

On the Burridge and Knopoff Model and the Theoretical Seismicity

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Abstract

In the mid-sixties of the 20th century, a seismicity one-dimensional model, based on a spring system with friction, was presented by the two authors mentioned in the title. Through this model was studied the viscosity behaviour involved in the stratum mutual displacement. However, in the analysis pertinent to the occasion, seeing the fig.13 ibid it can be observed an error related to the viscosity function (low speeds) limits, which determines its “performance” during the time - lapse of the stratum breaking. From the height of the present knowledge, using the Elementary Catastrophes Theory developed in later times by René Thom and Zeeman, we have reached certain conclusions which, although do not refute the results of said article, they can significantly improve some aspects of the model and in general of the earthquakes physics.

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17 **1 Introduction**

18 Starting in the first half of the 20th century, the data coming from seismic obser-
 19 vations lead to the (although loin) possibility of arriving at earthquakes forecasts based
 20 on a better understanding of their mechanisms. In the article by Burridge, R., & Knopoff,
 21 L. (1967) a study of telluric phenomena through simulations and experiments is proposed,
 22 beginning from an introduced there one-dimensional model. After this first attempt to
 23 analyze the details of the mechanisms of earthquakes, the results of other investigations
 24 have come to light. The report presented by Dahl, P. R. (1968) is a reference that is cited
 25 by different authors when considering a friction model; however, this work does not delve
 26 into the dependence in relation to relative velocity and neither in the behavior that oc-
 27 curs at critical points. In the article by Haessig Jr, D. A., & Friedland, B. (1991), the
 28 friction model and its dependence on relative velocity are presented, however, even in
 29 its approach there is no clarity when considering the critical points. De Sousa Vieira,
 30 M. (1999) assumes the earthquake model from a different perspective that considers Chaos
 31 theory, nevertheless, there is no type of consideration regarding the friction dependence
 32 on the relative velocity. In the article by Carlson, J. M., & Langer, J. S. (1989) taking
 33 into account the characteristics presented in the Burridge, R., & Knopoff, L. (1967) model,
 34 the dependence of the relative velocity with respect to the friction force leads to simi-
 35 lar considerations that leave aside the critical points. Later, Karnopp, D. (1985) improves
 36 the experiments and simulations made by the previous authors, reaching results that are
 37 in some sense inconclusive. However, it should be noted that their conclusions are based
 38 on an empirical approach and that is why it is difficult to agree or rule out the validity
 39 of the different reasonings.

40 A later article by Galvanetto, U. (2002) reconsiders the characteristic form of fric-
 41 tion but the motion equation leaves aside the physical phenomenon around critical points.
 42 More recently Awrejcewicz, J., & Olejnik, P. (2005) reviewed the different friction mod-
 43 els, including the one presented in Burridge, R., & Knopoff, L. (1967), however, the dis-
 44 continuity around the critical points are neglected.

45 In these circumstances, taking into account the high complexity estimating any type
 46 of viscosity, we have opted for the application of a tool of a completely theoretical na-
 47 ture. Our intention is to use the Elementary Catastrophes Theory in order to obtain rig-
 48 orously the friction behaviour within the Newton's viscosity segment of the relative ve-
 49 locity domain where, in our modest opinion, unfortunately, an important part of more
 50 objective information has remained lacking or lost. All this came to the surface when
 51 we during the analysis of the fig.13 of the first "founding" article have detected a non-
 52 compliance error at the limits of the resistance function values. The catastrophes the-
 53 ory is part of the (semi-)qualitative theory of the nonlinear complex systems (Thom, R.
 54 (1975), *Mathematical Models of Morphogenesis*, (1974)). There it is considered how mi-
 55 nuscule (even infinitesimal) external perturbations change the behavior of a specific sys-
 56 tem (no matter its nature) within certain critical circumstances (Zeeman, E. C. (1976)).

57 The strong conviction that the Catastrophes Theory in this case is applicable, is
 58 based on an epistemological point of view, similar to the justification used when it is about
 59 the Statistical Mathematics handling i.e., when one has to deal with large numbers. In
 60 others words we agree to take that, the mathematics, in general, reflects the objective
 61 laws of the nature and therefore, if we have a situation where a certain classification the-
 62 orem allows only a strongly determinate set of options, the last is also valid in the “real
 63 life” which in the present context is the geological and geophysical sciences.

64 2 The total (“all-inclusive”) friction formula

65 In the Burrige, R., & Knopoff, L. (1967) article it is used the viscosity function
 66 $F(v)$ that includes all the effects of the between stratum resistance (although immedi-
 67 ately below fig. 13 in the text a symbol F^* is inserted apart from F). Following the con-
 68 cepts of this article, we will split the function’s graph by three branches (the symbols
 69 $-\infty$ and ∞ signify its natural finite limits)

$$F(v) = \begin{cases} F_-(v) \\ f(v) \\ F_+(v). \end{cases} \quad (1)$$

70 Which are as follows: the left part that matches the expression used in the mentioned
 71 article (with its argument between $-\infty$ and $-H$, see our fig.1):

$$F_-(v) = \frac{B}{1 - A(v + H)} - Ev, \quad (2)$$

72 the function $f(v) = f(v; \alpha) = v^3 - \alpha v$ that represents the simplest, so-called fold catas-
 73 trophie (Bröcker, T. (1978)) and which velocity domain is $[-H, H]$ being α a positive num-
 74 ber and finally, the right part, which also matches the expression used in the cited ar-
 75 ticle (of course, its argument belongs to the interval locked between H and ∞ , see again
 76 our fig.1)

$$F_+(v) = -\frac{B}{1 + A(v - H)} - Ev. \quad (3)$$

77 Here A , B , E , and H are certain positive constants which values can be found empiri-
 78 cally. In that, is important underline the same branches dimensions of F : speed cubed.
 79 Further, always when we refer to the two figures in this article we will use the adjective
 80 “our”, without mentioning anything else - this are figures in the Burrige, R., & Knopoff,
 81 L. (1967) article while, the graphics, coming from other sources, will be cited normally.

82 3 The Catastrophe Theory correction

83 According to the top graph of the fig.11 it is presented the simple law of friction
 84 without seismic radiation, with instability at zero velocity where there is an “instanta-
 85 neous” and directly jump from B to $-B$. As this is unrealistic, the bottom part of the
 86 same figure presents a situation where the jump is already indirect: at $v = 0$ the fric-
 87 tion first goes up (to B), then goes down (to $-B$), and, finally, it rises to a certain value.
 88 But this is even less acceptable than the above. Let us remember this behaviour to com-
 89 pare it later with the borderline case results to which it leads our proposal.

90 In the above mentioned Burrige, R., & Knopoff, L. (1967) article (and in other
 91 ones too), one can read about the potential energy: “. . . when all the springs are stretched
 92 so that the masses are all on the verge of instability. . . any trigger applied to the system
 93 will cause a large shock to take place. . .” which means, that there, there are all the con-
 94 ditions to be applied the Elementary Catastrophes Theory - a sudden response during
 95 the little by little continuous perturbations increasing. We are going to take advantage

96 of the well-known principle of simplicity by using the simplest known elementary catas-
 97 trophe – the fold catastrophe (Bröcker, T. (1978)). Graphically, after all the three branches
 98 of the total resistant function (1) have been assembled, we obtain the following figure:

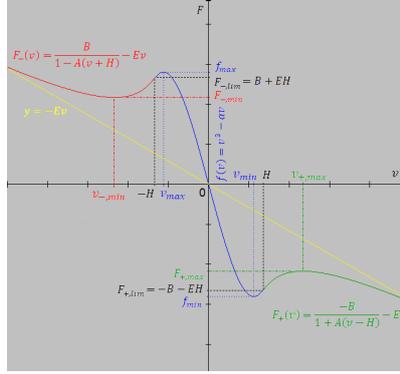


Figure 1. Example of a total resistant function graph , with $A = 1$, $B = 2$, $E = 0.5$, $H = 1.3247$, $\alpha = 3.76449027$.

99 In order to achieve the analytical construction let us sew the three branches of $F(v)$
 100 at its two joining points $-H$ and H , see the equation (1), (2), and (3):

$$F_-(H) = f(-H), \quad f(H) = F_+(H). \quad (4)$$

101 And after to executing the calculations we obtain:

$$B + EH = -H^3 + \alpha H \Leftrightarrow H^3 - \alpha H = -B - EH. \quad (5)$$

102 Because in Classical Physics it is settled that all the auxiliary expressions with phys-
 103 ical sense are differentiable at least once time, we iron the two joints of the graph (see
 104 our fig.1):

$$F'_-(-H) = f'(-H), \quad f'(H) = F'_+(H), \quad (6)$$

105 and this gives us respectively:

$$BA - E = 3H^2 - \alpha \Leftrightarrow 3H^2 - \alpha = BA - E. \quad (7)$$

106 From the equation (7), we can express directly our unique co-variable i.e. the per-
 107 turbation $\alpha = 3H^2 + E - BA$. Despite, it easy to see, that, the total resistance func-
 108 tion $F = F(v)$ has a continuous first derivative, it cannot be polished more, however,
 109 it is not necessary either according to the well-known theorem of existence and unicity
 110 of the ordinary differential equations.

111 Multiplying the equation (7) by H and subtracting it from the equation (5) we get:

$$-2H^3 = -B - BAH. \quad (8)$$

112 That is, for H we have the following cubic equation:

$$2H^3 - BAH - B = 0. \quad (9)$$

113 Its unique real solution according to the Cardano procedure is:

$$H = \frac{1}{2} \left[\sqrt[3]{2B \left(1 + \sqrt[2]{1 - \frac{2BA^3}{27}} \right)} + \sqrt[3]{2B \left(1 - \sqrt[2]{1 - \frac{2BA^3}{27}} \right)} \right] > 0. \quad (10)$$

114 The other two roots are conjugated complex numbers and that is why we are going to
 115 abandon them. It is obvious that there are limitations on A and B (the discriminants
 116 of a real square root are non-negative):

$$BA^3 \leq \frac{27}{2}. \quad (11)$$

117 The above cubic equation allows us to deduce that in this case, the parameter α
 118 will always be positive; in fact $B > 0$ then from eq.(9) $BA - 2H^2 < 0$ and hence $3H^2 >$
 119 BA . Using the existence of first derivatives we can easily find the coordinates of the four
 120 extremities of our curve. These are

$$v_{\pm, \min} = \mp \left(H - \frac{1}{A} + \sqrt{\frac{B}{AE}} \right), \quad F_{\pm, \min} = F(v_{\pm, \min}) = \pm \left[E \left(H - \frac{1}{A} \right) + 2\sqrt{\frac{BE}{A}} \right], \quad (12)$$

121 and for the “global” extremities (watch carefully our figure 1):

$$v_{\min}^{max} = \mp \sqrt{\frac{\alpha}{3}}, \quad F_{\min}^{max} = F(v_{\min}^{max}) = f(v_{\min}^{max}) = \pm 2\sqrt{\left(\frac{\alpha}{3}\right)^3}. \quad (13)$$

122 A look at the fig. 11 of the analyzed article, leads us to the following speculation around
 123 the error that we observed in the fig.13 ibid. If on our fig.1 we carry out a horizontal com-
 124 pression eliminating thus the velocity interval presented by the fold catastrophe, we will
 obtain the following graph (see fig.2):

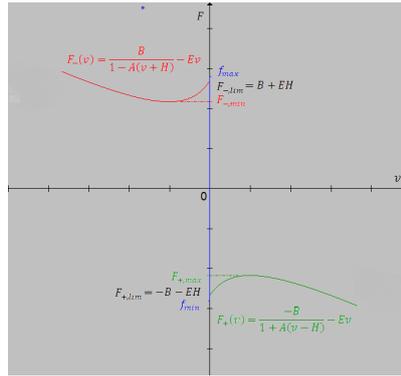


Figure 2. The borderline case of our fig.1 (the argument here is unimportant).

125

126 The one that explains the not elegant appearance of the vertical in $v = 0$ segment
 127 between f_{max} and $F_{-,lim}$ (as well as the one between $F_{+,lim}$ and f_{min}) in other authors.
 128 So in the diploid form of the graph (fig.13) the maximal value (at low speeds) is B (the

error) and not $B+EH$; in all probability, it is (implicitly) admitted by the authors that our f_{max} (they do not introduce it but, there must be one like this) coincide with $B+EH$ of the fig.11 and for some reason must be removed.

Here is the place to find out what happens in the bifurcation values surroundings $\alpha = 0$. The point is that, because the catastrophe function behaviour is local, the branches F_- and F_+ are not always presented by the formulae (2) and (3). Actually, the eq.(11) indicates the region of its validity. Anyway, the respective curve of the total resistance “distributes” itself along the abscissa.

4 Conclusions

The advantage of our approach consists of the possibility to consider the real behavior of the seismic viscosity around the critical points without resorting to any kind of (particular) modeling. It is very significant that, if we do a velocity (the last one is relative) inversion in the freehand graphs of the Karnopp, D. (1985) article, namely a) and b), we will see exactly a caricature of our graphs, 2 and 1 respectively, obtained based on a rigorous theory.

In this order of ideas, we dare to say that maybe the time has come to redo Burridge and Knopoff’s experiment, but this time using the latest in 21st-century technology: multichannel NATIONAL INSTRUMENTS interfaces with the powerful software LabVIEW RT. This would allow to carry out an appropriate acquisition of data in real-time which could later be “delayed in time” and through a computer to compare our graph with the experiment in the low speeds domain.

However, we think the best benefit that could be obtained through this approach consists of the real possibility of predicting earthquakes based on the real time - behavior - analysis of the friction between $v_{-,min}$ and v_{max} as well as between v_{min} and $v_{+,max}$.

Acknowledgments

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