Cloud patterns have four interpretable dimensions

Martin Janssens¹, Jordi Vilà-Guerau de Arellano², Marten Scheffer³, Coco Antonissen⁴, Pier Siebesma⁴, and Franziska Glassmeier⁵

¹Wageningen University & Research
²Wageningen University Research
³Wageningen University
⁴TU Delft
⁵TU Delft

November 22, 2022

Abstract

Shallow cloud fields over the subtropical ocean exhibit many spatial patterns. The frequency of occurrence of these patterns can change under global warming. Hence, they may influence subtropical marine clouds' climate feedback. While numerous metrics have been proposed to quantify cloud patterns, a systematic, widely accepted description is still missing. Therefore, this paper suggests one. We compute 21 metrics for 5000 satellite scenes of shallow clouds over the subtropical Atlantic Ocean and translate the resulting dataset to its principal components (PCs). This yields a unimodal, continuous distribution without distinct classes, whose first four PCs explain 82% of all 21 metrics' variance. The PCs correspond to four interpretable dimensions: *Characteristic length, void size, directional alignment* and horizontal *cloud-top height variance*. These dimensions span a space in which an effective pattern description can be given, which may be used to better understand the patterns' underlying physics and feedback on climate.

Cloud patterns have four interpretable dimensions

Martin Janssens¹, Jordi Vilà-Guerau de Arellano¹, Marten Scheffer¹, Coco Antonissen², A. Pier Siebesma²³, Franziska Glassmeier²

¹Wageningen University & Research ²Delft University of Technology ³Royal Netherlands Meteorological Institute

Key Points:

1

2 3

4 5 6

7

13

8	• Shallow cloud field patterns in satellite observations are quantified by 21 metrics and
9	follow a unimodal, continuous distribution.
10	• Most existing metrics are redundant; 4 principal components capture 82% of the
11	variance of 21 metrics.
12	• Characteristic length, void size, directional alignment and cloud-top height variance
13	combine to effectively describe the patterns.

Corresponding author: Martin Janssens, martin.janssens@wur.nl

14 Abstract

Shallow cloud fields over the subtropical ocean exhibit many spatial patterns. The frequency 15 of occurrence of these patterns can change under global warming. Hence, they may influence 16 subtropical marine clouds' climate feedback. While numerous metrics have been proposed 17 to quantify cloud patterns, a systematic, widely accepted description is still missing. There-18 fore, this paper suggests one. We compute 21 metrics for 5000 satellite scenes of shallow 19 clouds over the subtropical Atlantic Ocean and translate the resulting dataset to its prin-20 cipal components (PCs). This yields a unimodal, continuous distribution without distinct 21 classes, whose first four PCs explain 82% of all 21 metrics' variance. The PCs correspond 22 to four interpretable dimensions: Characteristic length, void size, directional alignment and 23 horizontal *cloud-top height variance*. These dimensions span a space in which an effective 24 pattern description can be given, which may be used to better understand the patterns' 25 underlying physics and feedback on climate. 26

27 Plain Language Summary

Satellite images show that clouds which develop in the lowest five kilometres of the 28 atmosphere organise into many visually distinct patterns. Because different patterns have 29 different radiative properties, a change in the relative occurrence of a pattern may influence 30 Earth's response to warming. To study this effect, the patterns must first be quantified; 31 numerous metrics have been developed for this task. In this paper, we compute 21 such 32 33 metrics for 5000 cloud fields observed by satellite over the Atlantic Ocean east of Barbados. We show that the information contained in the 21 metrics can already very accurately be 34 described by only 4 derived metrics, which capture a cloud field's typical cloud size, the size 35 of connected clear sky patches, the clouds' degree of directional alignment and variance in 36 cloud-top height. Combinations of these 4 metrics do not identify the existence of typical 37 patterns, as previously suggested. However, they form a new, effective and interpretable 38 pattern description, which can be used to better understand how cloud fields organise and 39 how this impacts the wider climate system. 40

41 **1** Introduction

Shallow cumulus clouds are the most abundant cloud type over the tropical oceans 42 (Johnson et al., 1999), but result from many interacting processes and scales. This combi-43 nation makes them the most uncertain aspect of how clouds will feed back onto a warming 44 climate (e.g. Bony & Dufresne, 2005; Schneider et al., 2017). Several mechanisms that gov-45 ern this response have recently been uncovered (Rieck et al., 2012; Bretherton, 2015; Klein 46 et al., 2017). However, the origins and sensitivity of the rich spectrum of spatial patterns ex-47 hibited by shallow cloud fields has remained rather unexplored (Nuijens & Siebesma, 2019). 48 Such spatial patterns alter precipitation distributions in cloud resolving simulations of deep 49 convection in warmer conditions (Muller & Held, 2012; Tobin et al., 2012); recent research 50 indicates that spatial patterning may influence the low cloud climate feedback too (Bony et 51 al., 2020). Establishing this contribution of shallow cloud-field patterns and its underlying 52 physics are therefore important research objectives. 53

The first step of such research is to classify or quantitatively measure any characteristic 54 of the horizontal dimension of a shallow cloud field. Two comprehensive, complementing 55 approaches were recently proposed: Expert visual inspection, which returns subjective, but 56 interpretable classes of patterns (Stevens et al., 2019) and unsupervised machine learn-57 ing, which is challenging to interpret, but gives more objectively inferred pattern measures 58 (Denby, 2020). A third, more traditional approach is to compute one or more human-59 defined metrics; these are both interpretable and objective and are therefore considered in 60 this paper. 61

Quantified patterns are often associated with a quantity called "organisation". This 62 term has consequently taken on numerous interpretations. It is often synonymous with 63 "aggregation" in studies of deep convection (Tobin et al., 2012; White et al., 2018; Hol-64 loway et al., 2017), sometimes characterised as the regular, random or clustered structure 65 of nearest neighbour distances of cloud objects (Weger et al., 1992; Seifert & Heus, 2013; 66 Tompkins & Semie, 2017), or connected to cloud scale (Neggers et al., 2019; Bony et al., 67 2020). However, cloud field organisation has also been defined by metrics of fractal analysis 68 (Cahalan & Joseph, 1989), directional alignment (Brune et al., 2018), subcritical perco-69 lation (Windmiller, 2017) or spatial variance (de Roode et al., 2004; Wood & Hartmann, 70 2006). While this makes it difficult to objectively define and discuss organisation, all these 71 interpretations share the same aim: Quantifying cloud field patterns. Hence, this diversity 72 can potentially also be harnessed to distinguish between different patterns. 73

The aim of this paper is therefore to systematically extract the independent information 74 encapsulated by the set of metrics associated with "cloud field organisation" in literature, 75 and to use this information to describe and interpret cloud field patterns as effectively as 76 possible. We first compute 21 diverse metrics for 5000 satellite observations of mesoscale 77 cloud fields in the trades and synthesise these in a multivariate distribution (section 2). Next, 78 we show that these metrics vary primarily along 4 principal components, allowing drastic 79 dimensionality reduction (section 3.1). Analysis of these main principal components results 80 in a pattern description that is remarkably effective, in addition to being interpretable and 81 objective (section 3.2). We then highlight several approaches to approximate the principal 82 components that balance the description's complexity and accuracy (section 3.3). Finally, we 83 demonstrate and discuss the ability of our description to characterise previously diagnosed 84 and novel regimes of characteristic patterns (section 3.4), before concluding (section 4). 85

⁸⁶ 2 Constructing a cloud field pattern distribution

2.1 Data

Following Stevens et al. (2019) and Bony et al. (2020), we concentrate on shallow, 88 subtropical clouds in the marine North Atlantic trades east of Barbados (20°-30°N, 48°-89 58° W), which are representative for the entire trades (Medeiros & Nuijens, 2016). Our 90 cloud fields stem from the MODIS instrument borne by NASA's Aqua and Terra satellites. 91 Specifically, we sample daytime overpasses during December-May 2002-2020 and directly use 92 the level 2 Cloud Water Path (CWP), Cloud-Top Height (CTH) and cloud mask products 93 at 1km resolution (Platnick et al., 2015) as basis for our metrics. Fig. S1 shows that 94 the results are not overly sensitive to resolution. We only interpret pixels classified as 95 "confidently cloudy" by the cloud mask algorithm as cloud. 96

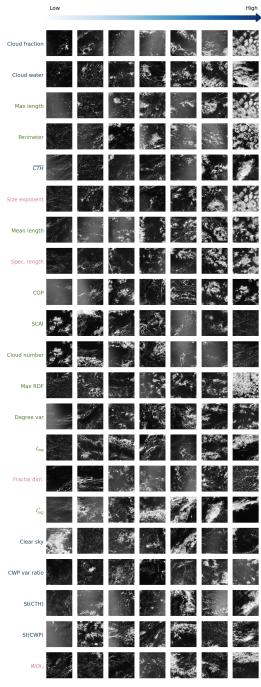
Our data points are scenes of cloud fields, which are $512 \text{km} \times 512 \text{km}$ subsets sampled 97 within the $10^{\circ} \times 10^{\circ}$ observation region. To boost the size of our dataset, scenes are allowed 98 to overlap 256km. We attempt to minimise the impact of errors and biases in remotely 99 sensed cloud products by rejecting scenes with i) high clouds such as cirrus wisps, if more 100 than 20% of the clouds' tops lie above 5km, ii) overly large sensor zenith angle, if this 101 angle exceeds 45°, following e.g. Wood and Field (2011) and iii) sunglint errors, manually 102 excluding scenes where these are visually found to influence the cloud mask. A set of 5004 103 scenes remains. 104

105

87

2.2 Metrics and dimensionality reduction

To appropriately capture the body of existing organisation metrics, we require them to meet either of the following two criteria: i) Are they perceived to capture a unique aspect of the patterns? or ii) do they frequently recur or recently first appear in literature? Additionally, they must be easy to interpret. This procedure (see tab. S1 for details) diagnoses 21 metrics, which broadly divide into three categories: Statistical moments of



- Field statistics - Object metrics - Scale decomposition metrics

Figure 1. Visual representation of scenes ordered by metrics derived from three categories (text colour) and sampled at linear intervals. Bright backgrounds stem from sunglint, which is accounted for in metric computations.

physical cloud field properties, object-based metrics, and attributes of scale decompositions.
 These metrics are briefly introduced below, visually presented in fig. 1 and further detailed
 in Text S1.

Statistical moments of cloud field properties comprise measures of typical cloud mass and area: The cloud mask's coverage fraction (Cloud fraction), the CWP's scene integral (Cloud water) and standard deviation (St(CWP)) and the variance ratio for "mesoscale aggregation" of moisture proposed by Bretherton and Blossey (2017) (CWP var. ratio), here applied only to cloud water. Furthermore, this class contains measures of the clouds' vertical extent: The mean and standard deviation of cloud-top height (CTH and St(CTH) respectively).

Object-based metrics measure size, shape and relative positioning of individual cloud 121 segments, which are identified from cloud mask fields using 4-connectivity labelling. To 122 avoid artefacts at the resolution scale, objects of a smaller dimension than four times the 123 instrument resolution are ignored. Our results are not sensitive to the chosen connectivity 124 scheme or minimum object size (see fig. S1). The resulting metrics further divide into 125 two categories: Scene statistics of individual object properties and measures of the spa-126 tial distribution of the objects. The first category includes the mean and maximum object 127 length (Mean length, Max length), the number of objects (Cloud number) and the mean ob-128 ject perimeter (Perimeter); the second comprises the Simple Convective Aggregation Index 129 (SCAI) (Tobin et al., 2012), Convective Organisation Potential (COP) (White et al., 2018), 130 the peak of the average radial distribution function (Rasp et al., 2018) (Max RDF), the 131 degree variance (Degree var) of the cloud objects' nearest-neighbour network representation 132 (Glassmeier & Feingold, 2017) and the Organisation Index (I_{org}) (Weger et al., 1992), of 133 which we include two versions. The first, most commonly applied form, compares the cloud 134 field nearest-neighbour cumulative density function (NNCDF) to a Weibull distribution. 135 The second variant (I_{org}^*) compares it to an inhibition NNCDF that accounts for object size 136 and therefore is less likely to erroneously predict regularity in the cloud fields (Benner & 137 Curry, 1998). This metric is similar to that introduced by Pscheidt et al. (2019). 138

We compute four metrics from scale decompositions: The size exponent of the cloud 139 object size distribution modelled as a power law (Size exponent), the box-counting dimension 140 of cloud boundaries in the cloud mask field (Fractal dim.), the Spectral length scale as defined 141 by Jonker et al. (1999) and the deviation of variance from the mean in the horizontal, vertical 142 or diagonal orientations of the cloud water field's stationary wavelet spectrum (WOI_3) 143 (Brune et al., 2018). In this paper, we use these metrics as discriminators between individual 144 cloud fields, not to measure their cumulative scaling properties. Finally, we introduce a novel 145 metric: A scene's largest, rectangular, contiguous cloud-free area (Clear sky), as a simple 146 measure of *lacunarity*, the degree to which continuous areas without clouds dominate a 147 scene. 148

We describe patterns as a linear combination of the computed metrics, which are standardised to weight them equally. Since many metrics in fig. 1 strongly correlate (see fig. S2), they are treated to a Principal Component Analysis (PCA, e.g. Abdi and Williams (2010)). This transforms the metrics to an orthogonal basis whose components (principal components - PCs) explain the maximum variance in the dataset. If a *small* number of PCs (orthogonal dimensions) can accurately capture the metric set's variance, these form an effective pattern description.

¹⁵⁶ **3** Describing patterns

157

3.1 A four-dimensional pattern distribution

Figure 2 shows uni- and bivariate kernel density estimates on planes spanned by the first four PCs of the metric distribution, annotated with the fractional variance of the dataset explained by each PC (explained variance ratio - EVR). It reveals that multiple

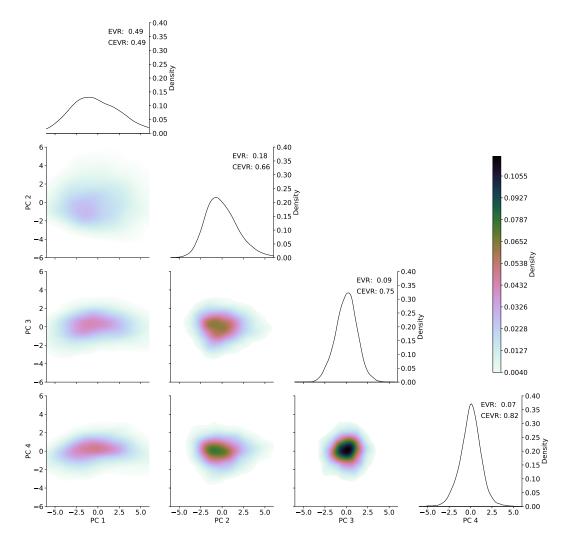


Figure 2. Univariate (diagonal, density on y-axis) and bivariate (off-diagonal, density in colour) Gaussian kernel density estimates of the first four principal components (PCs) of the pattern distribution. The annotations EVR and CEVR denote the individual and cumulative explained variance ratio of each PC, respectively. Bandwidths for the Gaussian kernels are computed using Scott's rule (Scott, 1992).

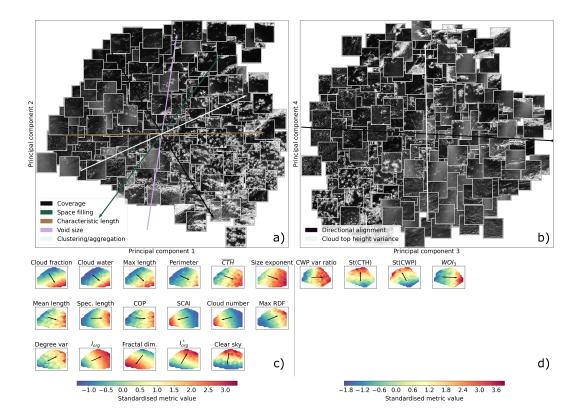


Figure 3. Top: Images of scenes projected onto planes spanned by the first and second (a) and third and fourth (b) PCs of the metric distribution, overlaid by arrows oriented along the mean gradient of several metric groups (see main text). Bottom: Filled contours of standardised metric values that have in excess of 50% of their variance explained by the first (c) and second (d) plane, constructed by piecewise linear barycentric interpolation and overlaid by an arrow pointing along the mean gradient. Subfigures b) and d) are rotated counter-clockwise by 49° in-plane to improve clarity of visualisation.

PCs (dimensions) are needed to capture the multivariate distribution's cumulative EVR 161 (CEVR) appropriately. However, the first PC is by far the most influential (EVR=0.49 -162 widest distribution). Furthermore, the CEVR of the first two PCs already rises to 0.66, 163 while including 3 and 4 of the 21 original dimensions explains 75% and 82% of the dataset's 164 variance, respectively. After the fourth PC, EVR quickly deflates (PCs 5-9 have EVRs 165 of 0.04, 0.03, 0.03, 0.02, 0.02), dropping below 0.01 after the tenth PC (fig. S3). These 166 statistics show that four PCs effectively capture the information in all 21 metrics. Therefore, 167 we reduce our 21-dimensional metric set to these four PCs. 168

Of course, truncating the PCA after precisely four components remains somewhat arbitrary. Yet, this choice strikes a useful balance between including enough dimensions to effectively describe patterns and sufficiently few dimensions to interpret them. This claim is visually supported by fig. 3 a) and b) (fig. S3 adds quantitative evidence): Combinations of PC1 and PC2 (fig. 3 a) consistently and coherently position visually similar (different) scenes close to (far from) each other. PC3 and PC4 (fig. 3 b) ably reveal further distinctions. Hence, linear combinations of these four PCs form an effective pattern description.

¹⁷⁶ **3.2** An interpretable pattern description

Our four-dimensional pattern description is not only effective; by relating the PCs to 177 their underpinning metrics, it can also be interpreted. This interpretation is facilitated by 178 fig. 3 c) and d), which show the in-plane gradient and mean direction of change of metrics 179 that predominantly vary in the planes depicted in fig. 3 a) and b) respectively. By averaging 180 the gradients of several similarly varying metrics, we identify a meaningful vocabulary that 181 labels several directions of change in the two planes (arrows in fig. 3 a and b). Using 182 this vocabulary, we name the principal components and relate them to several common 183 interpretations of organisation. 184

Strikingly, 17/21 metrics mainly describe variations in the first two PCs (fig. 3 c, see also fig. S4). These metrics derive from all three categories (field statistics, objects and scale decompositions) and point in a rather continuous spectrum of directions, offering a remarkable number of interesting choices for interpreting fig. 3 a):

- 1. Coverage (Arrow in fig. 3 a represents the mean gradient of Cloud fraction, Max length and Cloud water)
- ¹⁹¹ 2. Space filling (Fractal dim., I_{org}^*)
 - 3. Characteristic length (Spectral length scale, Size exponent, Mean length)
- ¹⁹³ 4. Void size (Clear sky)

192

¹⁹⁴ 5. Aggregation or clustering (I_{org} , SCAI, Cloud number, Max RDF), as commonly as-¹⁹⁵ sociated with deep convective organisation (Tompkins & Semie, 2017; Tobin et al., ¹⁹⁶ 2012).

We adopt the two directions that best align themselves with the PCs as names for our pattern description's first two dimensions: *Characteristic length* and *void size*. We find it both intuitive and beautiful that these two dimensions, which respectively measure the typical scale of clouds and the complementary clear sky space between them, naturally emerge from our approach.

Linear combinations of the PCs can construct different terms in the quintet above. For instance, *clustering/aggregation* differs only subtly from *characteristic length*, assigning slightly more importance to voids between cloud clusters. *Space filling* weights voids even more heavily. Finally, *coverage* distinguishes itself from *void size* by assigning marginally more importance to characteristic length. Hence, the same aspects of the patterns in fig. 3 a) can be described with different pairs of terms.

Several such pairs are already indirectly recognised as central traits of "organisation". 208 For instance, Seifert and Heus (2013) suggest that both a spectral length scale (character-209 istic length) and I_{org} (clustering) may be needed to discriminate between various modes 210 of organisation; Neggers et al. (2019) identify organisation as a combination of maximum 211 cloud size (coverage) and typical nearest-neighbour distances between smaller clouds (space 212 filling); chapter 5 of van Laar (2019) distinguishes "cloud field characteristics" (cloud frac-213 tion, maximum cloud size - coverage) from "organisation parameters" (I_{org} , SCAI, COP -214 clustering) and Bony et al. (2020) span their planar description of organisation with mean 215 length (characteristic length) and I_{org} (clustering). The arrows in Figure 3 relate all these 216 interpretations to each other. 217

However, our four-dimensional pattern description goes beyond these common, twodimensional interpretations of organisation. Figure 3 d) shows that the third and fourth PC distinguish patterns with different directional alignment (WOI_3) of the scene's larger scales (CWP var ratio) and those with different horizontal variance of vertical cloud development (St(CTH)). Hence, variations in PC 3 and PC4 can be understood as combinations of *directional alignment* and *cloud-top height variance*. Summarising, we propose to think of cloud field patterns, as described by organisation metrics, as a linear combination of the 4 PCs: *Characteristic length, void size, directional alignment* and *cloud-top height variance*, each term contributing a dimension that is uncorrelated to the others. However, many valid interpretations exist, especially of the first two dimensions. Therefore, the consistency with which "organisation" is understood can be considerably advanced by using the relationships between the various interpretations established in this section.

3.3 Selecting metric subsets

While four PCs describe patterns remarkably well, they still require input from all 232 metrics. If added interpretability or less computation is desired, one might approximate 233 the PCs with a subset of metrics. This approach challenges each chosen metric to do 234 considerably more work than merely inspiring an interpretation of the PCs, as in the previous 235 section, since no metric subset is fully orthogonal or optimally variance-capturing. Moreover, 236 it is often not obvious that a given metric is much better suited to approximate a PC than 237 a similarly varying one. This problem is illustrated by applying sparse PCA (Zou et al., 238 2006) to our data. Despite optimising a cost function that explicitly balances the accuracy 239 of the approximate PCs with how many metrics contribute to them, this technique cannot 240 241 robustly indicate metric subsets (see fig. S5).

One practical way to compose a subset nonetheless is choosing one metric that most 242 closely correlates to each PC (Cadima & Jolliffe, 1995). This approach selects the Spectral 243 length scale, Clear sky, WOI_3 and St(CTH) (CEVR=0.59) and is a reasonable approxi-244 mation of the PC description (CEVR=0.82). If one's primary interest is in the first two 245 dimensions of the pattern distribution, several roughly orthogonal metric pairs competently 246 estimate the plane in fig. 3 a). Examples include Spectral slope and Clear sky (CEVR=0.31), 247 Cloud fraction and Fractal dim. (CEVR=0.31) or Perimeter and I_{org}^* (CEVR=0.30). All 248 three pairs sacrifice explained variance compared to two PCs (CEVR=0.66). Yet, they cap-249 ture far more information than various metric combinations considered in literature, e.g. 250 Cloud number and I_{org} (Bony et al., 2020, CEVR=0.18), I_{org} and Fractal dim. (Denby, 251 2020, CEVR=0.20), Spec. Length and I_{org} (Seifert & Heus, 2013, CEVR=0.19) or I_{org} , 252 SCAI, COP and Max RDF (van Laar, 2019, CEVR=0.26). Therefore, we recommend to 253 always assess the orthogonality and EVR of one's metrics with a PCA, before optionally 254 selecting a metric subset that approximates their desirable properties appropriately. 255

3.4 Regimes of patterns

256

Asking how many dimensions cloud field patterns possess is not equal to asking how many fundamental types of cloud patterns exist. Dividing clouds into distinct classes (e.g. cumulus or cirrus) is a classical approach, which recently inspired efforts to also classify shallow cloud field patterns, using both the human eye (Stevens et al., 2019) and metrics (Bony et al., 2020). We compare our pattern description to these classes ("sugar", "gravel", "fish" and "flowers") by identifying seven k-means clusters in the four-dimensional PC distribution (fig. 4).

Scenes arguably dominated by "sugar" and "gravel" reside in clusters 5 (brown) and 264 3 (maroon). These patterns should, in the terminology from section 3.2, be understood as 265 small-scale with rather small voids (or disaggregated/unclustered); "gravel" distinguishes 266 itself through its higher cloud-top height variance and low directional alignment (see also 267 left side of fig. 3 b). Cluster 1 (navy) comprises i.a. "fish", which shares gravel's void size, 268 cloud-top height variance and low degree of directional alignment, only at larger scales. 269 Finally, one may see "flowers" in cluster 7 (blue), as large-scale, aggregated structures with 270 little directional alignment and low cloud-top height variance. 271

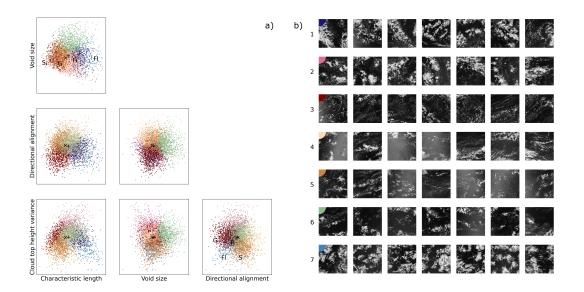


Figure 4. Seven regimes of the 4D pattern description, identified as k-means clusters of different colour: a) Scenes scattered over planes defined by the first four PCs, each normalised to unit variance, named using the convention from section 3.2; b) seven examples of scenes in each regime. Pluses and crosses indicate the distribution's mean and mode, respectively. S, G, Fi and Fl suggest typical locations for the "sugar", "gravel", "fish" and "flowers" patterns diagnosed by Stevens et al. (2019), in the two planes shown in fig. 3, determined by eye.

The natural emergence of these regimes from our systematic metric analysis is encour-272 agingly consistent with human pattern identification (Stevens et al., 2019) and solidifies 273 Bony et al. (2020)'s conclusion that these patterns can be objectively identified. However, 274 even in an unrealistic scenario where all scenes in these four regimes could unambiguously 275 be labelled sugar, gravel, fish or flowers, they would contain only 52% of the observations in 276 our dataset. Figure 4 indicates several other regimes that differ in important regards. For 277 instance, many scenes possess vast voids (cluster 6, sea green). In this regime, clouds likely 278 affect the region's climate much less than sugar, gravel, fish or flowers, which all have higher 279 cloud cover. Analyses of the patterns' climate sensitivity must probably consider this and 280 other different regimes explicitly. 281

In fact, pattern classification is itself an approximation. The pattern distribution is *uni*-282 modal and continuous (fig. 2), and therefore does not inherently possess multiple "classes", 283 "clusters" or "modes". Breaking the continuum into clusters neglects subtly different pat-284 terns within a cluster. For instance, the band-like sub-regime at high directional alignment 285 in fig. 3 b) falls within cluster 4 (peach) in fig. 4, even if this sub-regime is visually distinct 286 from all displayed scenes in cluster 4. To capture such subtleties, we recommend shifting 287 focus from regimes, classes or clusters of patterns to a more fitting, continuous representa-288 tion. 289

Finally, many of the human-identified patterns (sugar, flowers) appear on our distribution's extremes (see also fig. 3 a) and b)). While this may explain why they are most easily distinguished by humans, they lie far from the distribution's statistical mean and mode (indicated by pluses and crosses respectively in fig. 4 a) and are thus not typical. Instead, the modal pattern is partial to smaller scales and voids, which characterise scenes with shallow, cold-pool dominated convection (clusters 3, 5) or processes on a wide range of scales (cluster 4); this space may be most relevant to the climatology of patterns.

²⁹⁷ 4 Conclusion and outlook

Research on the climate feedback of patterns in shallow trade-wind cloud fields requires 298 a consistently understood description of those patterns. In this paper, we have systemati-299 cally developed such a description for square, 500 km^2 satellite-observed cloud fields east of 300 Barbados. By projecting one new and 20 previously developed organisation metrics onto a 301 set of PCs, we show that cloud patterns can be effectively described as a 4-dimensional, lin-302 ear combination of characteristic length, void size, directional alignment and cloud-top height 303 variance. This description is objective and interpretable, in contrast to direct unsupervised 304 machine learning (objective, not usually interpretable) or human pattern identification (in-305 terpretable, not objective). It also demonstrates that patterns follow a continuous, unimodal 306 distribution without distinct classes and that visually striking patterns are extreme, rather 307 than typical. Future studies of the physics behind and climate impact of shallow cloud field 308 patterns can therefore rely either on our PCs or, if accuracy is less important, on metrics 309 that correlate closely to them. 310

The effectiveness of our approach may well extend to descriptions of deep convective 311 organisation. Many relationships between our metrics are consistent with those found for 312 deep convective cloud fields (Rempel et al., 2017; Brueck et al., 2020), suggesting that an 313 effective, low-dimensional description of deep convective organisation is attainable. Our 314 pattern description could also be used for forecast verification (Jolliffe & Stephenson, 2012), 315 using the pattern distribution's dimensions as matching criteria between model and obser-316 vation in similar fashion to e.g. the criteria developed by Wernli et al. (2008). In turn, 317 the forecast verification community may offer useful insights to descriptions of cloud field 318 patterns. 319

Finally, our approach can itself be refined in several regards. First, using predefined 320 metrics to describe patterns leaves potentially undiscovered information from the descrip-321 tion. Therefore, it may be fruitful to compare our approach to more unsupervised machine 322 learning (e.g. Denby, 2020). However, the completeness of a pattern description should ide-323 ally be assessed in terms of how fully the underlying processes are separated. This requires 324 process-resolving numerical simulations and/or temporally evolving observations, which link 325 the evolution of the pattern continuum to that of the atmospheric state. Next, our con-326 clusions are tied to our observation scales (1-500 km), meaning that we may inadequately 327 capture this scale window's extremes. Furthermore, we treat this scale window in an inte-328 gral sense and ignore patterns that appear on one scale, but may be cancelled by another 329 (Nair et al., 1998). Hence, a further refinement could be to consider pattern distributions 330 on a per-scale basis. Lastly, some subjectivity will likely remain in how different researchers 331 interpret "organisation". This attests the richness of the underlying patterns, which we 332 hope remains appreciated. 333

334 Acknowledgments

MJ and FG acknowledge support by The Branco Weiss Fellowship - Society in Science, 335 administered by ETH Zürich; FG also acknowledges a Veni grant from the Dutch Re-336 search Council (NWO). A.P.S acknowledges funding by the European Union's Horizon 337 2020 research and innovation programme under grant agreement no. 820829 (CONSTRAIN 338 project). The Aqua and Terra MODIS Clouds 5-Min L2 Swath 1km products, which contain 339 the cloud mask, cloud water path and cloud-top height products used in this study, were 340 extracted from NASA's Level-1 and Atmosphere Archive & Distribution System (LAADS) 341 Distributed Active Archive Center (DAAC) (http://dx.doi.org/10.5067/MODIS/MYD06 342 343 _L2.061; http://dx.doi.org/10.5067/MODIS/MOD06_L2.061). All download, preprocessing, metric computation and analysis was done using Python and its Numpy (van der Walt 344 et al., 2011), Pandas (McKinney, 2010) and Scipy (Virtanen et al., 2020) libraries. Field 345 processing and segmentation is done with Scikit Image (van der Walt et al., 2014), PCA and 346 clustering with Scikit Learn (Pedregosa et al., 2011), sparse PCA with Ristretto (Erichson et 347 al., 2020) and plots with Matplotlib (Hunter, 2007) and Seaborn (Waskom & the seaborn de-348

velopment team, 2020). The integrated code, along with detailed instructions on how to run it, are available in a living GitHub repository (https://github.com/martinjanssens/

cloudmetrics) and its frozen image at the time of submission (https://doi.org/10.6084/ m9.figshare.12687302.v1).

353 **References**

387

388

- Abdi, H., & Williams, L. J. (2010). Principal component analysis. Wiley interdisciplinary reviews: computational statistics, 2(4), 433–459.
- Benner, T. C., & Curry, J. A. (1998). Characteristics of small tropical cumulus clouds
 and their impact on the environment. *Journal of Geophysical Research: Atmospheres*, 103(D22), 28753–28767.
- Bony, S., & Dufresne, J.-L. (2005). Marine boundary layer clouds at the heart of tropical cloud feedback uncertainties in climate models. *Geophysical Research Letters*, 32(20).
- Bony, S., Schulz, H., Vial, J., & Stevens, B. (2020). Sugar, gravel, fish and flowers: Dependence of mesoscale patterns of trade-wind clouds on environmental conditions.
 Geophysical Research Letters.
- Bretherton, C. (2015). Insights into low-latitude cloud feedbacks from high-resolution
 models. Philosophical Transactions of the Royal Society A: Mathematical, Physical
 and Engineering Sciences, 373(2054), 20140415.
- Bretherton, C., & Blossey, P. (2017). Understanding mesoscale aggregation of shallow
 cumulus convection using large-eddy simulation. Journal of Advances in Modeling
 Earth Systems, 9(8), 2798–2821.
- Brueck, M., Hohenegger, C., & Stevens, B. (2020). Mesoscale marine tropical precipitation
 varies independently from the spatial arrangement of its convective cells. *Quarterly Journal of the Royal Meteorological Society*, 146(728), 1391–1402.
- Brune, S., Kapp, F., & Friederichs, P. (2018). A wavelet-based analysis of convective organ ization in ICON large-eddy simulations. *Quarterly Journal of the Royal Meteorological* Society, 144 (717), 2812–2829.
- Cadima, J., & Jolliffe, I. T. (1995). Loading and correlations in the interpretation of principle compenents. *Journal of applied Statistics*, 22(2), 203–214.
- Cahalan, R. F., & Joseph, J. H. (1989). Fractal statistics of cloud fields. Monthly weather review, 117(2), 261–272.
- Denby, L. (2020). Discovering the importance of mesoscale cloud organization through unsupervised classification. *Geophysical Research Letters*, 47(1), e2019GL085190.
- de Roode, S. R., Duynkerke, P. G., & Jonker, H. J. (2004). Large-eddy simulation: How large is large enough? *Journal of the atmospheric sciences*, 61(4), 403–421.
- Erichson, N. B., Zheng, P., Manohar, K., Brunton, S. L., Kutz, J. N., & Aravkin, A. Y.
 (2020). Sparse principal component analysis via variable projection. SIAM Journal on Applied Mathematics, 80(2), 977–1002.
 - Glassmeier, F., & Feingold, G. (2017). Network approach to patterns in stratocumulus clouds. Proceedings of the National Academy of Sciences, 114 (40), 10578–10583.
- Holloway, C. E., Wing, A. A., Bony, S., Muller, C., Masunaga, H., L'Ecuyer, T. S., ...
 Zuidema, P. (2017). Observing convective aggregation. Surveys in Geophysics, 38(6), 1199–1236.
- Hunter, J. D. (2007). Matplotlib: A 2d graphics environment. Computing in Science & Engineering, 9(3), 90–95. doi: 10.1109/MCSE.2007.55
- Johnson, R. H., Rickenbach, T. M., Rutledge, S. A., Ciesielski, P. E., & Schubert, W. H. (1999). Trimodal characteristics of tropical convection. *Journal of climate*, 12(8), 2397–2418.
- Jolliffe, I. T., & Stephenson, D. B. (2012). Forecast verification: a practitioner's guide in atmospheric science. John Wiley & Sons.
- Jonker, H. J., Duynkerke, P. G., & Cuijpers, J. W. (1999). Mesoscale fluctuations in scalars generated by boundary layer convection. *Journal of the atmospheric sciences*, 56(5), 801–808.

Klein, S. A., Hall, A., Norris, J. R., & Pincus, R. (2017). Low-cloud feedbacks from cloud-402 controlling factors: a review. In Shallow clouds, water vapor, circulation, and climate 403 sensitivity (pp. 135–157). Springer. 404 McKinney, W. (2010). Data Structures for Statistical Computing in Python. In Stéfan 405 van der Walt & Jarrod Millman (Eds.), Proceedings of the 9th Python in Science 406 Conference (p. 56 - 61). doi: 10.25080/Majora-92bf1922-00a 407 Medeiros, B., & Nuijens, L. (2016). Clouds at Barbados are representative of clouds across 408 the trade wind regions in observations and climate models. Proceedings of the National 409 Academy of Sciences, 113(22), E3062–E3070. 410 Muller, C. J., & Held, I. M. (2012). Detailed investigation of the self-aggregation of con-411 vection in cloud-resolving simulations. Journal of the Atmospheric Sciences, 69(8), 412 2551 - 2565.413 Nair, U., Weger, R., Kuo, K., & Welch, R. (1998). Clustering, randomness, and regularity 414 in cloud fields: 5. the nature of regular cumulus cloud fields. Journal of Geophysical 415 Research: Atmospheres, 103(D10), 11363–11380. 416 Neggers, R., Griewank, P., & Heus, T. (2019). Power-law scaling in the internal variability of 417 cumulus cloud size distributions due to subsampling and spatial organization. Journal 418 of the Atmospheric Sciences, 76(6), 1489–1503. 419 Nuijens, L., & Siebesma, A. P. (2019). Boundary layer clouds and convection over subtrop-420 ical oceans in our current and in a warmer climate. Current Climate Change Reports, 421 5(2), 80-94.422 Pedregosa, F., Varoquaux, G., Gramfort, A., Michel, V., Thirion, B., Grisel, O., ... Duch-423 esnay, E. (2011). Scikit-learn: Machine learning in Python. Journal of Machine 424 Learning Research, 12, 2825–2830. 425 Platnick, S., Ackerman, S., King, M., Menzel, P., Wind, G., & Frey, R. (2015).426 MODIS Atmosphere L2 Cloud Product (06-L2). Goddard Space Flight Center, USA: 427 NASA MODIS Adaptive Processing System. doi: http://dx.doi.org/10.5067/MODIS/ 428 MYD06_L2.061 429 Pscheidt, I., Senf, F., Heinze, R., Deneke, H., Trömel, S., & Hohenegger, C. (2019). How 430 organized is deep convection over germany? Quarterly Journal of the Royal Meteoro-431 logical Society, 145(723), 2366–2384. 432 Rasp, S., Selz, T., & Craig, G. C. (2018). Variability and clustering of midlatitude sum-433 mertime convection: Testing the Craig and Cohen theory in a convection-permitting 434 ensemble with stochastic boundary layer perturbations. Journal of the Atmospheric 435 Sciences, 75(2), 691–706. 436 Rempel, M., Senf, F., & Deneke, H. (2017). Object-based metrics for forecast verification of 437 convective development with geostationary satellite data. Monthly Weather Review, 438 145(8), 3161-3178.439 Rieck, M., Nuijens, L., & Stevens, B. (2012). Marine boundary layer cloud feedbacks in a 440 constant relative humidity atmosphere. Journal of the Atmospheric Sciences, 69(8), 441 2538 - 2550.442 Schneider, T., Teixeira, J., Bretherton, C. S., Brient, F., Pressel, K. G., Schär, C., & 443 Siebesma, A. P. (2017). Climate goals and computing the future of clouds. Nature 444 Climate Change, 7(1), 3–5. 445 Scott, D. W. (1992). Multivariate density estimation: theory, practice, and visualization. 446 John Wiley & Sons. 447 Seifert, A., & Heus, T. (2013). Large-eddy simulation of organized precipitating trade wind 448 cumulus clouds. Atmos. Chem. Phys, 13(11), 5631–5645. 449 Stevens, B., Bony, S., Brogniez, H., Hentgen, L., Hohenegger, C., Kiemle, C., ... others 450 (2019). Sugar, gravel, fish and flowers: Mesoscale cloud patterns in the trade winds. 451 Quarterly Journal of the Royal Meteorological Society, 1–12. 452 Tobin, I., Bony, S., & Roca, R. (2012). Observational evidence for relationships between the 453 degree of aggregation of deep convection, water vapor, surface fluxes, and radiation. 454 Journal of Climate, 25(20), 6885-6904. 455 Tompkins, A. M., & Semie, A. G. (2017). Organization of tropical convection in low 456

457	vertical wind shears: Role of updraft entrainment. Journal of Advances in Modeling
458	Earth Systems, $9(2)$, 1046–1068.
459	van der Walt, S., Colbert, S. C., & Varoquaux, G. (2011). The numpy array: a structure for
460	efficient numerical computation. Computing in Science & Engineering, 13(2), 22–30.
461	van der Walt, S., Schönberger, J. L., Nunez-Iglesias, J., Boulogne, F., Warner, J. D., Yager,
462	N., the scikit-image contributors (2014, 6). scikit-image: image processing in
463	Python. PeerJ, 2, e453. Retrieved from https://doi.org/10.7717/peerj.453 doi:
464	10.7717/peerj.453
465	van Laar, T. W. (2019). Spatial patterns in shallow cumulus cloud populations over a
466	heterogeneous surface (Doctoral dissertation, University of Cologne). Retrieved from
467	http://kups.ub.uni-koeln.de/id/eprint/10221
468	Virtanen, P., Gommers, R., Oliphant, T. E., Haberland, M., Reddy, T., Cournapeau, D.,
469	Contributors, S (2020). SciPy 1.0: Fundamental Algorithms for Scientific
470	Computing in Python. Nature Methods, 17, 261–272. doi: https://doi.org/10.1038/
471	s41592-019-0686-2
472	Waskom, M., & the seaborn development team. (2020, September). mwaskom/seaborn.
473	Zenodo. Retrieved from https://doi.org/10.5281/zenodo.592845 doi: 10.5281/
474	zenodo.592845
475	Weger, R., Lee, J., Zhu, T., & Welch, R. (1992). Clustering, randomness and regular-
476	ity in cloud fields: 1. theoretical considerations. Journal of Geophysical Research:
477	Atmospheres, 97(D18), 20519-20536.
478	Wernli, H., Paulat, M., Hagen, M., & Frei, C. (2008). SAL — a novel quality measure
479	for the verification of quantitative precipitation forecasts. Monthly Weather Review,
480	136(11), 4470-4487.
481	White, B., Buchanan, A., Birch, C., Stier, P., & Pearson, K. (2018). Quantifying the effects
482	of horizontal grid length and parameterized convection on the degree of convective
483	organization using a metric of the potential for convective interaction. Journal of the
484	Atmospheric Sciences, $75(2)$, $425-450$.
485	Windmiller, J. M. (2017). Organization of tropical convection (Doctoral dissertation,
486	Ludwig-Maximilian University of Munich). Retrieved from https://edoc.ub.uni
487	-muenchen.de/21245/
488	Wood, R., & Field, P. R. (2011). The distribution of cloud horizontal sizes. Journal of
489	$Climate, \ 24 \ (18), \ 4800 - 4816.$
490	Wood, R., & Hartmann, D. L. (2006). Spatial variability of liquid water path in marine low
491	cloud: The importance of mesoscale cellular convection. Journal of Climate, $19(9)$,
492	1748–1764.
493	Zou, H., Hastie, T., & Tibshirani, R. (2006). Sparse principal component analysis. Journal
494	of computational and graphical statistics, $15(2)$, $265-286$.

494

Supporting Information for "Cloud patterns have four interpretable dimensions"

DOI: 10.1002/

Martin Janssens¹, Jordi Vilà-Guerau de Arellano¹, Marten Scheffer¹, Coco

Antonissen², A. Pier Siebesma²³, Franziska Glassmeier²

 $^1\mathrm{Wageningen}$ University & Research

²Delft University of Technology

 $^3\mathrm{Royal}$ Netherlands Meteorological Institute

Contents of this file

- 1. Text S1 $\,$
- 2. Figures S1 to S5
- 3. Table S1 $\,$

Corresponding author: M. Janssens, Departments of Meteorology & Air Quality and Aquatic Ecology & Water Quality Management, Wageningen University & Research, Wageningen, Lumen Building, Droevendaalsesteeg 3a, 6708 PB Wageningen, The Netherlands, (martin.janssens@wur.nl)

Introduction

This supplement contains further descriptions of the metrics that we use to characterise our cloud field pattern distribution (Text S1). Specifically, we elaborate upon details of and justify choices made in their computation. Code that evaluates these metrics given input scenes of cloud mask, cloud water path and cloud top height can be found in our accompanying GitHub repository (https://github.com/martinjanssens/cloudmetrics) and Figshare copy of this repository at the time of publication https://figshare.com/ projects/Cloud_field_organisation_description_with_metrics/86303. The supplement also contains five figures, that quantify i) the sensitivity of our metric distribution to field resolution, object segmentation strategy and minimum cloud size, ii) the absolute Pearson correlation between all metrics, iii) the fraction of variance in each metric explained by every PC, iv) an estimate of the quality of our metric-based approach to approximating cloud field patterns and v) the sensitivity to free parameters of approximating principal components with a subset of metrics through sparse principal component analysis.

Text S1. - Details of metrics

Statistical moments of cloud field properties

We quantify several statistics of the extracted cloud field products. Some of these are straightforward computations that do not feature design choices (cloud fraction, total cloud water, standard deviation of cloud water over cloudy pixels). The other metrics require further qualification.

Mean and standard deviation of cloud top height ($\overline{\text{CTH}}$ and St(CTH) respectively) i) explicitly ignore clouds higher than 5km, as cirrus wisps were found to disproportionately affect the results otherwise and ii) only consider cloudy pixels. Higher-order moments of these fields were small and are therefore not included.

Cloud water variance ratio R (CWP var. ratio) is directly adopted from Bretherton and Blossey (2017), but instead of being applied to the total, vertically integrated moisture field, it is here only applied to the cloud water:

$$R = \frac{Std\left(\overline{CWP_b} - \overline{CWP}\right)}{Std(CWP)} \tag{1}$$

In this relation, $\overline{\cdot}$ denotes a domain average and CWP_b indicates the cloud water contained in blocks of 16x16 pixels.

Object-based metrics

Object-based metrics follow from segmenting the cloud mask field into N_o objects according to their 4-connectivity. To avoid artefacts at the grid scale, we only consider objects with areas larger than 4 pixels. Each extracted object covers an area A_i , such that a typical length scale for that object is $l_i = \sqrt{A_i}$.

Mean object size is defined as $\frac{1}{N_o} \sum_i l_i$

Max object size is defined as $\max l_i$.

Mean perimeter is derived by extracting the perimeter of each object P_i and defining the mean perimeter $\overline{P} = \frac{1}{N_o} \sum_i P_i$.

The Simple Convective Aggregation Index (SCAI) (Tobin et al., 2012) is defined as:

$$SCAI = \frac{N_o D_0}{N_{max}} \tag{2}$$

Where N_{max} is the number of pixels in a scene, $D_0 = \sqrt[N_p]{\prod_i^{N_p} d_i}$ is the geometric mean of Euclidian pairwise distance between all object centroids d_i and $N_p = N_o(N_o - 1)/2$.

The Convective Organisation Potential (COP) (White et al., 2018) is:

$$COP = \frac{1}{N_p} \sum_{i=0}^{N_o} \sum_{j=i+1}^{N_o} \frac{l_i + l_j}{\sqrt{\pi} d_{ij}}$$
(3)

Where d_{ij} now explicitly represents the distance between two object centroids.

Max RDF is the maximum value of the radial distribution function RDF(r) as proposed in Rasp, Selz, and Craig (2018):

$$RDF(r) = \frac{1}{N_i} \sum_{i} \frac{\sum_{r \le r_i < r+dr} 1}{L\left(\pi \left(r+dr\right)^2 - r^2\right)}$$
(4)

Where r_i are pairwise distances from the ith centroid to all other centroids, dr denotes the width of a radial annulus over which we sum such distances, L is the length of the scene's side, and N_i are the number of centroids that lie within a distance r_{max} from the domain edges. We only consider coordinates within a radius r_{max} from any original centroid. We

Degree variance of nearest-neighbour network representations of the scenes are quantified by constructing a Voronoi tessellation from the computed object centroids and measuring the variance in the degree (number of neighbours) distribution of the identified Voronoi cells.

 I_{org} (Weger et al., 1992) is included in two flavours. The first is the original metric, which integrates the area under the curve defined by the NNCDF, the cumulative density function of nearest neighbour distances d_N between object centroids (y axis) and the corresponding Weibull distribution (x axis):

$$W = 1 - \exp\left(\frac{N_o}{L^2}\pi d_N^2\right) \tag{5}$$

If the object centroids are scattered as a Poisson point process, they should follow W exactly, resulting in $I_{org} = 0.5$. $I_{org} < 0.5$ if they are regularly spaced; if they appear in clusters, $I_{org} > 0.5$. As pointed out by Benner and Curry (1998), this overestimates the regularity of the cloud field, because in reality separate cloud objects are inhibited from forming within the area covered by another object. To account for this, we also include a second version of I_{org} , which we name I_{org}^* . This metric compares the cloud field NNCDF to an inhibition NNCDF, which is constructed by randomly scattering N_o objects throughout the scene, provided that they do not fall within the circular area of an object that has already been placed. The computer-generated random positions of this approach are less robust than the Weibull distribution (Weger et al., 1992), but we

find that repeating the computations 3 times does not impact the resulting I_{org}^* below the third significant digit.

Scale decomposition metrics

Size exponent b is computed by counting all cloud sizes N_c in bins of exponentially increasing width, and fitting the resulting cloud size distribution with a power law:

$$\log N_c \propto b \log l \tag{6}$$

The average coefficient of determination R^2 of fitting this relation to all scenes is good: 0.923. We also investigated a fit according to subcritical percolation theory that incorporates an exponential term. However, undersampling of large cloud structures make such fits quite unrealistic on a per-scene basis, even though the fit converges when sampling a large number of scenes at similar cloud fraction (not shown). It is therefore likely that these cloud fields obey the rules of subcritical percolation. Yet, the parameters of the corresponding fit cannot reliably be identified on a per scene basis.

The box-counting dimension D_f (fractal dim.) of each cloud mask field is derived by counting the number of square boxes N_c of dimension l_b that are neither fully cloudfree nor fully cloudy (i.e. boxes that contain cloud borders). D_f is then computed by least-squares fitting the following relation over a range of l_b :

$$\log N_c \propto D_f \log l_b \tag{7}$$

The average R^2 of this fit is 0.997, indicating an excellent goodness of fit.

The Spectral Length Scale (Spectral length) Λ is derived from the field's Fourier transform. Computing this value requires several design choices. First, the scenes are tiltcompensated by subtracting a scene's best-fit plane. Next, one would normally apply a radially symmetric window function to account for the scenes' aperiodicity. However, we find that the application of such a function occludes so much spatial information that our scenes are ordered much less coherently. Hence, we refrain from applying window functions. Next, we Fourier transform the scenes and construct their 1D PSD S(k) by averaging the transform's power signals over shells of radial wavenumber k. The validity of this approach rests on the assumption that the satellite scenes are spatially isotropic, which they are often not. Yet, we find that on a scale from 0-1 (0 representing a 2D PSD where the power is equally distributed over the azimuthal direction and 1 representing the case where all power is concentrated in a single direction), the average anisotropy of all scenes is 0.104. We judge that this justifies the use of the 1D PSD. Finally, Λ is computed from the distribution's first moment, as suggested in Jonker, Duynkerke, and Cuijpers (1999):

$$\Lambda^{-a} = \frac{\int_0^{k_{Ny}} k^a S(k) dk}{\int_0^{k_{Ny}} S(k) dk}; \quad a \neq 0$$
(8)

Where k_{Ny} is the Nyquist wavenumber and we choose to set a = 1.

We compute Wavelet-based Organisation Indices (WOIs) following Brune, Kapp, and Friederichs (2018). These metrics are based on the domain-averaged, squared coefficients of the 2D stationary wavelet transform (SWT) of each scene's cloud water path (CWP) field, E_{CWP} . We use the Haar wavelet as our basis. E_{CWP} contains a scale decomposition

X - 8 JANSSENS ET AL.: CLOUD PATTERNS HAVE 4 INTERPRETABLE DIMENSIONS

over three (horizontal, vertical, diagonal) directions, with each scale representing a power of 2 that exactly fits the 512 pixel field. Using E_{CWP} , we derive the metrics proposed by (Brune et al., 2018):

$$WOI_1 = \frac{E_{CWP}^l}{\overline{E_{CWP}}} \tag{9}$$

$$WOI_2 = \frac{E_{CWP}}{N_c} \tag{10}$$

$$WOI_3 = \frac{1}{3} \sqrt{\sum_d \left(\frac{E_{CWP_d}^l - \overline{E_{CWP}^l}}{\overline{E_{CWP}^l}}\right)^2 + \left(\frac{E_{CWP_d}^s - \overline{E_{CWP}^s}}{\overline{E_{CWP}^s}}\right)^2}$$
(11)

Where \cdot^{l} and \cdot^{s} indicate total energy contained in the large scales (resolution $2^{1}-2^{5}$) and small scales (resolution $2^{6}-2^{9}$) respectively, $\bar{\cdot}$ indicates averaging over all three directions and N_{c} is the number of cloudy pixels in a scene. These metrics measure the fraction of cloud water contained in the scene's large scales (WOI_{1}), the average cloud water in cloudy pixels (WOI_{2}) and the anisotropy in the spectrum's three directions (WOI_{3}). Since WOI_{1} and WOI_{2} are almost exact mirrors of R (eq. 1) and cloud water variance in cloudy pixels respectively, respectively, we choose to only include WOI_{3} in our analysis.

Our simple *Clear Sky* metric extracts the scene's largest rectangular area spanned by the horizontal and vertical lines drawn through any cloud-free pixel whose ends are the first cloudy pixel encountered along those lines. This rectangle is normalised by the domain size, to arrive at a fraction that represents the largest, contiguous, clear sky area.

References

Benner, T. C., & Curry, J. A. (1998). Characteristics of small tropical cumulus clouds and their impact on the environment. *Journal of Geophysical Research: Atmospheres*,

- Bretherton, C., & Blossey, P. (2017). Understanding mesoscale aggregation of shallow cumulus convection using large-eddy simulation. Journal of Advances in Modeling Earth Systems, 9(8), 2798–2821.
- Brueck, M., Hohenegger, C., & Stevens, B. (2020). Mesoscale marine tropical precipitation varies independently from the spatial arrangement of its convective cells. *Quarterly Journal of the Royal Meteorological Society*, 146(728), 1391–1402.
- Brune, S., Kapp, F., & Friederichs, P. (2018). A wavelet-based analysis of convective organization in ICON large-eddy simulations. *Quarterly Journal of the Royal Mete*orological Society, 144 (717), 2812–2829.
- Denby, L. (2020). Discovering the importance of mesoscale cloud organization through unsupervised classification. *Geophysical Research Letters*, 47(1), e2019GL085190.
- Erichson, N. B., Zheng, P., Manohar, K., Brunton, S. L., Kutz, J. N., & Aravkin, A. Y. (2020). Sparse principal component analysis via variable projection. *SIAM Journal* on Applied Mathematics, 80(2), 977–1002.
- Glassmeier, F., & Feingold, G. (2017). Network approach to patterns in stratocumulus clouds. Proceedings of the National Academy of Sciences, 114(40), 10578–10583.
- Jonker, H. J., Duynkerke, P. G., & Cuijpers, J. W. (1999). Mesoscale fluctuations in scalars generated by boundary layer convection. *Journal of the atmospheric sciences*, 56(5), 801–808.
- Neggers, R., Griewank, P., & Heus, T. (2019). Power-law scaling in the internal variability of cumulus cloud size distributions due to subsampling and spatial organization.

- X 10 JANSSENS ET AL.: CLOUD PATTERNS HAVE 4 INTERPRETABLE DIMENSIONS Journal of the Atmospheric Sciences, 76(6), 1489–1503.
- Rasp, S., Selz, T., & Craig, G. C. (2018). Variability and clustering of midlatitude summertime convection: Testing the Craig and Cohen theory in a convection-permitting ensemble with stochastic boundary layer perturbations. *Journal of the Atmospheric Sciences*, 75(2), 691–706.
- Tobin, I., Bony, S., & Roca, R. (2012). Observational evidence for relationships between the degree of aggregation of deep convection, water vapor, surface fluxes, and radiation. *Journal of Climate*, 25(20), 6885–6904.
- van Laar, T. W. (2019). Spatial patterns in shallow cumulus cloud populations over a heterogeneous surface (Doctoral dissertation, University of Cologne). Retrieved from http://kups.ub.uni-koeln.de/id/eprint/10221
- Weger, R., Lee, J., Zhu, T., & Welch, R. (1992). Clustering, randomness and regularity in cloud fields: 1. theoretical considerations. *Journal of Geophysical Research: Atmospheres*, 97(D18), 20519–20536.
- White, B., Buchanan, A., Birch, C., Stier, P., & Pearson, K. (2018). Quantifying the effects of horizontal grid length and parameterized convection on the degree of convective organization using a metric of the potential for convective interaction. *Journal of the Atmospheric Sciences*, 75(2), 425–450.
- Zou, H., Hastie, T., & Tibshirani, R. (2006). Sparse principal component analysis. Journal of computational and graphical statistics, 15(2), 265–286.

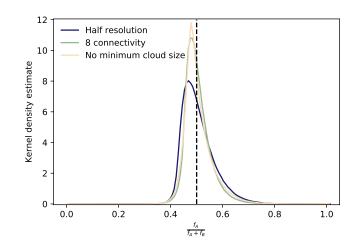


Figure S1. Gaussian kernel density estimates of the ratio $D = \frac{f_A}{f_A - f_B}$, which is constructed from high-dimensional kernel density estimates of the reference metric distribution used in the main text (f_A) and three separately perturbed metric distributions (f_B). An identical distribution to the original would yield a Dirac pulse centred at 0.5 (dashed line); deviations from this line quantify the contrast between the original and perturbed distributions. Sensitivities are quantified with respect to i) scenes that are downsampled to half the original resolution (most sensitive), ii) object segmentation based on 8-connectivity rather than 4-connectivity and iii) not including a lower bound to the minimum cloud size that is considered an object (least sensitive). All perturbed distributions are narrow and have an expected value around 0.5, indicating the robustness of the distribution presented in the main text. Furthermore, the visual relation between metrics is largely unaffected (not shown).

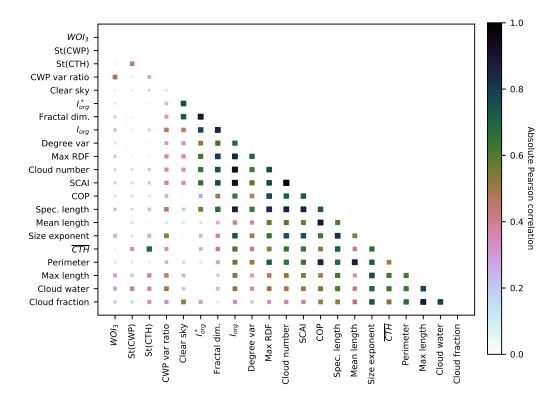
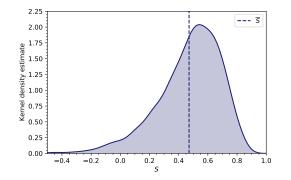


Figure S2. Standardised metric correlation matrix, with squares sized and coloured according to absolute Pearson correlation between a metric pair. Many metrics closely correlate, indicating that their cumulative information can be captured by a smaller number of effective indicators. Several closely correlating metrics follow well-known relationships, e.g. perimeter and mean length (any combination yields approximately constant fractal dim.), or cloud number and numerous aggregation metrics (this relation is similar for deep convective organisation (Brueck et al., 2020)). Others follow rather trivial ones, e.g. max length and cloud fraction, or the Spectral Length and size exponent. Several strong correlations are at first sight not trivial. For instance, I_{org} (both versions) and Fractal dim. are highly similar (up to a factor -1). Hence, highly concentrated shallow cloud clusters in rather empty scenes (high I_{org}^*) tend towards "lines" (low fractal dimension, approaching 1 from above); $I_{org}^* = 0.5$ and fractal dim.=2 both indicate random scattering of points. Finally, while some effort has been invested in contrasting and improving aggregation/clustering measures (e.g. SCAI, I_{org} and max RDF (van Laar, 2019)), these are extremely similar. Instead, shifting focus to metrics that are comparatively *uncorrelated* might September 26, 2020, 12:36pm be more more fruitful to further develop our understanding of shallow cloud field organisation.



Gaussian kernel density estimate of $S = 1 - \frac{\|x_{a_i} - x_{n_i}\|_2}{\|x_{a_i} - x_{r_i}\|_2}$ compiled from $0 \leq 1$ Figure S3. i < 3951 scenes, where x_{a_i} is the vector of the metrics for an "anchor scene", x_{n_i} are the metrics of a "neighbour scene" that overlaps with half the area of the anchor scene and x_{r_i} are the average metrics of 100 randomly sampled scenes. S measures how much the metrics minimise the Euclidian distance between an anchor and its half-overlapping scene, relative to the average Euclidian distance to a randomly sampled scene. If S = 0, the metrics estimate that a half-overlapping scene is equally similarly organised as a randomly sampled scene; if S = 1, the anchor and half-overlapping scene are estimated to be identically organised. Since half-overlapping scenes share numerous spatial features, they should usually be more similarly organised than random scenes (S > 0) - a feature we expect the metrics to capture. As 96% of the distribution exceeds S = 0, this inspires confidence in this ability. The dashed line indicates the mean, $\overline{S} = 0.47$. While this lies significantly below 1, we expect the desired upper bound of S to also lie below 1, since half-overlapping scenes are (by visual inspection) rarely *identically* organised. Estimating this bound requires knowing how far a typical pattern extends beyond a scene's boundaries; this demands a better characterisation of the relation between the measurement scale ("scene") and the true scale of a pattern. However, even without an explicit upper bound on S < 1, this distribution shows that our metrics on average come closer to that bound than to being random. Proficiency of a cloud field description can also be assessed by comparing S across approaches. A version of S already served as cost function for a machinelearned pattern description (Denby, 2020). One could also compile statistics on how similar humans find half-overlapping scenes compared to random scene pairs. Comparing both resulting S to our metrics could more objectively assess which approach to pattern description (human, metrics or machine) is best.

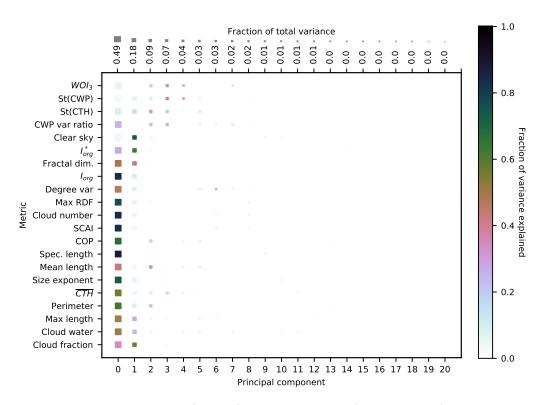


Figure S4. Fraction of variance (colour) in each metric (vertical axis) explained by each PC (horizontal axis). Sizes of squares are scaled by the total dataset's explained variance fraction in each PC (top horizontal axis). 17/21 metrics have more than 70% of their variance captured by the first two PCs; the remaining 4/21 metrics reach this threshold after four PCs.



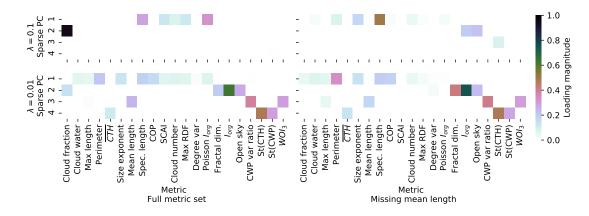


Figure S5. Sensitivity of Sparse Principal Component Analysis (SPCA, Zou et al., 2006). SPCA encourages sparsity in the weighting of metrics that form each of the four main, approximate PCs ("loadings") by casting the PCA as a regression problem, whose cost function contains at least i) a least squares error term of the PCA fit and ii) a penalty (in the L_0 or L_1 norm) on the magnitude of the regression coefficients (the loadings). This penalty is weighted by a regularisation parameter λ . We solve the resulting non-convex optimisation problem using the approach developed by Erichson et al. (2020) and refer to that paper for further details. This figure shows the optimal sparsity structure in the loadings identified by SPCA under four combinations of two free parameters: The magnitude of the sparsity penalty λ (top row vs bottom row) and the omission of a single, seemingly redundant, metric (SCAI, left column vs right column). Unfortunately, the optimal sparsity structure i) is rather sensitive and ii) reacts relatively unpredictably to changes in these free parameters. This is true also when other metrics are excluded, when a different sparsity-inducing algorithm is used or when the sparsity penalty is in the L_0 norm, rather than the L_1 norm as displayed here. These considerations curb SPCA's utility for metric selection and prevent us from recommending its use.

X - 16 JANSSENS ET AL.: CLOUD PATTERNS HAVE 4 INTERPRETABLE DIMENSIONS

Table S1. Metrics quantified for initial analysis. Selection for paper is guaranteed by meeting either criteria 1 or 2, and separately meeting criterion 3, as presented in section 2.2. This excludes the lower portion of the table. Two metrics in the table's middle section meet the criteria, but are still excluded: WOI_1 , WOI_2 (see Text S1). Metrics annotated with (*) are not included in

the coded library.

Metric	Criterion 1	Criterion 2	Criterion 3
	Unique	Recurrent/recent	Interpretable
Cloud fraction	No	Yes	Yes
Cloud water	Yes	Yes	Yes
Max length	No	Yes	Yes
Perimeter	No	Yes	Yes
\overline{CTH}	Yes	Yes	Yes
Size exponent	Yes	Yes	Yes
Mean length	No	Yes	Yes
Spectral length scale	No	Yes	Yes
COP	No	Yes	Yes
SCAI	No	Yes	Yes
Cloud number	No	Yes	Yes
Max RDF	No	Yes	Yes
Degree var.	Yes	Yes	Yes
I_{org}	No	Yes	Yes
Fractal dimension	Yes	Yes	Yes
I_{org}^*	Yes	No	Yes
Open sky	Yes	No	Yes
CWP var ratio	Yes	Yes	Yes
St(CTH)	Yes	Yes	Yes
St(CWP)	Yes	Yes	Yes
WOI_3	Yes	Yes	Yes
WOI_1	No	Yes	Yes
WOI_2	No	Yes	Yes
Multifractality index (*)	Yes	Yes	No
Multifractal intermittency (*)	Yes	Yes	No
Object eccentricity	No	No	Yes
Covariance-based orientation	No	No	Yes
Raw moment-based orientation	No	No	Yes
b_{org} in small clouds (Neggers et al., 2019)	Yes	Yes	No
Skewness/kurtosis of CTH, CWP	Yes	Yes	No
Geometric mean nearest neighbour distance	No	No	Yes
Variance of CTH, CWP in largest cloud	No	No	Yes
1D PSD slope	No	No	Yes
Variance in azimuthal PSD	No	No	Yes
Aboav-Wearie fit (Glassmeier & Feingold, 2017)	Yes	Yes	No