3D magnetotelluric inversion reveals the superposition of tectonic systems in the northern Songliao Basin

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Abstract

The creation and evolution of the Songliao Basin is closely related to the closure of the Paleo-Asian Ocean and the Mongolia-Okhotsk Ocean, and the subduction of the Paleo-Pacific Ocean. In an attempt to demonstrate the various attributes of the lithospheric structure under structural superimposition and transformation of the Songliao Basin, this work used full impedance 3D inversion to obtain a 3D electrical structural model of the northern Songliao Basin for the first time. The results showed there are NE-trending high-resistance anomalies and sporadic low-resistance anomalies at the depth of less than 10 km. At 15-30 km, there are several NE- and NW-trending high-conductivity anomalies and there is a large area of SN-oriented high conductor at 50 km. The outcomes of the investigation demonstrate: 1) The high-resistance anomalies found in the NE direction of the upper crust are consistent with the position of shallow faults and rifts. They constitute the main components of the upper crust of the Songliao Basin; 2) The high conductivity anomalies occurring at the intersection of the central depression zone and the north plunge zone represent the thinnest position of the lithosphere in the Songliao Basin; 3) The high-conductor anomalies in crust-mantle can be attributed to the paleo-shear zones which were generated in the collage of micro-continents during the Paleo-Asian Ocean closure, and reactivated under the continuous transformation of the Mongolia-Okhotsk and the Paleo-Pacific tectonic systems in the later period.

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10	Key Points:						
11	• 3D modeling of Songliao Basin via magnetotelluric array dataset						
12 13	• Discovering conductor in the northern Songliao Basin hints at the extreme point of lithospheric thinning						
14 15	• The superposition of tectonic systems is responsible for the current crust-mantle structure.						
16							

17 Abstract

18 The creation and evolution of the Songliao Basin is closely related to the closure of the Paleo-Asian Ocean and the Mongolia-Okhotsk Ocean, and the subduction of the Paleo-Pacific Ocean. 19 In an attempt to demonstrate the various attributes of the lithospheric structure under structural 20 superimposition and transformation of the Songliao Basin, this work used full impedance 3D 21 inversion to obtain a 3D electrical structural model of the northern Songliao Basin for the first 22 23 time. The results showed there are NE-trending high-resistance anomalies and sporadic lowresistance anomalies at the depth of less than 10 km. At 15-30 km, there are several NE- and 24 NW-trending high-conductivity anomalies and there is a large area of SN-oriented high 25 conductor at 50 km. The outcomes of the investigation demonstrate: 1) The high-resistance 26 27 anomalies found in the NE direction of the upper crust are consistent with the position of the volcanic rocks predicted by the reflection earthquake, whereas the low-resistance anomalies 28 correspond to the distribution of shallow faults and rifts. They constitute the main components of 29 the upper crust of the Songliao Basin; 2) The high conductivity anomalies occurring at the 30 31 intersection of the central depression zone and the north plunge zone represent the thinnest position of the lithosphere in the Songliao Basin; 3) The high-conductor anomalies in crust-32 mantle can be attributed to the paleo-shear zones which were generated in the collage of micro-33 continents during the Paleo-Asian Ocean closure, and reactivated under the continuous 34 transformation of the Mongolia-Okhotsk and the Paleo-Pacific tectonic systems in the later 35 period. 36

37 **1 Introduction**

The Songliao Basin is a large Mesozoic-Cenozoic sedimentary basin, superimposed on 38 the Paleozoic basement and located in northeastern China. The precise location of the Songaliao 39 Basin is the eastern section of the Central Asian Orogenic Belt (CAOB). The main body of the 40 41 regional structure is in the Songnen Block which lies adjacent to the Xing'an block, by the 42 Nenjiang fault in the west and connected with the Jiamusi block in the east by the Yilan-Yitong fault. The dynamic background of Songliao Basin's is relatively complicated, as it is located at 43 the intersection of multiple periods of structural systems. During the Paleozoic era, the Songliao 44 Basin experienced the evolution of the Paleo-Asian ocean tectonic system. With the closure of 45 the Paleo-Asian ocean, many micro-continent blocks in the Northeast (E'ergun block, Xing'an 46 block, Songnen block, Zhangguangcailing block, Jiamusi block, and Xingkai Block) were 47 combined into a composite land block (Xiao et al., 2003; Miao et al., 2008; Jian et al., 2008). 48 During the Mesozoic, the "scissor" closure of Mongolia Okhotsk Ocean from west to East and 49 the westward subduction of the Paleo-Pacific slab placed the Songliao Basin under the combined 50 51 action of these two dynamic systems (Xu et al., 2009; Wu et al., 2011).

52 The dynamic yet complicated background restricts the formation and evolution of 53 magmatism, stratigraphic deposits, and structural frameworks of the Songliao Basin. The closure 54 and extinction of the Paleo-Asian Ocean and the Mongolia-Okhotsk Ocean, and the westward

subduction of the Paleo-Pacific Ocean left a widespread of granite along the Solon-Xar Moron 55 River-Changchun area, towards the west of the Da Hingan Mountains, and east of the Korean 56 Peninsula-Songliao Basin (Wang et al., 2017). Under the action of different structural systems, 57 there are regional unconformities at the top of the Yingcheng Formation and Nenjiang Formation 58 59 in the Songliao Basin as well as their upper and lower structural frameworks, hence the stratigraphic distribution and direction, and rock combinations are significantly different (Wang 60 et al., 2007). Zircon U-Pb age statistics of volcanic rocks revealed that there were multiple 61 periods of volcanic activity in the Songliao Basin during the Mesozoic era. Volcanic rocks of 62 different periods recorded the process of transition from the Mongolia-Okhotsk tectonic domain 63 to the Pacific tectonic domain. (Xu et al., 2013). Under the joint action of the Mongolia-Okhotsk 64 tectonic system and the Pacific tectonic system, the hydrocarbon seismic stack section in the 65 Binbei area of the Songliao Basin presents a thrust nappe fault system having a two-way form, 66 showing uneven strength longitudinally and laterally (Shan et al., 2009). Various phenomena 67 indicate that the current structural morphology of the Songliao Basin is a result of the joint action 68 of different structural systems. Although there is a continuous superimposition of late structural 69 systems the traces of pre-existing structural systems can still be observed from various angles. 70

71 Besides, the subduction of the Pacific plate into Eurasia resulted in the thinning of the lithosphere in eastern China (Wu et al., 2005; Tao et al., 2014; Zhang et al., 2014). Thus, in 72 comparison to the Da Hingan Mountains and Zhangguangcai Mountains flanking on both sides, 73 74 the depth of lithospheric mantle bottom interface of Songliao Basin is as thin as 60-100 km (Wang et al., 2016; He and Santosh, 2016; Han et al., 2018). The maximum geothermal gradient 75 is 62°C/km (Liu et al., 2017). The thinning of the lithosphere is accompanied by crustal thinning 76 to a considerable extent. The seismic profile revealed that the depth of the Moho in the Songliao 77 Basin is 29~38 km (Yang et al., 2003; Wang et al., 2016). The shallowest position is on the 78 Mingshui-Anda-Changling line, and the variation characteristics show a trend of gradually 79 getting deeper from the middle to the east and west sides. 80

Detection and imaging of complex and deep processes that take place in the lithosphere 81 by the effect of fluids are carried out by the Magnetotagnetic field (MT). Such detection is 82 83 possible by the magnetotactic field is possible because resistivity is essentially a transmission property of the medium and is particularly sensitive to the presence of low resistivity phases of 84 interconnections, such as partial melts or aqueous fluids (Wei et al., 2001; Wannamaker et al., 85 2008). With the rapid development of 3D magnetotelluric inversion (Siripunvaraporn et al., 86 2005; Egbert and Kelbert, 2012), it has been used and is playing a major role in the study of the 87 88 basin's structure (Peacock et al., 2015; Zeng et al., 2015) and volcanic mechanism (Heise et al., 2008; Aizawa et al., 2014; Zhang et al., 2016; Li et al., 2020). Earlier MT studies in this area 89 were mostly conducted on isolated profiles, thus making it difficult to conduct a comprehensive 90 study of large structures. There is, however, still a major lack of high-resolution electrical 91 92 evidence for the transformation of the lithospheric structure of the Songliao Basin by three major 93 tectonic systems. This paper uses 157 broadband magnetotelluric sounding data in the northern

- 94 Songliao Basin (Fig 1) to perform 3D magnetotelluric inversion giving the first electrical
- structure of crust and mantle in the area and explains the characteristics of the lithosphere under
 the superposition and transformation of the multi-phase structural system.

97 2 Research Method

98 2.1 Data collection and processing

99 The research area is located at an expanse of 120,000 square kilometers in northeastern China, including Heilongjiang Province, Jilin Province, and Liaoning Province (about 100 123°-127° east longitude and about 45°-49° north latitude). The basin is divided into five 101 primary structural units according to the distribution and development characteristics of 102 the fault depressions, namely; the central depression zone (I), the southeast uplift zone 103 (II), the northeast uplift zone (III), the northern plunge zone (IV), and the west slope zone 104 (V). Multiple secondary fault depressions and fault uplifts lie inside. This work uses a 105 total of 157 broadband magnetotelluric sounding data points, of which 122 survey points 106 were from the exploration work of Daging Oilfield Research Institute from 1994 to 1999, 107 and the remaining 35 survey points were from Jilin University from June 2017 to 108 September 2018. MTU-5 magnetotelluric instruments produced by Phoenix Geophysics 109 of Canada were used for field collection of magnetotelluric sounds and the effective 110 frequency band was 320~0.0005 Hz. The poles were arranged according to the tensor 111 112 measurement method. Each measuring point measures 5 components, 3 of them being magnetic field components while the other 2 are mutually orthogonal horizontal electric 113 114 field components. The signal acquisition time was about 20 hours and the average point distance was 10 km. 115

During data processing SSMT2000 software was used to perform fast Fourier transform 116 on the original time series, transforming the time domain signal into frequency domain 117 data, thus obtaining high-quality impedance tensor information via processing techniques 118 such as "Robust" estimation and power spectrum selection. Following a series of 119 processing steps, the apparent resistivity and phase curves of all measuring points were 120 121 acquired. The data quality of most measuring points was found to be good, and any 122 sudden jump point in the frequency band was deleted before participating in the inversion. 123



Figure 1. The location map of the magnetotelluric sounding line in the north of Songliao Basin

To verify the reliability of the data, three wells around the study area (see Fig1 for 126 locations) (Linshen 2, Zhao 12 and Chaoshen 1 wells with depths of 3450 m, 3217 m, 127 and 3052 m respectively) were utilized to calibrate the shallow electrical structure (Fig 128 2). The borehole resistivity logging results revealed that, except for a small range uplift of 129 resistivity at a depth of 400 m, the other shallow electrical characteristics of 1200 m all 130 showed low resistance of about a few $\Omega \cdot m$; the resistivity of 1200-3000m increased and 131 was found to be directly proportional to the depth and grew rapidly to approximately 132 $1000\Omega \cdot m$. At around the same time, the geometric mean apparent resistivity of the 133 nearest magnetotelluric sounding (MT) point around the well was selected for one-134 dimensional OCCAM inversion. The inversion results were found to be consistent with 135 the trend of resistivity logging curves, showing "sudden jump-low-high resistance". The 136 changing trend is sufficient evidence of the reliability of MT data. 137

Depth (m) LLD R25 R25 Zhao12 (Ω-m) 1000 1 Linshen2 (Ω-m) 1000 1 Zhao12 (Ω-m) 1000 1 Chaoshen1 (Ω-m) 1000 1 1 (Ω-m) 1000 1 (Ω-m) 1000 1 (Ω-m) 1000 1 (Ω-m) 1000 1 (Ω-m) 1000 1 SSA-042S SSA-042S 1 SSA-042S SSA-042S 1 SSA-042S 1 SSA-04		Well Linshen2	Well Zhao12	Well Chaoshen I	Relative resistivity ratio	Resistivity	of near-well MT bat	hymetric (Ω·m)
	Depth (m)	LLD 1 (Ω·m) 1000	R25 1 (Ω·m) 1000	R25 1 (Ω·m) 1000	Zhao12 1 (Ω·m) 1000 Chaoshen1 1 (Ω·m) 1000 Linshen2	$\frac{\text{Linshen2}}{1 (\Omega \cdot m) 1000}$ $\frac{\text{SSD-041A}}{1 (\Omega \cdot m) 1000}$	Zhao12 1 (Ω·m) 1000 SSA-032S 1 (Ω·m) 1000	Chaoshen1 1 (Ω·m) 1000 SSA-045S 1 (Ω·m) 1000
	3000 2000 1000	الإدام المالية المالية المركزة المالية المحدة مستسمعة والمنافضة معتمانية والمحدين المحافظة ومناسرة فرارة المسادر المرم	كالماليان المالية المقادية المطاملين المتحاصرة المتكارسة معهدة ومعارفة والمعاريق والمحاصر ومعادية ومحاسبة والمعالية والمحالية وال	المجامعة من المحدث من المحدث معقولها المحالية المؤالين بالمحالة معتقدا المحالية من المحالية وفيا بعصورا لمحالية وفي المحالية المحالية المحالية والمحالية المحالية المحالية والمحالية والمحالية والمحالية والمحالية والمحالية والمحالية المحالية والمحالية المحالية المحالية والمحالية والم				

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Figure 2. Comparison of the results of different borehole resistivity logs and near-hole magnetotelluric sounding

(In addition to the depth track, columns 1-3: Linshen2, Zhao12, Chaoshen1 resistivity logging
 curves; column 4: relative resistivity values; columns 5-7 show the joint mode of resistivity

- logging and near-well earth electromagnetic sounding point 1D OCCAM inversion result
 comparison)
- 145 2.2 Dimensional analysis

To construct an electrical model that is an accurate reflection of the geological structure, 146 it is imperative to perform data analysis before inversion. The phase tensor has an 147 interesting characteristic of being less affected by the current distortion effect caused by 148 the non-uniformity of the near-surface, and it has become a valuable parameter for 149 analyzing the dimensionality of the region. Using 3D model studies, it has been 150 illustrated that the direction of the principal axis of the phase tensor reflects the lateral 151 changes in the conductivity structure of the underlying region (Caldwell et al., 2004). 152 When the 2D deviation is 0 and the ellipse is approximately circular, the underground 153 medium is considered to have a 1D characteristic whereas when the 2D deviation is 0 and 154 the long and short axes of the ellipse are not equal, the underground medium is 155 considered to have 2D characteristic; when the degree is non-0, it is a 3D medium, and 156 3D characteristic is proportional to the magnitude of 2D deviation. In general, when the 157 absolute value of 2D deviation rises above 3°, it is expressed as a strong 3D characteristic 158 (Caldwell et al., 2004). 159



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Figure 3. Horizontal distribution map of phase tensor of all frequency points in northern Songliao Basin

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Fig 3 illustrates the sub-band phase tensor level distribution of all measurement points. 163 The color of ellipse represents the value of the 2D deviation β , and ellipse azimuth of 164 most measurement points in 4 periods (T) is N 20°E —N 40°E, particularly the azimuth 165 angle of the main axis of the ellipse around the structural zone is basically along the 166 structural direction. When T=10s, the 2D deviation of most measuring points was small 167 and showed 1D and 2D characteristics. As the period increased, the 2D deviation value 168 began showing an increasing trend, indicating that the part having depth has strong 3D 169 characteristics. The direction of the major axis of the ellipse started becoming 170 increasingly diverse, including NE, NW, and near EW directions, revealing the complex 171 deep structure pattern of the area. The black arrow represents the real induction vector, its 172 direction pointing to the high conductor (Parkinson, 1959), whereas the lateral 173 unevenness is represented by the length of the arrow. From a macro perspective, the 174 direction of the induction vector is chaotic, and the length is varying, especially in the 175 middle of the central depression area. Perhaps, the good stratification in the shallow part 176 of this area affected the orientation of the induction vector, thus pointing to the structural 177 complexity of the Songliao Basin. Nevertheless, the direction of the induction vector 178 mostly points to the location of the nearby fracture. The direction of the fracture is not 179 unique but exists at least in two groups. 180

181 The color of the ellipse represents the two-dimensional deviation value, the black arrow 182 represents the real induction vector, the thin black dotted line represents the division of 183 the tectonic units within the basin (Fig. 1), the thick red dotted line represents the location 184 of the known deep fault whereas the missing points represent the eliminated frequency



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Figure 4. The results of phase tensor analysis of Line F in northern Songliao Basin

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Fig. 4 represents the results of the phase tensor analysis of the Line F. Line F passes from 187 south to north through the central depression zone and the northern plunge zone (Fig. 1). 188 There is a larger period between 10 s and 1000 s on the southern side of Line F and 2D 189 deviation degree and has a large ellipticity. The characteristics near the north gradually 190 191 weakened, indicating that the deepest part of the central depression area has a strong 3D geological background which only weakens when it extends northward. Hence it is 192 evident from the phase tensor analysis that the crust and mantle of the Songliao Basin 193 have an obvious 3D characteristic, particularly in the central depression zone. Therefore, 194 only 3D inversion can provide a more reliable electrical model. 195

196 2.3 Data inversion

Mod3DMT (Egbert & Kelbert, 2012) based on the nonlinear conjugate gradient (NLCG) 197 algorithm was used to perform 3D inversion of the full impedance data of 157 measuring 198 points. A uniform half-space of $100\Omega \cdot m$ is selected as the initial model, and the 199 frequency lies in 1000-0.0001Hz range, grid division in X direction: grid length 5km, 70 200 grids, 5 grids on both sides expanded by a factor of 2.5; grid division in Y direction: grid 201 length 5km, 90 grids, with 5 grids expanded on both sides by a factor of 2.5; grid division 202 in the Z direction: first layer 50m, layer increment factor: 1.23, a total of 35 grids, and 203 borders expanded by a factor of 2.5 5 grids. Adding 5% as noise, a total of 107 inversions 204 were performed, and the RMS decreased from 44.18 to 2.95. Finally, a 3D geoelectric 205 model of the lithosphere scale of the northern Songliao Basin was obtained. 206

207 2.4 Model validation

The distribution of total impedance RMS of 157 measuring points is illustrated in Fig.5. Except for s large fitting error of individual measuring points, the fitting error of most measuring points was concentrated between 1-3, which lies close in value to the overall fitting error. There is no over-fitting and under-fitting hence, proving the credibility of the 3D inversion model.



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Figure 5. Fitting situation of all impedances of all sounding points of magnetotelluric sounding
 in northern Songliao Basin

The gray line indicates the provincial boundary; the dotted line indicates the division of the structural units within the basin (see Fig. 1), while the black circle indicates the location.

218 **3 Electrical Structure Characteristics**

219 To analyze the electrical characteristics of slices at different depths, the resistivity slices with depths of 5 km, 10 km, 15 km, 20 km, 30 km, and 50 km were selected corresponding 220 respectively to the depth of shallow layer, the top surface of the basement (upper crust), upper 221 and lower crust boundary, lower crust. Moho surface and upper mantle in the northern Songliao 222 Basin. Electrical gradient bands and low-resistance anomalies are often the most common signs 223 224 to identify the location of faults. While the interface of the resistance level conversion is often indicative of them being two different substances and that faults are prone to occur between 225 them, on the other hand, faults or alteration zones provide a channel for magma to flow upwards, 226 and there are generally broken and loose low-resistance media where the fault zone develops. 227 Hence low-resistance anomalies appear in the electrical structure. After precisely combining the 228 location of surface faults and electrical structural characteristics, the locations of the Binzhou 229 fault zone, the Nehe-Suihua fault zone, Yi'an-Tongyu fault zone, Sunwu-Shuangliao fault zone, 230 and Nenjiang-Balihan fault zone were determined (Fig.6). 231

Some anomalous complexities are exhibited by the 5 km underground slice. Only towards the north of the boundary between the northern plunge zone (IV) and the central depression zone (I),

a connected high-conductivity anomaly C1 appears, with broken high-conductors in the NE-SE

and NW-SE directions. Preliminarily, it was inferred that C1 is the intersection of NE and NW 235 trending faults/fault depressions in the shallow part. The high resistance anomalies in this depth 236 range reflect the placement of the volcanic rock aggregates. The horizontal slice having a depth 237 of 10 km, has low-resistance anomalies that are found to be distributed in the western slope zone 238 239 (V), the northern plunge area (IV), the central depression zone (I), and the northeast uplift zone (III), with a resistivity lower than 15Ω . M, the high-conductivity anomaly at this depth initially 240 spreads in the NW and NE directions compared to the electrical structure of shallower slices. At 241 depths of 15 km, 20 km, and 30 km, the range of high-conductivity anomalies is comparatively 242 wide, and there were band-shaped high-conductivity anomalies in NW, NE, and EW direction 243 with a resistivity value of about 10Ω m. The multitropism of the depth anomaly suggests that 244 the northern Songliao Basin has experienced the superposition of many tectonic systems, rather 245 than a single structural realm. In the study of the Man-Sui section, a consensus was reached on 246 the existence of low-velocity, low-density, and high-conductivity layers at a depth of 15-25 km, 247 and it is believed that the plastic middle-lower crust is fluid rich. So, the high conductivity 248 anomalies in the NW and NE directions are most likely a fluid-rich fracture system. The contours 249 of low- and medium-resistance anomalies below 30Ω ·m at a depth of 50 km continue the 30-km 250 high-conductivity anomaly. It is worth noting that a large region of near-south high conductivity 251 anomaly lies at the junction of the central depression zone and the northern subduction zone, 252 having a longitudinal expanse of several tens of kilometers, which is presumably a manifestation 253 of substance upheaval in the deep asthenosphere in this area. 254





The white inverted triangle is the location of the measuring point; the white square is the focal
location of the earthquake in the area (average focal depth at C2 is 8.625km) (data from China
Seismic Network); the white circle is the location of the well that encountered CO₂ (according to
Zhang et al., 2009); the black dotted line demarcates the boundary of the structural unit; the
black dense line is fault zone ①Binzhou fault zone ②Nehe-Suihua fault zone ③Yi'an-Tongyu
fault zone ④Sunwu-Shuangliao fault zone ⑤Nenjiang-Balihan fault zone

263 4 Discussion

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264 4.1 Upper crustal material composition

Underneath the Cenozoic sedimentary caprock in the Songliao Basin, there is a relatively complete Cretaceous stratum. The early Cretaceous layer formed fault depressions after experiencing the evolutionary history of the fault depression period, including Yingcheng, Shahezi, and Huoshiling formations. The seismic cross-section through the Songliao Basin shows that the burial depth of the faulted layer is 2~5 km, and it has obvious characteristics of being deep in the middle and shallow on both sides. During this period, the asthenosphere upwelled due to the subduction of the Pacific plate, the lithosphere squeezed and stretched, the crust straightened and cracked, several crust faults appeared, and a fairly good number of magma intrusions and eruptions developed along the faults. Brittle deformation occurred and the land of the upper crust sank, forming a series of faulted basins distributed in the NNE direction (Wang et al., 2016).

- The 5 km electrical slice showed an obvious crisscrossing high-conductivity anomaly C1, 276 seemingly caused by the superposition of two-way faults/fault depressions, which exhibit 277 high-conductivity features that cross each other. Existing data show that the NW and NE-278 trending faults in the basin form an orthogonal grid-like extensional fault system, 279 dividing the Songliao Basin into east-west zones and the north-south block structure, 280 which is consistent with the electrical structure investigated and mentioned in this work. 281 Due to the frequent occurrence of volcanic magmatism in the Songliao Basin since the 282 Mesozoic (Wu et al., 2005; Zhang et al., 2010), volcanic rocks are quite widespread. 283 Drilling, seismic, gravity, and magnetic data show that deep volcanic rocks in the 284 Songliao Basin are widely distributed in the Jurassic Huoshiling Formation and Early 285 Cretaceous Yingcheng Formation (Wei et al., 2019). The distribution map of aggregates 286 of volcanic rock predicted by regional earthquakes (Zhang et al., 2010), is found to be 287 consistent with the NE-direction high-resistance anomaly position, so the high-resistance 288 anomaly within a depth of 5 km is inferred as volcanic rock. In the 10 km depth slice, 289 there are two NW-direction high conductivity anomalies, C2 and C3, respectively 290 corresponding to the seismically dense area and the location of the well, where CO_2 is 291 drilled in the region (Fig. 6) hence proving that C2 and C3 are avenues for the occurrence 292 293 of earthquakes and the migration of CO_2 for the fracture system inside the crust at this depth. 294
- Viewed together, the geological background and the electrical structure of 5km and 10km slices, it can be estimated that locally distributed volcanic rock bodies, NE, and NWoriented faults/fault depressions together constitute the key structural elements of the upper crust of the Songliao Basin.
- 4.2 Lithosphere thinning in Songliao Basin

A large volume of evidence suggests a thinning of the lithosphere in the Songliao Basin. Surface wave imaging studies showed that the lower part of the Songliao Basin is a thin lithospheric caprock, with a thickness much less than the average thickness of the continental lithosphere (Ren et al., 2002; Meng., 2003; Guo et al., 2014). The images released by the broadband seismic stations in Northeast China show that the average depth of the Lithospheric Asthenosphere Boundary (LAB) in the Songliao Basin is about

80 km. The shallowest part lies below the center of the basin, and geological cross-306 section data showed that the thickness of the lithosphere near Anda in the Songliao Basin 307 is only 60 km (Cheng and Wu, 1994). The thinning of the lithosphere was generally 308 believed to be related to the extension of the Mesozoic lithosphere, and its direct stress is 309 derived from the undercutting of magma. Studies have found that the uppermost surface 310 of the asthenosphere below the basin presents an upwardly convex arc (Guo et al., 2014). 311 The largest part of the depression of rift valley in the Songliao Basin corresponds to the 312 upheaving region of the asthenosphere (Han et al., 2018), but earlier researches have not 313 pointed out the extreme point of lithospheric thinning, that is, the center of the upwelling 314 of the asthenosphere area. In this investigation, which is based on a 50 km mantle-scale 315 horizontal slice, it has been found that there is a sizeable region of high conductivity 316 anomaly C4 lying under the Lindian fault depression located at the junction of the central 317 depression zone and the northern plunge zone, and covering an area of tens of kilometers 318 in the SN direction in the longitudinal slice. It has an uplifted shape with a top interface 319 of 45 km, which can be inferred to be the center of the asthenosphere upwelling. It 320 authenticates the existence of a relatively thin lithosphere in the Songliao Basin as well as 321 the continuous erosion of the bottom interface of the lithosphere due to the under-322 intrusion of the asthenosphere (Kuritani et al., 2011; Liu et al., 2016), thus forming an 323 uplift of the asthenosphere and gradually transforming the lithosphere into new-born 324 rocks with a high conductivity characteristics ring. It is worth mentioning that the center 325 of the upwelling of the asthenosphere is located below the Lindian fault depression, and 326 the shallowest part is 45 km deep. 327

4.3 The remaining traces of the three major structural systems in the Songliao Basin

The evolution of the Songliao Basin has undergone superimposed transformation of three 329 major structural domains of the Paleo-Asian Ocean, Mongolia-Okhotsk, and Paleo-330 Pacific (Wang et al., 2017; Liu et al., 2017; Wu et al., 2011; Xu et al., 2013). Despite the 331 superposition and transformation of the later structural domains, the remnants of the pre-332 existing structural system can still be observed. There was a deep seismic reflection 333 profile in the same direction as SN in the vicinity of Line F, which recorded the tectonic 334 events of south-to-north subduction during the closure of the Paleo-Asian Ocean and the 335 subduction residue of a block during the closure of the eastern Mongolian-Okhotic Ocean 336 (M3 and M4 in Fig. 7). The east-west sloping lower crust reflections and mantle 337 reflections (M1 and M2 in Fig. 7), existing in the seismic profile on the same location as 338 line D are considered to be the result of the overlap of each other, subduction 339 compression and strong extension that occurred later between Mongolia-Okhotsk tectonic 340 341 domain and the paleo-Pacific tectonic domain in the deep part of the Songliao Basin (Fu et al., 2019). 342

Early subduction residues and bidirectional convergence splice marks revealed by deep 343 seismic reflections verified that the Songliao Basin was once in a compression 344 environment under the action of the tectonic system. Due to the continuous impact of 345 external stress, the crust-mantle was vulnerable to damage, resulting in a structurally 346 weak zone or a patchy unstable zone. The Mongolia-Okhotsk tectonic system and the 347 Paleo-Pacific tectonic system are related with respect to time and space, and the two 348 intersect during the transformation of the Jurassic and Cretaceous (Fu et al., 2019), 349 through the combination of volcanic rocks. The distribution characteristics in time and 350 space found that the spatial scope of the influence of the Pacific Rim tectonic system in 351 the East Asian continent lies mainly in the Songliao Basin and its east whereas the spatial 352 scope of the Mongolia-Okhotsk structure system is primarily lying in the west of the 353 Songliao Basin and the edge of North China (Xu et al., 2013). Hence, the Songliao Basin 354 belongs to the joint action area of the two tectonic systems. The closure of the Mongolia-355 Okhotsk Ocean and the dehydration of the Pacific slab trapped in the mantle transition 356 zone, set off asthenospheric upwelling, filling the weak structural zone or the unsteady 357 assembly zone that had been left behind, leading to its activation and becoming a partial 358 melt, thus providing a channel for the upward migration. Therefore, the structural system 359 should have a corresponding performance on the geoelectric structure. The fact that the 360 remains of the tectonic system correspond to the magma channel is not accidental. The 361 joint seismic section (Li et al., 2014) passing through the northern margin of North China 362 and the southern part of the CAOB also found, low-velocity gradient, and channel-like in 363 lower crustal. It was speculated that it may have been an active continental margin where 364 the Paleo-Asian Ocean subducted and died southward, and evolved into a channel for 365 magma intrusion in an extensional environment following the collision. 366







FD stands for fault depression; FU stands for fault uplift; M stands for mantle reflection during 371 the earthquake, M1 designates residual traces of north-south subduction tectonic events during 372 the closure of the Paleo-Asian Ocean, M2 designates residual subduction of a block during the 373 closure of the eastern Mongolia-Ekhotsk Ocean, M3 designates subduction traces under the 374 action of the Mongolian-Okhotsk Ocean, M4 designates subduction traces under the action of the 375 ancient Pacific Ocean (Fu et al., 2019); the thin black line is the tectonic line in the seismic crust; 376 377 the thick line is the depth of the Moho surface (according to Sun, 2019); the dotted line in the 378 middle and lower crustal interface for seismic identification (according to Wang, 2015); thin

white lines indicate the fault zone (1)Binzhou fault zone (2)Nehe-Suihua fault zone (3)Yi'an Tongyu fault zone (4)Sunwu - Shuangliao fault zone (5)Nenjiang- Balihan fault zone.

This investigation selected the longitudinal slices in the same direction and close to the 381 seismic profiles for analysis, which include the Line D in the NW direction and the Line 382 F in the near SN direction. The depth of the Moho is about 30 km (Wang et al., 2016), 383 and the ups and downs of the Moho are drawn according to the Songshen deep seismic 384 reflection profiles (Yang et al., 2004). In the lower crust, there are low-resistance 385 anomalies C5, C6, C7, and C8 that cross the Moho, and the Moho at the bottom is 386 missing and overlap to varying degrees. The electrical results show that the high-387 conductivity layer in the crust originates from the upper mantle, so the absence and 388 overlap of the Moho may be caused by the upwelling of asthenospheric materials to the 389 interception of the lower crust by the brittle-ductile shear zone. It is worth noting that a 390 large number of high-conductivity layers in the mantle are filled below 30 km, with a 391 resistivity value of more than ten Ω ·m, and a slight uplift in the middle which is quite 392 prominent on both sides, and are connected with high-conductivity anomalies under the 393 crust of the Gulong and Yian fault depressions, Meilis and Suihua fault depressions 394 respectively. Although the traces of subduction residues cannot be distinguished in the 395 electrical section in terms of relative position and burial depth, the upwelling high 396 conductivity anomalies C5, C6, C7, and C8 on both sides of the Line D and Line F may 397 be the paleo-shear zones formed during the microcontinental assemblage when the Paleo-398 Asian Ocean was closed. (Yang et al., 2001). The combination of the closure of the 399 Mongolian-Okhotsk Ocean and the subduction of the Paleo-Pacific Ocean reactivated the 400 paleo-shear zones. Due to the dehydration of the Pacific plate lying in the mantle 401 transition zone (Kimura et al., 2018; Kuritani., 2011, 2019), the paleo-shear zones were 402 filled with upwelling asthenosphere material, thus showing the high conductivity, 403 characteristic of the connection with the upper mantle. 404

405 Therefore, we propose a schematic cartoon of the geoelectrical model presented above (Fig. 8). During the tectonic regime of the Paleo-Asian Ocean, the Paleo-Asian Ocean 406 was continuously consumed, the Songnen block, where the Songliao Basin is located, 407 was closing to the adjacent microcontinental blocks one after another. The intense 408 squeezing action resulted in several palaeo-shear zones in the lower part of the Songliao 409 Basin (Fig. 8a). The period when the Mongolian-Okhotsk and the paleo-Pacific tectonic 410 systems worked together, the Mongolia-Okhotsk Ocean and the Paleo-Pacific Ocean 411 respectively acted on the northwest and east sides of Songliao Basin. The two-way 412 converging ocean slabs reactivated the paleo-shear zones. The later dehydration of the 413 414 Pacific slab stimulated the asthenospheric material to rise along these weak zones allowing the paleo-shear zones to be "marked" in the geoelectrical model (Fig. 8b). 415



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417	Figure 8. Schematic cartoon of the geoelectrical model. (a) represents the period when the
418	Paleo-Asian ocean tectonic system was in effect (according to Li, et al., 2016, 2017). (b)
419	represents the period when the Mongolian-Okhotsk and the paleo-Pacific tectonic systems
420	worked together (according to Cheng, et al., 2018).

421 **5 Conclusions**

In this work, the full-impedance magnetotelluric 3D inversion of the magnetotelluric sounding
profile is carried out in the northern Songliao Basin, and a 3D electrical structure model is
established. Based on the characteristics of the electrical structure of the slices electrical
structure, we arrive at the following conclusion:

426 1) Based on the distribution and electrical structural characteristics of volcanic rocks
 427 predicted by earthquakes, it has been found that the volcanic rock clump is consistent
 428 with the location of NE-trending high-resistance anomalies, while the scattered low 429 resistances distributed in the NE and NW directions correspond to the location of faults

- 430 /fault depressions. It has been inferred that the upper crust of the northern Songliao Basin
 431 is composed of volcanic rocks and faults /fault depressions;
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 2) Mesozoic underplating set off the extension of the lithosphere in the Songliao Basin, resulting in an uplifted shape on the top of the asthenosphere in the Songliao Basin. The most prominent part of the uplift is located at the junction of the central depression zone and the northern submerged area (Below the Lindian Fault Depression). The buried depth is 45 km, which is the center of the upwelling of the asthenosphere hot material;
- 437
 3) It has been found that there is a set of symmetrically distributed high conductivity
 438 regions under the crust of the Gulong and Yian fault depressions, Meilis and Suihua fault
 439 depressions respectively in the Songliao Basin. This peculiarity can be attributed to the
 440 weak tectonic zone produced by the "Paleo-shear" effect during the closure of the Paleo441 Asian Ocean. Afterwards they were reactivated by the continuous transformation of the
 442 Mongolia-Okhotsk tectonic domain and the Paleo-Pacific tectonic system in the later
 443 period.

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448 Several figures are plotted using the MtPy software (Krieger and Peacock, 2014; Kirkby et al.,

- 449 2019). The data archiving is underway in https://www.pangaea.de/, and the copy data has been
- 450 uploaded in supporting information (SI).

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