

BURST GEOMAGNETIC PULSATIONS AS INDICATORS OF SUBSTORM EXPANSION ONSETS DURING SUPERSTORMS

Battuulai Tsegmed¹, Vladimir Mishin², Julia Klibanova³, and Marina Kurikalova⁴

¹Institute of Astronomy and Geophysics, Mongolian Academy of Sciences

²Institute of Solar-Terrestrial Physics Siberian Branch of Russian Academy of Sciences

³A. A. Ezhevsky Federal State Budgetary Educational Institute of Higher Education
Irkutsk State Agrarian University

⁴Institute of Solar-Terrestrial Physics of Siberian Branch of Russian Academy of Sciences,
Irkutsk

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Abstract

We report on the dynamics of field-aligned currents (FACs), broadband geomagnetic pulsations, and airglow obtained from the Irkutsk (IRK), Mondy (MND), and Borok (BOX) midlatitude geomagnetic observatories and the Tory (TOR) optical Observatory during superstorm substorms. For the first time, using the short duration, $[?]t < 0.5$ min, high-frequency component of the burst pulsations (Pi1B), we determined the substorm double expansion phase (EP) onsets < 5 min apart, which is hardly possible by means of the low frequency (periods of 2–5 min) Psc/PiB pulsations. We argue that the observed burst pulsations are the result of prompt changes in the solar wind dynamic pressure and/or the current circuit related to the westward electrojet. Each pulsed source can excite short bursts of broadband electromagnetic modes of the ionospheric Alfvén resonator in the range of short-period pulsations with a periodic resonance structure of the spectrum characteristic of the observed Pi1B/Psc pulsations

1 Vladimir Vilenovich Mishin, ORCID: 0000-0002-2729-2862, vladm@iszf.irk.ru

2 Battuulai Tsegmed, ORCID: 0000-0002-4828-8424, tseg@iag.ac.mn

3 Yuliya Yurevna Klibanova, ORCID: 0000-0001-7151-3629, malozemova81@mail.ru

4 Marina Aleksandrovna Kurikalova, ORCID: 0000-0002-0710-8917, kurikalova@iszf.irk.ru

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6 **BURST GEOMAGNETIC PULSATIONS AS INDICATORS OF SUBSTORM**
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9 V.V. Mishin¹, B. Tsegmed², Yu.Yu. Klibanova³, M.A. Kurikalova¹

10 ¹Institute of Solar-Terrestrial Physics of Siberian Branch of Russian Academy of Sciences,
11 Irkutsk, Russia (ISTP SBRAS).

12 ²Institute of Astronomy and Geophysics, Mongolian Academy of Sciences, Ulaanbaatar,
13 Mongolia.

14 ³A. A. Ezhevsky Federal State Budgetary Educational Institute of Higher Education Irkutsk
15 State Agrarian University, Irkutsk, Russia.

16
17 **Corresponding author:** Tsegmed Battuulai, Email: tseg@iag.ac.mn

18 **Postal address:** Institute of Astronomy and Geophysics, Mongolian Academy of Sciences,
19 Ulaanbaatar-51, POB-152, Mongolia 13343

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24 pulsations, and airglow obtained from the Irkutsk (IRK), Mondy (MND), and Borok (BOX)
25 midlatitude geomagnetic observatories and the Tory (TOR) optical Observatory during superstorm
26 substorms. For the first time, using the short duration, $\Delta t < 0.5$ min, high-frequency component of
27 the burst pulsations (Pi1B), we determined the substorm double expansion phase (EP) onsets <5 min
28 apart, which is hardly possible by means of the low frequency (periods of 2–5 min) Psc/PiB
29 pulsations. We argue that the observed burst pulsations are the result of prompt changes in the solar
30 wind dynamic pressure and/or the current circuit related to the westward electrojet. Each pulsed
31 source can excite short bursts of broadband electromagnetic modes of the ionospheric Alfvén

32 resonator in the range of short-period pulsations with a periodic resonance structure of the spectrum
33 characteristic of the observed Pi1B/Psc pulsations.

34 **Key points**

35 **We explored** the dynamics of field-aligned currents, geomagnetic burst pulsations, and
36 airglow during superstorm substorms.

37 Short duration of the high frequency part of Pi1B pulsation trains allows us to determine the
38 double expansion onsets.

39 **We suggest that prompt changes in the solar wind pressure or/and the westward electrojet**
40 **generate the two types of the observed pulsations**

41 **Key words**

42 substorm, expansion onset, superstorm, field-aligned currents, burst pulsations, auroral
43 emissions

44 **Plain Language Summary**

45 **We explored geomagnetic and optical midlatitude observations during superstorms. In**
46 **addition to the common low frequency (periods of 2–5 min) Psc/PiB pulsations, we used the short**
47 **period part (Pi1B) of burst pulsations with $T < 10$ s. That allowed us to determine not only isolated**
48 **substorm expansion onsets but also double substorm onsets and series of onsets of short substorm**
49 **activations or pseudobreakups during storms. This is hardly possible with the commonly used long-**
50 **period pulsations (Pi2, Pi3) is problematic due to their long duration**

52 **1. Introduction**

53 **Magnetic storms are a global phenomenon that greatly disturbs the space environment near**
54 **the Earth. A typical major storm lasts for several days. During this disturbed period, occurring**
55 **intermittently magnetospheric substorms create abrupt intensifications of optical (auroral) emissions**
56 **and electric fields and currents in the high- and mid-latitude ionosphere (Akasofu, 1964; Tinsley,**
57 **1986; Brunelli, 1988; Rassoul et al., 1993; Mikhalev, 2013; V.V. Mishin et al., 2018, Klibanova et**
58 **al., 2019). A typical substorm comprises the growth phase, explosive (active or expansion) phase,**
59 **and recovery phase (V.M. Mishin et al., 1979; Bazarzhapov et al., 1979; McPherron, 1979). The late**
60 **growth phase sometimes features "pseudobreakups", i.e., localized auroral disturbances with a train**
61 **of irregular burst pulsations but without the global dipolarization and explosive phase development.**
62 **Such disturbances may be related to solar wind (SW) sudden impulses.**

63 **Unlike the latter, storm sudden commencement (SSC) impulses are followed by storms and**

64 substorms (Nishida, 1978). In the substorm theory, a close focus is on the onset of the substorm
65 expansion phase (EP). The EP onset or simply onset, also known as substorm breakup, features the
66 dipolarization, the geomagnetic tail contraction related to the release of the magnetic energy
67 accumulated during the growth phase, and the fast decrease in the polar cap magnetic flux. These
68 processes result in rapidly growing field-aligned currents (FACs) and the intensified auroral activity
69 (V.M. Mishin et al., 2017; Stephens et al., 2019). The EP onset is normally determined by observing
70 the start of the geomagnetic pulsation (Pi2) train at high and mid latitudes, as well as by the auroral
71 intensification (Troitskaya & Guglielmi, 1967; Pudovkin, 1976; Kangas et al., 1998). In addition,
72 broadband burst pulsations or PiBs, with periods $T = (0.2\text{--}600)$ s, intensify during the onset. These
73 pulsations involve a short-period part, Pi1B with $T = (0.3\text{--}40)$ s, and a long-period part, Pi2-Pi3 with
74 $T = (40\text{--}600)$ s (see Table S1 in the Supporting Information). PiBs correlate well with the auroral
75 intensity, X-ray bursts, and ionospheric absorption of space radio noise (Troitskaya, 1961; Heacock,
76 1967; Undiedt et al, 1978; Bösinger et al., 1981; Bösinger & Yakhnin, 1987; Kangas et al., 1998).

77 During SSCs followed by the storm commencement, burst type pulsations named Psc are
78 generated. They also have a spiky character, the broadband spectrum between 0.2 and 600 s, and the
79 short-period, Psc1,2,3 (0.3–40) s, and long-period, Psc 4,5 (40–600) s, parts (Saito, 1969; Nishida,
80 1978, see also Table S1). In order to explore the PiB/Psc short-period component, one needs high-
81 frequency resolution data ($\Delta f \geq 10$ Hz, $\Delta t \leq 0.1$ s) whereas most observatories employ
82 magnetometers with a 1-10 s sampling rate. The latter allows studying only longer period Pi2-Pi3
83 pulsations capable of determining a single EP onset.

84 However, substorms with double EP onsets are not seldom (Russell, 2000; V.M. Mishin et
85 al., 2001). Those are not resolved in the data of pulsations with periods about and greater than the
86 time interval, τ , between the two onsets. For example, at the start of the 27 August 2001 substorm
87 (Baker et al., 2002; V.M. Mishin et al., 2013, V.M. Mishin et al., 2017), this interval was $\tau = 2$ min.
88 The onsets could be specified using the short period ($T < 10$ s) component of Pi1B with a shorter
89 duration. It is worth to note concerning the short period component of Pi1B, that the long period (10
90 –60 s) counterpart has the physical properties similar to Pi2 pulsations (Rae et al., 2011).

91 In this paper, we analyze the substorm events during the 6 April 2000 and 21 October 2001
92 superstorms, when the equatorward boundary of the auroral zone was near the city of Irkutsk (CGM:
93 $47.33^\circ \Phi$, $177.24^\circ \Lambda$). The effects of substorms were recorded both in geomagnetic and optical data.
94 For stormtime substorms with multiple successive intensifications of the AE (AL) index, the model
95 of an isolated substorm is hardly justified. Rather, quasi-periodic saw-tooth disturbances appear
96 during major storms (Troshichev et al., 2011). The duration of such disturbances may be short,

97 about 15–40 min (V.V. Mishin & Karavaev, 2017). During the 6 April 2000 storm, we also detected
98 a substorm with the growth phase started prior to the SSC at 16:10 UT. The *AE-index*
99 *intensifications observed later in the course of the storm are referred to as substorm activations.*

100 **2. Database**

101 During the 6 April 2000 and 21 October 2001 storms, we collected geomagnetic data from
102 Mondy (MND; CGM: 47.5° Φ , 177.5° Λ) and Borok (BOX; CGM: 53.9° Φ , 114° Λ) with search-
103 coil magnetometers at a $\Delta f = 10$ Hz frequency resolution. During the above events, both stations
104 were located on the night side. In addition, thanks to the Kyoto data center ([http://wdc.kugi.kyoto-
106 u.ac.jp](http://wdc.kugi.kyoto-
105 u.ac.jp)), we obtained fluxgate magnetometer data from the low-latitude Kakioka (KAK) station
107 (CGM: 29.25° Φ , 211.7° Λ) with a 1 Hz frequency resolution. Airglow at 557.7 and 630 nm was
108 observed by zenith photometers applying interference oscillating light filters ($\Delta\lambda$ 1/2 ~1–2 nm, V.V.
109 Mishin et al., 2018). The Tory geophysical observatory (TOR) of the ISTP SB RAS is located to the
110 south of Lake Baikal (CGM: 47° Φ , 177° Λ , MLT = UT + 7) about 75 km north-west from Mondy.

111 We employ the magnetogram inversion technique (MIT) developed at the ISTP SB RAS
112 more than 40 years ago (Bazarzhapov et al., 1979; V.M. Mishin et al., 1979; V.M. Mishin, 1990)
113 and has been upgraded (Lunyshkin & Pensikh, 2019). The MIT method uses the dipolar
114 geomagnetic coordinates: geomagnetic latitude, Φ , and local magnetic time (MLT). Using 1-min
115 data from the network of ground-based magnetometers, we obtained a sequence of Φ -MLT maps of
116 the field-aligned current (FAC) distribution in the ionosphere.

117 These maps identify the boundaries of the Iijima and Potemra (1978) FAC regions and the
118 values of the magnetic flux, Ψ , from the solar wind through the tail lobes into the magnetosphere
119 (see the definition of Ψ in Section 3). Accordingly, we determine the main onset of the substorm EP
120 at the start of the abrupt decrease of Ψ . The difference between the geomagnetic latitude in the
121 dipole system, Φ , and the latitude Φ_{cor} in the corrected system of geomagnetic coordinates (CGM)
122 used in satellite data is insignificant in the polar cap, but increases equatorward. For Irkutsk, it
123 yields: $\Delta\Phi = \Phi_{\text{cor}} - \Phi \approx 48^\circ - 42^\circ \approx 6^\circ$.

124 **3. Observations of geomagnetic disturbances and airglow**

125 **3.1. The 6–7 April 2000 superstorm**

126 The SSC on 6 April 2000 was observed at 16:40 UT. At this time, the solar wind ram
127 pressure (P_d) increased from 1 to 14 nPa, the interplanetary magnetic field (IMF) B_z turned

128 southward to -6 nT, and the IMF B_y component increased to ≈ 20 nT (Figure 1). The AE -index
129 increased from 250 nT to 1250 nT; and the D_{st} -index reached its minimum (-287 nT) near 23 UT
130 (V. M. Mishin et al., 2010a; V.V. Mishin et al., 2013; V.V. Mishin & Karavaev, 2017).

131 On can see in the H -component variations at Irkutsk a few bays caused by substorm
132 activations. We will mainly focus on two strongest bays during $\sim 16:40 - 18:10$ UT and $19:55 -$
133 $20:55$ UT indicated in Figure 1 by the red vertical lines. According to the Shue et al, (1998) model,
134 the enhanced SW pressure forced the subsolar magnetopause to $x = 7 R_E$. Then, the first substorm
135 EP onset occurred, as indicated by the sharply increased AE index and, as typical for isolated
136 substorms, by an abnormal increase of the variable part of the magnetic flux, Ψ_1 . We denote the
137 difference, $\Psi - \Psi_0$, between the total polar cap magnetic flux, Ψ , and the pre-substorm value, Ψ_0 , as
138 Ψ_1 . The total magnetic flux is $\Psi = \int \mathbf{B}(\mathbf{r}) \cdot d\mathbf{S}$, where $\mathbf{B}(\mathbf{r})$ is the dipolar geomagnetic field at 115 km
139 and S is the polar cap (R_0) area (V.M. Mishin et al., 2001, 2017).

140 The first EP onset occurred at the end of the growth phase, which started before the SSC
141 (V.M. Mishin et al., 2010a, 2010b). We infer that the polar cap expansion indicated by the increase
142 of Ψ_1 is a direct consequence of the magnetosphere compression. The second EP onset at $\sim 16:46$
143 UT (as well as the second substorm activation at $\sim 19:55$ UT) is indicated by sharply increasing AE
144 and decreasing Ψ_1 due to the magnetotail contraction (dipolarization) after reconnection. The SSC
145 and EP onsets are accompanied by burst pulsations (green arrows in Figure 1).

146 During $16:40 - 21$ UT, MND and BOX recorded several short bursts of broadband irregular
147 Psc/PiB pulsations, whereas TOR recorded bright emissions at 557.7 and 630 nm with intensities of
148 $\sim 100-2500$ R after the SSC front arrival (Figure 2). Near $16:40$ UT, one can see two double
149 substorm activations. Images from the POLAR satellite (Supplemental Information, Figure S2) also
150 show a weak auroral activation at $16:39$ UT and then the real substorm breakup at $16:46$ UT in the
151 Northern Hemisphere (MLAT $\sim 40^\circ - 90^\circ$).

152 Figure 3 shows the Psc/PiB bursts recorded at the mid-latitude MND and BOX and the low-
153 latitude KAK stations that indicate the double onset of the first substorm activation during the SSC.
154 The first Psc_{1,2} burst (the first onset) at $16:39:40$ UT was recorded during the SSC at periods $T > 0.4$
155 s at MND and BOX (with a smaller amplitude) and at KAK ($T > 7$ s). The PiB second burst with $T > 4$
156 s (the second onset) was concurrently recorded by all stations at $16:45:40$ UT with the greatest
157 amplitude at MND.

158 Figure 4 shows isocontours of the downward (upward) FAC density in the Region 1 (R1),
159 Region 2 (R2), and the polar cap – Region 0 (R0). During the SSC, the observatories were $10-15$
160 degrees equatorward of the FAC R2 southern boundary, which descended to the stations in ~ 5 min

161 and remained there for the EP interval. Although TOR and MND were distant from the westward
162 electrojet (WEJ) maximum, they detected the related intensifications of the green line and burst
163 pulsations.

164 At 17:10 UT and 18:08 UT, when BOX was near the southern FAC R2+ boundary and MND
165 was ten degrees to the south from the FAC R2- boundary, the *AE* index sharply intensified. This
166 might be consequent to fast variations in the SW pressure and IMF B_z (Figure 1). The increase in the
167 *AE* index to 2600 nT was coincident with the increased FACs in the R1, R2 regions to $I_{R2-} = 6$ MA,
168 $I_{R1+} = 8$ MA (Figures 1, 4). These two moments were coincident with the intensifications of
169 pulsations at MND and BOX, and 557.7 nm emissions at TOR on the night side (Figures 2, 5).

170 After 19:00 UT, the observatories were again within the FAC R2 region that caused a strong
171 increase of the airglow intensities (Figure 2). A similar behavior was observed during the 20
172 November 2003 superstorm (V.V. Mishin et al., 2018). During the second substorm activation
173 interval (19:55–20:55 UT), MND and BOX simultaneously recorded two PiB bursts with $T > 4$ s at
174 20:00 and 20:13 UT. Their amplitudes were 4 times greater at BOX (Figure 6). At that time, BOX
175 was at the R1/R2 boundary near the WEJ maximum and the focus of the vortex of the upward FAC
176 R2, which is the region of the maximum electron precipitation (Figure 7). Both bursts in the short
177 period range (0.5–10 s) at 19:52 UT and 21:05 UT were detected only at MND.

178

179 **3.2. The 21 October 2001 superstorm**

180 Figure 8 shows a moderate activity ($AE \sim 300\text{--}400$ nT) and slightly enhanced green line
181 emissions (~ 600 R) before the SSC at 16:48 UT. At the SSC, P_d increased from 2 nP to 20 nP, the
182 southward IMF B_z increased from -5 nT to -21 nT, the IMF B_y turned from $+5$ nT to -10 nT, and
183 the *AE*-index increased from ~ 280 nT to 1300 nT. According to the Shue et al. (1998) model, the
184 subsolar magnetopause shifted from $x \approx 10 R_E$ to $x \approx 5.3 R_E$. The minimum *SYM-H* (-192 nT) was
185 recorded at 20:25 UT.

186 The *H*-component variation at IRK, likewise the 6 April 2000 storm, indicated substorm
187 activations (*AE* increases) caused by the amplified SW pressure and the IMF variations at $\sim 16:48$
188 UT, $\sim 18:35$ UT, and ~ 20 UT. At the SSC, the southern boundary of the FAC R2 was far from IRK
189 and BOX. However, after 18:30 UT, and then around 20:00 UT, following the SW pressure
190 increases the boundary moved equatorward of these stations accompanied by two series of PiB
191 pulses and increases in the emission intensities.

192 The first Psc burst at $T > 4$ s was recorded at MND at 16:48 UT and also during the first
193 substorm activation (Figures 8 and 9). During the substorm activations after the SSC, there occurred

194 two PiB bursts at 18:18:04 and 18:30:26 UT, and then three PiB/PiC bursts after 20 UT with $T > 0.4$ s
195 and intensified green and red line emissions (Figure 9). At that time, MND and TOR were inside the
196 FAC R2 region.

197

198 **4. Discussion**

199 Section 3 describes several substorm activations accompanied by broadband PiB/PiC bursts and
200 airglow in the nightside mid-latitude atmosphere during the 6 April 2000 and 21 October 2001
201 events. Further, we discuss only the April 6, 2000 superstorm because the second event is very
202 similar.

203 **4.1. Short- vs. long-period pulsations**

204 First, we would like to justify the advantage of the short period and wavelength Pi1B pulsations for
205 timing EP onsets. These pulsations vanish faster in space and time than their longer period and
206 wavelength counterparts. Obviously, to resolve Pi1Bs, one needs records of a ≥ 10 Hz frequency
207 resolution rarely available at geomagnetic observatories. Precisely, the high time resolution
208 magnetic data from the MND and BOX stations allowed us to determine the EP double onset at $t_1 =$
209 16:39:40 UT and $t_2 = 16:45:40$ UT ($\tau = t_2 - t_1 = 6$ min) during the 6 April 2000 superstorm. The same
210 onsets (t_1, t_2) are revealed from the dynamic spectrum of the geomagnetic field oscillations at
211 periods in the range of 8–40 s obtained at the KAK station (Figure 3).

212 We note that a small duration of the Pi1B train (< 1 min) makes possible not only tracing the
213 EP onsets but also picking out the double onsets. This is practically impossible with the PiB's long-
214 period part (Pi2/Pi3 pulsations). Namely, when their period (T_{long}) is close to the interval, τ , between
215 the onsets, one pulsation train turns continuously into another. The latter is evident in the dynamic
216 spectra with 40–60 s periods from MND, BOX, and KAK (Figure 3). It is relevant to note that V.V.
217 Mishin et al. (2003, Figures 5 and 6) found from high resolution oscillograms of pulsations excited
218 by SW pulses that the first pulsation dies out prior to the start of the second one only if the interval
219 between SW pulses is sufficiently large: $\tau \gg T_{long}$.

220 Note that Cheng et al. (2018) attempted to determine the double substorm onset by the
221 decrease in the Pi2 period. They attributed the change to the tail shortening during the EP onset. In
222 fact, however, such a change of the Pi2/Pi3 period is dubious as there were only two periods
223 between the studied onsets, i.e., $\tau = 2T_{long}$. Such period decrease is characteristic of only a special
224 type of short period pulsations called IPDP (Kangas et al., 1998).

225 We will illustrate a more adequate determination of the duration, $\Delta\tau$, of Psc1,2/PiB bursts,
226 which depends on the frequency resolution, Δf , using the 16:40 UT event on 6 April 2000 as an
227 example. Figure 10 shows the Psc1 dynamic spectra obtained at MND with the sampling rate of 10
228 Hz and at KAK with 1 s data. It is seen that the burst lasted for $\Delta\tau \leq 20$ s at MND, which is in a
229 good agreement with an estimate $\Delta\tau = 15$ s from a detailed study by Parkhomov et al. (2010). At
230 KAK the burst lasted for $\Delta\tau = 2\text{--}3$ min, i.e., $\Delta\tau$ decreased with the increase of Δf . Clearly, timing the
231 isolated substorm EP onsets and substorm activations during storms is more accurate with high-
232 frequency sampling rates ($\Delta f > 10$ Hz).

233 As an example, let us note the CARISMA network data. Particularly, Rae et al. (2009, 2011)
234 using wavelet analysis of the CARISMA data with a 1 s resolution obtained a number of important
235 results on the relationship between the most equatorward auroral arc and the time of the expansion
236 onset determined from Pi1 pulsations. Presently, the network employs induction magnetometers
237 with a 20 Hz sampling rate. This will certainly improve the accuracy of determining the time of the
238 EP onset in pulsations and auroras. The new capability of the CARISMA network may also help to
239 clarify the features of PIBs in the short period range discussed below.

240 **4.2. Generation mechanism**

241 We now turn to discuss possible generation mechanisms of the observed pulsations and
242 optical emissions. The Psc generation is usually attributed to the arrival of the interplanetary shock
243 front (Nishida, 1978). In the studied event, the SW plasma density (pressure) increased ahead of the
244 front twice at 16:31:20 UT and at 16:35:16 UT (V.V. Mishin et al., 2013, Figure 8). That occurred
245 during the substorm growth phase initiated by the IMF turning southward about 20 min prior to the
246 SSC (Figure 1). The P_d pulses caused weak auroral emissions at 16:33:43 UT that slightly
247 intensified during the first onset at 16:39:51 UT (Figure S2). As shown earlier by V.V. Mishin et al.
248 (2013), the SW pressure enhancements resulted in MHD disturbances that propagated during 16:36–
249 16:40 UT from the dayside to the nightside. After that, MHD waves coming in the opposite
250 direction emerged from a possible nighttime substorm source during 16:41–16:44 UT.

251 The burst of pulsations related to the SSC results from the impact of the interplanetary shock
252 front on the magnetopause. Pulsations appear at mid- and low-latitude stations in accordance with
253 the source's global character (Parkhomov et al., 2017). According to Mishin et al. (2013), we
254 attribute the first burst of pulsations at MND and BOX to the Psc type caused by the daytime source.
255 At the same time (Figure 3), the effect of the substorm development is indicated by the fact that the
256 pulsations at MND (premidnight) were more enhanced than at BOX (duskside).

257 We ascribe the other bursts to the PiB type, though three of them (at 16:46 UT, 17:10 UT,

258 and 18:08 UT) occurred during the period of rapid variations of the SW pressure. Nonetheless, these
259 bursts were coincident with sharp intensifications of the *AE* index on the top of its greatly elevated
260 mean level. The bursts' location in the near midnight region is natural for the development of the
261 substorm current wedge (V.M. Mishin et al., 2010 a, b; 2011). What is more, the bursts after the
262 SSC were coincident with prompt changes of the near-midnight upward R1 and R2 FACs that
263 maximized at the center of R2- vortices near the boundary of the R1+ zone. Here, the maximum of
264 the westward electrojet (WEJ) is located, which corresponds to energetic electron precipitation and
265 auroras (e.g. V.M. Mishin, 1990; Korth et al., 2014, Carter et al., 2016). It is also relevant to note
266 that the Pi1B maximum intensity correlates with the WEJ maximum (Parkhomov & Rakhmatulin,
267 1975; Despirak, et al., 2020).

268 At the time of the second (main) substorm EP onset at 16:46 UT, the *AE* and *AL* indices,
269 FACs, and optical emissions intensified. However, there was only slight increase of pulsations at
270 BOX, most likely because the station was too far from the R2-/R1+ and auroral oval boundaries
271 (Figures 1 – 5 and Figures S1, S2). The next two PiBs were also detected during pressure pulses and
272 strong geomagnetic activity (*AE* = 2600 nT and FACs R1.2 = 6-8 MA). The Pi1B amplitude at
273 MND at 17:10 UT was much larger than at BOX, which was farther than MND from the R1-
274 boundary (Figure 5).

275 At 18:08 UT, when BOX was close to the R2+/R1- boundary, the Pi1B amplitude increased
276 by more than an order of magnitude relative to that at 17:10 UT (Figures 4 and 5). This bears
277 witness of the dependence of the PiB magnitude on the distance from the upward FAC regions. This
278 conjecture agrees with the data of the second substorm activation. Particularly, when the R2-
279 boundary approached MND, the amplitude of pulsations increased by 1.5 times still remaining
280 weaker than at BOX by a factor of two because of the greater distance to the maximum of the
281 upward R2- FAC.

282 At about 20 UT, both stations were located near or inside Region 2-, but BOX was at the
283 boundary R2-/R1 +, i.e., much closer to the upward FAC maximum. Accordingly, enhanced $T > 4$ s
284 pulsations observed at BOX were 4 times greater than at MND. At the same time, a sharp
285 attenuation was seen in the range $T < 4$ s (Figures 6 and 7). In addition, the maximum airglow
286 intensity was observed from TOR during 19–20 UT. The latter, together with the amplification of
287 the upward FAC, indicates that the electron precipitation was moving toward MND and TOR.

288 Therefore, we conclude that PiB pulsations were excited by rapid increases in the substorm
289 FACs' intensity indicated by the WEJ and *AE* index (mainly determined by *AL*, see Figure S1). The
290 amplitude of the core spectrum of pulsations increased when the upward R2- FAC boundary on the

291 nightside approached the stations. However, at the R2–/R1+ border, near the focus of FAC R2– and
292 hence electron precipitation, the amplitude of the pulsations in the range $T < 4$ s significantly
293 decreased. In particular, near local midnight (at 00:10 MLT and 01:08 MLT), their intensity at MND
294 dropped almost threefold (Figure 3). When BOX approached midnight (23 MLT), the Pi1B at 0.2–4
295 s also weakened, barely exceeding the noise level. However, the longer period part of pulsations ($>$
296 4 s) increased at that time.

297 We assume that this behavior is related to the change of the conductivity in the premidnight
298 (superstorm auroral/subauroral) ionosphere near Irkutsk. Indeed, around 20 UT the BOX station was
299 near the R1/R2 boundary, slightly northward of the maximum of the upward R2– FAC, that is, the
300 maximum of energetic electron precipitation (e.g., V.M. Mishin, 1990; Korth et al., 2014; Carter et
301 al., 2016). This agrees well with the intensification of optical emissions from TOR during 19–20
302 UT. Surely, energetic electrons increase the E-region conductivity and hence the electromagnetic
303 impedance. As a rule, this results in weakening of penetration of the short period pulsations into the
304 lower ionosphere (Lyatsky & Maltsev., 1983; Klibanova et al., 2008).

305 In other words, the impermeability of the lower wall of the ionospheric Alfvén resonator
306 (IAR, see shortly) increases. For longer period, $T > 4$ s, oscillations with wavelengths greater than the
307 height of the E layer ($\lambda > h_E \sim 120$ km) such a screening effect is unimportant. Therefore, the
308 oscillations with $T > 4$ s amplified when the FAC R2– and the WEJ maximum approached the BOX
309 station. The presence of the short period pulsations at MND at that time was because the station,
310 most likely, remained in the subauroral region equatorward of the electron precipitation where the
311 conductivity was low.

312 From the above analysis, we conclude that the impact of the interplanetary shock front was
313 responsible for bursts of Psc pulsations observed near the SSC. On the other hand, Nishida (1978)
314 and Parkhomov et al. (2017) attributed Psc to the development of ion cyclotron instability caused by
315 the transverse anisotropy on the outer L shells ($L > 6$) where the proton gyrofrequency, $\omega_{Bp} =$
316 $eB/m_p \approx 3 \cdot 10^3 / L^3$ (s^{-1}), falls into the range of short-period pulsations Psc1,2. The data from the
317 LANL 94 satellite, which was at the longitude of Irkutsk (Figure S3), give evidence in favor of this
318 mechanism at the beginning of the 6 April 2000 storm. Prior to the storm, the transverse anisotropy,
319 characteristic of the trapped plasma, practically disappears after SSC.

320 One of the main tasks of the Pi1B/Psc theory is to explain a continuous broadband spectrum
321 with the resonant structure. The latter is characteristic of the spectrum of the IAR model (Polyakov
322 & Rapoport, 1981). The lower wall of the resonator is due to the maximum conductivity in the E

323 layer. The upper wall is formed at height $h \sim 3000$ km by the inflection of the Alfvén velocity
324 profile related to the decrease of the plasma density. A typical theoretical spectrum features diffuse
325 resonance maxima on the top of enhanced quasi-continuous 0.1 Hz – 2 Hz noise with a noticeable
326 decrease in amplitude at frequencies greater than 0.5 Hz (Lysak, 1988). Therefore, it is reasonable to
327 suggest that the continuous spectrum of the Pi1B and Psc1,2 pulsations is associated with their
328 passage through the IAR (cf. Lysak, 1988; Parkhomov et al., 2017).

329 Further, recall the similarity of the observed pulsations, such as a sequence of a few
330 oscillations containing the periodic resonant structure and rapidly decaying just after they emerged.
331 This behavior suggests that they resulted from a pulsed electromagnetic impact on the IAR similar
332 to a lightning discharge, seismic oscillations, solar wind pressure pulses (Dovbnya et al., 2014) or
333 even injection of ion beams into the IAR region for a few seconds (Volokitin & Drozdov, 1993).
334 Lysak (1988) surmised that PiBs may be excited by a strong pulsed change of the field - aligned
335 currents. This is consistent with the observations of PiBs during auroral current enhancements
336 (Untiedt et al., 1978; Opgenoorth et al., 1980).

337 This is also consistent with our observations of Pi1B bursts after fast increases in the *AE*
338 index, magnetic flux, Ψ , and corresponding FACs - the characteristic feature of the EP onset
339 according to MIT (V.M. Mishin, 1990). Recall, for example, the second EP onset at 16:46 UT with
340 the sharp increase in the *AE* index and fast equatorward motion to mid latitudes of the FAC regions
341 with the several-fold increase in the FAC intensity (Figures 1, 4).

342 A final remark is in order regarding the ionospheric feedback instability (IFI). A large
343 number of works invoke the IFI to explain processes during the substorm EP onset. The IFI
344 develops in the presence of a sufficiently large convection electric field, E_c , due to deceleration of
345 the plasma convection and formation of a strong vertical shear in the convection flow in the
346 ionosphere (Trakhtengertz & Feldstein, 1981, 1991; Lysak, 1988; Lysak, 1991). Precipitation of
347 energetic electrons and the presence of depleted conductivity frequently occurring in the FAC R2–
348 are favorable factors for the IFI development (Streltsov & Mishin, 2018). However, in the events
349 studied in this paper, we did not find in situ data clearly supporting the IFI development near the
350 MND and BOX observatories.

351

352 **5. Conclusion**

353 The main goal of this work is to show that using short-period burst pulsations allows
354 determining the substorm expansion onsets more accurately than using conventional Pi2 pulsations.
355 In addition, this method is well suited for picking out double onsets and stormtime substorm

356 activations inaccessible for Pi2. In particular, we have shown that

357 1. Intensifications of mid-latitude broadband geomagnetic pulsations (Psc/PiBs) and airglow
358 during two strong magnetic storms mark the substorm expansion onsets or the onset of short
359 substorm activations.

360 2. Detecting the Psc/PiB short period part (Psc1,2/Pi1B) makes possible to identify a single
361 onset and to separate double substorm onsets, as well as series of onsets of short substorm
362 activations during storms, with the accuracy better than 0.5 min.

363 3. Burst pulsations with a broadband spectrum containing the periodic resonant structure are
364 excited in the ionospheric Alfvén resonator impacted by a pulsed change of the field-aligned
365 currents in the substorm current wedge circuit.

366 4. We attribute the near-midnight minimum of pulsations within the 0.2–4 sec periods to the
367 weakening of their penetration due to enhanced precipitation of energetic electrons

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380 and Japan magnetic observatories for providing magnetic data used in this study. Available
381 geomagnetic data are obtained directly from [http://www.intermagnet.org/data-donnee/download-](http://www.intermagnet.org/data-donnee/download-eng.php)
382 [eng.php](http://www.intermagnet.org/data-donnee/download-eng.php) (INTERMAGNET), <http://space.fmi.fi/image/> (IMAGE),
383 <http://space.augsburg.edu/maccs/request.jsp> (MACCS), <https://www.asf.alaska.edu/magnetometer>
384 (GIMA), and <http://www.carisma.ca> (CARISMA). The SW key parameters and optical observations
385 of auroras from POLAR were obtained through the NASA Space Science Data Coordinated Archive
386 (<http://cdaweb.gsfc.nasa.gov/>). The AE index was obtained through the SuperMAG website
387 (<http://supermag.jhuapl.edu>).

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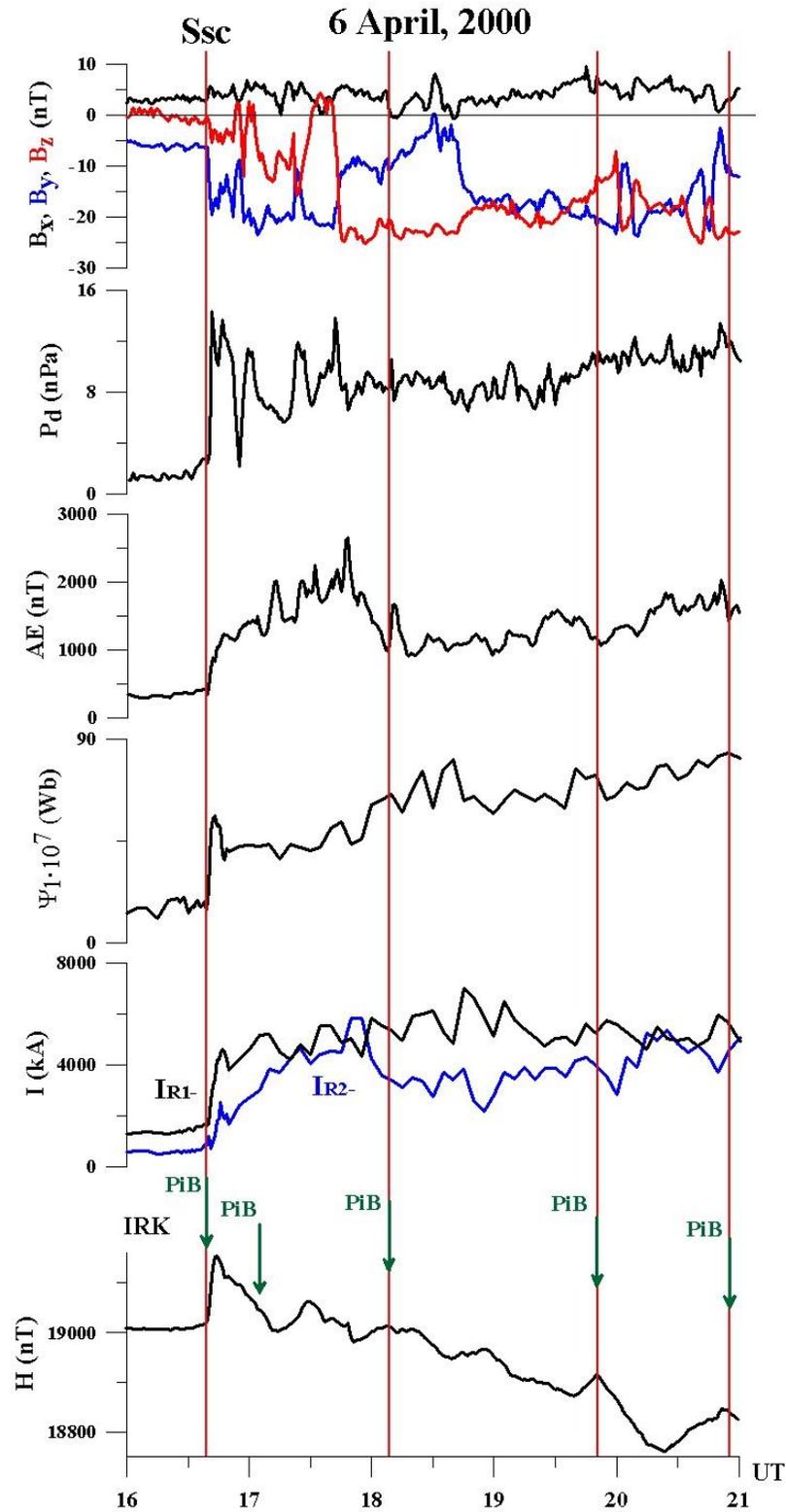
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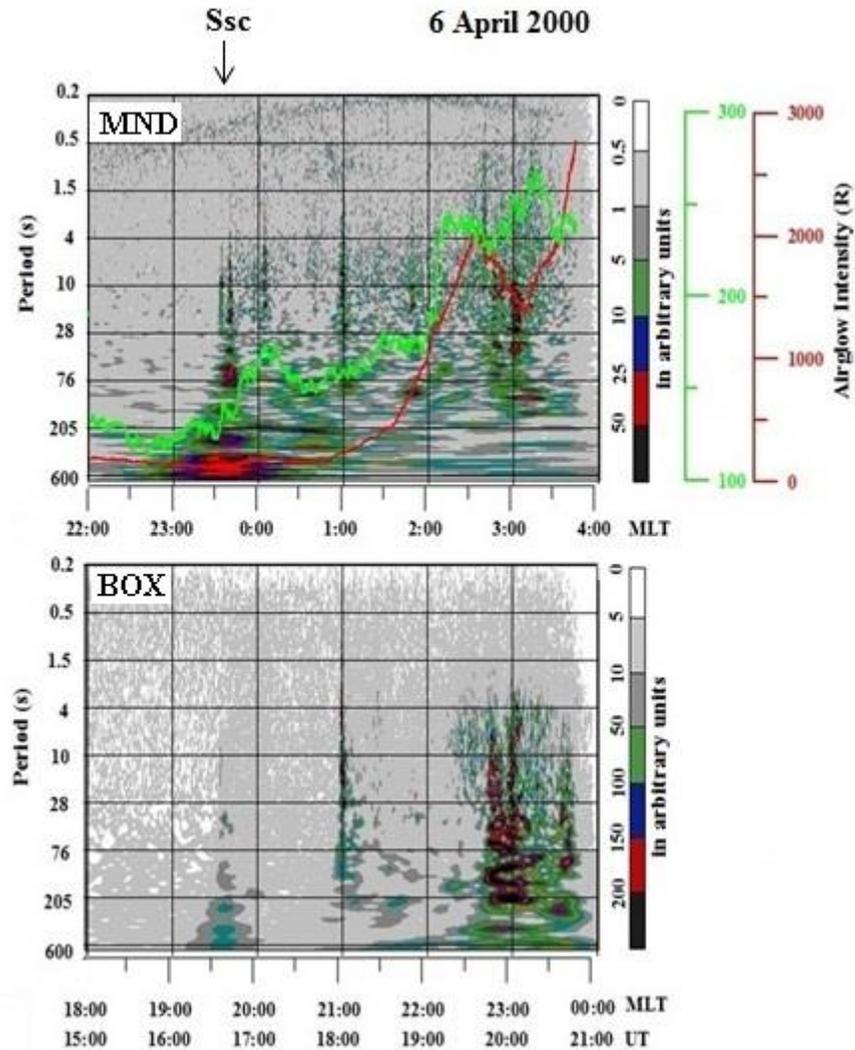
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Figure 1. The 6-7 Apr 2000 superstorm. **From top to bottom:** Variations in the IMF components, solar wind **ram** pressure (P_d), **AE-index**, magnetic flux through the polar cap, Ψ_1 , **FAC intensities in the R1-, R2- regions**, and the **H-component at IRK**. Red lines show the boundaries of two intervals of substorm activations.

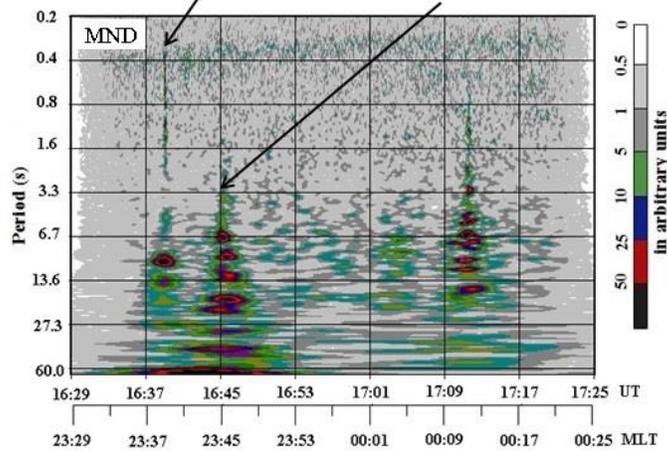


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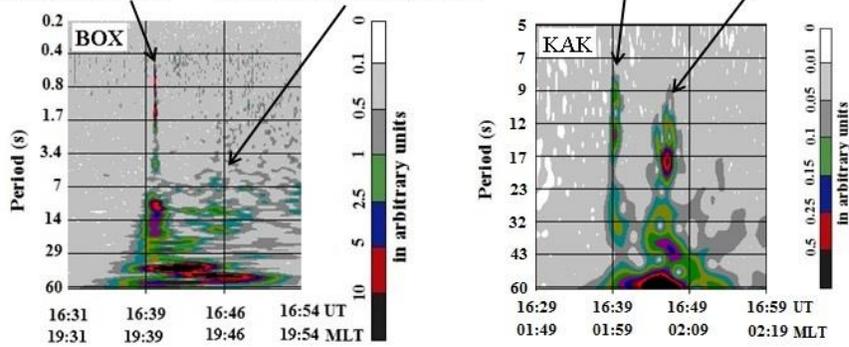
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560 Figure 2. Dynamic spectra of geomagnetic pulsations from (top) MND and (bottom) BOX
 561 during the 6-7 Apr 2000 superstorm: The amplitude variation in arbitrary units as a function of the
 562 period and universal time (UT) or magnetic local time (MLT). The green (red) line shows the
 563 intensity in Rayleigh of the 557.7 nm (630.0 nm) emission from TOR (intensity scales are on right,
 564 top panel).

First impulse in pulsations Psc 1,2 at 16:39:40 UT (first substorm EP onset at Ssc) Second impulse in pulsations Pi1B at 16:45:40 UT (second substorm EP onset)



First pulse in pulsations Psc1,2 at 16:39:40 UT Second impulse in pulsations Pi1B (weak) at 16:45:40 UT 16:39:40 UT 16:45:40 UT



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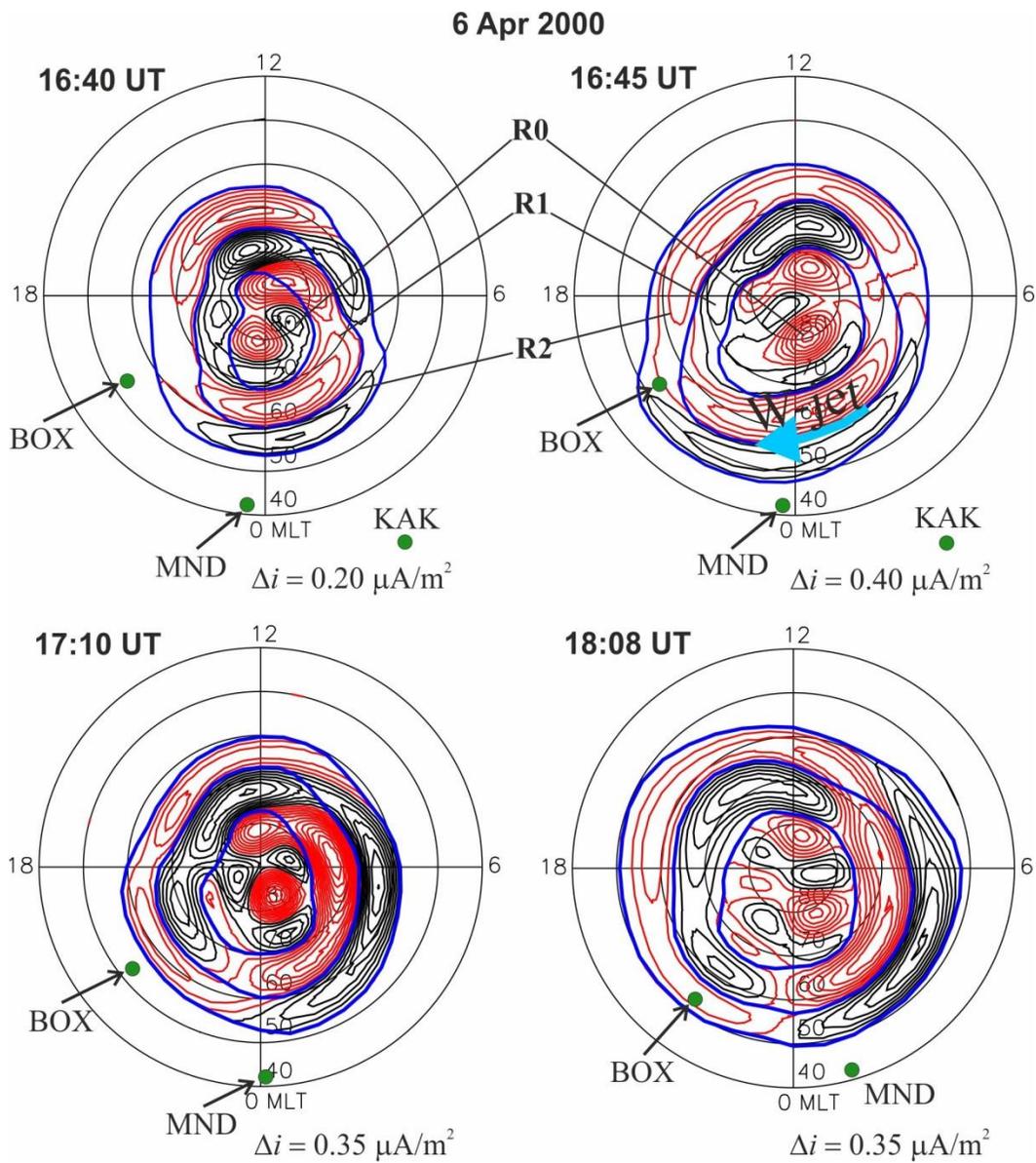
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Figure 3. Dynamic spectra of the Psc/PiB geomagnetic pulsations from the MND (CGM: 47.5° Φ, 177.5° Λ), BOX (CGM: 53.9° Φ, 114° Λ), and KAK (CGM: 29.25° Φ, 211.7° Λ) during the 6 April 2000 first substorm activation.



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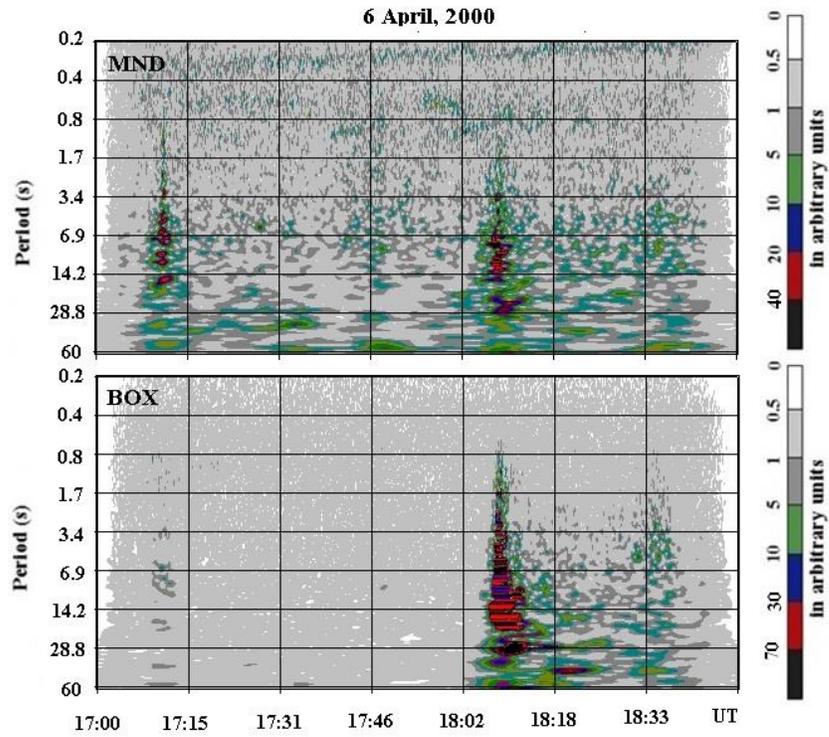
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Figure 4. Magnetogram Inversion Technique MLAT-MLT maps of the FAC density distribution at (top panel) 16:40 and 16:45 UT (double EP onset at the start of the 6 April 2000 superstorm and first substorm activation) and (bottom) at 17:10 and 18:08 UT. Blue thick lines show the boundaries of the FAC R1, R2, and of the polar cap, R0. Red (black) isolines are isocontours of the downward (upward) FAC density. Green bold dots show the locations of the observatories.



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Figure 5. Dynamic spectra of geomagnetic pulsations from the MND and BOX during the 6-7 Apr 2000 superstorm. Shown are: the amplitude variations depending on the period (ordinate), universal time (UT). The relative amplitude values feature a color scale in relative units.

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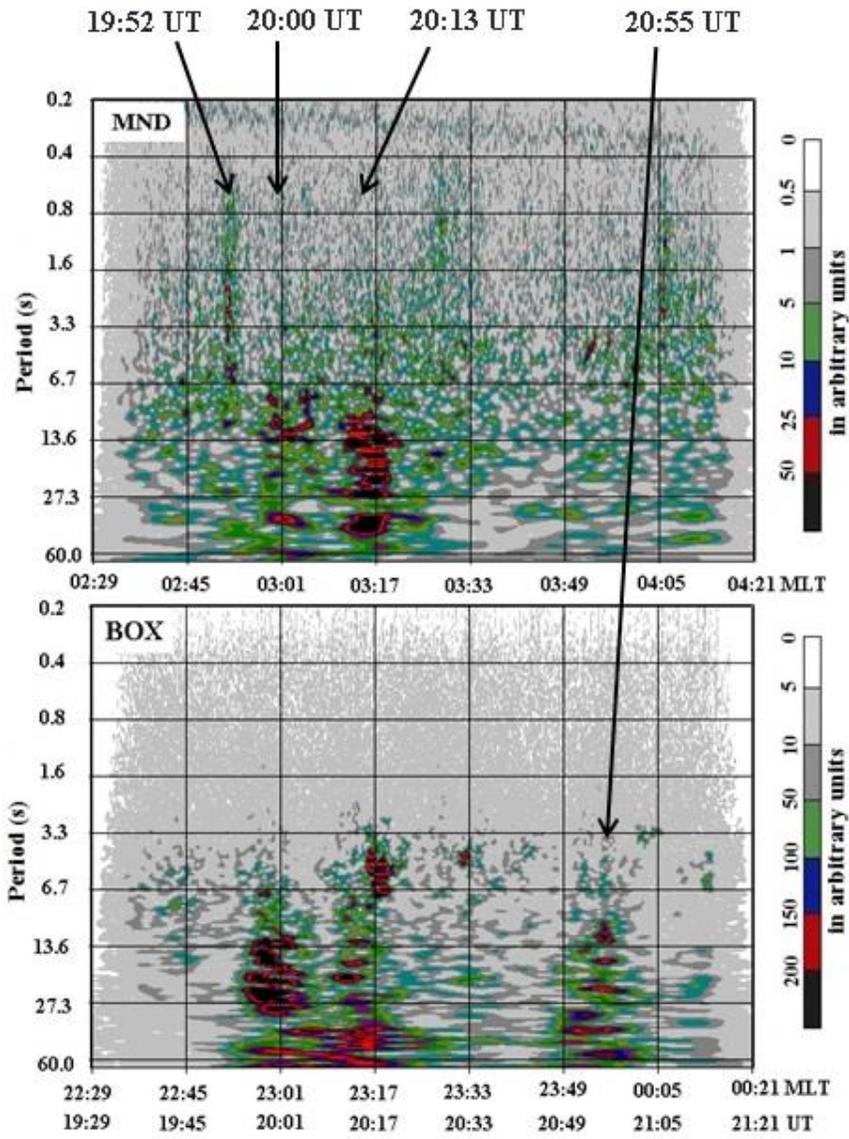
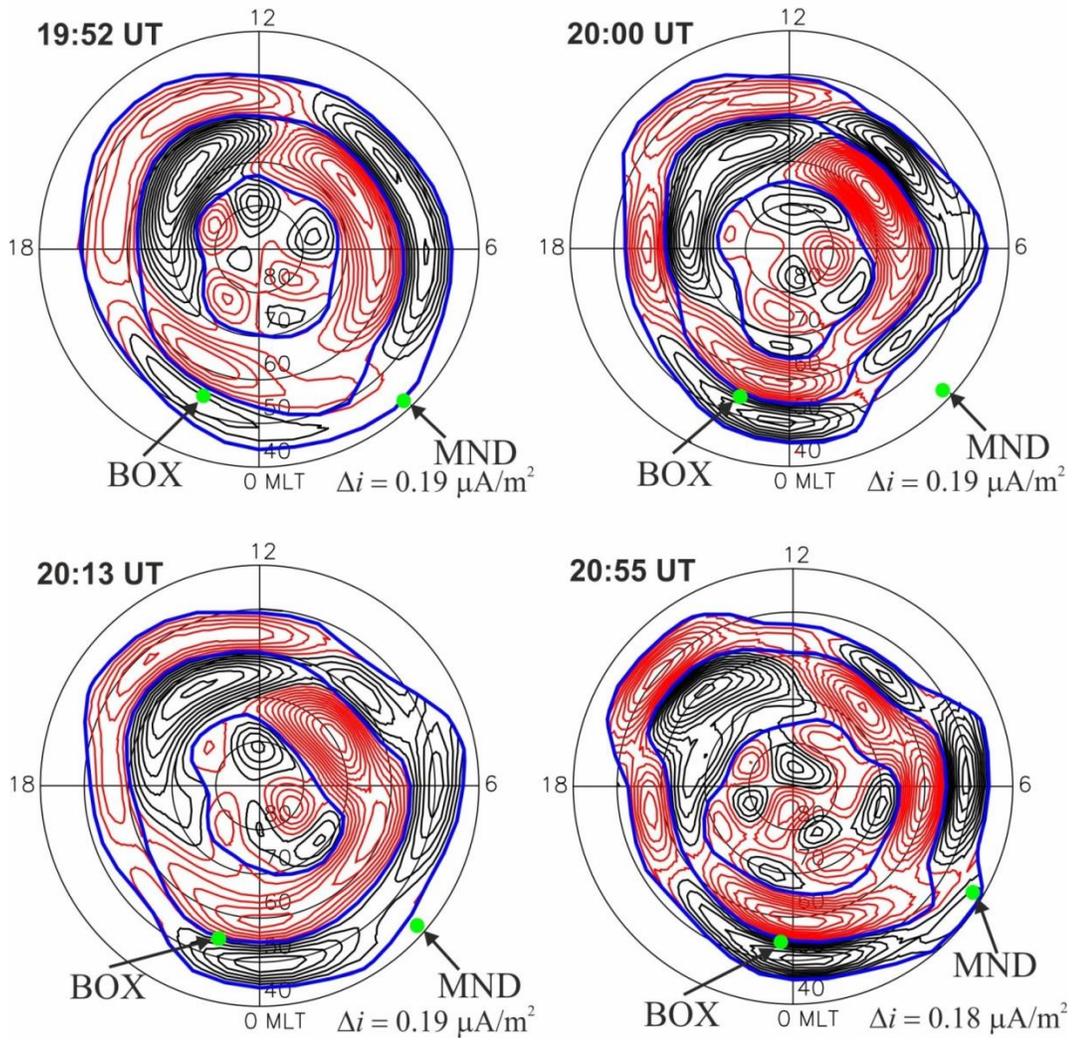


Figure 6. PiB dynamic spectra from MND (CGM: $47.5^\circ \Phi$, $177.5^\circ \Lambda$) and BOX (CGM: $53.9^\circ \Phi$, $114^\circ \Lambda$) during the substorm second activation.

6 Apr 2000



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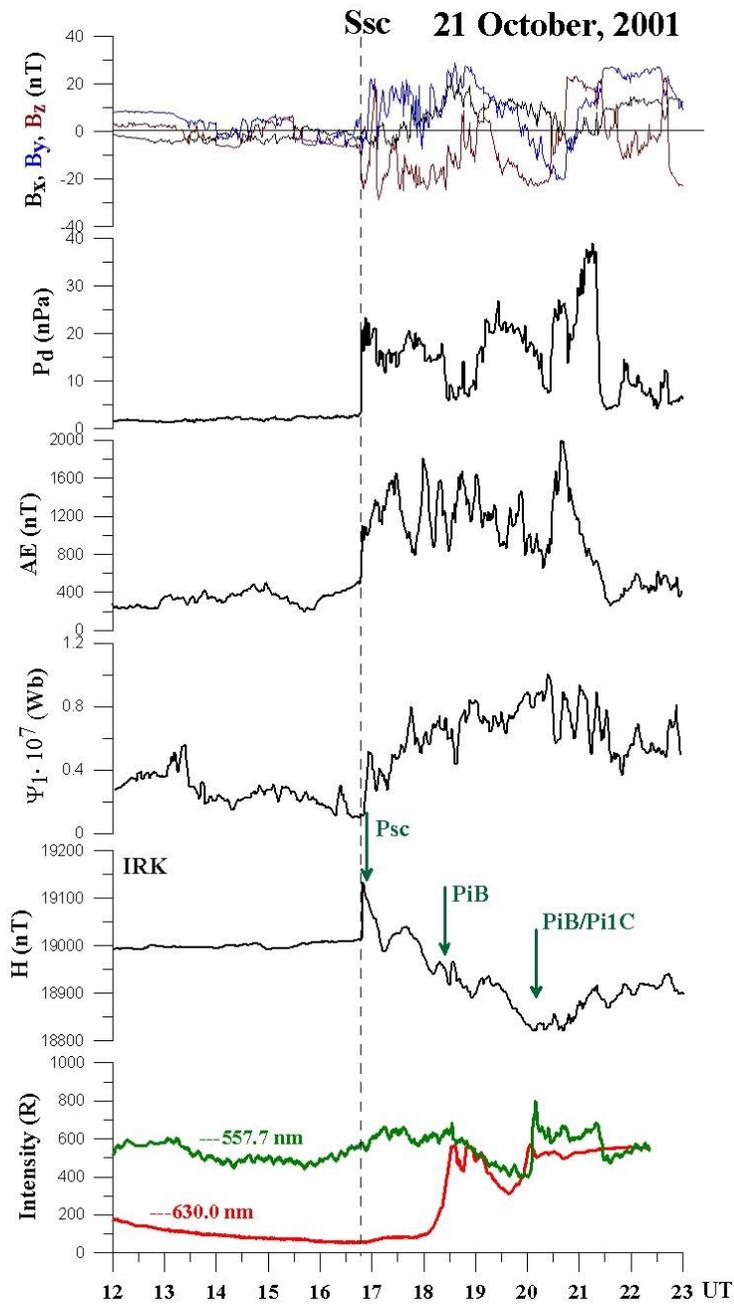
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640 Figure 7. Four moments during the substorm second activation interval during the 6 April 2000
641 superstorm. The same format as in Figure 4.

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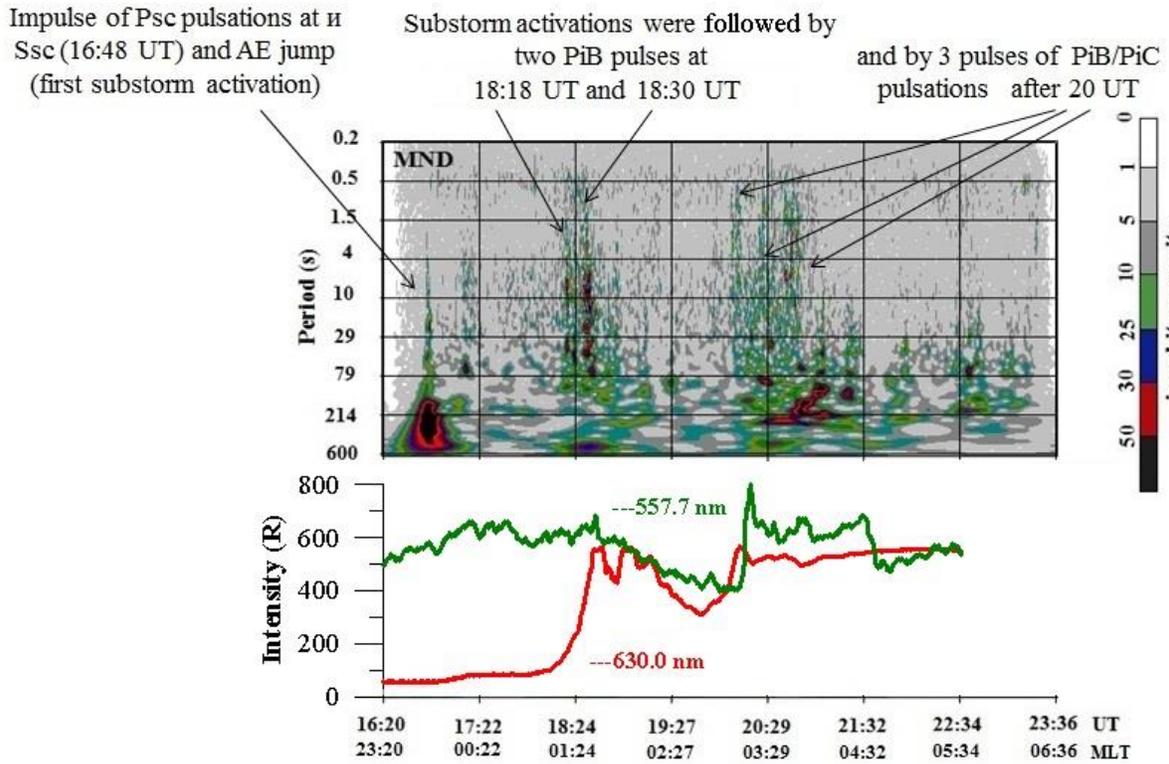
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Figure 8. The 21 October 2001 storm. Variations in the IMF components, SW ram pressure (P_d), AE-index, magnetic flux through the polar cap, Ψ_1 , the H-component of the geomagnetic field from IRK, and intensities of the 557.7 nm (green) and 630.0 nm (red) emissions from TOR.



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Figure 9. (top) Dynamic spectra of geomagnetic pulsations from MND during the 21 October 2001 storm vs. the pulsation period and UT/MLT. Color codes in arbitrary units for the wave spectrum are given to the right of the spectrogram. (bottom) The intensity of the 557.7 nm (the green line) and 630.0 nm (red) from TOR as a function of UT.

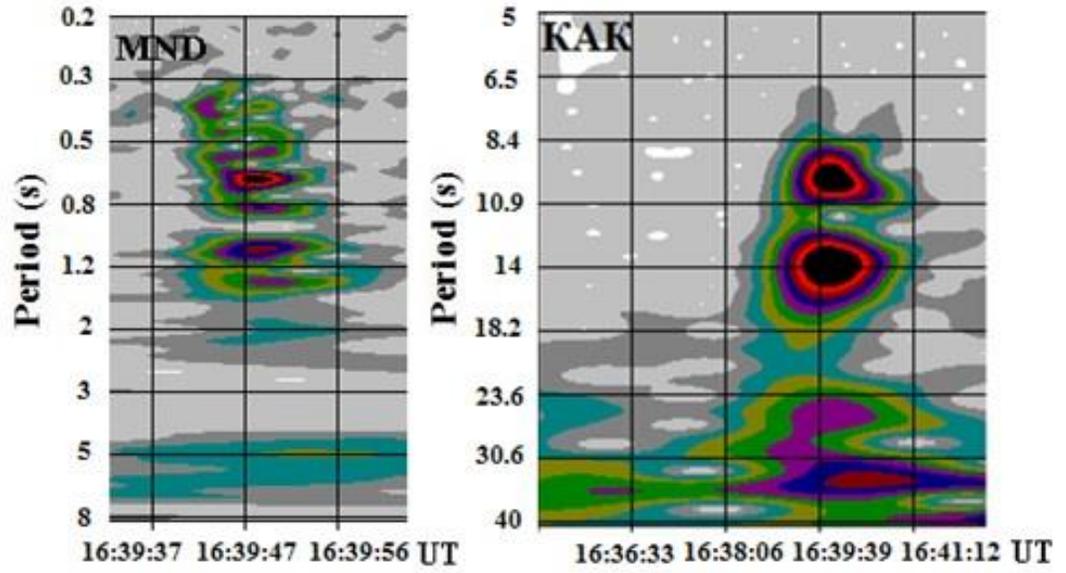
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Figure 10. The 6 April 2000 storm. The duration, $\Delta\tau$, of the Psc1, 2 pulsation depending on the sampling rate at MND (10 Hz) and KAK (1 Hz).