# Dynamic Bayesian networks for evaluation of Granger causal relationships in climate reanalyses

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## Abstract

We apply a Bayesian structure learning approach to study interactions between global teleconnection modes, illustrating its use as a framework for developing process-based diagnostics with which to evaluate climate models. Homogeneous dynamic Bayesian network models are constructed for time series of empirical indices diagnosing the activity of major tropical, Northern and Southern Hemisphere modes in the NCEP/NCAR and JRA-55 reanalyses. The resulting probabilistic graphical models are comparable to Granger causal analyses that have recently been advocated. Reversible jump Markov Chain Monte Carlo is employed to provide a quantification of the uncertainty associated with the selection of a single network structure. In general, the models fitted from the NCEP/NCAR reanalysis and the JRA-55 reanalysis are found to exhibit broad agreement in terms of associations for which there is high posterior confidence. Differences between the two reanalyses are found that involve modes for which known biases are present or that may be attributed to seasonal effects, as well as for features that, while present in point estimates, have low overall posterior mass. We argue that the ability to incorporate such measures of confidence in structural features is a significant advantage provided by the Bayesian approach, as point estimates alone may understate the relevant uncertainties and yield less informative measures of differences between products when network-based approaches are used for model evaluation.

# Dynamic Bayesian networks for evaluation of Granger causal relationships in climate reanalyses

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# Key Points:

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6	•	Bayesian structure learning provides a principled approach to quantifying uncer-
7		tainty in estimated network structures for relationships between teleconnections
8	•	Dynamic Bayesian networks estimated from NCEP/NCAR and JRA-55 reanal-
9		ysis data show broad overall consistency
10	•	Structural differences in high posterior credibility associations may be indicative

11 of biases relevant for subsequent model evaluation

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#### 12 Abstract

We apply a Bayesian structure learning approach to study interactions between global 13 teleconnection modes, illustrating its use as a framework for developing process-based 14 diagnostics with which to evaluate climate models. Homogeneous dynamic Bayesian net-15 work models are constructed for time series of empirical indices diagnosing the activity 16 of major tropical, Northern and Southern Hemisphere modes in the NCEP/NCAR and 17 JRA-55 reanalyses. The resulting probabilistic graphical models are comparable to Granger 18 causal analyses that have recently been advocated. Reversible jump Markov Chain Monte 19 Carlo is employed to provide a quantification of the uncertainty associated with the se-20 lection of a single network structure. In general, the models fitted from the NCEP/N-21 CAR reanalysis and the JRA-55 reanalysis are found to exhibit broad agreement in terms 22 of associations for which there is high posterior confidence. Differences between the two 23 reanalyses are found that involve modes for which known biases are present or that may 24 be attributed to seasonal effects, as well as for features that, while present in point es-25 timates, have low overall posterior mass. We argue that the ability to incorporate such 26 measures of confidence in structural features is a significant advantage provided by the 27 Bayesian approach, as point estimates alone may understate the relevant uncertainties 28 and yield less informative measures of differences between products when network-based 29 approaches are used for model evaluation. 30

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# Plain Language Summary

To produce reliable forecasts and projections, climate models should accurately re-32 produce the observed behavior of different processes that play a role in Earth's climate, 33 including the relationships between them. Statistical methods can be used to describe 34 these interactions in models and in observations, which can then be compared to eval-35 uate how well a given model captures the observed relationships. However, networks ob-36 tained from estimates of the true historical state of the climate, known as reanalyses, will 37 also be affected by the properties of the systems used to create these estimates, as well 38 as random variability, and hence may have significant uncertainties. Using what are known 39 as Bayesian statistical methods, we estimate the uncertainties associated with particu-40 lar interactions in two widely used reanalyses. Interactions that are found to be very likely 41 to be present in one reanalysis but not the other are suggested to be due to systematic 42 differences in the two reanalysis systems and need to be kept in mind when these state 43

estimates are used to evaluate climate models. Therefore, it is important to account for
the uncertainty associated with each relationship when analyzing state estimates and further employing them to evaluate climate models.

#### 47 **1** Introduction

The behavior of the Earth's climate system, from day-to-day changes in weather 48 to longer-term variations in climate, arises as a result of the interactions of diverse pro-49 cesses within the coupled land-ocean-atmosphere-cryosphere system over a vast range 50 of spatiotemporal scales (see, e.g., Ghil and Lucarini (2020) for a recent review). Devel-51 oping an accurate understanding of the underlying processes, including their response 52 to external forcing and the interactions between different components, is an important 53 first step in the development of realistic numerical forecasting models. Inevitably, even 54 with a good understanding of the main processes being simulated, any given model will 55 still be limited in its ability to represent the climate system, e.g., due to deficiencies in 56 the parameterization of unresolved processes. These limitations manifest as systematic 57 biases in the output of the model compared to observations. By comparing the repre-58 sentation of a particular component in the model with its observed counterpart, short-59 comings in the model implementation can be identified for improvement, at least sub-60 ject to the severe limitation that the globally contiguous observational record of the cli-61 mate system, and in particular the subsurface ocean, extends only over a few decades. 62

This process-oriented approach, in which attention is focused on a comparatively 63 small set of physical processes, has been widely applied in the climate science commu-64 nity. On the one hand, the complexity of the full coupled climate system means that it 65 is typically only feasible to focus on specific subsystems in any given analysis, or oth-66 erwise necessitates some form of dimension reduction. For instance, specific targeted process-67 oriented diagnostics (Maloney et al., 2019) permit the representation of these processes 68 in models to be evaluated against observations in order to drive model development (Eyring 69 et al., 2019). At the same time, the existence of distinct teleconnections, i.e., recurrent 70 large-scale modes of variability, has long been recognized (Ångström, 1935), motivating 71 simplified models of the climate in terms of a small number of interacting modes. For 72 example, atmospheric teleconnection patterns such as the Arctic Oscillation (AO) (Thompson 73 & Wallace, 1998), the North Atlantic Oscillation (NAO) (Walker, 1923; van Loon & Rogers, 74 1978), the Pacific North American (PNA) (Wallace & Gutzler, 1981; Horel & Wallace, 75

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1981; Barnston & Livezev, 1987) and Pacific South American patterns (PSA) (Mo & Ghil, 76 1987; Lau et al., 1994; O'Kane et al., 2017), and the Southern Annular Mode (SAM) (Rogers 77 & van Loon, 1982; Thompson & Wallace, 2000) constitute important sources of low-frequency 78 variability (Hannachi et al., 2017) and are associated with wide-ranging impacts (see, 79 e.g., Leathers et al., 1991; Thompson & Wallace, 1998; Mo & Paegle, 2001; Hurrell et 80 al., 2003; Gillett et al., 2006). On interannual time-scales, modes that express in the trop-81 ical oceans such as the El Niño Southern Oscillation (ENSO) (Walker, 1924; Bjerknes, 82 1969) and the Indian Ocean Dipole (IOD) (Saji et al., 1999) emerge as dominant sources 83 of variability, with important implications for global weather and climate (see, e.g., Schott 84 et al., 2009; McPhaden et al., 2021). As these large-scale modes tend to be relatively per-85 sistent, understanding their evolution and dynamics may enable more skillful forecast-86 ing over longer time-scales (Goddard et al., 2001; A. G. Marshall et al., 2014; Hannachi 87 et al., 2017). Thus, beyond simply being convenient for reducing the dimension of the 88 problem in model evaluation studies, simplified models based on these physically observed 89 modes provide an approach for better understanding the properties and interactions of 90 key sources of climate variability. 91

Even when focusing on only a few individual processes, it is generally difficult to 92 directly attribute model biases to problems in the representation of a single process (Eyring 93 et al., 2019), or to understand the behavior of an individual teleconnection mode in iso-94 lation from other processes. To capture the intrinsically coupled nature of the system 95 in simplified models, network based approaches have become increasingly popular (Tsonis 96 & Roebber, 2004; Tsonis et al., 2006; Donges et al., 2009b; Steinhaeuser et al., 2011). 97 In this framework, the climate system is represented in terms of a set of nodes, corre-98 sponding to appropriately defined subsystems or processes of interest, and edges describ-99 ing the interactions between these nodes. The subsystems, for example, may be iden-100 tified with one or more spatial gridpoints (Bello et al., 2015; Fountalis et al., 2018), or 101 pre-defined modes characterized by empirical indices (Tsonis et al., 2007). In either case, 102 the full climate system is then modeled in terms of a collection of individual, non-linear 103 dynamical systems interacting with their neighbors in the constructed network. Such net-104 works have variously been applied to study synchronization and climate shifts (Tsonis 105 et al., 2007; G. Wang et al., 2009), to investigating the collective spatial structure of the 106 statistical relationships between fields and changes over time (Tsonis & Swanson, 2008; 107 Tsonis et al., 2008; Gozolchiani et al., 2008; Yamasaki et al., 2008; Donges et al., 2011; 108

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Gozolchiani et al., 2011; Berezin et al., 2012; Guez et al., 2012; Steinhaeuser et al., 2012; Radebach et al., 2013; Y. Wang et al., 2013) and as tools for automatic mode identification (Bello et al., 2015) or dimension reduction (Fountalis et al., 2018; Falasca et al., 2019) via community detection methods.

When attempting to build a network representation of the climate, the structure 113 of the network is generally not known beforehand and must be inferred by some means. 114 A typical approach is to add or remove edges from the network on the basis of the level 115 of statistical interdependence of pairs of nodes, quantified by, e.g., the correlation (Tsonis 116 & Roebber, 2004) or mutual information (Donges et al., 2009a) between time series as-117 sociated with each node. As these measures of association will almost always be non-118 zero in finite samples, some level of thresholding or pruning must also be applied in or-119 der to exclude edges corresponding to weak or spurious associations. In the simplest case, 120 the result is an undirected network; that is, the presence of an edge between two nodes 121 indicates some level of mutual association, but does not provide information on any pos-122 sible directionality in the relationship. 123

Networks constructed in this way have proven to be very useful but do have some 124 important limitations. In particular, while correlation graphs allow for comparisons be-125 tween modeled and observed associations (Falasca et al., 2019), often it is of interest whether 126 a particular set of variables is causally related to another set. If there is a causal mech-127 anism, we may also wish to quantify the magnitude of the effects of those causal factors; 128 commonly, climate networks based on the above approaches do not provide direct ac-129 cess to measures of effect. In practice, to answer these sorts of questions it is necessary 130 to imbue the networks with additional structure. This can be naturally achieved by iden-131 tifying the original graph with an underlying statistical model, that is, by working in the 132 context of a (probabilistic) graphical model (Koller & Friedman, 2009). In a graphical 133 model, the graph encodes, in a well-defined way, the set of qualitative independence re-134 lationships between the random variables, corresponding to nodes, in the model (Jordan, 135 2004). Given a particular functional form for the interactions between variables in terms 136 of the joint probability density function (PDF), quantitative questions can also be for-137 mulated and addressed using standard algorithms (Pearl, 1982, 1988; Dechter, 1999; Koller 138 & Friedman, 2009). 139

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Graphical models may utilize undirected or directed edges, or even a mixture of 140 both, corresponding to different ways of constructing the joint PDF of the model. Mod-141 els based on directed acyclic graphs (DAGs), in which all edges also have an associated 142 direction and the graph does not contain any directed closed loops, provide an intuitive 143 representation of complex systems in terms of conditional independence relationships be-144 tween quantities (Pearl, 1988; Spiegelhalter et al., 1993; Jordan, 2004). Compared to undi-145 rected networks, directed graphs have the advantage that edges in the graph can often 146 be interpreted causally. As a result, they provide a powerful tool for studying causal re-147 lationships (Pearl, 1995), and form a useful basis for other forms of causal inference (Greenland 148 & Brumback, 2002). Bayesian, or belief, networks (BNs), as such models are usually called, 149 have therefore received increasing attention from the climate science community (Ebert-150 Uphoff & Deng, 2012a), having variously been used for forecasting and risk assessment 151 based on expert systems (e.g., Abramson et al., 1996; Catenacci & Giupponi, 2009; Pe-152 ter et al., 2009; Catenacci & Giupponi, 2013; Leonard et al., 2014; Boneh et al., 2015), 153 for learning independence relationships and possible causal interactions in observations 154 (Ebert-Uphoff & Deng, 2012b; Runge et al., 2014; Runge, 2015; Runge et al., 2015; Kretschmer 155 et al., 2016, 2017; Horenko et al., 2017; Li et al., 2018; Runge, 2018a, 2018b; Runge, Nowack, 156 et al., 2019; Runge, Bathiany, et al., 2019; Samarasinghe et al., 2019, 2020; Saggioro et 157 al., 2020; Pfleiderer et al., 2020; Di Capua et al., 2020) and models (Deng & Ebert-Uphoff, 158 2014; Ebert-Uphoff & Deng, 2017), and, most recently, for model evaluation (Vázquez-159 Patiño et al., 2020; Nowack et al., 2020). 160

One approach for using BNs as tools for model evaluation is to learn the network 161 structure in a model and in observations and assess the agreement between the two. Learn-162 ing the structure of climate networks, in the absence of expert knowledge, has primar-163 ily been achieved by utilizing constraint-based algorithms (Spirtes & Glymour, 1991; Colombo 164 & Maathuis, 2014) in which the set of edges is determined starting from a series of con-165 ditional independence tests (Ebert-Uphoff & Deng, 2012a; Runge, Bathiany, et al., 2019). 166 Constraint-based methods can flexibly incorporate linear or non-linear conditional in-167 dependence tests (Hlinka et al., 2013) together with predefined constraints, allowing for 168 non-linear dependence structures to be estimated from data. However, the inclusion of 169 edges on the basis of an (initially arbitrary) significance level together with multiple test-170 ing adjustments makes assessing the level of confidence in the inferred networks difficult. 171

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Usually, sensitivity analyses are necessary to get some handle on the robustness of theresulting networks.

Ideally though, to apply BNs as tools for model evaluation we would not only iden-174 tify differences between modeled and observed networks, but also assess whether any mis-175 match is likely due to a model bias or simply sampling variability. In other words, an 176 additional level of uncertainty quantification is required. While a great deal of insight 177 can be derived from the structure of the estimated networks, quantifying both the sign 178 and magnitude of the interaction between nodes has generally been achieved by perform-179 ing a second stage of fitting a parametric model, conditional on the inferred structure. 180 This, again, leads to complications in determining the uncertainties in the estimated in-181 teraction strengths, which (depending on the approach used) may be insufficiently con-182 servative (Draper, 1995). For the purposes of model evaluation, it is usually of interest 183 whether a model captures both the existence of a link and with the correct strength. Un-184 reliable estimates for the uncertainty in fitted interaction strengths may result in an in-185 ability to determine if an interaction is present but differs significantly in the model com-186 pared to observations. On top of this, the additional complexity involved in implement-187 ing such two-stage fitting procedures appears to have discouraged their use (McGraw & 188 Barnes, 2018) compared to simpler model-based analyses framed in terms of Granger causal-189 ity (Granger, 1969). 190

In this article, we investigate possible solutions to the above limitations through 191 the use of Bayesian methods (Uusitalo, 2007). In the Bayesian framework for learning 192 the structure and effect measures (Spiegelhalter & Lauritzen, 1990; Dawid & Lauritzen, 193 1993; Madigan et al., 1995), it is natural to use the posterior probability of a given net-194 work and its associated parameters as a score to measure model fitness (Buntine, 1991; 195 Cooper & Herskovits, 1992; Geiger & Heckerman, 1994; Heckerman et al., 1995). Ex-196 isting knowledge and constraints may be incorporated through the use of suitable prior 197 distributions, although in practice this must be balanced against computational feasi-198 bility. Bootstrap (Friedman et al., 1999) or sampling-based (Madigan & Raftery, 1994; 199 Madigan et al., 1995; Godsill, 2001) methods provide some measure of the uncertainties 200 in model selection, as well as allowing the predictions of multiple models to be combined 201 via model averaging (Madigan & Raftery, 1994). Consequently, the Bayesian approach 202 provides a principled quantification of the uncertainties associated with estimation of the 203 network structure and parameters. Models typically used for testing for Granger causal 204

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relationships, that is, linear autoregressive models, can be straightforwardly expressed
in graphical terms (Arnold et al., 2007; Lèbre, 2009), and, at least under a choice of conjugate priors, analyzed efficiently using closed-form expressions for the desired posterior
densities (Geiger & Heckerman, 1994). In this sense, the model specification is largely
familiar, reducing the barriers to use in climate applications. While there is additional
complexity associated with the inference scheme, we believe that this is justified by the
need for some level of uncertainty quantification.

To illustrate the utility of Bayesian methods for structure learning, we consider the 212 application of fitting BNs to a set of teleconnection indices derived from two reanaly-213 sis datasets. Our purpose in doing so is two-fold. Firstly, as the score-based approach 214 has not been widely used in analyses of this type, we wish to investigate the suitability 215 of the method for networks of the size encountered in realistic data. Our second, and more 216 important, aim is to perform a comparison of different reanalysis products on the ba-217 sis of the fitted networks. In addition to allowing for possible biases and differences in 218 the reanalyses at the level of interactions between modes to be studied, differences in the 219 networks from different reanalyses give some additional indication of uncertainties in the 220 observed networks that model runs are evaluated against. This, in turn, must be taken 221 into account when deciding what level of disagreement between models and observations 222 can be taken to indicate clear model biases. 223

The remainder of this paper is structured as follows. In the next section we provide a brief review of Bayesian network models and the inference methods used to fit such models. In Section 3 we describe the datasets and diagnostics that we study, together with our choice of prior distributions. In Section 4 we present the results of fitting the network models. Finally, we summarize our findings and discuss their implications for follow-up comparisons in Section 5.

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# 2 Dynamic Bayesian networks

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# 2.1 Structure learning

As noted above, a graphical model is simply a statistical model that has associated with it a graph encoding the relationships between the variables in the model. Each random variable in the model is represented by a node in the graph, with the allowed conditional dependence relationships between variables indicated by edges between nodes.

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Under suitable assumptions (see, e.g., Koller and Friedman (2009) for a review), translating between graphs and joint PDFs that constitute a model can be achieved using a
prescribed set of rules.

For BNs and other graphical models based on DAGs, the graph structure implies a factorization of the joint PDF into conditional density functions. Given a set of random variables  $Y^1, \ldots, Y^n$  and a DAG G, we denote by  $pa_G(Y^i)$  the set of nodes in Gthat have a directed edge connecting to  $Y^i$ ,

$$pa_G(Y^i) = \{Y^j | G \text{ contains an edge from } Y^j \text{ to } Y^i\}.$$
(1)

The graph G then provides a representation of the joint PDF  $P(Y^1, ..., Y^n)$  if the PDF admits a factorization of the form

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$$P(Y^1, \dots, Y^n) = \prod_{i=1}^n P(Y^i | \operatorname{pa}_G(Y^i), \theta_i), \qquad (2)$$

where  $\theta_i$  denotes any parameters required to characterize the conditional density. For example, when all of the variables in  $\{Y^i\} \cup \operatorname{pa}_G(Y^i)$  are discrete, the conditional density  $P(Y^i|\operatorname{pa}_G(Y^i), \theta_i)$  is the conditional probability table (CPT) summarizing the probability of observing each level of  $Y^i$  for each combination of values of the parents  $\operatorname{pa}_G(Y^i)$ ; the  $\theta_i$  are simply the values of each probability in the table. For continuous variables, the  $\theta_i$  are any parameters required to fully specify the corresponding continuous PDF.

Generally in geophysical applications, the random variables of interest exhibit non-253 trivial spatial and temporal correlations. In our case, these variables are a collection of 254 (continuous) teleconnection indices, some of which (e.g., ENSO) show substantial au-255 tocorrelation. Feedback loops or temporal dependence of this form cannot be represented 256 in a BN with a single node for each index  $Y^i$ , due to the requirement for the graph G 257 to be acyclic (Uusitalo, 2007). To handle this, the set of nodes is expanded to consist 258 of the values of the random variables  $Y^i$  at the current time t,  $Y_t^i$ , as well as the values 259 of the variables  $Y_{t-\tau}^i$  at previous times  $t-\tau$  (Kjærulff, 1995; Friedman et al., 1998; K. Mur-260 phy & Mian, 1999; K. P. Murphy & Russell, 2002), up to some maximum lag  $\tau_{\rm max}$ . Tem-261 poral dependencies are described by edges between the nodes corresponding to, say,  $Y_i^t$ 262 and  $Y_{t-\tau}^i$ , and a full time series of observations is described by a graph at each point in 263 time relating the values of the variables at that time point to those in the previous time-264 slices. Similar graphical models for multivariate time series have also been introduced 265 as time series graphs (Eichler, 2012). The resulting model is referred to as a dynamic 266 Bayesian network (DBN). 267

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In the simplest case, the structure of the graph and the associated parameters re-268 mains the same across all time-slices, so that the full time series is modeled by repeat-269 ing the graph at each time t. The corresponding DBN is said to be (time-)homogeneous 270 (K. P. Murphy & Russell, 2002). The assumption of homogeneity is often violated in geo-271 physical applications, however. Interactions between spatially separated processes, for 272 example, may be seasonally dependent, while on longer time-scales the possibility of cli-273 mate regime changes, e.g., in association with tipping points (Lenton et al., 2008), in re-274 sponse to anthropogenic forcing has recently become a key concern. Non-homogeneous 275 DBNs, in which either the graph structure, parameters or both simultaneously, are al-276 lowed to change over time, admit the possibility of modeling features such as secular trends 277 and regime changes (Wu et al., 2018), at the cost of a significant increase in complex-278 ity in terms of model specification and inference. Here we focus on the simpler case of 279 homogeneous models for the purposes of investigating the usefulness of Bayesian meth-280 ods for assessing model uncertainty; the more complicated case of non-homogeneous mod-281 els will be described in a separate study. 282

Fitting a homogeneous DBN to an observed time series  $D = \{y_1, \dots, y_T\}$ , where  $y_t$  denotes the values of the random variables  $Y_t = (Y_t^1, \dots, Y_t^n)^T$  at time t, requires learning (in general) both the structure of the graph G and the values of the corresponding parameters. Since

$$P(\theta, G|D) = P(\theta|G, D)P(G|D), \tag{3}$$

where  $\theta$  denotes the collection of all parameters for the conditional PDFs, the learning 288 process can be conveniently divided into two steps. In the first, structure learning stage, 289 the structure of the graph G is sought, independent of specific values of the parameters. 290 Structure learning methods for BNs can be roughly categorized as constraint-based or 291 score-based. The former set of methods attempt to reconstruct the graph structure on 292 the basis of conditional independence tests and available prior knowledge and constraints, 293 using, e.g., the PC-algorithm (Spirtes & Glymour, 1991) and its extensions (e.g., Colombo 294 & Maathuis, 2014; Runge, 2018a; Runge, Nowack, et al., 2019). As discussed in Section 1, 295 most recent examples in climate science have made use of constraint-based algorithms 296 to learn an initial structure, followed in some cases by a separate parameter learning step. 297 In contrast, in score-based approaches the graph G is estimated based on maximizing 298 a suitable score function (Cooper & Herskovits, 1992; Geiger & Heckerman, 1994; Heck-299 erman et al., 1995), such as the marginal likelihood P(D|G) or an information criterion. 300

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However, when there may be significant uncertainty associated with the selection of a

single model, rather than finding a single optimal model, it may be preferable to attempt

to account for this model uncertainty by sampling from the full posterior distribution

of possible graphs P(G|D) (Madigan et al., 1995). Estimates for derived quantities of

interest  $\Delta$  may then be computed by averaging over the posterior distribution (Madigan

<sup>306</sup> & Raftery, 1994; Draper, 1995),

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$$\Pr(\Delta|D) = \sum_{G \in \mathcal{G}} \Pr(\Delta|G, D) P(G|D) \approx \frac{1}{S} \sum_{s=1}^{S} \Pr(\Delta|G^{(s)}, D),$$
(4)

where  $\mathcal{G}$  is the space of allowed structures and  $\{G^{(s)}\}_{s=1}^{S}$  is a sample of size S from the posterior distribution P(G|D). In particular, structural uncertainties may be quantified by taking  $\Delta$  to be an indicator function for the presence of a given edge, with Eq. (4) quantifying the posterior probability of the existence of that edge, given the chosen model class and observed data.

Directly sampling the posterior P(G|D) can be achieved in the case that the marginal likelihood

$$P(D|G) = \int d\theta P(D|G,\theta) P(\theta|G)$$
(5)

can be evaluated, where  $P(\theta|G)$  denotes a set of priors for the full set of node PDF parameters  $\theta$  conditional on the structure of the graph, and we have used the shorthand  $\int d\theta$  to denote marginalization. The factor  $P(D|G, \theta)$  is simply the likelihood under the model,

$$P(D|G,\theta) = \prod_{t=1}^{T} \prod_{i=1}^{n} P(Y_t^i | \operatorname{pa}_G(Y_t^i), \theta_i);$$
(6)

note that the inner multiplication follows from the assumption that the joint PDF can 321 be factored according to G, and the outer multiplication from the assumption of homo-322 geneity. For simplicity, we work with the conditional likelihood assuming a sufficiently 323 large set of pre-sample values are available to condition on. We restrict our attention to 324 structures in which the parent set of a variable  $Y_t^i$  is not allowed to contain variables at 325 the same time t, that is, we do not allow contemporaneous dependencies among variables. 326 Excluding models with instantaneous links ensures that the structures we allow natu-327 rally satisfy structural modularity, such that the parent set of a variable  $Y_t^i$  may be cho-328 sen independently of the parent set for any other variable (Friedman & Koller, 2003). 329 However, it should be kept in mind that doing so prevents any interactions that occur 330 on time-scales that are shorter than the data sampling frequency from being directly han-331 dled, as these would manifest as instantaneous dependencies or feedback loops at  $\tau =$ 332

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0. Assuming that the priors  $P(\theta|G)$  satisfy the properties of parameter independence (Heckerman et al., 1995),

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$$P(\theta|G) = \prod_{i=1}^{n} P(\theta_i|G), \tag{7}$$

and modularity (that is, for any two graphs G and G', if  $Y_t^i$  has the same parent set in

G and G', then  $P(\theta_i|G) = P(\theta_i|G')$ , the marginal likelihood may be written as the prod-

uct of local marginal likelihoods  $\Psi_i(D,G)$  (Grzegorczyk & Husmeier, 2011):

$$P(D|G) = \prod_{i=1}^{n} \int d\theta_i \prod_{t=1}^{T} P(Y_t^i | \text{pa}_G(Y_t^i), \theta_i) P(\theta_i | G) \equiv \prod_{i=1}^{n} \Psi_i(D; G).$$
(8)

340 For structurally modular priors of the form

$$P(G) = \prod_{i=1}^{n} P(\operatorname{pa}_{G}(Y_{t}^{i}))$$
(9)

the posterior over graphs also factorizes,

$$P(G|D) = \frac{P(D|G)P(G)}{P(D)} = \frac{1}{P(D)} \prod_{i=1}^{n} \Psi_i(D;G)P(\operatorname{pa}_G(Y_t^i)),$$
(10)

<sup>344</sup> so that each factor can be computed independently, up to an overall normalization.

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## 2.2 Choice of conditional densities

The conditional densities  $P(Y^i|\text{pa}_G(Y^i), \theta_i)$  can, in principle, be chosen to model 346 arbitrary relationships between the random variables in the graph, consistent with the 347 independence assumptions embodied by the graph structure. In practice, this prevents 348 marginalizing out the graph parameters (i.e., evaluating  $\Psi_i(D;G)$  analytically) to sam-349 ple from the marginal posterior distribution P(G|D) directly. It is then necessary to con-350 struct a Markov Chain Monte Carlo (MCMC) sampler that samples from the joint pos-351 terior  $P(\theta, G|D)$  using, e.g., reversible jump MCMC (RJMCMC) (Green, 1995) or re-352 lated methods (Carlin & Chib, 1995; Godsill, 2001). A summary of samplers that we use 353 is given in Appendix A. 354

In special cases, the necessary integrals can be evaluated in closed form, allowing for the posterior P(G|D) to be efficiently sampled after marginalizing out the conditional PDF parameters. For continuous data, a widely used example for which this is possible is the linear Gaussian regression model (Punskaya et al., 2002; Lèbre et al., 2010). In this model, which can be regarded as a specialization of the BGe model for continuous data (Geiger & Heckerman, 1994), each  $Y_t^i$  is assumed to be conditionally Gaus361 sian distributed,

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$$Y_{t}^{i}|\mathrm{pa}_{G}(Y_{t}^{i}), \tau_{i}^{2} \sim N(\mu_{t}^{i}, \tau_{i}^{-2}),$$

$$\mu_{t}^{i} = \beta_{0}^{i} + \sum_{j=1}^{p_{i}} \beta_{(k_{j}, \tau_{j})}^{i} Y_{t-\tau_{j}}^{k_{j}}$$
(11)

with mean  $\mu_t^i = E[Y_t^i| \operatorname{pa}_G(Y_t^i)]$  given by a linear function of the parent variables  $\operatorname{pa}_G(Y_t^i) = \{Y_{t-\tau_j}^{k_j} | j = 1, \ldots, p_i\}$ . The local marginal likelihoods  $\Psi_i(D; G)$  and posterior distributions for the parameters of a given graph can be analytically evaluated provided that conjugate normal-gamma priors are assumed for the conditional precision  $\tau_i^2$  and coefficients  $\beta_{(k_j,\tau_j)}^i$ ,

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 $\tau_i^2 \sim \text{Gamma}(a_{\tau}, b_{\tau}),$   $\beta_{(k_j, \tau_j)}^i | \tau_i^2, \text{pa}_G(Y_t^i) \sim N\left(0, \frac{\nu_i^2}{\tau_i^2}\right), \quad j = 1, \dots, p_i,$ (12)

where  $a_{\tau}$ ,  $b_{\tau}$ , and  $\nu_i^2$  are prior hyperparameters. Similar linear models have previously 369 been applied (Kretschmer et al., 2016, 2017; Saggioro et al., 2020; Di Capua et al., 2020) 370 to perform estimation of the interaction strengths after an initial stage of constraint-based 371 structure learning. Using the underlying generative model, Eq. (11), posterior estimates 372 for both the structural features and model parameters can be obtained within a single 373 sampling scheme. Appropriate choices for the hyperparameters  $a_{\tau}$ ,  $b_{\tau}$ , and  $\nu^2$  allow vary-374 ing levels of regularization to be imposed so as to yield more reliable, if more conserva-375 tive, estimates given relatively short and noisy time series. Alternatively, they may be 376 allowed to vary and another level of priors specified for the unknown hyperparameters. 377

In the presence of significant non-linearity, linear models of this form may no longer 378 be appropriate. In this case, one strategy to capturing the underlying non-linear rela-379 tionship is to first discretize the original data and employ models for discrete data, e.g., 380 the analytically tractable BDe model (Buntine, 1991; Cooper & Herskovits, 1992; Heck-381 erman et al., 1995). However, this comes at the cost of necessarily losing some informa-382 tion, generally yielding only a coarse approximation of the original continuous distribu-383 tion (Friedman & Goldszmidt, 1996). Additionally, choosing an appropriate discretiza-384 tion scheme is in general difficult, as there is an inevitable trade-off between having suf-385 ficient resolution to describe the data versus the exponential growth in the number of 386 parameters that are required to specify the CPT of each variable. Hence, in the follow-387 ing we focus on the case of continuous data. 388

For a given choice of model, after performing any initial pre-processing, fitting the homogeneous DBN model to the observed indices data then consists of sampling from

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the posterior distribution P(G|D) using the known marginal likelihood P(D|G) in order to derive a set of candidate networks. For a given graph drawn from P(G|D), the posterior distribution for the parameters, and hence summary statistics and credible intervals, can be computed analytically or obtained via standard within-model MCMC methods.

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# 2.3 Toy model example

To make the above discussion more concrete, it is helpful to consider a simple toy 397 example. A standard problem in climate studies is to determine the direction of the re-398 lationships, if any, between some set of variables. For example, during the positive (El 399 Niño) phase of ENSO, anomalously warm Pacific sea surface temperatures (SSTs) drive 400 elevated mean surface temperatures over North and South America (McGraw & Barnes, 401 2018). As a simplified example of this sort of driver-response relationship, McGraw and 402 Barnes (2018) considered a two-dimensional system consisting of two observables  $D_t$  and 403  $R_t$  that evolve according to 404

$$D_t = \alpha D_{t-1} + \sqrt{1 - \alpha^2} \epsilon_t^D,$$

$$R_t = D_{t-\tau} + \gamma \epsilon_t^R.$$
(13)

The innovations  $\epsilon_t^D$  and  $\epsilon_t^R$  are taken to be independent Gaussian noise drawn from a standard normal distribution. Typically in climate analyses, the relationships between the system variables would be studied by regressing the postulated response (e.g., surface temperature anomalies or  $R_t$ ) on lagged values of the driver (SST anomalies or  $D_t$ ),

$$R_t = c_0 + \sum_{j=1}^{\kappa} c_j D_{t-j},\tag{14}$$

and vice versa to test for the possibility of the reversed relationship. However, this approach is susceptible to detecting spurious relationships when one or both processes exhibit substantial autocorrelation. As noted above, this often occurs in climate applications, where, e.g., the driver may correspond to relatively slowly varying boundary conditions such as SST driving an atmospheric response, as in the ENSO-surface temperature example. To account for this, autoregressive models of the form

$$R_t = a_0 + \sum_{j=1}^k a_j R_{t-j} + \sum_{j=1}^k b_j D_{t-j},$$
(15)

(and similarly for the dependence of  $D_t$  on  $R_t$ ), may be used instead to test for the presence of Granger causal links between processes.

Depending on the system, however, there can be considerable uncertainty associ-420 ated with selecting one of these models over the other. We can apply the sampling based 421 approach described in the previous sections to better quantify this. The two models, Eq. (14) 422 and Eq. (15), correspond to particular choices of parent set under the linear Gaussian 423 regression model. Thus, to illustrate the method we consider the results of learning the 424 structure of a DBN describing the system Eq. (13) under a linear Gaussian model. For 425 given values of the system parameters, we generate a random realization of the system 426 and fit a DBN after standardizing the input data to zero mean and unit variance. For 427 the prior hyperparameters, we take, for example,  $a_{\tau} = 1.5$ ,  $b_{\tau} = 10$ , and  $\nu^2 \approx 43.3$ , 428 yielding a weakly informative t prior with 3 degrees of freedom for each coefficient, with 429 90% prior credible intervals of  $-4 \leq \beta \leq 4$ . The prior 1% and 99% percentiles for the 430 conditional precision (variance) are 0.57 (0.02) and 56.7 (1.74), respectively. Alternative 431 choices with a heavier tailed distribution for the coefficients and much broader priors for 432 the conditional precision, e.g.,  $a_{\tau} = 0.5, b_{\tau} = 10$ , and  $\nu^2 \approx 2$ , tend to yield similar 433 parameter estimates but less sparse models. The set of models considered includes all 434 lags of up to 6 time steps and at most 4 parent nodes for each variable. Posterior sam-435 ples are obtained using the  $MC^3$  algorithm described in Appendix A, with the choices 436 for the structure priors and proposal densities as described in Section 3.3. 437

Graphical summaries of the results of fitting this Gaussian DBN to realizations of 438 the system Eq. (13) for  $(\alpha, \gamma, \tau) = (0.2, 1, 2)$ , (0.5, 10, 2), and (0.9, 4, 2) are shown in Fig-439 ure 1. The true system, Eq. (13), corresponds to the graph given in Figure 1(a), with 440 the edges labeled by the values of the standardized regression coefficients. The remain-441 ing panels show the estimated posterior probabilities  $\hat{\pi}$  for each edge, computed as in 442 Eq. (4), for each realization of the system; for clarity, only edges for which  $\hat{\pi} > 0.5$  are 443 shown. Where an edge also appears in the maximum a posteriori (MAP) estimate for 444 the structure, the posterior 95% highest density interval (HDI) is also shown for the cor-445 responding coefficient. 446

For low levels of autocorrelation and noise, the model recovers the correct edges with virtual certainty; in this case, the MAP structure consists of the (true) parent sets pa $(D_t) = \{D_{t-1}\}$  and pa $(R_t) = \{D_{t-2}\}$ . Recovery of the true dependence structure is more difficult in the presence of large amounts of noise (e.g.,  $\gamma = 10$ ) or high autocorrelation (e.g.,  $\alpha = 0.9$ ). McGraw and Barnes (2018) note that, in the latter case, a bivariate Granger causality analysis exhibits reduced power for detecting the  $D_{t-\tau} \rightarrow$ 

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Figure 1. DBNs corresponding to (a) the true model given by Eq. (13), and network structures inferred using a linear Gaussian model for a single realization of the process with (b)  $\alpha = 0.2, \gamma = 1, \tau = 2$ ; (c)  $\alpha = 0.5, \gamma = 10, \tau = 2$ ; and (d)  $\alpha = 0.9, \gamma = 4, \tau = 2$ . For each of the fitted networks, all edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown, with thickness indicating the value of  $\hat{\pi}$ . Where an edge also occurs in the MAP model, the approximate posterior 95% HDI for the corresponding coefficient is also shown.

 $R_t$  edge while lagged regressions tend to too frequently identify spurious wrong-way re-453 lationships in which R drives D. For high levels of noise, both methods show reduced 454 power for detecting the driver-response relationship, although accounting for autocor-455 relation successfully controls the false-positive rate for the reversed relationship. This 456 difficulty in identifying the true underlying model is clearly evident in the lower panels 457 of Figure 1 from the reduced estimated posterior probability of a dependence of  $R_t$  on 458  $D_t$ . In both cases, although the MAP structure suggests that  $R_t$  depends on lagged val-459 ues of  $D_t$ , the lag at which this occurs is misidentified, with  $pa(R_t) = \{D_{t-1}\}$ . Across 460 all sampled models, the estimated posterior probability for lagged values of  $D_t$  to  $R_t$  re-461 flects greater uncertainty in the model structure, and hence in the selection of a single 462 optimal model. For instance, while the MAP parent set for  $R_t$  when  $\alpha = 0.5$ ,  $\gamma = 10$ 463 is found to contain only  $D_{t-1}$ , the next most likely parent sets (for this particular sam-464

ple, pa $(R_t) = \emptyset$  and pa $(R_t) = \{D_{t-2}\}$  also have non-negligible posterior probabilities, so that in practice it may be important to take this model uncertainty into account. A similar uncertainty measure is considerably more difficult to construct and interpret in the testing based approach of McGraw and Barnes (2018). As time series with long memory or large noise levels are frequently encountered in the climate analyses that we are interested in, the ability to provide some formal estimate of model uncertainty is an important advantage of the approach adopted here.

# **3** Data and methods

We now describe an application of the above methods to sets of reanalysis-derived teleconnection indices. In addition to uncertainties due to model and parameter selection, in practice the reanalysis datasets that we use have associated uncertainties as well. Comparison of the networks derived from different products over a common timespan allows the role of these differences to be investigated and provides a baseline against which free-running model simulations can be compared.

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#### 3.1 Data

The data that we analyze are obtained from the Japanese 55-year Reanalysis (JRA-55) and the National Centers for Environmental Prediction/National Center for Atmospheric Research (NCEP/NCAR) Reanalysis 1 (NNR1).

The NCEP/NCAR Reanalysis 1 (Kalnay et al., 1996) is an atmospheric reanaly-483 sis covering the years 1948 to present. The data assimilation system employs a global 484 spectral model with a T62 resolution on 28 vertical levels, and assimilates surface and 485 atmospheric observational data. While a fixed analysis and forecast system is used for 486 the duration of the reanalysis, changes in observing systems still have an impact and, 487 consequently, the reanalysis is less reliable in the first decade than at later times (Kistler 488 et al., 2001). NNR1 represents a first generation reanalysis providing a multidecadal record 489 of the atmospheric state, albeit with several known errors (Kistler et al., 2001) and bi-490 ases, particularly in data-sparse regions in the high latitudes and the Southern Hemi-491 sphere (SH) (see, e.g., Hines et al., 2000; G. J. Marshall & Harangozo, 2000; G. J. Mar-492 shall, 2002; Bromwich & Fogt, 2004; Greatbatch & Rong, 2006; Hertzog et al., 2006; Bromwich 493 et al., 2007; Lindsay et al., 2014). For the purposes of our analysis, global fields of daily 494

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mean 500 hPa geopotential height ( $Z_{q500 \text{ hPa}}$ ), zonal winds at 850 hPa and 200 hPa ( $u_{850 \text{ hPa}}$ ) 495 and  $u_{200 \text{ hPa}}$ , mean sea level pressure (MSLP), and surface zonal and meridional winds 496  $(u_{\rm sfc} \text{ and } v_{\rm sfc})$  are obtained on the provided  $2.5^{\circ} \times 2.5^{\circ}$  latitude-longitude grid. Daily 497 mean top-of-atmosphere outgoing longwave radiation (OLR) fields are provided on a T62 498 Gaussian grid and are subsequently regridded to a  $2.5^{\circ} \times 2.5^{\circ}$  latitude-longitude grid 499 using a bilinear interpolation scheme. To compute indices of tropical variability based 500 on SST data for NNR1, we use version 1.1 of the HadISST SST dataset (Rayner et al., 501 2003), which provides monthly global SST on a  $1^{\circ} \times 1^{\circ}$  latitude-longitude grid from 1870 502 to present. 503

The JRA-55 reanalysis (Kobayashi et al., 2015), covering the period from 1958 to 504 present, is a more recent atmospheric reanalysis product that aims to correct issues found 505 in previous reanalyses. As for the NNR1 reanalysis, a frozen analysis system is employed 506 and atmospheric and surface observations are assimilated. The assimilation system used 507 for JRA-55 employs a TL319 resolution operational system with 60 vertical levels. The 508 representation of the atmospheric circulation has been found to be greatly improved com-509 pared to previous generation reanalyses, although there remain known biases (Harada 510 et al., 2016). Daily mean  $Z_{q500 \text{ hPa}}$ ,  $u_{850 \text{ hPa}}$ ,  $u_{250 \text{ hPa}}$ ,  $u_{\text{sfc}}$ ,  $v_{\text{sfc}}$ , MSLP, and OLR fields 511 are obtained on a  $1.25^{\circ} \times 1.25^{\circ}$  latitude-longitude grid. For SST fields, the model sur-512 face brightness temperature provided on a  $1.25^{\circ} \times 1.25^{\circ}$  latitude-longitude grid is used. 513 Where required by the definition of the index as noted below, we regrid the initial fields 514 to a  $2.5^{\circ} \times 2.5^{\circ}$  latitude-longitude grid using a bilinear interpolation method. 515

## 3.2 Indices

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From the full gridded fields we compute a set of indices diagnosing the activity of 517 a selection of major global teleconnections, which will form the nodes in the fitted graph-518 ical models. In a fully data-driven approach, the definitions of such indices might be au-519 tomatically determined by using community detection methods (Steinhaeuser et al., 2011; 520 Bello et al., 2015). This has the advantage of accounting for differences between datasets 521 or models in the representation of particular modes, e.g., due to shifts in the geographic 522 centers of action. While these approaches have been employed in studies of causal effect 523 networks (Kretschmer et al., 2017), here we use a set of fixed, empirical definitions for 524 the teleconnection indices. We do so for two reasons. Firstly, this allows for a simpler 525 evaluation of the performance of the models defined in Section 2, as the features in the 526

fitted networks can be directly compared to well-studied relationships between traditional
indices. Such a comparison would be more difficult to perform when using automatically
extracted indices, as some differences might arise solely from the definition of the index.
Additionally, positional shifts and other differences in the expression of particular modes
with respect to a predefined diagnostic are themselves of interest from the point of view
of comparing and evaluating models, and so we prefer to use a single set of reference definitions.

We choose a set of indices that provides reasonably comprehensive coverage of the dominant teleconnection processes active on intraseasonal through to interannual timescales. Where anomalies are required in the definition of an index, for consistency across the different datasets we compute all anomalies as differences from the daily or monthly climatology calculated with respect to the reference period 1 January 1979 to 30 December 2001.

As measures of tropical variability, we include an updated version (Zhang et al., 540 2019) of the multivariate ENSO index (MEI) (Wolter & Timlin, 1993, 1998, 2011), the 541 dipole mode index (DMI) to characterize IOD activity (Saji et al., 1999), and the Wheeler-542 Hendon Madden-Julian oscillation (MJO) index (Wheeler & Hendon, 2004), denoted be-543 low as RMM1 and RMM2. For both the MEI and the RMM1 and RMM2 indices, all 544 of the input fields are evaluated on a common  $2.5^{\circ} \times 2.5^{\circ}$  latitude-longitude grid. Where 545 required, monthly MJO indices are defined as the monthly mean of the corresponding 546 daily index. 547

In the extratropical atmosphere, we include indices of the AO, the SAM, the PNA, 548 the PSA1 and PSA2 modes, and a set of modes associated with blocking in the North 549 Atlantic and western Europe. We define the AO and SAM as the leading empirical or-550 thogonal functions (EOFs) (Lorenz, 1956) of anomalies of monthly mean  $Z_{g500 \text{ hPa}}$  pole-551 ward of  $20^{\circ}$ N and  $20^{\circ}$ S, respectively. All anomalies are weighted by the square root of 552 the cosine of the gridpoint latitude when computing the EOFs. Corresponding AO and 553 SAM indices are calculated by projecting the (area-weighted) daily or monthly anoma-554 lies onto the leading EOF and normalizing by the standard deviation of the monthly lead-555 ing principal component (PC). 556

The PNA pattern is taken to be the leading mode obtained after performing a VARI-MAX rotation (Kaiser, 1958) of the first 10 EOF modes of monthly-standardized anoma-

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lies of monthly mean  $Z_{g500 \text{ hPa}}$  polewards of 20°N during boreal winter, taken to be De-559 cember, January and February (DJF). The PNA index is then the projection of the stan-560 dardized height anomalies onto the resulting pattern, standardized by the monthly mean 561 and standard deviation within the climatology reference period. The analogous modes 562 in the SH, PSA1 and PSA2, are defined as the second and third modes in an EOF anal-563 ysis of year-round anomalies of daily mean  $Z_{q500 \text{ hPa}}$  polewards of 20°S, projecting onto 564 each mode and normalizing by the standard deviation of the corresponding PC over the 565 reference period to obtain associated indices. 566

Following the method presented in Straus et al. (2017), we define a set of Euro-Atlantic 567 circulation regimes via a k-means clustering analysis of the leading 24 PCs of boreal win-568 ter anomalies in daily mean  $Z_{q500 \text{ hPa}}$  in the sector 20°N - 80°N, 90°W - 30°E, after ap-569 plying a 10 day running mean smoothing. This method has the advantage of better cap-570 turing spatial asymmetries present in opposing phases of the NAO. Using k = 4 clus-571 ters, we obtain a set of patterns that correspond to the positive and negative NAO phases. 572 NAO<sup>+</sup> and NAO<sup>-</sup>, as well as two clusters associated with blocking events in the Atlantic 573 and western Europe, denoted AR and SCAND, respectively. Indices for each of the 4 clus-574 ters are obtained by projecting daily or monthly height anomalies onto the composites 575 associated with each cluster and standardizing by the monthly mean and standard de-576 viations of the monthly index over the reference period. 577

For both reanalyses, monthly time series of the chosen indices are computed for the period 1 January 1960 to 30 November 2005, and this full time period is used for fitting DBNs. As some indices (e.g., SAM) exhibit significant trends over this period, we estimate and remove a linear trend for every index beforehand. Each index time series is then standardized to have zero mean and unit variance over the fitting period.

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# 3.3 Priors and sampler settings

As noted in Section 2, when excluding the possibility of contemporaneous edges,

- it is natural to choose structurally modular priors for the parent sets  $pa_G(Y_t^i)$ , such that
- the prior for a graph G decomposes into independent priors for the parent sets  $pa_G(Y_t^i)$
- of each of the *n* indices included in the model. We fix a maximum allowed lag  $\tau_{\rm max}$ , such
- that the  $n(\tau_{\max} + 1)$  nodes in the network at a given time are  $\bigcup_{i=1}^{n} \{Y_t^i, \dots, Y_{t-\tau_{\max}}^i\}$ .
- For networks based on monthly indices, we take  $\tau_{\rm max} = 6$  months, corresponding to the

approximate *e*-folding time of the MEI. To enforce some degree of sparsity in the networks, we also impose a constraint on the maximum parent set size for each index,  $|pa_G(Y_t^i)| \le p_{\max} = 10$ . Subject to these constraints, in the absence of additional information we adopt uniform priors over the set of possible parent sets for each index, i.e.,

$$P(\operatorname{pa}_{G}(Y_{t}^{i})) = \begin{cases} \left[\sum_{j=0}^{p_{\max}} \binom{n\tau_{\max}}{j}\right]^{-1}, & |\operatorname{pa}_{G}(Y_{t}^{i})| \le p_{\max}, \\ 0, & \text{otherwise.} \end{cases}$$
(16)

The MCMC sampling schemes described in Appendix A also require an appropriate proposal density for proposing updates to the structure of the model. We choose to adopt a uniform proposal density on graphs G' in the neighborhood of the current graph G,

<sup>599</sup> 
$$q_G(G';G) = \begin{cases} \frac{1}{|\mathrm{nhd}(G)|}, & G' \in \mathrm{nhd}(G), \\ 0, & \mathrm{otherwise.} \end{cases}$$

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The neighborhood nhd(G) of a graph G is defined as the set of graphs that can be reached 600 from that structure by a single move in a predefined move set. Note that, when an ex-601 plicit distinction is made between structure and parameter updates, as in the basic RJM-602 CMC scheme in Appendix A, we do not include the current graph itself in the neighbor-603 hood. We take the set of possible moves to consist of addition of a single edge, deletion 604 of a single edge, or an exchange of two edges (Grzegorczyk & Husmeier, 2011). The neigh-605 borhood of a graph contains those graphs that can be reached from it by performing one 606 of these three moves, subject to the imposed condition on the maximum parent set size. 607 A structure update move thus consists of determining the neighborhood of the current 608 graph based on the available moves before selecting with equal probability one graph from 609 this set. As the models considered here allow for the node parameters to be marginal-610 ized out, we use the  $MC^3$  sampling algorithm (Algorithm 2 in Appendix A) with the above 611 proposal for fitting the model. For each index, posterior samples were obtained by run-612 ning 8 chains of length  $10 \times 10^6$  samples. The number of samples to be retained for anal-613 ysis was chosen based on the estimated convergence rate from short initial runs follow-614 ing Brooks et al. (2003), where we required the thinning to be such that the resulting 615 dependence between samples was reduced by a factor of 100 compared to the dependence 616 between successive samples. To provide some assessment of chain convergence, homo-617 geneity of the distribution of parent sets within chains was monitored using  $\chi^2$  and Kolmogorov-618 Smirnov tests (Brooks et al., 2003) for each index, although we note that, as always, these 619 tests are not sufficient alone to determine approximate convergence (Sisson, 2005). In 620

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the following, both the posterior distribution over possible models and for individual features is of interest. Thus, in addition to applying these tests to the distribution of overall model indicators, we also monitored the results for the case when each model was labelled by a binary indicator for the presence of each possible edge, so as to check if inferences for the individual posterior edge probabilities were homogeneous across chains. Where required, samples for the node parameters were drawn from the conditional posterior distributions by Gibbs sampling.

The associations between the continuous valued indices are modeled using the lin-628 ear Gaussian conditional density, Eq. (11). For each index, we take the hyperparame-629 ter values  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu_i^2 = 3$ , for  $i = 1, \ldots, n$ . This corresponds to inde-630 pendent  $t_3$  marginal priors for the regression coefficients, with 95% prior HDI  $-1 \leq \beta \leq$ 631 1. The 1% and 99% percentiles for the conditional precision are respectively 1.1 and 113.4. 632 Prior simulations suggest that this choice of priors yields a reasonable scale for the prior 633 predictive distribution for the 1-step ahead forecast values, while not being so heavily 634 regularized that relatively large values of the coefficients and indices are excluded. Sim-635 ilarly, the conditional precision hyperparameters are chosen to yield typical monthly in-636 novation variances of order 1 or less. As these choices lead to somewhat informative pri-637 ors, to assess the sensitivity of the results to the hyperparameter values we also perform 638 fits with the much more weakly informative choices of  $a_{\tau} = 0.5$ ,  $b_{\tau} = 10$ ,  $\nu_i^2 \approx 2$  (cor-639 responding to a 90% prior HDI of  $-4 \leq \beta \leq 4$  and prior 1% and 99% percentiles for 640  $\tau^2$  of 7.6×10<sup>-4</sup> and 33.2, respectively), which we find lead to qualitatively similar re-641 sults (see supporting information). 642

#### 643 4 Results

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# 4.1 Full year networks for monthly indices

We first consider the results of applying the above methods to year-round monthly indices for the two reanalyses. To better display the spatial structure of interactions, we separate the full network into subgraphs consisting of those indices corresponding to Northern Hemisphere (NH) extratropical, tropical, and SH extratropical modes, together with their inferred parent sets. From the posterior samples over possible parent sets for each index, approximate posterior probabilities  $\hat{\pi}$  for the presence of each edge in the network are computed as the sample average, Eq. (4), of the corresponding indicator function.

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The results are summarized in Figure 2 and Figure 3 for the possible parent sets of trop-652 ical indices, in Figure 4 for the NH extratropical indices, and in Figure 5 for the SH ex-653 tratropical indices. For each subset, edges are shown with weights corresponding to their 654 estimated posterior probability; for clarity, only those edges for which this value is at least 655 0.5 are shown. Overall, the learned structures for the two reanalyses are similar. In par-656 ticular, there is broad agreement in terms of those edges for which the estimated pos-657 terior probability is high. These edges largely correspond, as is expected, to relationships 658 in which there is a strong association between the parent and child node, either through 659 strong autocorrelation or Granger causal links. 660

Strong evidence of long-range dependence on lagged values, reflecting the expected 661 high levels of autocorrelation, is apparent for indices representing ENSO and the MJO, 662 with the fitted models for both products featuring posterior probabilities near one for 663 dependence of the MEI and both RMM indices on lags of up to four months (Figure 2 664 and Figure 3). In both cases, the preferred lags are consistent with the expected time-665 scales for these modes, noting that a maximum lag of 6 months is imposed when fitting 666 the model. Purely atmospheric modes, on the other hand, exhibit little memory beyond 667 time horizons of several weeks to a month. This is apparent in the fitted networks for 668 both models in the absence of strong evidence for dependence of the NH extratropical 669 (Figure 4) and SH extratropical indices (Figure 5) on lagged values of themselves beyond 670 lags of one month. For instance, in both NNR1 and JRA-55 there is found to be little, 671 or at most relatively weak, evidence for serial dependence on monthly time-scales for the 672 AR and SCAND indices, which respectively represent blocking in the Atlantic and Scan-673 dinavia with a characteristic time-scale of 7 - 10 days. Although the DMI diagnoses In-674 dian Ocean SST variability, the partial autocorrelation structure of the monthly mean 675 DMI is consistent with an AR(1) process, and hence only the edge  $DMI_{t-1} \rightarrow DMI_t$ 676 is found to have appreciable posterior mass. 677

Where a high posterior probability edge is inferred between distinct indices, the edge generally matches with well-known associations or teleconnections among the modes included in the fit. This is highlighted, for example, in Figure 6, where the posterior probabilities for edges entering the AO, NAO<sup>+</sup>, and NAO<sup>-</sup> nodes in both reanalyses are shown. The close association between the AO and NAO is clearly evident in both reanalyses, and the estimated dependence relationships involving just these two indices are in essentially exact agreement. For both NNR1 and JRA-55, the presence of an edge in the

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Figure 2. Subgraphs corresponding to the fitted parent sets of the tropical indices in NNR1 based on full-year data for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.

network from the AO at lag 1 to each of the child nodes  $AO_t$ ,  $NAO_t^+$ , and  $NAO_t^-$  is in-685 ferred with very high confidence. The same is true for the edges  $\mathrm{NAO}^-_{t-1} \to \mathrm{AO}_t$  and 686  $\text{NAO}_{t-1}^- \rightarrow \text{NAO}_t^+$ . This clear interdependence of the AO and NAO is, of course, very 687 well established, to the extent that the existence of the former as a distinct physical mode 688 has been debated (Deser, 2000; Ambaum et al., 2001). The appearance of this relation-689 ship in the learned structures does, however, provide a useful check that the method re-690 covers the expected relationships among particular modes. That these features agree be-691 tween the two reanalyses also suggests that the relationships between these modes are 692 consistent in the separate datasets. 693

Similarly, the fitted parent sets for the monthly PNA index in NNR1 and JRA-55, 694 shown in Figure 7, both include with posterior probability greater than 0.5 dependence 695 on the value of the PNA index at a lag of one month ( $\hat{\pi} \approx 0.95$  and  $\hat{\pi} \approx 0.91$  in NNR1 696 and JRA-55, respectively) and on the MEI at lag one ( $\hat{\pi} \approx 0.71$  and  $\hat{\pi} \approx 0.59$  for NNR1 697 and JRA-55, respectively). In this case, the results of the fit indicate that in both re-698 analyses there is reasonable evidence that the PNA is associated with tropical forcing, 699 here captured by the lagged value of the MEI, consistent with previous observational and 700 modeling studies (e.g., Hoskins & Ambrizzi, 1993; Trenberth et al., 1998; Franzke et al., 701 2011). The posterior mass for this edge is similar in NNR1 and JRA-55, suggesting that, 702 as for the AO and NAO, this relationship is also consistent in the two products, in the 703 sense that there is comparable evidence for the association across the two datasets. 704

Although the structures estimated from the NNR1 and JRA-55 datasets agree well 705 in the above examples, there are notable differences in the estimated edge probabilities 706 as well. Moreover, several of these differences involve features that are assigned appre-707 ciable posterior probability (i.e.,  $\hat{\pi} \geq 0.5$ ) based on the data from one reanalysis and 708 not the other. That is, the differences are not limited only to associations that are weakly 709 supported in both datasets. For example, in Figure 7, the estimated structure for NNR1 710 contains an edge from the lag one monthly mean RMM1 index to the monthly PNA in-711 dex at lag 0 with high posterior probability. The same feature in JRA-55 is found to have 712 a posterior probability  $\hat{\pi} \approx 0.38$  that is less than 40% of that found for NNR1. Given 713 the observed evidence for interactions between tropical convection and extratropical modes 714 such as the PNA on intraseasonal time-scales (Lau & Phillips, 1986), in this case the weaker 715 evidence for an MJO-PNA relationship in the JRA-55 data may arise due to the sup-716 pressed MJO observed in the JRA-55 reanalysis (Harada et al., 2016). It should be noted, 717 however, that directly attributing the difference to this bias is not straightforward; for 718 example, other known associations between the MJO and other modes of variability are 719 identified in the JRA-55 fits, discussed below. Confirming whether there is a difference 720 in this particular feature, and the underlying source of the difference, would require a 721 closer examination of the representation of the two processes in the reanalysis, which is 722 left for further studies. Nevertheless, this does highlight that comparison of the learned 723 parent sets between the two reanalyses allows differences in the captured interactions be-724 tween modes to be identified, particularly when we may have confidence that such a dif-725 ference is in fact present. By adopting a Bayesian approach to structure learning, the 726

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<sup>727</sup> level of confidence in this difference can be estimated: while the edge  $\text{RMM1}_{t-1} \rightarrow \text{PNA}_t$ <sup>728</sup> may be found in individual models for the time-evolution of the PNA index in both re-<sup>729</sup> analyses, the existence of this feature is approximately 2.5 times more likely in NNR1 <sup>730</sup> than in JRA-55 given the class of models and possible parent nodes that we consider. <sup>731</sup> The presence of this edge with reasonable confidence in one product and not the other <sup>732</sup> is suggestive of an underlying bias in one reanalysis, rather than simply being due to sam-<sup>733</sup> pling variability.

Other differences in edges with large posterior mass, however, do not have quite 734 as clear possible sources in specific underlying biases. In both reanalyses, the parent set 735 of the NAO<sup>t</sup> node is estimated to contain the monthly PSA2 index at lag one, but with 736 a posterior probability that is approximately 1.8 times larger in NNR1 ( $\hat{\pi} \approx 0.89$  com-737 pared to  $\hat{\pi} \approx 0.50$ , see Figure 6). Similarly, the posterior probability for the edge  $PSA2_{t-1} \rightarrow$ 738  $NAO_t^-$  is  $\approx 4.7$  times larger in NNR1, but in this case for both NNR1 and JRA-55 the 739 estimated probability is less than 0.5, while an edge  $PSA_{t-1} \rightarrow AR_t$  is also present in 740 both reanalyses. O'Kane et al. (2016) and O'Kane et al. (2017) concluded that the PSA 741 modes predominantly reflect dynamics localized to within the SH waveguide, such that 742 a direct physical interaction between the PSA2 mode and the NAO is unlikely. Thus, 743 in contrast to the  $\text{RMM1}_{t-1} \rightarrow \text{PNA}_t$  feature, one might expect that the differing level 744 of confidence in this edge between the two reanalyses may be due to differences in rel-745 evant factors that are omitted from this simple analysis. In particular, as the fit shown 746 in Figure 6 is based on data from all seasons, the presence of such a feature may reflect 747 seasonal covariations in the extratropical circulation in both hemispheres that is other-748 wise unaccounted for here. While we consider the effects of seasonality in Section 4.2 be-749 low, ultimately direct determination of the root of these differences must be based on 750 detailed evaluation of the two products. Here we seek only to highlight the use of the 751 fitted networks for learning possible dependence relationships and hence their utility for 752 guiding comparative analyses on the basis of identifying relationships that can be inferred 753 to be present or absent with reasonable confidence. 754

Features for which there is lower confidence tend to differ more between the two reanalyses, although in these cases it becomes less clear as to whether they indicate substantive differences. This is most evident in the subgraphs corresponding to the parent sets for the SH extratropical indices in Figure 5. While there is agreement between the models fitted to NNR1 and JRA-55 in features such as the one month memory in the

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SAM and the presence of an association (even if not a direct interaction) between the 760 PSA1 and tropical forcing captured by the MEI, a larger number of features with pos-761 terior mass  $\hat{\pi} > 0.5$  are found using the NNR1 data, involving a larger set of indices 762 as parents. This greater disagreement between reanalyses, and overall lower confidence 763 in the inferred non-independence relationships, may in part be due to the disparate res-764 olutions and configurations of the respective atmospheric models in combination with 765 relatively sparse observations in the SH prior to the satellite period. Spurious associa-766 tions may also arise as a result of omitted factors, such as seasonal changes in the cir-767 culation, and low signal-to-noise ratios outside of the SH winter. For instance, the pos-768 terior distributions over parent nodes for the SAM index in the two reanalyses are sum-769 marized in Figure 8. In both NNR1 and JRA-55, non-zero associations are found with 770 the MEI at lags of 4 and 6 months, albeit with somewhat higher posterior probabilities 771 in JRA-55, and in NNR1 an edge from  $AR_{t-5} \rightarrow SAM_t$  is identified with  $\hat{\pi} \approx 0.52$ . 772 The same edge in JRA-55 is found to have  $\hat{\pi} \approx 0.42$ . Noting that a similar, compar-773 atively low confidence relationship is found between  $NAO^+$  and the PSA1 at a lag of four 774 months in NNR1, and absent a mechanism for such interactions, it appears likely that 775 in this case the association is an artifact arising from the use of year-round data; we con-776 firm that this is the case in the following section. It is worth noting, however, that the 777 fact that there is overall low posterior weight for these edges in both reanalyses allows 778 identifying them as lacking robustness. This in turn is of use for the purposes of guid-779 ing model evaluation, where differences to observations in low probability relationships 780 are of potentially less relevance. 781

It is important to bear in mind that the structures discussed above summarize the 782 marginal posterior probability for the presence or absence of each individual edge over 783 all possible models, rather than the presence or strength of an association between two 784 indices within a single model. In addition to inspecting the estimated marginal distri-785 butions for each feature, given a sample from the approximate posterior distribution it 786 may also be of interest to consider aspects of the sample that involve either the joint oc-787 currence of one or more edges, as well as the posterior distribution over complete mod-788 els. In particular, point-estimates for the parent set, analogous to those obtained using 789 790 constraint-based approaches, can be obtained as the MAP estimate with the largest posterior probability. As all parent sets are assigned equal prior probability, this structure 791 is simply the one that maximizes the marginal log-likelihood. Conditional on a partic-792

<sup>793</sup> ular MAP structure, estimates for the parameters under the model may be simply ob-

- tained by sampling from the conditional distributions  $P(\theta|G, D)$ , which in the case of
- the linear Gaussian model can be evaluated in closed form. Thus, in the sampling-based

approach we may also obtain estimates of the strength of an association, conditioned on
 a particular model, in addition to the above estimates for the probability of the presence

<sup>798</sup> of the corresponding edge in the structure.

<sup>799</sup> For example, the MAP parent sets for the NH extratropical modes in each reanal-

ysis are summarized in Table 1, where we show both the estimated posterior probabil-

ity for each edge in the parent set, as well as the posterior mean and 95% HDI for the

corresponding coefficient in Eq. (11). In general, edges present in the MAP parent sets

**Table 1.** MAP parent sets for monthly NH extratropical teleconnection indices across all seasons for NNR1 and JRA-55 for fits with  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ , showing the estimated posterior probability  $\hat{\pi}$  of the edge, the mean parameter value  $\hat{\beta}$  conditional on the MAP structure, and the 95% posterior HDI. Dashes indicate a node that is not in the MAP parent set for a given reanalysis.

		JRA	A-55	NNR1		
Parent node	$\hat{\pi}$	$\hat{eta}$	95% HDI	$\hat{\pi}$	$\hat{eta}$	95% HDI
$\overline{AO_t}$						
$AO_{t-1}$	1.00	0.45	(0.35, 0.56)	1.00	0.46	(0.35, 0.57)
$NAO_{t-1}^{-}$	0.99	0.24	(0.13, 0.35)	0.99	0.25	(0.13, 0.36)
$\overline{\mathrm{AR}_t}$						
$NAO_{t-6}^+$	0.44	0.12	(0.03, 0.19)	0.32	0.12	(0.04, 0.20)
$PNA_{t-1}$	0.58	0.14	(0.06, 0.22)	0.36	0.14	(0.06, 0.22)
$PSA1_{t-1}$	0.97	-0.17	(-0.25, -0.09)	0.67	-0.14	(-0.22, -0.06)
$\text{RMM2}_{t-5}$	0.11	-	_	0.43	-0.12	(-0.20, -0.04)
$\overline{NAO_t^+}$						
$AO_{t-1}$	1.00	0.35	(0.23, 0.46)	1.00	0.37	(0.24, 0.48)
$NAO_{t-1}^+$	0.66	0.17	(0.05, 0.29)	0.69	0.17	(0.06, 0.29)
NAO <sub>t</sub>	0.98	0.31	(0.19, 0.43)	1.00	0.34	(0.22, 0.46)
$PSA2_{t-1}^{t-1}$	0.50	-0.11	(-0.19, -0.03)	0.89	-0.14	(-0.22, -0.06)
$\overline{NAO_t^-}$						
$AO_{t-1}$	1.00	-0.52	(-0.65, -0.39)	1.00	-0.52	(-0.65, -0.39)
$NAO_{t-1}^{-}$	0.78	-0.22	(-0.34, -0.10)	0.70	-0.23	(-0.35, -0.11)
$PNA_{t-1}$	0.73	-0.17	(-0.26, -0.07)	0.55	-0.15	(-0.24, -0.06)
$PSA2_{t-1}$	0.10	-	-	0.47	0.12	(0.04, 0.20)
PNA <sub>t</sub>						
$MEI_{t-1}$	0.58	0.21	(0.12, 0.29)	0.71	0.21	(0.12, 0.29)
$PNA_{t-1}$	0.91	0.15	(0.07, 0.24)	0.95	0.16	(0.08, 0.24)
$\text{RMM1}_{t-1}$	0.38	—	-	0.99	0.17	(0.09, 0.25)
$SCAND_t$						
$MEI_{t-6}$	0.34	-0.13	(-0.21, -0.05)	0.28	-	-
$SCAND_{t-1}$	0.53	0.13	(0.04, 0.21)	0.31	-	
$MEI_{t-5}$	0.23	-	-	0.45	-0.13	(-0.22, -0.05)
$AR_{t-1}$	0.28	-	-	0.53	-0.13	(-0.21, -0.05)

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tend to have at least moderately high posterior probabilities, that is, they correspond

to non-independence relationships that are also present in a large fraction of the sam-804 pled models. Edges found in the MAP parent set are not necessarily found in the ma-805 jority of possible models, however, with several edges having posterior probabilities of 806 order  $\sim 0.3$  in Table 1. For these edges, while their inclusion leads to a good fit for this 807 particular sample, the data do not provide especially strong evidence for their presence 808 compared to other edges in the MAP structure, given the set of all other possible pre-809 dictors. To some extent this also reflects the fact that multiple models with parent sets 810 that do not include these edges may still provide a reasonable fit to the observed data, 811 despite not maximizing the marginal likelihood, and hence can account for non-negligible 812 posterior mass. Considering only the single most probable structure may therefore fail 813 to take into account relevant model uncertainty. Notably, differences between the MAP 814 structures for the two reanalyses are again not only restricted to edges with low poste-815 rior mass, suggesting that at least some differences may be due to systematic effects rather 816 than as a result of minor differences in the two samples. For example, the absence of the 817 edge  $\text{RMM}_{t-1} \rightarrow \text{PNA}_t$  in the MAP parent set from JRA-55 indicates that not only is 818 there weak evidence for this relationship in the dataset, but that it is also not required 819 in order to produce a good fit to the observed time series. Where an edge is present in 820 the MAP parent set for both reanalyses, there is good agreement between the two prod-821 ucts in terms of the estimated coefficient. For the NH modes, the two reanalyses yield 822 very similar estimates for the strength of each association, while there are somewhat larger 823 differences for the tropical and SH extratropical modes (see supporting information). 824

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#### 4.2 Seasonal networks for monthly indices

Overall, the fitted networks based on full year data show good agreement between 826 the two reanalyses. Some of the differences noted in the previous section may arise due 827 to confounding or spurious associations generated by, among other things, the omission 828 of relevant variables in the fit (i.e., failure of causal sufficiency, in the case that the mod-829 els are interpreted as being causal). An obvious possible factor is the seasonal variation 830 in relationships between nodes that arises as a result of seasonal changes in the back-831 ground flow and hence in the available pathways for propagation of disturbances (e.g., 832 Hoskins & Ambrizzi, 1993; Ambrizzi et al., 1995). To account for this seasonal depen-833 dence, one possibility would be to include a seasonal indicator as a node within the graph 834 itself. For simplicity, however, to investigate the impact of seasonal variation we consider 835

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the results of repeating the above analysis restricted to data in the three month winter season for each hemisphere. Note that, as lags of up to six months are still allowed, observations entering into these fits also include lagged values of the indices during the previous season. For brevity, we restrict our attention to only a subset of the major teleconnections within each hemisphere.

In Figure 9, the estimated posterior probabilities for the parent sets of the  $NAO^+$ 841 and NAO<sup>-</sup> indices during DJF are shown. As for the full year fits, in both reanalyses 842 there is strong evidence for an association between the AO and the two phases of the NAO, 843 with approximately equal posterior mass assigned to each edge in NNR1 and JRA-55. 844 Compared to the full year networks, the edge from  $PSA2_{t-1} \rightarrow NAO_t^+$  is no longer found 845 to have appreciable posterior mass ( $\hat{\pi} < 0.1$  in both reanalyses), consistent with this 846 edge arising as a result of the use of full-year data. In both NNR1 and JRA-55, a rela-847 tionship between the MJO, via the value of the RMM1 index at a lag of two months, with 848 the positive phase of the NAO is found in a large fraction of sampled models ( $\hat{\pi} \approx 0.67$ 849 in NNR1 and  $\hat{\pi} \approx 0.83$  in JRA-55). Interactions between the MJO and the NAO have 850 previously been reported in winter season observations (e.g., Lin et al., 2009) arising from 851 known dynamical mechanisms (e.g., Frederiksen & Frederiksen, 1993). The inferred pres-852 ence of this relation, with moderately high confidence, indicates that both products are 853 consistent in capturing this association, although it should be noted that monthly mean 854 data is being used here in contrast to the more usual daily or pentad data. In both NNR1 855 and JRA-55, additional edges are also found with somewhat lower posterior probabil-856 ities, including the edges  $NAO_{t-5}^+ \rightarrow NAO_t^-$ ,  $PSA1_{t-5} \rightarrow NAO_t^-$ , and  $PNA_{t-3} \rightarrow NAO_t^+$ 857 with estimated probabilities  $\hat{\pi} \approx 0.55, 0.51$ , and 0.54 in JRA-55, respectively, and  $\hat{\pi} \approx$ 858 0.49, 0.43, and 0.45 in NNR1. As these features are present in both reanalyses with rel-859 atively similar posterior weights, this suggests that there is some, if comparatively weak, 860 evidence for these associations from both products, and the two reanalyses appear to thus 861 be consistent. More notably, the feature  $DMI_{t-6} \rightarrow NAO_t^-$  is estimated to have a pos-862 terior probability of  $\hat{\pi} \approx 0.52$  based on the indices computed using NNR1 and HadISST 863 data, while when fitted to JRA-55 the same feature is assigned a posterior mass of only 864  $\hat{\pi} \approx 0.09$ , implying substantially weaker evidence is found for the presence of this edge 865 in the JRA-55 data. 866

For fits based on austral winter (June-July-August, JJA) data, similar results are found in that several apparently spurious associations cease to be present in a large frac-

-30-

tion of the sampled structures. In Figure 10, the estimated posterior probabilities for par-869 ent nodes of the PSA1 index during JJA are shown for the two reanalyses. Compared 870 to the previous, full-year fits in Figure 5, the set of edges with high posterior mass from 871 NNR1 data no longer includes long-range dependence on the positive phase of the NAO, 872 or on the PSA2 index. Instead, an edge  $\text{RMM1}_{t-5} \rightarrow \text{PSA1}_t$  is found with posterior prob-873 ability  $\hat{\pi} \approx 0.60$ , with the same feature being obtained from the JRA-55 data with  $\hat{\pi} \approx$ 874 0.77. The JRA-55 fit also contains an edge  $\text{RMM2}_{t-6} \rightarrow \text{PSA1}_t$  with estimated poste-875 rior probability  $\hat{\pi} \approx 0.94$ ; in NNR1, the corresponding feature is found to have  $\hat{\pi} \approx$ 876 0.42. Associations between the MJO and the PSA modes on intraseasonal time-scales 877 during winter have previously been noted (Mo & Paegle, 2001), although the fitted mod-878 els here assign greater posterior weight to dependence at longer lags. While the poste-879 rior probability for the presence of RMM1 as a predictor of the monthly mean PSA1 is 880 roughly consistent between NNR1 and JRA-55, the approximate factor of two difference 881 in the value of the sampled posterior probability for the RMM2 edge is more sizable and 882 may provide weak evidence of a difference between the two reanalyses in terms of this 883 relationship. A similar statement may be made for the  $MEI_{t-1} \rightarrow PSA1_t$  edge, for which 884 the fitted posterior probabilities are  $\hat{\pi} \approx 0.59$  in JRA-55 and  $\hat{\pi} \approx 0.22$  in NNR1. In-885 terestingly, approximately half ( $\hat{\pi} \approx 0.49$ ) of the sampled structures in NNR1 instead 886 contain an edge  $PNA_{t-1} \rightarrow PSA1_t$ , which may reflect the effects of a common depen-887 dence on tropical forcing. Overall, there are a larger number of features with high pos-888 terior mass in the JRA-55 fits. As usual, the precise underlying reasons for these differ-889 ences are not determined by the fits alone. Possible biases, such as a poorer represen-890 tation of the wintertime SH circulation in one product compared to the other, would re-891 quire further detailed follow-up. Our purpose here has been to highlight the use of Bayesian 892 structure learning as a tool to identify possible differences, and to estimate the level of 893 uncertainty associated with each. 894

# <sup>895</sup> 5 Summary

Probabilistic graphical models provide a natural and intuitive framework with which to describe the complicated interactions between climate processes. As a result, they are increasingly being applied for the purposes of studying potential causal relationships and for model evaluation (Vázquez-Patiño et al., 2020; Nowack et al., 2020). In this latter application, models are generally evaluated on the basis of structural comparisons be-

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tween graph structures inferred from observations and from model runs. Additionally,
the strength of associations may be compared by performing a second stage of model fitting, conditional on the inferred structure. As is, this constitutes a powerful lens for examining differences between models and observations.

That being said, existing approaches still have some important limitations when 905 used as tools for model evaluation. In particular, the most widely used strategy of first 906 learning a suitable structure using a constraint-based learning algorithm, followed op-907 tionally by fitting a regression model conditioned on this structure, does not lend itself 908 to easily estimating the level of confidence in the obtained model. While sensitivity anal-909 yses enable some determination of the robustness of particular features, in general the 910 sampling properties associated with this procedure are difficult to assess (Madigan & Raftery, 911 1994; Draper, 1995). This can present a challenge for using learned structures as tools 912 for model evaluation, where some measure of significance of observed differences is usu-913 ally desirable so as to assess whether they are due to model biases or sampling variabil-914 ity. 915

This limitation can in principle be overcome by employing a Bayesian approach to 916 structure learning. By learning a posterior distribution over possible structures, rather 917 than selecting a single graph, overall model uncertainty can be quantified and accounted 918 for. This can be particularly important where multiple different structures may all be 919 nearly equally well supported by the data, in which case selection of a single model may 920 overestimate the confidence warranted in particular features. Given a sample from the 921 model posterior distribution, by averaging over the set of possible models the posterior 922 credibility of given features may instead be estimated in the Bayesian approach to iden-923 tify edges that are well supported by the data. Subsequent estimation of the model pa-924 rameters conditional on a given structure is straightforward, and provides a basis for com-925 paring the magnitude of relationships between different processes. 926

The result of the sampling-based structure learning algorithms is a sample from the set of possible models, from which posterior probabilities for particular features can be derived. In this way, robust features for which there is high confidence may be identified, and the set of such edges may in turn form the focus of model comparisons. To illustrate this approach, we have applied an MCMC based approach to learn DBNs describing associations between teleconnections in two different reanalyses, with the goals

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of identifying the robust structural differences between the two products and to estab-

933 934

lish a set of baseline estimates for subsequent model evaluation studies.

In general, features in the networks derived from NNR1 and JRA-55 data that have 935 high estimated posterior probabilities agree reasonably well. While not surprising, as both 936 reanalyses attempt to provide an estimate of the same climate state, this provides some 937 reassurance that consistent results are obtained using a sampling algorithm, and that 938 the expected associations and dependence structures, such as long memory in oceanic 939 modes and close correspondence between the AO and NAO, are recovered with reason-940 able levels of certainty. Differences between the models estimated from the two reanal-941 yses are not only limited to edges with low posterior mass, however. In some cases, these 942 differences involve modes for which there are known biases in one reanalysis or the other; 943 the lack of evidence for a dependence of the PNA on the MJO in JRA-55 is one such ex-944 ample. In other cases, these features may arise as a result of the effects of omitted con-945 founding effects, such as seasonal cycles and other common drivers, that nevertheless dif-946 fer between the two reanalyses. Some evidence for this is found by considering networks 947 fitted from data restricted to only the winter season in each hemisphere. In these fits, 948 apparently spurious cross-equatorial dependence present in fits to year-round data are 949 no longer found to have strong evidence to support their presence. A greater number of 950 differences between the networks derived from the two reanalyses are found for edges that 951 have lower posterior probabilities, with this being particularly noticeable for the SH modes. 952 This may in part be due to greater differences in the representation of these modes, as 953 well as lower signal-to-noise ratios, at least outside of the austral winter. 954

It is important to note that in this study our aim has not been to perform a de-955 tailed evaluation of the differences between the two reanalysis products. Extensive char-956 acterization of the performance and biases of both NNR1 and JRA-55 has been done in 957 the past, and further investigations of the differences found here would require additional 958 detailed study of the involved processes in each product. Rather, our purpose has been 959 to demonstrate the applicability of the Bayesian approach to structure learning in de-960 riving graphical models suitable for use as tools for process-based model evaluation. We 961 argue that an important aspect of this application is the need to account for inevitable 962 model uncertainty in order to identify differences that are likely to be robust and hence 963 represent genuine model biases. Independent of the context of the analysis, this can be 964 naturally achieved in the Bayesian approach. By considering a relatively straightforward 965

initial application to reanalysis data, for which there exists (at least some) consensus on
 the interactions between modes, we have shown that a score-based sampling approach
 recovers the expected relationships, while also providing additional benefits over constraint based approaches in the form of estimates for the posterior distribution over models and
 features.

Given this, we have primarily focused on a single analysis of year-round monthly 971 mean data, or data within a single season. While sufficient to illustrate the relevant fea-972 tures of the results, in order to utilize this approach in the context of causal discovery 973 it would be necessary to further extend the analysis presented here. In addition to care-974 ful selection of the relevant variables, time periods, and temporal resolution, further con-975 sideration should be given to the form of the likelihood and priors used in defining the 976 model. In this work, the simplest case of a linear model with conjugate priors on the pa-977 rameters defining the conditional PDFs has been used, together with priors on the struc-978 tures to ensure structural modularity. No prior restriction has been enforced to ensure 979 stationarity of the resulting autoregressive model. Additionally, no attempt has been made 980 to incorporate pre-existing or expert knowledge into the definition of the chosen priors. 981 More complex forms for the conditional PDF, as well as the use of non-conjugate pri-982 ors and inclusion of additional constraints, may be directly handled using a generic re-983 versible jump MCMC instead of the more specialized MC<sup>3</sup> used here. 984

Whether or not the resulting model has analytic structure, care must be taken in 985 the design of the sampler and in assessing the approximate convergence of the simula-986 tion to the target posterior distribution. For the results presented here, the non-parametric 987 convergence diagnostics of Brooks et al. (2003) have been used to monitor (non-)convergence, 988 but in general assessing convergence for trans-dimensional MCMC remains challenging 989 (Sisson, 2005). In our case, trace plots and convergence diagnostics applied to individ-990 ual edge indicators suggested that (model-averaged) estimates derived from individual 991 chains were consistent and stable, but there was evidence for non-homogeneity in the pos-992 terior distribution over structures across chains based on  $10 \times 10^6$  samples. For this rea-993 son, we have avoided making definitive statements with respect to Bayes factors and other 994 model-specific quantities, recognizing that further sampling would be required for these 995 to be reliably determined. Poor mixing and multi-modal posterior distributions over mod-996 els may also be problematic, with the sampled chains remaining trapped in the vicin-997 ity of a single mode. Given the large model space considered here, this is a concern as 998

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ensuring the chains have adequately explored the full posterior distribution is unlikely 999 to be feasible, raising the possibility that the fits have converged to a local mode in the 1000 model space. For generating causal hypotheses, it may be more suitable to restrict at-1001 tention to smaller systems of variables. On the other hand, the relatively close agree-1002 ment between the fits for the two reanalyses shown here suggests that the simulations 1003 explore sufficiently similar regions of the model space for useful qualitative comparisons 1004 to still be performed. As in this paper only two datasets were compared, direct inspec-1005 tion of the individual fitted networks, amounting to visual inspection of the sampled pos-1006 terior distributions, was sufficient. In more extensive evaluation studies, this may be ex-1007 tended by the use of appropriate summary measures computed on the fits for different 1008 models. 1009

The DBN models obtained for the NNR1 and JRA-55 reanalyses form a set of ground 1010 truth results against which free-running models may be compared; we are currently per-1011 forming such a comparison over the historical period. However, the homogeneous mod-1012 els considered here cannot model one of the most notable features of the observed cli-1013 mate over this period, namely, the existence of secular trends in the behavior of partic-1014 ular modes. Key open questions remain as to whether there is evidence for accompany-1015 ing changes in the underlying interaction structures, and whether models indicate the 1016 possibility of regime shifts under future forcing scenarios. Addressing these questions in 1017 this framework will require the use of non-homogeneous network models in which either 1018 the model parameters or structure are allowed to vary. In a forthcoming article, we per-1019 form just such a comparison against the reanalysis-derived networks presented here for 1020 a subset of the CMIP5 ensemble, as well as relaxing the time-homogeneity assumptions 1021 discussed above to study the existence of regime transitions in projections. Assessing the 1022 evidence for, say, the existence of sudden structure changes versus slow variation in the 1023 underlying network parameters does not entail any additional conceptual changes, mak-1024 ing the approach presented here well-suited for investigating such questions. 1025

1026 Acknowledgments

The HadISST SST dataset is provided by the UK Met Office Hadley Centre as described in Rayner et al. (2003), and may be accessed at https://www.metoffice.gov .uk/hadobs/hadisst/ (last access: 29 April 2019). The NCEP/NCAR reanalysis output used is provided by the NOAA/OAR/ESRL PSL, Boulder, Colorado, USA, described

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in Kalnay et al. (1996), and may be accessed at https://psl.noaa.gov/data/reanalysis/
reanalysis.shtml (last access: 10 May 2019). The JRA-55 reanalysis output used is
made available through the JRA-55 project and may be accessed following the procedures and access conditions described in Kobayashi et al. (2015) and at https://jra
.kishou.go.jp/JRA-55/index\_en.html (last access: 12 April 2019).

Regridding of the reanalysis fields was performed using the Climate Data Oper-1036 ators software suite (Schulzweida, 2019), while the analysis code was implemented us-1037 ing the Python libraries NumPy (Oliphant, 2006; Van Der Walt et al., 2011), SciPy (Virtanen 1038 et al., 2020), pandas (Wes McKinney, 2010), scikit-learn (Pedregosa et al., 2011), and 1039 xarray (Hoyer & Hamman, 2017). Plots were generated using the Python package Mat-1040 plotlib (Hunter, 2007). All source code and the generated indices used to perform the 1041 analyses presented in this study may be found at https://doi.org/10.5281/zenodo 1042 .4331149. 1043

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### <sup>1048</sup> Appendix A MCMC sampling methods

In this appendix we summarize the MCMC methods used for structure learning. We employ the composite parameter space formulation described in detail in Godsill (2001). For the models considered in the main text, in the absence of same-time edges the networks are structurally modular and the parent set associated with each index can be inferred independently. However, in general a given structure  $G \in \mathcal{G}$  will describe the joint distribution of all of the indices simultaneously at a given time. To accommodate such models, we keep the notation relatively general.

The observed data takes the form of a time series of one or more indices  $D = \{y_1, \ldots, y_T\}$ . The allowed set of models  $\mathcal{G}$  to describe these data are taken to be specified a priori. Let  $\theta$  denote the collection of all such parameters across all of the possible models in  $\mathcal{G}$ . For example, for the set of all linear Gaussian models, the vector  $\theta$  would contain all of the possible regression coefficients  $\beta^i_{(j,\tau)}$  associated with the possible edges  $Y^j_{t-\tau} \to Y^i_t$  as well as the conditional precisions  $\tau^2_i$  and intercept parameters  $\beta^i_0$ . Any given model G will only use a small subset of all of the possible parameters. The subset of the parameters used by a structure G will be written  $\theta_{\mathcal{I}(G)}$ ; here  $\mathcal{I}(G)$  is an appropriate index set indicating which components of the complete parameter vector  $\boldsymbol{\theta}$  are required by G. The remaining parameters not used by G will be denoted by  $\theta_{-\mathcal{I}(G)}$ . A particular model for the indices  $\{Y^1, \ldots, Y^n\}$  is described by the pair  $(G, \boldsymbol{\theta}) \in \mathcal{G} \times \Theta$ , where  $\mathcal{G} \times \Theta$  is the composite model space (Godsill, 2001).

To infer the network structure and parameters given data D, we aim to sample from the posterior distribution

1070 
$$P(G, \theta|D) \propto P(D|G, \theta)P(G, \theta).$$

The likelihood  $P(D|G, \theta)$  is taken to depend only on the subset of parameters associated with G,

$$P(D|G, \boldsymbol{\theta}) \equiv P(D|G, \boldsymbol{\theta}_{\mathcal{I}(G)})$$

while the prior distribution  $P(G, \theta)$  is taken to be of the form

$$P(G, \boldsymbol{\theta}) = P(\boldsymbol{\theta}_{\mathcal{I}(G)}|G)P(\boldsymbol{\theta}_{-\mathcal{I}(G)}|\boldsymbol{\theta}_{\mathcal{I}(G)}, G)P(G).$$

The factor  $P(\boldsymbol{\theta}_{-\mathcal{I}(G)}|\boldsymbol{\theta}_{\mathcal{I}(G)}, G)$  corresponds to a set of proper pseudo-priors (Carlin & Chib, 1995) for the parameters not used by G, and may otherwise be chosen essentially freely. The samplers that we use correspond to Metropolis-Hastings schemes in the composite model space (Godsill, 2001) with a proposal density of the form

$$q(G', \boldsymbol{\theta}'; G, \boldsymbol{\theta}) = q_1(G'; G)q_2(\boldsymbol{\theta}'_{\mathcal{I}(G')}; \boldsymbol{\theta}_{\mathcal{I}(G)})P(\boldsymbol{\theta}'_{-\mathcal{I}(G')}|\boldsymbol{\theta}'_{\mathcal{I}(G')}, G'),$$
(A1)

for a move from  $(G, \theta)$  to  $(G', \theta')$  with corresponding acceptance probability

$$\alpha = \min\left\{1, \frac{q_1(G; G')}{q_1(G'; G)} \frac{q_2(\boldsymbol{\theta}_{\mathcal{I}(G)}; \boldsymbol{\theta}'_{\mathcal{I}(G')})}{q_2(\boldsymbol{\theta}'_{\mathcal{I}(G')}; \boldsymbol{\theta}_{\mathcal{I}(G)})} \frac{P(G', \boldsymbol{\theta}'_{\mathcal{I}(G')}|D)}{P(G, \boldsymbol{\theta}_{\mathcal{I}(G)}|D)}\right\}.$$
(A2)

When the class of models considered does not admit analytic evaluation of any of the required integrals, we make use of the simple reversible jump MCMC scheme given in Algorithm 1. At each iteration, either an update to the parameter associated with the structure *G* is chosen, with probability  $j_{\theta}(G, \boldsymbol{\theta}_{\mathcal{I}(G)})$ , or an update to the current structure is proposed. In the case of a parameter update, the model structure is left unchanged, G' = G, and a new set of parameter values  $\boldsymbol{\theta}'_{\mathcal{I}(G)}$  is drawn from a proposal density  $q_{\theta}(\boldsymbol{\theta}'_{\mathcal{I}(G)}; \boldsymbol{\theta}_{\mathcal{I}(G)})$ . The updated state  $(G, \boldsymbol{\theta}'_{\mathcal{I}(G)})$  is accepted with probability

$$\alpha = \min\left\{1, \frac{j_{\theta}(G, \boldsymbol{\theta}'_{\mathcal{I}(G)})}{j_{\theta}(G, \boldsymbol{\theta}_{\mathcal{I}(G)})} \frac{q_{\theta}(\boldsymbol{\theta}_{\mathcal{I}(G)}; \boldsymbol{\theta}'_{\mathcal{I}(G)})}{q_{\theta}(\boldsymbol{\theta}'_{\mathcal{I}(G)}; \boldsymbol{\theta}_{\mathcal{I}(G)})} \frac{P(G, \boldsymbol{\theta}'_{\mathcal{I}(G)}|D)}{P(G, \boldsymbol{\theta}_{\mathcal{I}(G)}|D)}\right\}.$$
(A3)

Note that, when the probability of a parameter update move is the same in both states, 1091 this is simply an ordinary Metropolis-Hastings update for a single model. If instead a 1092 structure update move is chosen, a new structure G' is proposed according to the pro-1093 posal distribution  $q_G(G'; G)$ . Any parameters that are common to both G' and G are 1094 held fixed at their current values, while any new parameters  $\theta'_{\mathcal{I}(G')\setminus\mathcal{I}(G)}$  are sampled from 1095 an additional proposal density  $\tilde{q}_{\theta}(\boldsymbol{\theta}'_{\mathcal{I}(G')\setminus\mathcal{I}(G)})$ ; if  $\mathcal{I}(G') \subset \mathcal{I}(G)$  we set  $\tilde{q}_{\theta} \rightarrow 1$ . Pa-1096 rameters that are either present in the initial structure but not in the proposed struc-1097 ture, or are not used by either, are left unchanged for simplicity. The acceptance ratio 1098 for this move is 1099

$$\alpha = \min\left\{1, \frac{j_G(G', \boldsymbol{\theta}'_{\mathcal{I}(G')})}{j_G(G, \boldsymbol{\theta}_{\mathcal{I}(G)})} \frac{q_G(G; G')}{q_G(G'; G)} \frac{\tilde{q}_{\theta}(\boldsymbol{\theta}_{\mathcal{I}(G) \setminus \mathcal{I}(G')})}{\tilde{q}_{\theta}(\boldsymbol{\theta}'_{\mathcal{I}(G') \setminus \mathcal{I}(G)})} \frac{P(G', \boldsymbol{\theta}'_{\mathcal{I}(G')}|D)}{P(G, \boldsymbol{\theta}_{\mathcal{I}(G)}|D)}\right\}.$$
 (A4)

This structure update move is just a particular case of the general reversible jump move (Green, 1995) with a unit Jacobian. For more general mappings from the current to proposed parameters a non-trivial Jacobian factor would remain following the change of variables in the proposal density Eq. (A1).

When the models considered allow for the conditional posterior distribution  $P(\theta_{\mathcal{A}(G)}|\theta_{\mathcal{I}(G)\setminus\mathcal{A}(G)},G,D)$ 1105 of some subset of the parameters  $\theta_{\mathcal{A}(G)}$ ,  $\mathcal{A}(G) \subseteq \mathcal{I}(G)$ , to be evaluated given G and 1106 any remaining parameters  $\theta_{\mathcal{I}(G)\setminus\mathcal{A}(G)}$ , we adopt a conditional Metropolis-Hastings scheme 1107 that takes better advantage of this structure. For simplicity, we assume that the set of 1108 parameters  $\theta_{\mathcal{I}(G)\setminus\mathcal{A}(G)}$ , if not empty, is shared across all of the possible structures. Up-1109 dates to the structure are proposed as before under the proposal  $q_G(G'; G)$ . The shared 1110 parameters  $\theta_{\mathcal{I}(G)\setminus\mathcal{A}(G)}$  are kept at their previous values, while the remaining parame-1111 ters are drawn from the exactly known posterior distribution  $P(\theta'_{\mathcal{A}(G')}|\theta'_{\mathcal{I}(G')\setminus\mathcal{A}(G')},G',D)$ . 1112 Under this proposal, the acceptance ratio simplifies to 1113

$$\alpha = \min\left\{1, \frac{q_G(G;G')}{q_G(G';G)} \frac{P(G'|\boldsymbol{\theta}'_{\mathcal{I}(G')\setminus\mathcal{A}(G')}, D)}{P(G|\boldsymbol{\theta}_{\mathcal{I}(G)\setminus\mathcal{A}(G)}, D)}\right\}.$$
(A5)

The parameters associated with the current structure can be updated via a standard Metropolis-Hastings or Gibbs step, as in the parameter update move for the simple reversible jump scheme. Parameter and structure updates can either be proposed randomly or performed

- in a fixed order. For the models presented in Section 4, the conditional posterior distri-
- bution for all of the parameters can be evaluated, i.e., the set  $\theta_{\mathcal{I}(G)\setminus\mathcal{A}(G)}$  is empty. In
- this case, the scheme reduces to the  $MC^3$  scheme of Madigan et al. (1995), with the de-
- pendence on the graph parameters dropping out entirely. The acceptance ratio for a struc-

## Algorithm 1 Simple RJMCMC sampler

**Require:** initial state  $x^{(1)} \equiv (G^{(1)}, \theta^{(1)})$ , observed data D, chain length 2S 1: for s = 2, ..., 2S do Draw  $u_1, u_2 \sim \text{Uniform}(0, 1)$ 2: if  $u_1 < j_{\theta}(x^{(s-1)})$  then 3: Set  $G' \leftarrow G^{(s-1)}$ 4: Draw new parameters  $\boldsymbol{\theta}'_{\mathcal{I}(G')}$  from  $q_{\boldsymbol{\theta}}(\boldsymbol{\theta}'_{\mathcal{I}(G')}; \boldsymbol{\theta}^{(s-1)}_{\mathcal{I}(G^{(s-1)})})$ 5:Calculate  $\alpha$  according to Eq. (A3) 6: 7: else Draw new parent set G' from  $q_G(G'; G^{(s-1)})$ 8: Draw  $\boldsymbol{\theta}'_{\mathcal{I}(G') \setminus \mathcal{I}(G^{(s-1)})}$  from  $\tilde{q}_{\boldsymbol{\theta}}(\boldsymbol{\theta}'_{\mathcal{I}(G') \setminus \mathcal{I}(G^{(s-1)})})$ 9: Calculate  $\alpha$  according to Eq. (A4) 10: if  $u_2 < \alpha$  then 11:  $x^{(s)} \leftarrow (G', \theta')$ 12:else 13: $x^{(s)} \leftarrow (G^{(s-1)}, \boldsymbol{\theta}^{(s-1)})$ 14: 15: Discard first S samples as warm-up 16: **return**  $\{x^{(S+1)}, \ldots, x^{(2S)}\}$ 

ture drawn according to  $q_G(G'; G)$  is in this case

1123 
$$\alpha = \min\left\{1, \frac{q_G(G;G')}{q_G(G';G)} \frac{P(D|G')}{P(D|G)} \frac{P(G')}{P(G)}\right\},\tag{A6}$$

where the likelihood can be written in terms of the local marginal likelihoods, as in Eq. (8). Particular choices of the proposal density and structure priors are described in Section 3.3. We summarize the resulting sampling scheme for the structures in Algorithm 2. For each sampling method, we run multiple chains, discarding the first half of each sample as burnin. For the MC<sup>3</sup> sampler, approximate convergence of the chains to the target distribution is assessed using the  $\chi^2$  and Kolmogorov-Smirnov tests proposed in Brooks et al. (2003).

# **Algorithm 2** $MC^3$ sampler

Aige	Sampler
Req	uire: initial state $G^{(1)}$ , observed data $D$ , chain length $2S$
1: <b>f</b>	or $s = 2, \ldots, 2S$ do
2:	Draw $u \sim \text{Uniform}(0, 1)$
3:	Draw new parent set $G'$ from $q_G(G'; G^{(s-1)})$
4:	Calculate $\alpha$ according to Eq. (A6)
5:	if $u < \alpha$ then
6:	$G^{(s)} \leftarrow G'$
7:	else
8:	$G^{(s)} \leftarrow G^{(s-1)}$
9: I	Discard first $S$ samples as warm-up
10: <b>r</b>	eturn $\{G^{(S+1)}, \dots, G^{(2S)}\}$

## Appendix B Expressions for marginal likelihoods

In this appendix we state the closed-form expressions for the prior and posterior densities for the parameters of the node conditional distributions, and the resulting marginal likelihoods or local scores, for the models used in the main text.

The linear Gaussian model given in Section 2 reads

$$\begin{aligned} \tau_i^2 \sim \operatorname{Gamma}(a_{\tau}, b_{\tau}), \\ \beta_0^i | \tau_i^2, \nu_i^2 \sim N\left(0, \frac{\nu_i^2}{\tau_i^2}\right), \\ \beta_{(k_j, \tau_j)}^i | \tau_i^2, \nu_i^2, \operatorname{pa}_G(Y_t^i) \sim N\left(0, \frac{\nu_i^2}{\tau_i^2}\right), \quad j = 1, \dots, |\operatorname{pa}_G(Y_t^i)|, \\ Y_t^i | \beta_0^i, \beta_{(k_j, \tau_j)}^i, \tau_i^2, \operatorname{pa}_G(Y_T^i) \sim N\left(\beta_0^i + \sum_{j=1}^{|\operatorname{pa}_G(Y_t^i)|} \beta_{(k_j, \tau_j)}^i Y_{t-\tau_j}^{k_j}, \frac{1}{\tau_i^2}\right), \end{aligned}$$

1135 For the prior on the conditional precision  $\tau_i^2$ , we adopt the convention

1136 
$$P(\tau_i^2 | a_{\tau}, b_{\tau}) = \frac{1}{\Gamma(a_{\tau})} \frac{(\tau_i^2)^{a_{\tau}-1}}{b_{\tau}^{a_{\tau}}} \exp\left(-\frac{\tau_i^2}{b_{\tau}}\right).$$

With this convention, the unconditional prior distribution for a given coefficient  $\beta$  is a generalized *t*-distribution with scale  $\hat{\sigma}^2 = \nu^2/(a_\tau b_\tau)$  and  $2a_\tau$  degrees of freedom,  $\beta \sim T_1(0, \nu^2/(a_\tau b_\tau), 2a_\tau)$ . For the models shown in Section 4,  $a_\tau$ ,  $b_\tau$ , and  $\nu_i^2$  are taken as fixed hyperparameters. In practice, the signal-to-noise  $\nu_i^2$  may be poorly known, in which case it is possible to also sample it from a conjugate inverse gamma prior using the sampling schemes for models with partial analytic structure discussed in Appendix A. For a given parent set, we write the vector of predictor variables as  $(p_i = |pa_G(Y_t^i)|)$ 

1144 
$$oldsymbol{x}_t^{iT} = (1, y_{t- au_1}^{k_1}, \dots, y_{t- au_{p_i}}^{k_{p_i}}),$$

at each time t = 1, ..., T and introduce the  $T \times (p_i + 1)$  design matrix

1146 
$$X_i = \begin{pmatrix} \boldsymbol{x}_1^{iT} \\ \vdots \\ \boldsymbol{x}_T^{iT} \end{pmatrix}$$

The likelihood for the observed values of the index  $Y^i$  then takes the simple form of a product of normal densities, and the local marginal likelihood  $\Psi_i(D;G)$  can be evaluated using standard conjugacy results. The marginal joint distribution for the observed index values  $\boldsymbol{y}_i^T = (y_1^i, \dots, y_T^i)$  under this model is a multivariate *t*-distribution, giving

$$\Psi_i(D;G) = \frac{\Gamma\left(\frac{T+2a_\tau}{2}\right)}{\Gamma\left(\frac{2a_\tau}{2}\right)\pi^{T/2}(2a_\tau)^{T/2}} \left(\det \Sigma_i\right)^{-1/2} \left(1 + \frac{1}{2a_\tau} \boldsymbol{y}_i^T \Sigma_i^{-1} \boldsymbol{y}_i\right)^{-\frac{T+2a_\tau}{2}}, \quad (B1)$$

1153 where

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1154

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$$\Sigma_i = \frac{1}{a_\tau b_\tau} \left( I_{T \times T} + \nu_i^2 X_i X_i^T \right),\tag{B2}$$

and 
$$I_{T \times T}$$
 is the  $T \times T$  identity matrix. For a given parent set, the regression coefficients  
and precision can be sampled from the conditional posterior distributions

$$\tau_i^2 | D, G, \nu_i^2 \sim \text{Gamma}\left(\frac{T + 2a_\tau}{2}, \frac{b_\tau}{1 + \frac{1}{2a_\tau}\boldsymbol{y}_i^T \boldsymbol{\Sigma}_i^{-1} \boldsymbol{y}_i}\right), \tag{B3}$$

$$\boldsymbol{\beta}_{i}|\tau_{i}^{2},\nu_{i}^{2},D,G\sim N\left(\tau_{i}^{2}\Sigma_{\beta_{i}}X_{i}^{T}\boldsymbol{y}_{i},\Sigma_{\beta_{i}}\right),$$
(B4)

where the posterior covariance matrix for the coefficients  $\beta_i^T \equiv (\beta_0^i, \beta_{(k_1, \tau_1)}^i, \dots, \beta_{(k_{p_i}, \tau_{p_i})}^i)$ is given by

$$\Sigma_{\beta_i} = \frac{\nu_i^2}{\tau_i^2} \left( I_{(p_i+1)\times(p_i+1)} + \nu_i^2 X_i^T X_i \right)^{-1}.$$

#### 1163 **References**

- Abramson, B., Brown, J., Edwards, W., Murphy, A., & Winkler, R. L. (1996). Hailfinder: A Bayesian system for forecasting severe weather. *International Journal*of Forecasting, 12(1), 57 71. doi: 10.1016/0169-2070(95)00664-8
- Ambaum, M. H. P., Hoskins, B. J., & Stephenson, D. B. (2001). Arctic Oscillation
   or North Atlantic Oscillation? Journal of Climate, 14(16), 3495-3507. doi: 10
   .1175/1520-0442(2001)014(3495:AOONAO)2.0.CO;2

1170	Ambrizzi, T., Hoskins, B. J., & Hsu, HH. (1995). Rossby Wave Propagation
1171	and Teleconnection Patterns in the Austral Winter. Journal of the Atmo-
1172	spheric Sciences, $52(21)$ , $3661-3672$ . doi: $10.1175/1520-0469(1995)052(3661:$
1173	$RWPATP \rangle 2.0.CO; 2$
1174	Arnold, A., Liu, Y., & Abe, N. (2007). Temporal Causal Modeling with Graph-
1175	ical Granger Methods. In Proceedings of the 13th ACM SIGKDD Inter-
1176	national Conference on Knowledge Discovery and Data Mining (p. 66–75).
1177	New York, NY, USA: Association for Computing Machinery. doi: $10.1145/$
1178	1281192.1281203
1179	Barnston, A. G., & Livezey, R. E. (1987). Classification, Seasonality and Persis-
1180	tence of Low-Frequency Atmospheric Circulation Patterns. Monthly Weather
1181	$Review,\ 115(6),\ 1083\text{-}1126. \text{doi:}\ \ 10.1175/1520\text{-}0493(1987)115\langle 1083\text{:}\text{CSAPOL}\rangle 2$
1182	.0.CO;2
1183	Bello, G. A., Angus, M., Pedemane, N., Harlalka, J. K., Semazzi, F. H. M., Kumar,
1184	V., & Samatova, N. F. (2015). Response-Guided Community Detection: Ap-
1185	plication to Climate Index Discovery. In Machine Learning and Knowledge
1186	Discovery in Databases (pp. 736–751). Cham: Springer International Publish-
1187	ing. doi: 10.1007/978-3-319-23525-7_45
1188	Berezin, Y., Gozolchiani, A., Guez, O., & Havlin, S. (2012). Stability of Climate
1189	Networks with Time. Scientific Reports, $2(1)$ , 666. doi: 10.1038/srep00666
1190	Bjerknes, J. (1969). Atmospheric Teleconnections from the Equatorial Pacific.
1191	Monthly Weather Review, 97(3), 163-172. doi: 10.1175/1520-0493(1969)
1192	$097\langle 0163: ATFTEP \rangle 2.3. CO; 2$
1193	Boneh, T., Weymouth, G. T., Newham, P., Potts, R., Bally, J., Nicholson, A. E.,
1194	& Korb, K. B. (2015). Fog Forecasting for Melbourne Airport Using a
1195	Bayesian Decision Network. Weather and Forecasting, $30(5)$ , 1218-1233.
1196	doi: 10.1175/WAF-D-15-0005.1
1197	Bromwich, D. H., & Fogt, R. L. (2004). Strong Trends in the Skill of the ERA-40
1198	and NCEP–NCAR Reanalyses in the High and Midlatitudes of the South-
1199	ern Hemisphere, 1958–2001. Journal of Climate, 17(23), 4603-4619. doi:
1200	10.1175/3241.1
1201	Bromwich, D. H., Fogt, R. L., Hodges, K. I., & Walsh, J. E. (2007). A tropospheric
1202	assessment of the ERA-40, NCEP, and JRA-25 global reanalyses in the po-

1203	lar regions. Journal of Geophysical Research: Atmospheres, 112(D10). doi:
1204	10.1029/2006 JD007859
1205	Brooks, S., Giudici, P., & Philippe, A. (2003). Nonparametric Convergence As-
1206	sessment for MCMC Model Selection. Journal of Computational and Graphical
1207	Statistics, $12(1)$ , 1-22. doi: 10.1198/1061860031347
1208	Buntine, W. (1991). Theory Refinement on Bayesian Networks. In
1209	B. D. D'Ambrosio, P. Smets, & P. P. Bonissone (Eds.), Uncertainty pro-
1210	ceedings 1991 (p. 52 - 60). San Francisco (CA): Morgan Kaufmann. doi:
1211	https://doi.org/10.1016/B978-1-55860-203-8.50010-3
1212	Carlin, B. P., & Chib, S. (1995). Bayesian Model Choice Via Markov Chain Monte
1213	Carlo Methods. Journal of the Royal Statistical Society: Series B (Methodolog-
1214	ical), 57(3), 473-484. doi: 10.1111/j.2517-6161.1995.tb02042.x
1215	Catenacci, M., & Giupponi, C. (2009). Potentials of Bayesian Networks to Deal
1216	with Uncertainty in Climate Change Adaptation Policies (SSRN Scholarly Pa-
1217	per No. ID 1526605). Rochester, NY: Social Science Research Network. doi:
1218	10.2139/ssrn.1526605
1219	Catenacci, M., & Giupponi, C. (2013). Integrated assessment of sea-level rise adap-
1220	tation strategies using a Bayesian decision network approach. Environmental
1221	Modelling and Software, 44, 87 - 100. doi: 10.1016/j.envsoft.2012.10.010
1222	Colombo, D., & Maathuis, M. H. (2014). Order-Independent Constraint-Based
1223	Causal Structure Learning. Journal of Machine Learning Research, 15(116),
1224	3921-3962.
1225	Cooper, G. F., & Herskovits, E. (1992). A Bayesian method for the induction of
1226	probabilistic networks from data. Machine learning, $9(4)$ , 309–347.
1227	Dawid, A. P., & Lauritzen, S. L. (1993). Hyper Markov Laws in the Statistical Anal-
1228	ysis of Decomposable Graphical Models. The Annals of Statistics, 21(3), 1272–
1229	1317.
1230	Dechter, R. (1999). Bucket elimination: A unifying framework for reasoning. Arti-
1231	ficial Intelligence, 113(1), 41 - 85. doi: https://doi.org/10.1016/S0004-3702(99)
1232	00059-4
1233	Deng, Y., & Ebert-Uphoff, I. (2014). Weakening of atmospheric information flow
1234	in a warming climate in the community climate system model. Geophysical Re-
1235	search Letters, 41(1), 193-200. doi: 10.1002/2013GL058646

1236	Deser, C. (2000). On the teleconnectivity of the "Arctic Oscillation". Geophysical
1237	Research Letters, 27(6), 779-782. doi: https://doi.org/10.1029/1999GL010945
1238	Di Capua, G., Runge, J., Donner, R. V., van den Hurk, B., Turner, A. G., Vel-
1239	lore, R., Coumou, D. (2020). Dominant patterns of interaction between
1240	the tropics and mid-latitudes in boreal summer: causal relationships and the
1241	role of timescales. Weather and Climate Dynamics, $1(2)$ , 519–539. doi:
1242	10.5194/wcd-1-519-2020
1243	Donges, J. F., Schultz, H. C. H., Marwan, N., Zou, Y., & Kurths, J. (2011). Investi-
1244	gating the topology of interacting networks. The European Physical Journal B,
1245	84(4),635651.doi: 10.1140/epjb/e2011-10795-8
1246	Donges, J. F., Zou, Y., Marwan, N., & Kurths, J. (2009a). The backbone of the cli-
1247	mate network. EPL (Europhysics Letters), 87(4), 48007. doi: 10.1209/0295
1248	-5075/87/48007
1249	Donges, J. F., Zou, Y., Marwan, N., & Kurths, J. (2009b). Complex networks in
1250	climate dynamics. The European Physical Journal Special Topics, 174(1), 157–
1251	179. doi: $10.1140/epjst/e2009-01098-2$
1252	Draper, D. (1995). Assessment and Propagation of Model Uncertainty. Journal of
1253	the Royal Statistical Society: Series B (Methodological), 57(1), 45-70. doi: 10
1254	.1111/j.2517-6161.1995.tb02015.x
1255	Ebert-Uphoff, I., & Deng, Y. (2012a). Causal Discovery for Climate Research Using
1256	Graphical Models. Journal of Climate, 25(17), 5648-5665. doi: 10.1175/JCLI
1257	-D-11-00387.1
1258	Ebert-Uphoff, I., & Deng, Y. (2012b). A new type of climate network based on prob-
1259	abilistic graphical models: Results of boreal winter versus summer. Geophysical
1260	Research Letters, $39(19)$ . doi: 10.1029/2012GL053269
1261	Ebert-Uphoff, I., & Deng, Y. (2017). Causal discovery in the geosciences—using syn-
1262	thetic data to learn how to interpret results. Computers and Geosciences, 99,
1263	50 - 60. doi: 10.1016/j.cageo.2016.10.008
1264	Eichler, M. (2012, Jun 01). Graphical modelling of multivariate time series. Probabil-
1265	ity Theory and Related Fields, 153(1), 233–268. doi: 10.1007/s00440-011-0345
1266	-8
1267	Eyring, V., Cox, P. M., Flato, G. M., Gleckler, P. J., Abramowitz, G., Caldwell, P.,
1268	Williamson, M. S. (2019). Taking climate model evaluation to the next

1269	level. Nature Climate Change, $9(2)$ , 102–110. doi: 10.1038/s41558-018-0355-y
1270	Falasca, F., Bracco, A., Nenes, A., & Fountalis, I. (2019). Dimensionality Reduc-
1271	tion and Network Inference for Climate Data Using $\delta\text{-MAPS:}$ Application to
1272	the CESM Large Ensemble Sea Surface Temperature. Journal of Advances in
1273	Modeling Earth Systems, 11(6), 1479-1515. doi: 10.1029/2019MS001654
1274	Fountalis, I., Dovrolis, C., Bracco, A., Dilkina, B., & Keilholz, S. (2018). $\delta$ -
1275	MAPS: from spatio-temporal data to a weighted and lagged network be-
1276	tween functional domains. Applied Network Science, $3(1)$ , 21. doi:
1277	10.1007/s41109-018-0078-z
1278	Franzke, C., Feldstein, S. B., & Lee, S. (2011). Synoptic analysis of the Pa-
1279	cific–North American teleconnection pattern. Quarterly Journal of the Royal
1280	Meteorological Society, 137(655), 329-346. doi: https://doi.org/10.1002/
1281	qj.768
1282	Frederiksen, J. S., & Frederiksen, C. S. (1993). Monsoon Disturbances, Intrasea-
1283	sonal Oscillations, Teleconnection Patterns, Blocking, and Storm Tracks of the
1284	Global Atmosphere during January 1979: Linear Theory. Journal of the Atmo-
1285	$spheric \ Sciences, \ 50(10), \ 1349-1372. \qquad \text{doi:} \ 10.1175/1520-0469(1993)050\langle 1349:$
1286	$MDIOTP \rangle 2.0.CO;2$
1287	Friedman, N., & Goldszmidt, M. (1996). Discretizing Continuous Attributes While
1288	Learning Bayesian Networks. In ICML 1996.
1289	Friedman, N., Goldszmidt, M., & Wyner, A. (1999). Data Analysis with Bayesian
1290	Networks: A Bootstrap Approach. In Proceedings of the Fifteenth Conference
1291	on Uncertainty in Artificial Intelligence (p. 196–205). San Francisco, CA,
1292	USA: Morgan Kaufmann Publishers Inc.
1293	Friedman, N., & Koller, D. (2003). Being Bayesian About Network Structure. A
1294	Bayesian Approach to Structure Discovery in Bayesian Networks. Machine
1295	Learning, $50(1)$ , 95–125. doi: 10.1023/A:1020249912095
1296	Friedman, N., Murphy, K., & Russell, S. (1998). Learning the Structure of Dynamic
1297	Probabilistic Networks. In Proceedings of the Fourteenth Conference on Uncer-
1298	tainty in Artificial Intelligence (p. 139–147). San Francisco, CA, USA: Morgan
1299	Kaufmann Publishers Inc.
1300	Geiger, D., & Heckerman, D. (1994). Learning gaussian networks. In Proceedings
1301	of the Tenth International Conference on Uncertainty in Artificial Intelligence

1302	(pp. 235–243).
1303	Ghil, M., & Lucarini, V. (2020). The physics of climate variability and climate
1304	change. Rev. Mod. Phys., 92, 035002. doi: 10.1103/RevModPhys.92.035002
1305	Gillett, N. P., Kell, T. D., & Jones, P. D. (2006). Regional climate impacts of the
1306	Southern Annular Mode. Geophysical Research Letters, $33(23)$ , L23704. doi:
1307	10.1029/2006 GL027721
1308	Goddard, L., Mason, S., Zebiak, S., Ropelewski, C., Basher, R., & Cane, M. (2001).
1309	Current approaches to seasonal to interannual climate predictions. Interna-
1310	tional Journal of Climatology, 21(9), 1111-1152. doi: 10.1002/joc.636
1311	Godsill, S. J. (2001). On the Relationship Between Markov chain Monte Carlo
1312	Methods for Model Uncertainty. Journal of Computational and Graphical
1313	Statistics, $10(2)$ , 230-248. doi: 10.1198/10618600152627924
1314	Gozolchiani, A., Havlin, S., & Yamasaki, K. (2011). Emergence of El Niño as an Au-
1315	tonomous Component in the Climate Network. Phys. Rev. Lett., 107, 148501.
1316	doi: 10.1103/PhysRevLett.107.148501
1317	Gozolchiani, A., Yamasaki, K., Gazit, O., & Havlin, S. (2008). Pattern of climate
1318	network blinking links follows El Niño events. EPL (Europhysics Letters),
1319	83(2), 28005. doi: 10.1209/0295-5075/83/28005
1320	Granger, C. W. J. (1969). Investigating Causal Relations by Econometric Models
1321	and Cross-spectral Methods. $Econometrica, 37(3), 424-438.$ doi: 10.2307/
1322	1912791
1323	Greatbatch, R. J., & Rong, Pp. (2006). Discrepancies between Different Northern
1324	Hemisphere Summer Atmospheric Data Products. Journal of Climate, $19(7)$ ,
1325	1261-1273. doi: 10.1175/JCLI3643.1
1326	Green, P. J. (1995). Reversible jump Markov chain Monte Carlo computation and
1327	Bayesian model determination. $Biometrika, 82(4), 711-732.$ doi: 10.1093/
1328	biomet/82.4.711
1329	Greenland, S., & Brumback, B. (2002). An overview of relations among causal mod-
1330	elling methods. International Journal of Epidemiology, $31(5)$ , 1030-1037. doi:
1331	10.1093/ije/31.5.1030
1332	Grzegorczyk, M., & Husmeier, D. (2011). Non-homogeneous dynamic Bayesian net-
1333	works for continuous data. Machine Learning, $83(3)$ , $355-419$ . doi: $10.1007/$
1334	s10994-010-5230-7

Guez, O., Gozolchiani, A., Berezin, Y., Brenner, S., & Havlin, S. (2012). Climate 1335 network structure evolves with North Atlantic Oscillation phases. EPL (Euro-1336 physics Letters), 98(3), 38006. doi: 10.1209/0295-5075/98/38006 1337 Hannachi, A., Straus, D. M., Franzke, C. L. E., Corti, S., & Woollings, T. (2017).1338 Low-frequency nonlinearity and regime behavior in the Northern Hemisphere 1339 Reviews of Geophysics, 55(1), 199-234. extratropical atmosphere. doi: 1340 10.1002/2015RG000509 1341 Harada, Y., Kamahori, H., Kobayashi, C., Endo, H., Kobayashi, S., Ota, Y., ... 1342 Takahashi, K. (2016). The JRA-55 Reanalysis: Representation of Atmospheric 1343 Circulation and Climate Variability. Journal of the Meteorological Society of 1344 Japan. Ser. II, 94(3), 269-302. doi: 10.2151/jmsj.2016-015 1345 Heckerman, D., Geiger, D., & Chickering, D. M. (1995).Learning Bayesian net-1346 works: The combination of knowledge and statistical data. Machine learning, 1347 20(3), 197-243.1348 Hertzog, A., Basdevant, C., & Vial, F. (2006).An Assessment of ECMWF and 1349 NCEP-NCAR Reanalyses in the Southern Hemisphere at the End of the Pre-1350 satellite Era: Results from the EOLE Experiment (1971–72). Monthly Weather 1351 Review, 134(11), 3367-3383. doi: 10.1175/MWR3256.1 1352 Hines, K. M., Bromwich, D. H., & Marshall, G. J. (2000).Artificial Sur-1353 face Pressure Trends in the NCEP-NCAR Reanalysis over the South-1354 ern Ocean and Antarctica. Journal of Climate, 13(22), 3940-3952. doi: 1355 10.1175/1520-0442(2000)013(3940:ASPTIT)2.0.CO;2 1356 Hlinka, J., Hartman, D., Vejmelka, M., Runge, J., Marwan, N., Kurths, J., & 1357 Paluš, M. (2013).Reliability of Inference of Directed Climate Networks 1358 Using Conditional Mutual Information. Entropy, 15(6), 2023–2045. doi: 1359 10.3390/e15062023 1360 Horel, J. D., & Wallace, J. M. (1981). Planetary-Scale Atmospheric Phenomena As-1361 sociated with the Southern Oscillation. Monthly Weather Review, 109(4), 813-1362 829. doi: 10.1175/1520-0493(1981)109(0813:PSAPAW)2.0.CO;2 1363 Horenko, I., Gerber, S., O'Kane, T. J., Risbey, J. S., & Monselesan, D. P. (2017).1364 On inference and validation of causality relations in climate teleconnections. 1365 In C. L. E. Franzke & T. J. O'Kane (Eds.), Nonlinear and stochastic climate 1366 dynamics (pp. 184–208). Cambridge University Press. 1367

1368	Hoskins, B. J., & Ambrizzi, T. (1993). Rossby Wave Propagation on a Realistic Lon-
1369	gitudinally Varying Flow. Journal of the Atmospheric Sciences, $50(12)$ , 1661-
1370	1671. doi: 10.1175/1520-0469(1993)050 (1661:RWPOAR)2.0.CO;2
1371	Hoyer, S., & Hamman, J. (2017). xarray: N-D labeled arrays and datasets in
1372	Python. Journal of Open Research Software, 5(1). doi: 10.5334/jors.148
1373	Hunter, J. D. (2007). Matplotlib: A 2d graphics environment. Computing in Science
1374	& Engineering, $9(3)$ , 90–95. doi: 10.1109/MCSE.2007.55
1375	Hurrell, J. W., Kushnir, Y., Ottersen, G., & Visbeck, M. (Eds.). (2003). The
1376	North Atlantic Oscillation: Climatic Significance and Environmental Impact
1377	(Vol. 134). American Geophysical Union.
1378	Jordan, M. I. (2004). Graphical models. Statist. Sci., $19(1)$ , 140–155. doi: 10.1214/
1379	08834230400000026
1380	Kaiser, H. F. (1958). The varimax criterion for analytic rotation in factor analysis.
1381	Psychometrika, 23(3), 187–200. doi: 10.1007/BF02289233
1382	Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L.,
1383	Joseph, D. (1996). The NCEP/NCAR 40-Year Reanalysis Project.
1384	Bulletin of the American Meteorological Society, 77(3), 437-472. doi:
1385	$10.1175/1520\text{-}0477(1996)077\langle0437\text{:}\mathrm{TNYRP}\rangle2.0.\mathrm{CO}\text{;}2$
1386	Kistler, R., Kalnay, E., Collins, W., Saha, S., White, G., Woollen, J., Fiorino,
1387	M. (2001). The NCEP–NCAR 50-Year Reanalysis: Monthly Means CD-ROM
1388	and Documentation. Bulletin of the American Meteorological Society, $82(2)$ ,
1389	247-268. doi: 10.1175/1520-0477(2001)082 (0247:TNNYRM>2.3.CO;2
1390	Kjærulff, U. (1995). dhugin: A computational system for dynamic time-sliced
1391	bayesian networks. International journal of forecasting, $11(1)$ , 89–111.
1392	Kobayashi, S., Ota, Y., Harada, Y., Ebita, A., Moriya, M., Onoda, H., Taka-
1393	hashi, K. $(2015)$ . The JRA-55 Reanalysis: General Specifications and Basic
1394	Characteristics. Journal of the Meteorological Society of Japan. Ser. II, $93(1)$ ,
1395	5-48. doi: 10.2151/jmsj.2015-001
1396	Koller, D., & Friedman, N. (2009). Probabilistic Graphical Models. Cambridge: The
1397	MIT Press.
1398	Kretschmer, M., Coumou, D., Donges, J. F., & Runge, J. (2016). Using
1399	Causal Effect Networks to Analyze Different Arctic Drivers of Midlati-
1400	tude Winter Circulation. Journal of Climate, $29(11)$ , 4069-4081. doi:

1401	10.1175/JCLI-D-15-0654.1
1402	Kretschmer, M., Runge, J., & Coumou, D. (2017). Early prediction of extreme
1403	stratospheric polar vortex states based on causal precursors. Geophysical Re-
1404	search Letters, $44(16)$ , 8592-8600. doi: 10.1002/2017GL074696
1405	Lau, KM., & Phillips, T. J. (1986). Coherent Fluctuations of Extratropical Geopo-
1406	tential Height and Tropical Convection in Intraseasonal Time Scales. Journal
1407	of the Atmospheric Sciences, $43(11)$ , 1164-1181. doi: $10.1175/1520-0469(1986)$
1408	$043\langle 1164: CFOFGH \rangle 2.0.CO; 2$
1409	Lau, KM., Sheu, PJ., & Kang, IS. (1994). Multiscale Low-Frequency Circulation
1410	Modes in the Global Atmosphere. Journal of the Atmospheric Sciences, $51(9)$ ,
1411	1169-1193. doi: 10.1175/1520-0469(1994)051(1169:MLFCMI>2.0.CO;2
1412	Leathers, D. J., Yarnal, B., & Palecki, M. A. (1991). The Pacific/North American
1413	Teleconnection Pattern and United States Climate. Part I: Regional Temper-
1414	ature and Precipitation Associations. Journal of Climate, $4(5)$ , 517-528. doi:
1415	$10.1175/1520\text{-}0442(1991)004\langle 0517\text{:}\mathrm{TPATPA}\rangle 2.0.\mathrm{CO}\text{;}2$
1416	Lèbre, S. (2009). Inferring Dynamic Genetic Networks with Low Order Independen-
1417	cies. Statistical Applications in Genetics and Molecular Biology, $8, 1-38$ . doi:
1418	10.2202/1544-6115.1294
1419	Lèbre, S., Becq, J., Devaux, F., Stumpf, M. P., & Lelandais, G. (2010). Statistical
1420	inference of the time-varying structure of gene-regulation networks. $BMC$ Sys-
1421	tems Biology, $4(130)$ , 1. doi: 10.1186/1752-0509-4-130
1422	Lenton, T. M., Held, H., Kriegler, E., Hall, J. W., Lucht, W., Rahmstorf, S., &
1423	Schnellnhuber, H. J. (2008). Tipping elements in the Earth's climate system.
1424	$Proceedings \ of \ the \ National \ Academy \ of \ Sciences, \ 105 (6), \ 1786-1793. \qquad {\rm doi:}$
1425	10.1073/pnas.0705414105
1426	Leonard, M., Westra, S., Phatak, A., Lambert, M., van den Hurk, B., McInnes,
1427	K., Stafford-Smith, M. $(2014)$ . A compound event framework for un-
1428	derstanding extreme impacts. $WIREs Climate Change, 5(1), 113-128.$ doi:
1429	10.1002/wcc.252
1430	Li, S., Wang, M., Bond, N. A., Huang, W., Wang, Y., Xu, S., Bai, Y. (2018).
1431	Precursors of September Arctic Sea-Ice Extent Based on Causal Effect Net-
1432	works. Atmosphere, $9(11)$ , 437. doi: 10.3390/atmos9110437

Lin, H., Brunet, G., & Derome, J. (2009). An Observed Connection between the

-49-

1434	North Atlantic Oscillation and the Madden–Julian Oscillation. Journal of Cli-
1435	mate, $22(2)$ , 364-380. doi: 10.1175/2008JCLI2515.1
1436	Lindsay, R., Wensnahan, M., Schweiger, A., & Zhang, J. (2014). Evaluation of Seven
1437	Different Atmospheric Reanalysis Products in the Arctic*. Journal of Climate,
1438	27(7), 2588-2606. doi: 10.1175/JCLI-D-13-00014.1
1439	Lorenz, E. N. (1956). Empirical Orthogonal Functions and Statistical Weather Pre-
1440	diction (Tech. Rep.). Cambridge: Massachusetts Institute of Technology.
1441	Madigan, D., & Raftery, A. E. (1994). Model Selection and Accounting for
1442	Model Uncertainty in Graphical Models Using Occam's Window. Jour-
1443	nal of the American Statistical Association, $89(428)$ , 1535-1546. doi:
1444	10.1080/01621459.1994.10476894
1445	Madigan, D., York, J., & Allard, D. (1995). Bayesian Graphical Models for Dis-
1446	crete Data. International Statistical Review / Revue Internationale de Statis-
1447	$tique, \ 63(2), \ 215-232.$ doi: 10.2307/1403615
1448	Maloney, E. D., Gettelman, A., Ming, Y., Neelin, J. D., Barrie, D., Mariotti, A.,
1449	$\ldots$ Zhao, M. (2019). Process-Oriented Evaluation of Climate and Weather
1450	Forecasting Models. Bulletin of the American Meteorological Society, $100(9)$ ,
1451	1665-1686. doi: 10.1175/BAMS-D-18-0042.1
1452	Marshall, A. G., Hudson, D., Wheeler, M. C., Alves, O., Hendon, H. H., Pook,
1453	M. J., & Risbey, J. S. (2014). Intra-seasonal drivers of extreme heat over Aus-
1454	tralia in observations and POAMA-2. Climate Dynamics, 43(7), 1915–1937.
1455	doi: 10.1007/s00382-013-2016-1
1456	Marshall, G. J. (2002). Trends in Antarctic Geopotential Height and Tempera-
1457	ture: A Comparison between Radiosonde and NCEP–NCAR Reanalysis Data.
1458	$Journal of Climate, \ 15(6), \ 659-674. \qquad {\rm doi:} \ \ 10.1175/1520-0442(2002)015\langle 0659:$
1459	$TIAGHA \rangle 2.0.CO; 2$
1460	Marshall, G. J., & Harangozo, S. A. $$ (2000). An appraisal of NCEP/NCAR reanal-
1461	ysis MSLP data viability for climate studies in the South Pacific. $Geophysical$
1462	Research Letters, $27(19)$ , 3057-3060. doi: 10.1029/2000GL011363
1463	McGraw, M. C., & Barnes, E. A. (2018). Memory Matters: A Case for Granger
1464	Causality in Climate Variability Studies. Journal of Climate, 31(8), 3289-3300.
1465	doi: 10.1175/JCLI-D-17-0334.1

1466 McPhaden, M. J., Santoso, A., & Cai, W. (Eds.). (2021). El Niño Southern Oscilla-

1467	tion in a Changing Climate (Vol. 253). American Geophysical Union.
1468	Mo, K. C., & Ghil, M. (1987). Statistics and Dynamics of Persistent Anoma-
1469	lies. Journal of the Atmospheric Sciences, 44(5), 877-902. doi: 10.1175/
1470	1520-0469(1987)044(0877:SADOPA)2.0.CO;2
1471	Mo, K. C., & Paegle, J. N. (2001). The Pacific–South American modes and their
1472	downstream effects. International Journal of Climatology, 21(10), 1211-1229.
1473	doi: 10.1002/joc.685
1474	Murphy, K., & Mian, S. (1999). Modelling gene expression data using dynamic
1475	Bayesian networks (Tech. Rep.). Berkely, CA: Computer Science Division,
1476	University of California.
1477	Murphy, K. P., & Russell, S. (2002). Dynamic Bayesian networks: representation,
1478	inference and learning. University of California, Berkeley Dissertation.
1479	Nowack, P., Runge, J., Eyring, V., & Haigh, J. D. (2020). Causal networks for cli-
1480	mate model evaluation and constrained projections. Nature Communications,
1481	11(1), 1415. doi: 10.1038/s41467-020-15195-y
1482	O'Kane, T. J., Monselesan, D. P., & Risbey, J. S. (2017). A multiscale re-
1483	examination of the Pacific South American pattern. Mon. Wea. Rev., $145(1)$ ,
1484	379–402. doi: 10.1175/MWR-D-16-0291.1
1485	O'Kane, T. J., Risbey, J. S., Monselesan, D. P., Horenko, I., & Franzke, C. L.
1486	(2016). On the dynamics of persistent states and their secular trends in the
1487	waveguides of the Southern Hemisphere troposphere. Climate Dynamics,
1488	46(11-12), 3567-3597.doi: 10.1007/s00382-015-2786-8
1489	Oliphant, T. E. (2006). A guide to numpy (Vol. 1). Trelgol Publishing USA.
1490	Pearl, J. (1982). Reverend Bayes on Inference Engines: A Distributed Hierarchical
1491	Approach. In Proceedings of the Second National Conference on Artificial In-
1492	telligence (AAAI 1982) (p. 133-136). AAAI Press.
1493	Pearl, J. (1988). Probabilistic Reasoning in Intelligent Systems. San Francisco, Cali-
1494	fornia: Morgan Kaufmann Publishers, Inc.
1495	Pearl, J. (1995). Causal diagrams for empirical research. Biometrika, 82(4), 669-688.
1496	doi: 10.1093/biomet/82.4.669
1497	Pedregosa, F., Varoquaux, G., Gramfort, A., Michel, V., Thirion, B., Grisel, O.,
1498	Duchesnay, E. (2011). Scikit-learn: Machine learning in Python. Journal of
1499	Machine Learning Research, 12, 2825–2830.

1500	Peter, C., Lange, W. d., Musango, J. K., April, K., & Potgieter, A. (2009). Applying
1501	Bayesian modelling to assess climate change effects on biofuel production. $Cli$ -
1502	mate Research, $40(2-3)$ , 249–260. doi: 10.3354/cr00833
1503	Pfleiderer, P., Schleussner, CF., Geiger, T., & Kretschmer, M. (2020). Robust
1504	predictors for seasonal Atlantic hurricane activity identified with causal
1505	effect networks. Weather and Climate Dynamics, $1(2)$ , $313-324$ . doi:
1506	10.5194/wcd-1-313-2020
1507	Punskaya, E., Andrieu, C., Doucet, A., & Fitzgerald, W. J. (2002). Bayesian curve
1508	fitting using MCMC with applications to signal segmentation. $\ \ IEEE \ \ Transac-$
1509	tions on Signal Processing, $50(3)$ , 747-758. doi: 10.1109/78.984776
1510	Radebach, A., Donner, R. V., Runge, J., Donges, J. F., & Kurths, J. (2013). Dis-
1511	entangling different types of El Niño episodes by evolving climate network
1512	analysis. Phys. Rev. E, 88, 052807. doi: 10.1103/PhysRevE.88.052807
1513	Ångström, A. (1935). Teleconnections of climatic changes in present time. $Ge$ -
1514	ografiska Annaler, 17(3-4), 242-258. doi: 10.1080/20014422.1935.11880600
1515	Rayner, N. A., Parker, D. E., Horton, E. B., Folland, C. K., Alexander, L. V., Row-
1516	ell, D. P., Kaplan, A. (2003). Global analyses of sea surface temper-
1517	ature, sea ice, and night marine air temperature since the late nineteenth
1518	century. Journal of Geophysical Research: Atmospheres, 108(D14). doi:
1519	10.1029/2002JD002670
1520	Rogers, J. C., & van Loon, H. (1982). Spatial Variability of Sea Level Pressure
1521	and 500 mb Height Anomalies over the Southern Hemisphere. Monthly
1522	Weather Review, $110(10)$ , 1375-1392. doi: $10.1175/1520-0493(1982)110(1375:$
1523	$SVOSLP$ $\geq 2.0.CO; 2$
1524	Runge, J. (2015). Quantifying information transfer and mediation along causal path-
1525	ways in complex systems. Phys. Rev. E, 92, 062829. doi: 10.1103/PhysRevE
1526	.92.062829
1527	Runge, J. (2018a). Causal network reconstruction from time series: From theoret-
1528	ical assumptions to practical estimation. Chaos: An Interdisciplinary Journal
1529	of Nonlinear Science, 28(7), 075310. doi: 10.1063/1.5025050
1530	Runge, J. (2018b). Conditional independence testing based on a nearest-neighbor
1531	estimator of conditional mutual information. In A. Storkey & F. Perez-Cruz
1532	(Eds.), Proceedings of the Twenty-First International Conference on Artificial

1533	Intelligence and Statistics (Vol. 84, pp. 938–947). Playa Blanca, Lanzarote,
1534	Canary Islands: PMLR.
1535	Runge, J., Bathiany, S., Bollt, E., Camps-Valls, G., Coumou, D., Deyle, E.,
1536	Zscheischler, J. (2019). Inferring causation from time series in Earth system
1537	sciences. Nature Communications, 10, 2553. doi: 10.1038/s41467-019-10105-3
1538	Runge, J., Nowack, P., Kretschmer, M., Flaxman, S., & Sejdinovic, D. (2019).
1539	Detecting and quantifying causal associations in large nonlinear time series
1540	datasets. Science Advances, 5(11). doi: 10.1126/sciadv.aau4996
1541	Runge, J., Petoukhov, V., Donges, J. F., Hlinka, J., Jajcay, N., Vejmelka, M.,
1542	Kurths, J. (2015). Identifying causal gateways and mediators in com-
1543	plex spatio-temporal systems. Nature communications, $6(1)$ , 1–10. doi:
1544	10.1038/ncomms9502
1545	Runge, J., Petoukhov, V., & Kurths, J. (2014). Quantifying the Strength and Delay
1546	of Climatic Interactions: The Ambiguities of Cross Correlation and a Novel
1547	Measure Based on Graphical Models. Journal of Climate, $27(2)$ , 720-739. doi:
1548	10.1175/JCLI-D-13-00159.1
1549	Saggioro, E., de Wiljes, J., Kretschmer, M., & Runge, J. (2020). Reconstruct-
1550	ing regime-dependent causal relationships from observational time series.
1551	(arXiv:2007.00267)
1552	Saji, N. H., Goswami, B. N., Vinayachandran, P. N., & Yamagata, T. (1999). A
1553	dipole mode in the tropical Indian Ocean. Nature, $401(6751)$ , $360-363$ . doi: 10
1554	.1038/43854
1555	Samarasinghe, S. M., Deng, Y., & Ebert-Uphoff, I. (2020). A Causality-Based
1556	View of the Interaction between Synoptic- and Planetary-Scale Atmospheric
1557	Disturbances. Journal of the Atmospheric Sciences, 77(3), 925-941. doi:
1558	10.1175/JAS-D-18-0163.1
1559	Samarasinghe, S. M., McGraw, M. C., Barnes, E. A., & Ebert-Uphoff, I. (2019).
1560	A study of links between the Arctic and the midlatitude jet stream us-
1561	ing Granger and Pearl causality. $Environmetrics, 30(4), e2540.$ doi:
1562	10.1002/env.2540
1563	Schott, F. A., Xie, SP., & McCreary Jr., J. P. (2009). Indian ocean circulation and
1564	climate variability. Reviews of Geophysics, $47(1)$ , RG1002. doi: https://doi
1565	.org/10.1029/2007 RG000245

1566	Schulzweida, U. (2019, October). CDO User Guide. Retrieved from https://doi
1567	.org/10.5281/zenodo.3539275 doi: 10.5281/zenodo.3539275
1568	Sisson, S. A. (2005). Transdimensional Markov Chains. Journal of the
1569	American Statistical Association, 100(471), 1077-1089. doi: 10.1198/
1570	01621450500000664
1571	Spiegelhalter, D. J., Dawid, A. P., Lauritzen, S. L., & Cowell, R. G. (1993).
1572	Bayesian Analysis in Expert Systems. Statistical Science, 8(3), 219–247.
1573	Spiegelhalter, D. J., & Lauritzen, S. L. (1990). Sequential updating of conditional
1574	probabilities on directed graphical structures. Networks, $20(5)$ , 579-605. doi:
1575	10.1002/net.3230200507
1576	Spirtes, P., & Glymour, C. (1991). An algorithm for fast recovery of sparse causal
1577	graphs. Social science computer review, $9(1)$ , $62-72$ .
1578	Steinhaeuser, K., Chawla, N. V., & Ganguly, A. R. (2011). Complex networks as a
1579	unified framework for descriptive analysis and predictive modeling in climate
1580	science. Statistical Analysis and Data Mining: The ASA Data Science Journal,
1581	4(5), 497-511. doi: 10.1002/sam.10100
1582	Steinhaeuser, K., Ganguly, A. R., & Chawla, N. V. (2012). Multivariate and multi-
1583	scale dependence in the global climate system revealed through complex net-
1584	works. Climate Dynamics, $39(3)$ , 889–895. doi: 10.1007/s00382-011-1135-9
1585	Straus, D. M., Molteni, F., & Corti, S. (2017). Atmospheric regimes: The link
1586	between weather and the large scale circulation. In C. L. E. Franzke $\&$
1587	T. J. O'Kane (Eds.), Nonlinear and stochastic climate dynamics (pp. 105–
1588	135). Cambridge University Press, Cambridge.
1589	Thompson, D. W. J., & Wallace, J. M. (1998). The Arctic oscillation signature
1590	in the wintertime geopotential height and temperature fields. $Geophysical Re-$
1591	search Letters, 25(9), 1297-1300. doi: 10.1029/98GL00950
1592	Thompson, D. W. J., & Wallace, J. M. (2000). Annular Modes in the Extratropi-
1593	cal Circulation. Part I: Month-to-Month Variability. Journal of Climate, 13(5),
1594	1000-1016. doi: 10.1175/1520-0442(2000)013 (1000:AMITEC)2.0.CO;2
1595	Trenberth, K. E., Branstator, G. W., Karoly, D., Kumar, A., Lau, NC., & Ro-
1596	pelewski, C. (1998). Progress during TOGA in understanding and mod-
1597	eling global teleconnections associated with tropical sea surface tempera-
1598	tures. Journal of Geophysical Research: Oceans, 103(C7), 14291-14324. doi:

-54-

1599	10.1029/97JC01444
1600	Tsonis, A. A., & Roebber, P. (2004). The architecture of the climate network. Phys-
1601	ica A: Statistical Mechanics and its Applications, 333, 497 - 504. doi: 10.1016/
1602	j.physa.2003.10.045
1603	Tsonis, A. A., Swanson, K., & Kravtsov, S. (2007). A new dynamical mechanism
1604	for major climate shifts. Geophysical Research Letters, $34(13)$ . doi: $10.1029/$
1605	2007GL030288
1606	Tsonis, A. A., & Swanson, K. L. (2008). Topology and Predictability of El Niño and
1607	La Niña Networks. Phys. Rev. Lett., 100, 228502. doi: 10.1103/PhysRevLett
1608	.100.228502
1609	Tsonis, A. A., Swanson, K. L., & Roebber, P. J. (2006). What Do Networks Have to
1610	Do with Climate? Bulletin of the American Meteorological Society, 87(5), 585–
1611	596. (Publisher: American Meteorological Society) doi: 10.1175/BAMS-87-5
1612	-585
1613	Tsonis, A. A., Swanson, K. L., & Wang, G. (2008). On the Role of Atmospheric
1614	Teleconnections in Climate. Journal of Climate, 21(12), 2990-3001. doi: 10
1615	.1175/2007 JCLI1907.1
1616	Uusitalo, L. (2007). Advantages and challenges of Bayesian networks in environ-
1617	mental modelling. Ecological Modelling, 203(3), 312 - 318. doi: 10.1016/j
1618	.ecolmodel.2006.11.033
1619	Van Der Walt, S., Colbert, S. C., & Varoquaux, G. (2011). The numpy array: a
1620	structure for efficient numerical computation. Computing in Science & Engi-
1621	neering, 13(2), 22.
1622	van Loon, H., & Rogers, J. C. (1978). The Seesaw in Winter Temperatures be-
1623	tween Greenland and Northern Europe. Part I: General Description. Monthly
1624	Weather Review, $106(3)$ , 296-310. doi: $10.1175/1520-0493(1978)106(0296:$
1625	TSIWTB $2.0.CO;2$
1626	Vázquez-Patiño, A., Campozano, L., Mendoza, D., & Samaniego, E. (2020). A
1627	causal flow approach for the evaluation of global climate models. International
1628	Journal of Climatology, $40(10)$ , 4497-4517. doi: 10.1002/joc.6470
1629	Virtanen, P., Gommers, R., Oliphant, T. E., Haberland, M., Reddy, T., Cournapeau,
1630	D., SciPy 1.0 Contributors (2020). SciPy 1.0: Fundamental Algorithms

doi:

 $Nature\ Methods,\ 17,\ 261-272.$ 

for Scientific Computing in Python.

1631

1632	10.1038/s41592-019-0686-2
1633	Walker, G. T. (1923). Correlation in Seasonal Variations of Weather, VIII, a prelimi-
1634	nary study of world weather. Memoirs of the India Meteorological Department,
1635	24, 75-131.
1636	Walker, G. T. (1924). Correlations in Seasonal Variations of Weather. I. A further
1637	study of world weather. Memoirs of the India Meteorological Department, 24,
1638	275–332.
1639	Wallace, J. M., & Gutzler, D. S. (1981). Teleconnections in the Geopotential Height
1640	Field during the Northern Hemisphere Winter. Monthly Weather Review,
1641	109(4), 784-812. doi: 10.1175/1520-0493(1981)109 (0784:TITGHF>2.0.CO;2
1642	Wang, G., Swanson, K. L., & Tsonis, A. A. (2009). The pacemaker of major climate
1643	shifts. Geophysical Research Letters, 36(7). doi: 10.1029/2008GL036874
1644	Wang, Y., Gozolchiani, A., Ashkenazy, Y., Berezin, Y., Guez, O., & Havlin, S.
1645	(2013). Dominant Imprint of Rossby Waves in the Climate Network. <i>Phys.</i>
1646	<i>Rev. Lett.</i> , 111, 138501. doi: 10.1103/PhysRevLett.111.138501
1647	Wes McKinney. (2010). Data Structures for Statistical Computing in Python. In
1648	Stéfan van der Walt & Jarrod Millman (Eds.), Proceedings of the 9th Python
1649	in Science Conference (p. 56 - 61). doi: 10.25080/Majora-92bf1922-00a
1650	Wheeler, M. C., & Hendon, H. H. (2004). An All-Season Real-Time Multivariate
1651	MJO Index: Development of an Index for Monitoring and Prediction. <i>Monthly</i>
1652	$Weather \ Review, \ 132 (8), \ 1917-1932. \qquad {\rm doi:} \ \ 10.1175/1520-0493 (2004) \\ 132 \langle 1917. \rangle \\ $
1653	$AARMMI \rangle 2.0.CO;2$
1654	Wolter, K., & Timlin, M. S. (1993). Monitoring ENSO in COADS with a Season-
1655	ally Adjusted Principal Component Index. In Proceedings of the 17th Climate
1656	Diagnostics Workshop (Vol. 57, pp. 52–57).
1657	Wolter, K., & Timlin, M. S. (1998). Measuring the strength of ENSO events: How
1658	does 1997/98 rank? Weather, $53(9)$ , 315-324. doi: 10.1002/j.1477-8696.1998
1659	.tb06408.x
1660	Wolter, K., & Timlin, M. S. (2011). El Niño/Southern Oscillation behaviour since
1661	1871 as diagnosed in an extended multivariate ENSO index (MEI.ext). $Inter-$
1662	national Journal of Climatology, 31(7), 1074-1087. doi: 10.1002/joc.2336
1663	Wu, P. PY., Julian Caley, M., Kendrick, G. A., McMahon, K., & Mengersen, K.
1664	(2018). Dynamic Bayesian network inferencing for non-homogeneous complex

-56-

1665	systems. Journal of the Royal Statistical Society: Series C (Applied Statistics),
1666	67(2), 417-434. doi: 10.1111/rssc.12228
1667	Yamasaki, K., Gozolchiani, A., & Havlin, S. (2008). Climate Networks around the
1668	Globe are Significantly Affected by El Niño. Phys. Rev. Lett., 100, 228501. doi:
1669	10.1103/PhysRevLett.100.228501
1670	Zhang, T., Hoell, A., Perlwitz, J., Eischeid, J., Murray, D., Hoerling, M., & Hamill,
1671	T. M. (2019). Towards Probabilistic Multivariate ENSO Monitoring. Geophys-
1672	ical Research Letters, 46(17-18), 10532-10540. doi: 10.1029/2019GL083946

1672



Figure 3. Subgraphs corresponding to the fitted parent sets of the tropical indices in JRA-55 based on full-year data for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



(a) NNR1 NH extratropical, all seasons (b) JRA-55 N

(b) JRA-55 NH extra tropical, all seasons

Figure 4. Subgraphs corresponding to the fitted parent sets of the NH extratropical indices in (a) NNR1 and (b) JRA-55 for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



Figure 5. Subgraphs corresponding to the fitted parent sets of the SH extratropical indices in (a) NNR1 and (b) JRA-55 for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



Figure 6. Subgraphs corresponding to the fitted parent set of the monthly AO and NAO indices in (a) NNR1 and (b) JRA-55, for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



Figure 7. Subgraphs corresponding to the fitted parent set of the monthly PNA index in (a) NNR1 and (b) JRA-55, for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



(b) JRA-55 SAM, all seasons

Figure 8. Subgraphs corresponding to the fitted parent set of the monthly SAM index in (a) NNR1 and (b) JRA-55, for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



Figure 9. Subgraphs corresponding to the fitted parent set of the monthly NAO indices in (a) NNR1 and (b) JRA-55, for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$  during DJF. All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



Figure 10. Subgraphs corresponding to the fitted parent set of the monthly PSA1 index in (a) NNR1 and (b) JRA-55, for  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$  during JJA. All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.

# Supporting Information for "Dynamic Bayesian networks for evaluation of Granger causal relationships in climate reanalyses"

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- 1. Figures S1 to S15
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## Introduction

This supporting information provides the patterns corresponding to the teleconnection indices analyzed in the main text, summaries of the edge posterior distributions for the alternative choice of hyperparameters, and summaries of the maximum a posteriori (MAP) structure estimates for the fits presented in the main text.

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**Figure S1.** AO loading pattern of 500 hPa geopotential height anomalies in NNR1 (left) and JRA-55 (right).



Figure S2. RMM EOFs in HadISST and NNR1 (left column) and JRA-55 (right column).



MEI.v2 MSLP patterns (base period 19790101 - 20011230)

**Figure S3.** Seasonal MSLP anomaly patterns contributing to the MEI in HadISST and NNR1 (odd numbered rows) and JRA-55 (even numbered rows).



MELv2 SST patterns (base period 19790101 - 20011230)

**Figure S4.** Seasonal SST anomaly patterns contributing to the MEI in HadISST and NNR1 (odd numbered rows) and JRA-55 (even numbered rows).



MEI.v2 $\mathit{u}\text{-wind}$  patterns (base period 19790101 - 20011230)

**Figure S5.** Seasonal zonal wind anomaly patterns contributing to the MEI in HadISST and NNR1 (odd numbered rows) and JRA-55 (even numbered rows).



MEI.v2v-wind patterns (base period 19790101 - 20011230)

**Figure S6.** Seasonal meridional wind anomaly patterns contributing to the MEI in HadISST and NNR1 (odd numbered rows) and JRA-55 (even numbered rows).



MEI.v2 OLR patterns (base period 19790101 - 20011230)

**Figure S7.** Seasonal OLR anomaly patterns contributing to the MEI in HadISST and NNR1 (odd numbered rows) and JRA-55 (even numbered rows).


Daily NH teleconnection patterns (base period 19790101 - 20011230)

**Figure S8.** Northern hemisphere cluster loading patterns of 500 hPa geopotential height anomalies in NNR1 (left column) and JRA-55 (right column).



Monthly PNA pattern (base period 19790101 - 20011230)

**Figure S9.** PNA loading pattern of 500 hPa geopotential height anomalies in NNR1 (left) and JRA-55 (right).



Daily PSA patterns (base period 19790101 - 20011230)

**Figure S10.** PSA loading patterns of 500 hPa geopotential height anomalies in NNR1 (left) and JRA-55 (right).



Figure S11. SAM loading pattern of 500 hPa geopotential height anomalies in NNR1 (left) and JRA-55 (right).



Figure S12. Subgraphs corresponding to the fitted parent sets of the tropical indices in NNR1 based on full-year data for  $a_{\tau} = 0.5$ ,  $b_{\tau} = 10$ , and  $\nu^2 \approx 2$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



Figure S13. Subgraphs corresponding to the fitted parent sets of the tropical indices in JRA-55 based on full-year data for  $a_{\tau} = 0.5$ ,  $b_{\tau} = 10$ , and  $\nu^2 \approx 2$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



(a) NNR1 NH extratropical, all seasons (b) JRA-55 NH extratropical, all seasons

Figure S14. Subgraphs corresponding to the fitted parent sets of the NH extratropical indices in (a) NNR1 and (b) JRA-55 for  $a_{\tau} = 0.5$ ,  $b_{\tau} = 10$ , and  $\nu^2 \approx 2$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.



(a) NNR1 SH extratropical, all seasons

(b) JRA-55 SH extratropical, all seasons

Figure S15. Subgraphs corresponding to the fitted parent sets of the SH extratropical indices in (a) NNR1 and (b) JRA-55 for  $a_{\tau} = 0.5$ ,  $b_{\tau} = 10$ , and  $\nu^2 \approx 2$ . All edges with an estimated posterior probability  $\hat{\pi}$  greater than 0.5 are shown.

**Table S1.** MAP parent sets for monthly tropical teleconnection indices across all seasons for NNR1 and JRA-55 for fits with  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ , showing the estimated posterior probability  $\hat{\pi}$  of the edge, the mean parameter value  $\hat{\beta}$  conditional on the MAP structure, and the 95% posterior HDI. Dashes indicate a node that is not in the MAP parent set for a given reanalysis.

	JRA-55			NNR1			
Parent node	$\hat{\pi}$	β	95% HDI	$\hat{\pi}$	β	95% HDI	
$DMI_t$		,			,		
$\overline{\mathrm{DMI}_{t-1}}$	1.00	0.63	(0.56, 0.69)	1.00	0.74	(0.68, 0.79)	
$SAM_{t-2}$	0.59	0.09	(0.03, 0.16)	0.33	_	_	
$MEI_{t-1}$	0.22	_		0.80	0.13	(0.06, 0.20)	
$MEI_{t-5}$	0.08	_	_	0.65	-0.14	(-0.21, -0.07)	
$AR_{t-1}$	0.04	_	_	0.94	0.10	(0.05, 0.16)	
$MEI_t$							
$MEI_{t-1}$	1.00	1.45	(1.37, 1.54)	1.00	1.43	(1.34, 1.51)	
$MEI_{t-2}$	1.00	-0.85	(-1.00, -0.70)	1.00	-0.73	(-0.87, -0.59)	
$MEI_{t-3}$	1.00	0.65	(0.49, 0.80)	1.00	0.38	(0.24, 0.52)	
$MEI_{t-4}$	1.00	-0.48	(-0.63, -0.31)	0.98	-0.15	(-0.23, -0.07)	
$MEI_{t-5}$	0.65	0.28	(0.13, 0.42)	0.17	_		
$MEI_{t-6}$	0.57	-0.14	(-0.22, -0.06)	0.10	-	-	
$NAO_{t}^{+}$	0.38	0.04	(0.01, 0.06)	0.17	_	_	
$AO_{t-2}$	0.06	_		0.46	0.05	(0.02, 0.07)	
$AR_{t-3}$	0.05	_	_	0.66	-0.04	(-0.07, -0.02)	
$RMM2_{t-3}$	0.34	_	_	0.41	0.04	(0.01, 0.06)	
RMM1 <sub>t</sub>							
MEI <sub>t-3</sub>	0.47	-0.37	(-0.49, -0.24)	0.37	_	-	
$MEI_{t-6}$	0.61	0.24	(0.13, 0.36)	0.33	_	_	
$AR_{t-1}$	0.54	-0.12	(-0.20, -0.04)	0.24	_	_	
$PSA1_{t-3}$	0.38	-0.14	(-0.22, -0.05)	0.11	_	_	
$SAM_{t-3}$	0.34	-0.13	(-0.20, -0.04)	0.09	_	_	
$RMM1_{t-1}$	0.40	-0.12	(-0.19, -0.04)	0.86	-0.14	(-0.21, -0.06)	
$RMM1_{t-3}$	0.99	-0.17	(-0.25, -0.09)	0.97	-0.16	(-0.24, -0.08)	
$RMM1_{t-4}$	1.00	-0.20	(-0.27, -0.12)	1.00	-0.19	(-0.27, -0.11)	
$RMM2_{t-1}$	0.99	-0.17	(-0.25, -0.09)	1.00	-0.20	(-0.27, -0.12)	
$MEI_{t-4}$	0.58	_		0.59	-0.45	(-0.68, -0.22)	
$MEI_{t-5}$	0.41	_	_	0.35	0.34	(0.11, 0.57)	
$RMM2_t$							
$AO_{t-1}$	0.94	0.18	(0.10, 0.26)	0.63	0.15	(0.07, 0.23)	
$DMI_{t-5}$	0.50	-0.12	(-0.20, -0.04)	0.06	-		
$MEI_{t-4}$	0.56	-0.15	(-0.23, -0.07)	0.52	-	_	
$AR_{t-1}$	0.44	0.13	(0.05, 0.21)	0.93	0.16	(0.08, 0.24)	
$\text{RMM1}_{t-1}$	0.71	0.13	(0.05, 0.21)	0.26	0.14	(0.06, 0.22)	
$RMM2_{t-1}$	0.61	-0.13	(-0.21, -0.05)	0.45	-	_	
$RMM2_{t-3}$	1.00	-0.24	(-0.32, -0.16)	0.98	-0.16	(-0.23, -0.08)	
$RMM2_{t-4}$	0.94	-0.17	(-0.25, -0.09)	0.94	-0.16	(-0.24, -0.09)	
$MEI_{t-1}$	0.10	-	_	0.22	-0.18	(-0.26, -0.10)	
$AR_{t-6}$	0.03	-	-	0.33	-0.11	(-0.19, -0.04)	
$PSA1_{t-3}$	0.05	-	-	0.59	-0.13	(-0.22, -0.05)	
$\text{RMM1}_{t-5}$	0.03	-	-	0.20	0.11	(0.03, 0.19)	
$RMM2_{t-2}$	0.02	-	_	0.16	0.13	(0.05, 0.20)	

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**Table S2.** MAP parent sets for monthly SH extratropical teleconnection indices across all seasons for NNR1 and JRA-55 for fits with  $a_{\tau} = 1.5$ ,  $b_{\tau} = 20$ , and  $\nu^2 = 3$ , showing the estimated posterior probability  $\hat{\pi}$  of the edge, the mean parameter value  $\hat{\beta}$  conditional on the MAP structure, and the 95% posterior HDI. Dashes indicate a node that is not in the MAP parent set for a given reanalysis.

	$_{\rm JR}$	A-55	NNR1			
$\hat{\pi}$	$\hat{\beta}$	95% HDI	$\hat{\pi}$	β	95% HDI	
0.94	-0.15	(-0.23, -0.07)	0.68	-0.13	(-0.21, -0.05)	
0.39	-0.11	(-0.20, -0.04)	0.60	-0.13	(-0.21, -0.05)	
0.49	0.12	(0.04, 0.21)	0.96	0.16	(0.08, 0.24)	
0.17	-		0.58	-0.12	(-0.21, -0.04)	
0.31	0.12	(0.03, 0.20)	0.30	-	_	
0.49	0.12	(0.04, 0.21)	0.34	-	-	
0.73	0.14	(0.05, 0.21)	0.43	-	_	
0.47	0.12	(0.04, 0.20)	0.41	-	-	
0.07	—	_	0.54	-0.14	(-0.23, -0.06)	
0.75	-0.30	(-0.45, -0.16)	0.51	-	_	
0.76	0.27	(0.13, 0.42)	0.54	-	_	
1.00	0.29	(0.21, 0.37)	1.00	0.32	(0.24, 0.40)	
	$ \hat{\pi} \\ 0.94 \\ 0.39 \\ 0.49 \\ 0.17 \\ 0.31 \\ 0.49 \\ 0.73 \\ 0.73 \\ 0.07 \\ 0.07 \\ 0.75 \\ 0.76 \\ 1.00 \\ \end{array} $	$\begin{array}{c c} & & & & \\ \hline \hat{\pi} & & \hat{\beta} \\ \hline 0.94 & -0.15 \\ 0.39 & -0.11 \\ 0.49 & 0.12 \\ 0.17 & - \\ \hline 0.31 & 0.12 \\ 0.49 & 0.12 \\ 0.73 & 0.14 \\ 0.73 & 0.14 \\ 0.73 & 0.12 \\ 0.07 & - \\ \hline 0.75 & -0.30 \\ 0.76 & 0.27 \\ 1.00 & 0.29 \\ \end{array}$	$\begin{array}{c c c c c c c c c c c c c c c c c c c $	$\begin{array}{c c c c c c c c c c c c c c c c c c c $	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	