

Dislocation creep of olivine: Low-temperature plasticity controls transient creep at high temperatures

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Abstract

Transient creep occurs during geodynamic processes that impose stress changes on rocks at high temperatures. The transient is manifested as evolution in the viscosity of the rocks until steady-state flow is achieved. Although several phenomenological models of transient creep in rocks have been proposed, the dominant microphysical processes that control such behavior remain poorly constrained. To identify the intragranular processes that contribute to transient creep of olivine, we performed stress-reduction tests on single crystals of olivine at temperatures of 1250–1300°C. In these experiments, samples undergo time-dependent reverse strain after the stress reduction. The magnitude of reverse strain is $\sim 10^{-3}$ and increases with increasing magnitude of the stress reduction. High-angular resolution electron backscatter diffraction analyses of deformed material reveal lattice curvature and heterogeneous stresses associated with the dominant slip system. The mechanical and microstructural data are consistent with transient creep of the single crystals arising from accumulation and release of backstresses among dislocations. These results allow the dislocation-glide component of creep at high temperatures to be isolated, and we use these data to calibrate a flow law for olivine to describe the glide component of creep over a wide temperature range. We argue that this flow law can be used to estimate both transient creep and steady-state viscosities of olivine, with the transient evolution controlled by the evolution of the backstress. This model is able to predict variability in the style of transient (normal versus inverse) and the load-relaxation response observed in previous work.

1 **Dislocation creep of olivine: Backstress evolution** 2 **controls transient creep at high temperatures**

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12 tests, mantle viscosity

13 **Abstract**

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15 temperatures. The transient is manifested as evolution in the viscosity of the rocks until steady-state flow
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19 tests on single crystals of olivine at temperatures of 1250–1300°C. In these experiments, samples undergo
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22 diffraction analyses of deformed material reveal lattice curvature and heterogeneous stresses associated
23 with the dominant slip system. The mechanical and microstructural data are consistent with transient
24 creep of the single crystals arising from accumulation and release of backstresses among dislocations.
25 These results allow the dislocation-glide component of creep at high temperatures to be isolated, and we
26 use these data to calibrate a flow law for olivine to describe the glide component of creep over a wide
27 temperature range. We argue that this flow law can be used to estimate both transient creep and
28 steady-state viscosities of olivine, with the transient evolution controlled by the evolution of the
29 backstress. This model is able to predict variability in the style of transient (normal versus inverse) and
30 the load-relaxation response observed in previous work.

31 **1 Introduction**

32 The evolution of viscosity during flow of mantle rocks at high temperatures is fundamental to a variety of
33 geodynamic processes. For example, transient creep of the upper mantle has been identified as a major
34 contributor to geodetically observed surface deformations during post-seismic creep (Pollitz 2005; Freed,
35 Hirth, and Behn 2012; Masuti et al. 2016; Qiu et al. 2018), for which the strains are typically $< 10^{-3}$, and
36 inferred viscosities are one to two orders of magnitude lower than the long-term, steady-state viscosity.
37 Because transient viscosities also continue to evolve during postseismic deformation, they likely cause a
38 time-dependent transfer of stresses to neighbouring faults, rather than the instantaneous transfer assumed
39 by popular calculations of Coloumb stress changes (e.g., Freed 2005). Although sophisticated earthquake
40 forecast models do incorporate time-dependent loading according to average plate motion rates (e.g.,
41 Field et al. 2015, 2017), they still do not incorporate variable loading rates that would occur due to

42 transient creep of the lithosphere. In addition, transient viscosities are expected to be important, although
43 they have not yet been thoroughly considered, in other small-strain processes including flexure of the
44 lithosphere near volcanic loads (Zhong and Watts 2013) or in subducting slabs near trenches (Hunter and
45 Watts 2016), during which the strains rarely exceed 10^{-2} .

46 Although by far the bulk of investigations into the rheological behavior of the upper mantle have focused
47 on measuring and predicting steady-state viscosities in olivine attained under constant applied stress (e.g.,
48 Hirth and Kohlstedt 2003), some experiments have been conducted to investigate transient creep in
49 olivine. These experimental data confirm that the initial viscosities at the onset of creep or just after other
50 abrupt changes in stress are typically several orders of magnitude lower than the eventual steady-state
51 viscosities (Post 1977; Smith and Carpenter 1987; Hanson and Spetzler 1994; Chopra 1997). These rapid
52 changes in viscosity occur over the initial strain increments of $<10^{-2}$, and the most likely microstructural
53 changes responsible for this evolution are changes in the distribution and arrangement of dislocations
54 (Hanson and Spetzler 1994; S. Karato and Spetzler 1990; Durham, Goetze, and Blake 1977). However,
55 there are certain cases in which the initial viscosity is higher at the onset of creep and decreases with
56 progressive deformation, which is referred to as an “inverse” transient. Hanson and Spetzler (1994)
57 observed that single crystals of olivine can exhibit either a normal or an inverse transient, depending on
58 the orientation of loading. In addition, Cooper et al. (2016) observed that the type of transient can change
59 in a single experiment as the applied stress is gradually increased. This variability in observations
60 highlights that the specifics of the microstructural evolution and its link to both transient creep and
61 steady-state creep remain poorly constrained.

62 Due to the complexity of, and uncertainty regarding, potential micromechanical processes involved,
63 experimental observations of transient creep have generally been fit by empirical models in which the
64 underlying physics is not investigated. The Burgers model is often employed (Smith and Carpenter 1987;

65 Chopra 1997), amongst several others (Post 1977; Hanson and Spetzler 1994). However, the lack of a
66 rigorous microphysical basis to these models limits confidence in extrapolating them to predict
67 deformation at natural conditions under which the key processes potentially differ from those in
68 experiments. Moreover, as these processes and associated microstructural changes are poorly constrained,
69 model predictions cannot easily be tested against the microstructures of natural rocks. Thus, there remains
70 a need to establish the fundamental microphysics of transient creep.

71 Distinguishing among different models of transient creep is difficult. In general, the experimentalist is
72 tasked with fitting models to strain-time curves or strain rate-time curves, for which differences among
73 models are subtle (Post 1977; Hanson and Spetzler 1994). Additional constraints can potentially be
74 gleaned from the microstructures of deformed materials (Hanson and Spetzler 1994). However, the small
75 strains involved can result in microstructural changes beyond the resolution of some common
76 observational techniques, such as standard electron backscatter diffraction (EBSD) (Thieme et al. 2018).

77 We alleviate these problems by using stress-reduction experiments. These tests involve applying a rapid
78 reduction in stress to a sample undergoing steady-state creep and have been applied extensively to the
79 study of engineering materials (Blum and Weckert 1987; Biberger and Gibeling 1995; Poirier 1977).
80 Stress reductions typically induce an ensuing period of transient creep. As each test requires only a small
81 strain interval and causes only limited modification of the microstructure, multiple tests can be performed
82 on a sample within one experiment (e.g., with different magnitudes of stress reduction). Unlike increases
83 in stress magnitude, which are always followed by ongoing creep in the original direction, reductions in
84 stress can also result in a period of anelastic reverse strain (Blum and Weckert 1987; Biberger and
85 Gibeling 1995; Poirier 1977). As such, stress-reduction tests provide a wide variety of characteristics
86 against which to test models of transient creep.

87 Despite their potential, stress-reduction tests have been underutilized for geological minerals.
88 Stress-reduction tests have been applied to ice (Meyssonier and Goubert 1994; Caswell, Cooper, and
89 Goldsby 2015) and garnet (Z. Wang, Karato, and Fujino 1996). Applications to olivine have been limited
90 primarily to torsional microcreep experiments on fine-grained aggregates by Gribb and Cooper (1998)
91 and Jackson et al. (2002; 2014). Chopra (1997) reported two experiments on natural dunites, and Cooper
92 et al. (2016) reported one experiment on a single crystal. However, the microphysical controls on the
93 behaviour exhibited in such tests have not been systematically investigated.

94 Here we report stress-reduction experiments on single crystals of olivine to isolate and investigate
95 transient creep arising from intragranular processes. We analyse the microstructures of the deformed
96 samples using high-angular resolution electron backscatter diffraction (HR-EBSD) (Wilkinson, Meaden,
97 and Dingley 2006; D. Wallis et al. 2019), which is ideally suited to detecting subtle microstructural
98 changes, including lattice rotations and stress heterogeneities associated with geometrically necessary
99 dislocations (GNDs). We combine the results with previous rheological data to constrain the
100 microphysical processes that control the transient viscosity of rocks after rapid stress changes and make
101 predictions about the variability in rheological behavior that may be observed during transient creep in
102 Earth.

103 **2 Methods**

104 ***2.1 Sample preparation***

105 Four samples were prepared from gem-quality single crystals of San Carlos olivine. Rough crystals were
106 oriented using EBSD. The crystals were cut into rectangular parallelepipeds, with lengths in the range of
107 2 to 5 mm and aspect ratios in the range of 2:1 to 3:1, that were free from visible cracks and inclusions.

108 Samples were cut with long axes corresponding to the $[110]_c$, $[101]_c$, and $[011]_c$ orientations, following
109 the notation of Durham and Goetze (1977), with the intention of maximizing the resolved shear stress on
110 the $(010)[100]$, $(001)[100]$ and $(100)[001]$, and $(010)[001]$ slip systems, respectively. The end faces of
111 samples were ground flat and parallel using 1500 grit SiC paper.

112 **2.2 Deformation experiments**

113 Deformation experiments were conducted in a 1-atm uniaxial creep apparatus in the Rock Rheology
114 Laboratory at the University of Oxford. Details of the apparatus design are provided by Cao et al. (2020).
115 Samples were placed in direct contact with alumina platens, which were in direct contact with SiC
116 pistons. Although San Carlos olivine and alumina are not in chemical equilibrium, reaction products were
117 not visible after experiments, indicating that any reaction was limited in extent. The loading column is
118 housed inside a vertical tube furnace. The hot zone of the furnace is >5 cm, and temperature was
119 controlled with a thermocouple placed within 1 cm of the sample. The temperature difference across the
120 samples was $<1^\circ\text{C}$. The oxygen fugacity was controlled with a mixture of CO and CO_2 , and the ratio of
121 gas flow rates was maintained constant to set the oxygen fugacity at a value of 10^{-8} to $10^{-7.5}$ atm, near the
122 Ni:NiO buffer at the temperatures of the experiments. A gas-tight seal was maintained on the moving SiC
123 piston by means of frictionless oil bellows. An axial load was applied to the loading column with a
124 piezoelectric actuator using a closed-loop servo-control system, and the load was measured and controlled
125 with a precision of ~ 1 N. Changes in sample length were assessed by measuring displacements of the SiC
126 piston relative to the loading frame. Displacements were measured with a linear encoder with resolution
127 of 10 nm and corrected for distortion of the apparatus using the complex compliance of the apparatus
128 calibrated by Cao et al. (2020).

129 Samples were initially subjected to a small compressive load (~ 10 N) using the piezoelectric actuator to
130 ensure that contact was maintained with pistons and the sample remained upright and aligned throughout

the heating procedure. Samples were heated to temperatures of either 1250 or 1300°C over the course of approximately 3 hours. Deformation was not initiated until the measured temperature reached and maintained a steady value within 1°C of the set point.

Once the temperature stabilized, the load was gradually increased at a constant rate to the desired value, corresponding to stresses of 77 to 189 MPa, over the course of ~120 s. An example of the stress and strain evolution for experiment OxR0009 is illustrated in Figure 1. The load was maintained at the desired value, and the displacement was monitored until an approximately steady-state displacement rate was achieved, which typically required 2 to 5% strain. At this point, the load was reduced to a new set point at the maximum rate achievable by the apparatus, which was on the order of ~10 ms for the largest stress reductions. In order to record this deformation with sufficient resolution, the data sampling rate was increased from 0.1 Hz to 100 kHz just prior to the load reduction and then returned to 0.1 Hz just after the load reduction. This new load was maintained and the displacement monitored until a positive displacement rate was clearly observed. The load was then gradually increased back to the original value at a similar loading rate as in the initial loading increment. The load was then held constant until a steady-state displacement rate was again achieved, which was typically within a factor of two of the steady-state rate achieved in the first loading step. This cycle of load reduction followed by a return to the original load was carried out multiple times in each experiment. In each cycle, the load was reduced a greater amount than in the previous cycle, but after each load reduction, the load was always returned to the original control value. For several experiments, this progression of load reductions was repeated to test the reproducibility of the sample response. An overview of the experimental conditions are provided in Table 1.

Table 1: Summary of experiment conditions and results from stress-reduction experiments.									
Sample # Experiment #	Loading direction	Temperature (°C)	Initial stress (MPa)	Total strain (%)	Steady-state strain rate (10^{-5} s^{-1})	Best-fit backstress (MPa)	Best-fit hardening coefficient, γ	Transient	Symbol in figures
OxR0002	[101] _c	1250	162	17%	3.76	84	14	inverse	■

170120c									
OxR0007 170323	[110] _c	1250	189	12%	5.81	102	28	inverse	▶
OxR0009 170616	[110] _c	1300	105	17%	1.95	64	12	unclear	◀
OxR0009 171018	[110] _c	1300	77	38%	1.31	39	5	inverse	△
OxR0010 171016	[011] _c	1300	116	13%	0.12	104	55	normal	○

152

153 Experiments were stopped after an increment of deformation at the original load value. This final load
154 value was maintained until steady-state displacement rates were obtained so that the preserved
155 microstructure best reflects the microstructure just prior to each load reduction. At this point, the load was
156 removed entirely from the sample and the furnace was shut off. Cooling to ambient temperature required
157 several hours, but temperatures fell below 800°C in ~3 min, which is rapid enough to preserve the
158 microstructures characteristic of the highest stresses applied (Bai and Kohlstedt 1992b; Goetze and
159 Kohlstedt 1973).

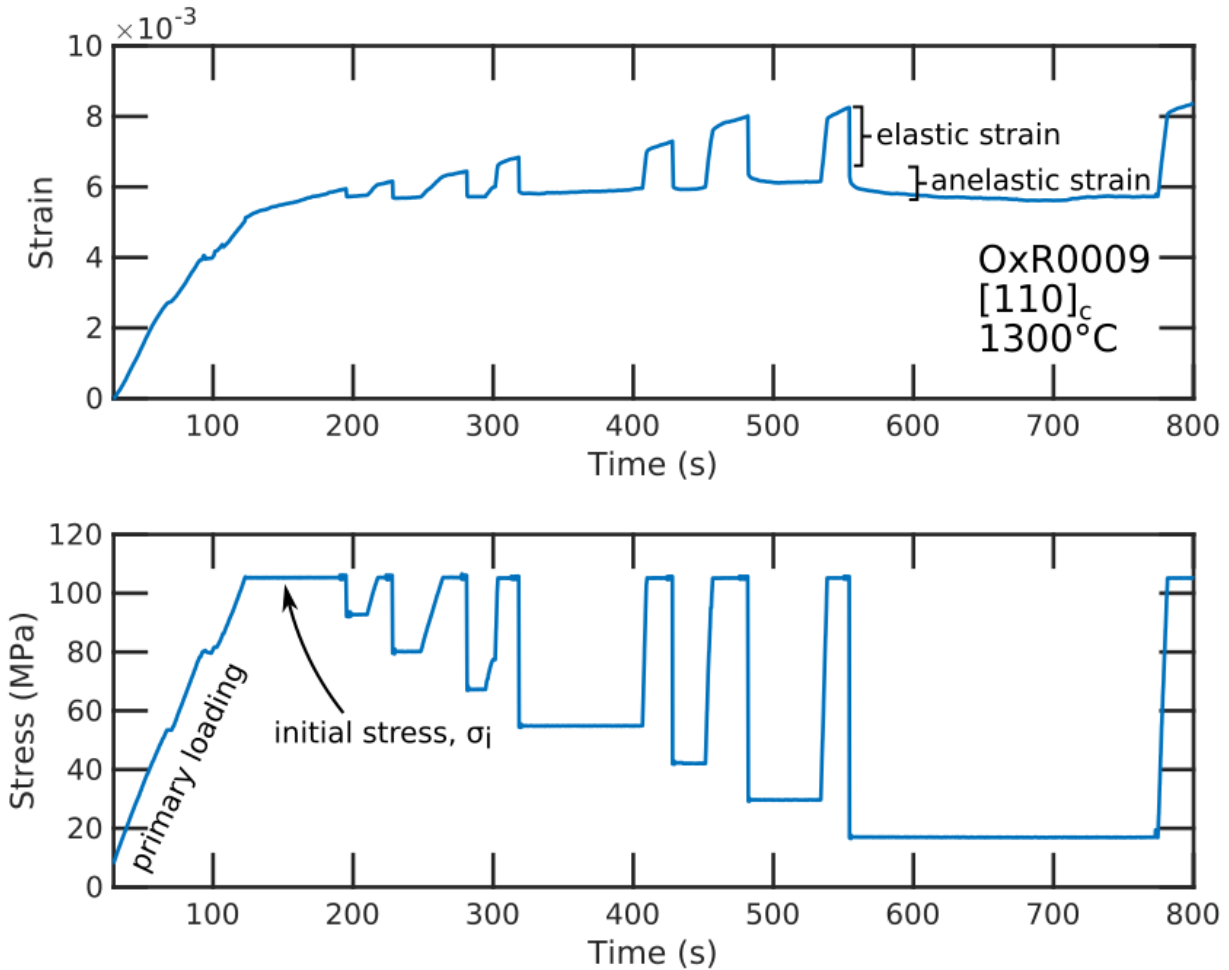


Figure 1: Axial strain and stress as functions of time for the first 800 s of deformation of sample OxR0009. Seven stress reductions are illustrated after the primary loading phase. The initial stress used to set the microstructure was 105 MPa. After each stress reduction, the reverse strain was composed of both an elastic and an anelastic component, as denoted by annotations. Elastic strains in this figure are a combination of the elastic distortion of both the apparatus and the sample.

160 2.3 Microstructural analysis

161 For one experiment (OxR0009, compressed along $[110]_c$), the microstructures within the sample were
 162 characterized after two segments of deformation. The experiment was initially interrupted at 7% strain. At
 163 this point, the face parallel to (001) was polished with successively finer diamond grits down to a grit size
 164 of 50 nm, providing a polish suitable for EBSD. An initial EBSD map was collected in an area of high

165 surface quality near the bottom of the polished vertical face of the sample. The sample was then reloaded
 166 into the apparatus and subjected to further increments of deformation. The experiment ended after a total
 167 of 38% strain. Two additional EBSD maps were collected on the same sample surface parallel to (001).
 168 One of these maps was located near the center of the sample and the other near a vertical edge.

169 The EBSD maps of sample OxR0009 were constructed using high-angular resolution electron backscatter
 170 diffraction (HR-EBSD). Data were acquired on an FEI Quanta 650 scanning electron microscope at the
 171 University of Oxford and a Philips XL30 scanning electron microscope at Utrecht University. Both
 172 instruments were equipped with an Oxford Instruments AZtec EBSD system and calibrated for HR-EBSD
 173 analysis by mapping undeformed Si standards following the procedure of Wilkinson et al. (2006). A
 174 routine to calibrate the chamber geometry, based on determining the pattern centre from diffraction
 175 patterns collected at a range of camera insertion distances, was performed before acquisition of each
 176 dataset. Conventions for reference frames in data acquisition and processing were validated using the
 177 approach of Britton et al. (2016). The details of each map are presented in Table 2.

Table 2: Details of HR-EBSD maps collected from sample OxR0009. The first map was collected after experiment 170616, and the subsequent maps were collected after experiment 171018. See Table 1 for details of the experiments.				
Total strain	Map location on sample surface	Step size (μm)	Number of points	Pixels in diffraction pattern
7%	Bottom edge	3.0	230 x 150	1344 x 1024
38%	Centre	0.3	300 x 250	1344 x 1024
38%	Side edge	0.5	230 x 160	1344 x 1024

178

179 Maps of lattice rotation and elastic strain heterogeneity were generated following the HR-EBSD
 180 post-processing method of Wilkinson et al. (2006) and Britton and Wilkinson (2011, 2012). 100 regions
 181 of interest in each diffraction pattern were cross correlated with a reference pattern chosen within each

map. Shifts in the regions of interest were used to estimate the deformation gradient tensor, which was decomposed into rotations and strains (assumed to be entirely elastic) of the diffraction pattern, at each point in the map. The pattern remapping approach of Britton and Wilkinson (2012) allows residual elastic strain measurements in the presence of lattice rotations of several degrees. Stresses were estimated from the elastic strain measurements following the approach of Britton and Wilkinson (2012) and employing the elastic constants for olivine at room temperature from Abramson et al. (1997). Measured elastic strains are relative to the unknown strain state of the reference point and therefore provide maps of stress heterogeneity. We rotated measured stress tensors from the reference frame of the EBSD maps into the reference frames of the measured crystal orientations to obtain the shear stresses resolved on each slip system (David Wallis et al. 2017). A comprehensive overview of all data acquisition and processing procedures for HR-EBSD is provided by Wallis et al. (2019).

3 Results

3.1 General creep behavior

Figure 1 presents stress and strain as a function of time for the first 800 s of experiment 170616 with sample OxR0009, which is a typical example of the response of all samples. In this particular experiment, the strain rate did not significantly evolve, but was instead approximately constant over the first ~1% of strain (120 to 200s in Figure 1). In other experiments, the strain rate after initial loading was observed to decrease or increase with progressive strain before a steady-state strain rate was reached, which we refer to as “normal” or “inverse” transients, respectively. Table 1 indicates the type of transient observed for each experiment. Figure 1 also illustrates that, once a steady-state strain rate was clearly achieved, stress reduction tests were executed. At the end of each reduction test, the stress was increased to approximately

the same value as the initial stress. During these high-stress intervals, the strain rate returned to a similar strain rate to that just prior to the initial stress reduction.

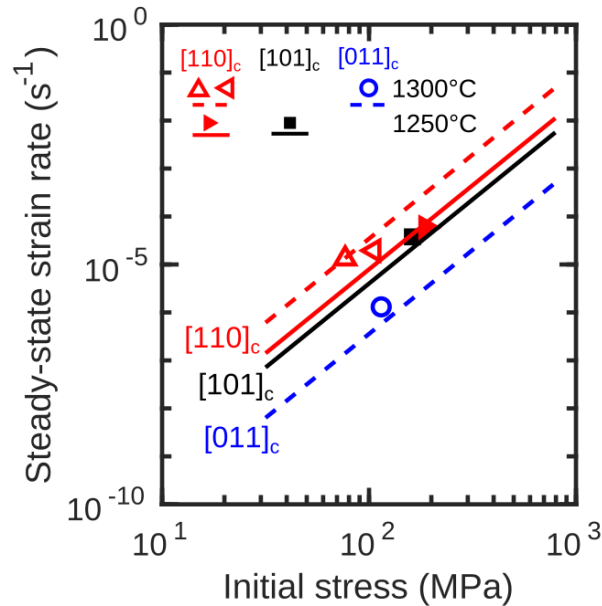


Figure 2: Comparison between strain rates from the steady-state portions of the experiments and published flow laws for single crystals of San Carlos olivine. Symbols represent data from this study (following Table 1), and lines represent flow laws from Bai and Kohlstedt (1992a). Color indicates crystal orientation and symbol indicates experiment. Open symbols and dashed lines indicate a temperature of 1300°C, and filled symbols and solid lines indicate a temperature of 1250°C.

Figure 2 compares the steady-state strain rate from each experiment to predictions of the flow laws for single crystals of San Carlos olivine from Bai and Kohlstedt (1992a). These data are also reported in Table 1. In general, samples loaded in the $[110]_c$, $[101]_c$, and $[011]_c$ orientations exhibit good agreement with the flow-law predictions for olivine deformed at 1250°C and 1300°C. The maximum differences between measured and predicted strain rates are on the order of a factor of two.

210 **3.2 Stress-reduction tests**

211 The results of individual stress-reduction tests are compiled in Figure 3. To obtain these data, we partition
 212 each time series of strain into ranges spanning the point just before the stress reduction to the point just
 213 before the stress was increased back to the initial stress. The strain at the beginning of this range (i.e., just
 214 before the reduction) was set to 0, and the strain data were then corrected for the compliance of the
 215 apparatus. This corrected strain after the stress reduction is presented as a function of time in the left
 216 column of Figure 3. Small stress reductions result in a small elastic lengthening (negative strain) of the
 217 sample that is immediately followed by continued shortening (positive strain). Large stress reductions
 218 result in correspondingly large elastic lengthening of the sample, but that lengthening is followed by an
 219 additional, time-dependent lengthening of the sample that we refer to as anelastic reverse strain. The
 220 period of anelastic lengthening lasts for 10s to 100s of seconds depending on the magnitude of the stress
 221 reduction.

222 We further analyze these data following the method of Blum and Finkel (1982). Many early analyses of
 223 stress-reduction experiments (e.g., Ahlquist and Nix 1971) relied on measuring strain rates after the
 224 anelastic reverse strain as a means to assess the “constant microstructure” mechanical behavior. Similar
 225 methods have been applied to some geological materials such as ice (Caswell, Cooper, and Goldsby
 226 2015). Unfortunately, strain rates are generally not constant during this increment of stress-reduction tests,
 227 and therefore identifying the best segment for calculating a strain rate is not trivial. In contrast, the
 228 method of Blum and Finkel (1982) does not rely on measuring strain rates but instead on measuring the
 229 magnitude of anelastic reverse strain after the stress reduction and prior to subsequent forward creep. We
 230 identify the total amount of reverse strain (both elastic and anelastic) using the minimum strain recorded
 231 after the stress reduction. Based on the noise in the strain data, this selection process may overestimate the
 232 total reverse strain by $\sim 10^{-4}$ (roughly 10% of the maximum anelastic reverse strain). We also identify the

233 elastic component of the reverse strain as the strain at the first data point after the reduced stress level is
234 reached.

235 The total and elastic reverse strains are presented for all experiments in the middle column of Figure 3.
236 For most experiments, the elastic strains are linearly correlated with the magnitude of the stress change
237 and in good agreement with the strains predicted given the Young's modulus in the specific crystal
238 orientation (dashed lines). Agreement with the elastic modulus provides confidence in the apparatus
239 correction and the method of identifying the elastic strain. However, for samples deformed at smaller
240 initial stresses, some elastic strains are smaller than predicted by the elastic modulus (e.g., OxR0009b and
241 OxR0010), which may indicate the apparatus compliance is overestimated at small loads.

242 We take the difference between the total reverse strain and the elastic reverse strain to represent the
243 anelastic reverse strain, which is plotted as a function of the magnitude of the stress reduction in the right
244 column of Figure 3. Crystals deformed in the $[110]_c$ and $[101]_c$ orientation exhibit maximum anelastic
245 reverse strains $>10^{-3}$, but significant anelastic reverse strain is only observed following stress reductions
246 greater than $\sim 40\%$ of the initial stress. In contrast, the crystal deformed in the $[011]_c$ orientation exhibits
247 anelastic reverse strains that gradually increase with the magnitude of the stress reduction, even for small
248 stress reductions.

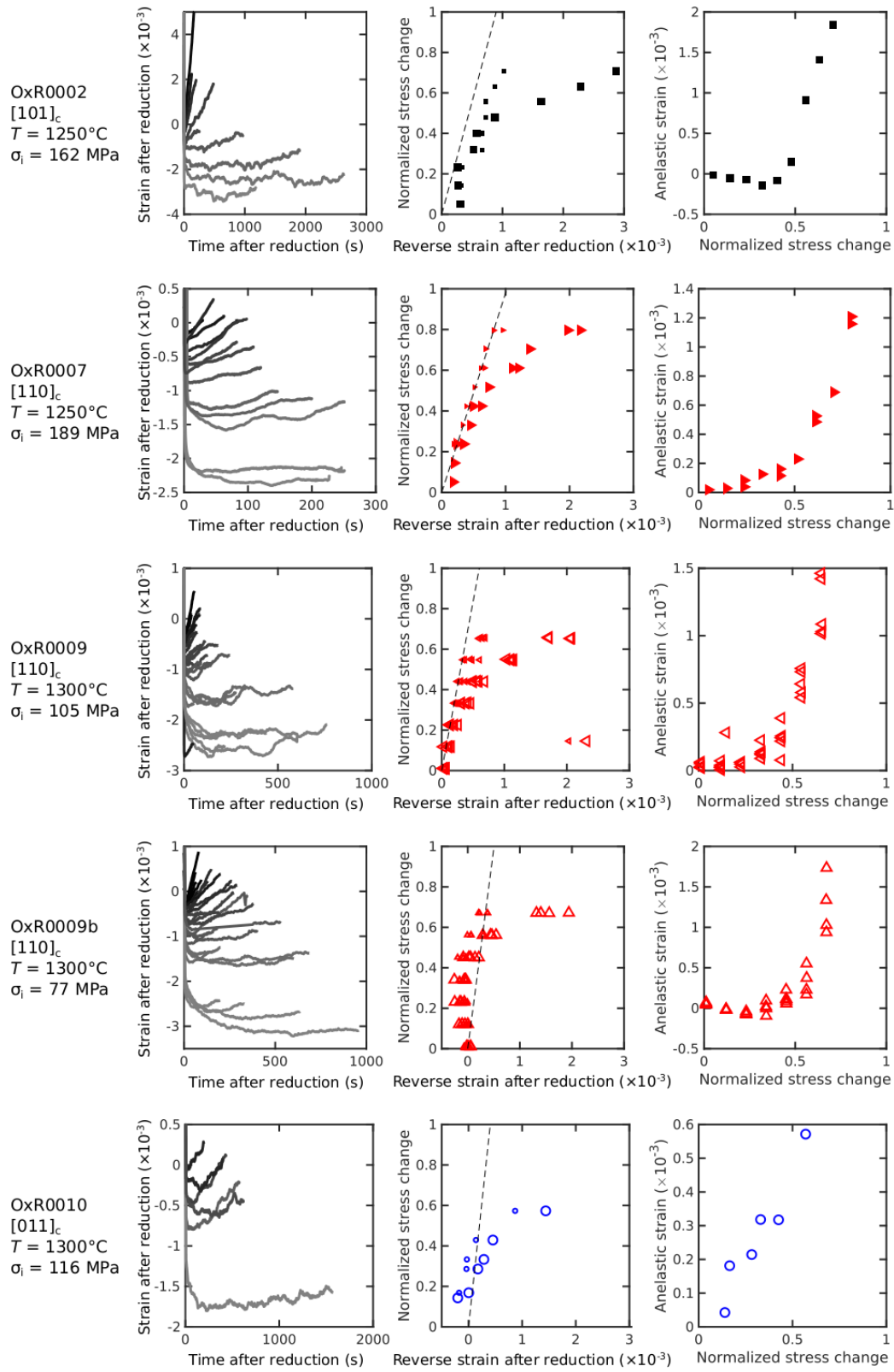


Figure 3: Compilation of data from all stress-reduction tests. (left) Strains after each reduction are aligned and plotted in grayscale. The grayscale indicates the order of the reductions, with darker colors indicating stress reductions earlier in the experiment. (middle) Magnitude of the stress reduction as a function of the strain after the reduction. Small markers indicate the elastic reverse strain, and larger markers indicate the total reverse strain. The dashed line indicates the Young's modulus for each crystal orientation based on the data of Abramson et al. (1997). (right) The anelastic reverse strain as function of the magnitude of the stress reduction. Anelastic reverse strains are calculated as the difference between large markers and small markers in the middle column. Marker shapes and colors are as designated in Table 1.

249 **3.3 Microstructures**

250 Figure 4 presents HR-EBSD maps of sample OxR0009. The dataset collected near the bottom edge of the
 251 sample after deformation to 7% strain exhibits lattice rotations dominated by the ω_{12} component. The map
 252 of this component exhibits a smooth gradient from the upper-right to lower-left, i.e., in the [100]
 253 direction. For the mapped crystal orientation, the ω_{12} component corresponds to rotations around the [001]
 254 axis, and large magnitudes of that component are therefore consistent with the presence of (010)[100]
 255 edge dislocations (David Wallis et al. 2016). Signal in the maps of the ω_{13} and ω_{23} components has
 256 magnitudes many times lower than that of ω_{12} , precluding major contributions from other dislocation
 257 types to the mapped lattice curvature. The map of shear stress resolved on the (010)[100] slip system (i.e.,
 258 σ_{ab}) also exhibits gradients in the [100] direction, varying by magnitudes on the order of 100 MPa over
 259 distances of a few hundred microns. The dataset collected from near the centre of the sample at the final
 260 strain of 38% exhibits lattice rotations similar to those in the dataset collected at lower strain. The stress
 261 distribution is homogeneous over the length scale of the map, which is notably smaller than that collected
 262 at 7% strain. However, the dataset collected near a side edge of the sample at 38% strain exhibits more
 263 discontinuous rotation and stress fields. Each component of the lattice rotation exhibits two linear
 264 discontinuities corresponding to subgrain boundaries. The boundaries have traces parallel to [010] and are
 265 most pronounced in the map of ω_{12} , consistent with them being dominated by (010)[100] edge
 266 dislocations. The boundaries also coincide with discontinuities in the map of σ_{ab} . These stresses have the

greatest magnitudes, i.e., are localised, adjacent to the boundaries. Across the boundaries, the stresses exhibit changes in sign and changes in magnitude on the order of 100 MPa. The orientation gradients between the upper-right and lower-left corners of each map of ω_{12} correspond to average GND densities in the range $4 \times 10^{11} \text{ m}^{-2}$ to $8 \times 10^{11} \text{ m}^{-2}$ (David Wallis et al. 2016). We emphasize that subgrain boundaries are essentially not observed at low strains, and at high strains are only observed in small portions of the sample.

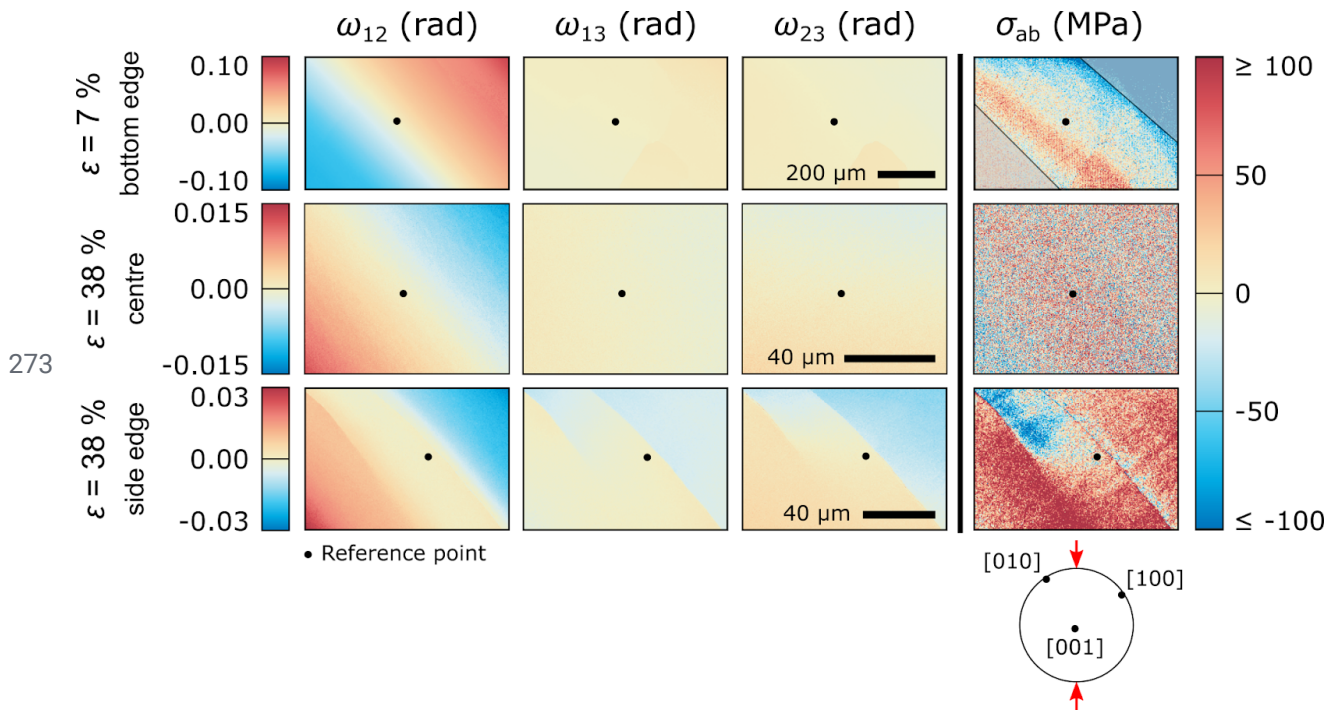


Figure 4: HR-EBSD maps of sample OxR0009, which was compressed in the $[110]_c$ orientation indicated in the pole figure, at different strains and locations in the sample annotated on the left. Lattice rotations (ω_{ij}) and shear stress on the $(010)[100]$ slip system (σ_{ab}) are relative to the orientation and stress state, respectively, of the reference points marked in black. The greyed-out areas in the top stress map indicate regions where the stress measurements are unreliable due to the large distances and large lattice rotations relative to the reference point.

280 **4 Discussion**

281 In this study we present results of stress-reduction tests carried out on single crystals of San Carlos
 282 olivine. The key observations are: 1) prior to stress reductions, strain rates reach a steady state in
 283 agreement with previously published flow laws for olivine, 2) after stress reductions of sufficient
 284 magnitude, an anelastic reverse strain is observed, 3) the magnitude and duration of anelastic reverse
 285 strain is a strong function of the magnitude of the stress reduction, 4) significant heterogeneity in the
 286 residual stress is correlated with gradients in lattice orientation (i.e., with GNDs). To interpret these
 287 observations, we first compare previously published interpretations of stress reduction tests to our data
 288 and then reanalyze our data in the context of recent work on the plastic deformation of olivine.

289 ***4.1 Comparison to previous interpretations of stress-reduction tests***

290 Stress-reduction tests have been a common technique for investigating the microphysics of creeping,
 291 crystalline materials. Two primary aspects of the microphysics have been investigated with stress
 292 reduction tests. The first is the concept of a backstress (or internal stress), which is related to the details of
 293 the microstructural evolution during deformation. The observation of backstress in single crystals
 294 indicates that it arises from the long-range elastic interactions among dislocations. Dislocation arrays can
 295 result in a net Burgers vector with a long-range internal stress field that opposes the macroscopically
 296 applied stress, hence the term backstress. As described in detail by Bayley et al. (2006), the net backstress
 297 associated with an array of dislocations is strongly dependent on the spatial gradients in dislocation
 298 density, rather than simply the number of dislocations, and therefore the backstress is sensitive to the full
 299 dislocation microstructure, including mobile dislocations, subgrain boundaries, and their spatial
 300 distributions. The macroscopic creep rate is assumed to result from the action of an effective stress, σ_e ,
 301 which is the difference between the externally applied stress and the net backstress, $\sigma_e = \sigma - \sigma_b$, where σ is

302 the former and σ_b is the latter. Macroscopic forward creep results from positive values of σ_e , and reverse
 303 creep results from negative values of σ_e .

304 Early experiments on pure metals by Gibbs (1966), solid-solution alloys by Ahlquist and Nix (1971), and
 305 ionic crystals by Menezes and Nix (1971) demonstrated that the average backstress could be determined
 306 with stress-reduction tests (for a review, see Takeuchi and Argon 1976). In this method, the microstructure
 307 and associated σ_b are presumed to remain constant during the nearly instantaneous stress reduction. Just
 308 after the reduction, the immediate strain rate is observed to be positive, negative, or approximately zero,
 309 depending on the magnitude of the stress reduction. These scenarios are presumed to indicate that the
 310 applied stress after the reduction is either greater than, less than, or equal to the backstress, respectively.
 311 Therefore, the reduced stress at which the subsequent strain rate is approximately zero is taken to be equal
 312 to the backstress. However, there is some controversy about the practical aspects of identifying a zero
 313 strain rate (for a review, see Nix and Ilshner 1979), and therefore we adopt the method of Blum and
 314 Finkel (1982), as described above. In this latter method, the reduced stress at which anelastic reverse
 315 strain becomes measurable is taken as equal to the average backstress.

316 These methods have been used to characterize net backstresses in a wide range of materials. Two
 317 well-cited studies, in particular, had remarkably consistent results, estimating backstresses at 52% of the
 318 initial stress in both pure metals (Evans, Roach, and Wilshire 1985) and solid-solution alloys (Blum and
 319 Finkel 1982). Many subsequent modeling efforts therefore have taken this value as a constant in the creep
 320 of metals (e.g., Biberger and Gibeling 1995). Similar values have been observed in some ceramics such as
 321 $\text{Gd}_3\text{Ga}_5\text{O}_{12}$ (Z. Wang, Karato, and Fujino 1996), which exhibited a backstress of 62% of the initial stress.
 322 However, broader compilations of backstress estimates (Takeuchi and Argon 1976; Argon and Takeuchi
 323 1981; Čadek 1987; Milička, Trojanová, and Lukáč 2007) demonstrate that the relationship between the
 324 backstress and initial stress varies systematically as a function of the initial stress. For instance, Menezes

325 and Nix (1971) reported values of the backstress in LiF that vary systematically between 21% and 83%.
326 In addition, Wang et al. (1993) reported backstresses in SrTiO₃ up to ~100%.

327 The second aspect of the microphysics of creep investigated with stress reduction tests is the rate limiting
328 process at constant microstructure. As noted above, the specific dislocation arrangements are considered
329 to be unchanged (neglecting unbowing of dislocations) during the stress reduction. Therefore, in the case
330 that the stress reduction is small and no anelastic strain occurs, the forward strain rate just after the
331 reduction is a function of the same microstructure as before the reduction. A variety of studies (for a
332 review, see Biberger and Gibeling 1995) have evaluated these types of tests and found that the strain rates
333 after a stress reduction are fit well by a flow law with an exponential dependence of strain rate on the
334 stress, which is one member of a class of equations used to describe plastic deformation limited by the
335 glide velocity of dislocations (e.g., Kocks, Argon, and Ashby 1975; Weertman 1957). Furthermore, the
336 dependence of the strain rate on the reduced stress can be related to the apparent activation area of gliding
337 dislocations (the area of slip plane traversed in a single activation event), which was shown to have a
338 direct relationship to the observed spacing of free dislocations in subgrain interiors. These results led the
339 authors (and authors cited therein) to conclude that the strain rates directly after a load reduction are
340 dominated by the kinetics of gliding dislocations rate limited by the lattice friction (i.e., the
341 Peierls-Nabarro mechanism). Similar conclusions based on stress-reduction tests have been drawn for LiF
342 (Müller, Biberger, and Blum 1992), SrTiO₃ (Z. Wang, Karato, and Fujino 1993), and Gd₃Ga₅O₁₂ (Z.
343 Wang, Karato, and Fujino 1996).

344 The observations described above have played a large part in informing the development of composite
345 models to describe both transient and steady-state creep properties of crystalline materials. Key early
346 models include those proposed by Hart (1976), Nix and Ilshner (1979), and Mughrabi (1980), although a
347 variety of elaborations have been proposed (e.g., Korhonen, Hannula, and Li 1987; Garmestani, Vaghar,

and Hart 2001; Bammann 1989). The key features of these models are that 1) plastic deformation across a wide range of conditions is effectively controlled by the glide of dislocations, 2) the glide velocity of those dislocations is moderated by the long-range stresses associated with heterogeneity in the state variables that describe the microstructure (what we refer to as a net backstress), and 3) the macroscopic behavior can be modified by different recovery mechanisms (such as dislocation climb and annihilation) that modify the backstress. We use these three concepts as a foundation for analyzing our data for olivine in the next section.

4.2 Linking stress reductions to backstress evolution and dislocation-glide kinetics

Our working hypothesis based on the discussion above is that deformation after a stress reduction is controlled by the glide of dislocations, and the anelasticity observed after larger stress reductions is controlled by the net backstresses associated with the long-range stress fields of the dislocation microstructure. Plastic deformation and the evolution of backstresses were recently investigated for olivine by Hansen et al. (2019). In that work, the authors used experiments conducted at temperatures as low as room temperature to examine the yield strength and hardening behavior of single crystals and polycrystalline samples of San Carlos olivine. Notably, the hardening behavior of single crystals is indistinguishable from the hardening behavior of polycrystalline aggregates, which we interpret to indicate that the dominant processes leading to strain hardening are intragranular. Furthermore, they observed a distinct Bauschinger effect, which supports the conclusion that the hardening is associated with long-range interactions of dislocations (see section 2.5 in Kassner 2015). This conclusion is further supported by microstructural correlations of the GND density field with residual stress fields in the same samples (David Wallis, Hansen, Kumamoto, et al. 2020). In support of the hypothesis that these long-range stress fields are also important during deformation at high temperatures, we see spatial

371 correlations between lattice rotation gradients (i.e., GND densities) and stresses in our samples (Figure 4)
 372 and the similar samples of Wallis et al. (2017). Furthermore, Wallis et al. (2020) highlighted the
 373 similarities between stress fields in aggregates of olivine deformed at temperatures of 1150–1250°C and
 374 those deformed at room temperature. These similarities include the magnitudes, the forms of the
 375 probability distributions, and the characteristic length scales of the stress heterogeneity, along with a
 376 causative link to GNDs in both temperature regimes.

377 4.2.1 Magnitude of anelastic reverse strain

378 To further analyze these data, we assess the magnitude of anelastic reverse strain after a stress reduction.
 379 We use the equations describing plasticity in olivine outlined by Hansen et al. (2019). The basic rate
 380 equation is derived following the classical treatment (e.g., Weertman 1957) and assumes the dislocation
 381 velocity is proportional to a Boltzman distribution ($\exp\left[\frac{\Delta G}{RT}\right]$, where ΔG is the Gibbs free energy, R is the
 382 gas constant, and T is the temperature) describing the probability of the dislocation overcoming the
 383 energy barrier to its motion. Under application of a stress, ΔG becomes stress dependent. Hansen et al.
 384 (2019) used a common simplification that ΔG is a linear function of the applied stress, although we
 385 recognize that a variety of nonlinear relationships have been proposed to describe energy barriers, such as
 386 the lattice friction, of different shapes (chapter 4 in Caillard and Martin 2003; chapter 6 in Suzuki,
 387 Takeuchi, and Yoshinaga 2013; p. 141 in Kocks, Argon, and Ashby 1975). Additionally, Hansen et al.
 388 (2019) used sinh instead of an exponential, which was originally proposed by Garofalo (1963), and later
 389 given physical justification by considering the potential for forward and backward fluctuations of the
 390 dislocation (Wu and Krausz 1994; Krausz 1968; Wilson and Wilson 1966). These considerations yield a
 391 flow law for the plastic strain rate from gliding dislocations, $\dot{\epsilon}_p$, of the form

$$\dot{\epsilon}_p = A_1 \exp\left(-\frac{\Delta F}{RT}\right) \sinh\left(\frac{\Delta F}{RT} \frac{\sigma - \sigma_b}{\Sigma}\right), \quad (1)$$

392 where A_1 is a constant with units of s^{-1} , ΔF is the activation energy, σ is the applied stress, σ_b is the
 393 backstress, and Σ is stress to overcome the barrier at 0 K. Equations of this form or similar can be used to
 394 describe a wide variety of barriers to dislocation motion, such as precipitates, forest dislocations, or grain
 395 boundaries in the general context of “low-temperature plasticity” (chapter 2 in Frost and Ashby 1982).
 396 However, in regards to creep of single crystals of olivine, we are primarily concerned with the barrier
 397 presented by the fundamental motion of dislocations through the lattice (aka, lattice friction or the
 398 Peierls-Nabarro mechanism) by, for instance, the formation and migration of kink pairs. In this case, Σ is
 399 referred to as the Peierls stress.

400 As described above, the effective stress ($\sigma - \sigma_b$) determines the glide velocity of dislocations, and
 401 therefore any evolution of the strain rate during a test at constant stress results from the evolution of σ_b . It
 402 is important to note that $\dot{\epsilon}_p$ will be negative if $\sigma_b > \sigma$. Hansen et al. (2019) provided an evolution equation
 403 for σ_b ,

$$\frac{d\sigma_b}{d\epsilon_p} = \gamma [\sigma_{b,max} - \text{sgn}(\dot{\epsilon}_p) \sigma_b], \quad (2)$$

404 where γ is a rate constant and $\sigma_{b,max}$ is the maximum backstress. That is, if there is zero initial backstress,
 405 the initial rate of increase is given by $\gamma\sigma_{b,max}$, and the steady-state backstress is $\sigma_{b,max}$. Although they
 406 implemented this equation in an empirical manner, there are physical meanings to each of these terms
 407 (e.g., Mecking and Kocks 1981). Assuming the backstress is related to the dislocation density (e.g.,
 408 Taylor 1934), $\gamma\sigma_{b,max}$ is related to the dislocation nucleation rate, and γ is related to the dislocation
 409 recovery rate associated with dynamic recovery, which operates at high stresses and low temperatures. At
 410 elevated temperatures, additional recovery terms should be included to account for effects such as
 411 dislocation climb or grain-boundary migration. The explicit links between backstress evolution and the
 412 dislocation microstructure in olivine will be explored in future work.

413 An analytical solution exists for equation 2, providing the evolution of the back stress as a function of
 414 strain,

$$\sigma_b(\varepsilon_p) = \text{sgn}(\dot{\varepsilon}_p) \sigma_{b,\max} + (\sigma_{b,0} - \text{sgn}(\dot{\varepsilon}_p) \sigma_{b,\max}) \exp(-\gamma \varepsilon_p \text{sgn}(\dot{\varepsilon}_p)), \quad (3)$$

415 where $\sigma_{b,0}$ is the initial back stress. This equation can be rearranged to predict the total plastic strain
 416 necessary for σ_b to reach a final value, $\sigma_{b,f}$

$$\varepsilon_p = [\text{sgn}(\dot{\varepsilon}_p) \gamma]^{-1} \ln \left[\frac{\sigma_{b,f} - \text{sgn}(\dot{\varepsilon}_p) \sigma_{b,\max}}{\sigma_{b,0} - \text{sgn}(\dot{\varepsilon}_p) \sigma_{b,\max}} \right]. \quad (4)$$

417 In the case of a large stress-reduction test, $\dot{\varepsilon}_p < 0$ just after the reduction. The backstress at the time of the
 418 reduction, $\sigma_{b,0}$, is some fraction, β , of the initial applied stress prior to the reduction, σ_i . After the
 419 reduction, the backstress will decrease until a value of $\sigma_{b,f}$ is reached. This final value cannot, by
 420 definition, be larger than the applied stress after the reduction, σ_r , and it will only reduce further than σ_r if
 421 some additional recovery process is involved, such as “static” annealing of the dislocation density (e.g.,
 422 annihilation by dislocation climb). Based on published static annealing rates for mobile dislocations in
 423 olivine (Toriumi and Karato 1978; S.-I. Karato, Rubie, and Yan 1993) as a proxy for relaxation of the
 424 backstress, estimates for the total recovery for the longest durations of anelastic strain (100s of seconds)
 425 range from <1% to ~10%. Therefore, we assume that static recovery is negligible during the anelastic
 426 portion of these experiments, and we expect that $\sigma_{b,f} = \sigma_r$ at the end of anelasticity and prior to subsequent
 427 forward creep. These considerations allow us to rewrite equation 4 for the maximum anelastic reverse
 428 strain after a stress reduction as

$$\varepsilon_p = -\frac{1}{\gamma} \ln \left(\frac{\sigma_r + \sigma_{b,\max}}{\beta \sigma_i + \sigma_{b,\max}} \right). \quad (5)$$

429 Note that the assumption made above is only valid for the case in which anelastic reverse strain occurs
430 (i.e., when $\sigma_r < \beta\sigma_i$). If the applied stress is not reduced below this threshold, then no anelastic reverse
431 strain occurs (i.e., $\varepsilon_p = 0$).

432 Equation 5 can be compared to our experimental data to find values for key parameters. We assume that
433 $\sigma_{b,max} = 1.8$ GPa according to the results of our experiments at low temperatures (Hansen et al. 2019). We
434 then compare the measured anelastic reverse strains for large stress reductions to equation 5 to find the
435 best-fit values of γ and β (and therefore the backstress prior to the stress reduction from equation 4).
436 Values for these parameters are presented in Table 1 and a comparison of equation 5 to anelastic reverse
437 strains is presented in Figure 5.

438

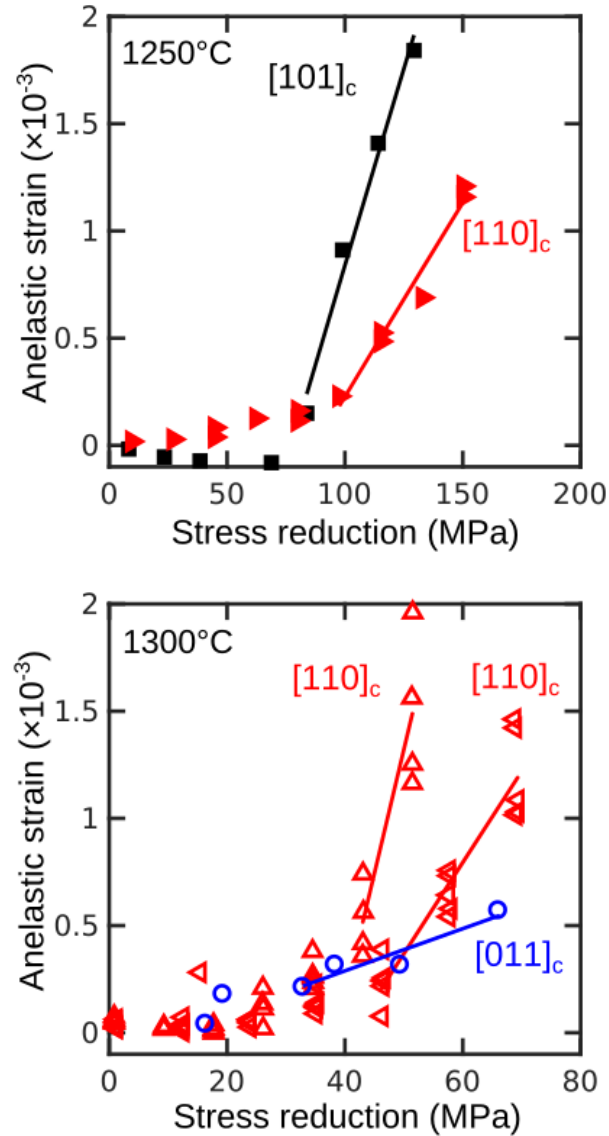


Figure 5: Anelastic reverse strain as a function of the magnitude of the stress reduction at two different temperatures. Marker shapes and colors are as designated in Table 1. Solid lines are fits to equation 5. Only data with clearly measurable anelastic reverse strain are included in fits.

439 These results indicate that, in olivine deformed in the $[110]_c$ and $[101]_c$ orientations, the magnitudes of
 440 anelastic reverse strain are consistent with backstresses between 48 and 61% of the initially applied stress,
 441 which are largely consistent with most previously reported values in other materials (see section 4.1

above). Interestingly, our results for olivine deformed in the $[011]_c$ orientation suggest that the backstress is nearly equal to the applied stress. This difference in backstress between different crystal orientations, and therefore between different activated slip systems, is likely a result of different rates of dislocation recovery, as will be explored later in this section.

The values of the best-fit rate constant, γ , also differ between crystal orientations. Olivine deformed in the $[110]_c$ and $[101]_c$ orientations is characterized by values of γ between 5 and 28. However, olivine deformed in the $[011]_c$ orientation is characterized by a value of γ of 55, suggesting some anisotropy in the hardening behavior. The former values are considerably smaller than the value of $\gamma = 75 \pm 20$ determined in low-temperature experiments (Hansen et al. 2019). However, the value of $\gamma = 55$ for the crystal deformed in the $[011]_c$ orientation is within error of that determined in previous experiments. Those previous experiments examined both polycrystalline aggregates and single crystals, but only single crystals in the $[011]_c$ and $[111]_c$ orientations were investigated, and both orientations should have a large part of the dislocation content on the (010)[001] slip system. Therefore, at least where comparisons can be made, this analysis of the anelastic reverse strain is consistent with the same mechanism of deformation being responsible for transient deformation at low and high temperatures, as suggested by the microstructural interpretations of Wallis et al. (2020). Furthermore, these results provide initial evidence that transient deformation on the strongest slip system may dominate the transient response of a polycrystalline sample, which is inconsistent with some recent theoretical treatments of transient creep in olivine suggesting the transient is controlled by the weakest slip system (Masuti et al. 2019).

4.2.2 Strain rates immediately after a stress change

If the physics of backstress associated with dislocation interactions (equation 2) can explain the magnitude of the anelastic reverse strain, then following from previous treatments for metals (e.g., Biberger and Gibeling 1995), we hypothesize the physics of dislocation glide can explain the strain rates

465 observed just after a change in stress. This hypothesis is similar to the suggestion by Cooper et al. (2016)
 466 that load relaxation experiments on olivine single crystals are rate limited by the glide of dislocations. We
 467 plot strain rates just after our stress reductions in Figure 6a as a function of the effective stress, σ_e , which
 468 was determined by subtracting the value of σ_b determined in Figure 5 from the reduced stress. In other
 469 words, we suggest these stresses and strain rates in Figure 6a reflect the rheological behavior of olivine
 470 when no backstress is present. Only strain rates for small stress reductions without any measurable
 471 anelastic reverse strain are presented. Strain rates from the crystal deformed in the $[011]_c$ direction are not
 472 plotted because all tests for that sample exhibit some anelasticity. The observed strain rates are
 473 consistently faster than the strain rates predicted for steady-state creep, and interestingly, the apparent
 474 power-law exponent is ~ 1 . This observation may initially seem at odds with the hypothesis that creep at
 475 “constant microstructure” is controlled by the glide of dislocations. However, as recently discussed in the
 476 context of olivine deformation by Cooper et al. (2016), the apparent stress exponent of an exponential
 477 flow law decreases with decreasing stress, and therefore, these data could indeed be explained by an
 478 exponential function. In fact, the sinh function employed by Hansen et al. (2019) and described in
 479 equation 1 results in a linear relationship between strain rate and effective stress at low stresses.

480 We also compare these strain rates to data from previously published experiments at high temperatures in
 481 Figure 6b. To do so, we need data from experiments in which the backstress, and therefore the effective
 482 stress, is known. We suggest that the strain rates just after initial loading of constant-stress experiments
 483 reasonably represent a case in which the backstress is zero since very little strain has accumulated to
 484 produce a backstress. Data of this type are available for single crystals of San Carlos olivine (Hanson and
 485 Spetzler 1994; Cooper, Stone, and Plookphol 2016) and polycrystalline aggregates of San Carlos olivine
 486 (Chopra 1997). We also include the yield stresses from single crystals deformed at room temperature at
 487 constant strain rate (Hansen et al. 2019). A complication in comparing data from the first increments of
 488 strain in constant-stress (or constant-strain-rate) tests is that the initial strain rate (or stress) is likely highly

489 sensitive to the density of mobile dislocations present in the starting material. Therefore, we normalize the
 490 strain rates in Figure 6b by the dislocation density at the beginning of the measurement. For our data, the
 491 initial dislocation density is calculated from the initial stress before the reduction using the
 492 dislocation-density piezometer of Bai and Kohlstedt (1992b). The initial values of dislocation density in
 493 the previous studies noted above are not available, but we estimate the starting density of single-crystal
 494 samples as 10^{10} m^{-2} based on observations in untreated San Carlos olivine (Toriumi and Karato 1978) and
 495 of polycrystalline samples as 10^{12} m^{-2} based on observations of hot-pressed Anita Bay dunite (S.-I.
 496 Karato, Paterson, and FitzGerald 1986).

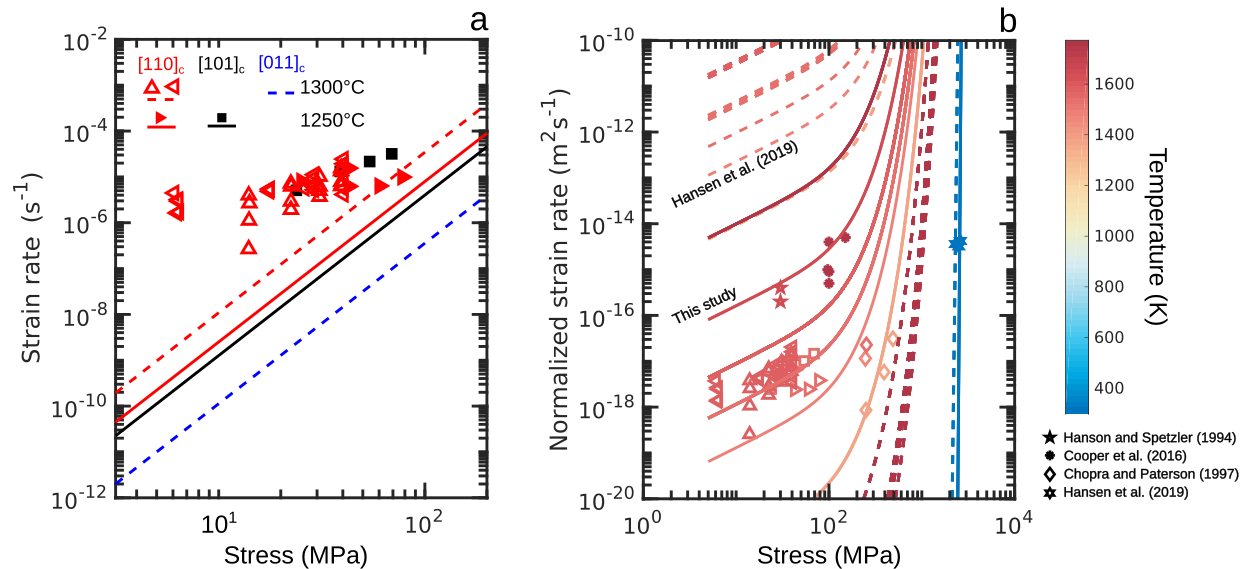


Figure 6: Analysis of strain rates measured immediately after stress reductions. (a) Strain rates measured after small reductions for which no anelastic reverse strain was detected. Stresses are effective stresses, that is, the applied stress minus the measured backstress given in Table 1. Marker shapes and colors are as designated in Table 1. Dashed and solid lines are flow laws for single crystals of olivine from Bai et al. (1991). (b) Strain rates as in (a) compared to previously published data. Strain rates from previous high-temperature creep tests (Hanson and Spetzler 1994; Cooper, Stone, and Ploophol 2016; Chopra 1997) were taken from the beginning of experiments in which the backstress is expected to be negligible. Data from Hansen et al. (2019) are yield stresses measured in constant-strain-rate experiments. Dashed lines represent the flow law from Hansen et al. (2019). Solid lines represent the flow law calibrated in this study.

497 The comparison in Figure 6b provides two key observations. First, the apparent stress exponent differs for
 498 each individual data set, systematically increasing with increasing stress. This observation is consistent
 499 with deformation just after a stress reduction being controlled by the glide of dislocations following an
 500 exponential or sinh flow law. Second, we also plot the flow law from Hansen et al. (2019) (equation 1 and
 501 dashed lines in Figure 6b). This flow law does well at predicting the strain rates in the room-temperature
 502 experiments on which the flow law is based, but does poorly at predicting the strain rates at high
 503 temperatures. The flow law overpredicts strain rates from this study by approximately six orders of
 504 magnitude. This discrepancy at high temperatures is not that surprising since the temperature-dependence
 505 is the most poorly constrained component of the flow law. Not only was the flow law primarily calibrated
 506 with data collected at room temperature, but data at higher temperatures were either subject to low
 507 resolution in temperature measurement or were taken from other studies conducted with different
 508 methods. Furthermore, the apparent temperature dependence of equation 1 depends strongly on the
 509 assumption that the activation energy for glide is linearly dependent on stress. A nonlinear dependence, or
 510 even a transition in the dependence as a function of temperature is possible (chapter 4 in Caillard and
 511 Martin 2003). Therefore, we suggest that (1) our data are *consistent* with transient deformation being
 512 controlled by the process of dislocation glide but that (2) existing flow laws for glide-controlled plasticity
 513 in olivine need to be recalibrated before extension to such high temperatures.

514 4.2.3 Recalibration of the flow law for dislocation glide

515 Here we recalibrate the flow law for dislocation glide in olivine using the data presented in Figure 6b.
 516 However, we must modify equation 1 to reasonably apply this flow law to the range of data in Figure 6b,
 517 which covers a temperature range from 20°C to 1300°C. Although the form of the flow law calibrated by
 518 Hansen et al. (2019) was derived from the Orowan equation (see their equation 6), the density of mobile
 519 dislocations that actually produce strain was assumed to be relatively constant and therefore was

520 incorporated into A_1 . This assumption is generally justified for applications to low-temperature,
 521 high-stress deformation because the stress dependence of the sinh term far outweighs the stress
 522 dependence of the mobile dislocation density. However, this is not necessarily the case at high
 523 temperatures or low stresses, and therefore we suggest a more robust form of the flow law as

$$\dot{\epsilon}_p = A_2 \rho \exp\left(-\frac{\Delta F}{RT}\right) \sinh\left(\frac{\Delta F}{RT} \frac{\sigma - \sigma_b}{\Sigma}\right), \quad (6)$$

524 where A_2 is a constant and ρ is the density of mobile dislocations. As with equation 1, this flow law
 525 inherently assumes a linear dependence of ΔG on the effective stress (see section 4.2.1). A nonlinear
 526 relationship may better characterize the fundamental physics of dislocation glide in olivine, but as
 527 demonstrated by Jain et al. (2017)

528 Figure 6b compares our compiled data for strain rates controlled by dislocation glide to a best fit to
 529 equation 6. In the applied fitting procedure, we fixed $\Sigma = 3.1$ GPa, which is well constrained by the
 530 previous calibration using low-temperature data, and conducted a grid search to find the optimum values
 531 of A_2 and ΔF . The best-fit values of these parameters are $A_2 = 10^{11.1 \pm 0.2} \text{ m}^2 \text{ s}^{-1}$ and $\Delta F = 827 \pm 20$ kJ/mol. This
 532 value for ΔF is much higher than most previous estimates for dislocation glide flow laws in olivine,
 533 which are generally between 300 and 600 kJ/mol (see the compilation by Demouchy et al. 2013). Our
 534 choice of a linear relationship between ΔG and effective stress likely biases the best-fit activation energies
 535 to high values. One outlier of previous work is an estimate of 1480 kJ/mol from Long et al. (2011)

536 Although our calibrated flow law uses a simplified form, at least empirically it is able to capture the
 537 observed behavior over a wide range in temperatures, reasonably predicting both the magnitudes of the
 538 strain rates as well as the change in the apparent stress exponent as a function of temperature.
 539 Furthermore, combining equation 6 with a parameterization of the backstress evolution, such as that in
 540 equation 2, allows the complete transient to be predicted. Different values of γ can be used to predict the

541 transient behavior of different slip systems, as discussed in relation to Figure 5. However, we note that
 542 the dislocation recovery mechanisms implicit in equation 2 are only calibrated for use at low
 543 temperatures. Therefore, prediction of the steady-state strain rates at elevated temperatures requires
 544 high-temperature recovery mechanisms (e.g., dislocation climb) to be accounted for. Incorporating these
 545 recovery mechanisms into a parameterization of dislocation creep in olivine will be a key topic addressed
 546 in future work.

547 We can, however, assess the magnitude of the backstresses at high temperature. To restate the discussion
 548 above, we suggest that equation 6 controls the macroscopic strain rate in both transient and steady-state
 549 creep, and the evolution between the two regimes reflects the evolution of the net backstress. In this
 550 framework, at steady state, the steady-state strain rate would be equal to equation 6, which allows us to
 551 solve for the backstress at steady state,

$$\sigma_b = \sigma - \text{asinh} \left[\frac{\dot{\epsilon}_{ss}}{A_2 \rho} \exp \left(\frac{F}{RT} \right) \right] \frac{RT\Sigma}{F}, \quad (7)$$

552 where $\dot{\epsilon}_{ss}$ is the steady-state strain rate. In Figure 7, we present values of the steady-state back stress
 553 estimated with equation 7. These values were calculated using steady-state strain rates predicted by the
 554 flow laws for single-crystal olivine from Bai and Kohlstedt (1992a) and dislocation densities predicted by
 555 the dislocation-density piezometer from Bai and Kohlstedt (1992b). These calculated backstresses
 556 compare favorably to those observed in stress-reduction tests, which provides further confidence in our
 557 calibration of equation 6 and in our suggestion that equation 6 can be used to predict the steady-state
 558 strain rate if the backstress is known.

559 Notably, the calculated curves in Figure 7 capture the orientation dependence of the backstress. Equation
 560 7 correctly predicts that crystals compressed along $[011]_c$ have a backstress nearly equal to the applied
 561 stress, and that crystals deformed in other orientations have a backstress that is approximately 50% of the

562 applied stress. Because we use the same values of A_2 , F , and Σ regardless of orientation, the orientation
 563 dependence of the calculated backstress results entirely from the different steady-state flow laws.
 564 Therefore, there is a clear link between the anisotropy of the steady-state creep rate and the anisotropy of
 565 the backstress evolution, which we suggest is primarily a reflection of anisotropic rates of dislocation
 566 recovery (e.g., L. Wang et al. 2016). We reiterate and emphasize that, although we propose dislocation
 567 creep during the transient and steady state is governed by equation 6, the steady-state strain rate can still
 568 be set by the kinetics of dislocation climb that counteract the buildup of additional backstress.

569

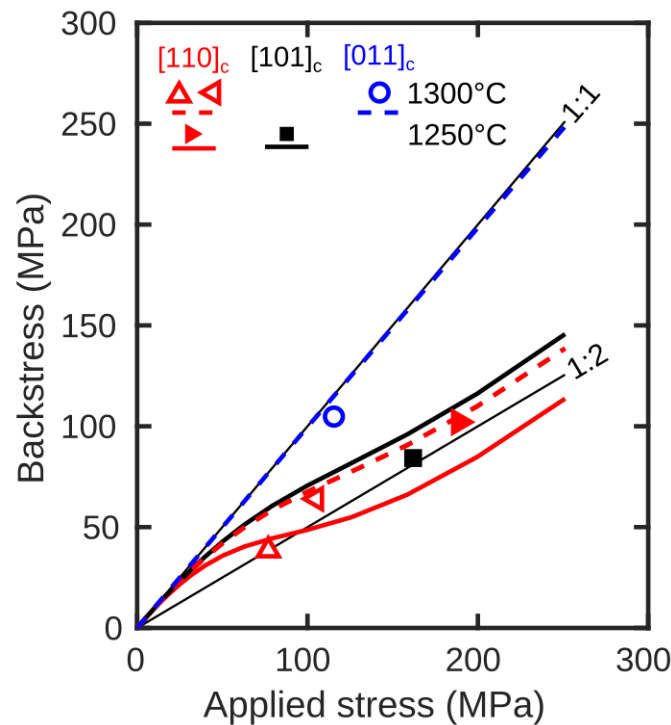


Figure 7: Measured and predicted backstresses as a function of the applied stress. Markers indicate backstresses determined using the analysis presented in Figure 5. The two thin black lines are guides for values of backstress equal to the applied stress (1:1) or half of the applied stress (1:2). The curves are backstresses predicted from equation 7.

570 4.2.4 Implications for the type of transient creep and load relaxation

571 Our hypothesis linking low-temperature plasticity, transient creep, steady-state creep, and the evolution of
572 backstress predicts several other phenomena. A key phenomenon is variation in the sign of the strain
573 acceleration during the transient. That is, whether or not the strain rate decreases (a “normal” transient) or
574 increases (an “inverse” transient) during the evolution to steady state. Hanson and Spetzler (1994)
575 observed that the type of transient depends on the crystal orientation, and Cooper et al. (2016) observed a
576 transition in the type of transient in the same experiment as the applied stress was increased in steps.
577 Figure 8a schematically describes how the model presented here can predict both types of transients. We
578 suggest that the strain rates for olivine will generally be described by the flow law for dislocation glide
579 (equation 6). After initial application of stress to the sample, the mobile dislocation density and backstress
580 evolve toward their steady-state values. This change in the microstructural parameters shifts the flow law
581 to higher or lower strain rates until a steady state is reached. Thus, we hypothesize that the apparent
582 steady-state power law is fundamentally defined by the relationship between the steady-state backstress
583 and the applied stress. At the onset of deformation, if the dislocation-glide flow law predicts faster strain
584 rates than the steady-state strain rate, then the initial strain rate will be fast, and the flow law will
585 gradually shift to slower rates (a normal transient) due to an increasing backstress. Conversely, if the
586 initial strain rate dictated by the flow law is slower than at steady-state, the flow law will shift to faster
587 rates (an inverse transient) due to an increase in the density of mobile dislocations. Therefore, we suggest
588 that the intersection of the dislocation-glide flow law and the steady-state power law can be used to
589 predict whether normal or inverse transients will occur. Figure 8b presents the type of transient predicted
590 by this method over a range of temperatures and pressures for all three orientations of olivine. The
591 predictions are in excellent agreement with observations in this study (Table 1), as well as with the
592 observations made by Hanson and Spetzler (1994) and Cooper et al. (2016) .

593 We additionally note that, while this analysis correctly predicts a normal transient for the initial loading of
594 olivine along [011], it also predicts that once steady state is reached, any subsequent stress increase
595 would lead to an inverse transient. Therefore, this model also correctly predicts the change in the type of
596 transient from normal to inverse observed by Cooper et al. (2016) in progressive load-stepping
597 experiments (see their Figure 4a).

598 We finally point out that our model is similar to that developed by Stone (1991) (in turn, fundamentally
599 inspired by the phenomenological theory of Hart (1970)) and applied to geological materials such as
600 halite (Stone, Plookphol, and Cooper 2004) and olivine (Cooper, Stone, and Plookphol 2016). Stone
601 (1991) describes a “master curve” or equation of state relating stress, strain rate, and the hardness, which
602 is a state variable related to the aspects of the microstructure that control the material response. The
603 “master curve” described by that model is analogous to the flow law presented here. The key difference
604 lies in the choice of microstructural variable that controls the hardness and translates the master curve in
605 stress–strain-rate space. Stone advocates that the self-similar nature of the microstructure in creeping
606 materials, which involves many different length scales, allows the essential aspects of that distribution to
607 be described by a single state variable. Stone uses the subgrain size as that variable. Instead, here we use
608 the net backstress as a representation of the full dislocation microstructure, which includes the stress
609 fields associated with subgrain boundaries. This distinction is important because the backstress is a more
610 general description of the microstructure and therefore can be applied to our samples, which generally do
611 not include appreciable subgrain boundaries.

612 An additional difference between the model of Stone (1991) and that presented here lies in the nature of
613 the relationship between the microstructural variable and the master curve. The master curve of Stone
614 (1991) represents, at high stresses, a plastic yield stress controlled by the subgrain size, and at low
615 stresses, a flow law describing Nabarro-Herring creep among subgrains. The transition of Stone’s master

616 curve to Nabarro-Herring creep at low stresses was inspired by strain rates during load relaxation that
 617 approached a linear relationship with stress. The flow law presented here results in a similar
 618 phenomenological response, which includes, at high stresses, an exponential dependence of strain rate on
 619 effective stress, and at low stresses, a linear dependence of strain rate on effective stress. In our model,
 620 this full range of behavior results from a single flow law due to the nature of the sinh term in equation 6.
 621 The model of Stone (1991) explains load-relaxation data for halite exceptionally well (Stone, Ploorkphol,
 622 and Cooper 2004), collapsing the data set to a master curve when scaled by the subgrain size. In contrast,
 623 in load relaxation experiments on single crystals of olivine in the $[011]_c$ orientation, Cooper et al. (2016)
 624 found that the master curve is not shifted by the subgrain size. In fact, they demonstrated that the
 625 load-relaxation response is unchanged by the initially applied stress and any associated changes in
 626 microstructure, always falling on the same stress-strain-rate curve. Therefore, Cooper et al. (2016)
 627 suggested that a different physical process controls the relaxation response of olivine, dislocation glide
 628 with kinetics described by an exponential flow law. Our data and analysis provides an alternative, but
 629 related, explanation that can be applied to both halite and olivine in load-relaxation experiments.

630 As described by equation 6, the backstress shifts the dislocation-glide flow law in stress-strain-rate space.
 631 Figure 7 illustrates that the backstress for crystals deformed in the $[011]_c$ orientation is essentially equal to
 632 the applied stress. This observation suggests that, during the initial evolution to steady state, the
 633 backstress reaches a value equal to the applied stress before the flow law can shift to lower rates to reach
 634 the steady-state rate. Therefore, during a load-relaxation experiment, the effective stress is initially zero,
 635 and forward strain only proceeds due to the reduction of the backstress associated with dislocation
 636 recovery. The need for recovery implies that load-relaxation in this context is not a constant
 637 microstructure test, and indeed Cooper et al. (2016) observed a reduction in dislocation density during
 638 relaxation. Continued forward strain and reduction of backstress would occur in lock step and simply
 639 follow the steady-state power law regardless of the initial stress, again consistent with the observation of

640 Cooper et al. (2016). In contrast, for halite (Gupta and Li 1970) and other orientations of olivine (Figure
 641 7), the backstresses are a smaller portion of the applied stress, and therefore the effective stress is both
 642 finite and a function of the applied stress. In this context, a stress relaxation experiment would follow the
 643 low-temperature plasticity flow law prior to any relaxation of the backstress. Therefore, we predict that
 644 stress relaxation experiments conducted on other orientations of olivine would exhibit a similar response
 645 as halite, following the flow law for dislocation glide and dependent on the magnitude of the initial stress.

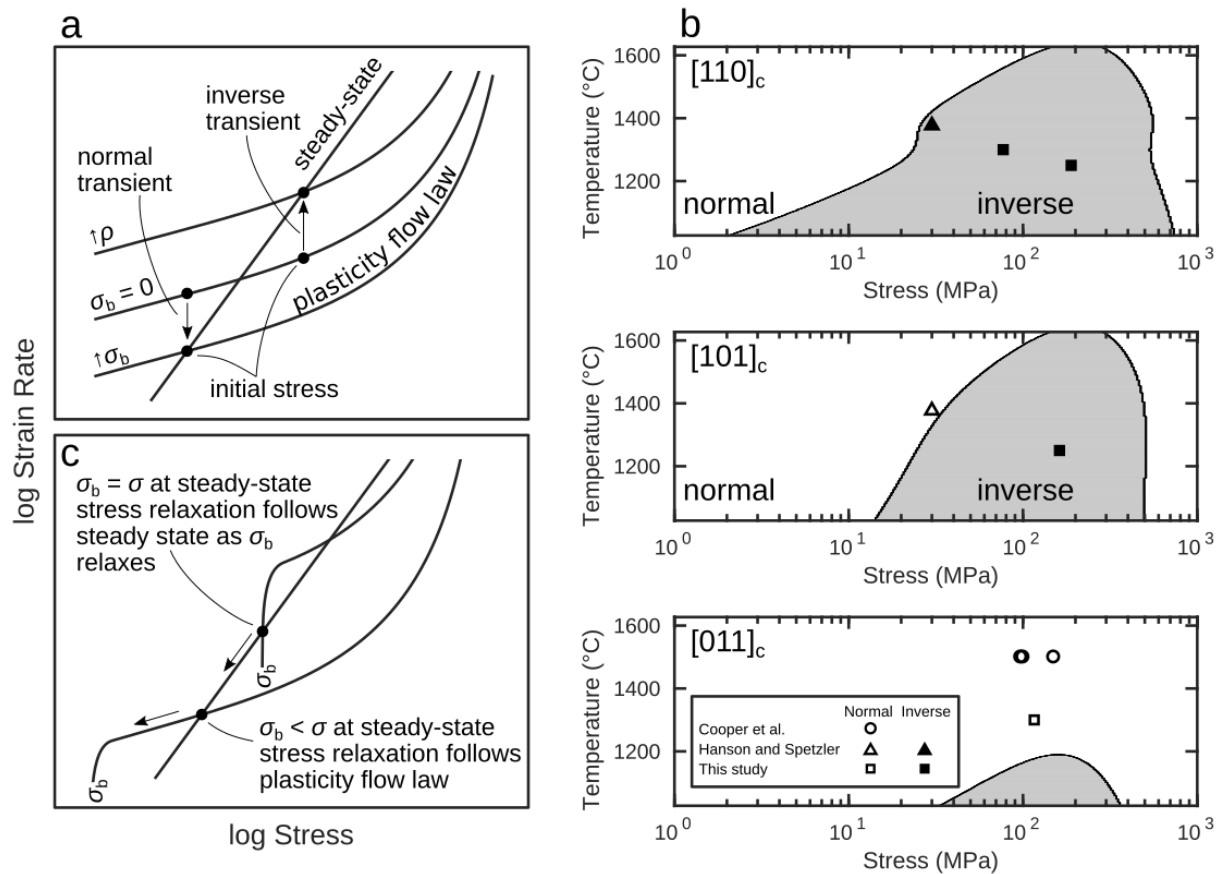


Figure 8: Evaluation of the proposed model in the contexts of the shape of the transient and of stress relaxation. (a) Schematic description of our model. The initial stress applied to a sample leads to a strain rate controlled by flow law for dislocation glide (equation 6) without any backstress. If this strain rate is higher than the steady-state strain rate, we predict that the transient will be normal and controlled by the increase in backstress. If this strain rate is lower than the steady-state strain rate, we predict that the transient will be inverse and controlled by the increase in density of mobile dislocations. In this schematic, the backstress is assumed to always be less than the minimum stress on the plot. (b) Conditions leading to inverse or normal transients calculated with the concept presented in (a). Data for comparison come from the initial strain rates at the start of experiments

presented by Cooper et al. (2016), Hanson and Spetzler (1994), and this study. (c) Schematic prediction of different responses during stress relaxation experiments. If steady state is reached before the backstress equals the applied stress, then subsequent relaxation will follow the flow law. If the backstress increases to equal the applied stress at the steady state, then subsequent load relaxation only results in forward strain as the backstress relaxes, following the steady-state power law.

646 ***4.3 Consequences for transient creep in the upper mantle***

647 Rheological models employed to analyze geodynamic scenarios vary widely in their form and behavior.
 648 In terms of models that capture transient behavior, by far the most common is the Burgers model, which
 649 contains two viscous elements, one describing the initial viscosity at the onset of loading and the other
 650 describing the steady-state viscosity. Burgers models (and related variations including multiple anelastic
 651 elements) have been calibrated for deformation of olivine dominated by diffusive mass transfer (Sundberg
 652 and Cooper 2010; e.g., Faul and Jackson 2005) and by dislocation motion (Chopra 1997). The former
 653 result in a linear viscoelastic model for which the viscosities do not depend on the level of stress applied,
 654 whereas the latter result in nonlinear viscoelastic models for which the viscosities are stress dependent.
 655 Linear viscoelastic models based on a Burgers model have been widely applied to a range of phenomena
 656 including glacial isostatic adjustment (Yuen et al. 1986; e.g., Rumpker and Wolff 1996), postseismic creep
 657 (e.g., Hetland and Hager 2005; Pollitz, Bürgmann, and Banerjee 2006), tidal dissipation (e.g., Nimmo,
 658 Faul, and Garnero 2012; Nimmo and Faul 2013; Bierson and Nimmo 2016), and seismic attenuation
 659 (Carcione et al. 2014; S.-I. Karato 2012). Linear Burgers models have likely seen application to such a
 660 wide range of processes because of the relative mathematical ease of their implementation and the
 661 recognition that nonlinear viscoelasticity is generally restricted to large stress and strain amplitudes,
 662 whereas most of these processes involve small stresses. However, nonlinear viscoelastic models have seen
 663 increasing application to scenarios with larger stress changes, especially postseismic creep (Freed, Hirth,
 664 and Behn 2012; Masuti et al. 2016). The viscosities in these nonlinear models are assumed to follow
 665 power laws, motivated by the power-law behavior of olivine during steady-state dislocation creep.

666 However, the stress dependence is assumed to be the same for both the initial and the steady-state
667 viscosity. Therefore, in both linear Burgers models and nonlinear variations for upper mantle rocks, the
668 ratio of the initial and steady-state viscosities is generally taken to be a constant.

669 Our results and analysis presented above demonstrate that the rheological behavior of olivine can be quite
670 different than the behavior implemented in previous modelling efforts. As demonstrated by our data in
671 Figure 6 and illustrated schematically in Figure 8a, during initial loading, we predict the strain rate of
672 olivine to be close to linearly dependent on the stress at low stresses, with increasingly higher sensitivities
673 to stress with increasing stress (Figure 6b also illustrates that the transition between linear and nonlinear
674 behavior is not particularly dependent on temperature). Therefore, we predict the nonlinearity of the
675 initial viscosity to be a function of the stress. This conclusion is particularly noteworthy because, for
676 example, the apparent viscosity associated with geodynamic phenomena on short timescales could
677 potentially follow linear viscoelastic behavior even when dislocation glide is the main strain producing
678 process. Furthermore, if the nonlinearity of the initial viscosity is dependent on stress and temperature,
679 then the ratio of the initial viscosity to the steady-state viscosity will also be a function of the stress.

680 We note that the transient deformation behaviour of rocks may be subject to additional effects not
681 captured in our experiments on single crystals, including grain-boundary strengthening (Hansen et al.
682 2019; Kumamoto et al. 2017), recovery from grain-boundary migration (e.g., Toriumi 1982), stress
683 transfer between grains (Duval, Ashby, and Anderman 1983; Masuti et al. 2019), additional deformation
684 mechanisms, and the effects of other minerals and/or fluids. Nonetheless, the dislocation-mediated
685 processes that we analyse here should be ubiquitous across relevant geological conditions regardless of
686 the occurrence of other effects. Therefore, the model that we present provides a fundamental starting point
687 for analysing transient creep that allows isolation of the contribution from intragranular processes, with

the benefit that departures from such behaviour in polycrystalline or polyminerallic materials may constrain contributions from additional processes.

5 Conclusions

Here we present a series of deformation experiments to investigate the microphysical mechanisms controlling transient creep in olivine. We conducted stress-reduction tests on olivine single crystals at elevated temperatures in a variety of crystal orientations. Our results reveal that stress-reductions are often accompanied by anelastic reverse strain. The magnitude of the reverse strain is dependent on the magnitude of the stress reduction. In certain crystal orientations, small stress reductions result in no reverse strain, indicating that a backstress is present in the crystals that is a fraction of the initially applied stress.

We recognise that deformation of olivine is fundamentally controlled by the glide velocity of dislocations, and therefore hypothesize that a flow law based on the kinetics of dislocation glide can be applied across a wide range of temperatures. The buildup of backstresses associated with strain hardening in low-temperature plasticity is quantitatively consistent with the magnitudes of observed reverse strains. This hypothesis also predicts that the strain rates just after a stress reduction can be predicted with a flow law calibrated at low temperatures, however, the existing flow law considerably overpredicts the strain rates observed here. Therefore, we recalibrate the activation energy of the flow law to be consistent with initial strain rates in our experiments, initial strain rates in previously published high-temperature experiments, and yield stresses in low-temperature experiments.

We suggest that this recalibrated flow law can be used to predict the transient and steady-state creep strain rates by taking into account the evolution of the backstress and dislocation density. By combining the recalibrated flow law with previously published flow laws for the steady-state creep of olivine single

crystals, we are able to correctly predict the magnitude of the observed backstress and its orientation dependence. We are also able to predict whether strain rates decrease (normal) or increase (inverse) during transient creep in our experiments and previously published experiments. This model also provides an explanation for the results of previous stress-relaxation experiments.

The results and analysis presented here indicate that the rheological behaviour of olivine during transient creep is fundamentally different from that incorporated into analyses of geodynamic processes. We suggest that the initial viscosity after a stress change, if controlled by dislocations, can range from linear to nonlinear, depending on the stress magnitude. Therefore, the ratio between the transient and steady-state viscosities is likely a function of the stress.

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References

- Abramson, E. H., J. M. Brown, L. J. Slutsky, and J. Zaug. 1997. "The Elastic Constants of San Carlos Olivine to 17 GPa." *Journal of Geophysical Research, [Solid Earth]* 102 (B6): 12253–63.
- Ahlquist, C. N., and W. D. Nix. 1971. "The Measurement of Internal Stresses during Creep of Al and Al-Mg Alloys." *Acta Metallurgica* 19 (4): 373–85.
- Argon, A. S., and S. Takeuchi. 1981. "Internal Stresses in Power-Law Creep." *Acta Metallurgica*. [https://doi.org/10.1016/0001-6160\(81\)90113-9](https://doi.org/10.1016/0001-6160(81)90113-9).
- Bai, Quan, and D. L. Kohlstedt. 1992a. "High-Temperature Creep of Olivine Single Crystals III. Mechanical Results for Unbuffered Samples and Creep Mechanisms." *Philosophical Magazine A* 66 (6): 1149–81.
- . 1992b. "High-Temperature Creep of Olivine Single Crystals, 2. Dislocation Structures." *Tectonophysics* 206 (1): 1–29.
- Bai, Quan, S. J. Mackwell, and D. L. Kohlstedt. 1991. "High-Temperature Creep of Olivine Single Crystals 1. Mechanical Results for Buffered Samples." *Journal of Geophysical Research, [Solid Earth]* 96 (B2): 2441–63.
- Bammann, D. J. 1989. "A Micro-Mechanically Motivated Model of Finite Deformation Plasticity." In *Advances in Plasticity 1989*, edited by Akhtar S. Khan and Masataka Tokuda, 283–86. Oxford: Pergamon.
- Bayley, C. J., W. A. M. Brekelmans, and M. G. D. Geers. 2006. "A Comparison of Dislocation Induced Back Stress Formulations in Strain Gradient Crystal Plasticity." *International Journal of Solids and Structures* 43 (24): 7268–86.
- Benjamin Britton, T., and Angus J. Wilkinson. 2012. "Stress Fields and Geometrically Necessary Dislocation Density Distributions near the Head of a Blocked Slip Band." *Acta Materialia* 60 (16): 5773–82.
- Biberger, M., and J. C. Gibeling. 1995. "Analysis of Creep Transients in Pure Metals Following Stress Changes." *Acta Metallurgica et Materialia*. [https://doi.org/10.1016/0956-7151\(95\)00052-w](https://doi.org/10.1016/0956-7151(95)00052-w).
- Bierson, C. J., and F. Nimmo. 2016. "A Test for Io's Magma Ocean: Modeling Tidal Dissipation with a Partially Molten Mantle." *Journal of Geophysical Research: Planets*. <https://doi.org/10.1002/2016je005005>.
- Blum, W., and A. Finkel. 1982. "New Technique for Evaluating Long Range Internal Back Stresses." *Acta Metallurgica* 30 (8): 1705–15.
- Blum, W., and E. Weckert. 1987. "On the Interpretation of the 'Internal Stress' Determined from Dip Tests during Creep of Al-5at.%Mg." *Materials Science and Engineering*. [https://doi.org/10.1016/0025-5416\(87\)90449-6](https://doi.org/10.1016/0025-5416(87)90449-6).
- Britton, T. B., J. Jiang, Y. Guo, A. Vilalta-Clemente, D. Wallis, L. N. Hansen, A. Winkelmann, and A. J. Wilkinson. 2016. "Tutorial: Crystal Orientations and EBSD — Or Which Way Is Up?" *Materials Characterization* 117: 113–26.
- Britton, T. B., and A. J. Wilkinson. 2011. "Measurement of Residual Elastic Strain and Lattice Rotations with High Resolution Electron Backscatter Diffraction." *Ultramicroscopy* 111 (8): 1395–1404.
- . 2012. "High Resolution Electron Backscatter Diffraction Measurements of Elastic Strain Variations in the Presence of Larger Lattice Rotations." *Ultramicroscopy* 114: 82–95.
- Čadek, Josef. 1987. "The Back Stress Concept in Power Law Creep of Metals: A Review." *International Journal of Green Nanotechnology: Materials Science and Engineering* 94 (October): 79–92.
- Caillard, D., and J. L. Martin. 2003. *Thermally Activated Mechanisms in Crystal Plasticity*. Elsevier.
- Cao, Ri, Lars N. Hansen, Christopher A. Thom, and David Wallis. 2020. "An Apparatus for Measuring

- Nonlinear Viscoelasticity of Minerals at High Temperature.” <https://doi.org/10.31223/X59W2X>.
- Carcione, J. M., F. Poletto, B. Farina, and A. Craglietto. 2014. “Simulation of Seismic Waves at the Earth’s Crust (brittle-Ductile Transition) Based on the Burgers Model.” *Solid Earth* 5 (2): 1001.
- Caswell, Tess E., Reid F. Cooper, and David L. Goldsby. 2015. “The Constant-Hardness Creep Compliance of Polycrystalline Ice.” *Geophysical Research Letters* 42 (15): 6261–68.
- Chopra, Prame N. 1997. “High-Temperature Transient Creep in Olivine Rocks.” *Tectonophysics*. [https://doi.org/10.1016/s0040-1951\(97\)00134-0](https://doi.org/10.1016/s0040-1951(97)00134-0).
- Cooper, Reid F., Donald S. Stone, and Thawatchai Plookphol. 2016. “Load Relaxation of Olivine Single Crystals.” *Journal of Geophysical Research: Solid Earth*. <https://doi.org/10.1002/2016jb013425>.
- Demouchy, Sylvie, Andréa Tommasi, Tiziana Boffa Ballaran, and Patrick Cordier. 2013. “Low Strength of Earth’s Uppermost Mantle Inferred from Tri-Axial Deformation Experiments on Dry Olivine Crystals.” *Physics of the Earth and Planetary Interiors* 220 (July): 37–49.
- Durham, W. B., and C. Goetze. 1977. “Plastic Flow of Oriented Single Crystals of Olivine: 1. Mechanical Data.” *Journal of Geophysical Research* 82 (36): 5737–53.
- Durham, W. B., C. Goetze, and B. Blake. 1977. “Plastic Flow of Oriented Single Crystals of Olivine: 2. Observations and Interpretations of the Dislocation Structures.” *Journal of Geophysical Research*. <https://doi.org/10.1029/jb082i036p05755>.
- Duval, P., M. F. Ashby, and I. Anderman. 1983. “Rate-Controlling Processes in the Creep of Polycrystalline Ice.” *The Journal of Physical Chemistry* 87 (21): 4066–74.
- Evans, R. W., W. J. F. Roach, and B. Wilshire. 1985. “Creep of Aluminium Following Stress Reductions.” *Scripta Metallurgica* 19 (8): 999–1003.
- Faul, Ulrich H., and Ian Jackson. 2005. “The Seismological Signature of Temperature and Grain Size Variations in the Upper Mantle.” *Earth and Planetary Science Letters* 234 (1): 119–34.
- Field, Edward H., Glenn P. Biasi, Peter Bird, Timothy E. Dawson, Karen R. Felzer, David D. Jackson, Kaj M. Johnson, et al. 2015. “Long-Term Time-Dependent Probabilities for the Third Uniform California Earthquake Rupture Forecast (UCERF3).” *Bulletin of the Seismological Society of America* 105 (2A): 511–43.
- Field, Edward H., Thomas H. Jordan, Morgan T. Page, Kevin R. Milner, Bruce E. Shaw, Timothy E. Dawson, Glenn P. Biasi, et al. 2017. “A Synoptic View of the Third Uniform California Earthquake Rupture Forecast (UCERF3).” *Seismological Research Letters* 88 (5): 1259–67.
- Freed, Andrew M. 2005. “Earthquake Triggering by Static, Dynamic, and Postseismic Stress Transfer.” *Annual Review of Earth and Planetary Sciences* 33 (1): 335–67.
- Freed, Andrew M., Greg Hirth, and Mark D. Behn. 2012. “Using Short-Term Postseismic Displacements to Infer the Ambient Deformation Conditions of the Upper Mantle.” *Journal of Geophysical Research* 117 (B1). <https://doi.org/10.1029/2011JB008562>.
- Frost, H. J., and M. F. Ashby. 1982. *Deformation Mechanism Maps: The Plasticity and Creep of Metals and Ceramics*. Oxford, UK: Pergamon Press.
- Garmestani, H., M. R. Vaghar, and E. W. Hart. 2001. “A Unified Model for Inelastic Deformation of Polycrystalline Materials—application to Transient Behavior in Cyclic Loading and Relaxation.” *International Journal of Plasticity* 17 (10): 1367–91.
- Garofalo, and F. 1963. “An Empirical Relation Defining the Stress Dependence of Minimum Creep Rate in Metals.” *Trans. AIME* 227: 351–56.
- Gibbs, G. B. 1966. “Creep and Stress Relaxation Studies with Polycrystalline Magnesium.” *The Philosophical Magazine: A Journal of Theoretical Experimental and Applied Physics* 13 (122): 317–29.
- Gibeling, J. C., and W. D. Nix. 1981. “Observations of Anelastic Backflow Following Stress Reductions during Creep of Pure Metals.” *Acta Metallurgica* 29 (10): 1769–84.
- Goetze, C., and D. L. Kohlstedt. 1973. “Laboratory Study of Dislocation Climb and Diffusion in Olivine.”

- Journal of Geophysical Research* 78 (26): 5961–71.
- Gibb, Tye T., and Reid F. Cooper. 1998. “Low-Frequency Shear Attenuation in Polycrystalline Olivine: Grain Boundary Diffusion and the Physical Significance of the Andrade Model for Viscoelastic Rheology.” *Journal of Geophysical Research: Solid Earth*. <https://doi.org/10.1029/98jb02786>.
- Gupta, I., and J. C. M. Li. 1970. “Stress Relaxation, Internal Stress and Work Hardening in LiF and NaCl Crystals.” *International Journal of Green Nanotechnology: Materials Science and Engineering* 6 (1): 20–26.
- Hansen, Lars N., Kathryn M. Kumamoto, Christopher A. Thom, David Wallis, William B. Durham, David L. Goldsby, Thomas Breithaupt, Cameron D. Meyers, and David L. Kohlstedt. 2019. “Low-temperature Plasticity in Olivine: Grain Size, Strain Hardening, and the Strength of the Lithosphere.” *Journal of Geophysical Research: Solid Earth*. <https://doi.org/10.1029/2018jb016736>.
- Hanson, David R., and Hartmut A. Spetzler. 1994. “Transient Creep in Natural and Synthetic, Iron-Bearing Olivine Single Crystals: Mechanical Results and Dislocation Microstructures.” *Tectonophysics*. [https://doi.org/10.1016/0040-1951\(94\)90191-0](https://doi.org/10.1016/0040-1951(94)90191-0).
- Hart, E. W. 1970. “A Phenomenological Theory for Plastic Deformation of Polycrystalline Metals.” *Acta Metallurgica* 18 (6): 599–610.
- . 1976. “Constitutive Relations for the Nonelastic Deformation of Metals.” *Journal of Engineering Materials and Technology* 98 (3): 193–202.
- Hetland, E. A., and B. H. Hager. 2005. “Postseismic and Interseismic Displacements near a Strike-Slip Fault: A Two-Dimensional Theory for General Linear Viscoelastic Rheologies.” *Journal of Geophysical Research, [Solid Earth]* 110 (B10). <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2005JB003689>.
- Hirth, Greg, and David Kohlstedt. 2003. “Rheology of the Upper Mantle and the Mantle Wedge: A View from the Experimentalists.” In *Geophysical Monograph Series*, edited by John Eiler, 138:83–105. Washington, D. C.: American Geophysical Union.
- Hunter, J., and A. B. Watts. 2016. “Gravity Anomalies, Flexure and Mantle Rheology Seaward of Circum-Pacific Trenches.” *Geophysical Journal International* 207 (1): 288–316.
- Jackson, Ian, John D. Fitz Gerald, Ulrich H. Faul, and Ben H. Tan. 2002. “Grain-Size-Sensitive Seismic Wave Attenuation in Polycrystalline Olivine.” *Journal of Geophysical Research: Solid Earth*. <https://doi.org/10.1029/2001jb001225>.
- Jackson, I., U. H. Faul, and R. Skelton. 2014. “Elastically Accommodated Grain-Boundary Sliding: New Insights from Experiment and Modeling.” *Physics of the Earth and Planetary Interiors* 228: 203–10.
- Jain, Chhavi, Jun Korenaga, and Shun-Ichiro Karato. 2017. “On the Yield Strength of Oceanic Lithosphere.” *Geophysical Research Letters* 44 (19): 9716–22.
- Karato, Shun-Ichiro. 2012. “On the Origin of the Asthenosphere.” *Earth and Planetary Science Letters* 321: 95–103.
- Karato, Shun-Ichiro, Mervyn S. Paterson, and John D. FitzGerald. 1986. “Rheology of Synthetic Olivine Aggregates: Influence of Grain Size and Water.” *Journal of Geophysical Research* 91 (B8): 8151–76.
- Karato, Shun-Ichiro, David C. Rubie, and Hong Yan. 1993. “Dislocation Recovery in Olivine under Deep Upper Mantle Conditions: Implications for Creep and Diffusion.” *Journal of Geophysical Research* 98 (B6): 9761–68.
- Karato, S., and H. A. Spetzler. 1990. “Defect Microdynamics in Minerals and Solid-State Mechanisms of Seismic Wave Attenuation and Velocity Dispersion in the Mantle.” *Reviews of Geophysics*. <https://doi.org/10.1029/rg028i004p00399>.
- Kassner, Michael E. 2015. *Fundamentals of Creep in Metals and Alloys*. Butterworth-Heinemann.
- Kocks, U. F., A. S. Argon, and M. F. Ashby. 1975. “Thermodynamics and Kinetics of Slip.” *Progress in Materials Science* 19.
- Korhonen, M. A., S-P Hannula, and Che-Yu Li. 1987. “State Variable Theories Based on Hart’s

- Formulation.” In *Unified Constitutive Equations for Creep and Plasticity*, 89–137. Springer.
- Krausz, A. S. 1968. “A Rate Theory of Dislocation Mobility.” *Acta Metallurgica* 16 (7): 897–902.
- Kumamoto, Kathryn M., Christopher A. Thom, David Wallis, Lars N. Hansen, David E. J. Armstrong, Jessica M. Warren, David L. Goldsby, and Angus J. Wilkinson. 2017. “Size Effects Resolve Discrepancies in 40 Years of Work on Low-Temperature Plasticity in Olivine.” *Science Advances* 3 (9): e1701338.
- Long, Hongbo, Donald J. Weidner, Li Li, Jiuhua Chen, and Liping Wang. 2011. “Deformation of Olivine at Subduction Zone Conditions Determined from in Situ Measurements with Synchrotron Radiation.” *Physics of the Earth and Planetary Interiors* 186 (1–2): 23–35.
- Masuti, Sagar, Sylvain D. Barbot, Shun-Ichiro Karato, Lujia Feng, and Paramesh Banerjee. 2016. “Upper-Mantle Water Stratification Inferred from Observations of the 2012 Indian Ocean Earthquake.” *Nature* 538 (7625): 373–77.
- Masuti, Sagar, Shun-Ichiro Karato, Jennifer Girard, and Sylvain D. Barbot. 2019. “Anisotropic High-Temperature Creep in Hydrous Olivine Single Crystals and Its Geodynamic Implications.” *Physics of the Earth and Planetary Interiors* 290 (May): 1–9.
- Mecking, H., and U. F. Kocks. 1981. “Kinetics of Flow and Strain-Hardening.” *Acta Metallurgica*. [https://doi.org/10.1016/0001-6160\(81\)90112-7](https://doi.org/10.1016/0001-6160(81)90112-7).
- Menezes, R. A., and W. D. Nix. 1971. “Average Internal Stresses in LiF Single Crystals during High Temperature Creep.” *Acta Metallurgica* 19 (7): 645–49.
- Meyssonier, J., and A. Goubert. 1994. “Transient Creep of Polycrystalline Ice under Uniaxial Compression: An Assessment of Internal State Variable Models.” *Annals of Glaciology*. <https://doi.org/10.1017/s0260305500010983>.
- Milička, Karel, Zuzanka Trojanová, and Pavel Lukáč. 2007. “Internal Stresses during Creep of Magnesium Alloys at 523 K.” *Materials Science and Engineering: A* 462 (1–2): 215–19.
- Mughrabi, H. 1980. “Strength of Metals and Alloys.” In *Proc. Intern. Conf. Metals and Alloys*, Ed. by P. Haasen, V. Gerold and G. Kostorz, Pergamon, Oxford, 1615.
- Müller, W., M. Biberger, and W. Blum. 1992. “Subgrain-Boundary Migration during Creep of LiF III. Stress Reduction Experiments.” *Philosophical Magazine A*. <https://doi.org/10.1080/01418619208201586>.
- Nimmo, F., and U. H. Faul. 2013. “Dissipation at Tidal and Seismic Frequencies in a Melt-Free, Anhydrous Mars.” *Journal of Geophysical Research: Planets* 118 (12): 2558–69.
- Nimmo, F., U. H. Faul, and E. J. Garnero. 2012. “Dissipation at Tidal and Seismic Frequencies in a Melt-Free Moon.” *Journal of Geophysical Research: Planets* 117 (E9). <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2012JE004160>.
- Nix, W. D., and B. Ilshner. 1979. “Mechanisms Controlling Creep of Single Phase Metals and Alloys.” In *Strength of Metals and Alloys*, edited by P. Haasen, V. Gerold, and G. Kostorz, 1503–30. Pergamon.
- Poirier, J. P. 1977. “Microscopic Creep Models and the Interpretation of Stress-Drop Tests during Creep.” *Acta Metallurgica*. [https://doi.org/10.1016/0001-6160\(77\)90178-x](https://doi.org/10.1016/0001-6160(77)90178-x).
- Pollitz, Fred F. 2005. “Transient Rheology of the Upper Mantle beneath Central Alaska Inferred from the Crustal Velocity Field Following the 2002 Denali Earthquake.” *Journal of Geophysical Research* 110 (B8). <https://doi.org/10.1029/2005jb003672>.
- Pollitz, Fred F., Roland Bürgmann, and Paramesh Banerjee. 2006. “Post-Seismic Relaxation Following the Great 2004 Sumatra-Andaman Earthquake on a Compressible Self-Gravitating Earth.” *Geophysical Journal International* 167 (1): 397–420.
- Post, Robert L., Jr. 1977. “High-Temperature Creep of Mt. Burnet Dunite.” *Tectonophysics* 42 (2–4): 75–110.
- Qiu, Qiang, James D. P. Moore, Sylvain Barbot, Lujia Feng, and Emma M. Hill. 2018. “Transient

- Rheology of the Sumatran Mantle Wedge Revealed by a Decade of Great Earthquakes.” *Nature Communications* 9 (1): 995.
- Rumpker, Georg, and Detlef Wolff. 1996. “Viscoelastic Relaxation of a Burgers Half-Space: Implications for the Interpretation of the Fennoscandian Uplift.” *Geophysical Journal International*. <https://doi.org/10.1111/j.1365-246x.1996.tb07036.x>.
- Smith, B. K., and F. O. Carpenter. 1987. “Transient Creep in Orthosilicates.” *Physics of the Earth and Planetary Interiors*. [https://doi.org/10.1016/0031-9201\(87\)90033-1](https://doi.org/10.1016/0031-9201(87)90033-1).
- Stone, D. S. 1991. “Scaling Laws in Dislocation Creep.” *Acta Metallurgica et Materialia* 39 (4): 599–608.
- Stone, D. S., T. Ploorkhol, and R. F. Cooper. 2004. “Similarity and Scaling in Creep and Load Relaxation of Single-Crystal Halite (NaCl).” *Journal of Geophysical Research, [Solid Earth]* 109 (B12). <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2004JB003064>.
- Sundberg, Marshall, and Reid F. Cooper. 2010. “A Composite Viscoelastic Model for Incorporating Grain Boundary Sliding and Transient Diffusion Creep; Correlating Creep and Attenuation Responses for Materials with a Fine Grain Size.” *Philosophical Magazine* 90 (20): 2817–40.
- Suzuki, Taira, Shin Takeuchi, and Hideo Yoshinaga. 2013. *Dislocation Dynamics and Plasticity*. Springer Science & Business Media.
- Takeuchi, S., and A. S. Argon. 1976. “Steady-State Creep of Single-Phase Crystalline Matter at High Temperature.” *Journal of Materials Science* 11 (8): 1542–66.
- Taylor, Geoffrey Ingram. 1934. “The Mechanism of Plastic Deformation of Crystals. Part I.—Theoretical.” *Proceedings of the Royal Society of London. Series A, Containing Papers of a Mathematical and Physical Character* 145 (855): 362–87.
- Thieme, M., S. Demouchy, D. Mainprice, F. Barou, and P. Cordier. 2018. “Stress Evolution and Associated Microstructure during Transient Creep of Olivine at 1000–1200 °C.” *Physics of the Earth and Planetary Interiors*. <https://doi.org/10.1016/j.pepi.2018.03.002>.
- Toriumi, Mitsuhiro. 1982. “Grain Boundary Migration in Olivine at Atmospheric Pressure.” *Physics of the Earth and Planetary Interiors* 30 (1): 26–35.
- Toriumi, Mitsuhiro, and Shun-Ichiro Karato. 1978. “Experimental Studies on the Recovery Process of Deformed Olivines and the Mechanical State of the Upper Mantle.” *Tectonophysics* 49 (1): 79–95.
- Wallis, David, Lars N. Hansen, T. Ben Britton, and Angus J. Wilkinson. 2016. “Geometrically Necessary Dislocation Densities in Olivine Obtained Using High-Angular Resolution Electron Backscatter Diffraction.” *Ultramicroscopy* 168 (September): 34–45.
- . 2017. “Dislocation Interactions in Olivine Revealed by HR-EBSD.” *Journal of Geophysical Research, [Solid Earth]* 122 (10): 7659–78.
- Wallis, David, Lars N. Hansen, Kathryn M. Kumamoto, Christopher A. Thom, Oliver Plümper, Markus Ohl, William B. Durham, et al. 2020. “Dislocation Interactions during Low-Temperature Plasticity of Olivine and Their Impact on the Evolution of Lithospheric Strength.” *Earth and Planetary Science Letters* 543 (August): 116349.
- Wallis, David, Lars N. Hansen, Angus J. Wilkinson, and Ricardo A. Lebensohn. 2020. “Dislocation Interactions in Olivine Control Postseismic Creep of the Upper Mantle.” *arXiv [cond-Mat.mtrl-Sci]*. arXiv. <http://arxiv.org/abs/2006.05195>.
- Wallis, D., L. N. Hansen, T. B. Britton, and A. J. Wilkinson. 2019. “High-Angular Resolution Electron Backscatter Diffraction as a New Tool for Mapping Lattice Distortion in Geological Minerals.” *Journal of Geophysical Research: Solid Earth*. <https://doi.org/10.1029/2019jb017867>.
- Wang, Lin, Stephan Blaha, Zsanett Pintér, Robert Farla, Takaaki Kawazoe, Nobuyoshi Miyajima, Katsuyoshi Michibayashi, and Tomoo Katsura. 2016. “Temperature Dependence of [100](010) and [001](010) Dislocation Mobility in Natural Olivine.” *Earth and Planetary Science Letters* 441 (May): 81–90.

- Wang, Zichao, Shun-Ichiro Karato, and Kiyoshi Fujino. 1993. "High Temperature Creep of Single Crystal Strontium Titanate (SrTiO₃): A Contribution to Creep Systematics in Perovskites." *Physics of the Earth and Planetary Interiors* 79 (3): 299–312.
- . 1996. "High Temperature Creep of Single Crystal Gadolinium Gallium Garnet." *Physics and Chemistry of Minerals*. <https://doi.org/10.1007/bf00202301>.
- Weertman, J. 1957. "Steady-State Creep of Crystals." *Journal of Applied Physics* 28 (10): 1185–89.
- Wilkinson, Angus J., Graham Meaden, and David J. Dingley. 2006. "High-Resolution Elastic Strain Measurement from Electron Backscatter Diffraction Patterns: New Levels of Sensitivity." *Ultramicroscopy* 106 (4-5): 307–13.
- Wilson, J. F., and N. K. Wilson. 1966. "Kinetics of Stress Relaxation in Metals." *Transactions of the Society of Rheology* 10 (1): 399–418.
- Wu, X. J., and A. S. Krausz. 1994. "A Kinetics Formulation for Low-Temperature Plasticity." *Journal of Materials Engineering and Performance* 3 (1): 169–77.
- Yuen, David A., Roberto C. A. Sabadini, Paolo Gasperini, and Enzo Boschi. 1986. "On Transient Rheology and Glacial Isostasy." *Journal of Geophysical Research*, Geophys. Monogr, 91 (B11): 11420.
- Zhong, Shijie, and A. B. Watts. 2013. "Lithospheric Deformation Induced by Loading of the Hawaiian Islands and Its Implications for Mantle Rheology." *Journal of Geophysical Research*, [Solid Earth] 118 (11): 2013JB010408.