Extreme events driving year-to-year differences in gross primary productivity across the US

Alexander Turner¹, Philipp Köhler², Troy Magney³, Christian Frankenberg², Inez Fung⁴, and Ronald Cohen⁴

¹University of Washington ²California Institute of Technology ³UC Davis ⁴University of California

November 23, 2022

Abstract

Solar-Induced chlorophyll Fluorescence (SIF) has previously been shown to strongly correlate with gross primary productivity (GPP), however this relationship has not yet been quantified for the recently launched TROPOspheric Monitoring Instrument (TROPOMI). Here we use a Gaussian mixture model to develop a parsimonious relationship between SIF from TROPOMI and GPP from flux towers across the conterminous United States (CONUS). The mixture model indicates the SIF-GPP relationship can be characterized by a linear model with two terms. We then estimate GPP across CONUS at 500-m spatial resolution over a 16-day moving window. We find that CONUS GPP varies by less than 4% between 2018 and 2019. However, we observe four extreme precipitation events that induce regional GPP anomalies: drought in west Texas, flooding in the midwestern US, drought in South Dakota, and drought in California. Taken together, these events account for 28% of the year-to-year GPP differences across CONUS.

Extreme events driving year-to-year differences in gross 1 primary productivity across the US 2

Alexander J. Turner^{1,2,3*}, Philipp Köhler⁴, Troy S. Magney⁵, Christian Frankenberg^{4,3}, Inez Fung¹, and Ronald C. Cohen^{1,2}

 ¹Department of Earth and Planetary Sciences, University of California, Berkeley, CA, 94720, USA.
 ²College of Chemistry, University of California, Berkeley, CA, 94720, USA.
 ³Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, 91109, USA.
 ⁴Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA, 91226, USA. ⁵Department of Plant Sciences, University of California, Davis, CA, 95616, USA.

Key Points:

3

5 6 8

10

11

12	٠	We estimate 500-m GPP from TROPOMI SIF over a moving 16-day window for
13		all of the conterminous United States.
14	•	There are two distinct relationships between TROMPOMI SIF and AmeriFlux GPP
15		across ecosystems.
16	•	Extreme precipitation events drive four regional GPP anomalies that account for
17		28% of the year-to-year differences across the US.

28% of the year-to-year differences across the US.

^{*}now at: Department of Atmospheric Sciences, University of Washington, Seattle, WA, 98195, USA

Corresponding author: Alexander J. Turner, turneraj@uw.edu

18 Abstract

Solar-Induced chlorophyll Fluorescence (SIF) has previously been shown to strongly cor-19 relate with gross primary productivity (GPP), however this relationship has not yet been 20 quantified for the recently launched TROPOspheric Monitoring Instrument (TROPOMI). 21 Here we use a Gaussian mixture model to develop a parsimonious relationship between 22 SIF from TROPOMI and GPP from flux towers across the conterminous United States 23 (CONUS). The mixture model indicates the SIF-GPP relationship can be characterized 24 by a linear model with two terms. We then estimate GPP across CONUS at 500-m spa-25 tial resolution over a 16-day moving window. We find that CONUS GPP varies by less 26 than 4% between 2018 and 2019. However, we observe four extreme precipitation events 27 that induce regional GPP anomalies: drought in west Texas, flooding in the midwest-28 ern US, drought in South Dakota, and drought in California. Taken together, these events 20

account for 28% of the year-to-year GPP differences across CONUS.

³¹ Plain Language Summary

Gross primary productivity is the total amount of CO_2 taken up by plants during 32 photosynthesis and represents one of the main drivers of variability in atmospheric CO_2 . 33 Plants emit a small amount of light during the process of photosynthesis, this is known 34 as "solar-induced chlorophyll fluorescence" (SIF). We can measure this SIF signal from 35 space and use it to study the biosphere. Here we build a high-resolution estimate of gross 36 primary productivity over the United States using satellite measurements of SIF from 37 2018 through 2019. We find the major drivers of variability in gross primary productiv-38 ity across the US were drought in west Texas, flooding in the midwestern US, drought 39 in South Dakota, and drought in California. 40

41 **1** Introduction

Terrestrial gross primary productivity (GPP) is the total amount of carbon diox-42 ide (CO_2) assimilated by plants through photosynthesis and represents one of the main 43 drivers of interannual variability in the global carbon cycle (Le Quéré et al., 2018). As 44 such, quantifying the spatiotemporal patterns of terrestrial GPP is critical to understand-45 ing how the carbon cycle will both respond to and influence climate. Work over the past 46 decade has shown satellite measurements of solar-induced chlorophyll fluorescence (SIF) 47 to correlate strongly with tower-based estimates of GPP (e.g., Frankenberg et al., 2011; 48 X. Yang et al., 2015; Sun et al., 2017; Turner et al., 2020; Wang et al., 2020) and are of-49 ten used as a remote-sensing proxy for GPP. 50

This relationship between SIF and GPP is typically expressed through a pair of light use efficiency models (Monteith, 1972) that relate GPP and SIF to the absorbed photosynthetically active radiation (APAR):

$$GPP = APAR \times \Phi_{CO_2} \tag{1}$$

$$SIF = APAR \times \beta \Phi_F \tag{2}$$

where Φ_{CO_2} is the light use efficiency of CO_2 assimilation, Φ_F is the fluorescence yield, and β is the probability of fluoresced photons escaping the canopy. Solving for APAR and substituting, we can rewrite GPP as:

$$GPP = \frac{\Phi_{CO_2}}{\beta \Phi_F} SIF.$$
(3)

The derivation follows from Lee et al. (2013), Guanter et al. (2014), Sun et al. (2017), and others.

This seemingly straight forward relationship between SIF and GPP has been widely used to infer GPP from measurements of SIF (e.g., Frankenberg et al., 2011; Parazoo et al., 2014; X. Yang et al., 2015; H. Yang et al., 2017; Sun et al., 2017, 2018; Magney et al., 2019; Turner et al., 2020) with some work showing that SIF captures more variability in GPP than APAR alone (e.g., X. Yang et al., 2015; H. Yang et al., 2017; Magney et al., 2019). However, there is much complexity encapsulated in the first term of Eq. 3 ($\Phi_{CO_2}/\beta\Phi_F$). There is an ongoing debate about what *exactly* SIF is telling us about GPP (e.g., Dechant et al., 2020; Marrs et al., 2020) and the spatio-temporal scales at which SIF and GPP correlate well.

Here we focus on the ecosystem-scale relationship between SIF and GPP, as that 68 is the relevant observable scale from space-borne instruments. We begin by character-69 izing the relationship between SIF from TROPOMI and GPP from flux towers. Follow-70 ing this, we use this ecosystem-scale relationship to infer GPP at a spatial resolution of 71 500-m using TROPOMI SIF measurements and identify drivers of interannual variabil-72 ity in GPP. Previous work has identified effects such as seasonal redistribution (Butterfield 73 et al., 2020), drought (e.g., Sun et al., 2015), and flooding (Yin et al., 2020) as impor-74 tant drivers of interannual variability in GPP. 75

⁷⁶ 2 Identifying distinct relationships between SIF and GPP

We build on our previous work (Turner et al., 2020) downscaling measurements of 77 SIF to 500-m spatial resolution. Briefly, the TROPOspheric Monitoring Instrument (TROPOMI; 78 Veefkind et al., 2012) is a nadir-viewing imaging spectrometer. TROPOMI has a 2,600 79 km swath with a nadir spatial resolution of 5.6 km along track and 3.5 km across track. 80 Köhler et al. (2018) presented the first retrievals of SIF from TROPOMI. As in Turner 81 et al. (2020), we apply a *post hoc* bias correction to ensure positivity of monthly aver-82 age values as systematically negative SIF values are non-physical. We downscale indi-83 vidual TROPOMI scenes using the near-infrared reflectance of vegetation index (NIR_v) 84 that was proposed by Badgley et al. (2017, 2019). We use the MCD43A4.006 (v06) MODIS 85 NBAR reflectances (Schaaf et al., 2002) to compute NIR_v. Two notable differences from 86 Turner et al. (2020) are: 1) the analysis is extended to cover all of CONUS and 2) we 87 now use a 16-day moving window, thus including a full orbit cycle in each averaging win-88 dow to minimize effects due to viewing-illumination geometry and noise. 89

The extension to CONUS facilitates comparison of TROPOMI SIF retrievals to flux 90 tower data over a more representative set of ecosystems and robustly infer the SIF-GPP 91 relationship. Specifically, there are 82 AmeriFlux sites (D. Baldocchi et al., 2001) within 92 CONUS that reported data in 2018, 2019, or 2020 whereas Turner et al. (2020) only in-93 cluded 11 sites and did not have data from forests. Figure 1 shows the location of these 94 82 AmeriFlux sites overlaid on the dominant landcover. These eddy covariance sites pro-95 vide a direct measure of net ecosystem exchange (NEE; CO₂ fluxes) (D. D. Baldocchi 96 et al., 1988). We use GPP that has been partitioned by the group operating the site. If 97 GPP is not provided we compute it using nighttime measurements of NEE as a proxy 98 for ecosystem respiration (Reichstein et al., 2005). The AmeriFlux sites used here cover 99 10 ecosystems as defined by the International Geosphere-Biosphere Programme: ever-100 green needleleaf forest, deciduous broadleaf forest, mixed forest, grassland, cropland, wet-101 land, woody savanna, savanna, open shrubland, and closed shrubland. 102

We characterize the relationship between TROPOMI SIF and AmeriFlux GPP by 103 plotting downscaled instantaneous SIF observations against the nearest AmeriFlux GPP 104 data in time. Specifically, the 6 steps we take here are: 1) apply the post hoc bias cor-105 rection to the TROPOMI SIF data, 2) find all TROPOMI scenes that cover an Amer-106 iFlux site, 3) downscale TROPOMI scenes to 500-m using MODIS NIR, 4) construct 107 a timeseries of SIF observations from the 500-m grid cell that contains the AmeriFlux 108 site, 5) construct a timeseries of AmeriFlux GPP data that are coincident in time with 109 the TROPOMI overpass, and 6) regress SIF on GPP with a bisquare regression. The bisquare 110 regression was chosen due to robustness against outliers. Additionally, we force the re-111



Figure 1. Dominant landcover over conterminous United States (CONUS). Colors show the dominant landcover over CONUS. Classification is based on the 2019 USDA CropScape database (USDA, 2018). Forests are shown in green croplands in yellow, and wetlands in blue. Location of 82 AmeriFlux sites used in this study are shown as yellow stars.

gression through the origin based on the physical constraint that GPP should be zero if SIF is zero. We observe a linear relationship between SIF and GPP when plotted against all ecosystems (Supplemental Figure S1) and when separated by ecosystem (Supplemental Figure S2). Notable exceptions are closed shrubland, open shrubland, and savanna ecosystems where SIF explains less than 10% of the variability in GPP for AmeriFlux sites in those ecosystems due, in part, to a low signal-to-noise ratio.

Many of the ecosystems exhibit a similar linear relationship between SIF and GPP, 118 which begs the question: "what ecosystems have a distinct SIF-GPP relationship?" To 119 address this, we bootstrap the bisquare regression for each ecosystem 2000 times. The 120 slopes from this bootstrap can be seen in Figure 2. The range of slopes vary from 13 to 121 $18 \ (\mu \text{mol}\,\text{m}^{-2}\,\text{s}^{-1}) \ / \ (\text{mW}\,\text{m}^{-2}\,\text{sr}^{-1}\,\text{s}^{-1})$ with grasslands at the low end and evergreen 122 needleleaf forests at the high end. We then use a two component Gaussian mixture model (see, 123 for example, Bishop, 2007) to identify clusters of ecosystems with a similar SIF-GPP re-124 lationship. The implementation of our Gaussian mixture model is adapted from Turner 125 and Jacob (2015). Parameters of the mixture model are obtained via an iterative expectation-126 maximization algorithm. A drawback of these mixture models is they often find local 127 minima. To address this, we repeat the fitting of the mixture model with multiple ini-128 tializations and use simulated annealing to search for a global minimum. We tested a 129 range of mixture model sizes and found a mixture of two Gaussians to be the most ro-130 bust. The resulting mixture model is overlaid on the histogram in Figure 2. 131

¹³² We observe a clustering of ecosystems with SIF-GPP relationships around 16.3 (μ mol m⁻² s⁻¹) / (mW m⁻² sr⁻¹ ¹³³ This grouping is the dominant weighting term for wetlands, evergreen needleleaf forest, ¹³⁴ deciduous broadleaf forest, mixed forest, cropland, and woody savanna. We refer to this ¹³⁵ cluster as the "Dominant Cluster" and assume that ecosystems not specifically mentioned ¹³⁶ elsewhere will have a response that is similar to this primary cluster. The other com-¹³⁷ ponent of the mixture model corresponds to grasslands. Table 1 lists the SIF-GPP re-¹³⁸ lationships for these two clusters. These relationships can be used to reconstruct GPP

Cluster	SIF-GPP relationship ^{a} (s_i)
Dominant Cluster	16.3 ± 0.4
Grassland	13.7 ± 0.1

 Table 1. SIF-GPP relationships for different groupings.

^{*a*}All SIF-GPP relationships have units of $(\mu \text{mol m}^{-2} s^{-1}) / (\text{mW m}^{-2} \text{sr}^{-1} \text{nm}^{-1})$. Uncertainty is the diagonal of the covariance matrix for the mixture model.

from TROPOMI SIF as: GPP = SIF \times ($\sum_i f_i s_i$) where s_i is the SIF-GPP relationship in Table 1 for the *i*th cluster and f_i is the fraction of a grid cell represented by that cluster.



Figure 2. Identifying distinct SIF-GPP relationships across ecosystems. Histogram shows the distribution of slopes that map SIF to GPP using a bisquare regression and a 2000 member bootstrap. Colors denote the different ecosystems and triangles at the bottom show the mean for that ecosystem. Gray distributions are from a two-member Gaussian Mixture Model and the stars indicate the mean for that component.

TROPOMI is in low earth orbit and only observes a snapshot in time. The equatorial overpass time at nadir is 13:30 local time. By assuming that GPP scales linearly with PAR (i.e., Eq. 1) we can compute a correction factor to estimate daily integrated GPP. More formally, we scale the instantaneous SIF by the ratio of the integral of the cosine of the solar zenith angle (SZA) over the day to cos (SZA) from the TROPOMI overpass time. Putting everything together, we estimate daily GPP from TROPOMI SIF ¹⁴⁸ observations as:

$$GPP(x, y, t) = SIF(x, y, t) \cdot \gamma \sum_{i} s_i f_i(x, y) \cdot \frac{\int_{t_0}^{t_f} \cos\left[SZA(x, y, t)\right] dt}{\cos\left[SZA(x, y, t_s)\right]}$$
(4)

where SIF(x, y, t) is the 500-m downscaled SIF using a 16-day moving window, γ is a unit conversion from μ mol to gC, s_i is the SIF-GPP relationship inferred from comparison with AmeriFlux GPP (see Table 1), $f_i(x, y)$ is the fraction of the grid cell represented by the *i*th cluster, SZA is the local solar zenith angle, t_0 is sunrise, t_f is sunset, and t_s is the hour corresponding to the TROPOMI overpass time.

¹⁵⁴ **3** Drivers of interannual variations in US gross primary productivity

Figure 3 shows annual mean GPP across CONUS inferred from TROPOMI SIF 155 measurements using Eq. 4. A number of prominent features are visible such as the Cen-156 tral Valley of California, the Snake River Valley in Idaho, and the Adirondack Moun-157 tains in upstate New York. California's Central Valley and Idaho's Snake River Valley 158 are both major agricultural regions in the western US (e.g., the Central Valley of Cal-159 iforia accounts for more than 15% of irrigated land in the US). The Adirondack Moun-160 tains are a roughly circular dome that rise above the surrounding lowlands, resulting in 161 a shorter growing season and lower annual mean GPP. This shortened growing season 162 can be seen in an animation of GPP over CONUS (Supplemental Movie S1). 163

We observe substantial GPP across the eastern US (delineated here by $98^{\circ}W$) with 164 annual mean values generally in excess of 5 ${\rm gC/m^2/day}$. This region accounts for less 165 than half of the land but more than 70% of the annual mean GPP. This delineation in 166 GPP roughly coincides with the location of drylands in CONUS that are more sensitive 167 to changes in precipitation; drylands are also projected to expand in future climate (Yao 168 et al., 2020). Most of the large vear-to-vear differences occur in these western US dry-169 lands (see Fig. 3c), a notable exception being a negative GPP anomaly in 2019 relative 170 to 2018 that extended across Illinois, Indiana, and Ohio. Here we highlight four precipitation-171 driven GPP anomalies, which taken together, account for 28% of the interannual GPP 172 variability across the United States: 1) 2018 drought in west Texas, 2) 2019 midwest-173 ern corn belt flooding, 3) 2018 drought in South Dakota, and 4) 2018 drought in Cal-174 ifornia. Figure 4 summarizes the interannual precipitation differences that we hypoth-175 esize are responsible for explaining these four GPP anomalies. 176

The largest positive GPP anomaly in 2019 relative to 2018 was observed across west-177 ern Texas. This single event accounted for 11% of the year-to-year difference in GPP across 178 CONUS. From Figure 4a, we observe 50% higher GPP in spring 2019 compared to spring 179 2018. This increase in GPP was driven by a lack of precipitation in spring 2018. The cu-180 mulative precipitation from October 2017 through June 2018 was 50% less than Octo-181 ber 2018 through June 2019 (500 mm vs 1000 mm). The other notable difference between 182 GPP in 2018 and 2019 was a second peak during fall 2018 that was not present in 2019. 183 This second peak coincided with a series of precipitation events beginning in early Septem-184 ber. This tight coupling between GPP and precipitation is expected for dryland systems 185 such as west Texas (e.g., Smith et al., 2019). The seasonal GPP dynamics inferred from 186 TROPOMI SIF are also present in the MODIS vegetation index NIR_v, albeit with slight 187 differences in magnitude, implying convergent responses in SIF and $NIR_{\rm y}$ for this ecosys-188 tem. 189

The second largest anomaly is the reduction in 2019 GPP relative to 2018 across the midwestern corn belt (defined here as Illinois, Indiana, and Ohio) that accounted for 7% of the year-to-year difference in CONUS GPP. We observe a decrease in the maximum GPP between 2019 and 2018 as well as a two week delay in the timing of the maximum. This anomaly was highlighted in recent work from Yin et al. (2020) who attribute the anomaly to flooding in the midwestern US. The flooding delayed planting of crops



Figure 3. Interannual variations in gross primary productivity across CONUS. Map of annual mean GPP for 2018 (panel a) and 2019 (panel b). (Panel c) Map of the difference in annual mean GPP between 2019 and 2018. Red indicates higher GPP in 2019 and red indicates higher GPP in 2018. Inset in bottom left corner shows a timeseries of the average GPP across CONUS for 2018 and 2019.



Figure 4. Major drivers of interannual variability in CONUS GPP. Black line shows the TROPOMI-derived GPP over Texas (a), the midwest corn belt (b), South Dakota (c), and California (d). Blue line shows the cumulative precipitation over the water year as measured by the GPM satellite. Green line is NIR_v from MODIS. Black and Green dotted lines are 2018 GPP and NIR_v superimposed on the 2019 timeseries.

by two weeks and resulted in decreased carbon uptake across the midwestern corn belt
and Mississippi Alluvial Valley, where we also observe a negative anomaly in Figure 3c.
Yin et al. (2020) provide a detailed discussion of these floods and their impacts on crop
productivity.

South Dakota exhibits a dipole with positive anomalies in 2019 in the west and neg-200 ative anomalies in the east, again relative to 2018. The negative anomalies in the east 201 are driven by the flooding events discussed above and in Yin et al. (2020). However, the 202 positive anomaly in western portion of the state is the dominant term. This positive anomaly 203 is driven by a series of summer precipitation events that served to extend the growing 204 season across the western plains. From Figure 4c, we can see three precipitation events 205 throughout the mid-to-late summer that coincide with pauses in senesence: mid-July, 206 early August, and mid-September. As with Texas, this highlights the tight coupling be-207 tween GPP and precipitation for dryland systems. In toto, these precipitation events served 208 to increase statewide GPP in 2019 relative to 2018. 209

The final notable anomaly is California's positive GPP anomaly in 2019. 2018 was 210 a mild drought in California with $\sim 80\%$ of the state being classified as abnormally dry; 211 2019 had 50% more precipitation during the water year than 2018 (Figure 4c). Two con-212 sequences of this drought in 2018 were: a delayed onset of photosynthesis and a mid-summer 213 senescence. The onset of photosynthesis in 2018 coincided with a series of atmospheric 214 rivers that delivered about a third of the total precipitation that year, indicating a wa-215 ter limitation up to that point. In contrast, 2019 had ample precipitation through the 216 winter and we observe both an earlier onset of photosynthesis and an extension of the 217 growing season into the fall. Evergreen forests are the main contributor to the SIF sig-218 219 nal during the summer and fall (Turner et al., 2020) and, as such, will be more sensitive to the accumulated precipitation. The spatial pattern of the differences in August-220 November GPP (Fig. S3) strongly correlate with evergreen forests. 221

In contrast to the anomalies presented earlier, the SIF-derived GPP and MODIS-222 based vegetation index (NIR_v) show divergent seasonal dynamics for California. NIR_v 223 shows small differences between 2018 and 2019 with a strong similarity to the 2019 SIF-224 derived GPP. Vegetation indices estimate *photosynthetic capacity* provided optimal soil 225 moisture, temperature, and PAR are known (Sellers, 1985). As such, this suggests that 226 we observed a down-regulation of photosynthesis from evergreen forests in response to 227 a water limitation during fall 2018, whereas these forests were close to photosynthetic 228 capacity in fall 2019 resulting in a similar seasonality to 2018 and 2019 NIR_{y} . Sims et 229 al. (2014) also report a low sensitivity of MODIS vegetation indices to drought stress in 230 forests. 231

4 Conclusions

We have developed a parsimonious relationship between measurements of SIF from 233 TROPOMI and GPP inferred from flux towers. This relationship allows for estimation 234 of GPP directly from TROPOMI SIF measurements. We combine this SIF-GPP rela-235 tionship with work downscaling TROPOMI data to 500-m spatial resolution to construct 236 estimates of GPP across the conterminous United States in 2018 and 2019. Our estimate of US GPP varies by less than 4% between 2018 and 2019. We do, however, observe large 238 regional anomalies that are driven by extreme precipitation events. Namely, west Texas, 239 South Dakota, and California experienced droughts in 2018 while the midwestern US corn 240 belt states (Illinois, Indiana, and Ohio) experienced flooding in 2019. Taken together, 241 these four events account for 28% of the year-to-year variability in GPP across the con-242 terminous United States. 243

The impact of the west Texas drought, South Dakota drought, and midwestern flooding are observed in other remote-sensing measures of photosynthetic capacity such as NIR_v while the California drought shows a divergent result using SIF; the divergent responses are driven by specific ecosystems such as evergreen forests. Our work suggests that SIF provides a measure of *photosynthetic activity* as opposed to *photosynthetic capacity*, and converge with other remote-sensing measures under non-stressed conditions. Future work investigating the response to extreme events across ecosystems may provide additional insight into these divergent responses in remote-sensing measurements related to photosynthesis.

253 Acknowledgments

We are grateful to the team that has realized the TROPOMI instrument, consisting of 254 the partnership between Airbus Defence and Space, KNMI, SRON, and TNO, commis-255 sioned by NSO and ESA. We acknowledge the following AmeriFlux sites for their data 256 records: US-ALQ, US-ARM, US-Bi1, US-Bi2, US-CF1, US-CF2, US-CF3, US-CF4, US-257 CS1, US-CS2, US-CS3, US-EDN, US-GLE, US-Hn2, US-Hn3, US-Ho1, US-JRn, US-Jo2, 258 US-KS3, US-Los, US-Me2, US-Me6, US-Men, US-Mpj, US-MtB, US-Myb, US-NC2, US-259 NC3, US-NC4, US-Rls, US-Rms, US-Ro4, US-Ro5, US-Ro6, US-Rwf, US-Rws, US-SRG, 260 US-SRM, US-Seg, US-Ses, US-Sne, US-Snf, US-Syv, US-Ton, US-Tw1, US-Tw4, US-Tw5, 261 US-UMd, US-Var, US-Vcm, US-Vcp, US-WCr, US-Whs, US-Wjs, US-Wkg, US-xAB, US-262 xBR, US-xCP, US-xDC, US-xDL, US-xHA, US-xJE, US-xJR, US-xKA, US-xKZ, US-263 xNG, US-xNQ, US-xRM, US-xSE, US-xSL, US-xSP, US-xSR, US-xST, US-xTE, US-xUK, 264 US-xUN, US-xWD, US-xWR, US-xYE. In addition, funding for AmeriFlux data resources 265 was provided by the U.S. Department of Energy's Office of Science. Funding: AJT was 266 supported as a Miller Fellow with the Miller Institute for Basic Research in Science at UC Berkeley. This research was funded by grants from the Koret Foundation and NASA 268 80NSSC19K0945 for support of the computational resources. Part of this research was 269 funded by the NASA Carbon Cycle Science program (grant NNX17AE14G). TROPOMI 270 SIF data generation by PK and CF is funded by the Earth Science U.S. Participating 271 Investigator program (grant NNX15AH95G). This research used the Savio computational 272 cluster resource provided by the Berkeley Research Computing program at the Univer-273 sity of California, Berkeley (supported by the UC Berkeley Chancellor, Vice Chancel-274 lor for Research, and Chief Information Officer). Author contributions: AJT wrote 275 the text with feedback from all authors. PK and CF performed the TROPOMI SIF re-276 trievals. AJT downscaled the SIF data, conducted the AmeriFlux analysis, and drafted 277 the figures. All authors contributed to the discussion and analysis. Competing inter-278 ests: The authors declare no competing interests. Data and materials availability: 279 Daily gridded 500-m TROPOMI SIF and GPP data from February 1, 2018 through June 280 15, 2020 is temporarily available on Google Drive here: "https://bit.ly/2GHEOOq", and 281 will be uploaded to ORNL DAAC at acceptance. 282

283 **References**

- Badgley, G., Anderegg, L. D. L., Berry, J. A., & Field, C. B. (2019). Terrestrial Gross Primary Production: Using NIR_V to Scale from Site to Globe. *Global change biology*. doi: 10.1111/gcb.14729
- Badgley, G., Field, C. B., & Berry, J. A. (2017). Canopy near-infrared reflectance
 and terrestrial photosynthesis. Sci Adv, 3(3), e1602244. doi: 10.1126/sciadv
 .1602244
- Baldocchi, D., Falge, E., Gu, L., Olson, R., Hollinger, D., Running, S., ... Wofsy,
 S. (2001). FLUXNET: A New Tool to Study the Temporal and Spatial Variability of Ecosystem–Scale Carbon Dioxide, Water Vapor, and Energy Flux
 Densities. Bulletin of the American Meteorological Society, 82(11), 2415-2434.
 doi: 10.1175/1520-0477(2001)082(2415:fantts)2.3.co;2
- Baldocchi, D. D., Hicks, B. B., & Meyers, T. P. (1988). Measuring Biosphere Atmosphere Exchanges of Biologically Related Gases with Micrometeorological

297	Methods. <i>Ecology</i> , 69(5), 1331-1340. doi: 10.2307/1941631
298	Bishop, C. M. (2007). Pattern Recognition and Machine Learning (1st ed.).
299	Springer.
300	Butterfield, Z., Buermann, W., & Keppel-Aleks, G. (2020). Satellite observations re-
301	veal seasonal redistribution of northern ecosystem productivity in response to
302	interannual climate variability. Remote Sensing of Environment, 242, 111755.
303	doi: 10.1016/j.rse.2020.111755
304	Dechant, B., Ryu, Y., Badgley, G., Zeng, Y., Berry, J. A., Zhang, Y., Moya, I.
305	(2020). Canopy structure explains the relationship between photosynthesis and
306	sun-induced chlorophyll fluorescence in crops. Remote Sensing of Environment.
307	241, 111733. doi: 10.1016/j.rse.2020.111733
308	Frankenberg, C., Butz, A., & Toon, G. C. (2011). Disentangling chlorophyll fluo-
309	rescence from atmospheric scattering effects in O ₂ A-band spectra of reflected
310	sun-light. Geophysical Research Letters, 38(3). doi: 10.1029/2010gl045896
311	Guanter, L., Zhang, Y., Jung, M., Joiner, J., Voigt, M., Berry, J. A.,, Griffis,
312	T. J. (2014). Global and time-resolved monitoring of crop photosynthe-
313	sis with chlorophyll fluorescence. Proceedings of the National Academy
314	of Sciences of the United States of America, 111(14), E1327-33. doi:
315	10.1073/pnas.1320008111
316	Köhler, P., Frankenberg, C., Magney, T. S., Guanter, L., Joiner, J., & Landgraf, J.
317	(2018). Global Retrievals of Solar-Induced Chlorophyll Fluorescence With
318	TROPOMI: First Results and Intersensor Comparison to OCO-2. <i>Geophysical</i>
319	Research Letters, 45(19), 10.456-10.463, doi: 10.1029/2018gl079031
320	Lee, J. E., Frankenberg, C., van der Tol, C., Berry, J. A., Guanter, L., Boyce, C. K.,
321	Saatchi, S. (2013). Forest productivity and water stress in Amazonia: ob-
322	servations from GOSAT chlorophyll fluorescence. <i>Proc. Biol. Sci.</i> , 280(1761).
323	20130171. doi: 10.1098/rspb.2013.0171
324	Le Quéré, C., Andrew, R. M., Friedlingstein, P., Sitch, S., Hauck, J., Pongratz, J.,
325	Zheng, B. (2018). Global Carbon Budget 2018. Earth System Science
326	Data, 10(4), 2141-2194. doi: 10.5194/essd-10-2141-2018
327	Magney, T. S., Bowling, D. R., Logan, B. A., Grossmann, K., Stutz, J., Blanken,
328	P. D., Frankenberg, C. (2019). Mechanistic evidence for tracking the
329	seasonality of photosynthesis with solar-induced fluorescence. Proceedings of
330	the National Academy of Sciences of the United States of America, 116(24),
331	11640-11645. doi: 10.1073/pnas.1900278116
332	Marrs, J. K., Reblin, J. S., Logan, B. A., Allen, D. W., Reinmann, A. B., Bombard,
333	D. M., Hutyra, L. R. (2020). SolarInduced Fluorescence Does Not Track
334	Photosynthetic Carbon Assimilation Following Induced Stomatal Closure.
335	Geophysical Research Letters, 47(15). doi: 10.1029/2020gl087956
336	Monteith, J. L. (1972). Solar Radiation and Productivity in Tropical Ecosystems.
337	Journal of Applied Ecology, $9(3)$, 747-766.
338	Parazoo, N. C., Bowman, K., Fisher, J. B., Frankenberg, C., Jones, D. B., Cescatti,
339	A., Montagnani, L. (2014). Terrestrial gross primary production inferred
340	from satellite fluorescence and vegetation models. Global change biology,
341	20(10), 3103-21. doi: 10.1111/gcb.12652
342	Reichstein, M., Falge, E., Baldocchi, D., Papale, D., Aubinet, M., Berbigier, P.,
343	Valentini, R. (2005). On the separation of net ecosystem exchange into as-
344	similation and ecosystem respiration: review and improved algorithm. Global
345	change biology, 11(9), 1424-1439. doi: 10.1111/j.1365-2486.2005.001002.x
346	Schaaf, C. B., Gao, F., Strahler, A. H., Lucht, W., Li, X., Tsang, T., Roy,
347	D. (2002). First operational BRDF, albedo nadir reflectance products
348	from MODIS. Remote Sensing of Environment, 83(1-2), 135-148. doi:
349	10.1016/s0034-4257(02)00091-3
350	Sellers, P. J. (1985). Canopy reflectance, photosynthesis and transpira-
351	tion. International Journal of Remote Sensing, $6(8)$, 1335-1372. doi:

352	10.1080/01431168508948283
353	Sims, D. A., Brzostek, E. R., Rahman, A. F., Dragoni, D., & Phillips, R. P. (2014).
354	An improved approach for remotely sensing water stress impacts on forest C
355	uptake. Global change biology, 20(9), 2856-2866. doi: 10.1111/gcb.12537
356	Smith, W. K., Dannenberg, M. P., Yan, D., Herrmann, S., Barnes, M. L., Barron-
357	Gafford, G. A., Yang, J. (2019). Remote sensing of dryland ecosystem
358	structure and function: Progress, challenges, and opportunities. Remote Sens-
359	ing of Environment, 233, 111401. doi: 10.1016/j.rse.2019.111401
360	Sun, Y., Frankenberg, C., Jung, M., Joiner, J., Guanter, L., Köhler, P., & Magney,
361	T. (2018). Overview of Solar-Induced chlorophyll Fluorescence (SIF) from
362	the Orbiting Carbon Observatory-2: Retrieval, cross-mission comparison, and
363	global monitoring for GPP. Remote Sensing of Environment, 209, 808-823.
364	doi: 10.1016/j.rse.2018.02.016
365	Sun, Y., Frankenberg, C., Wood, J. D., Schimel, D. S., Jung, M., Guanter, L.,
366	Yuen, K. (2017). OCO-2 advances photosynthesis observation from
367	space via solar-induced chlorophyll fluorescence. Science, 358(6360). doi:
368	10.1126/science.aam5747
369	Sun, Y., Fu, R., Dickinson, R., Joiner, J., Frankenberg, C., Gu, L.,, Fernando,
370	N. (2015). Drought onset mechanisms revealed by satellite solar-induced
371	chlorophyll fluorescence: Insights from two contrasting extreme events.
372	Journal of Geophysical Research: Biogeosciences, 120(11), 2427-2440. doi:
373	10.1002/2015jg003150
374	Turner, A. J., & Jacob, D. J. (2015). Balancing aggregation and smoothing errors
375	in inverse models. Atmos Chem Phys. 15(12), 7039-7048. doi: 10.5194/acp-15
376	-7039-2015
377	Turner, A. J., K"hler, P., Magney, T. S., Frankenberg, C., Fung, I., & Cohen,
378	R. C. (2020). A double peak in the seasonality of California's photo-
379	synthesis as observed from space. $Biogeosciences, 17(2), 405-422.$ doi:
380	10.5194/bg-17-405-2020
381	USDA. (2018). National Agricultural Statistics Service Cropland Data Layer:
382	Published crop-specific data layer. Retrieved 15/3/2019, from https://
383	nassgeodata.gmu.edu/CropScape/
384	Veefkind, J. P., Aben, I., McMullan, K., Förster, H., de Vries, J., Otter, G.,
385	Levelt, P. F. (2012). TROPOMI on the ESA Sentinel-5 Precursor: A
386	GMES mission for global observations of the atmospheric composition for
387	climate, air quality and ozone layer applications. <i>Proc SPIE</i> , 120, 70-83. doi:
388	10.1016/j.rse.2011.09.027
389	Wang, X., Dannenberg, M. P., Yan, D., Jones, M. O., Kimball, J. S., Moore,
390	D. J. P., Smith, W. K. (2020). Globally Consistent Patterns of Asynchrony
391	in Vegetation Phenology Derived From Optical, Microwave, and Fluorescence
392	Satellite Data. Journal of Geophysical Research: Biogeosciences, 125(7). doi:
393	10.1029/2020jg 005732
394	Yang, H., Yang, X., Zhang, Y., Heskel, M. A., Lu, X., Munger, J. W., Tang, J.
395	(2017). Chlorophyll fluorescence tracks seasonal variations of photosynthe-
396	sis from leaf to canopy in a temperate forest. Global change biology, $23(7)$,
397	2874-2886. doi: 10.1111/gcb.13590
398	Yang, X., Tang, J., Mustard, J. F., Lee, JE., Rossini, M., Joiner, J., Richard-
399	son, A. D. (2015). Solar-induced chlorophyll fluorescence that correlates
400	with canopy photosynthesis on diurnal and seasonal scales in a temper-
401	ate deciduous forest. Geophysical Research Letters, $42(8)$, 2977-2987. doi:
402	10.1002/2015gl 063201
403	Yao, J., Liu, H., Huang, J., Gao, Z., Wang, G., Li, D., Chen, X. (2020). Accel-
404	erated dryland expansion regulates future variability in dryland gross primary
405	production. Nature communications, $11(1)$. doi: $10.1038/s41467-020-15515-2$
406	Yin, Y., Byrne, B., Liu, J., Wennberg, P. O., Davis, K. J., Magney, T., Franken-

407	berg, C.	(2020).	${\rm Cropland}$	Carbon	Uptake	Delayed	and Reduced by 2019
	A (* 1) T 1	1 4 0 17		1(1) 1		000 10010	000140

408 Midwest Floods. AGU Advances, 1(1). doi: 10.1029/2019av000140