Denoising surface waves extracted from ambient noise using three-station interferometry: Methodology and application to 1-D linear array

Hongrui Qiu¹, Fenglin Niu¹, and Lei Qin²

¹Rice University ²University of Southern California

November 23, 2022

Abstract

We develop an automatic workflow for denoising the fundamental mode surface wave from ambient noise cross correlations (ANCs) calculated for a dense linear array. The surface wave signal traveling between each station pair is first enhanced through three-station interferometry. Then, phase travel times at different periods are determined in the frequency domain. The proposed array-based method is applied to a 1.6-km-long dense linear nodal array crossing surface traces of the San Jacinto fault near Anza, California. Surface wave signals in ANCs of the nodal array are significantly enhanced after denoising, particularly at high frequencies (> 2 Hz). Phase travel times are extracted reliably in the period ranges of 0.3-1.3 s and 0.3-1.6 s for Rayleigh and Love waves, respectively. The corresponding period-dependent phase velocity profiles derived from the eikonal equation reveal high-resolution details of fault zone internal structures beneath the array. A broad (500-1000 m) low-velocity zone that narrows with increasing period is observed, illuminating a flower-shaped structure of the San Jacinto fault damage zone.

1 2 3	Denoising surface waves extracted from ambient noise using three-station interferometry: Methodology and application to 1-D linear array		
4 5	Hongrui Qiu ¹ , Fenglin Niu ¹ , and Lei Qin ²		
6	¹ Department of Earth, Environmental and Planetary Sciences, Rice University, Houston, TX,		
7	USA		
8	² Department of Earth Sciences, University of Southern California, Los Angeles, CA, USA		
9			
10	Corresponding author: Hongrui Qiu (hq7@rice.edu; qiuhonrui@gmail.com)		
11			
12			
13	Key points:		
14 15	• Surface waves from ambient noise cross correlations are significantly enhanced at high frequencies using three-station interferometry		
16 17	• Phase travel times are extracted reliably between 0.3-1.6 s for a 1.6-km-long linear array and are used to perform surface wave tomography		
18 19	• Phase velocity models of Rayleigh and Love waves derived via eikonal tomography reveal high-resolution fault zone images		

20 Abstract

21 We develop an automatic workflow for denoising the fundamental mode surface wave 22 from ambient noise cross correlations (ANCs) calculated for a dense linear array. The 23 surface wave signal traveling between each station pair is first enhanced through three-24 station interferometry. Then, phase travel times at different periods are determined in the 25 frequency domain. The proposed array-based method is applied to a 1.6-km-long dense 26 linear nodal array crossing surface traces of the San Jacinto fault near Anza, California. 27 Surface wave signals in ANCs of the nodal array are significantly enhanced after 28 denoising, particularly at high frequencies (> 2 Hz). Phase travel times are extracted 29 reliably in the period ranges of 0.3-1.3 s and 0.3-1.6 s for Rayleigh and Love waves, 30 respectively. The corresponding period-dependent phase velocity profiles derived from 31 the eikonal equation reveal high-resolution details of fault zone internal structures 32 beneath the array. A broad (500-1000 m) low-velocity zone that narrows with increasing 33 period is observed, illuminating a flower-shaped structure of the San Jacinto fault damage 34 zone.

35 Plain Language Summary

36 Properties of fault damage zone (width of 100-1000's meters), such as its geometry and 37 velocity reduction compared to the surrounding host rock, can have a profound impact on 38 our understandings of earthquake ruptures and the long-term behavior of the fault. 39 Several dense nodal arrays with 10-100 m spacing and aperture of a few kilometers were 40 deployed crossing surface traces of major faults, to provide high-resolution images of the 41 fault zone internal structures. Surface waves propagate between every two sensors with 42 frequency-dependent speeds are extracted from ambient noise cross correlations. By 43 measuring the relation between velocity and frequency, we can infer structures at depth 44 as surface waves are more sensitive to shallow structures at a higher frequency. However, 45 surface waves extracted from ambient noise at high frequencies (> 2 Hz) that are essential 46 to image fault zone in the top 100's meters are often very noisy. Here, we develop a new 47 method that utilizes three-station interferometry to suppress signals that are not traveling 48 on the surface. The quality of surface waves is significantly improved after the denoising,

especially at high frequencies (> 2 Hz), providing more reliable measurements and better
constraints on fault zone internal structures at shallow depth.

51 **1. Introduction**

52 Noise-based surface wave tomography has been widely used to resolve crustal 53 structures at various scales (e.g., Lin et al., 2009; Qiu et al., 2019; Wang et al., 2019; 54 Zigone et al., 2019). Analysis of high frequency (e.g., > 1 Hz) surface waves provides 55 crucial information on the shallow (top 10s to 100s of meters) materials with 56 unprecedented spatial resolution and thus improves our understanding of the local 57 seismic hazard. In contrast to high-quality signals at long periods (e.g., > 2 s), extraction 58 of surface waves from ambient noise cross correlations (ANCs) calculated at high 59 frequencies (e.g., > 1 Hz) remains a challenging topic due to its low signal to noise ratio (SNR). Previous studies that utilize ANCs at high frequencies must first enhance the 60 61 surface wave signals by performing preprocessing and/or postprocessing steps that are 62 often ad hoc and may only work well for a specific dataset.

63 In the present paper, we develop a simple workflow, based on the idea of three-station 64 interferometry proposed by Zhang et al. (2020), that effectively enhances surface waves, 65 particularly at high frequencies, for ANCs of a 1-D linear array. We apply this method to 66 data recorded by a dense linear array deployed at the Ramona Reservation (RR) site 67 across surface traces of the San Jacinto fault, near Anza (Fig. 1). Seismic waveforms from the RR array have been analyzed for fault zone internal structures in Qin et al. 68 69 (2020). ANCs were computed for each station pair of the RR array in Wang et al. (2019). 70 To enhance surface waves with low SNR in ANCs at high frequencies, Wang et al. 71 (2019) first applied a period-dependent tapering window and then applied double-72 beamforming tomography to derive Rayleigh wave phase velocities for periods from 0.3 73 s to 0.8 s beneath the array.

Here, we first describe station configuration and ANC data of the RR array in section 2. Then, in section 3, following the flow chart illustrated in Figure 2, we present the theoretical formulation for three-station interferometry using ANCs of a 1-D linear array and illustrate the surface wave denoising process with a subset of ANC data computed from the linear segment of the RR array. In section 4, surface wave phase travel times are

manuscript submitted to Journal of Geophysical Research: Solid Earth

first extracted from the denoised wavefield and then inverted for phase velocity dispersion models via the eikonal equation. Discussion of the denoising method and comparison between the resulting phase velocity profiles and fault zone images from previous studies (Qin et al., 2020; Wang et al., 2019) are presented in section 5.

83 2. Data

84 The RR array (red triangle in Fig. 1b) is located at north of Anza (blue square in Fig. 85 1b), California, and crosses surface traces of the Clark segment of the San Jacinto fault (Fig. 1a). The array consists of 94 three-component 5-Hz Fairfield geophones (balloons 86 87 in Fig. 1a) that were set to record continuously for a month with a sampling rate of 500 88 Hz. ANC is obtained by first computing cross correlations of ambient noise data in 5-min 89 windows, and then stacking them over the entire recording period for each station pair 90 (Wang et al., 2019). The positive and negative time lags of the monthlong stacked ANC 91 are fold and averaged to suppress the effects of the asymmetric noise source distribution. 92 We use ANCs of a sub-array RR01-RR47 (yellow, blue, and red balloons in Fig. 1a) to 93 demonstrate the surface wave denoising process (Fig. 2a) developed in this study. The 94 sub-array has 47 stations with an average station spacing of ~30 m and an aperture of 95 ~1.6 km.

96 We project stations in the sub-array to the straight line connecting RR01 and RR47 97 (cyan dashed line in Fig. 1a) and compare interstation distances calculated using station 98 locations before and after the projection. The comparison yields negligible differences (< 99 1%) suggesting that the sub-array RR01-RR47 is in a 1-D linear configuration (later 100 referred to as "the linear RR array"). In Wang et al. (2019), a period-dependent velocity 101 threshold is applied to taper off the contamination of body waves or potential higher-102 mode surface waves. In this study, however, we use a tapering window, between a 103 moveout velocity range of 2 km/s and 0.1 km/s (white dashed lines in Figs. 1c-d), to ANCs of Transverse-Transverse (TT) and Vertical-Vertical (ZZ) components. The 104 105 tapered ANCs are then filtered between 0.2 Hz and 10 Hz and depicted as colormaps in Figures 1c-d. Before performing surface wave denoising (Section 3) and tomography 106 107 (Section 4) illustrated in Figure 2, a series of Gaussian narrow bandpass filters centered on consecutive periods between 0.3 s and 1.6 s are applied to the ANCs shown in Figures 108

109 1c-d and then each filtered waveform is normalized by its corresponding maximum. The

110 Gaussian narrow bandpass filters are generated following section 3.1 of Qiu et al. (2019).

111 **3. Surface Wave Denoising**

112 Let $G_{i,j}(t)$ be the positive lag of ANC for the station pair of *i*-th (virtual source) and *j*-113 th (virtual receiver) sensors in the linear RR array (yellow, green, and red triangles in Fig. 114 1a), we can expand it as

$$G_{i_j}(t) = S_{i_j}(t) + B_{i_j}(t) + N_{i_j}(t),$$
(1)

where $S_{i,j}(t)$ and $B_{i,j}(t)$ represent signals traveling on the surface (i.e. surface waves) and at depth (i.e. body waves) between the source *i* and receiver *j*, respectively. $N_{i,j}(t)$ is the residual (later referred to as "background noise"). This section aims to develop a denoising process that preserves $S_{i,j}(t)$ while suppressing $B_{i,j}(t)$ and $N_{i,j}(t)$ in equation 1.

119 **3.1 Three-station interferometry for a 1-D linear Array**

120 Since surface waves are dispersive, let $\tilde{G}_{i_j}(\omega)$ be the Fourier transform of $G_{i_j}(t)$ at 121 the angular frequency ω , we can rewrite equation 1 in the frequency domain

$$\tilde{G}_{i_{j}}(\omega) = A_{G_{i}i_{j}} \cdot e^{i \cdot \varphi_{G_{i}i_{j}}} = \tilde{S}_{i_{j}}(\omega) + \tilde{B}_{i_{j}}(\omega) + \tilde{N}_{i_{j}}(\omega)$$

$$= \sum_{S} A_{S_{i}i_{j}} \cdot e^{-i\left(\omega \cdot T_{i_{j}}^{S} + \varphi_{S}\right)} + \sum_{B} A_{B_{i}i_{j}} \cdot e^{-i\left(\omega \cdot T_{i_{j}}^{B} + \varphi_{B}\right)} + \tilde{N}_{i_{j}}(\omega), \qquad (2a)$$

where A_{S_ij} and T_{ij}^S are amplitude spectrum and phase travel time of surface wave signals in ANC at the angular frequency ω that propagate between the *i*-th and *j*-th stations, while A_{B_ij} and T_{ij}^B represent those of body wave signals. φ_s and φ_B are initial phases of surface- and body-wave signals in the ANC, respectively, and dependent on the distribution of ambient noise sources (e.g., Lin et al., 2008). Since the fundamental mode surface wave, $\tilde{F}_{i_j}(\omega)$, is often the dominant signal in ANC, by assuming the highermode surface waves are negligible, we, therefore, can simplify equation 2a as:

$$\tilde{G}_{i_j}(\omega) = \tilde{F}_{i_j}(\omega) + \tilde{O}_{i_j}(\omega) = A_{F_ij} \cdot e^{-i\left(\omega \cdot T_{ij}^F + \varphi_F\right)} + \tilde{O}_{i_j}(\omega),$$
(2b)

where the symbol or subscript *F* stands for the fundamental mode surface wave. $\tilde{O}_{i_j}(\omega) = \tilde{B}_{i_j}(\omega) + \tilde{N}_{i_j}(\omega)$, that consists of signals from body waves and background noise, is the term we want to suppress in the denoising process. It is interesting to note 132 that $\varphi_F = \pi/4$ for an azimuthally homogenous ambient noise source distribution 133 (Snieder, 2004), whereas $\varphi_F = 0$ when noise sources are only present in line with the 134 station pair *i* and *j* (Lin et al., 2008).

For surface waves of a certain (e.g., fundamental) mode traveling between three stations i < j < k in a 1-D linear array, the travel times satisfy the following relation

$$T_{ik}^S = T_{ij}^S + T_{jk}^S, (3a)$$

137 whereas

$$T_{ik}^B < T_{ij}^B + T_{jk}^B, \tag{3b}$$

for body waves. Therefore, we introduce a third station k and perform three-station interferometry following Zhang et al. (2020):

$$\tilde{I}_{i_{j}}(\omega;k) = \begin{cases} \tilde{G}_{i_{k}}^{*}(\omega) \cdot \tilde{G}_{j_{k}}(\omega), & k < i \\ \tilde{G}_{i_{k}}(\omega) \cdot \tilde{G}_{j_{k}}(\omega), & i < k < j. \\ \tilde{G}_{i_{k}}(\omega) \cdot \tilde{G}_{j_{k}}^{*}(\omega), & k > j \end{cases}$$

$$(4a)$$

In equation 4a, we cross correlate $G_{i_k}(t)$ and $G_{j_k}(t)$ in the time domain, when k < i or k > i (later referred to as "outer-source zone"). The interferometry becomes equivalent to the convolution of $G_{i_k}(t)$ and $G_{j_k}(t)$ in the time domain for station k located within the two virtual sources (i.e., i < k < j; later referred to as "inter-source zone"). For the case k = i or j, we define $\tilde{I}_{i_j}(\omega; k) = A_{G_ij}^2 \cdot e^{i \cdot \varphi_{G_ij}}$ that approximates the convolution of $G_{i_i}(t)$ and $G_{i_i}(t)$ or $G_{j_j}(t)$, by assuming the amplitude spectrum of the auto-correlation $G_{i_i}(t)$ or $G_{j_j}(t)$ is similar to that of $G_{i_j}(t)$, i.e., $A_{G_ii} \approx A_{G_jj} \approx A_{G_ij}$.

147 Combining equations 2b, 3a, and 4a, if the fundamental mode surface wave is the 148 dominate signal in ANC (i.e., \tilde{O}_{i_j} in Equation 2b is negligible), phase term of the 149 interferogram $\tilde{I}_{i_j}(\omega; k) = A_{i_j k} \cdot e^{i \cdot \varphi_{i_j k}}$ is given by

$$\varphi_{ij_k}(\omega) = \begin{cases} -\omega \cdot T_{ij}^F - 2\varphi_F, & i < k < j \\ -\omega \cdot T_{ij}^F - \varphi_F, & k = i \text{ or } j. \\ -\omega \cdot T_{ij}^F, & k > i \text{ or } k < j \end{cases}$$
(4b)

150 φ_F and T_{ij}^F denote the initial phase and phase travel time of the fundamental mode surface 151 wave signal (Equation 2b) extracted from the ANC of station pair *i* and *j*. Equation 4b 152 suggests that the interferograms within either the inter- or outer-source zones share the 153 same phase, whereas interferograms from different zones are only aligned in phase when 154 φ_F is zero. In cases when the term \tilde{O}_{i_j} is significant, we can divide the interferogram 155 $\tilde{I}_{i_j}(\omega;k)$ into two components: $\tilde{I}_{i_j}^F(\omega;k)$ and $\tilde{I}_{i_j}^O(\omega;k)$. $\tilde{I}_{i_j}^F(\omega;k)$ represents the 156 interferogram that only involves the fundamental mode surface wave signal, i.e., when 157 \tilde{O}_{i_j} is set to zero. The phase of $\tilde{I}_{i_j}^F(\omega;k)$, given by equation 4b, is independent of k 158 when φ_F is zero. In contrast, $\tilde{I}_{i_j}^O(\omega;k)$ engages contributions from body waves and 159 background noise, and thus has a phase that varies significantly with k.

160 **3.2 Surface wave denoising of the linear RR Array**

161 As described in section 2, we first apply a series of Gaussian narrow bandpass filters 162 centered on periods between 0.3 s and 1.6 s to the ANCs of the linear RR array (Figs. 1c-163 d). The period range is determined based on the station spacing and array aperture. Then 164 the denoising process (Fig. 2a) is performed on the filtered ANCs for each period and 165 component separately. Figure 3 shows results of the three-station interferometry applied 166 to ANCs of TT component filtered at 0.3 s (Figs. S1a-c for 0.8 s and Figs. S2a-c & S3a-c for ZZ component) for an example station pair RR10 (i = 10) and RR40 (j = 40). The 167 168 ANCs with RR10 and RR40 as the virtual source, $G_{i,k}(t)$ and $G_{i,k}(t)$ filtered at 0.3 s, are shown in Figures 3a and 3b, respectively. Black and blue waveforms denote ANCs with 169 170 station k inside the outer-source (k < 10 or k > 40) and inter-source (10 < k < 40) zones, 171 respectively, whereas the ANC of station pair RR10 and RR40 is depicted in red (i.e., k is 172 the y-axis of Fig. 3).

173 Figure 3c demonstrates the resulting interferograms in the time domain computed 174 following equation 4a (blue for convolution, black for cross correlation). Both the phase 175 and envelope functions of the interferograms vary significantly with station k. Such k-176 value-dependent interferograms are also observed using ZZ component ANCs filtered at 177 0.3 s (e.g., Fig. S2c). This is because of the low SNR for the fundamental mode surface wave (i.e., $\tilde{O}_{i,i}$ in Equation 2b is non-negligible) in the ANCs filtered at 0.3 s (e.g., Figs. 178 179 3a-b and S2a-b). On the other hand, the interferograms are coherent and aligned well in phase for all k values at 0.8 s (e.g., Figs. S1c and S3c), when high-quality signals of the 180 181 fundamental mode surface wave are observed in the filtered ANCs (e.g., Figs. S1a-b and S3a-b). This suggests that the initial phase φ_F is zero, i.e., $\varphi_{ij,k}(\omega) = -\omega \cdot T_{ij}^F$ (Equation 182 4b), for ANCs of the linear RR array filtered at 0.8 s. We also verified that the initial 183

phase φ_F is zero for all periods (from 0.3 s to 1.6 s) analyzed in this study, by estimating the systematic phase difference between interferograms in the inter- and outer-source zones computed for the example station pair RR10 and RR40 (not shown here).

187 Therefore, we can simply stack the interferograms $\tilde{I}_{i,j}(\omega; k) = A_{ij,k} \cdot e^{i \cdot \varphi_{ij,k}}$, defined 188 in equation 4a, over all available station k to enhance the contribution from the 189 fundamental mode surface wave, $\tilde{I}_{i,j}^F(\omega; k)$, which has a phase that is independent of k 190 (Equation 4b with $\varphi_F = 0$). The other component, $\tilde{I}_{i,j}^O(\omega; k)$, that involves contributions 191 from body waves and background noise is suppressed through stacking, as its phase 192 varies significantly with k (Section 3.1). The denoised waveform obtained through linear 193 stacking is given by:

$$\tilde{C}_{i_j}^{3}(\omega) = \sum_{k=1}^{N} e^{i \cdot \varphi_{ij_k}} \sqrt{A_{ij_k}} / N.$$
(5a)

Here, *N* is the number of stations in the 1-D linear array. We take the square root of the amplitude spectrum A_{ij_k} in Equation 5a to suppress the effect of source spectra multiplication introduced in the three-station interferometry (Equation 4a). This is based on the assumption that amplitude spectra of the filtered ANCs are similar for all station pairs, i.e., $A_{ij_k} = A_{G_ik} \cdot A_{G_jk} \approx \hat{A}_G^2$.

In this study, we perform phase weighted stacking rather than linear stacking, as phase weighted stacking is more efficient in suppressing incoherent patterns and has negligible effects on phase measurements (e.g., Fig. S4; Schimmel and Paulssen, 1997). Let $\hat{C}_{i_{-}j}^{3}(t; \omega_{c})$ be the waveform in time domain denoised through linear stacking using ANCs filtered at the angular frequency ω_{c} , the corresponding waveform denoised through phase weighted stacking in the time domain is given by

$$C_{i_j}^3(t;\omega_c) = W_{i_j}(t;\omega_c) \cdot \hat{C}_{i_j}^3(t;\omega_c),$$
(5b)

where $W_{i_j}(t; \omega_c)$ is a weighting function that indicates the phase coherence of interferograms in the time domain (e.g., Fig. 3c) averaged over all k values.

207 3.3 Results

Figure 4a shows the comparison between the TT component ANC (black) of the example station pair RR10 and RR40 filtered at 0.3 s and the corresponding denoised 210 waveforms, computed through linear stacking (blue; Equation 5a) and phase weighted 211 stacking (bottom red curve; Equation 5b). Although coda waves in the denoised 212 waveform $C_{i,i}^{3}(t; \omega_c)$ are greatly suppressed, the fundamental mode surface wave is still not the dominant signal (e.g., large-amplitude wavelets before the surface wave). This is 213 214 likely due to the poor SNR of surface waves in the ANCs filtered at 0.3 s (Figs. 3a-b). 215 Therefore, we further enhance the surface wave signal by repeating the denoising process 216 (Fig. 2a): first self-normalize the output wavefield of the current iteration, and then use 217 the normalized wavefield as the input for the next iteration. The number of iterations is 218 determined so that the difference between input and output wavefields of the last iteration 219 is negligible (Fig. 2a).

We use symbol $C_{i,j}^{2+n}(t;\omega_c)$ to represent the waveform of station pair *i* and *j*, after 220 applying $n \geq 1$ iterations of the denoising process described in section 3.2 to ANCs 221 222 filtered at angular frequency ω_c . Figure 4a suggests that four more iterations (red 223 waveforms; n = 5) are needed to obtain surface waves with sufficient quality from the ANCs filtered at 0.3 s, i.e., the difference between the input, $C_{i_j}^6(t; \omega_c)$, and output, 224 $C_{i_j}^7(t;\omega_c)$, waveforms is visually negligible. Although the SNR gradually increases in 225 226 waveforms from bottom to top (Fig. 4a), the surface wave signal is always coherent and 227 aligned in phase. As the SNR is much higher for surface waves in ANCs filtered at low 228 frequencies (> 0.6 s), the number of iterations used to extract good quality surface waves is smaller (n = 3; e.g., 0.8 s in Fig. S1d). 229

Comparison between the TT component ANC data $G_{i,j}(t; \omega_c)$ filtered at 0.3 s and the 230 denoised waveforms $C_{i,j}^7(t; \omega_c)$ is illustrated in Figures 4b-c for all station pairs. 231 232 Although coherent fundamental mode surface waves are seen propagating at a group 233 velocity slightly slower than 0.5 km/s in the filtered ANC data (Fig. 4b), wavelets with 234 large amplitudes are observed before and after the surface wave signals (e.g., black 235 waveform in Fig. 4b). The large amplitude waves arriving prior to the surface wave are likely related to the tapering window at short interstation distances (e.g., < 1 km), as the 236 237 apparent moveout velocity is 2 km/s (the upper limit velocity of our tapering window; 238 Section 2). At long interstation distances, the precursor wavelets have an apparent 239 moveout velocity of ~ 1 km/s and may represent the contribution from body wave energy.

These waves that we want to suppress in the denoising process are sometimes even larger than the surface wave signal in the filtered ANCs (e.g., black waveform in Fig. 4b).

242 After five iterations of the denoising process for ANCs filtered at 0.3 s, surface wave 243 signals are well preserved whereas the other signals are greatly suppressed (Fig. 4c). The 244 amplitude spectra averaged over all station pairs for data before (in black) and after (in 245 red) denoising are demonstrated in Figure 4d. The observation of a smoother mean 246 amplitude spectrum after denoising (red curve in Fig. 4d) suggests that the difference 247 between amplitude spectra of every two station pairs is much smaller. Since SNR of the 248 fundamental mode surface wave is much larger in ANCs filtered at 0.8 s (Figs. S1a-b and 249 S3a-b), the difference between waveforms before and after denoising is still noticeable 250 but much smaller (Figs. S1e-g and S3e-g). As the fundamental mode surface wave is the 251 dominant signal after denoising (e.g., Figs. 4c and S1f-S3f), all results in section 4 refer 252 to the fundamental mode surface wave.

253 **4. Surface Wave Tomography**

254 In this section, we use waveforms of TT and ZZ components denoised at each period 255 (e.g., Figs. 4c and S1f-S3f) to infer phase velocity structures of Love and Rayleigh waves 256 beneath the array, respectively. Following the flow chart shown in Fig. 2b, we first 257 determine cycle-skipped phase travel times of surface waves propagating between all 258 available station pairs at each period (e.g., Fig. 5a) in the frequency domain, which is 259 much simpler than measuring in the time domain but requires high SNR (Section 4.1). 260 Second, we infer phase velocity structures beneath the linear RR array, using travel time 261 measurements after cycle-skipping correction from section 4.1, via the eikonal equation 262 in section 4.2 (e.g., Fig. 5b). The aim of this section is to demonstrate that robust surface 263 wave phase velocity models can be resolved from the denoised waveforms.

4.1 Determination of phase travel time

Frequency time analysis (FTA) is widely used in previous studies to determine phase travel time of surface wave in ANC (e.g., Bensen et al., 2007; Lin et al., 2008; Qiu et al., 2019). First, Gaussian narrow bandpass filters centered on a series of consecutive frequencies are applied to the ANC, then the phase travel time dispersion is measured using the envelope and phase functions of the filtered ANC in the time domain. The advantage of FTA is that reliable phase travel times can still be extracted when SNR is low at high frequencies. However, ad hoc criteria and thresholds are required to automate the FTA. Additional details on the FTA method can be found in section 3 of Qiu et al. (2019). Since our goal is to verify that the signals after denoising are representative of surface waves and high SNR is achieved for all frequencies, we thus measure phase travel times from the denoised waveforms in the frequency domain, which is much simpler than the FTA method.

277 Although surface wave is the dominant signal in the denoised waveform, we still 278 observe waves with small amplitudes before and after the surface wave (e.g., black 279 waveform in Fig. 4c). This is because we can only suppress rather than remove signals 280 that are not surface waves. Here, we apply a tapering window (e.g., black dashed lines in 281 Figs. S5d-S8d) centered on the surface wave to further remove these background signals. 282 Width of the tapering window is set to four times the dominant period of the array-mean 283 amplitude spectrum (e.g., red curve in Fig. 4d), whereas the center is determined by the 284 average phase and group velocities of the array (e.g., red and cyan stars in Figs. S5b). 285 Details of the tapering window can be found in Text S1.

Assuming these background signals are negligible after denoising and tapering (e.g., blue waveform in Fig. 5a), combining equation 2b and $\varphi_F = 0$, we have

$$\tilde{C}_{i_j}^{DT}(\omega;\omega_c) = A_{F_ij}(\omega;\omega_c)e^{-i\cdot\omega T_{i_j}^F(\omega)},$$
(6)

where $\tilde{C}_{i_j}^{DT}(\omega; \omega_c)$ is the spectrum of the tapered waveform that is denoised at the angular frequency ω_c for station pair *i* and *j*. Equation 6 suggests that we can extract cycleskipped phase travel time from the phase spectrum of the tapered waveform. Therefore, we measure the wrapped phase (i.e., between -2π and 0) of the spectrum $\tilde{C}_{i_j}^{DT}(\omega; \omega_c)$ at the peak frequency f_{max} (3.2 Hz; e.g., dashed line in Fig. 4d) of the array-mean amplitude spectrum, where the array-mean SNR of the surface wave is the highest. Then, the cycleskipped phase travel time is computed as the wrapped phase divided by $-2\pi f_{\text{max}}$.

Figure 5a shows the cycle-skipped phase travel time (white circles) measured for surface waves filtered at 0.3 s from station pairs associated with a common virtual source RR10 (y-axis of 0 km). To obtain the actual phase travel time, we perform a simple cycle-skipping correction as follows: (1) Similar to Figure 5a, we first extract all the cycle-skipped phase travel times for
surface waves of a virtual shot gather and arrange them as a function of the location
to the virtual source.

- 302 (2) We perform cycle-skipping correction for surface waves traveling NE (toward RR47)
 303 and SW (toward RR01) separately.
- 304 (3) For surface waves traveling in the same direction, the principle of the cycle-skipping
 305 correction is to ensure that the travel time of any virtual receiver is larger than those
 306 of receivers that are closer to the virtual source after the correction.
- 307 (4) In practice, we examine measurements T_i and T_{i+1} of every two adjacent virtual 308 receivers with the *i*-th station being closer to the virtual source. If $T_i \ge T_{i+1}$, we use T_i 309 as the reference and add N/f_{max} (N is an integer) to T_{i+1} so that $T_{i+1}+N/f_{\text{max}} > T_i \ge$ 310 $T_{i+1}+(N-1)/f_{\text{max}}$. The correction is performed for closer-to-source pairs first.

Travel times, for the virtual shot gather of RR10, after the correction are illustrated as red stars in Figure 5a. We note that a more sophisticated cycle-skipping correction (e.g., using phase velocity structure inferred at a longer period as the reference) is needed when station spacing is larger than one wavelength.

315 **4.2 1-D eikonal tomography**

We use the eikonal equation to derive phase velocity structures using travel time measurements of all station pairs in the linear RR array (Section 4.1). First, we project all stations to the straight line connecting RR01 and RR47 (cyan dashed line in Fig. 1a). Second, travel time measurements associated with each virtual source *i* at the target frequency f_{max} are extracted and interpolated (e.g., black curve in Fig. 5a) with a regular grid size of Δ =50 m. Since variations in topography (Fig. 2b of Qin et al., 2020) have a negligible effect (< 0.5%) on the results, the eikonal tomography can be simplified as:

$$\tilde{v}_i(x; f_{\max}) = 2 \cdot \Delta / [T_i(x + \Delta; f_{\max}) - T_i(x - \Delta; f_{\max})],$$
(7)

where $\tilde{v}_i(x; f_{\text{max}})$ and $T_i(x; f_{\text{max}})$ are the local phase velocity and interpolated phase travel time, respectively, of the grid cell at location *x*. Since the local phase velocity \tilde{v}_i only varies with the grid cell location, it is independent of virtual source *i*. Thus, we can average the 1-D phase velocity profiles resolved from all available virtual sources at the same frequency f_{max} to achieve a more reliable phase velocity model:

$$\bar{v}(x; f_{\max}) = \sum_{i=1}^{N_x} \tilde{v}_i(x; f_{\max}) / N_x, \qquad (8a)$$

328 and estimate the corresponding uncertainty as the standard deviation:

$$\delta(x; f_{\max}) = \sqrt{\sum_{i=1}^{N_x} [\tilde{v}_i(x; f_{\max}) - \bar{v}(x; f_{\max})]^2 / N_x},$$
(8b)

329 where N_x is the number of virtual sources available for stacking at location x.

330 In surface wave studies, phase velocities derived at near-virtual-source grid cells are 331 often excluded to satisfy the far-field approximation (e.g., Bensen et al., 2007). The size 332 of the exclusion zone is usually multiples of the analyzed wavelength (e.g., one 333 wavelength in Wang et al., 2019). Here, however, we set an exclusion zone with a fixed 334 size of 100 m, i.e., discard phase velocities derived at the four grid cells closest to the 335 virtual source. Figure 5b shows the 1-D phase velocity profile, in white dots, derived 336 using measurements associated with the virtual source RR10 (black curve in Fig. 5a) for 337 Love waves at 3.2 Hz (~0.3 s), whereas phase velocity profiles resolved from all virtual 338 sources are illustrated in gray curves and as the colormap. The average phase velocity 339 and uncertainty profiles are calculated via equation 8 and demonstrated as red stars and 340 error bars, respectively, in Figure 5b. Results for Love waves at a lower frequency (~0.8 341 s) and those of Rayleigh waves are shown in Figures S9 and S10-S11, respectively.

342 Figure 6 shows phase velocity models resolved at periods ranging from 0.3 s to 1.6 s 343 for Love waves (Fig. 6a) and 0.3 s to 1.3 s for Rayleigh waves (Fig. 6b), together with the 344 corresponding uncertainty estimations (Figs. 6c-d). The period range in the plot is 345 determined so that the resolved maximum uncertainty is smaller than 0.1 km/s. In general, the uncertainties are smaller than 0.03 km/s for both Rayleigh and Love waves at 346 347 all analyzed periods, indicating the resolved phase velocity structures are robust and 348 reliable. This also justifies our choice of an exclusion zone with a 100-m-radius, as one 349 wavelength at low frequency (e.g., ~900 m for Rayleigh wave at ~0.8 s; Fig. S11b) is 350 much larger than 100 m. We note that results of Love waves at ~0.4 s are excluded here 351 due to anomalous large uncertainties (Fig. S12b). This is due to the observation of non-352 negligible signals, likely representing higher-mode surface waves, traveling at a different 353 phase speed compared to that of the fundamental mode Love waves (Figs. S13-S14). We note that reliable phase travel times can still be measured for the fundamental mode Love waves at ~ 0.4 s if FTA is used.

356 Phase velocity models of both Love and Rayleigh waves show a ~500- to 1000-m-357 wide low-velocity zone at low frequencies (e.g., > 0.8 s) that gradually narrows with the 358 period. Combined with the fact that phase velocity at lower frequency is more sensitive to 359 structures at greater depth, this observation likely indicates a flower-shaped (i.e., width 360 decreases with depth) fault damage zone beneath the linear RR array. We also see several \sim 100-m-wide narrow zones, that are close to the mapped fault surface traces (black 361 362 dashed lines in Figs. 6a-b), with extremely low phase velocities (< 500 m/s) at high 363 frequencies (e.g., 0.3-0.6 s). However, the shape and location of these low-velocity zones 364 are different between Figures 6a and 6b. In addition, complicated patterns of low-velocity 365 anomalies at various scales in both phase velocity models (Figs. 6a and 6b) suggest high 366 degrees of heterogeneity in shear wave velocity structures beneath the linear RR array. 367 Structure patterns that are inconsistent between models of Love and Rayleigh waves may 368 indicate the existence of radial anisotropy or complicated structures of Vp/Vs ratio.

369 **5. Discussion**

370 We compare the Rayleigh wave phase velocity models derived from this study and 371 Wang et al. (2019) in the overlapping period (0.3-0.8 s) and spatial (RR01-RR47) ranges 372 (Fig. 7). The same ANC dataset of ZZ component is utilized in Wang et al. (2019) to 373 derive the phase velocity model (Fig. 7b) for Rayleigh waves. In their study, double 374 beamforming technique is used, where the local phase velocity is obtained through grid 375 search: first sum all ANCs of two sub-arrays (three nearby stations) via slant-stacking 376 with different velocity values, then determine the local phase velocity of each sub-array 377 based on the maximum amplitude of the envelope function for the stacked waveform. 378 Through slant-stacking, Wang et al. (2019) was able to enhance the low SNR surface 379 waves in ANCs at high frequencies and generate a robust phase velocity model.

In general, consistent velocity values and structural patterns are observed in both phase velocity models, such as an ultra-low velocity (< 0.4 km/s) zone on the NE side of the middle fault surface trace (F2 in Figs. 7a-b). This indicates that the enhanced signals in ANCs after denoising are surface waves. It is also important to note that our Rayleigh wave phase velocity model covers a wider period range (0.3-1.3 s; Fig. 6b) compared to that of Wang et al. (2019) (0.3-0.8 s). There are two reasons for that: first, we only exclude data associated with station pairs that are less than 100 m apart, whereas the size of exclusion zone is one wavelength in Wang et al. (2019); second, phase velocities inferred from slant-stacking have much larger uncertainties than those derived from travel times, as the peak of envelope function is much sensitive to noise than phase function (e.g., Section 3.3 of Qiu et al., 2019), at longer periods.

391 The uncertainties estimated from this study (Fig. 7c) are significantly smaller than 392 those of Wang et al. (2019) (Fig. 7d), suggesting the phase velocity structure is better 393 constrained in this study, particularly near the fault surface traces (e.g., NE of F3 and 394 between F1 and F2 in Fig. 7) and at high frequencies (e.g., < 0.6s). The larger uncertainty 395 values observed in Wang et al. (2019) are likely due to the low SNR of surface waves in 396 ANCs at high frequencies and should be significantly reduced if the denoised waveforms 397 (e.g., Fig. 4c) are used. By comparing phase velocity models resolved from this study and 398 Wang et al. (2019), we demonstrate that fundamental mode surface waves in ANCs are 399 successfully enhanced through the proposed denoising method. Besides, phase velocity 400 structures are better constrained (i.e., much smaller uncertainties) using the denoised 401 waveforms in this study than those inferred from the raw ANCs in Wang et al. (2019).

402 The derivations presented in section 3.1 are based on two assumptions: higher-mode 403 surface waves are negligible, and stations are aligned in a straight line. After performing 404 the denoising process on ANCs of the linear RR array filtered at different periods 405 between 0.3 s and 1.6 s for TT and ZZ components, we do not observe higher-mode 406 surface waves except for TT component data filtered at 0.4 s (Figs. S13-S14). Even when 407 higher-mode surface waves are present, we further demonstrate that the denoising process 408 based on three-station interferometry will only enhance surface waves of each mode 409 separately in Text S2. Regarding the effect of array geometry, we do not investigate in 410 detail the bias in surface wave travel times measured from the denoised waveforms when 411 stations are not perfectly aligned in a straight line. Assuming the velocity structure is 412 homogenous perpendicular to the strike direction of the array, the station-configuration 413 error is proportional to the difference between interstation distances calculated using 414 station locations before and after projecting the array to a straight line (e.g., Fig. 1a).

415 For a rough estimation, the mean and maximum differences between interstation 416 distances of the linear RR array calculated before and after the linear projection (Section 417 2) are $\sim 0.1\%$ and 1%, respectively. The uncertainties estimated from eikonal tomography 418 (Figs. 6c-d) suggest $\sim 0.5-1\%$ and $\sim 3\%$ for the mean and maximum perturbations in the 419 resolved phase velocities, which is comparable and larger than the bias in array geometry. 420 Therefore, we conclude that this denoising method is still robust when the station-421 configuration error is less than the allowable uncertainty of the resulting phase velocity 422 model (e.g., mean and maximum of 1% and 3% in this study). Surface wave denoising of 423 ANCs computed from arrays in a 2-D configuration (e.g., include all stations in the RR 424 array) will be the subject of a future study.

425 Although we only illustrate the denoising method using ANCs computed following 426 the preprocessing steps described in Wang et al. (2019), it can be applied to any noisy 427 wavefield that consists of surface waves propagating within a linear array. Therefore, this 428 method can be used as a routine postprocessing procedure for improving the quality of 429 surface waves in ANCs of a linear array with any type of preprocessing steps (e.g., coda wave C^3 or C^N ; Froment et al., 2011). After denoising, we may be able to perform surface 430 431 wave studies on data from linear array deployments that were previously considered to be 432 too short in recording duration or too "noisy" (i.e., deployed during high seismic activity 433 period) to retrieve good quality surface wave signals from ANCs.

434 We note that Lin et al. (2009) used the term "apparent phase velocity" for the phase 435 velocity resolved from the eikonal equation (Equation 8a). They concluded that the 436 apparent phase velocity is accurate when the wave frequency is sufficiently high or the 437 wave amplitude is varying smoothly in space. Otherwise, a correction term must be 438 added (Lin & Ritzwoller, 2011). Considering the linear RR array is dense (average 439 station spacing of ~ 30 m) and deployed crossing fault damage zones that can 440 significantly amplify seismic motions (Qin et al., 2020), both the "high frequency" and 441 "smooth amplitude variation" approximations are likely violated for phase velocity 442 models derived in section 4.2. Therefore, we propose performing travel-time based full 443 waveform tomography (e.g., Zhang et al., 2018) on the denoised surface waves, rather 444 than inverting the "apparent" phase velocity models (Figs. 6a-b), to infer shear wave 445 velocity structures beneath the linear RR array as the subject of a future study.

446 In the present study, we only show the denoised ANCs filtered at frequencies up to ~ 3 447 Hz (Figs. 4 and S2d-g) by requiring a minimum wavelength of 150 m (i.e., the maximum 448 station spacing) for eikonal tomography (Section 4). However, this denoising method can 449 be applied to ANCs at even higher frequencies when surface wave signals are present 450 (e.g., ~5 Hz in Figs. S15-S16). Using high-quality surface waves denoised from ANCs of 451 linear arrays crossing major fault zones for a wide range of frequencies (e.g., 0.5-5 Hz in 452 this study; 2-40 Hz in Zigone et al., 2019), shear wave velocity model that extends to 453 both shallow (top few tens of meters) and deep (top few kilometers) structures can be 454 derived. Such fault zone velocity model with unprecedented high-resolution will 455 complement the qualitative model inferred from traditional fault zone analyses (e.g., Qin 456 et al., 2018, 2020; Qiu et al., 2017, 2020; Share et al., 2017, 2019). An integration of both 457 quantitative and qualitative fault zone models can have significant implications for 458 seismic hazard evaluations (e.g., Ben-Zion & Shi, 2005; Spudich & Olsen, 2001) and 459 long-term behavior of the fault (e.g., Thakur et al., 2020).

460 **6.** Conclusions

We demonstrate the effectiveness and robustness of the three-station-interferometrybased surface wave denoising method using ANCs of the linear RR array, particularly at high frequencies (e.g., > 2 Hz). The proposed surface wave denoising method can be applied to a wide range of topics in the future using data recorded by 1-D linear arrays:

- 465 1. Reduce the minimum duration of ambient noise recording and preprocessing steps466 needed to achieve high-quality surface waves from ANCs.
- 2. Obtain good quality surface waves from ANCs computed for fault zone linear arrays
 deployed right after a major earthquake (e.g., fault zone arrays in Ridgecrest area;
 Catchings et al., 2020) without applying any sophisticated preprocessing steps.
- 470 3. Provide high-quality surface wave signals both at high (> 2 Hz) and low frequencies
 471 (< 1 Hz) for better constraints of shallow (top 10s to 100s of meters) materials
 472 through full-waveform surface wave tomography.
- 473 4. Investigate the initial phase φ_F (Equation 2b) for a wider frequency band and arrays 474 at different locations, and its relation to the ambient noise source distribution

475

Δ	7	6
-+	1	υ

477 Acknowledgements

We thank Fan-Chi Lin and Elizabeth Berg for providing the ambient noise cross correlation data and Rayleigh wave phase velocity model of Wang et al. (2019), and Nori Nakata and Benxin Chi for useful discussions. Continuous seismic recordings of the RR array are available through the International Federation of Digital Seismography Networks (Allam, 2015; https://www.fdsn.org/networks/detail/9K_2015/). This work was supported by Rice University and the National Science Foundation (Award EAR-1251667).

485

486

487 **References**

Allam, A.A. (2015), San Jacinto damage zone imaging arrays. International Federation of
Digital Seismograph Networks. Dataset/Seismic Network. Doi:

490 https://doi.org/10.7914/SN/9K_2015

- 491 Ben-Zion, Y., & Shi, Z. (2005). Dynamic rupture on a material interface with
- 492 spontaneous generation of plastic strain in the bulk. *Earth and Planetary Science*493 *Letters*. https://doi.org/10.1016/j.epsl.2005.03.025
- 494 Bensen, G. D., Ritzwoller, M. H., Barmin, M. P., Levshin, A. L., Lin, F.-C., Moschetti,
- 495 M. P., et al. (2007). Processing seismic ambient noise data to obtain reliable broad-
- 496 band surface wave dispersion measurements. Geophysical Journal International,
- 497 169(3), 1239–1260. https://doi.org/10.1111/j.1365-246X.2007.03374.x
- Catchings, R. D., Goldman, M. R., Steidl, J. H., Chan, J. H., Allam, A. A., Criley, C. J.,
 et al. (2020). Nodal Seismograph Recordings of the 2019 Ridgecrest Earthquake
- 500 Sequence. Seismological Research Letters. https://doi.org/10.1785/0220200203
- 501 Curtis, A., Nicolson, H., Halliday, D., Trampert, J., & Baptie, B. (2009). Virtual
- seismometers in the subsurface of the Earth from seismic interferometry. Nature
 Geoscience. https://doi.org/10.1038/ngeo615
- 504 Froment, B., Campillo, M., & Roux, P. (2011). Reconstructing the Green's function
- 505 through iteration of correlations. Comptes Rendus Geoscience.
- 506 https://doi.org/10.1016/j.crte.2011.03.001

507	Lin, FC., & Ritzwoller, M. H. (2011). Apparent anisotropy in inhomogeneous isotropic
508	media. Geophysical Journal International. https://doi.org/10.1111/j.1365-
509	246X.2011.05100.x
510	Lin, FC., Moschetti, M. P., & Ritzwoller, M. H. (2008). Surface wave tomography of
511	the western United States from ambient seismic noise: Rayleigh and Love wave
512	phase velocity maps. Geophysical Journal International.
513	https://doi.org/10.1111/j.1365-246X.2008.03720.x
514	Lin, FC., Ritzwoller, M. H., & Snieder, R. (2009). Eikonal tomography: Surface wave
515	tomography by phase front tracking across a regional broad-band seismic array.
516	Geophysical Journal International. https://doi.org/10.1111/j.1365-
517	246X.2009.04105.x
518	Qin, L., Ben-Zion, Y., Qiu, H., Share, P. E., Ross, Z. E., & Vernon, F. L. (2018). Internal
519	structure of the san jacinto fault zone in the trifurcation area southeast of anza,
520	california, from data of dense seismic arrays. Geophysical Journal International,
521	213(1), 98–114. https://doi.org/10.1093/gji/ggx540
522	Qin, L., Share, PE., Qiu, H., Allam, A. A., Vernon, F. L., & Ben-Zion, Y. (2020).
523	Internal structure of the San Jacinto fault zone at the Ramona Reservation, north of
524	Anza, California, from dense array seismic data. Geophysical Journal International.
525	https://doi.org/10.1093/gji/ggaa482
526	Qiu, H., Ben-Zion, Y., Ross, Z. E., Share, P. E., & Vernon, F. L. (2017). Internal
527	structure of the San Jacinto fault zone at Jackass Flat from data recorded by a dense
528	linear array. Geophysical Journal International, 209(3), 1369-1388.
529	https://doi.org/10.1093/gji/ggx096
530	Qiu, H., Lin, FC., & Ben-Zion, Y. (2019). Eikonal tomography of the Southern
531	California plate boundary region. Journal of Geophysical Research: Solid Earth,
532	2019JB017806. https://doi.org/10.1029/2019JB017806
533	Qiu, H., Ben-Zion, Y., Catchings, R., Goldman, M., Allam, A., & Steidl, J. (2020).
534	Detailed seismic imaging of the Mw 7.1 Ridgecrest earthquake rupture zone from
535	data recorded by dense linear arrays. https://doi.org/10.1002/essoar.10504611.1
536	Schimmel, M., & Paulssen, H. (1997). Noise reduction and detection of weak, coherent
537	signals through phase-weighted stacks. Geophysical Journal International.

- 538 https://doi.org/10.1111/j.1365-246X.1997.tb05664.x
- 539 Share, P. E., Ben-Zion, Y., Ross, Z. E., Qiu, H., & Vernon, F. L. (2017). Internal
- 540 structure of the San Jacinto fault zone at Blackburn Saddle from seismic data of a
- 541 linear array. *Geophysical Journal International*, *210*(2), 819–832.
- 542 https://doi.org/10.1093/gji/ggx191
- 543 Share, P. E., Allam, A. A., Ben-Zion, Y., Lin, F.-C., & Vernon, F. L. (2019). Structural
- 544 Properties of the San Jacinto Fault Zone at Blackburn Saddle from Seismic Data of a
- 545 Dense Linear Array. *Pure and Applied Geophysics*, *176*(3), 1169–1191.
- 546 https://doi.org/10.1007/s00024-018-1988-5
- 547 Snieder, R. (2004). Extracting the Green's function from the correlation of coda waves: A
 548 derivation based on stationary phase. Physical Review E Statistical Physics,
- 549 Plasmas, Fluids, and Related Interdisciplinary Topics.
- 550 https://doi.org/10.1103/PhysRevE.69.046610
- Spudich, P., & Olsen, K. B. (2001). Fault zone amplified waves as a possible seismic
 hazard along the Calaveras fault in central California. *Geophysical Research Letters*.
 https://doi.org/10.1029/2000GL011902
- Thakur, P., Huang, Y., & Kaneko, Y. (2020). Effects of Low-Velocity Fault Damage
 Zones on Long-Term Earthquake Behaviors on Mature Strike-Slip Faults. *Journal of Geophysical Research: Solid Earth*. https://doi.org/10.1029/2020JB019587
- 557 Wang, Y., Allam, A. A., & Lin, F.-C. (2019). Imaging the Fault Damage Zone of the San
- 558Jacinto Fault Near Anza With Ambient Noise Tomography Using a Dense Nodal
- Array. Geophysical Research Letters. https://doi.org/10.1029/2019GL084835
- Zhang, C., Yao, H., Liu, Q., Zhang, P., Yuan, Y. O., Feng, J., & Fang, L. (2018). Linear
 Array Ambient Noise Adjoint Tomography Reveals Intense Crust-Mantle
- Interactions in North China Craton. *Journal of Geophysical Research: Solid Earth.*https://doi.org/10.1002/2017JB015019
- Zhang, S., Feng, L., & Ritzwoller, M. H. (2020). Three-station interferometry and
 tomography: coda versus direct waves. Geophysical Journal International.
 https://doi.org/10.1093/gji/ggaa046
- 567 Zigone, D., Ben-Zion, Y., Lehujeur, M., Campillo, M., Hillers, G., & Vernon, F. L.
- 568 (2019). Imaging subsurface structures in the San Jacinto fault zone with high-

frequency noise recorded by dense linear arrays. *Geophysical Journal International*,
217(2), 879–893. https://doi.org/10.1093/gji/ggz069

- 571
- 572

573 Figure captions

574

575 Figure 1. (a) Google map for the RR array (colored balloons) deployment that crosses 576 surface traces of the San Jacinto fault (colored lines). The stations colored in white are 577 not analyzed in this study, whereas the green balloons denote three sensors closest to 578 each corresponding fault surface trace. Surface wave denoising procedure is 579 demonstrated for an example station pair (red balloons). (b) Zoom out map of the San 580 Jacinto fault zone. The background colors indicate topography. The red star and blue 581 square denote locations of the RR array and the town of Anza. The black lines illustrate surface traces of major faults in this area. EF – Elsinore Fault; SAF – San Andreas Fault; 582 583 SJF – San Jacinto Fault. (c) Ambient noise cross correlations at TT component of all 584 station pairs for the sub-array RR01-RR47. The cross correlations are arranged according 585 to interstation distance with red and blue colors representing positive and negative values. 586 All the waveforms are first tapered using a velocity range of 2 km/s and 0.1 km/s (dashed 587 lines), and then bandpass filtered between 0.2 and 10 Hz. (d) Same as (c) for the ZZ 588 component.

589 Figure 2. (a) Flow chart of the surface wave denoising and imaging procedure 590 developed in this study. The dashed box outlines the part of the diagram that performs 591 surface wave tomography. The workflow adopted in this study for surface wave 592 tomography is shown in (b). ANC – Ambient Noise Cross-correlation.

Figure 3. (a) Ambient noise cross correlations (ANCs) of TT component narrow bandpass filtered at 0.3 s associated with the virtual source RR10 (red star). Waveforms are arranged by the station number of the virtual receiver. (b) Same as (a) for virtual source RR40. Waveforms in black and blue represent ANCs of virtual receivers in the outer- and inter-source zones, respectively, while the red waveform denotes the ANCs of the station pair of the two virtual sources RR10 and RR40 (red balloons in Fig. 1a). (c) Interferograms calculated via three-station interferometry (Equation 4a). 600 Figure 4. (a) Comparison between the ANC (red waveforms in Figs. 3a-b), linear 601 stacked (LS) C3 (Equation 5a), and phase weighted stacked (PWS) C3 – C7 (Equation 602 5b) of the example station pair RR10 and RR40 at TT component. (b) ANCs of TT 603 component (Fig. 1c) narrow bandpass filtered at 0.3 s with red and blue colors 604 representing positive and negative values. The three white dashed lines illustrate moveout 605 velocities of 2 km/s, 0.5 km/s, and 0.1 km/s. The black waveform denotes ANC of the 606 example pair RR10 and RR40, i.e. bottom black waveform in (a). (c) PWS C7 of all 607 station pairs at 0.3 s. The white dashed line denotes a moveout velocity of 0.5 km/s. The 608 black waveform denotes PWS C7 of the example station pair RR10 and RR40, i.e. top 609 red waveform in (a). (d) Array-mean amplitude spectra of waveforms (b) before and (c) 610 after denoising are depicted in black and red, respectively. The peak frequency of the red 611 amplitude spectrum is illustrated by the red dashed line and labeled in the top right of the 612 panel (c).

613 Figure 5. (a) Love waves associated with the virtual source RR10 after denoising at 614 0.3 s and tapering. Red and blue colors represent positive and negative values, 615 respectively. White circles denote cycle-skipped phase travel time measurements, 616 whereas red stars indicate travel times after cycle skipping correction (Section 4.1). The 617 black curve illustrates the corrected phase travel time after interpolation using a grid size 618 of 50 m. (b) Phase velocity profiles resolved for Love waves at 0.3 s underneath the 1-D 619 linear array. White circles depict the 1-D phase velocity profile derived via eikonal 620 equation using travel time measurements shown as the black curve in (a). The colormap 621 illustrates phase velocity profiles obtained using different stations as virtual sources (x-622 axis), with white space representing the near-source exclusion zone. The phase velocity 623 profile averaged over results of all virtual sources are depicted as red stars, with the error 624 bar representing the corresponding standard deviation. The black vertical dashed line 625 denotes the array-mean phase velocity estimated at 0.3 s (red star in Fig. S4b).

Figure 6. Phase velocity dispersion profiles for (a) Love and (b) Rayleigh waves beneath the RR array. The vertical dashed lines denote locations of the mapped fault surface traces (Fig. 1a), while the horizontal arrow outlines the group of stations that record fault zone trapped waves (Qin et al., 2020). Uncertainties of the resolved phase

22

velocity profiles are shown in (c) and (d) for Love and Rayleigh waves, respectively.Phase velocity and uncertainty profiles for Love wave at 0.4 s are excluded here due to

632 large uncertainties (Figs. S12).

Figure 7. Comparison of (a) phase velocity and (b) uncertainty profiles of this study and those, (b) and (d), of Wang et al. (2019) in the overlapping period range. The white space indicates the area not covered by the final model. Note, the definition of phase velocity uncertainty is different in this study (standard deviation) compared to that of Wang et al. (2019) (standard deviation of the mean). A conversion is made to transform the uncertainties shown in Wang et al. (2019) into the uncertainty profiles of (d). Figure 1.



Figure 2.



Figure 3.



Figure 4.



Figure 5.



Figure 6.



Figure 7.





Journal of Geophysical Research: Solid Earth

Supporting Information for

Denoising surface waves extracted from ambient noise using three-station interferometry: Methodology and Application to 1-D linear array

Hongrui Qiu¹, Fenglin Niu¹, and Lei Qin²

¹Department of Earth, Environmental and Planetary Sciences, Rice University, Houston, TX, USA ²Department of Earth Sciences, University of Southern California, Los Angeles, CA, USA

Contents of this file

Text S1 to S2 Figures S1 to S16

Introduction

This supporting information provides additional details that are complementary to the main article.

Text S1.

Determination of the tapering window

In order to select the accurate time window that outlines the surface wave signal (e.g., black dashed lines in Fig. S5d), we first estimate the array-mean phase and group velocity dispersions (e.g., Fig. S5b) by stacking the denoised waveforms (Fig. S5c) shifted with respect to different moveout velocities (Fig. S5a). Figure S5a shows the procedure of grid searching the array-mean phase and group velocities at 0.3 s:

- 1. For each analyzed moveout velocity V_0 (x-axis of Fig. 5a), we shift the denoised waveform of each station pair with interstation distance of Δ towards the negative correlation time direction with $dt = \Delta/V_0$.
- 2. We directly stack the denoised waveforms after the alignment and measure the maximum amplitude of the stacked waveform (red dots in Fig. 5a). The arraymean phase velocity is then determined as the velocity V_0 when the grid search curve (red dashed curve in Fig. 5a) reaches the maximum.
- 3. To estimate the array-mean group velocity, we stack envelope functions of the shifted waveforms. The result for grid search of array-mean group velocity is shown in cyan color.

After obtaining the array-mean phase and group velocities at each period, \overline{V}_{ph} and \overline{V}_{gp} , we set the width of the tapering window as four times the dominate period T_{max} of the array-mean amplitude spectrum (e.g., Fig. 4d). The center of the window is set to $N \cdot T_{max}$, where N is a integer that satisfies $N \cdot T_{max} \ge \Delta \cdot (1/\overline{V}_{ph} - 1/\overline{V}_{gp}) \ge (N-1) \cdot T_{max}$, for the station pair with interstation distance of Δ .

Text S2.

Three-station interferometry with the presence of higher-mode surface waves

The aim of this section is to investigate the derivation for denoising using threestation interferometry if higher-mode surface waves are present. For simplicity, we first only add the first overtone surface wave, $\tilde{M}_{i,i}(\omega)$, to equation 2b:

$$\widetilde{G}_{i_{j}}(\omega) = \widetilde{S}_{i_{j}}(\omega) + \widetilde{O}_{i_{j}}(\omega) = \widetilde{F}_{i_{j}}(\omega) + \widetilde{M}_{i_{j}}(\omega) + \widetilde{O}_{i_{j}}(\omega)
= A_{F_{ij}} \cdot e^{-i\left(\omega \cdot T_{ij}^{F} + \varphi_{F}\right)} + A_{M_{ij}} \cdot e^{-i\left(\omega \cdot T_{ij}^{M} + \varphi_{M}\right)} + \widetilde{O}_{i_{j}}(\omega),$$
(S1)

where T_{ij}^{M} and φ_{M} are phase travel time and initial phase of the first overtone surface wave, respectively.

By performing three-station interferometry, we plug equation S1 into equation 4a:

$$\tilde{I}_{i_j}(\omega;k) = \tilde{I}^F_{i_j}(\omega;k) + \tilde{I}^M_{i_j}(\omega;k) + \tilde{I}^{FM}_{i_j}(\omega;k) + \tilde{I}^O_{i_j}(\omega;k),$$
(S2)

where each component is given by:

$$\tilde{I}_{i_{j}}^{F}(\omega;k) = A_{F_{ik}} \cdot A_{F_{jk}} \cdot e^{i\varphi_{i_{j}k}} = \begin{cases} \tilde{F}_{i_{k}k}^{*}(\omega) \cdot \tilde{F}_{j_{k}k}(\omega), & k < i \\ \tilde{F}_{i_{k}k}(\omega) \cdot \tilde{F}_{j_{k}k}(\omega), & i < k < j, \\ \tilde{F}_{i_{k}k}(\omega) \cdot \tilde{F}_{j_{k}k}^{*}(\omega), & k > j \end{cases}$$
(S3a)

$$\tilde{I}_{i_{-}j}^{M}(\omega;k) = \begin{cases} \widetilde{M}_{i_{-}k}^{*}(\omega) \cdot \widetilde{M}_{j_{-}k}(\omega), & k < i \\ \widetilde{M}_{i_{-}k}(\omega) \cdot \widetilde{M}_{j_{-}k}(\omega), & i < k < j, \\ \widetilde{M}_{i_{-}k}(\omega) \cdot \widetilde{M}_{j_{-}k}^{*}(\omega), & k > j \end{cases}$$
(S3b)

$$\tilde{I}_{i_j}^{FM}(\omega;k) = \begin{cases} \tilde{F}_{i_k}^*(\omega) \cdot \tilde{M}_{j_k}(\omega) + \tilde{M}_{i_k}^*(\omega) \cdot \tilde{F}_{j_k}(\omega), & k < i \\ \tilde{F}_{i_k}(\omega) \cdot \tilde{M}_{j_k}(\omega) + \tilde{M}_{i_k}(\omega) \cdot \tilde{F}_{j_k}(\omega), & i < k < j, \\ \tilde{F}_{i_k}(\omega) \cdot \tilde{M}_{j_k}^*(\omega) + \tilde{M}_{i_k}(\omega) \cdot \tilde{F}_{j_k}^*(\omega), & k > j \end{cases}$$
(S3c)

and

$$\tilde{I}_{i_j}^{O}(\omega;k) = \begin{cases} \tilde{G}_{i_k}^{*}(\omega) \cdot \tilde{O}_{j_k}(\omega) + \tilde{O}_{i_k}^{*}(\omega) \cdot \tilde{S}_{j_k}(\omega), & k < i \\ \tilde{G}_{i_k}(\omega) \cdot \tilde{O}_{j_k}(\omega) + \tilde{O}_{i_k}(\omega) \cdot \tilde{S}_{j_k}(\omega), & i < k < j. \\ \tilde{G}_{i_k}(\omega) \cdot \tilde{O}_{j_k}^{*}(\omega) + \tilde{O}_{i_k}(\omega) \cdot \tilde{S}_{j_k}^{*}(\omega), & k > j \end{cases}$$
(S3d)

Combined with equation 3a, the terms $\tilde{I}_{i_j}^F(\omega; k)$ and $\tilde{I}_{i_j}^M(\omega; k)$ are only related to the fundamental mode and the first overtone surface waves, respectively, and therefore are independent on k value in phase as illustrated in equation 4b. On the other hand, the term $\tilde{I}_{i_j}^O(\omega; k)$ is associated with non-surface wave signals $\tilde{O}(\omega)$, and therefore varies significantly with k value in phase.

For the term $\tilde{I}_{i,j}^{FM}(\omega; k)$ that represents the interference between the two modes of surface waves, we can further expand equation S3c by plugging in equation S1: $\tilde{I}_{i,j}^{FM}(\omega; k)$

$$= \begin{cases} A_{FM_{k}} \cdot e^{-i\omega \cdot \left(T_{ik}^{F} - T_{jk}^{M}\right)} + A_{MF_{k}} \cdot e^{-i\omega \cdot \left(T_{ik}^{M} - T_{jk}^{F}\right)}, & k < i \\ A_{FM_{k}} \cdot e^{-i\left[\omega \cdot \left(T_{ik}^{F} + T_{jk}^{M}\right) + \varphi_{FM}\right]} + A_{MF_{k}} \cdot e^{-i\left[\omega \cdot \left(T_{ik}^{M} + T_{jk}^{F}\right) + \varphi_{FM}\right]}, & i < k < j \\ A_{FM_{k}} \cdot e^{-i\omega \cdot \left(T_{ik}^{M} - T_{jk}^{F}\right)} + A_{MF_{k}} \cdot e^{-i\omega \cdot \left(T_{ik}^{F} - T_{jk}^{M}\right)}, & k > j \end{cases}$$
(S4)

where $A_{FM_k} = A_{F_ik} \cdot A_{M_jk}$, $A_{MF_k} = A_{M_ik} \cdot A_{F_jk}$, and $\varphi_{FM} = \varphi_F + \varphi_M$. The travel time differences of $T_{ik}^F - T_{jk}^M$ and $T_{ik}^M - T_{jk}^F$ for k < i are dependent on the choice of k, and it is also true for cases when k < j and i < k < j. Therefore, the terms $\tilde{I}_{i_j}^{FM}(\omega; k)$ and $\tilde{I}_{i_j}^O(\omega; k)$ are suppressed through the stacking introduced in section 3.2 (Equation 5b). This is equivalent to that, after stacking, only the terms $\tilde{I}_{i_j}^F(\omega; k)$ and $\tilde{I}_{i_j}^M(\omega; k)$ are preserved, i.e., the denoised waveform will only enhance the fundamental mode and first overtone surface waves but suppress the contributions from non-surface wave signals and the interference pattern between the two modes. Similar formulations can easily be derived for cases when more than two modes of surface waves are present.



Figure S1. (a)-(c) Same as Fig. 2 for TT component data narrow bandpass filtered at 0.8 s. (d)-(g) Same as Fig. 3 for TT component data narrow bandpass filtered at 0.8 s. Signal to noise ratio of the fundamental mode surface waves is much higher in the filtered ANCs, (a)-(b), than that of Fig. 2. As illustrated in panel (c), the systematic time shift between black and blue waveforms is zero, suggesting φ_F (Equation 2b) is zero. Comparison between panels (e) and (f) suggests that the denoising process mainly suppresses coda waves of the filtered ANCs at low frequency (0.8 s).



Figure S2. Same as Fig. S1 for ZZ component data narrow bandpass filtered at 0.3 s. Similar to Fig. 3, signal to noise ratio of the fundamental mode surface waves is low.



Figure S3. Same as Fig. S2 for the denoising of ZZ component data narrow bandpass filtered at 0.8 s.



Figure S4. Synthetic test for phase weighted stacking...



Figure S5. (a) Phase (red) and group (cyan) velocity are estimated as the velocity of the peak of the curve using data at TT component narrow bandpass filtered and denoised at 0.3 s (shown in Fig. 3c), assuming a homogenous velocity structure beneath the array (Appendix I). The black dashed line is used to estimate uncertainty of the array-mean velocity. (b) Array-mean phase (red) and group (cyan) velocity dispersion curves, with error bars indicating the estimated uncertainty. Stars depict the array-mean velocities obtained in (a). (c) Same as Fig. 4c with the red and cyan dashed lines illustrating the moveout of array-mean phase and group velocities. Peak frequency of the array-mean amplitude spectrum (dashed lines in Fig. 4d) is labeled at the bottom right. (d) Waveforms aligned with respect to the travel time predicted by the array-mean phase velocity (red dashed line). The waveforms are further tapered using time windows outlined by the black dashed lines (Appendix I). The cyan dashed line indicates the array-mean group velocity moveout after the alignment.



Figure S6. Same as Fig. S5 for TT component data narrow bandpass filtered at 0.8 s.



Figure S7. Same as Fig. S5 for ZZ component data narrow bandpass filtered at 0.3 s.



Figure S8. Same as Fig. S5 for ZZ component data narrow bandpass filtered at 0.8 s.



Figure S9. Same as Fig. 5 for denoised data of TT component narrow bandpass filtered at 0.8s.



Figure S10. Same as Fig. 5 for denoised data of ZZ component narrow bandpass filtered at 0.3 s.



Figure S11. Same as Fig. 5 for denoised data of ZZ component narrow bandpass filtered at 0.8s.



Figure S12. Same as Fig. 5 for denoised data of TT component narrow bandpass filtered at 0.4 s. Denoised waveforms after tapering are less coherent in (a) compared to results at other periods. Phase velocity models inferred from different virtual shot gather (gray curves and colormap) are in general inconsistent and thus yield large uncertainty values.



Figure S13. Same as Fig. 3 for the denoising of TT component data narrow bandpass filtered at 0.4 s. Coherent waves that travel at a slightly faster speed with weaker amplitude than the fundamental surface waves are observed in (c), the wavefield after denoising. In addition, the amplitude spectrum after denoising shows 2-3 peaks compared, which is consistent with the observation that there is likely non-negligible higher-mode surface waves with a slightly different dominate frequency in the denoised wavefield.



Figure S14. Same as Fig. S5 for the denoised data of TT component narrow bandpass filtered at 0.4 s. Different from Fig. S5, two peaks are found in the grid search curve for array-mean Love wave phase velocity shown in (a), suggesting there are at least two modes of surface waves in the final denoised wavefield shown in (c).



Figure S15. Same as Fig. 4 for the denoising of TT component data narrow bandpass filtered at 0.2 s. Compared to the filtered ANCs in (b), the fundamental mode surface waves are significantly enhanced after the denoising in (c). In addition, similar to Fig. S13, both (c) and (d) may indicate the existence of weak higher mode surface waves in the denoised wavefield at 0.2 s.



Figure S16. Same as Fig. 4 for the denoising of ZZ component data narrow bandpass filtered at 0.2 s. Fundamental mode surface waves are significantly enhanced after the denoising process from (b) to (c).