Plain-scale sustained changes in well water levels following a large earthquake: possible evidence of permeability decrease in a shallow groundwater system

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Abstract

Observation of earthquake-induced changes in well water levels provides an opportunity to study the effects of seismic activity on the groundwater system. In this study, we used data from a plain-scale well network in the alluvial Canterbury Plain of New Zealand's South Island to document sustained and complex changes in well water levels, following the 2010 Darfield earthquake. We interpret that the sustained increases in well water levels in the midstream area, as well as sustained decreases in the downstream area, resulted from decreases in plain-scale permeability. These decreases in permeability were caused by the consolidation of the liquefied sediments following the earthquake. These results may provide a better understanding of the effects of large earthquakes on groundwater systems and resources, especially in liquefaction areas.

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8	Key points:
9	• Plain-scale sustained changes occurred in well water levels in the Canterbury Plain
10	after the Darfield earthquake.
11	• These sustained changes indicate that there should be plain-scale permeability
12	decreases in the shallow groundwater system.
13	• The permeability decreases could be attributed to the consolidation of sediment in the
14	Canterbury Plain.
15	Abstract: Observation of earthquake-induced changes in well water levels provides an
16	opportunity to study the effects of seismic activity on the groundwater system. In this
17	study, we used data from a plain-scale well network in the alluvial Canterbury Plain of
18	New Zealand's South Island to document sustained and complex changes in well water
19	levels, following the 2010 Darfield earthquake. We interpret that the sustained increases
20	in well water levels in the midstream area, as well as sustained decreases in the
21	downstream area, resulted from decreases in plain-scale permeability. These decreases in

22 permeability were caused by the consolidation of the liquefied sediments following the earthquake. These results may provide a better understanding of the effects of large 23 24 earthquakes on groundwater systems and resources, especially in liquefaction areas. Keywords: earthquake; well water levels; permeability decrease; liquefaction; 25 consolidation; New Zealand 26 **Plain Language Summary** 27 The Mw 7.1 Darfield earthquake that occurred in the eastern part of New Zealand's 28 South Island on September 4th, 2010. It generated widespread hydrological responses in 29 the Canterbury Plain, such as changes in well water levels. Although many studies have 30 31 focused on coseismic hydrological responses to earthquakes, few have studied the effects of large earthquake on the sustained change in groundwater systems. In this study, we 32 33 analyze data from a plain-scale well network, and propose a new model to explain the sustained changes observed in well water levels after the earthquake. In this model, the 34 35 decrease in permeability of the shallow groundwater system was induced by seismic 36 sediment consolidation in the Canterbury Plain. This in turn resulted in sustained changes in the hydraulic gradient of the shallow groundwater system, which increased 37 38 groundwater resources in midstream areas but showed a decrease in downstream areas.

39 **1. Introduction**

The change in groundwater level, a widespread hydrogeological phenomenon that occurs during and after earthquakes, is of interest because it allows for the study of the potential effects of earthquakes on hydrological systems and groundwater resources. Several

43	research studies show that earthquakes, especially large ones, can cause changes in well
44	water levels by changing the pore pressure and hydrogeologic parameters of the
45	groundwater system through static stress changes and seismic waves (dynamic stress)
46	during the earthquakes (e.g., Brodsky et al., 2003; Cox et al., 2012; Gulley et al., 2013;
47	Roeloffs et al., 1998; Shi et al., 2018; Shi & Wang, 2017; Wang & Manga, 2015).
48	Coseismic changes in the well water level have been studied extensively by many
49	hydrogeologists and geophysicists since the 1960s. Several mechanisms such as static
50	strain (Ge and Stover, 2000; Jonsson et al., 2003; Roeloffs, 1996; Wakita, 1975),
51	consolidation (Cox et al., 2012; Wang et al., 2001; Wang and Chia, 2008), new additional
52	recharge (Wang et al., 2004; 2015), and permeability increase (Brodsky et al., 2003;
53	Roeloffs et al., 1998) were proposed to explain this hydrological phenomenon. However,
54	there are currently few research studies on the sustained changes in well water levels after
55	earthquakes.

The Greendale Fault in New Zealand has produced a series of earthquakes from 2010 56 through the present. This includes the Mw 7.1 Darfield earthquake (main shock) on 57 September 4, 2010, which affected well water levels in the Canterbury Plain (Cox et al., 58 2012). A detailed analysis of the coseismic changes in well water levels in the Canterbury 59 Plain was performed by Cox et al. (2012). They showed that the groundwater levels 60 increased coseismically via the consolidation of unconsolidated sediments during the 61 earthquake. In this study, we revisited the Darfield earthquake and focused on the 62 63 sustained changes in well water level following the earthquake. An analysis of summer irrigation on the shallow groundwater system in the Canterbury Plain showed that it had 64 no effect on groundwater levels. This event allows for the investigation of permeability 65

changes in the shallow groundwater systems, as well as their effects on well water levels
following earthquake-induced consolidation and liquefaction of sediments in the
Canterbury Plains.

69 2. Observation

70 The Canterbury Plain includes a large (~2,000 km²) alluvial fan extending from the eastern margin of the Southern Alps to the South Pacific Ocean (Fig. 1a). This alluvial 71 plain is composed of Quaternary glacial outwash from the Southern Alps; it overlies late 72 Cretaceous-Tertiary sediments and a Permian-Jurassic Torlesse greywacke basement 73 74 (Cox et al., 2012; Rutter et al., 2016). The groundwater system in this region consists of 75 thick (300–600 m), unconfined or semiconfined gravel aquifers in the proximal area, as 76 well as confined aquifers of gravel beds interlayered with estuarine and shallow marine beds near the coast (Fig. 1b). Groundwater in the unconfined and semiconfined aquifers 77 78 flows southeastward from the foothills toward the coast. In the confined aquifers, 79 groundwater flows upward and is discharged through springs along the coast and offshore, 80 as a result of the vertical hydraulic gradient (Cox et al., 2012).

A large-scale network of hydrological monitoring wells was built for monitoring the
groundwater resources, and continuously monitored throughout the plain by the
Canterbury Regional Council of New Zealand (Environment Canterbury) (Fig. 1a). We
obtained water level data, collected from 2009 to 2012, by pressure transducers for 112
wells from the Environment Canterbury website (https://ecan.govt.nz/). Data used in this
study were obtained from wells monitored at 15-min intervals at depths ranging from 5 to
405 m. All data were corrected for barometric pressure such that the data reflected the

- 88 water pressure, and manual measurements were taken in the field and adjusted where
- required to remove instrument errors. 89



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Fig. 1. Hydrogeology of the study area (modified from Cox et al., 2012 & Rutter et al., 91 2016). (a) Distribution of observational wells in the Canterbury Plain. Piezometric 92 contours indicate groundwater flow from the foothills to the coast. The shallow 93 94

groundwater system in the alluvial plain is divided into the following areas: downstream

(piezometric head < 25 m), midstream (25 m < piezometric head < 125 m), and upstream
(piezometric head > 125 m), according to hydrogeological conditions. (b) Simplified
cross-section along line W-E (Fig. 1a). Blue arrows indicate the direction of groundwater
flow in the shallow groundwater system.

99 **3. Results**

We used three-day data to plot **Fig. 2** as the water levels in these wells returned to new 100 101 'balanced' levels after several days. Additionally, this makes it easier to indicate the changes caused by the Darfield earthquake in a short-term time sequence. Changes in 102 water levels in these observational wells, including coseismic change, post-seismic 103 104 change, and sustained change, were quantified based on the following definitions: coseismic change is a step-like change in water levels during the earthquake; post-seismic 105 change is the change in water levels from the time of the earthquake to when water levels 106 107 reach their new post-seismic levels; and finally, a sustained change is the total change in water levels caused by the earthquake. The sustained change is the sum of the coseismic 108 109 change and the post-seismic change. Additionally, if these changes were smaller than the 110 amplitude of the tidal response, the wells were recorded as "no response," whereas if there was no data or the data was considered anomalous during and after the earthquake, 111 112 the wells were recorded as "no data."



Fig. 2. Primary changes in water levels in two observation wells induced by the Darfield earthquake. The "coseismic step-like" increases in well water levels are followed by "post-seismic" decreases, as well as (a) a sustained higher water level and (b) sustained lower water level, which represent the typic types (types II and III, respectively) of changes shown in Fig. 3.

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The well water levels in the Canterbury Plain changed during and after the Darfield earthquake. The water levels in 94 (not including the well in which the record of water level was missing during the earthquake) of the 112 studied wells increased abruptly during the earthquake, followed by a gradual decrease (Fig. 2). Most wells recovered to a level either above or below the pre-earthquake level after 2-3 days, indicating sustained changes in well water levels (Fig. 3). Sustained increases (type II) were more common in the central region of the plain, where well water levels (piezometric head) before the 126 earthquake were > 25 m. Sustained decreases (type III) were common near the coast,

127 where well water levels (piezometric head) before the earthquake were < 25 m.



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Fig. 3. Change in well water levels in the Canterbury Plain related to the 2010 Darfield
earthquake. There are six typic patterns of sustained change in water levels with respect
to coseismic, post-seismic, and sustained change.

132 Well water levels in groundwater systems usually become higher with increased

133 proximity to upstream areas. Therefore, the groundwater level data can be used to

134 determine the part of whole groundwater system in which the observed well is present,

- and the change in well water level at different parts can be conveniently analyzed through
- 136 the correlation between the change in well water level and well water level. As shown in
- 137 Fig. 4, The maximum value of sustained increase in the midstream area was

approximately 4 m, and the minimum value of sustained decrease in the downstream area
was approximately -2 m. This indicates a sustained increase in the hydraulic gradient by
as much as 5% between the midstream and the downstream areas of the shallow
groundwater system.



Fig. 4. Spatial and temporal changes in well water levels in the shallow groundwater system in the Canterbury Plain: (a) coseismic change, (b) post-seismic change, and (c) sustained change versus pre-seismic water levels in the upstream, midstream, and downstream area. It shows that there are sustained decreases in the well water level in the downstream area, sustained increases in the midstream area, and no changes in the upstream area.

149 **4. Discussion**

As mentioned in the **Introduction section**, four conceptual models are proposed as 150 possible mechanisms for such effects. However, none of these models can explain the 151 152 sustained increase in the hydraulic gradient observed after the Darfield earthquake. First, 153 the pattern of well water levels sustained changes in the Canterbury Plain was not related to static stress change induced by the faulting of the Greendale Fault. This is because the 154 155 change in static stress does not show such a pattern. Second, the changes cannot be explained purely by increased recharge, as that would imply that well water levels in the 156 157 entire system should be higher than they were before the earthquake, which was clearly 158 not the case. Third, the sustained changes cannot be attributed to an increase in 159 permeability, as this would have caused the well water levels in the recharge area to 160 decrease, whereas those in the discharge area to increase (Rojstaczer and Michel, 1995). 161 Lastly, the sustained decrease in well water levels in the downstream area contradicts the 162 hypothesis of permanent strain produced by consolidation. However, the coseismic 163 increase in water levels in the alluvial plain, away from the coast, could be attributed to 164 consolidation (Cox et al., 2012). Therefore, a new model is needed to properly explain

the sustained increase in the hydraulic gradient shown by our data for the CanterburyPlain.

167	Coseismic, step-like increases in water levels occurred on a regional scale across the
168	Canterbury Plain in the near field of the Darfield earthquake. Such increases are usually
169	attributed to coseismic consolidation of unconsolidated sediments (Cox et al., 2012;
170	Wang et al., 2001). Therefore, we hypothesize that the sustained increase in the hydraulic
171	gradient between the downstream and midstream areas indicates a decrease in
172	permeability. This decrease was caused by coseismic consolidation of the shallow
173	groundwater system in the downstream and midstream area, which in turn reduced
174	groundwater flow through the shallow groundwater system and produced the pattern of
175	sustained changes in well water levels seen in our results.
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However, a small number of confined wells in the downstream area (liquefaction area)
have experienced sustained decrease in water levels, which cannot be explained by the

decreased permeability. We suggest that this phenomenon may be caused by the
liquefaction-induced disruption of the groundwater system. In the downstream area, due
to failure of the weak permeable layer, the confined aquifer began to leak under the
vertical hydraulic gradient, which eventually led to the decrease of water levels in these
confined wells.

The change in well water levels, as well as changes in permeability, could be used to study the effects of changes in permeability on groundwater resources in the Canterbury Plain. Groundwater resources in the midstream area increased after the earthquake, as a result of decreased discharge to the downstream area. Meanwhile, groundwater resources in the downstream area decreased after the earthquake, because of decreased recharge from the midstream area. Further research is required to better understand the effect of the Darfield earthquake on groundwater resources in the Canterbury Plain.

199 **5.** Conclusions

200 In this study, we analyze data acquired from a large-scale well network and propose a 201 new model to explain the sustained changes in the hydraulic gradient documented after 202 the 2010 Mw 7.1 Canterbury earthquake. Our results show sustained increases in well 203 water levels in the midstream area and sustained decreases in well water levels in the downstream area after the earthquake. The sustained increase in the pressure gradient 204 205 indicates that there should be permeability decreases in the shallow groundwater system. This could possibly be attributed to the consolidation of sediment during the earthquake. 206 The decrease in permeability is used to explain that the sustained changes in the hydraulic 207 gradient had a positive effect on the groundwater resources in the midstream area but a 208

209 negative effect in the downstream area. This study may provide a better understanding of

210 the effects of large earthquakes on hydrological responses and groundwater resources,

211 especially in liquefaction areas.

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- 221 <u>https://ecan.govt.nz/data/well-search/welldetails/</u> with well number).

222 References

- 223 Brodsky, E. E., Roeloffs, E., Woodcock, D., Gall, I., & Manga, M. (2003). A mechanism
- for sustained groundwater pressure changes induced by distant earthquakes. *Journal of*
- 225 Geophysical Research, 108, 2390. doi:10.1029/2002JB002321
- 226 Cox, S. C., Rutter, H. K., Sims, A., Manga, M., Weir, J. J., Ezzy, T., et al. (2012).
- 227 Hydrological effects of the Mw 7.1 Darfield (Canterbury) earthquake, 4 September 2010,
- New Zealand. New Zealand Journal of Geology and Geophysics, 55(3), 231–247.
- 229 doi:10.1080/00288306.2012.680474

- 230 Elkhoury, J. E., Brodsky, E. E., & Agnew, D. C. (2006). Seismic waves increase
- 231 permeability. *Nature*, 441, 1135–1138. doi:10.1038/nature04798
- 232 Ge, S., & Stover, S. C. (2000). Hydrodynamic response to strike- and dip-slip faulting in
- a half-space. Journal of Geophysical Research, 105(B11), 25513–25524.
- 234 doi:10.1029/2000JB900233
- 235 Gulley, A. K., Dudley Ward, N. F., Cox, S. C., & Kaipio, J. P. (2013). Groundwater
- responses to the recent Canterbury earthquakes: a comparison. Journal of Hydrology, 504,
- 237 171–181. <u>doi:10.1016/j.jhydrol.2013.09.018</u>
- Jonsson, S., Segall, P., Pedersen, R., & Bjornsson, G. (2003). Post-earthquake ground
- 239 movements correlated to pore-pressure transients. *Nature*, 424, 179–183.
- 240 <u>doi:10.1038/nature01776</u>
- Liao, X., Wang, C. Y., & Liu, C. P. (2015). Disruption of groundwater systems by
- earthquakes. *Geophysical Research Letters*, 42, 9758–9763. doi:10.1002/2015GL066394
- 243 Manga, M., & Wang, C. Y. (2015). Earthquake Hydrology. doi:10.1016/b978-0-444-
- 244 53802-4.00082-8
- 245 Manga, M., Beresnev, I., Brodsky, E. E., Elkhoury, J. E., Elsworth, D., Ingebritsen, S. E.,
- et al. (2012). Changes in permeability caused by transient stresses: Field observations,
- experiments, and mechanisms. *Reviews of Geophysics*, 50, RG2004.
- 248 <u>doi:10.1029/2011RG000382</u>

- 249 Roeloffs, E. A. (1996). Poroelastic methods in the study of earthquake-related hydrologic
- 250 phenomena, in: Dmowska, R. (Ed.), Advances in Geophysics. Academic Press, San Diego.
- 251 Roeloffs, E.A., Sneed, M., Galloway, D. L., Sorey, M. L., Farrar, C. D., Howle, J. F., &
- Hughes, J. (2003). Water-level changes induced by local and distant earthquakes at Long
- 253 Valley caldera, California. Journal of Volcanology and Geothermal Research, 127, 269-
- 254 303. <u>doi:10.1016/S0377-0273(03)00173-2</u>
- 255 Rutter, H. K., Cox, S. C., Dudley Ward, N. F., & Weir, J. J. (2016). Aquifer permeability
- 256 change caused by a near-field earthquake, Canterbury, New Zealand. Water Resources
- 257 Research, 52(11), 8861–8878. doi:10.1002/2015WR018524
- 258 Shi, Z., & Wang, G. (2017). Evaluation of the permeability properties of the Xiaojiang
- 259 Fault Zone using hot springs and water wells. *Geophysical Journal International*, 209(3),
- 260 1526–1533. doi:10.1093/gji/ggx113
- 261 Shi, Z., Zhang, S., Yan, R., & Wang, G. (2018). Fault zone permeability decrease
- 262 following large earthquakes in a hydrothermal system. *Geophysical Research Letters*, 45,
- 263 1387–1394. <u>doi:10.1002/2017GL075821</u>
- 264 Wakita, H. (1975). Water wells as possible indicators of tectonic strain. *Science*, 189,
- 265 553–555. <u>doi:10.1126/science.189.4202.553</u>
- 266 Wang, C.-Y., Cheng, L.-H., Chin, C.-V., & Yu, S.-B. (2001). Co-seismic hydrologic
- response of an alluvial fan to the 1999 Chi-Chi earthquake, Taiwan. Geology, 29(9), 831-
- 268 834. doi:10.1130/0091-7613(2001)029%3C0831:CHROAA%3E2.0.CO;2

- 269 Wang, C.-Y., Wang, C.-H., & Manga, M. (2004). Co-seismic release of water from
- 270 mountains: Evidence from the 1999 (Mw = 7.5) Chi-Chi, Taiwan, earthquake. *Geology*,
- 271 *32*, 769–772. <u>doi:10.1130/G20753.1</u>
- 272 Wang, C.-Y., & Chia, Y. (2008). Mechanism of water levels changes during earthquakes:
- 273 Near field versus intermediate field. *Geophysical Research Letters*, 35, L12402.
- 274 <u>doi:10.1029/2008GL034227</u>
- 275 Wang, C.-Y., & Manga, M. (2015). New streams and springs after the 2014 Mw 6.0
- 276 South Napa earthquake. *Nature Communications*, 6, 7597. doi:10.1038/ncomms8597
- 277 Weaver, K. C., Doan, M. L., Cox, S. C., Townend, J., & Holden, C. (2019). Tidal
- 278 Behavior and Water-Level Changes in Gravel Aquifers in Response to Multiple
- 279 Earthquakes: A Case Study From New Zealand. Water Resources Research, 55(2): 1263-
- 280 1278. https://doi.org/10.1029/2018WR022784
- 281 Zhang, Y., Fu, L. Y., Huang, F., & Chen, X. (2015). Coseismic water-level changes in a
- well induced by teleseismic waves from three large earthquakes. *Tectonophysics*, 651-
- 283 652, 232–241. https://doi.org/10.1016/j.tecto.2015.02.027
- 284