. This alignment occurs 518 519 twice during the 23 Sept.-9 Dec. 2018 time period, on 10 Oct. (day 283) and 20 Nov. (day 324). 520 Aside from these brief alignment periods lasting only a few days, Spire and Swarm data sets are sampling vastly different sectors of the globe. Table 2 shows that IDEA succeeds in reducing the 521 522 $\mu(m/o)$ and overall *RMSe*(m/o) with respect to the free-running TIEGCM-GPI simulation but at the expense of an increased $\sigma(m/o)$. In the comparison with Swarm-C data, HASDM 523 performs the best in all three metrics. However, it should be noted that HASDM has a clear 524 advantage in this comparison over the other models because HASDM assimilates data from ~75 525 satellites from across the globe while IDEA only assimilates within a single orbit. 526

Sutton (2018) used accelerometer-derived neutral densities from a single satellite to drive 527 the IDEA technique. In the previous study, the technique showed high proficiency for estimating 528 neutral densities in local times away from where the assimilated data resided. In the current 529 530 work, using effective neutral densities from a single orbit plane at an approximate 2–2.5 hour cadence per satellite, the comparison with data from other local times deteriorates. While this is 531 532 not all that surprising, it does provide some insight into the specificity required to apply adequate 533 corrections to the external drivers. The observability of these corrections depends on features of 534 the observations, including the global coverage, spatial resolution, temporal cadence, measurement error, and measurement type (e.g., mass density from accelerometers vs. chemical 535 536 composition from a mass spectrometer). In essence, the impact that each external driver has on the observation must be distinguishable from the impact caused by other drivers. With 537 accelerometer data, this is satisfied to some extent because measurements are of high cadence 538 and sample two distinct local time and all latitudes over the course of an orbit. Any change in 539 geomagnetic activity will first impact high-latitude thermosphere before influencing lower 540 latitudes, while changes in solar flux affect the thermosphere in a much less localized manner. 541 Likewise, the model's day-to-night ratio of density will decrease as geomagnetic activity 542

543 increases yet is only slightly affected by variations in solar flux (Waldron, 2020). Both of these

signals can be discerned with accelerometer data but not with orbit-averaged data. Several

545 questions remain: Can this issue of observability with POD-inferred densities be overcome by

546 including data from multiple local-time orbital planes and/or with reduced averaging? And in

547 terms of assimilation, which characteristic of a density data set is more valuable, spatial

resolution in the latitudinal direction or sampling of multiple local time planes?

549 **Table 2.** Performance metrics of each model with respect to the assimilated Spire Global

- 550 CubeSat data and independent Swarm-C ACC data, calculated over the entire interval spanning
- days 266–343, 2018. Orbital averages of the Swarm-C data (as well as the corresponding model
- output) have been taken to minimize the effect of spurious errors in the accelerometer data.

	TIEGCM-GPI	HASDM	IDEA					
			Prior	Posterior				
Assimilated Spire CubeSat POD Data								
$\mu(m/o)$	1.33	1.10	1.04	1.04				
$\sigma(m/o)$	48.7%	40.6%	34.6%	30.7%				
RMSe(m/o)	62.7%	42.6%	34.9%	31.1%				
Independent Swarm-C ACC Orbit-Averaged Data								
$\mu(m/o)$	1.40	1.12	1.20	1.19				
$\sigma(m/o)$	24.9%	8.2%	37.2%	32.4%				
RMSe(m/o)	49.6%	15.2%	43.8%	39.4%				

553 **5 Summary and Conclusions**

The increases in satellite and debris populations in LEO necessitates improvements in 554 how we detect, track, and catalog orbiting objects. Additionally, if we are to avoid catastrophic 555 collisions in LEO, we must also be able to reliably predict the trajectories of satellites multiple 556 days in advance, giving satellite operators sufficient lead time to plan safe and effective 557 maneuvers. With the variability of the space environment, particularly thermospheric mass 558 density, being the largest uncertainty in the orbit prediction chain, this study investigates new 559 ways to monitor the upper atmosphere. In this notoriously data-starved region, the 560 instrumentation commonly carried on recently launched LEO SmallSats and CubeSats, 561

particularly GNSS receivers, can be used to improve the accuracy of physics-based neutral
density specifications. Notably, the amount of data available from this new category of
observation should continue to scale with the crowdedness of LEO, whereas the current groundbased tracking database remains limited in quantity and resolution.

In the current work, we have applied a post-processing method to the timeseries of POD 566 ephemeris from three CubeSats in Spire's constellation. This has allowed us to track the time 567 evolution of orbital energy of each CubeSat over an orbit arc. Further application of a satellite-568 surface force model converts this information into a timeseries of in situ atmospheric mass 569 570 density. By analyzing 78 days' worth of data from late 2018, we were able to observe the impact of minor and moderate fluctuations in geomagnetic activity during the prevailing solar minimum 571 conditions. We also found good agreement with HASDM, one of the only sources of 572 thermospheric data currently available for comparison. While the resulting timeseries from a 573 574 single satellite may be prone to errors, identified here simply as a discrepancy between density timeseries derived from co-orbiting CubeSats, this can be mitigated by assimilating timeseries 575 576 from multiple data sets into a physics-based model of the thermosphere.

Additionally, with more advanced processing methods, it may be possible to lower the 577 error for timeseries of individual CubeSats. The POD solutions used here were not specifically 578 tailored to the application of measuring density. One potential complication is that overfitting of 579 parameters or insufficient arc size may have led to significant noise in the inferred densities. 580 Future work will focus on improving the POD solutions to reduce error and finding the optimal 581 582 size of the POD fitting window as a function of altitude, phase of the solar cycle, satellite geometry characteristics, and GNSS instrument precision and errors. With these improvements 583 in place, it may even be possible to attain higher cadences than a single data point per orbit. This 584 has been demonstrated when using a state-of-the-art geodetic GNSS receiver (van den IJssel & 585 Visser, 2007), but extending this technique to GNSS-equipped SmallSat constellations would 586 provide much needed global coverage of thermospheric observations. When paired with a 587 suitable assimilative, physics-based models of the thermosphere, there is great potential to lower 588 the uncertainty of orbit predictions across the LEO catalog, improve the accuracy of conjunction 589 assessments, and increase the efficacy of STM activities. 590

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