Rocket Observation of sub-relativistic electrons in the quiet dayside auroral ionosphere

Taku Namekawa¹, Takefumi Mitani², Kazushi Asamura³, Yoshizumi Miyoshi⁴, Keisuke Hosokawa⁵, Yasunobu Ogawa⁶, Shinji Saito⁷, Tomoaki Hori⁴, Shin Sugo⁸, Oya Kawashima⁸, Satoshi Kasahara⁸, Reiko Nomura⁹, Naoshi Yagi¹⁰, Mizuki Fukizawa¹⁰, Takeshi Sakanoi¹⁰, Yoshifumi Saito¹¹, Ayako Matsuoka¹², Iku Shinohara⁹, Yury V. Fedorenko¹³, Alexander Nikitenko¹³, and Christopher Koehler¹⁴

¹University of Tokyo
²ISAS/JAXA
³The Institute of Space and Astronautical Science
⁴Institute for Space-Earth Environmental Research, Nagoya University
⁵University of Electro-Communications
⁶National Institute of Polar Research
⁷National Institute of Information and Communications Technology
⁸The University of Tokyo
⁹Japan Aerospace Exploration Agency
¹⁰Tohoku University
¹¹Institute of Space & Astronautical Science
¹²Kyoto University
¹³Polar Geophysical Institute
¹⁴University of Colorado Boulder

November 24, 2022

Abstract

An energy spectrum of electrons from 180 keV to 550 keV precipitating into the dayside polar ionosphere is observed for the first time by the HEP instrument onboard the RockSat-XN sounding rocket under geomagnetically quiet condition (AE [?]100 nT) at Andøya, Norway. The observed energy spectrum of precipitating electrons follows a power law of -4.86 and the electron flux does not vary much over the observation period (~274.4 seconds). A few minutes before the RockSat-XN observation, POES18 / MEPED observed precipitating electrons, which suggest chorus wave activities at the location close to the rocket trajectory. A ground-based VLF receiver observation at Lovozero, Russia also supports the presence of chorus waves during the rocket observation. A test-particle simulation for wave-particle interactions based on the Arase satellite data shows a similar energy spectrum of precipitating electrons, consistent with the RockSat-XN observation. These results suggest that the precipitation observed by RockSat-XN is likely to be caused by the wave-particle interactions between chorus waves and sub-relativistic electrons.

Rocket Observation of sub-relativistic electrons in the quiet dayside auroral 1 ionosphere 2

T. Namekawa^{1,2}, T. Mitani¹, K. Asamura¹, Y. Miyoshi³, K. Hosokawa⁴, Y. Ogawa⁵, S. Saito⁶, T. Hori³, S. Sugo², O. Kawashima², S. Kasahara², R. Nomura¹, N. Yagi⁷, M. 3

4

Fukizawa⁷, T. Sakanoi⁷, Y. Saito¹, A. Matsuoka⁸, I. Shinohara¹, Y. Fedorenko⁹, A. 5

Nikitenko⁹, C. Koehler^{10,11} 6

¹Japan Aerospace Exploration Agency, Japan, ²The University of Tokyo, Japan, ³Nagoya 7

University, Japan, ⁴The University of Electro-Communications, Japan, ⁵National Institute of 8

Polar Research, Japan, ⁶National Institute of Information and Communications Technology, 9

Japan, ⁷Tohoku University, Japan, ⁸Kyoto University, Japan, ⁹Polar Geophysical Institute, 10

Russia, ¹⁰Colorado Space Grant Consortium, USA, ¹¹University of Colorado at Boulder, USA 11

- 12
- Corresponding author: Taku Namekawa (namekawa@stp.isas.jaxa.jp) 13

Key Points: 14

- A sounding rocket observed an energy spectrum of sub-relativistic electron precipitation 15 • in the dayside polar ionosphere for the first time. 16
- Ground-based and satellite observations and simulation show that the observed electron 17 precipitation was caused by chorus waves. 18
- From the relation between the chorus wave and the aurora, the observed electrons likely 19 to coincide with dayside pulsating/diffuse aurora. 20
- 21

22 Abstract

- 23 An energy spectrum of electrons from 180 keV to 550 keV precipitating into the dayside polar
- 24 ionosphere is observed for the first time by the HEP instrument onboard the RockSat-XN
- 25 sounding rocket under geomagnetically quiet condition (AE ≤100 nT) at Andøya, Norway. The
- observed energy spectrum of precipitating electrons follows a power law of -4.86 and the
- 27 electron flux does not vary much over the observation period (~274.4 seconds). A few minutes
- 28 before the RockSat-XN observation, POES18 / MEPED observed precipitating electrons, which
- suggest chorus wave activities at the location close to the rocket trajectory. A ground-based VLF
- 30 receiver observation at Lovozero, Russia also supports the presence of chorus waves during the 31 rocket observation. A test-particle simulation for wave-particle interactions based on the Arase
- 32 satellite data shows a similar energy spectrum of precipitating electrons, consistent with the
- 33 RockSat-XN observation. These results suggest that the precipitation observed by RockSat-XN
- is likely to be caused by the wave-particle interactions between chorus waves and sub-relativistic
- 35 electrons.
- 36

37 Plain Language Summary

38 Sub-relativistic electrons precipitating into the Earth's dayside polar ionosphere are observed by

a sounding rocket under geomagnetically quiet conditions. An energy spectrum of these

40 electrons in an energy range from 180 keV to 550 keV is, for the first time, reported. A possible

41 mechanism for generating this precipitation is indicated as a resonance scattering of electrons by

42 chorus waves, based on satellite, ground-based observations, and a test particle simulation.

43

44 **1 Introduction**

45 Precipitation of sub-relativistic or relativistic electrons is often observed in the auroral

ionosphere of the Earth [e.g. *Clilverd et al.*, 2006; *Dietrich et al.*, 2010; *Lorentzen et al.*, 2001;

47 Blake et al., 1996; Kurita et al., 2016]. It is suggested that this phenomenon is related to the loss

- 48 mechanism of the outer radiation belt electrons such as pitch angle scattering due to
- 49 magnetospheric plasma waves [Horne & Thorne, 2003; Kennel & Petscheck, 1966]. Such
- 50 electron precipitation is observed in the nightside and dawnside ionosphere with auroral
- activities. Besides, precipitation of sub-relativistic or relativistic electrons occurs also on the
- 52 dayside ionosphere, even when it is quiet. The SAMPEX satellite observed microbursts of the
- relativistic electron precipitation with a time scale of less than 1 second when AE was less than
- 54 100 nT and its peak occurrence rate of one microburst is every 18.9 min at L = 5.5 and ~10 MLT
- 55 [*Douma et al.*, 2017]. The POES satellites also observed > 30 keV electron precipitation at $AE \le$
- ⁵⁶ 100 nT and its peak is the magnetic pre-noon sector [*Lam et al.*, 2010]. However, there are no
- 57 direct observations of the energy spectra of the precipitating electrons at the dayside ionosphere
- under the quiet condition, and most of the features of these precipitations have not been
- ⁵⁹ understood yet. Characteristic features of the energy spectra of precipitating sub-relativistic or
- 60 relativistic electrons in the dayside ionosphere, such as slopes of the energy spectra provide
- 61 important information to verify the relationship between the precipitations and the loss
- 62 mechanism of the outer radiation belt electrons by comparison with numerical simulations [e.g.
- 63 Miyoshi et al., 2015a].

- 64 The dayside sub-relativistic or relativistic electron precipitation may be associated with
- 65 processes that drive daytime auroral activities [e.g. *Kurita et al.*, 2015]. The diffuse aurora is
- 66 commonly observed from the morning to noon sector at the equatorward edge of the auroral oval
- and has been considered to originate from the precipitation of tens of keV electrons [*Sandholt et*
- *al.*, 2002; *Newell et al.*, 2009], which is driven by a chorus wave in the magnetosphere [*Li et al.*,
- 69 2009; *Ni et al.*, 2011a, 2011b, 2014; *Shi et al.*, 2012; *Nishimura et al.*, 2013]. However, there
- ⁷⁰ have been few studies about dayside diffuse aurora and their characteristics have not been well
- understood. For example, they occur actively when the magnetic activity is low as Kp = 0, and
- some of them occur with pulsation [*Han et al.*, 2015].
- 73 In this paper, we report on the first in-situ observation of an energy spectrum of sub-relativistic
- electrons precipitating into the dayside polar ionosphere made by the RockSat-XN sounding
- 75 rocket experiment. The rocket was launched under the geomagnetically quiet condition. We also
- analyze data from the low-altitude POES satellite and ground-based VLF antenna data, which
 can be used for a proxy of the chorus waves. We also perform a computer simulation for wave-
- 78 particle interactions and compare calculated energy spectra with the RockSat-XN observations.
- 79 We describe the instrumental setup in the next section, which is followed by observational
- results in Sections 3 and 4. Discussion and summary are provided in Sections 5 and 6,
- 81 respectively.
- 82

83 2 Instrumentation



- 85 Figure 1. Schematic view of the High Energy Particle (HEP) detector onboard the RockSat-XN
- sounding rocket. A sample trajectory of an incoming electron is shown.
- 87
- 88 The RockSat-XN mission is an international student rocket program led by NASA and Colorado
- 89 Space Grant Consortium (COSGC). It carried 8 observation packages. One of the packages is
- 90 called PARM (Pulsating AuroRa and Microburst) [Sugo et al., 2020] which includes a high-

- 91 energy electron detector (HEP: High Energy Particle detector) and a magnetometer (AFG: ASIC-
- based FluxGate magnetometer). Pitch angles of observed electrons were calculated by using
- 93 AFG data in this study.
- 94 HEP was designed to measure energy spectra of electrons with energies ranging from 300 keV to
- 95 2 MeV. This sensor consists of a mechanical collimator and seven-layered silicon semiconductor
- detectors (SSDs) which measure energies of incident particles. Figure 1 shows the schematic
- view of HEP. The thickness of each SSD is 600 μm leading to the total thickness of all the SSDs
- of 4.2 mm. The SSDs of this thickness enable us to measure the electron's energy up to 1.7 MeV,
- according to the ESTAR web-database of electron stopping powers and ranges, which is
- provided by the National Institutes of Standards and Technology (NIST). The field of view of
- HEP is $22.6^{\circ} \times 22.6^{\circ}$ as geometrically defined by the center of the aluminum shield and the edge of the collimator. A 125 µm-thick aluminum sheet was installed just in front of the SSDs, to
- block electrons with energies less than 300 keV. As a result, 18 % and 70 % of electrons pass
- 104 through the sheet in the cases of 200 keV and 300 keV electrons, respectively. HEP was mounted
- 105 on the top of the rocket in such a way that the center of the field of view of HEP was parallel to
- 106 the thrust axis of the rocket. The pulse height of output signals generated by each SSD layer was
- 107 digitized by analog-to-digital converters in parallel and sent to the ground.
- 108 Figure 2 shows electron energy spectra obtained by the ground calibrations when HEP is
- 109 irradiated by monoenergetic electron beams. HEP can measure electrons between 300 keV and
- 110 2 MeV with energy resolution higher than 22 % ($\Delta E/E$, full width at half maximum). Table 1
- summarizes the performance and specifications of HEP. Note that the signal processing time of
- 112 HEP for each particle detection event is $\sim 5 \ \mu s$.
- 113



Figure 2. The electron energy spectra measured by HEP under the irradiation of monoenergetic electron beams (300, 500, 700 keV, 1, 1.5, 2 MeV). The total count for each energy spectra is

- 117 equal.
- 118

119	Table 1.	Performance	and s	specifications	of HEP.
-----	----------	-------------	-------	----------------	---------

Parameter	Value
Energy range	300 keV–2 MeV
Energy resolution (ΔE/E)	22%(FWHM) at 300 keV, 5%(FWHM) at 1 MeV
Field of view	$22.6^{\circ} \times 22.6^{\circ}$ pyramid
Geometric factor	$0.81 \text{ cm}^2 \text{ sr}$
Time resolution	Each event is time-tagged with 100 ns resolution
Mass	2.7 kg
Power	$6.7 \text{ W} (28 \text{V} \times 0.24 \text{ A})$

121 **3 Observation**

122 RockSat-XN was launched from Andøya Space Center, Norway (69.3°N, 16.0°E in the

geographic, 67.4°N 112.3°E in the geomagnetic coordinate system) at 09:13:00 UT on January

124 13, 2019. Figure 3 (a) shows the position of the launch site and the rocket trajectory. The apex

altitude of the trajectory was 184.1 km at 219.6 s from the launch. HEP measured incident

particles between 99.6s and 374s from the launch, which correspond to altitudes of 117.2 km at

upleg and 73.2 km at downleg, respectively. During the flight, HEP detected 4493 events in total.

The zenith angle of the center of the field of view of HEP was between 21.9° and 33.7° from the

local magnetic field. The center of the field of view of HEP pointed toward the direction of local

pitch angles between 33.8° and 45.5° . Since the field of view of HEP is $22.6^{\circ} \times 22.6^{\circ}$, the

131 observed event includes the electrons with pitch angles between 11.2° and 68.1° .

Figure 3 (b) shows altitude profiles of electron density measured by the EISCAT VHF radar
located in Tromsø, Norway (69.60° N, 19.20° E in geographic, and 67.20° N 115.25° E in the

geomagnetic coordinate system) on January 13, 2019. The radar pointed to the vertical direction

(an elevation angle of 90°). Note that the density enhancement above an altitude of 150 km after

136 8 UT was due to the photo-ionization caused by the sunlight. There was a weak enhancement of

electron density at altitudes above 85 km although the geomagnetic condition during the

observation was relatively quiet (AE index was less than 100 nT from 9 UT to 10 UT).

139

140 Unfortunately, no satellite was located right on the geomagnetic field line passing through the

rocket during the observation. However, there are ground-based and satellite-based observations

that suggest the possible appearance of the chorus waves on the field line passing through the

rocket. Figure 3 (c) shows precipitating electron fluxes with energies higher than 30 keV

observed by the 0-deg. telescope of POES18 / MEPED [*Evans & Greer*, 2000]. Here, an energy

channel with central energy of 40 keV (the blue line of Figure 3 (c)) corresponds to electrons

with energies higher than 30 keV. MEPED observed electron precipitation from 09:10:20 UT to

147 09:13:40 UT during which POES18 passed close to Andøya as indicated by the red line of Figure

148 3 (a). Precipitating electron fluxes with energies higher than 30 keV observed by MEPED can be

used as a proxy for the power of chorus waves occurring at the magnetically conjugate locations
 in the magnetosphere [*Chen et al.*, 2014]. Thus, the observed electron precipitation suggests that

chorus waves were present in the magnetospheric site magnetically connecting to the location of

152 POES18 from 09:10:20 UT to 09:13:40 UT. While POES18 passed close to Andøya at 1.8 to 1

minute before the start of the RockSat-XN / HEP observation, previous studies showed that

dayside chorus wave intensification persists for at least 1.5 hours in the same location [*Keika et al.*, 2012]. This suggests that the chorus wave was present simultaneously with the electron

precipitation over Andøya during the RockSat-XN / HEP observation. Note that POES18 had

moved to locations at 440 - 2200 km southwest from Andøya during the RockSat-XN / HEP

observation (9:14:39.6 UT - 9:19:14 UT), where the observed precipitating electron flux was

small. The observed decrease of the precipitating flux at the timing of the RockSat-XN / HEP

160 observation may be due to the spatial distribution of an active region of the chorus waves.

161

162 Figure 3 (d) shows the frequency-time spectrogram of magnetic field fluctuations obtained by

the VLF receiver at Lovozero in Russia. The latitudinal and longitudinal differences between

164 Andøya and Lovozero are 4° and 14° respectively in the geomagnetic coordinate system. The

hiss and bursty VLF emissions were observed from 9:10 UT to 9:20 UT, coinciding with the

166 RockSat-XN / HEP observations, and the daytime chorus wave was observed from 9:20 UT to

167 10:00 UT after the launch. It has been demonstrated that the dayside uniform magnetic field zone

can be a source region of dayside chorus waves especially under steady solar wind and quiet

geomagnetic conditions [*Keika et al.*, 2012]. The dayside uniform magnetic field zone (DFZ) is
 the transition region between the near-Earth dipole zone and the compressed, off-equatorial

the transition region between the near-Earth dipole zone and the compressed, off-equatorial
 double-minimum field configuration closer to the magnetopause (called the dayside outer zone)

172 [*Tsurutani et al.*, 2009]) and distributed over a wide MLT range on the dayside of the

magnetosphere. The presence of chorus waves at Lovozero may support the possible appearance

of chorus waves in the magnetosphere magnetically connected to Andøya.



- 177 Figure 3. (a) Trajectories of RockSat-XN and POES18, and the locations of Andøya, Lovozero,
- 178 Tromsø, and Longyearbyen. The color on the RockSat-XN trajectory indicates altitude. The time
- (UT) shown along the POES18 trajectory indicates the position of the satellite. The red line on
- POES18 orbit corresponds to the period when the electron precipitation is observed (9:10:20 UT to 9:13:40 UT). (b) Electron density around the launch measured by EISCAT VHF radar at
- to 9:13:40 UT). (b) Electron density around the launch measured by EISCAT VHF radar at
 Tromsø. The red line indicates the launch time, and the solid line part corresponds to the altitude
- observed by the RockSat-XN sounding rocket. (c) Fluxes of precipitating electrons observed by
- the 0-deg. telescope of POES18 / MEPED. Energy labels of the precipitating electrons are center
- energies of each energy channel of MEPED obtained using the bow tie method [*Green*, 2013b]
- and 40 keV (blue), 130 keV (green), 287 keV (red), and 612 keV (black). (d) Frequency-time
- 187 spectrogram of plasma waves observed at Lovozero. The interval between the two black lines
- 188 corresponds to the period of the RockSat-XN observation.
- 189
- 190 **Error! Reference source not found.** (a, black line) shows an energy spectrum of particles
- detected by HEP from 99.6 s to 374 s from the launch. Here the energy is the sum of deposited
- energies in all SSDs. The detected energy of incident particles ranges from ~50 keV to ~30 MeV
- with two peaks at 130 keV and 1.5 MeV. The raw observation data contain background noise
- such as Galactic Cosmic Rays (GCRs) and electrons originated from Cosmic Ray Albedo
- Neutron Decay (CRAND). Therefore, we have to evaluate and then remove these noises
- components from the data. In the next section, we describe how we remove the noises due toGCRs and CRAND.
- 198







- 202 The incident flux of GCR is calculated by the model of Lal [1985] ($\phi = 300$ MeV). (b) Setup of
- the numerical simulations for estimating the noise level due to GCR and CRAND. (b-1) GCR
- protons and alpha particles were injected isotropically from the region of polar angles from 0° to

²⁰⁵ 120° on a spherical surface $(3\pi \text{ sr})$. (b-2) A pencil beam of electrons is injected toward the center ²⁰⁶ of SSDs vertically for the simulation of CRAND. (c) The histogram of the number of SSD layer

- in which particles (black: the observation, red: the estimated GCR particles) deposit their
- energies more than 31 keV. Counts of the seventh SSD layer includes penetrating particles. (d)
- The energy spectra of (black) observed events, the estimated background (red) GCR protons and
- alpha particles, and (blue) CRAND electrons which deposit energies only on the first SSD layer.
- 211

212 **4 Noise reduction and estimation of precipitating electron energy spectra**

4.1 Noise reduction due to GCR

We investigated the energy spectra of galactic cosmic ray protons and alpha particles using the models of Castagnoli & Lal [1980], which have a parameter ϕ indicating the influence of

- models of Castagnon & Lat [1980], which have a parameter ϕ indicating the influence of modulation due to solar activity. To estimate the background noise due to GCR, we calculate the
- response of HEP to GCRs with the Geant4 toolkit [*Agostinelli et al.*, 2003; *Allison et al.*, 2006:

response of HEP to GCRs with the Geant4 toolkit [*Agostinelli et al.*, 2003; *Allison et al.*, 2006:
 Allison et al., 2016]. A schematic view of the simulation setup is shown in Figure 4 (b-1)Error!

Reference source not found. Here, GCR particles were assumed to be uniformly injected from

- an area with a solid angle of 3π sr (see Figure 4 (b-1)). Since GCRs are shielded by the earth, we
- assume that GCRs do not enter HEP from the bottom.
- 222 The green and blue lines in Figure 4 (a) show the estimated energy spectra of GCR protons and
- alpha particles detected by HEP during the observation period. These spectra are calculated by the numerical simulations with the model parameter ϕ of GCR as 300 MeV. The red line in
- the numerical simulations with the model parameter ϕ of GCR as 300 MeV. The red line in Figure 4 (a) shows the sum of the blue and green lines. As evident from the figure, the higher
- energy component (> 1 MeV) of the observed spectrum is consistent with the estimated noise
- spectrum due to GCR. On the other hand, the estimated GCR counts are significantly smaller
- than the observed particle counts in the energy range between 40 keV and several hundred keV.
- Figure 4 (c) shows histograms of the number of hit SSDs. Here, the hit threshold is set to 31 keV
- for energy deposit at each layer. To evaluate particles that seem to be appropriate as incidents
- from the collimator, these histograms includes only particles that sequentially hit from the first
- layer SSD. For example, we excluded particles that do not hit on the first layer and those which
- hit on the first and third layers but not on the second from these histograms. The excluded
- particles account for only 1.6 % of the total number observed. Most of the observed particles
- stop at the first SSD layer or reach the seventh layer (black line). The latter cases are consistent
- with that of estimated galactic cosmic ray protons and alpha particles (red line) since most of
- them deposit their energies on all the SSD layers and their energies are very high. However, the
- former feature cannot be explained by GCRs. In the following analysis, we deal with only
- 239 particles that stopped in the first SSD layer since the other components are consistent with the
- noise level due to GCRs. Note that we subtract the estimated proton and alpha particles counts
- 241 which deposit energies only on the first layer from observed particle counts.
- 242

243 4.2 Noise reduction due to CRAND

An electron flux due to CRAND is estimated by integrating the spatial distribution of the 244 electron production rate due to neutron decay [Hess et al., 1961] along the trajectory of electrons 245 with an equatorial pitch angle of 1.6° which corresponds to the average center of the field of 246 view of HEP for the dipole magnetic field line passing through Andøya (L= 6.77). Details of this 247 calculation are shown by Lencheck et al. [1961]. Here, we integrate the electron production rate 248 from the altitude of 100 km in the southern hemisphere to the altitude of 100 km in the northern 249 hemisphere. We consider the electrons in the loss cone. The noise due to CRAND electrons is 250 then estimated by a Geant4 calculation with the electron fluxes of the CRAND origin. The 251 simulation setup is briefly shown in Figure 4 (b-2)Error! Reference source not found.. We 252 consider CRAND electrons to enter HEP in the normal direction of SSDs. The blue line in 253 Figure 4 (d) shows the estimated energy spectrum observed by HEP due to CRAND, where we 254 select particles that deposited their energies more than 31 keV only on the first SSD layer. The 255 estimated counts due to CRAND are much smaller than those detected during the observation. 256

Finally, we evaluate the energy spectrum of electrons of the magnetospheric origin by

subtracting the GCR protons, alpha particles, and CRAND electron counts shown in Figure 4 (d)

from the observed counts. The electron flux can be measured up to the maximum value of the

260 energy bin for which significant counts remain after subtracting the GCRs and CRAND

electrons. The subtracted flux indicates that the observed particles with energies less than 500

keV may contain precipitating electrons generated by plasma processes in the magnetosphere.

263

4.3 Estimation of the energy spectrum of precipitating electrons

To obtain an energy spectrum of the precipitating electron flux, we need to consider instrumental 265 effects such as the aluminum shield located just in front of SSDs (see Figure 1), and electrons 266 which deposit their energies partially and run away from the first SSD layer. Figure 5 shows the 267 estimated probability distribution of detected energies at the first SSD layer as a function of 268 incident electron energy in the case that the particles deposit their energies only on the first SSD 269 layer, as calculated by the Geant4 toolkit. Incident electrons with monochromatic energy are 270 detected as electrons with energies lower than the incident energy following certain probability 271 distribution. The total detection efficiency is ≤ 10 % for ≤ 180 keV electrons and 1 % for ≤ 150 272 273 keV. It is thus difficult to discuss the incident energy spectrum of electrons below 180 keV. Electrons with 550 keV that enter HEP are most likely to be observed as 500 keV, which is the 274 maximum value of the electron energy detected in this observation, as discussed above. 275 Therefore, we consider that the effective energy range of the present observation is from 180 keV 276

277 to 550 keV.



Figure 5. The probability distribution of detected energies as a function of incident electron 280 energies deposited on the first SSD layer only. Uncertainty of detected energies coming from the 281 282 energy resolution of HEP is taken into account.

283

284 The black line in Figure 6 (a) shows the energy-spectrum of observed electrons where signals above 31 keV are detected only by the first SSD layer during 99.6 s to 374 s from the launch. 285 The number of the detected electrons in the *i*-th energy bin f_i (the index *i* denotes the energy bin 286 number) is given by a product of the energy spectrum of the incident electrons s_i (the index j 287 denotes the energy bin number of the incident electrons) and a matrix of the probability

288

distribution of the HEP detector response R_{ii} , as follows; 289

$$f_i = \sum_j R_{ij} s_j. \tag{1}$$

- The matrix element R_{ij} is the probability that an incident electron in the *j*-th energy bin is 290
- detected in the *i*-th energy bin. We assumed that s_i follows power-law form Ax^B , where x is 291
- incident energy, A and B are parameters to be estimated. We derived a coefficient A and 292
- exponent B of the power-law model that minimizes chi-square with Levenberg-Marquardt least-293 squares method [*Press et al.*, 1992]. Here, the chi-square is calculated as the sum of the squares 294 of the differences between the number of the electrons detected by HEP in the RockSat-XN
- 295 observation f_i^{obs} and that calculated from the model incident spectrum $f_i^{model} = \sum_i R_{ij} s_i$ 296
- divided by the measurement error of f_i^{obs} . 297
- The red line in Figure 6 (b) shows the estimated energy spectrum of incident electrons s_i that 298
- give the minimum chi-square, and the blue line in Figure 6 (a) shows f_i^{model} corresponding to 299
- the red line s_i . 300



Figure 6. (a) The energy spectrum obtained by removing GCRs and CRAND electrons from the particles deposited only on the first SSD layer f_i^{obs} (black) and the energy spectrum calculated from the incident electron spectrum and the HEP response matrix f_i^{model} (blue). (b) The estimated energy spectrum of incident electrons s_j that denotes $3.16 \times 10^{10} x^{-4.83}$ ($\chi^2 = 35.0$). (c) Energy spectra of incident electrons calculated every 50 seconds. Their fluctuations are ± 30 % and ± 60 % at 180 keV and 550 keV, respectively. Note that the black line shows the energy spectra averaged from 99.6 s to 374 s from launch (same as the red line in Figure 6 (b)).

310

Figure 6 (c) shows the energy spectra of incident electrons s_j estimated every 50 s, starting from

120s after launch. The black line in Figure 6 (c)**Error! Reference source not found.** is the

- estimated s_i averaged from 99.6 s to 374 s from launch (same as the red line in Figure 6 (b)).
- 314 These calculated electron fluxes are more or less stable in time, where fluctuations are \pm 30 %
- and \pm 60 % at 180 keV and 550 keV, respectively.

317 **5 Discussion**

In the RockSat-XN experiment, we have observed a significant number of sub-relativistic 318 electrons precipitating into the dayside polar ionosphere with energies of several hundred keV 319 under the quiet magnetospheric condition. A likely explanation for these electrons is pitch angle 320 scattering of electrons in the dayside magnetosphere due to wave-particle interactions with 321 322 chorus waves, which are evident from the complementary ground and satellite observations during the RockSat-XN observation (Figure 3 (c, d)). Chorus waves can interact with high-323 energy electrons when they propagate to high latitude because the resonance energy increases 324 with increasing the background magnetic field strength [Horne & Thorne, 2003, Miyoshi et al., 325 2010, 2015a]. The equatorial chorus wave amplitude distribution for AE \leq 100 nT is highest in 326 the dawn MLT sector (7–13 MLT) [Li et al., 2009] and these dayside chorus waves show little 327 328 dependence of occurrence on geomagnetic activity [Tsurutani & Smith, 1977; Spasojević & Inan, 2010]. Also, Bunch et al. [2012] showed that the time-averaged chorus wave power in the 329

- dayside outer magnetosphere can exceed 10 pT for latitudes up to 45° off the equator.
- 331

In addition to the present observations, we estimate the flux of sub-relativistic electron

precipitation by using a computer simulation about wave-particle interactions (GEMSIS-RBW *Saito et al.*, 2012]). In the simulation, electron precipitation is caused by pitch-angle scattering

[*Saito et al.*, 2012]). In the simulation, electron precipitation is caused by pitch-angle scattering of magnetospheric electrons by chorus waves. Synthesized chorus wave packets are given at the

- magnetic equator and propagate toward the high latitude along the dipole geomagnetic field. For
- the source magnetospheric electron population in this simulation, we used an energy pitch
- angle distribution of electron flux observed by the HEP instrument [*Mitani et al.*, 2018] on board the Arase satellite [*Miyoshi et al.*, 2018b]. The energy - pitch angle distribution, which is given
- for each satellite spin (~8 s), was averaged between 9:00 UT and 9:30 UT on January 13, 2019,
- during which the Arase satellite was located at $L^* = 5.57 5.50$ and MLT = 18.21 18.52. Here,
- L* is the value of L shell defined by Roederer [1970]. Although Arase was located on the
- duskside, the difference in L* between Andøya (L* = 6.23) and Arase locations are only 0.66 0.73. Note that the angular resolution of Arase / HEP is larger than the local loss cone angle of
- electrons at Arase. We thus used observed electron flux in the pitch angle bin covering 0 -

11.25° as a flux just outside the loss cone. In case of lack of data for that pitch angle bin, the flux

347 value was extrapolated from the other bins of the same energy by assuming that the pitch angle 348 distribution follows a cosine function. In this simulation, we modeled the rising tones of chorus

- distribution follows a cosine function. In this simulation, we modeled the rising tones of
 waves with an amplitude of magnetic fluctuations as 17 pT based on POES18 / MEPED
- observations [*Chen et al.*, 2014] and a frequency band of $0.2 0.5 |\Omega_{e}|$ based on the VLF
- measurement at Lovozero (Figure 3 (c)), where Ω_{ρ} is the electron cyclotron frequency. The

background magnetic field strength at the magnetic equator is set as 125 nT by using the TS05

model [*Tsyganenko and Sitnov.*, 2005], and the background electron density in the

magnetosphere is set as 3.78 cm^{-3} based on the model by Sheeley et al. [2001]. Furthermore, we

- assume that the chorus waves propagate up to the magnetic latitude of 40° along the field line, where the resonance energy of electrons with the chorus waves with a frequency of 0.3 $|\Omega_a|$ is
- 640 keV at L* = 6.23. The repetition period of each chorus element is set as 3 Hz which is a

typical modulation period of chorus waves [e.g., *Miyoshi et al.*, 2015a].

A simulated energy spectrum of precipitating electrons is shown with the red line in Figure 7. It

361 shows a similar spectrum profile with the RockSat-XN / HEP observation in the energy range

from 200 keV to 500 keV. The absolute values of electron flux by the GEMSIS-RBW simulation

shows good agreement with those of the RockSat-XN / HEP observation despite their different
 MLT positions. These results strongly support the interpretation that the sub-relativistic electron

precipitation observed by RockSat-XN is generated through pitch angle scattering of

366 magnetospheric electrons with the chorus waves.

367



368

Figure 7. Energy spectra of electrons precipitating into the ionosphere. The flux (blue) observed by RockSat-XN / HEP and (red) calculated by GEMSIS-RBW simulation using the electron measurement by Arase satellite. The black line shows the electron energy spectrum in the pitch angle bin covering $0 - 11.25^{\circ}$ of the energy - pitch angle distribution of electron flux observed by RockSat-XN / HEP and that calculated by GEMSIS-RBW simulation have a similar powerlaw index.

375

Both electron precipitations observed by POES18 / MEPED and RockSat-XN / HEP may be 376 related to dayside pulsating/diffuse auroral activity. POES18 / MEPED observed electron 377 precipitation from 72.6°N to 68.4°N in the geomagnetic coordinate system. On the other hand, 378 the dayside pulsating aurora was observed at Longyearbyen (Figure 3 (a)), located at 75.1°N in 379 the geomagnetic coordinate system until 8 UT. Auroral emission was not detected due to the 380 cloudy weather after 8 UT, but dayside pulsating aurora activity was not surprising to be present 381 during the RockSat-XN / HEP observation because the dayside pulsating aurora frequently 382 occurs at 8-13 MLT [Han et al., 2015]. Since the electron precipitation observed by POES18 383 continued up to the region close to Andøya, there may be also pulsating/diffuse auroral activity 384 at Andøya during the RockSat-XN / HEP observation. The speculation is also supported by the 385 continuous increase of electron density at altitudes between 80 and 110 km measured with the 386

387 EISCAT VHF radar in Tromsø. The speculation also suggests the possible relationship between

the dayside pulsating/diffuse auroral activity and chorus waves since both electron precipitations observed by POES18 and RockSat-XN may be related to dayside chorus waves.

390

The sub-relativistic electron microbursts occur with a peak occurrence rate at L = 5.5 and ~10

MLT on the dayside ionosphere at AE ≤ 100 nT [*Douma et al.*, 2017] and one of the possible

mechanisms for generating microburst is pitch angle scattering of electrons due to wave-particle

interactions [*Horne & Thorne*, 2003; *Hikishima et al.*, 2010]. Furthermore, their relationship to

³⁹⁵ pulsating auroras has been suggested [*Saito et al.*, 2012, *Miyoshi et al.*, 2015a, b]. These features

- may also be present in the sub-relativistic electrons observed by RockSat-XN / HEP. However,
 we were unable to discriminate the temporal characteristics of the microbursts less than 1 second
- because the observed particle count rates of RockSat-XN / HEP was not sufficient.

399

400 6 Conclusion

401 The energy spectrum of precipitating electrons in the sub-relativistic energy range, from 180 keV

to 550 keV, in the quiet dayside auroral ionosphere has been obtained by in-situ observation by

403 RockSat-XN / HEP for the first time. The energy spectrum of the observed precipitation follows

a power law of -4.86 and temporal variations of the fluxes are \pm 30 % and \pm 60 % at 180 keV and 550 keV, respectively.

406

A few minutes before the RockSat-XN observation, POES18 / MEPED observed precipitating 407 electrons related to the power of chorus waves at the location close to Andøya. The VLF receiver 408 observation at Lovozero also supports the presence of chorus waves during the RockSat-XN / 409 HEP observation. Furthermore, the GEMSIS-RBW simulation using Arase satellite data as input 410 parameters shows an energy spectrum of precipitating electrons in good agreement with that 411 observed by RockSat-XN. All of the observations and the simulation suggest that the 412 precipitation observed by RockSat-XN was caused by the wave-particle interactions between 413 chorus waves and sub-relativistic electrons and was also likely to be accompanied by dayside 414 pulsating/diffuse auroras. 415

416

417 Acknowledgments, Samples, and Data

418 We are grateful for the support of NASA Wallops Flight Facility's (WFF) and Andøya Space

419 Center (ASC). This research was financially supported by the Grants-in-Aid for Scientific

420 Research (15H05815, 15H05747, 16H06286, 17H00728, 18KK0100, 20H01959, 20H01955) by

the Ministry of Education, Science, Sports and Culture, Japan. This study was supported by JSPS

Bilateral Open Partnership Joint Research Projects. We thank Dr. Janet Green and the NOAA's

423 National Geophysical Data Center (NGDS) for providing NOAA data. The rocket and

424 instruments' data used in this study are available at the UTokyo Repository

(http://hdl.handle.net/2261/00079467). The data of the VLF receiver are available at the site

426 (http://aurora.pgia.ru/erg-pgi/case1.html). The AE index used in this paper was provided by the

427 WDC for Geomagnetism, Kyoto (http://wdc.kugi.kyoto-u.ac.jp/wdc/Sec3.html). The magnetic

- 428 coordinate system is calculated by the IGRF-13 model
- 429 (https://www.ngdc.noaa.gov/IAGA/vmod/igrf.html). EISCAT is an international association
- 430 supported by research organizations in China (CRIRP), Finland (SA), Japan (NIPR and ISEE),
- 431 Norway (NFR), Sweden (VR), and the United Kingdom (UKRI). The EISCAT data used in this
- 432 study are processed with the help of the EISCAT staff and published on the web
- 433 (http://pc115.seg20.nipr.ac.jp/www/eiscatdata/). Science data of the ERG (Arase) satellite were
- 434 obtained from the ERG Science Center (ERG-SC) operated by ISAS/JAXA and ISEE/Nagoya
- 435 University (https://ergsc.isee.nagoya-u.ac.jp/index.shtml.en, Miyoshi et al., 2018a). The present
- study analyzed MGF-L2 v03_03 data, HEP-L2 v03_01 data, and orbit L3 v02 data obtained by
- ERG. Part of the work of YM and TH was done at ERG-SC. The present study used SPEDAS
- 438 for data analysis [Angelopoulos et al., 2019]
- 439

440 **References**

- 441 Agostinelli, S., Allison, J., Amako, K., Apostolakis, J., Araujo, H., Arce, P., et al. (2003).
- 442 GEANT4 A simulation toolkit. *Nuclear Instruments and Methods in Physics Research*,
- 443 Section A: Accelerators, Spectrometers, Detectors and Associated Equipment, 506(3), 250–
 444 303. https://doi.org/10.1016/S0168-9002(03)01368-8
- Allison, J., Amako, K., Apostolakis, J., Araujo, H., Dubois, P. A., Asai, M., et al. (2006). Geant4
 developments and applications. *IEEE Transactions on Nuclear Science*, *53*(1), 270–278.
 https://doi.org/10.1109/TNS.2006.869826
- Allison, J., Amako, K., Apostolakis, J., Arce, P., Asai, M., Aso, T., et al. (2016). Recent
- developments in GEANT4. Nuclear Instruments and Methods in Physics Research, Section
- 450 *A: Accelerators, Spectrometers, Detectors and Associated Equipment, 835, 186–225.*
- 451 https://doi.org/10.1016/j.nima.2016.06.125
- Angelopoulos, V., Cruce, P., Drozdov, A., Grimes, E. W., Hatzigeorgiu, N., King, D. A., et al.
 (2019). The Space Physics Environment Data Analysis System (SPEDAS). Space Science Reviews (Vol. 215). The Author(s). https://doi.org/10.1007/s11214-018-0576-4
- Blake, J. B., Looper, M. D., Baker, D. N., Nakamura, R., Klecker, B., & Hovestadt, D. (1996).
 New high temporal and spatial resolution measurements by SAMPEX of the precipitation of relativistic electrons. *Advances in Space Research*, *18*(8), 171–186.
- 458 https://doi.org/10.1016/0273-1177(95)00969-8
- Bunch, N. L., Spasojevic, M., & Shprits, Y. Y. (2012). Off-equatorial chorus occurrence and
 wave amplitude distributions as observed by the Polar Plasma Wave Instrument. *Journal of Geophysical Research: Space Physics*, *117*(4), 1–14. https://doi.org/10.1029/2011JA017228
- Castagnoli, G., & Lal, D. (1980). Solar Modulation Effects in Terrestrial Production of Carbon 14. *Radiocarbon*, 22(2), 133–158. https://doi.org/10.1017/s0033822200009413

464 465 466	Chen, Y., Reeves, G. D., Friedel, R. H. W., & Cunningham, G. S. (2014). Global time-dependent chorus maps from low-Earth-orbit electron precipitation and Van Allen Probes data. <i>Geophysical Research Letters</i> , 41(3), 755–761. https://doi.org/10.1002/2013GL059181
467	Clilverd, M. A., Rodger, C. J., & Ulich, T. (2006). The importance of atmospheric precipitation
468	in storm-time relativistic electron flux drop outs. <i>Geophysical Research Letters</i> , <i>33</i> (1), 1–5.
469	https://doi.org/10.1029/2005GL024661
470	Dietrich, S., Rodger, C. J., Clilverd, M. A., Bortnik, J., & Raita, T. (2010). Relativistic
471	microburst storm characteristics: Combined satellite and ground-based observations.
472	<i>Journal of Geophysical Research: Space Physics</i> , 115(12), 1–10.
473	https://doi.org/10.1029/2010JA015777
474	Douma, E., Rodger, C. J., Blum, L. W., & Clilverd, M. A. (2017). Occurrence characteristics of
475	relativistic electron microbursts from SAMPEX observations. <i>Journal of Geophysical</i>
476	Research: Space Physics, 122(8), 8096–8107. https://doi.org/10.1002/2017JA024067
477	Evans, D. S., & Greer, M. S. (2000). Polar orbiting environmental satellite space environment
478	monitor - 2: Instrument descriptions and archive data documentation, <i>NOAA Technical</i>
479	<i>Memorandum</i> , Boulder, Colorado OAR SEC 93, 93, version 1.4, January 2004.
480	Green, J. C. (2013a). External users manual, POES/MetOp SEM-2 processing, version 1.0, 20.,
481	NOAA National Geophysical Data Center. Available at
482	http://www.ngdc.noaa.gov/stp/satellite/poes/documentation.html.
483	Green, J. C. (2013b). MEPED Telescope Data Processing Algorithm Theoretical Basis
484	Document, version 1.0, 77., NOAA National Geophysical Data Center. Available at
485	http://www.ngdc.noaa.gov/stp/satellite/poes/documentation.html.
486	Han, D., Chen, X., Liu, J., Qiu, Q., Keika, K., Hu, Z., et al. (2015). An extensive survey of
487	dayside diffuse aurora based on optical observations at Yellow River Station. <i>Journal of</i>
488	<i>Geophysical Research: Space Physics</i> , 120(9), 7447–7465.
489	https://doi.org/10.1002/2015JA021699
490	Hess, W. N., Canfield, E. H., & Lingenfelter, R. E. (1961). Cosmic-ray neutron demography.
491	<i>Journal of Geophysical Research</i> , 66(3), 665–677. https://doi.org/10.1029/jz066i003p00665
492 493 494	 Hikishima, M., Omura, Y., & Summers, D. (2010). Microburst precipitation of energetic electrons associated with chorus wave generation. <i>Geophysical Research Letters</i>, 37(7), 1–5. https://doi.org/10.1029/2010GL042678
495 496 497	 Horne, R. B., & Thorne, R. M. (2003). Relativistic electron acceleration and precipitation during resonant interactions with whistler-mode chorus. <i>Geophysical Research Letters</i>, 30(10), 3–6. https://doi.org/10.1029/2003gl016973
498 499	Keika, K., Spasojevic, M., Li, W., Bortnik, J., Miyoshi, Y., & Angelopoulos, V. (2012). PENGU In/AGO and THEMIS conjugate observations of whistler mode chorus waves in the dayside

500 501	uniform zone under steady solar wind and quiet geomagnetic conditions. <i>Journal of Geophysical Research: Space Physics</i> , <i>117</i> (7), 1–15. https://doi.org/10.1029/2012JA017708
502 503	Kennel, C. F., & Petschek, H. E. (1966). Limit on stably trapped particle fluxes. Journal of Geophysical Research, 71(1), 1–28. https://doi.org/10.1029/jz071i001p00001
504 505 506 507	Kurita, S., Kadokura, A., Miyoshi, Y., Morioka, A., Sato, Y. & Misawa, H. (2015), Relativistic electron precipitations in association with diffuse aurora: Conjugate observation of SAMPEX and the all-sky TV camera at Syowa Station, <i>Geophysical Research Letters</i> , 42(12), 4702–4708. https://doi.org/10.1002/2015GL064564
508 509 510 511	 Kurita, S., Miyoshi, Y., Blake, J. B., Reeves, G. D., & Kletzing, C. A. (2016). Relativistic electron microbursts and variations in trapped MeV electron fluxes during the 8-9 October 2012 storm: SAMPEX and Van Allen Probes observations. <i>Geophysical Research Letters</i>, 43(7), 3017–3025. https://doi.org/10.1002/2016GL068260
512 513 514 515	Lal, D. (1985). Theoretically Expected Variations in the Terrestrial Cosmic-Ray Production Rates of Isotopes, <i>Proc. of the Inter. School of Physics</i> , Course XCV, Solar-Terrestrial Relationships and the Earth Environment in the last Millennia, G. C. Castagnoli ed., 216- 233.
516 517 518 519	Lam, M. M., Horne, R. B., Meredith, N. P., Glauert, S. A., Moffat-Griffin, T., & Green, J. C. (2010). Origin of energetic electron precipitation >30 keV into the atmosphere. <i>Journal of Geophysical Research A: Space Physics</i> , <i>115</i> (A4), 1–15. https://doi.org/10.1029/2009JA014619
520 521 522	Lenchek, A. M., Singer, S. F., & Wentworth, R. C. (1961). Geomagnetically trapped electrons from cosmic ray albedo neutrons. <i>Journal of Geophysical Research</i> , 66(12), 4027–4046. https://doi.org/10.1029/JZ066i012p04027
523 524 525	Li, W., Thome, R. M., Angelopoulos, V., Bortnik, J., Cully, C. M., Ni, B., et al. (2009). Global distribution of whistler-mode chorus waves observed on the THEMIS spacecraft. <i>Geophysical Research Letters</i> , 36(9), 1–5. https://doi.org/10.1029/2009GL037595
526 527 528	Lorentzen, K. R., Looper, M. D., & Blake, J. B. (2001). Relativistic electron microbursts during the GEM storms. <i>Geophysical Research Letters</i> , 28(13), 2573–2576. https://doi.org/10.1029/2001GL012926
529 530 531	Matsuoka, A., Teramoto, M., Nomura, R., Nosé, M., Fujimoto, A., Tanaka, Y., et al. (2018). The ARASE (ERG) magnetic field investigation. <i>Earth, Planets and Space</i> , 70(1), 1–16. https://doi.org/10.1186/s40623-018-0800-1
532 533 534	Mitani, T., Takashima, T., Kasahara, S., Miyake, W., & Hirahara, M. (2018). High-energy electron experiments (HEP) aboard the ERG (Arase) satellite. <i>Earth, Planets and Space</i> , 70(1), 1–14. https://doi.org/10.1186/s40623-018-0853-1
535 536	Miyoshi, Yoshizumi, Katoh, Y., Nishiyama, T., Sakanoi, T., Asamura, K., & Hirahara, M. (2010). Time of flight analysis of pulsating aurora electrons, considering wave-particle

- interactions with propagating whistler mode waves. *Journal of Geophysical Research: Space Physics*, *115*(10), 1–7. https://doi.org/10.1029/2009JA015127
- Miyoshi, Y., Oyama, S., Saito, S., Kurita, S., Fujiwara, H., Kataoka, R., et al. (2015a). Energetic
 electron precipitation associated with pulsating aurora: EISCAT and Van Allen Probe
 observations. *Journal of Geophysical Research: Space Physics*, *120*(4), 2754–2766.
 https://doi.org/10.1002/2014JA020690
- Miyoshi, Y., Saito, S., Seki, K., Nishiyama, T., et al., (2015b). Relation between fine structure of
 energy spectra for pulsating aurora electrons and frequency spectra of whistler mode chorus
 waves. *Journal of Geophysical Research: Space Physics*, *120*(9), 7728-7736.
 https://doi.org/10.1002/2015JA021562
- Miyoshi, Yoshizumi, Hori, T., Shoji, M., Teramoto, M., Chang, T. F., Segawa, T., et al. (2018a).
 The ERG Science Center. *Earth, Planets and Space*, 70(1). https://doi.org/10.1186/s40623018-0867-8
- Miyoshi, Yoshizumi, Shinohara, I., Takashima, T., Asamura, K., Higashio, N., Mitani, T., et al.
 (2018b). Geospace exploration project ERG. *Earth, Planets and Space*, 70(1).
 https://doi.org/10.1186/s40623-018-0862-0
- Newell, P. T., Sotirelis, T., & Wing, S. (2009). Diffuse, monoenergetic, and broadband aurora:
 The global precipitation budget. *Journal of Geophysical Research: Space Physics*, *114*(9),
 1–20. https://doi.org/10.1029/2009JA014326
- Ni, B., Thorne, R. M., Shprits, Y. Y., Orlova, K. G., & Meredith, N. P. (2011a). Chorus-driven
 resonant scattering of diffuse auroral electrons in nondipolar magnetic fields. *Journal of Geophysical Research: Space Physics*, *116*(6), 1–12. https://doi.org/10.1029/2011JA016453
- Ni, B., Thorne, R., Liang, J., Angelopoulos, V., Cully, C., Li, W., et al. (2011b). Global
 distribution of electrostatic electron cyclotron harmonic waves observed on THEMIS.
 Geophysical Research Letters, 38(17), 4–8. https://doi.org/10.1029/2011GL048793
- Ni, B., Li, W., Thorne, R. M., Bortnik, J., Ma, Q., Chen, L., et al. (2014). Resonant scattering of
 energetic electrons by unusual low-frequency hiss. *Geophysical Research Letters*, 41(6),
 1854–1861. https://doi.org/10.1002/2014GL059389
- Nishimura, Y., Bortnik, J., Li, W., Thorne, R. M., Ni, B., Lyons, L. R., et al. (2013). Structures
 of dayside whistler-mode waves deduced from conjugate diffuse aurora. *Journal of Geophysical Research: Space Physics*, *118*(2), 664–673.
 https://doi.org/10.1029/2012JA018242
- Press, W. H., Teukolsky, S. A., Vetterling, W. T., Flannery, B. P. (1992). Numerical Recipes in
 C: The Art of Scientific Computing, Second Edition. *Cambridge: Cambridge University Press*
- Roederer, J. G. (1970). Dynamics of Geomagnetically Trapped Radiation. *New York: Springer*.
 https://doi.org/10.1007/978-3-642-49300-3

Saito, S., Miyoshi, Y., & Seki, K. (2012). Relativistic electron microbursts associated with 574 575 whistler chorus rising tone elements: GEMSIS-RBW simulations. Journal of Geophysical Research: Space Physics, 117(10), 1-9. https://doi.org/10.1029/2012JA018020 576 Sandholt, P. E., Denig, W. F., Farrugia, C. J., Lybekk, B., & Trondsen, E. (2002). Auroral 577 structure at the cusp equatorward boundary: Relationship with the electron edge of low-578 latitude boundary layer precipitation. Journal of Geophysical Research: Space Physics, 579 107(A9), 1-9. https://doi.org/10.1029/2001JA005081 580 Sheeley, B. W., Moldwin, M. B., Rassoul, H. K., & Anderson, R. R. (2001). An empirical 581 582 plasmasphere and trough density model: CRRES observations. Journal of Geophysical Research: Space Physics, 106(A11), 25631–25641. https://doi.org/10.1029/2000JA000286 583 Shi, R., Han, D., Ni, B., Hu, Z. J., Zhou, C., & Gu, X. (2012). Intensification of dayside diffuse 584 auroral precipitation: Contribution of dayside Whistler-mode chorus waves in realistic 585 magnetic fields. Annales Geophysicae, 30(9), 1297-1307. https://doi.org/10.5194/angeo-30-586 1297-2012 587 Spasojevic, M., & Inan, U. S. (2010). Drivers of chorus in the outer dayside magnetosphere. 588 *Journal of Geophysical Research A: Space Physics*, 115(A4), 1–12. 589 https://doi.org/10.1029/2009JA014452 590 Sugo, S., Kawashima, O., Kasahara, S., Asamura, K., Nomura, R., Miyoshi, Y., Energy-resolved 591 detection of precipitating electrons of 30-100 keV by a sounding rocket associated with 592 593 dayside chorus waves. Earth and Space Science Open Archive, https://doi.org/10.1002/essoar.10503830.1 594 Thébault, E., Finlay, C. C., Beggan, C. D., Alken, P., Aubert, J., Barrois, O., et al. (2015). 595 International geomagnetic reference field: The 12th generation international geomagnetic 596 reference field - The twelfth generation. Earth, Planets and Space, 67(1). 597 https://doi.org/10.1186/s40623-015-0228-9 598 Thorne, R. M., O'Brien, T. P., Shprits, Y. Y., Summers, D., & Horne, R. B. (2005). Timescale 599 for MeV electron microburst loss during geomagnetic storms. Journal of Geophysical 600 Research: Space Physics, 110(A9), 1–7. https://doi.org/10.1029/2004JA010882 601 Tsurutani, B. T., & Smith, E. J. (1977). Two types of magnetospheric ELF chorus and their 602 substorm dependences. Journal of Geophysical Research, 82(32), 5112–5128. 603 604 https://doi.org/10.1029/ja082i032p05112 605 Tsurutani, B. T., Verkhoglyadova, O. P., Lakhina, G. S., & Yagitani, S. (2009). Properties of dayside outer zone chorus during HILDCAA events: Loss of energetic electrons. Journal of 606 607 Geophysical Research: Space Physics, 114(3), 1–19. https://doi.org/10.1029/2008JA013353 Tsyganenko, N. A., & Sitnov, M. I. (2005). Modeling the dynamics of the inner magnetosphere 608 during strong geomagnetic storms. Journal of Geophysical Research: Space Physics, 609 110(A3), 1–16. https://doi.org/10.1029/2004JA010798 610 611

Figure 1.



Figure 2.



Figure 3.



Figure 4.



Figure 5.



Figure 6.



Figure 7.

