

# Seasonality of Tropical Photosynthesis: A Global Map of Drivers and Comparison to Model Outputs

Maria del Rosario Uribe<sup>1</sup>, Carlos A. Sierra<sup>2</sup>, and Jeffrey Dukes<sup>3</sup>

<sup>1</sup>University of California, Irvine

<sup>2</sup>Max-Planck-Institute for Biogeochemistry

<sup>3</sup>Purdue University

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## Abstract

Tropical ecosystems strongly influence Earth's climate and weather patterns. Most tropical ecosystems remain warm year-round; nonetheless, their plants undergo seasonal cycles of carbon and water exchange. Previous research has shown the importance of water and light as drivers of the seasonality of photosynthetic activity in the tropics. Although data are scarce, field-based studies have found that seasonal cycles at a handful of tropical forest sites do not match those in land surface model simulations. A comprehensive understanding and model comparison of how seasonal variations in tropical photosynthetic activity relate to climate is lacking. In this study, we identify the seasonal relationships of precipitation and light availability with satellite-based photosynthetic activity. Three dominant and spatially distinct seasonal relationships emerge between photosynthetic activity and these two environmental drivers: photosynthetic activity that is positively correlated with both drivers (36% of tropical pixels), activity that increases following rain but decreases with light (28%), and activity that increases following bright seasons but decreases with rain (14%). We compare distributions of these observed relationships with those simulated by land surface models. In general, model simulations of gross primary productivity (GPP) overestimate the extent of positive correlations of photosynthetic activity with water and underestimate positive correlations with light. The largest discrepancies between simulations and observations are in the representation of the regions where photosynthetic activity increases with light and decreases with rain. Our clear scheme for representing the relationship between climate and photosynthetic activity can be used to benchmark tropical seasonality of GPP in land models.

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2 **Seasonality of Tropical Photosynthesis: A Global Map of Drivers and Comparison to**  
3 **Model Outputs**

4 **M. R. Uribe,<sup>1,2</sup> C. A. Sierra,<sup>3</sup> and J. S. Dukes<sup>1,2,4</sup>**

5 <sup>1</sup> Department of Forestry and Natural Resources, Purdue University, West Lafayette, IN.

6 <sup>2</sup> Purdue Climate Change Research Center, Purdue University, West Lafayette, IN.

7 <sup>3</sup> Max Planck Institute for Biogeochemistry, Jena, Germany.

8 <sup>4</sup> Department of Biological Sciences, Purdue University, West Lafayette, IN.

9 Corresponding author: Maria del Rosario Uribe (uribem@purdue.edu)

10 **Key Points:**

- 11 • In tropical ecosystems, photosynthetic activity responds strongly to seasonal variation in  
12 water and light.
- 13 • In most regions, photosynthesis varies with seasonal fluctuations in both resources, but  
14 relationships with a single driver are also common.
- 15 • The relationship of climate with photosynthetic activity is characterized by water  
16 availability and its synchronization with light.
- 17 • Land surface models overestimate the extent of positive correlations with water and  
18 underestimate those with light.
- 19 • The relationships of photosynthetic activity with climate identified from remote sensing  
20 data closely agree with global datasets such as VPM and GOSIF.  
21

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24 ecosystems remain warm year-round; nonetheless, their plants undergo seasonal cycles of carbon  
25 and water exchange. Previous research has shown the importance of water and light as drivers of  
26 the seasonality of photosynthetic activity in the tropics. Although data are scarce, field-based  
27 studies have found that seasonal cycles at a handful of tropical forest sites do not match those in  
28 land surface model simulations. A comprehensive understanding and model comparison of how  
29 seasonal variations in tropical photosynthetic activity relate to climate is lacking. In this study,  
30 we identify the seasonal relationships of precipitation and light availability with satellite-based  
31 photosynthetic activity. Three dominant and spatially distinct seasonal relationships emerge  
32 between photosynthetic activity and these two environmental drivers: photosynthetic activity that  
33 is positively correlated with both drivers (36% of tropical pixels), activity that increases  
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36 with those simulated by land surface models. In general, model simulations of gross primary  
37 productivity (GPP) overestimate the extent of positive correlations of photosynthetic activity  
38 with water and underestimate positive correlations with light. The largest discrepancies between  
39 simulations and observations are in the representation of the regions where photosynthetic  
40 activity increases with light and decreases with rain. Our clear scheme for representing the  
41 relationship between climate and photosynthetic activity can be used to benchmark tropical  
42 seasonality of GPP in land models.

## 43 **Plain Language Summary**

44 Tropical ecosystems strongly influence Earth's atmosphere and climate through their high rates  
45 of photosynthesis and transpiration. These rates vary seasonally, but tropical seasonal cycles and  
46 their drivers are incompletely characterized. Here, we identify and characterize the three most  
47 common relationships between seasonality of photosynthetic activity and two main climatic  
48 drivers, water and light. Each type of relationship is associated with specific climate properties.  
49 In short, when seasonal cycles of water and light are synchronized, vegetation responds  
50 positively to both drivers. When the cycles are not synchronized, drier sites respond positively to  
51 water and negatively to light, while wetter sites respond positively to light and negatively to  
52 water. Ecosystem models generally lack the mechanisms that lead to this latter relationship. Our  
53 improved understanding of these relationships can guide model improvement efforts for tropical  
54 ecosystems.

## 55 **1 Introduction**

56 Tropical ecosystems are sometimes called the "lungs of the planet," because their high  
57 photosynthetic rates drive large fluxes of carbon and water. Tropical forests alone account for  
58 about 60% of global terrestrial photosynthesis (Mitchard, 2018) and influence precipitation  
59 patterns, even at the continental scale (Lawrence & Vandecar, 2015). Collectively, tropical  
60 ecosystems disproportionately influence Earth's climate and weather patterns (Malhi et al.,  
61 2008).

62 At the intra-annual or seasonal scale, climate patterns determine ecosystems' metabolism,  
63 phenological patterns, and vegetation distribution. Ecosystem metabolism, in turn, affects the  
64 climate system through photosynthesis and the associated carbon, water, and energy feedbacks to

65 the atmosphere. But this ecosystem-atmosphere interaction is being altered by ongoing changes  
66 in the climate system. Forecasting the effects of these climatic changes on tropical ecosystems,  
67 and the subsequent consequences for biosphere-atmosphere interactions and climate at regional  
68 and global scales requires accurate estimates of current photosynthetic rates in tropical  
69 ecosystems and an understanding of their relationship with climate. While photosynthetic rates  
70 go through clear and well-understood seasonal cycles in temperate regions, seasonality of  
71 photosynthesis in the tropics is less well understood (Wu et al., 2016). Across most ecosystems  
72 in these consistently warm regions, both the patterns of seasonality and the drivers of those  
73 patterns remain largely uncharacterized (Restrepo-Coupe et al., 2017; Saleska et al., 2003).

74 Marked seasonal patterns in vegetation activity, although sometimes weaker or less  
75 defined in comparison to those of the temperate zones, have been observed in both field- and  
76 satellite-based measurements in the tropics. Estimates from eddy covariance towers show strong  
77 seasonal patterns in net ecosystem exchange and gross primary productivity (GPP) in most sites  
78 where data are available (Restrepo-Coupe et al., 2013; Saigusa et al., 2008). Such sites include  
79 tropical rain forests and savannas from the Amazon and Asia. Satellite-based measurements of  
80 proxies of phenology and photosynthetic activity such as leaf area index (LAI), enhanced  
81 vegetation index (EVI) and solar-induced fluorescence (SIF), often show similar seasonal  
82 patterns to those observed in the field (Bertani et al., 2017; Bradley et al., 2011a; Guan et al.,  
83 2015; Myneni et al., 2007; Xu et al., 2015). These studies demonstrate that seasonality extends  
84 across the tropics, with only a small portion of the region not showing any type of seasonality.

85 Land surface models, however, are unable to characterize the observed seasonal cycles,  
86 as shown for some specific sites in the Amazon (Restrepo-Coupe et al., 2017). At individual  
87 study sites, models simulate either constant GPP or opposite seasonal patterns to the ones  
88 observed in the field (Restrepo-Coupe et al., 2017). Yet, these models are a major component of  
89 Earth System Models (ESMs) and constitute the main tool scientists currently rely on for future  
90 projections of climate, ecosystems and their interrelationship. Models that represent seasonal  
91 cycles in the tropics more accurately would be able to estimate how changes in climate  
92 seasonality (e.g., timing or length of wet and dry seasons) could affect intra-annual carbon fluxes  
93 and, subsequently, the annual carbon budgets of tropical ecosystems (Saleska et al., 2003).  
94 Accurate simulation of terrestrial water cycling, including the effects of tropical vegetation on  
95 regional to global precipitation patterns, also depends on realistic simulations of photosynthetic  
96 activity.

97 In order to accurately represent the seasonality of tropical photosynthetic activity in land  
98 surface models, we need to understand the climatic drivers of this seasonality. This involves  
99 recognizing how they vary from one region in the tropics to another, and the potential  
100 mechanisms and delayed responses involved in the climate-vegetation relationship. Water and  
101 light availability are the main drivers of intra-annual variation in vegetation activity in the  
102 tropics, and within the region there is wide variation in the responses to these two drivers  
103 (Nemani et al., 2003; Seddon et al., 2016). Previous studies attribute the regional differences in  
104 vegetation seasonality to water stress (Guan et al., 2015; Wagner et al., 2017). In this sense,  
105 photosynthetic activity follows precipitation cycles in drier ecosystems, such as pastures,  
106 deciduous forests or degraded forests (Bradley et al., 2011a; Huete et al., 2006). In ecosystems  
107 with higher mean annual rainfall or a shorter dry season, like evergreen forests, photosynthetic  
108 activity is either less seasonal or more closely associated with light availability (Guan et al.,  
109 2015; Nemani et al., 2003). The mechanisms leading to these different vegetation-climate

110 relationships are still being studied and, therefore, are more challenging to represent in land  
111 surface models. Moreover, it is common for the relationship of photosynthetic activity with  
112 climate to be lagged depending on the climatic factor and vegetation types (Bradley et al., 2011a;  
113 D. Wu et al., 2015). Identifying the importance of these lagged correlations in different parts of  
114 the tropics can also guide future research and inclusion of underlying mechanisms in models.

115         Despite the large variation in seasonal patterns across the tropical region and the  
116 discrepancy found between GPP from field measurements and models, most satellite-based  
117 studies of tropical vegetation seasonality have focused on the Amazon basin. Moreover, model  
118 performance has only been tested at the site level, also within the Amazon. A global analysis of  
119 the drivers of photosynthesis seasonality in the tropics and how they differ in models would  
120 unveil large-scale patterns. These discoveries could help experiments and models target weakly  
121 represented regions and ecosystems. Here, we investigated how the seasonality of photosynthetic  
122 activity across the entire tropical region relates to the two most important regional-level climate  
123 drivers; precipitation and radiation. These two variables also provide the main forcing data used  
124 by land surface models to simulate most vegetation processes. In contrast to previous studies,  
125 rather than identifying a single climate predictor of photosynthetic activity, we sought to  
126 characterize its relationship (including direction and magnitude) with both precipitation and  
127 radiation. This approach allowed us to identify where in the tropics photosynthesis is positively  
128 or negatively associated with each of the two climate drivers, including lagged responses to these  
129 drivers. We then identified the climatic variables most commonly associated with each of the  
130 various relationships between photosynthesis and the climatic drivers. We expected annual  
131 precipitation and variability to explain the relationship with both water and light; for instance,  
132 photosynthetic activity in extremely wet regions should be positively correlated with light and  
133 negatively correlated with water. In arid regions, photosynthetic activity should be positively  
134 correlated with water and negatively correlated with light. In regions that are neither extremely  
135 wet nor dry, and with more evenly distributed precipitation throughout the year, we expected  
136 photosynthesis to be positively correlated with both water and light. We compared results from  
137 remote sensing data with three ecosystem models and two commonly used GPP datasets. We  
138 expected the most widespread pattern in these models and datasets to be a positive relationship  
139 between photosynthetic activity and water and a negative relationship with light.

## 140 2 Materials and Methods

### 141 2.1 Climate data

142 Monthly mean precipitation and net radiation were retrieved for the period 2000-2017.  
143 Precipitation was obtained from the TRMM (TMPA/3B43) Rainfall V7 product with a spatial  
144 resolution of 0.25-degree x 0.25-degree (TRMM, 2011). This product is the best estimate of an  
145 algorithm that uses multi-satellite data from two instruments, the Precipitation Radar and the  
146 TRMM Microwave Imager (Huffman et al., 2007). Data of incoming shortwave radiation at the  
147 surface were obtained from the Energy Balanced and Filled (EBAF) Surface data product  
148 Edition 2.8 from the NASA Clouds and the Earth's Radiant Energy System (CERES) experiment  
149 at a 1-degree x 1-degree spatial resolution (Loeb et al., 2009). This product is the output of  
150 radiative transfer models that use a series of satellite-based observations of top-of-atmosphere  
151 radiation and cloud physical and radiative properties to calculate the surface data (Kato et al.,  
152 2013).

153 For the characterization of site-specific climate variables, we calculated MAP, mean  
154 radiation, mean temperature, mean dry-season length, and precipitation seasonality index. All of  
155 the climate variables correspond to the average for the period 2000-2016 for each pixel. MAP  
156 and mean radiation were estimated using the datasets above. Mean temperature was estimated for  
157 the period of study from the Climatic Research Unit (CRU) Time-Series (TS) Version 4.02 of  
158 High-Resolution Gridded Data of Month-by-month Variation in Climate (Harris et al., 2014).  
159 The data are in a 0.5 x 0.5-degree grid and are produced based on observational data from  
160 national and external meteorological agencies. Dry-season length was obtained from the Rainy  
161 and Dry Seasons (RADS) dataset (Bombardi et al., 2019). This dataset uses global gridded daily  
162 precipitation datasets to provide several characteristics of precipitation seasonality at 0.25- x  
163 0.25-degree spatial resolution. In RADS, seasons are calculated at the local scale based on the  
164 accumulated precipitation anomalies of each grid point. Accumulated precipitation anomalies are  
165 calculated by comparing daily precipitation against the long-term mean daily precipitation.  
166 Calculations of the accumulated anomalies start every year in the dry season, which is estimated  
167 as the first minimum harmonic of the precipitation mean annual cycle. The start and end of the  
168 wet and dry seasons correspond to inflection points in the accumulated anomalies curve for each  
169 cycle. More details of the algorithm and assumptions for these calculations are provided in  
170 Bombardi et al. (2019). The precipitation seasonality index was calculated using the Walsh and  
171 Lawler equation (Walsh & Lawler, 1981). This index uses the total annual and monthly  
172 precipitation for each year within the period of study to characterize the distribution of  
173 precipitation throughout the year. Small values indicate less seasonality or equal distribution of  
174 the precipitation throughout the year, while higher values indicate higher concentration of  
175 precipitation in fewer months.

176 While some areas of the tropics underwent extensive changes in land cover during the  
177 study period (Hansen et al., 2013), the models in this study used static land cover data. Because  
178 of the contrast between the coarse resolution of the data used here and the fine resolution at  
179 which land cover change occurs, we were unable to include this information in our analyses. The  
180 effects of land cover change on seasonality of tropical GPP and its response to precipitation and  
181 radiation should be examined in future studies.

### 182 2.2 Satellite-based vegetation data

183 We used three independent datasets to estimate photosynthetic activity in this study. Two  
184 of these were satellite-based proxies: Solar Induced Fluorescence (SIF) and the Multi-Angle  
185 Implementation of Atmospheric Correction Enhanced Vegetation Index (MAIAC EVI). The  
186 third is a remote sensing-derived product based on SIF, known as GOSIF. SIF data came from  
187 the GOME2\_F data products V27 (Level 3)  
188 ([https://avdc.gsfc.nasa.gov/pub/data/satellite/MetOp/GOME\\_F/](https://avdc.gsfc.nasa.gov/pub/data/satellite/MetOp/GOME_F/), accessed on April 2018) (Joiner  
189 et al., 2013). This SIF monthly dataset is available at 0.5 x 0.5-degree resolution since 2007. The  
190 GOSIF dataset is based on SIF retrievals from OCO-2 in addition to a predictive model and other  
191 MODIS remote sensing and meteorological reanalysis datasets  
192 (<http://data.globalecology.unh.edu/data/GOSIF/>, accessed on March 2019) (Li & Xiao, 2019).  
193 GOSIF is available monthly at 0.05 x 0.05-degree resolution since 2000. MAIAC EVI was  
194 obtained from calibrated and geometrically corrected MODIS Collection 6 Level 1B satellite  
195 images (<https://portal.nccs.nasa.gov/datashare/maiac/DataRelease/Global-VI-8day-0.05degree/>,  
196 accessed on September 2018) (A. Lyapustin et al., 2018). The monthly MAIAC EVI product is  
197 at 0.05 x 0.05-degree resolution and since 2000.

198 Each of the satellite-based vegetation datasets has advantages and disadvantages. SIF  
199 data from the Global Ozone Monitoring Experiment–2 on MetOp-A and -B (GOME 2) should  
200 represent photosynthetic activity well, but the temporal period available, from 2007 to 2018 is  
201 not an ideal match with available model output, and sensor degradation for GOME-2 has been a  
202 concern (Zhang et al., 2018). GOSIF data, developed from Orbiting Carbon Observatory-2  
203 (OCO-2) measurements, should approximate the SIF data and covers a longer period, from 2000  
204 to 2018. However, the derived data in the GOSIF product are more removed from the direct  
205 observations than SIF. MAIAC EVI has a similarly long time series as GOSIF and has been  
206 frequently used in tropical ecosystem studies; it has proved to be a better proxy of photosynthetic  
207 activity in tropical rainforests than other vegetation indices such as Normalized Difference  
208 Vegetation Index (NDVI) or MODIS EVI (Maeda et al., 2016). However, MAIAC EVI is still a  
209 vegetation index that estimates vegetation greenness and not photosynthetic activity directly,  
210 which is more accurately estimated with SIF (Joiner et al., 2011).

211 Both the SIF and MAIAC EVI datasets have undergone an advanced cloud screening and  
212 filtering process, which has made them advantageous to use in the tropics compared to other  
213 remote sensing products. In SIF, the main problem of clouds for SIF retrievals is a shielding  
214 effect, as, contrary to vegetation indices, the SIF spectral signature is not affected by clouds  
215 (Joiner et al., 2014). Therefore, in the SIF dataset used here, cloud filtering is done by removing  
216 pixels with effective cloud fractions of >30%. This filter threshold has been proved to maintain  
217 spatial and temporal patterns of SIF without altering the sample size and the noise resulting from  
218 reduced coverage (Joiner et al., 2013). MAIAC EVI has a sophisticated cloud and aerosol  
219 screening correction algorithm. In this later product, pixels with atmospheric contamination are  
220 not excluded from the dataset and are not included in our study (A. I. Lyapustin et al., 2012). As  
221 part of our time series analysis, we filled data gaps of a maximum of three months using spline  
222 interpolation. Pixels with gaps longer than three months in the time series were excluded from  
223 the analysis.

### 224 2.3 Modeled GPP

225 To examine tropical seasonality exhibited in land surface models (LSMs), simulated GPP  
226 data were obtained from the TRENDY (Trends and drivers of the regional scale sources and

227 sinks of carbon dioxide) project (Sitch et al., 2015). Most TRENDY models are LSMs  
228 commonly coupled with ESMs and used for climate projections. Here, we used TRENDY v5 S2  
229 simulations from CLM4.5 (Oleson et al., 2013), JULES (Best et al., 2011), and LPJ-GUESS  
230 (Smith et al., 2001). In TRENDY, each model is run globally with different spatial scales and  
231 land cover types, but with the same forcing data. Land cover data for the simulations is fixed and  
232 provided by each modeling group. In addition to the LSM-simulated GPP, two global GPP  
233 products, Fluxcom (Jung & Team, 2016; Tramontana et al., 2016) and VPM (Zhang et al., 2017),  
234 were also analyzed. These datasets are derived using field observations, satellite-based  
235 measurements, and reanalysis meteorological data, in combination with interpolation or machine  
236 learning techniques. The final products are global-scale gridded GPP estimates with long  
237 temporal coverage and high spatial resolution (see Table S1 for details).

## 238 2.4 Data analysis

239 We used time series analysis to identify relationships between climate variables (i.e.,  
240 precipitation and radiation) and various estimates of photosynthetic activity (i.e., SIF, GOSIF,  
241 MAIAC EVI, and simulated GPP) across the entire tropics. We used cross-correlation function  
242 (CCF) analysis (Box et al., 2015), to examine time series of monthly data for each pixel in the  
243 tropics (20°N - 20°S) from 2000 to 2015 (except SIF and GOSIF, which were analyzed for 2007-  
244 2017 and 2000-2017, respectively). Although the SIF dataset covered different years than the  
245 other datasets, the 11 years of SIF data provide a robust basis for our seasonal analyses, and the  
246 SIF analyses can be broadly validated with the longer and independent GOSIF dataset. We  
247 excluded from all analyses all pixels with a mean EVI of less than 0.1. Those low EVI pixels,  
248 which correspond to barren lands or extremely low vegetation cover, were removed to avoid  
249 noise in the photosynthetic activity data and the subsequent calculations. Prior to the analysis,  
250 each pair of climate and photosynthetic activity variables was resampled to a common spatial  
251 resolution in order to enable time series analysis at the pixel level.

252 We used CCF analysis to calculate direct (i.e., same month) and lagged correlations  
253 between each climate variable (i.e., independent variable) and each photosynthetic activity  
254 variable (i.e., response variable). We quantified lagged correlations between one and four months  
255 because a variety of physiological and ecological mechanisms can potentially delay responses of  
256 photosynthetic activity to climate (D. Wu et al., 2015). Our lagged correlations analysis allowed  
257 us to determine the strongest immediate or lagged correlation between precipitation or radiation  
258 and photosynthetic activity.

259 In CCF analysis, temporal dependencies or high autocorrelation in the independent  
260 variable (in this case our climate variables) can hide true relationships or suggest false ones.  
261 "Prewhitening" can remove autocorrelations by extracting the "white noise" from the  
262 independent variable and applying the same transformation to the response variable, in this case,  
263 photosynthetic activity (Cryer & Chan, 2008). We used prewhitening to counteract  
264 autocorrelation, allowing us to analyze the actual linear relationship between the two time series.  
265 The data were pre-whitened by first finding an autoregressive integrated moving average  
266 (ARIMA) model for the climate variable time series, and then fitting the photosynthetic activity  
267 time series to the ARIMA model. ARIMA models are built using information contained in the  
268 time series and are commonly used in forecasting, but in the case of this study, and prewhitening  
269 in general, they are used to filter the original series. The CCF was finally performed on the  
270 climate series model residuals and the filtered photosynthetic activity time series (Box et al.,

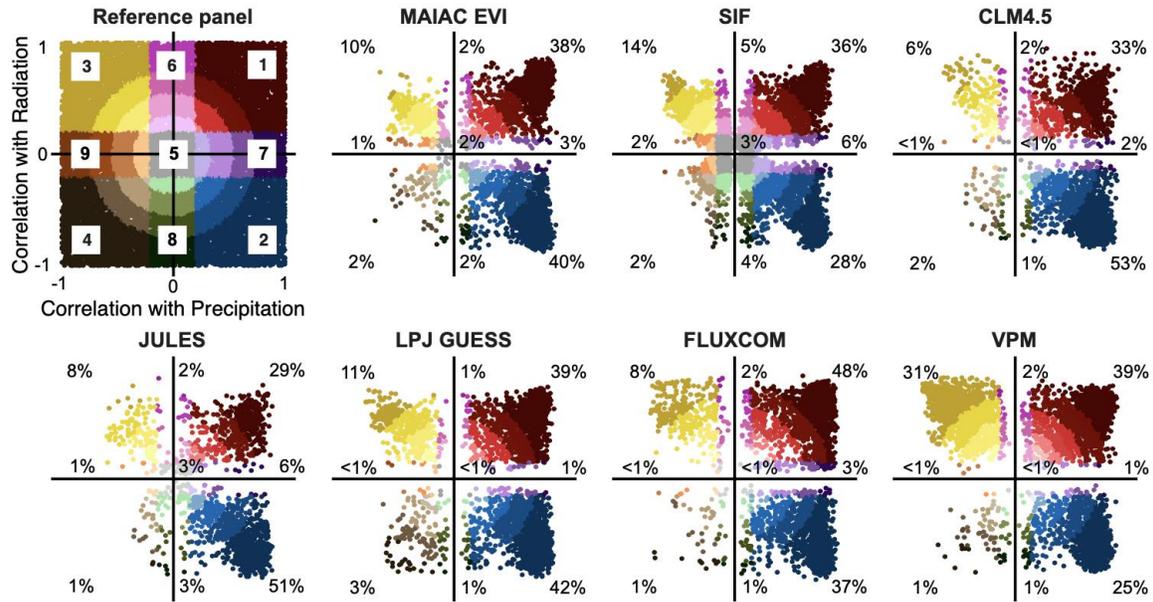
271 2015; Probst et al., 2012). An example of the CCF analysis is available at  
272 [https://github.com/rosariouribed/ccf\\_tropics](https://github.com/rosariouribed/ccf_tropics).

273         Given that precipitation and radiation are two highly correlated variables, we were  
274 interested in an analysis that considered the relationship of photosynthetic activity with the two  
275 variables individually and in combination. We determined the sign of the maximum direct or  
276 lagged correlation coefficient with each variable and then classified the type of relationship  
277 based on the signs and strengths of the two coefficients. For instance, photosynthetic activity  
278 could be correlated positively with precipitation and negatively with radiation or have a non-  
279 significant relationship with precipitation and positive with radiation, and so on (Fig. 1, upper-  
280 right reference panel). We created a map for each of the photosynthetic activity datasets showing  
281 the type of relationship with climate for each pixel. We assessed the agreement between the  
282 results from satellite and model data using the Kappa coefficient ( $\kappa$ ) for map agreement (Cohen,  
283 1960). The Kappa coefficient compares the agreement between two maps against a hypothetical  
284 scenario of randomly assigned values. Coefficients range from -1 to 1, where 0 indicates that the  
285 evaluated map is as good as if random values were selected, negative values suggest the map is  
286 worse than random values, and positive values suggest the evaluated map matches the reference  
287 map better than random values. Values closer to 1 indicate a better agreement between the maps.  
288 We also calculated the overall difference ( $D$ ) as a second measurement of agreement of our  
289 results. Overall difference has been suggested to provide a more reliable comparison between  
290 maps than the Kappa coefficient (Pontius & Santacruz, 2014). Calculations of  $D$  take into  
291 account how well maps agree on (1) the number of pixels classified in each category (i.e.,  
292 quantity difference) and (2) the location of the pixels in each category, given the number of  
293 pixels in each of the categories (i.e., allocation difference) (Pontius & Santacruz, 2014). Larger  
294  $D$  values indicate greater disagreement between the maps, either because of under- or  
295 overestimation of pixels in the different categories, or because of inaccurate spatial allocation of  
296 the pixels in each of the categories. Comparisons between maps were also performed at the  
297 biome level in order to identify models' biome-specific shortcomings.

298         We then used the classification of type of relationship and additional climate properties  
299 of each pixel to identify climate properties most closely associated with a specific type of  
300 relationship. Only the most common types of relationships (i.e., more than 10% of the pixels)  
301 were included in the comparison.

302         The additional climate properties included MAP, mean radiation, mean temperature, dry  
303 season length, precipitation seasonality index, and the correlation between monthly precipitation  
304 and radiation. We used Kruskal-Wallis nonparametric tests (Kruskal & Wallis, 1952) to identify  
305 significant differences in the climate properties among the type of relationships. If the Kruskal-  
306 Wallis result was significant ( $\alpha=0.05$ ) for a particular variable, we followed up with a pairwise  
307 comparison among all types of relationships through the Dunn test with Bonferroni correction  
308 (Dunn, 1964). This nonparametric *post hoc* test can be used for independent groups with non-  
309 normal distributions and different sizes. All analyses were performed in R 3.4.0 (R Core Team,  
310 2017), including the forecast v8.2 (Hyndman, 2017), TSA v1.01 (Chan & Ripley, 2012), FSA  
311 v0.8.22 (Ogle et al., 2018) and diffeR (Pontius Jr. & Santacruz, 2019) packages.

312



313

314 **Fig. 1.** Scatterplots showing the maximum correlation coefficient from the CCF analysis for  
 315 vegetation productivity from each of the datasets with precipitation (x axis) and radiation (y  
 316 axis). (Top-left/Reference panel): the nine colors and numbers correspond to each of the types of  
 317 relationships; the rings indicate the strength of the correlation with both drivers (distance from  
 318 the origin). Regions 1-4 (red, blue, yellow, and brown) indicate significant correlations with both  
 319 drivers. Region 5 (gray color) indicates non-significant relationships with any driver. Regions 6-  
 320 9 (pink, purple, green, and orange indicate non-significant correlations with one of the drivers.  
 321 (Other panels): the numbers indicate the percentage of pixels with the type of relationship where  
 322 the number is located.

### 323 3 Results

#### 324 3.1 Relationship of photosynthetic activity with precipitation and radiation in satellite 325 data

326 Based on the direction of the correlations between photosynthetic activity and the two  
 327 climate drivers, most sites in the tropics can be classified into three categories, with  
 328 photosynthetic activity that is either (1) positively correlated with both precipitation and  
 329 radiation, (2) positively correlated with precipitation but negatively correlated with radiation, or  
 330 (3) positively correlated with radiation but negatively with precipitation (i.e., regions 1, 2 and 3  
 331 in Fig. 1 upper-left reference panel). Hereafter, we refer to these three types of relationships as  
 332 cosynchronous, rain-following and light-following, respectively. For MAIAC EVI and SIF, these  
 333 three types of relationships together account for 78-88% of tropical pixels, where the  
 334 cosynchronous pixels correspond to 36-38%, the rain-following 28-40% and the light- following  
 335 10-14% of pixels for the two datasets, respectively (Fig. 1). Examples of the seasonal patterns of  
 336 precipitation, radiation and photosynthetic activity from SIF in pixels from the three main types  
 337 of relationships are displayed in the Supplementary Materials (Fig. S1). While both datasets  
 338 agree on the dominance of these three types of relationships, there are some differences between  
 339 the two datasets in the proportion of pixels that show each of these three dominant types (Fig. 1).  
 340 In the SIF dataset, more pixels are cosynchronous and light-following and fewer are rain-

341 following than in the MAIAC EVI product. The spatial distribution of the three types of  
 342 relationships is moderately consistent between the two datasets, with a kappa coefficient of  
 343 agreement of 0.5 and overall difference of 35.12% (Table 1, Fig. 2). The results from GOSIF  
 344 agree closely with those from SIF and EVI ( $\kappa = 0.52$  and  $0.65$ ,  $D = 33.42$  and  $22.58\%$ ,  
 345 respectively), providing stronger confidence in our results (Fig. S2).

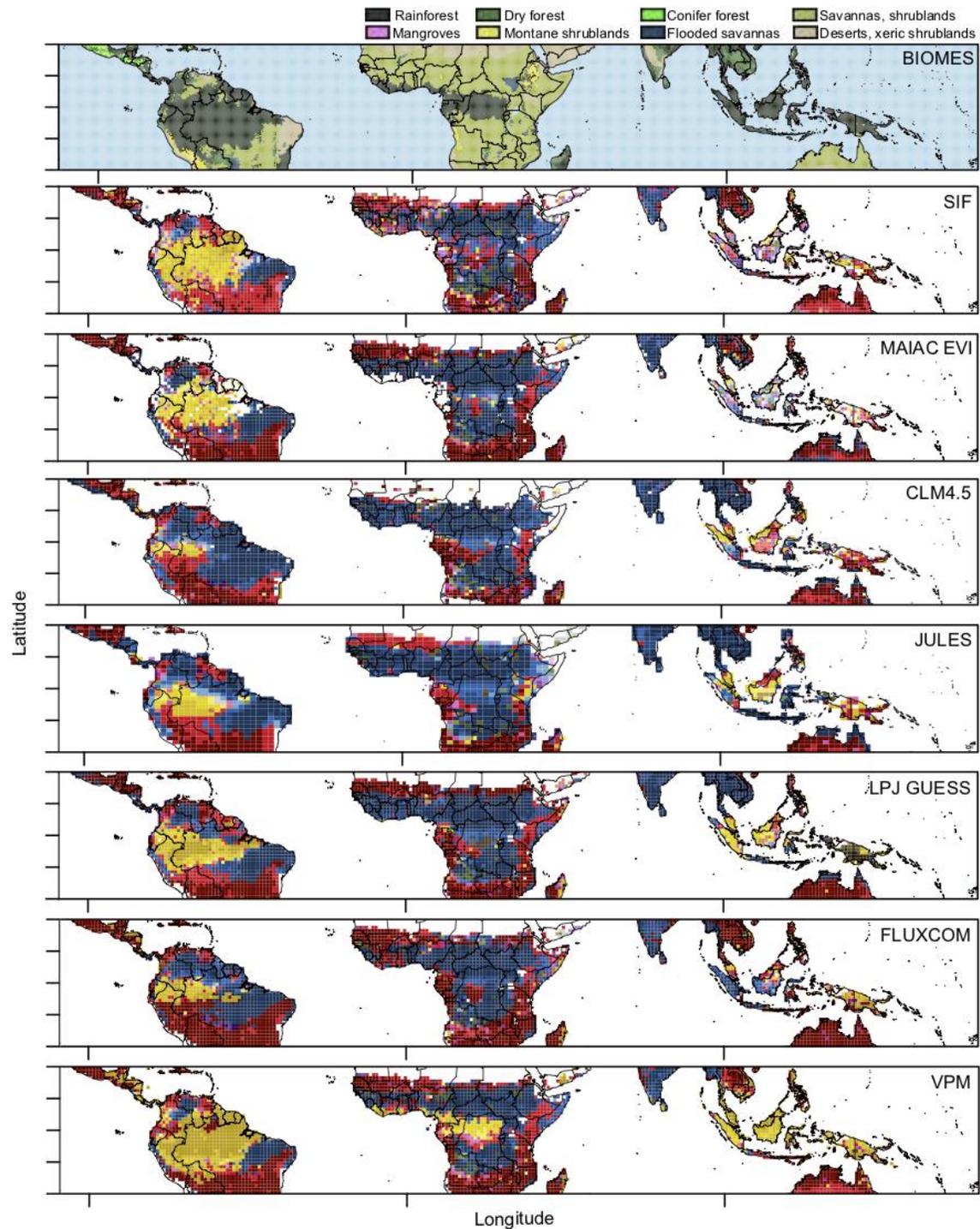
346 Based on the SIF results, cosynchronous pixels (i.e., positive correlations with both  
 347 precipitation and light) are located across all biomes, but make up the largest fraction (45-86%)  
 348 of savannas and shrublands, dry broadleaf forests, flooded savannas, montane shrublands, and  
 349 conifer forest (Fig. S3). The rain-following relationship (28% across all biomes) is most common  
 350 in savannas and shrublands, dry forests, flooded savannas, and xeric shrublands. Light-following  
 351 seasonality is mostly clustered in rainforests, specifically those of the Amazon basin and  
 352 southeast Asia.

353 Other types of relationships are not common (<22% in total), but include negative  
 354 correlations with both drivers (<3% of pixels) and weak correlations with one (<6% of pixels) or  
 355 both (<3% of pixels) drivers. For the SIF dataset, these other types of correlations occur mostly  
 356 in the rainforest in South America, central Africa and southeast Asia. A more detailed look at  
 357 areas with the weakest relationships shows a prevalence of pixels with very low intra-annual  
 358 variability in photosynthetic activity. These less distinct types of relationships could also be  
 359 explained by a higher diversity of vegetation cover or land cover change during the study period.

360 The seasonal peak of photosynthesis most commonly occurred within two months of the  
 361 seasonal peak of precipitation, while lags with radiation had a much wider range of variation,  
 362 from 0 to 4 months (Fig. 3, Fig. S4). These lags also varied among the types of relationships. In  
 363 regions with a cosynchronous relationship, peaks of photosynthesis typically followed peak  
 364 precipitation by 0-2 months, and peak radiation by 3-4 months (Fig. 3, Fig. S4a). In regions with  
 365 a rain-following relationship, photosynthetic activity lagged peaks of precipitation (and lows of  
 366 radiation) by 0-2 months (Fig. 3, Fig. S4b). In regions with a light-following relationship,  
 367 photosynthetic peaks lagged minimum precipitation by a wide range of 0-4 months, and  
 368 commonly lagged radiation peaks by 0 to 3 months, although with large variation (Fig. S4c). The  
 369 relatively few pixels that have the longer time lags with precipitation occur in the tropical  
 370 rainforests of South America, which is mostly a light-following region. A large proportion of the  
 371 longer time lags with radiation coincides with cosynchronous regions in the higher tropical  
 372 latitudes of Africa and Australia.

### 373 3.2 Climate properties of the different types of relationships of photosynthetic seasonality 374 with precipitation and radiation

375 Each of the three dominant climate-seasonality relationships was associated with distinct  
 376 climatic properties (Fig. 4A-E). Areas with a cosynchronous relationship tended to have  
 377 relatively low annual rainfall (Median = 1217 and Interquartile Range = 26 mm), high mean  
 378 daily radiation ( $\bar{x} = 227.9$  and  $IQR = 1.12 \text{ W m}^{-2}$ ), a long dry season ( $\bar{x} = 221.9$  and  $IQR = 1.4$   
 379 days), relatively low mean temperature ( $\bar{x} = 25.7$  and  $IQR = 0.1 \text{ }^\circ\text{C}$ ), and high rainfall seasonality  
 380 ( $\bar{x} = 0.8$  and  $IQR = 0.01$ ). Rain-following pixels had low MAP ( $\bar{x} = 1214$  and  $IQR = 24 \text{ mm}$ ),

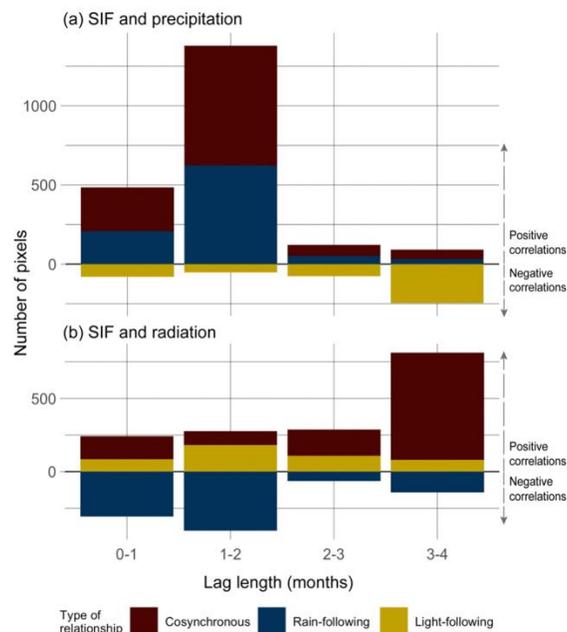


381

382 **Fig. 2.** Maps of biomes and results from the CCF analysis. (Top panel): biomes of the tropics  
 383 (WWF). (Other panels): Maps showing the spatial distribution of the maximum correlation  
 384 coefficient from the CCF analysis for vegetation productivity from each of the datasets with  
 385 precipitation and radiation. Colors in the map correspond to the colors and numbers in the  
 386 reference panel Fig. 1, which contain information about the type of combined relationship and  
 387 strength of the correlations. White pixels correspond to water bodies or pixels with scarce data  
 388 for CCF analysis.

389 high mean daily radiation ( $\bar{x} = 227.9$  and  $\text{IQR} = 0.9 \text{ W m}^{-2}$ ), an intermediate-length dry season ( $\bar{x}$   
 390  $= 203.1$  and  $\text{IQR} = 1.6$  days), higher mean temperature ( $\bar{x} = 26.1$  and  $\text{IQR} = 0.1 \text{ }^\circ\text{C}$ ), and  
 391 intermediate rainfall seasonality ( $\bar{x} = 0.8$  and  $\text{IQR} = 0.01$ ). Light-following pixels had high MAP  
 392 ( $\bar{x} = 2466$  and  $\text{IQR} = 40 \text{ mm}$ ), low mean daily radiation ( $\bar{x} = 201.5$  and  $\text{IQR} = 0.7 \text{ W m}^{-2}$ ), a short  
 393 dry season ( $\bar{x} = 189.9$  and  $\text{IQR} = 1.1$  days), high mean temperature ( $\bar{x} = 26.4$  and  $\text{IQR} = 0.1 \text{ }^\circ\text{C}$ ),  
 394 and low precipitation seasonality ( $\bar{x} = 0.5$  and  $\text{IQR} = 0.01$ ). The climates of areas with the light-  
 395 following relationship stood out as distinct from those in other areas across four of the five  
 396 variables – only temperature, by not differing from regions with the rain-following relationship,  
 397 opposed this trend. Cosynchronous and rain-following pixels had similar MAP and mean daily  
 398 radiation, but cosynchronous pixels had a longer dry season, lower mean temperature, and larger  
 399 precipitation seasonality index.

400 The correlation between precipitation and radiation also provided meaningful information  
 401 about the relationship between photosynthetic activity and climate. In areas with a  
 402 cosynchronous relationship, light and precipitation typically were abundant during the same or  
 403 similar times of year (Fig. 4F). In these areas, photosynthesis was greatest during the times of  
 404 year that were both wet and bright. In contrast, in areas with the other two dominant types of  
 405 relationships, the rainiest times of year were the darkest, and the driest seasons were brightest. In  
 406 these areas, GPP responded most positively to the climate factor that was most limiting. Thus,  
 407 areas that have lower MAP and a longer dry are water-limited and showed a positive correlation  
 408 with rainfall, and consequently a negative correlation with radiation. In areas with higher MAP



409

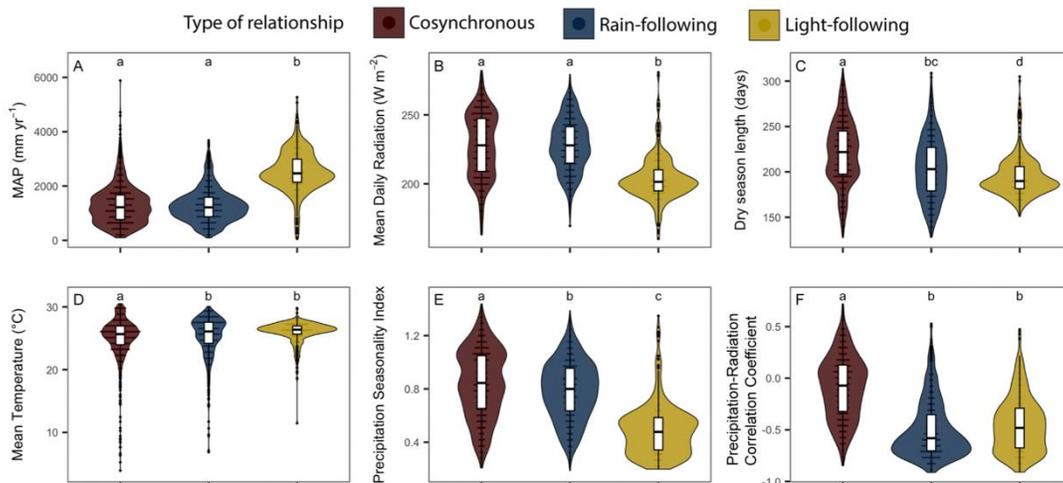
410 **Fig. 3.** Frequency (number of pixels) of the length of lags in the strongest correlations between  
 411 SIF and (a) precipitation and (b) radiation for the three most common types of relationships  
 412 ( $n=2533$ ). Colors correspond to the colors of the types of relationships shown in the reference  
 413 panel in Fig. 1. In order to differentiate the direction of the correlation, the count of pixels with  
 414 positive correlation coefficients between the climate driver and SIF is shown upward and the  
 415 count with negative correlation coefficients is shown downward. Note the difference in scales  
 416 between (a) and (b).

417 and a shorter dry season, GPP is limited by light; photosynthesis was greatest during the  
 418 brightest times of year and slower during the wetter (and darker) times of year.

### 419 3.3 Relationship of photosynthetic activity with water and light in ecosystem models

420 In land surface models, GPP showed the same three dominant types of relationships with  
 421 climate that we found in the satellite data. However, the proportion of vegetation with each type  
 422 of relationship, and the spatial distributions of these relationships, differed from observations  
 423 (Figs. 1, 2 and S5, S6). In the simulated data, the rain-following relationship was more  
 424 widespread than any other (51-53%), and always more common than in the satellite-based  
 425 datasets. In contrast, the proportion of cosynchronous pixels was typically lower than that shown  
 426 by satellite data (20-39%), as was also the case for the light-following relationship (6-11%).  
 427 Kappa coefficients and overall differences between the models' results and the MAIAC EVI and  
 428 SIF datasets (Table 1) ranged from 0.28 to 0.54 and 35.05 to 50.23%, respectively, with LPJ-  
 429 GUESS showing the most similarity to both satellite-based datasets. Overall, the models  
 430 produced more similar GPP predictions to one another than to the satellite-based observations  
 431 (Table 1).

432 Other vegetation products analyzed in this study also tended to show different patterns  
 433 from those in models and observations (Figs. 1 and 2). Fluxcom GPP had a high percentage of  
 434 pixels with cosynchronous and rain-following relationships (48% and 37%), while light-  
 435 following relationships were less common than in the satellite data (8%). VPM GPP showed a lot  
 436 more diversity in the type of relationships, with the greatest proportion of the tropics having  
 437 cosynchronous and light- following relationships (39% and 31%), and slightly less area covered  
 438 by rain-following pixels (25%). The Kappa coefficients and overall differences indicate that  
 439 VPM output agreed more closely with both satellite-based datasets and with the models than



440

441 **Fig. 4.** Climate characteristics of the three most common types of relationships inferred from  
 442 SIF. The shape of the violin plots and the horizontal lines inside of them represent number of  
 443 pixels; wider shapes and longer horizontal lines are more pixels at that level of that specific  
 444 variable. The boxplot shows the median and the quartiles. Significant differences between pairs  
 445 of groups are represented through letters; groups with same letters are not significantly different.  
 446 Colors and numbers of the types of relationships (legend) correspond to the colors and numbers  
 447 in the reference panel in Fig. 1.

448 Fluxcom (Table 1). Similar to the satellite observations, the model most similar to these two  
449 GPP products was LPJ-GUESS.

450 In our biome-level comparisons (Table S2), the best agreement between SIF and most  
451 models was found in montane grasslands and shrublands ( $\kappa = 0.46-0.57$ ,  $D = 18-25\%$ ), flooded  
452 grasslands and savannas ( $\kappa = 0.38-0.44$ ,  $D = 28.6-34.2\%$ ), and grasslands and savannas ( $\kappa =$   
453  $0.33-0.46$ ,  $D = 30.7-41.8\%$ ), in that order. In contrast, low agreement was found in mangroves ( $\kappa$   
454  $= -0.17-0.28$ ,  $D = 52.2-72.7\%$ ), rainforest ( $\kappa = 0.19-0.34$ ,  $D = 45.4-62.7\%$ ) and deserts and xeric  
455 shrublands ( $\kappa = 0.12-0.33$ ,  $D = 31.5-41.5\%$ ). In dry forests, Fluxcom and VPM showed closer  
456 agreement with SIF than the land surface models did.

#### 457 **4 Discussion**

458 Our results identify three primary relationships between photosynthetic activity and  
459 seasonal patterns of water and light availability across the entire tropical region. In addition to  
460 mapping these patterns, our classification provides a clear scheme for model comparison.  
461 Finally, the analysis shows where, and for which climate variables, lagged correlations occur and  
462 need to be accounted for in ecosystem models.

463 Previous research has shown the importance and the spatial distribution of water and light  
464 as drivers of photosynthetic activity at the seasonal scale in the tropics (Guan et al., 2015;  
465 Nemani et al., 2003; Seddon et al., 2016). Our results are in agreement with those studies; we  
466 found the prevalence of positive correlations (1) with water in most of the drier regions in the  
467 tropics and (2) with light in the wet tropical forests. In addition to finding these positive  
468 correlations, we also considered negative correlations with each driver as part of our analyses.  
469 These results revealed that a positive correlation with one driver does not necessarily imply a  
470 negative correlation with the other driver, particularly when both climate variables are highly  
471 correlated with each other. In fact, positive correlations with both drivers, a cosynchronous  
472 relationship, is the most common type of relationship in the tropics.

473 The cosynchronous relationship is mainly found in regions with a low MAP, high mean  
474 daily radiation, long dry season, and high precipitation seasonality. These regions correspond to  
475 water-limited biomes located in higher latitudes where precipitation and radiation are not  
476 negatively correlated, meaning water and light are abundant at approximately the same time.  
477 Rain-following relationships are found in regions with similar precipitation and radiation  
478 characteristics and are also common in similar biomes. However, this relationship is more  
479 common at lower latitudes where precipitation and radiation are negatively correlated.  
480 Vegetation from cosynchronous and rain-following regions is likely water-stressed for part of the  
481 year, and therefore responds positively to water availability. What differentiates the regions with  
482 these two types of relationships is the timing of radiation with respect to precipitation. In the  
483 cosynchronous relationship, vegetation is adapted to take advantage of seasonal availability (up  
484 to 4 months) of both water and light. In the rain-following relationship, vegetation is adapted to  
485 periods of increased rainfall that do not coincide with the most abundant light.

486 The third most common type of relationship, the light-following relationship, is found in  
487 regions with high MAP, low mean daily radiation, a short dry season, low precipitation  
488 seasonality, and a negative correlation between precipitation and radiation. This relationship and  
489 climate characteristics are very specific to the rainforests of South America where previous field  
490 (Restrepo-Coupe et al., 2013; Saleska et al., 2003) and satellite (Bertani et al., 2017; Bradley et

491 al., 2011b; Wagner et al., 2017) findings show light-driven seasonal photosynthetic activity.  
492 Similarly, the spatial distributions of the rain- and light- following relationships found here are  
493 comparable to the wet- and dry-season greening regions for tropical rainforests, with those  
494 forests in the central Amazon and southeast Asia greening in the dry season, when radiation  
495 peaks (Doughty et al., 2019; Guan et al., 2015).

496 The mechanisms by which plant species have adapted to (or been selected by) seasonal  
497 fluctuations in precipitation and light could determine vegetation responses to any future changes  
498 in these patterns. Recognizing the specific correlation of photosynthetic activity in these regions  
499 with the availability of water and light may lead to studies that identify relevant adaptive traits  
500 and selective processes, and to better-informed predictions.

501 Rather than establishing direct causality between water, light, and photosynthetic activity,  
502 we explored the nature of observed relationships. The widespread lagged correlations identified  
503 here should guide experimental and mechanistic representation of photosynthetic activity. These  
504 lagged correlations are consistent with analyses of Bradley et al. (2011) and Wu et al. (2015).  
505 The observed time lags suggest that responses of vegetation to rain and light are influenced by  
506 intermediate or indirect processes. Enhanced vegetation growth and leaf development may  
507 continue long after the seasonal peaks in delivery of water or light, such that the ecosystem  
508 reaches maximum biomass of leaves, or of the most photosynthetically-efficient leaves, later in  
509 the season (Duff et al., 1997; Rivera et al., 2002; J. Wu et al., 2016). This continued growth  
510 could cause delayed peaks in photosynthetic activity relative to resource delivery. We found a  
511 high frequency of lagged correlations in cosynchronous regions, where correlations with  
512 precipitation were strongest for short time lags, but correlations with radiation tended to be  
513 strongest after longer lags. The longer radiation lags mostly occur in the higher tropical latitudes.  
514 We speculate that this delayed correlation is the result of plant growth and leaf development  
515 continuing throughout the summer season (i.e., continuing after peak radiation), with increased  
516 precipitation later in the season leading to a spike in photosynthetic activity. However,  
517 phenology and increase in foliage can be highly variable in tropical ecosystems, even within the  
518 same biome (Bie et al., 1998; Guan et al., 2013; Ma et al., 2013; Monasterio & Sarmiento, 1976;  
519 Moore et al., 2018). Particularly in savannas, phenology and vegetation growth are influenced by  
520 plant composition (i.e., tree and grass fraction covers) and fire seasonality (Guan et al., 2014;  
521 Williams et al., 2005).

522 The light-following relationship is common in tropical forests, where trees have been  
523 shown to increase photosynthetic activity throughout the dry season. This decoupling of  
524 photosynthesis from precipitation is made possible by the deep-rooted trees, which can obtain  
525 water from deep soil layers when shallow soils are dry (Nepstad et al., 1994; Restrepo-Coupe et  
526 al., 2013). Thus, soil water storage and adaptive leaf development and demography likely  
527 regulate the relationship between precipitation and photosynthetic activity in these regions. Both  
528 mechanisms have previously been demonstrated to mediate the relationship between climate and  
529 photosynthetic activity in tropical forests (Guan et al., 2015; Restrepo-Coupe et al., 2013; Wu et  
530 al., 2016, 2017), and likely play a strong role in our results. For instance, some of these forests'  
531 photosynthetic activity peaks happen in the late dry season, months after precipitation declined;  
532 similarly, they reach low photosynthetic activity levels in the late wet season months after  
533 precipitation has peaked. However, site-specific differences in variables such as soil texture and  
534 plant traits also likely contribute to the observed geographical variation in these time lags. Our  
535 results suggest that it is important for models to consider mechanisms such as these, which can

536 affect the relationships of tropical ecosystems' photosynthetic activity with precipitation and  
 537 light. Our results show the extent of the area of different types of relationships and the associated  
 538 time lags with each driver.

539 In general, land surface models accurately represent most of the regional variation in the  
 540 type of relationship of photosynthesis with precipitation and radiation. However, the models  
 541 have two primary deficiencies: they tend to overestimate the extent of positive relationships with  
 542 precipitation, and underestimate the extent of positive relationships with radiation. These two  
 543 issues result in models overestimating the number of rain-following pixels and underestimating  
 544 the number of cosynchronous and light-following pixels. Differences between models and SIF  
 545 results for the cosynchronous relationships are located in scattered regions and biomes across the  
 546 tropics, with one focus area in mainland Southeast Asia. This constitutes a highly deforested  
 547 region widely covered by irrigated crops (Sen et al., 2013). The differences between models and  
 548 remote sensing found in this region could therefore be due to the lack of representation of land  
 549 cover change in the simulations analyzed here.

550 **Table 1.** Kappa coefficients ( $\kappa$ ) and overall difference (D) between the maps resulting from the  
 551 CCF analysis (Fig. 2). Larger  $\kappa$  values indicate closer agreement between the results of two  
 552 datasets. Larger D values indicate larger differences between the results of two datasets.

	MAIAC EVI	CLM4.5	JULES	LPJ-GUESS	FLUXCOM	VPM
SIF	$\kappa = 0.5$	0.28	0.3	0.39	0.4	0.48
	D = 35.12	50.23	49.31	43.16	41.15	37.26
MAIAC EVI		0.45	0.45	0.55	0.5	0.55
		35.05	35.31	29.02	31.96	31.24
CLM4.5			0.52	0.54	0.44	0.31
			28.2	29.03	34.81	49.2
JULES				0.55	0.4	0.32
				28.42	38.6	48.18
LPJ-GUESS					0.43	0.47
					36.42	36.6
FLUXCOM						0.43
						39.01

553 The largest differences between SIF and models are in rainforest regions. Model  
 554 misrepresentation of the seasonality of photosynthetic activity was previously demonstrated  
 555 using GPP data from eddy flux towers in the Amazon (Restrepo-Coupe et al., 2017). Our results  
 556 show the extent of this misrepresentation across the entire Amazon and other smaller rainforest  
 557 areas in Asia. The misrepresentation of rainforests, one of the largest biomes in the tropics, and  
 558 their characteristic light-following relationship is of major concern. As mentioned above,  
 559 incorporating increased rooting depth, leaf demography, or other processes could improve the  
 560 representation of seasonality in simulations (Poulter et al., 2009; J. Wu et al., 2017). The VPM  
 561 GPP dataset was able to reproduce the light-following relationship, something that neither the  
 562 models nor Fluxcom GPP did. VPM's capacity to represent this type of relationship might be  
 563 associated with the light-based model it uses to estimate GPP, which is not used by the other  
 564 datasets analyzed in this study. This light-use efficiency (LUE) GPP type of model not only

565 gives high importance to radiation but also uses EVI to calculate the fraction of PAR absorbed  
566 by chlorophyll, which helps to improve the representation of the seasonal variation of  
567 photosynthetic capacity (Zhang et al., 2017). Among the three models and the two global GPP  
568 datasets studied, LPJ-GUESS and VPM showed the closest agreement with our satellite-based  
569 results.

570 Land cover change has undoubtedly influenced the types of relationships observed in SIF  
571 and MAIAC EVI. The model output analyzed here uses a static land cover type and could,  
572 therefore, misrepresent some of the observed dynamics between climate and SIF and MAIAC  
573 EVI. Therefore, we recommend a careful reading of our comparison in regions with high levels  
574 of land cover change.

575 We identified clear, distinct patterns at large spatial and temporal scales in the observed  
576 relationships of tropical photosynthetic activity with precipitation and radiation, despite wide  
577 physiological, microclimatic and ecological variation. The large-scale nature of this study, at  
578 relatively low resolution, necessarily ignores important ecological factors and confounding  
579 variables associated with both climate and photosynthetic activity. Nutrient availability,  
580 microclimate, topography, soil texture, plant community dynamics, and anthropogenic  
581 disturbances are all important factors that influence photosynthetic activity and are not  
582 considered here. Yet, the broad spatial patterns identified here can guide research on the  
583 predominant mechanistic processes driving photosynthetic activity. Our evaluation of model and  
584 GPP products revealed the types of climate-vegetation relationships that are least accurately  
585 simulated by models. We expect that our classification of photosynthetic activity relationships  
586 will simplify model comparison and benchmarking for tropical ecosystems. Our characterization  
587 of the time lags shows the existence and extent of important delayed relationships of  
588 photosynthetic activity with each of the climate drivers in specific vegetation types and regions  
589 in the tropics. These results should guide modeling and experimental studies about the potential  
590 processes that determine seasonality of photosynthetic activity in the tropics. We believe this  
591 improved understanding could lead to more realistic predictions of how tropical vegetation will  
592 respond to ongoing climate change, with implications for the climate system.

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595 thank the TRENDY CLM4.5, JULES, and LPJ-GUESS modeling teams for contributing their  
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597 help with SIF data collection. We thank the Statistical Consulting Service at Purdue University  
598 for their assistance with the cross-correlation function analysis.

### 599 **Data Availability Statement**

600 Rainfall (TRMM, 2011) and radiation (Loeb et al., 2009) data were downloaded from GES DISC  
601 (<https://disc.gsfc.nasa.gov/>). Temperature data is available through Harris et al. (2014) and can  
602 be downloaded from CRU (<https://crudata.uea.ac.uk/cru/data/hrg/>). The Rainy and Dry Seasons  
603 (RADS) dataset is available through Bombardi et al. (2019). SIF data were downloaded from the  
604 GDFC data center ([https://avdc.gsfc.nasa.gov/pub/data/satellite/MetOp/GOME\\_F/](https://avdc.gsfc.nasa.gov/pub/data/satellite/MetOp/GOME_F/)). The GOSIF  
605 dataset is available through Li and Xiao (2019) and can be downloaded from the Global Ecology  
606 Group repository (<http://data.globalecology.unh.edu/data/GOSIF/>). MAIAC EVI data are  
607 available through Lyapustin et al. (2018) and can be downloaded from the NCCS Dataportal

608 (<https://portal.nccs.nasa.gov/datashare/maiac/DataRelease/Global-VI-8day-0.05degree/>). Model  
609 output from the TRENDY project is available through Sitch et al. (2015).

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*JGR: Biogeosciences*

Supporting Information for

**Seasonality of Tropical Photosynthesis: A Global Map of Drivers and Comparison to Model Outputs**

M. R. Uribe,<sup>1,2</sup> C. A. Sierra,<sup>3</sup> and J. S. Dukes<sup>1,2,4</sup>

<sup>1</sup> Department of Forestry and Natural Resources, Purdue University, West Lafayette, IN.

<sup>2</sup> Purdue Climate Change Research Center, Purdue University, West Lafayette, IN.

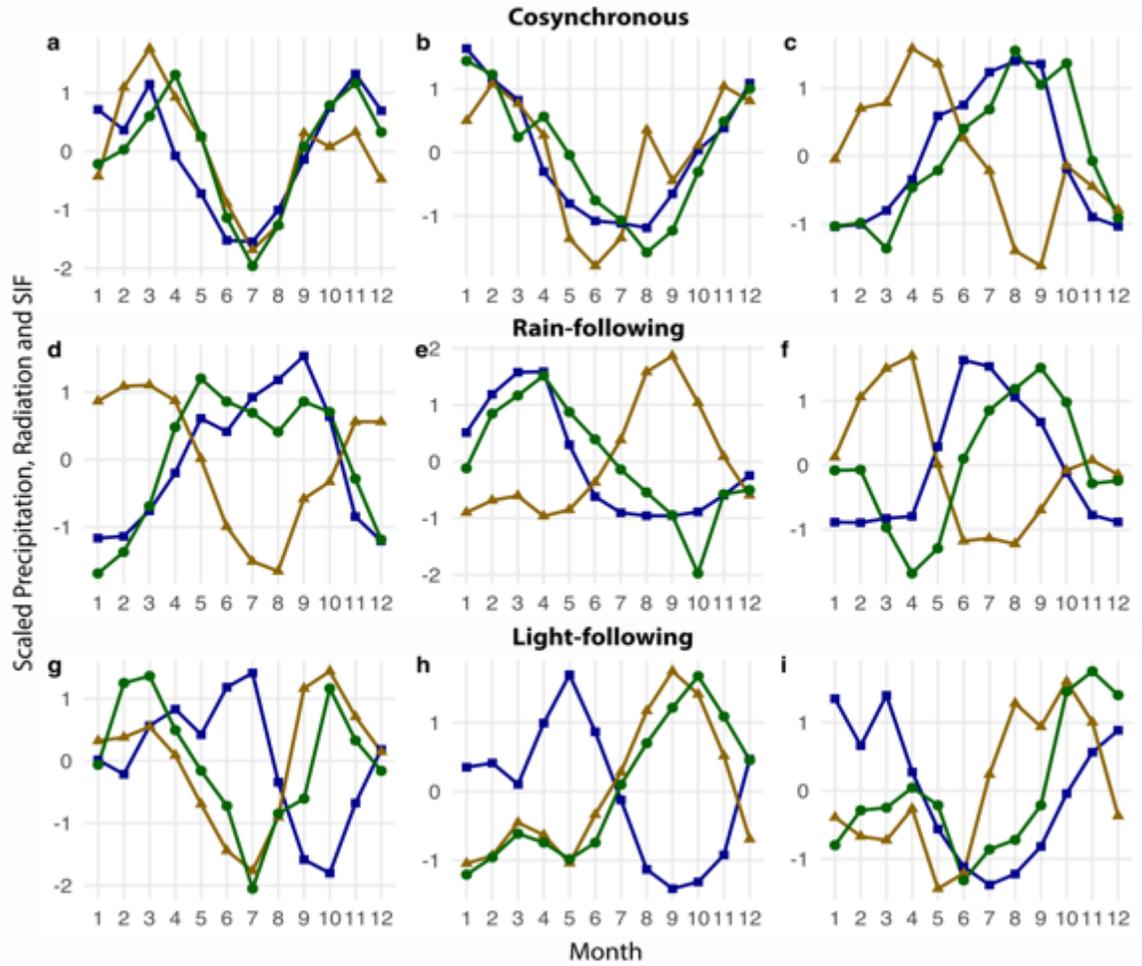
<sup>3</sup> Max Planck Institute for Biogeochemistry, Jena, Germany.

<sup>4</sup> Department of Biological Sciences, Purdue University, West Lafayette, IN.

**Contents of this file**

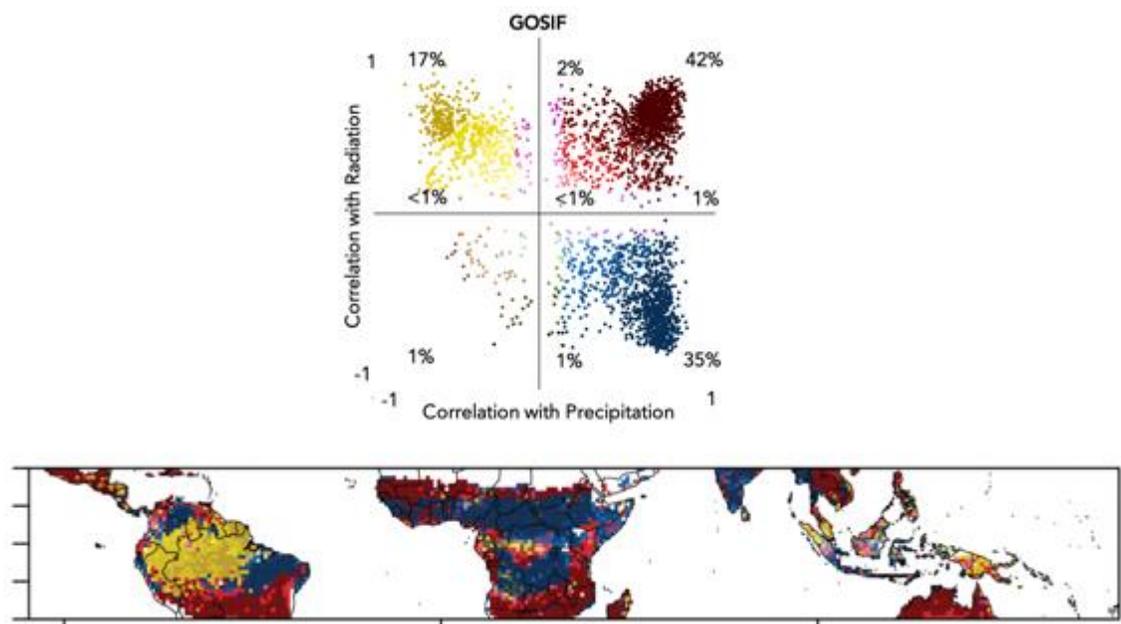
Figures S1 to S6

Tables S1 to S2

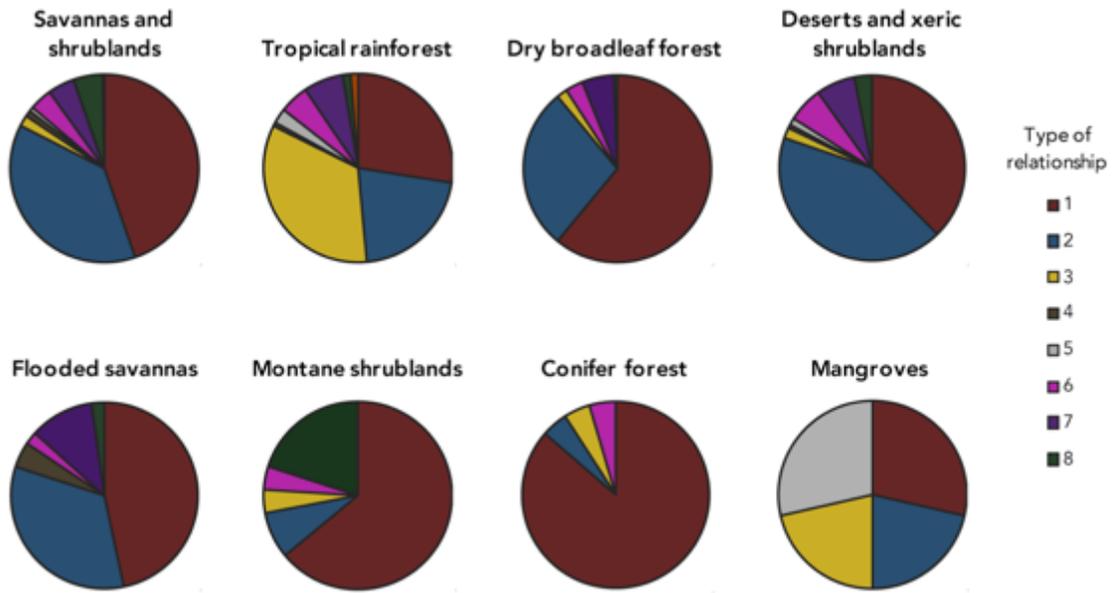


Relationship type	Plot	Lon	Lat	Biome	Location	Precipitation		Radiation	
						r	Lag	r	Lag
<b>Cosynchronous</b>	a	25.5	-2.5	Rainforest	Africa	0.59	0	0.55	0
	b	-54.5	-16.5	Savanna	S. America	0.58	1	0.48	2
	c	104.5	15.5	Dry Forest	Asia	0.72	1	0.58	4
<b>Rain-following</b>	d	21.5	7.5	Savanna	Africa	0.78	0	-0.66	0
	e	-45.5	-4.5	Xeric	S. America	0.47	1	-0.53	1
	f	95.5	17.5	Rainforest	Asia	0.66	1	-0.22	2
<b>Light-following</b>	g	121.5	-1.5	Rainforest	Asia	-0.29	1	0.54	0
	h	-52.5	4.5	Rainforest	S. America	-0.57	1	0.57	1
	i	-70.5	-8.5	Rainforest	S. America	-0.56	4	0.49	2

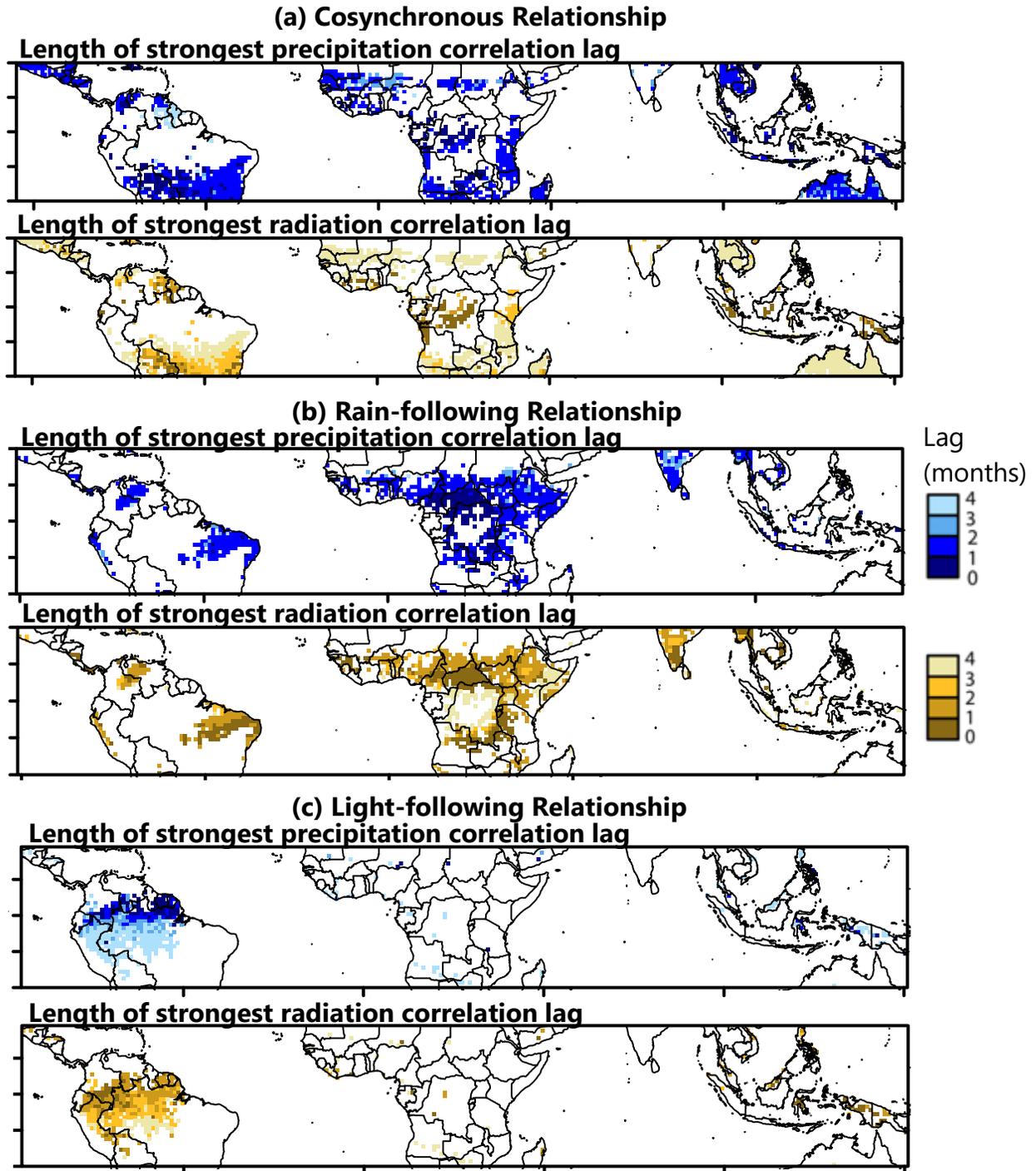
**Figure S1.** Seasonality profiles of Precipitation (blue), Radiation (yellow) and SIF (green) for the three main types of relationships. All data are scaled to fit and be comparable in the same plot. Relevant information of each site is provided in the accompanying table.



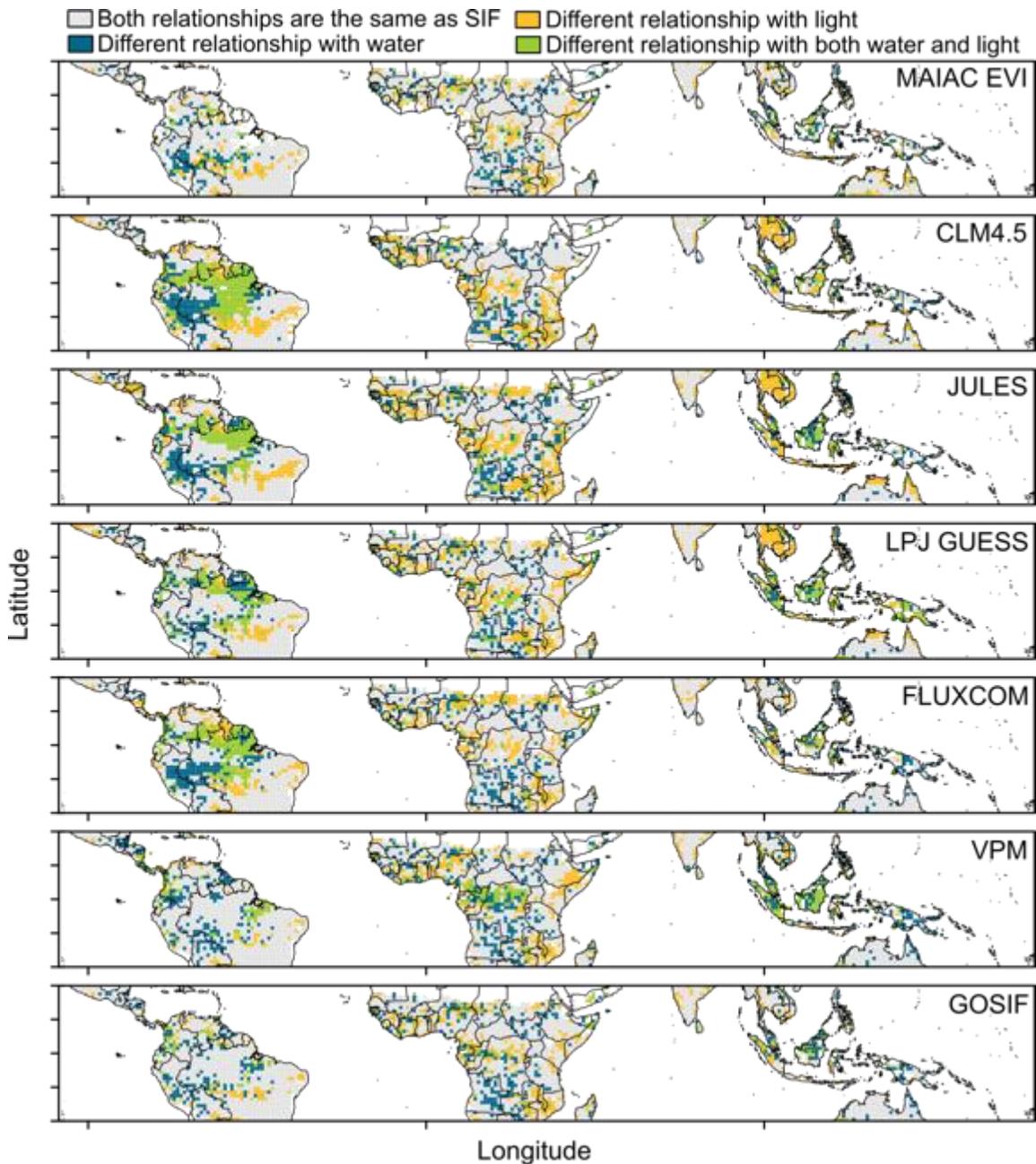
**Figure S2.** Scatterplot (upper panel) and map (lower panel) for GOSIF data, showing the maximum correlation coefficient from the CCF analysis for vegetation productivity from GOSIF with precipitation (x axis) and radiation (y axis). The numbers in the scatterplot indicate the percentage of pixels corresponding to the type of relationship where the number is located.



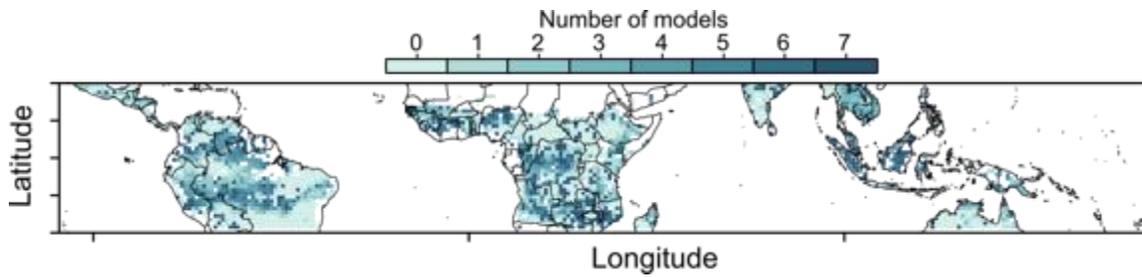
**Figure S3.** Types of relationships by biome (based on SIF results). The pie charts show the proportion of pixels with each type of relationship in each biome. Colors and numbers of the types of relationships (legend) correspond to the colors and numbers in the reference panel in Fig. 1.



**Figure S4.** Spatial distribution of the most strongly and significant correlated length of lag between SIF and precipitation and radiation, shown for the three most common types of relationships (cosynchronous, rain-following and light-following). Only lags are shown; that is, lags in which peaks in SIF follow peaks or troughs in the climate variable by 0-4 months. Precipitation and radiation lags are plotted separately for each type of relationship.



**Figure S5.** Comparison of correlation of photosynthetic activity with water and light for each dataset against SIF. For each pixel we show if the corresponding dataset or model agrees with SIF in the correlation between each driver and photosynthetic activity and the two drivers combined.



**Figure S6.** Number of models or datasets showing a different type of relationship than the one shown in SIF for each pixel. Zero indicates that all models or datasets are in agreement with SIF; seven indicates that all models disagree with SIF.

DATASET	ORIGINAL SPATIAL RESOLUTION (DEGREES)
MAIAC EVI	0.05 x 0.05
SIF	0.5 x 0.5
CLM 4.5	1.25 x 0.94
JULES	1.875 x 1.25
LPJ-GUESS	1 x 1
FLUXCOM	0.5 x 0.5
VPM	0.5 x 0.5
GOSIF	0.05 x 0.05

**Table S1.** Spatial resolutions of the datasets analyzed.

	MAIAC EVI	CLM4.5	JULES	LPJ-GUESS	FLUXCOM	VPM
<b>Rainforest</b>	$\kappa = 0.45$	0.19	0.24	0.30	0.28	0.34
	$D = 41.72$	62.78	58.56	52.62	53.78	45.36
<b>Dry forest</b>	0.53	0.26	0.23	0.33	0.48	0.50
	41.72	45.68	46.63	39.76	22.89	24.26
<b>Grasslands and savannas</b>	0.45	0.35	0.33	0.40	0.46	0.46
	31.64	40.29	41.75	34.91	32.85	30.7
<b>Flooded grasslands and savannas</b>	0.44	0.38	0.52	0.43	0.46	0.44
	28.57	34.15	26.19	30.61	28.57	30.61
<b>Montane grasslands and shrublands</b>	0.53	0.57	0.53	0.48	0.41	0.46
	18.0	23.08	22.45	22.64	25.0	20.75
<b>Deserts and xeric shrublands</b>	0.31	0.33	0.30	0.12	0.29	0.27
	33.75	34.38	41.48	35.63	39.24	31.48
<b>Mangroves</b>	0.31	-0.17	-0.08	0.00	0.28	0.15
	40.0	72.73	70.0	69.57	52.17	60.87

**Table S2.** Biome specific Kappa coefficients ( $\kappa$ ) and overall difference ( $D$ , %) between SIF and each of the other photosynthetic activity datasets (Fig. 2). Larger numbers indicate closer agreement between the results of two datasets. Larger  $D$  values indicate larger differences between the results of two datasets.