Upscaling of solute plumes in periodic porous media through a trajectory based spatial Markov model

Bianchi Janetti Emanuela¹, Sherman Thomas², Guedon Gael Raymond³, Bolster Diogo², and Porta Giovanni Michele³

¹Politecnico di Milano, Italy ²University of Notre Dame ³Politecnico di Milano

November 16, 2022

Abstract

We propose an approach to upscale solute transport in spatially periodic porous media. Our methodology relies on pore scale information to predict large scale transport features, including explicit reconstruction of the solute plume, breakthrough curves at fixed distances, and spatial spreading transverse to the main flow direction. The proposed approach is grounded on the recently proposed trajectory-based Spatial Markov model (tSMM), which upscales transport based on information collected from advective-diffusive particle trajectories across one periodic element. In previous works, this model has been applied solely to one-dimensional transport in a single periodic pore geometry. In this work we extend the tSMM to the prediction of multidimensional solute plumes. This is obtained by analyzing the joint space-time probability distribution associated with discrete particles, as yielded by the tSMM. By comparing numerical results from fully resolved simulations and predictions obtained with the tSMM over a wide range of Péclet numbers, we demonstrate that the proposed approach is suitable for modeling transport of conservative and linearly decaying solute species in a realistic pore space and showcase the applicability of the model to predict steady state solute plumes. Additionally, we evaluate the model performance as a function of numerical parameters employed in the tSMM parameterization.

¹ Upscaling of solute plumes in periodic porous media ² through a trajectory based spatial Markov model

Emanuela Bianchi Janetti^a, Thomas Sherman^b, Gaël Raymond Guédon^c,
 Diogo Bolster^b, Giovanni Porta^a

 ^aDipartimento di Ingeneria Civile ed Ambientale, Politecnico di Milano, Milano, Italy
 ^bDept. of Civil and Environmental Engineering and Earth Sciences, University of Notre Dame, IN, USA
 ^cDipartimento di Energia, Politecnico di Milano, Milano, Italy

9 Abstract

We propose an approach to upscale solute transport in spatially periodic porous media. Our methodology relies on pore scale information to predict large scale transport features, including explicit reconstruction of the solute plume, breakthrough curves at fixed distances, and spatial spreading transverse to the main flow direction. The proposed approach is grounded on the recently proposed trajectory-based Spatial Markov model (tSMM), which upscales transport based on information collected from advective-diffusive particle trajectories across one periodic element. In previous works, this model has been applied solely to one-dimensional transport in a single periodic pore geometry. In this work we extend the tSMM to the prediction of multidimensional solute plumes. This is obtained by analyzing the joint space-time probability distribution associated with discrete particles, as yielded by the tSMM. By comparing numerical results from fully resolved simulations and predictions obtained with the tSMM over a wide range of Péclet numbers, we demonstrate that the proposed approach is suitable for modeling transport

Preprint submitted to Elsevier

July 20, 2020

of conservative and linearly decaying solute species in a realistic pore space and showcase the applicability of the model to predict steady state solute plumes. Additionally, we evaluate the model performance as a function of numerical parameters employed in the tSMM parameterization.

¹⁰ Keywords: Upscaling, Porous Media, Transport, Spatial Markov Model,

¹¹ Transverse dispersion

12 **1. Introduction**

Solute transport in porous media is a fundamental problem across many 13 disciplines, including subsurface geological systems and the performance op-14 timization of engineered materials such as filtration membranes. A key chal-15 lenge in this context is to obtain accurate predictions at spatial scales much 16 larger than the ones associated with individual pores without having to re-17 solve the physical and chemical processes taking place within complex pore 18 spaces. This is achieved by upscaled formulations that embed pore scale 19 features into effective parameters and therefore can be employed to predict 20 large scale behaviors. To this end, a classical approach is resorting to a 21 continuum-scale advection-dispersion equation (ADE) [1, 2]. In such a for-22 mulation mechanical dispersion induced by pore scale velocity gradients is 23 modelled through a Fickian-like dispersion term, parameterized via a fourth-24 rank dispersivity tensor. The definition of the dispersivity tensor purely 25 based on pore scale properties presents significant challenges. From a theo-26 retical perspective, the solution of three closure problems is required to fully 27 parameterize solute transport based on pore scale information through vol-28 ume averaging [3]. These separate closures are necessary to isolate and char-29

acterize the separate effects of diffusion and advection on transport. However, 30 even such a detailed approach may not yield reliable predictions due to a lack 31 of separation of scales, violating the assumptions required by the volume 32 averaging method. In such a case, non-Fickian transport features emerge, 33 particularly at relatively short times and distances [4, 5]. Formally, these 34 effects can still be represented with Eulerian nonlocal (integro-differential) 35 models. In principle these models can be derived by applying upscaling ap-36 proaches, such as volume averaging, that can relate pore scale geometry and 37 fluid velocities with the emerging transport dynamics through a set of clo-38 sure differential equations [6]. However, it is often found that resorting to 39 such approaches leads to formidable mathematical and numerical complexity 40 [7, 8], which is associated with i) the numerical resolution of various closure 41 problems and ii) the approximation of integro-differential equations to obtain 42 the desired large scale outputs. 43

A specific problem in the context of solute transport upscaling is posed 44 by the modeling of solute plumes, which correspond to the explicit spatial 45 reconstruction of the solute spatial spread at a given time, or at steady state 46 (i.e., under steady state boundary conditions, such as continuous injection). 47 For instance, the analysis of transverse spreading and mixing of steady state 48 solute plumes has great practical relevance in bioremediation and reactive 49 transport scenarios at field and laboratory scales [9, 10, 11]. In these ap-50 plications the target process is the spreading and mixing of a solute in the 51 direction transverse to a steady flow field characterized by a prevalent direc-52 tion. Following classical ADE-based descriptions, transport in the transverse 53 direction is typically modeled by introducing a dispersivity parameter. This

standard definition typically considers dispersion to be uniquely proportional 55 to advective velocity [1]. This formulation was successfully employed, for ex-56 ample to interpret transport and mixing in microfluidic systems characterized 57 by relatively simple geometries [12]. However, studies performed in the last 58 decade have demonstrated the impact of molecular diffusion on transverse 59 dispersion through experiments and numerical simulations [10, 13, 14, 15]. 60 Such results can be qualitatively linked with analytical and numerical stud-61 ies showing that the dispersion tensor becomes asymmetric in advection-62 dominated scenarios [16, 17]. These studies show that full parameterization 63 of the dispersion tensor can become a troublesome task, particularly in media 64 characterized by a complex and multi-scale pore structure. Additional levels 65 of complexity are introduced when reactive processes are also considered on 66 top of pore scale advective-diffusive transport. 67

Over the last decade it has been recognized that pore-to-continuum up-68 scaling of solute transport can often be conveniently obtained by considering 60 solute velocities (or associated travel times) over fixed spatial increments by 70 means of a Markov chain. This led to the formulation of various flavours 71 of so-called Spatial Markov models [e.g. 18, 19, 20, 21]. The SMM is based 72 on the calculation of the travel time across a fixed distance and a one step 73 correlation existing between successive travel times. By including correla-74 tion the SMM is able to employ information available on a limited portion 75 of the system to predict transport across much larger distances. Notably 76 such an approach is effective in the presence of advection-dominated scenar-77 ios that become challenging to upscale with classical Eulerian approaches. 78 The advantages of employing a spatial Markov approach to obtain the solute 70

breakthrough curve (or first passage time) at a given longitudinal distance has 80 been demonstrated in a number of previous works, relying on both numerical 81 and laboratory scale experimental datasets [e.g. 18, 21, 22, 23]. Several re-82 cent works have discussed methodologies that employ Lagrangian SMM-like 83 approaches to predict solute particles' space-time locations at various scales 84 of observations [24, 25]. Yet, to the best of our knowledge, this approach has 85 not been applied to the explicit space-time reconstruction of solute plumes 86 starting from pore scale properties. In this work we present a methodology to 87 upscale transport of solute plumes in the longitudinal and transverse direc-88 tion via a SMM. We consider periodic media, which are routinely considered 89 as model porous media in theoretical approaches and are employed in engi-90 neering systems across a wide range of applications [26, 27, 28, 29]. Our work 91 starts from a recently proposed trajectory-based SMM (here labelled tSMM) 92 to upscale transport, mixing and surface reactions across porous media made 93 up of periodic elements [30, 31]. To date, the tSMM has been constrained to 94 a highly idealized setting, a periodic wavy channel, which while it displays 95 some characteristics of real porous media cannot represent their full com-96 plexity, such as transverse flow or complex pore size distributions. Building 97 on this, we analyze longitudinal and transverse transport within a realis-98 tic two-dimensional porous domain. Briefly, the specific objectives of this 99 contribution are to i) extend the tSMM to the case of a multi-dimensional 100 unsteady solute transport and ii) yield an efficient and accurate representa-101 tion of transient and steady state solute plumes in porous media based on the 102 multi-dimensional tSMM. This second goal entails a specific methodological 103 challenge, as steady state plumes are typically computationally expensive to 104

¹⁰⁵ simulate with Lagrangian particle-based approaches. Our objective is the
¹⁰⁶ development of an upscaled, parsimonious and computationally affordable
¹⁰⁷ particle-based model for this specific configuration.

108 2. Methodology

109 2.1. Pore scale setting

We consider transport of a passive solute in a periodic two-dimensional porous medium. The medium is composed of a collection of periodic unit cells, whose geometry is represented in Figure 1. The unit cell is artificially created using a stochastic generation procedure as detailed in [21, 32, 33]. The cell properties are reported in Table 1. We assume here to deal with

$$\frac{\text{Porosity}}{0.631} \quad \frac{L_x \,[\text{m}]}{4.096 \times 10^{-3}} \quad \frac{L_y \,[\text{m}]}{4.096 \times 10^{-3}} \quad \frac{\Delta \,[\text{m}]}{2 \times 10^{-6}} \quad \frac{\ell_s \,[\text{m}]}{1 \times 10^{-4}}$$

Table 1: Geometrical characteristics of the unit cell

114

a semi-infinite periodic porous domain, i.e. $x \in [0, \infty)$ and $y \in (-\infty, +\infty)$. Because the medium is composed by an infinite number of identical cells, for convenience we define

$$\hat{x} = x - L_x \operatorname{floor}\left(\frac{x}{L_x}\right), \quad \hat{y} = y + \frac{L_y}{2} - L_y \operatorname{floor}\left(\frac{y}{L_y} + \frac{1}{2}\right)$$
(1)

as a coordinate system referenced to the unit cell, $\hat{x} \in [0, L_x]$ and $\hat{y} \in [0, L_y]$, where L_x and L_y define the dimensions of the unit cell in the x- and ydirections, respectively. The cell is discretized into square pixels of side $\Delta = 2 \times 10^{-6}$ m, which for our example results in a unit cell composed of 2048 × 2048 pixels. The solid and fluid phases are identified by an indicator



Figure 1: Geometry (left) and normalized velocity magnitude $|\mathbf{u}|/|\mathbf{u}|$ (right) in the unit cell.

field I, with I = 1 associated with pore space and I = 0 with solid pixels. The correlation length of the indicator field I is taken as a representative length scale for the pore space and is denoted as ℓ_S . Transport is described by the standard advection-diffusion equation with no flux boundary conditions at the solid-fluid interface

$$\frac{\partial C(\mathbf{x},t)}{\partial t} + \boldsymbol{\nabla} \cdot \left[\mathbf{u}(\mathbf{x})C(\mathbf{x},t) \right] = \boldsymbol{\nabla} \cdot \left[D\boldsymbol{\nabla}C(\mathbf{x},t) \right] \quad \forall \mathbf{x} \in \Gamma_{fluid}, t > t_0$$
$$D\frac{\partial C(\mathbf{x},t)}{\partial n} = 0 \quad \forall \mathbf{x} \in \Sigma_{surface}, t > t_0 \qquad (2)$$
$$C(\mathbf{x},t_0) = C_0$$

where $\mathbf{u} = [u, v]$ is the fluid velocity, C is solute concentration, D is the diffusion coefficient, C_0 is the initial concentration distribution. The velocity \mathbf{u} is obtained by numerically solving the Navier-Stokes equations with

OPENFOAM(R), release v1712 [34] and diffusion is assumed to be known and 131 constant. The velocity is computed assuming periodic boundary conditions 132 on the cell boundaries and the no slip condition on the fluid-solid interface. 133 We impose a uniform pressure gradient along the x-direction, labeling x, y134 as longitudinal and transverse directions, respectively. The Péclet number 135 associated with transport is calculated as $Pe = U\ell_S/D$, where U is the av-136 erage fluid velocity. In our simulation we set $D = 10^{-9} \text{ m}^2/\text{s}$ and we adjust 137 Pe by setting U to the desired value. Note that this is acceptable as our 138 simulations are in a Stokes regime, where inertial effects are negligible with 139 respect to viscous ones. Transport is solved numerically using a Lagrangian 140 particle based random walk method, where the solute plume is discretized 141 into a finite number of N particles. Each particle displaces according to 142

$$x_i^{n+1} = x_i^n + u_i dt + \xi_i \sqrt{2Ddt} \qquad i = 1, ..., N , \qquad (3)$$
$$y_i^{n+1} = y_i^n + v_i dt + \eta_i \sqrt{2Ddt}$$

where dt is a time step that is constant, ξ_i , η_i are independent identically distributed random numbers drawn from normal distributions with zero mean and unit variance. We define a reference time step dt^* according to the following criterion $|\mathbf{d}_{max}| \leq 0.5\Delta$ where

$$|\mathbf{d}_{max}| = \max\left(|\mathbf{u}|\right)dt + 2\sqrt{2Ddt^*} \tag{4}$$

¹⁴⁷ is an estimate of the maximum displacement. No flux boundary conditions¹⁴⁸ at the fluid-solid boundary are imposed as elastic reflections.

149 2.2. Spatial Markov Model

We upscale transport using the framework of the trajectory-based spatial Markov model proposed in [30]. The methodology is based upon pore scale transport trajectories and their associated travel times. In the following, we first describe the pore scale trajectories simulations and then how these are used to parameterize the tSMM.

155 2.2.1. Pore scale trajectories

Parameterization of the model is grounded on the pore scale simulation of 156 a set $S = \{s_1, \ldots, s_{N_s}\}$ of N_s advective-diffusive trajectories, for a specific Pe. 157 These particle trajectories are simulated by solving Eq. (3) across a single 158 cell in the longitudinal direction, i.e., between the inlet location x = 0 and 159 the outlet location $x = L_x$. Figure 2 represents a sample of 100 trajectories 160 across the considered unit cell selected from a flux weighted initial condition 161 and setting initial location of particles distributed along the entire unit cell 162 cross section. 163

For each trajectory s_i we record the travel time τ needed to travel across a 164 distance L_x in the longitudinal direction and the y positions (y_{in}, y_{out}) of the 165 particle at the inlet and outlet as the particle enters and exits the domain. 166 Particles are injected at locations $x = 0, y_{in} \in [-L_y/2, L_y/2]$. Particles may 167 cross into adjacent cells along the y direction, but due to the periodicity 168 of the cell geometry each location y_{out} can be mapped to a corresponding 169 \hat{y}_{out} using Eq. (1). Therefore, the coordinate $y_{out}(s_i)$ can be determined as 170 $y_{out}(s_i) = \hat{y}_{out}(s_i) + \Delta_C(s_i)L_y$ where $\Delta_C(s_i)$ is an integer that indicates the 171 net number of cell transitions in the transverse direction observed for a given 172 s_i trajectory path. We can then compute $\Delta y(s_i) = y_{out}(s_i) - y_{in}(s_i)$. The 173

trajectories are subdivided into N_B equiprobable bins that are assigned by 174 considering the starting locations $y_{in}(s_i)$ in ascending order. This implicitly 175 defines a discretization of the \hat{y} axis in terms of the binning of the trajectories. 176 To exemplify this binning, the trajectories in Figure 2 are subdivided into 10 177 bins, indicated by different colours. The trajectories s_i consider all simulated 178 pathways between the locations $x = 0, y_{in} \in [-L/2, L_y/2]$ and $x = L_x, y_{out} \in$ 179 $(-\infty, +\infty)$. We observe that some trajectories may even travel backwards 180 along x close to the inlet section before traveling downstream, as indicated 181 in the highlighted parts in Figure 2. These effects are due to the combined 182 action of advection and diffusion and are present for both the considered Pe. 183 The comparison between the two considered cases allows for identification 184 of the effects of diffusion on the pore scale trajectory paths. In particular, 185 for Pe = 100 particles explore a wider portion of the pore space than for 186 Pe = 1000.187

188 2.2.2. The tSMM parameterization

The information collected in the parameterization step is then used to build the following trajectory-based Spatial Markov model (tSMM)

$$\begin{aligned} x_i^{k+1} &= (k+1)L_x \\ y_i^{k+1} &= y_i^k + \Delta y \left[s_i^k | \hat{y}_{out}(s_i^{k-1}) \right] \\ t_i^{k+1} &= t_i^k + \tau \left[s_i^k | \hat{y}_{out}(s_i^{k-1}) \right] \end{aligned}$$
(5)

where both y_i^{k+1} and τ_i^{k+1} are determined through a Markov chain, which is related to the transverse location assigned to the particle in the periodic cell



Figure 2: Sample of 100 trajectories employed for model parameterization for (a), (b) Pe = 100 and (c), (d) Pe = 1000, dashed red lines indicate the unit cell boundaries, the trajectories are binned in different colours as a function of the y_{in} location. The two middle panels represent a zoom on the region highlighted in red for the two cases.

during successive steps k, leveraging the information given by the trajectories 193 in s_i . The innovative feature of the model in Eq. (5) with respect to previous 194 implementations [30, 21] is that it allows for predictions of transverse spread-195 ing over successive Markov steps. This is achieved by considering y as a con-196 tinuous variable, i.e., the Markov chain has a longitudinal fixed spatial step 197 L_x while transport along y is considered through the $\Delta y(s_i)$ obtained from 198 the trajectories s_i recorded during the parameterization stage. In essence, at 199 the beginning of the simulation (step k = 0) each particle *i* is assigned to an 200 initial location y_i^0 corresponding to a selected initial or boundary condition 201 (e.g., flux weighted or uniform distribution on the inlet boundary). From 202 this information we select a trajectory s_i^1 , randomly sampling from those 203 whose $y_{in}(s_i)$ lies in the same bin interval as y_i^0 . By selecting the trajectory 204

we also obtain a given travel time $\tau(s_i^1)$ and transverse displacement $\Delta y(s_i^1)$, from which we evaluate $\hat{y}_{out}(s_i^1)$. The latter can be then used to select a new trajectory s_i^2 for the next transition and the procedure can then be repeated for any arbitrary step number k > 0.

209 2.3. Model outputs

To analyze the outputs of our tSMM model, our analysis relies on the following dimensionless space-time coordinates

$$\tilde{x} = \frac{x^k}{L}, \qquad \tilde{y} = \frac{y_c}{L}, \qquad \tilde{t} = \frac{tU}{L}$$
(6)

where $L = L_x = L_y$, x^k corresponds to the longitudinal spatial coordinate of k^{th} Markov step (see Eq. (5)) and $y_c(x) = y(x) - \overline{y}(x)$, i.e., is the transverse location centered with respect to the average transverse position $\overline{y}(x)$ observed at a given x. The value of $\overline{y}(x)$ is not constant with x because the average transverse velocity component is not exactly equal to zero. While this component is only approximately 1% of the longitudinal mean velocity U, it still induces plume migration along y after a number of cells.

The key output of the tSMM is the joint probability distribution $P(\tilde{x}, \tilde{y}, \tilde{t})$. In this distribution, the variable \tilde{x} can only assume discrete values, while the \tilde{y}, \tilde{t} are continuous.

Physically meaningful information related to the plume can then be extracted from this joint probability by considering conditional and marginal distributions. In our analysis we will consider the conditional distributions $P(\tilde{t}, \tilde{y}|\tilde{x})$ for a given dimensionless downstream distance \tilde{x} , or $P(\tilde{x}, \tilde{y}|\tilde{t})$ for a given dimensionless time, \tilde{t} . Examples of these conditional distributions ²²⁷ $P(\tilde{t}, \tilde{y}|\tilde{x})$ for $\tilde{x} = 5$, 10, 25 and 50 and $P(\tilde{x}, \tilde{y}|\tilde{t})$ for $\tilde{t} = 20$, obtained from the ²²⁸ tSMM are shown in Figure 3a and b for Pe = 100 and 1000, respectively.

In addition, we consider the marginal probability distributions $P(\tilde{x}, \tilde{y})$, 229 and $P(\tilde{t}|\tilde{x}), P(\tilde{y}|\tilde{x})$ conditional to a given dimensionless downstream distance. 230 These distributions have a clear physical meaning: the marginal distribution 231 $P(\tilde{x}, \tilde{y})$ represents the steady state distribution of the particle plume, while 232 $P(\tilde{t}|\tilde{x})$ corresponds to the breakthrough curve, i.e., the first passage time 233 probability distribution at distance \tilde{x} . Finally, the probability distribution 234 $P(\tilde{y}|\tilde{x})$ provides the probability distribution associated with transverse po-235 sition at a control plane and is related solely to transport in the transverse 236 direction. 237

To produce benchmark data against which to test the tSMM we run a 50 cell high resolution random walk direct numerical simulation (DNS) using transport Eq. (3). The accuracy of the tSMM defined in section 2.2 will be tested by comparing the above mentioned probability distributions with their analogs obtained from the DNS. For each of the above defined distributions we provide a quantitative evaluation of the mismatch between DNS and tSMM using the Hellinger distance [35]

$$HD[F_1, F_2] = \frac{1}{\sqrt{2}} \sqrt{\sum_{i=1}^{N} \left(\sqrt{f_{1,i}} - \sqrt{f_{2,i}}\right)^2}$$
(7)

where F_1 is any of the above-mentioned marginal or conditional distributions predicted by DNS and F_2 corresponds to its counterpart obtained with the tSMM. These distributions are approximated through N discrete bins and $f_{1,i}$, $f_{2,i}$ are the values of the distributions in the i^{th} bin.



Figure 3: Conditional joint distributions $P(\tilde{t}, \tilde{y}|\tilde{x})$ for $\tilde{x} = 5$, 10, 25, 50 and $P(\tilde{x}, \tilde{y}|\tilde{t})$ for $\tilde{t} = 20$ predicted by the tSMM considering (a) Pe = 100 and (b) Pe = 1000

The HD metric quantifies the distance between two probability measures and it is a proper distance metric in the mathematical sense, by satisfying the properties of non-negativity, symmetry, and triangle inequality. HD is also bounded between 0 and 1, where 0 means that the two distributions are indiscernible and 1 that they are maximally distant.

254 3. Results

First we show results obtained considering parameterization of the tSMM 255 with $N_s = 10^6$, $N_B = 100$, $dt^* = 10^{-5}$ s and 10^{-6} s for Pe = 100 and 256 Pe = 1000, respectively. The time step dt^* indicates the values of dt eval-257 uated according to Eq. (4). Then, in section 3.3 we analyze the impact of 258 parameters N_B and dt on the accuracy of the tSMM. For all cases, including 259 the DNS reference simulation and tSMM, we impose a flux weighted bound-260 ary conditions. Note that In the reference DNS simulation the dt parameter 261 is kept constant and equal to dt^* . 262

²⁶³ 3.1. Model performance as a function of Pe

The tSMM is able to replicate the shape of the reference conditional dis-264 tribution $P(\tilde{t}, \tilde{y}|\tilde{x})$ obtained from the DNS for both investigated Péclet num-265 bers. Figure 4a shows the joint distribution $P(\tilde{t}, \tilde{y}|\tilde{x})$ for $\tilde{x} = 25$ and Pe = 100266 from the high resolution direct numerical simulations and corresponding re-267 sults obtained with the tSMM (Figure 4b). The agreement between the two 268 solutions is significant for all transverse coordinates, \tilde{y} , and dimensionless 269 travel times, \tilde{t} . Analogous results are obtained for Pe = 1000 (see Figure 4c 270 and d) and for all other investigated Markov steps (not shown). For both Pe271 values the maximum value of the probability distributions is found at $\tilde{y} \approx 0$. 272 Note that \tilde{y} locations associated with zero probability across the whole time 273 window correspond to the occurrence of solid along the considered transverse 274 section. 275

To quantify the accuracy of the tSMM outputs with respect to the reference DNS, Figure 5 shows the metric $HD[P_{DNS}(\tilde{y}, \tilde{t}|\tilde{x}); P_{tSMM}(\tilde{y}, \tilde{t}|\tilde{x})]$ defined in Eq. (7) for both investigated Pe numbers and all Markov steps. We note that the distance between the DNS and the tSMM distributions slightly increases with \tilde{x} and is generally larger for Pe = 100 than for Pe = 1000. This result is likely due to the fact that the effect of noise in low probability values increases with the strength of diffusion.

Figure 6 depicts conditional joint distribution $P(\tilde{x}, \tilde{y}|\tilde{t})$ for $\tilde{t} = 20$, corresponding to the time dependent solute plume. DNS and tSMM predictions are shown for Pe = 100 (see Figure 6a, b) and for Pe = 1000 (see Figure 6c, d). Again, the tSMM is able to capture all essential features displayed by the fully resolved simulations. Note that tSMM allows for predictions



Figure 4: Conditional joint distributions $P(\tilde{t}, \tilde{y}|\tilde{x})$ for $\tilde{x} = 25$ and Pe = 100 obtained with (a) DNS, (b) tSMM and for Pe = 1000 predicted by (c) DNS and (d) tSMM.



Figure 5: $HD[P_{DNS}(\tilde{y}, \tilde{t}|\tilde{x}); P_{tSMM}(\tilde{y}, \tilde{t}|\tilde{x})]$ as a function of the downstream location \tilde{x} for Pe = 100 (red) and 1000 (blue).

of the evolution of the plume in the longitudinal and transverse directions employing a significantly smaller computational effort than the DNS. As an example, the computational time for running the tSMM is approximately 1% of that one needed for the DNS results when considering 50 Markov steps. Note that this percentage decreases for simulation across higher numbers of unit cells (i.e., the computational gain increases with the dimension of the system of interest).

The marginal distributions $P(\tilde{x}, \tilde{y})$ are depicted in Figure 7a and b for Pe = 100 and 1000 respectively. As mentioned above, these distributions identify the steady-state plume for a nonreactive solute. These distributions are here obtained at no additional computational cost with respect to the transient case, which for a DNS would not be the case and significant additional cost would be required.

Figure 8 displays breakthrough curves $P(\tilde{t}|\tilde{x})$ considering travel distances $\tilde{x} = 5, 10, 25$ and 50 from the injection location and provides a quantitative comparison between the reference DNS and the tSMM results. We note that



Figure 6: Conditional distributions $P(\tilde{x}, \tilde{y}|\tilde{t})$ for $\tilde{t} = 20$ and Pe = 100 obtained with (a) DNS (b) tSMM and for Pe = 1000 predicted by (c) DNS and (d) tSMM



Figure 7: Marginal distribution $P(\tilde{x}, \tilde{y})$ for (a) Pe = 100 and (b) Pe = 1000.

the tSMM can reproduce the breakthrough curves across a wide range of distances and both Péclet numbers. This result shows that the trajectorybased upscaled model accurately predicts arrival times in a porous medium made of periodic unit cells displaying a disordered geometry and is in line with those obtained within simpler geometrical settings [30].

The comparison between tSMM and DNS marginal distribution of transverse locations $P(\tilde{y}|\tilde{x})$ is shown in Figure 9 for two selected distances from the injection ($\tilde{x} = 5$ and $\tilde{x} = 25$) and for both investigated *Pe* numbers. To compare the spreading of the particle plume over all Markov steps we consider the standard deviation of distribution $P(\tilde{y}|\tilde{x})$ as a function of \tilde{x} (see Figure 10).

Results obtained through the tSMM are in close agreement with those yielded by the DNS. We observe that the change in Pe has marked effects on transverse spreading of the solute, as has been previously observed in laboratory and numerical studies [10, 13]. In particular, the standard deviation σ_y



Figure 8: Breakthrough curves $P(\tilde{t}|\tilde{x})$ obtained at control planes located at distances $\tilde{x} = 5, 10, 25, 50$ unit cells for (a) Pe = 100 and (b) Pe = 1000. Symbols and lines represent the DNS and tSMM results, respectively.



Figure 9: Comparison between the marginal distributions $P(\tilde{y}|\tilde{x})$ as given by direct numerical simulation (black lines) and tSMM (red dots) for $\tilde{x} = 5$ and a) Pe = 100, b) Pe = 1000, for $\tilde{x} = 25$ and c) Pe = 100 and d) Pe = 1000.



Figure 10: Comparison of the standard deviation, σ_y , of the distribution $P(\tilde{y}|\tilde{x})$ for the DNS and tSMM as a function of \tilde{x} .

continuously increases as a function of the longitudinal distance traveled for 319 Pe = 100. On the contrary, σ_y is approximately constant up to $\tilde{x} = 20$ for 320 Pe = 1000 and then starts increasing. This result implies that for such an 321 advection-dominated situation we only observe significant transverse spread-322 ing after particles have traveled a distance of 20 cells. This result is due to 323 the converging-diverging nature of advective streamlines in two-dimensional 324 fields, and this particular behavior might be different if investigated in three-325 dimensions. Yet, the tSMM is able to predict these different dynamics based 326 on the simulation of transport across a single unit cell. Note also that the 327 methodology can extended to three dimensions, upon relying on the same 328 procedure described in Section 2.2. 329

330 3.2. Steady state plumes with first order degradation

As a showcase application of the capabilities of the tSMM, we also evaluate the influence of a first order reaction on pinching off the steady-state plume. We assume in this application that the solute undergoes degrada-

tion following linear kinetics. This is accounted for in a straightforward 334 manner in the tSMM framework. Starting from the conservative plume re-335 sults (see Figure 7), for each particle we define the probability of reaction as 336 $R_i(t,\lambda) = 1 - e^{-\lambda t}$, where λ is the kinetic degradation rate. Then we compare 337 R_i with a random number, U_i , drawn from a standard uniform distribution. 338 If $U_i \ge P_i$ no reaction occurs while if $U_i < P_i$ the particle is removed from 339 the system. In our examples λ is chosen based on obtaining specific values 340 of Damkhöler numbers, $Da = \frac{\ell_S^2 \lambda}{D}$. 341

We compute marginal distributions $P(\tilde{x}, \tilde{y})$ to represent the steady state 342 plume for this reactive scenario. The results obtained for Da = 1 and 5 343 (corresponding to $\lambda = 0.1$ and 0.5) and Pe = 100 are depicted in Figures 11a 344 and b while Figures 11c and d show the case associated with Da = 5 and 10 345 (corresponding to $\lambda = 0.5$ and 1) and Pe = 1000. These results portray the 346 ability of our proposed tSMM to predict transport in longitudinal and trans-347 verse direction while also accounting for a reactive solute undergoing a first 348 order reaction process. Note that the results are obtained at negligible addi-340 tional computational cost with respect to the unsteady state, conservative, 350 transport simulations. 351

352 3.3. Error analysis

To provide a quantitative description of the influence of parameters N_B and dt on the accuracy of the proposed spatial Markov model we evaluate $HD[F_1, F_2]$, see Eq. 7, choosing as F_1 and F_2 the marginal distributions of travel times, $P(\tilde{t}|\tilde{x})$, or transverse positions, $P(\tilde{y}|\tilde{x})$, conditional to a given downstream location \tilde{x} from the injection point, evaluated with DNS and tSMM respectively. Note that the time step is kept constant and equal to



Figure 11: Marginal distribution $P(\tilde{x}, \tilde{y})$ for a reactive contaminant undergoing a degradation following a linear kinetics for Pe = 100 (a) Da = 1, (b) Da = 5, and Pe = 1000 (c) Da = 5, (d) Da = 10.

 dt^* in the DNS, while we consider in following different values for the simulation of the trajectories employed to parameterize the tSMM, To simplify the notation we introduce here the following indicators

$$HD_t = HD[P_{DNS}(\tilde{t}|\tilde{x}); P_{SMM}(\tilde{t}|\tilde{x})]$$
(8)

$$HD_y = HD[P_{DNS}(\tilde{y}|\tilde{x}); P_{SMM}(\tilde{y}|\tilde{x})]$$
(9)

to assess the model errors. The analysis is performed considering both 353 Pe = 100 and Pe = 1000 and considering different values of N_B and dt354 for the parameterization of the tSMM, while, as mentioned above, in the ref-355 erence DNS the dt is constant and equal to dt^* . In order to provide an overall 356 assessment of the impact of parameters dt and N_B we focus on the average 357 of HD_{α} (with $\alpha = t, y$) across all 50 investigated Markov steps (the averag-358 ing operator is denoted by the symbol $\langle \cdot \rangle$). Figures 12a and b show $\langle HD_t \rangle$ 359 and $\langle HD_y \rangle$ as a function of dt/dt^* and N_B , respectively. Continuous lines 360 correspond to Pe = 100, while dashed lines depict results associated with 361 Pe = 1000. Red and blue colors are related to arrival time and transverse 362 location distributions, respectively. Note that for Pe = 100 both $\langle HD_t \rangle$ and 363 $\langle HD_y \rangle$ are not very sensitive to the choice of parameter dt (see Figure 12a). 364 On the contrary for Pe = 1000 we observe a sharp increase of $\langle HD_t \rangle$ and 365 $\langle HD_y \rangle$ for $dt > 2dt^*$. This is probably due to the fact that advective particle 366 displacements depend linearly on dt, as opposed to the diffusive ones which 367 scale with $dt^{0.5}$. The variation of $\langle HD_t \rangle$ and $\langle HD_y \rangle$ as a function of N_B are 368 displayed in Figure 12b for the two investigated Pe numbers. These results 369 show that the quality of model predictions deteriorates for decreasing num-370



Figure 12: $\langle HD_{\alpha} \rangle$ for Pe = 100 and Pe = 1000 as a function of the Markov step number (cell number) for different dt and numbers of bins N_B employed in the parameterization step.

ber of bins associated with the tSMM parameterization. Note that $N_B = 1$ 371 corresponds to considering the particle trajectories as totally uncorrelated 372 across successive Markov steps, while setting $N_B > 1$ in the tSMM param-373 eterization allows consideration of correlation between particle trajectories 374 belonging to the same bin. We note that considering uncorrelated particle 375 trajectories does not provide good agreement between DNS and tSMM dis-376 tributions, as indicated by high values of the HD_{α} metrics. Also in this case 377 correlation effects at a fixed downstream distance become stronger as Péclet 378 number increases, in line with the results of previous studies [22, 31]. 379

380 4. Discussion and conclusions

Our study proposes a methodology for upscaling solute plumes in periodic porous media through a trajectory based spatial Markov Model. We extend the work of [30] to the case of a multi-dimensional unsteady solute transport and exemplify our approach considering a two-dimensional porous medium with a disordered geometry. Our framework is based on the simulation of advection-diffusion random walk particle trajectories across a single periodic flow cell with the aim of predicting transport over a much larger scale. In particular, our analysis explicitly includes the evaluation of the joint space-time probability distributions associated with solute plumes providing an efficient and accurate representation of both transient and steady state transport in porous media for different *Pe* numbers. Our work leads to the following major conclusions:

1. From a comparison with high resolution direct numerical simulations 393 we show that the proposed tSMM accurately predicts spatial and tem-394 poral distributions of a conservative solute plume using information 395 collected from a single cell simulation. The current framework is not 396 restricted to a simplified geometry setting or a particular flow condition, 397 but can be employed to accurately predict multi-dimensional transport 398 in a realistic two-dimensional pore space once the flow field has been 399 evaluated. Note that, in principle our methodology can be used with 400 different type of initial injection condition, e.g. flux-weighted or uni-401 formly distributed, pulse or continuous injection, and can be extended 402 to a three-dimensional setup. 403

2. Our model is able to predict different transport dynamics, particularly
regarding the influence of *Pe* on transverse plume spreading. Our results are in line with previous laboratory and numerical studies [10, 13].
In particular, the change in *Pe*, due to a change in the diffusion coefficient of the compound, had marked effects on transverse spreading of
the solute and this is well captured by our upscaling approach.

410 3. Particle trajectories and associated travel times, which were simulated

with the proposed tSMM to predict conservative transport, can be nat-411 urally extended to simulate reactive transport processes with negligible 412 additional computational cost. As an example, in this work we analyse 413 the influence of a first order kinetic reaction on a steady state plume. 414 To do so our model has been coupled with a probabilistic representation 415 of a linear degradation reaction and applied for several Da numbers. 416 Note that the methodology is already fully compatible with the analy-417 sis of other types of reactions (e.g. sorption/desorption) as was shown 418 by [31] for an idealized benchmark problem. 419

4. The effect of tSMM parameterization (in particular the parameters 420 dt and N_B) was studied for the two analysed Pe. As expected, the 421 difference between tSMM and DNS distribution both in space and time 422 increases with increasing dt and decreasing N_B respectively. A marked 423 effect of parameterization was observed for Pe = 1000 with respect 424 to Pe = 100 due to fact that (i) the contribution of the advective 425 transport, which increases with Pe, is more affected by the choice of 426 dt employed in the tSMM parameterizazion and (ii) the relevance of 427 particle trajectory correlation increases with Pe. 428

For all the above points the simulation time needed for the tSMM is negligible if compared to the one required for high resolution direct numerical simulation. This is one of the great advantages of the proposed tSMM which allows predicts of multi-dimensional transport across large distances (for both conservative and reactive solutes) without the burden of excessive computational resources. Note that, in its current form, the model can be applied to a spatially periodic domain. This restriction is similar to the fact that

solving a closure problem on a periodic unit cell is required for many well es-436 tablished upscaling procedures. At the same time, broadening the scenarios 437 of interest, for example extending the methodology to upscaling transport of 438 a conservative and/or reactive solute in a disordered non periodic porous do-439 main would constitute additional elements of interest which are compatible 440 with the approach we rest upon. A first attempt in this direction has been 441 provided by [21, 36] obtaining promising results which can be advanced in 442 the context of future investigations. 443

444 Acknowledgments

445 Data Availability Statement

446 Data sets for this research are available online at

447 https://data.mendeley.com/datasets/rzg53tn963/draft

448 References

- [1] A. Scheidegger, Statistical hydrodynamics in porous media, Journal of
 Applied Physics 25 (8) (1954) 994–1001. doi:10.1063/1.1721815.
- [2] A. H.-D. Cheng, J. Bear, Modeling Groundwater Flow and Contaminant
 Transport, Springer Publishing Company, 2016.
- [3] F. Valdés-Parada, D. Lasseux, F. Bellet, A new formulation of the dispersion tensor in homogeneous porous media, Advances in Water Resources 90 (2016) 70-82. doi:10.1016/j.advwatres.2016.02.012.
- [4] J. Salles, J.-F. Thovert, R. Delannay, L. Prevors, J.-L. Auriault,
 P. Adler, Taylor dispersion in porous media. determination of the dispersion tensor, Physics of Fluids A 5 (10) (1992) 2348–2376.
- [5] B. Berkowitz, A. Cortis, M. Dentz, H. Scher, Modeling non-fickian transport in geological formations as a continuous time random walk, Rev.
 Geophys. 44 (2).
- [6] B. Wood, F. Valdés-Parada, Volume averaging: Local and nonlocal closures using a green's function approach, Advances in Water Resources
 51 (2013) 139–167. doi:10.1016/j.advwatres.2012.06.008.
- [7] Y. Davit, B. D. Wood, G. Debenest, M. Quintard, Correspondence between one-and two-equation models for solute transport in two-region
 heterogeneous porous media, Transport in porous media 95 (1) (2012)
 213–238.

- [8] G. Porta, G. Ceriotti, J.-F. Thovert, Comparative assessment of
 continuum-scale models of bimolecular reactive transport in porous media under pre-asymptotic conditions, Journal of contaminant hydrology
 185 (2016) 1–13.
- [9] O. Cirpka, A. Valocchi, Two-dimensional concentration distribution for
 mixing-controlled bioreactive transport in steady state, Advances in Water Resources 30 (6-7) (2007) 1668–1679. doi:10.1016/j.advwatres.
 2006.05.022.
- [10] G. Chiogna, C. Eberhardt, P. Grathwohl, O. Cirpka, M. Rolle, Evidence
 of compound-dependent hydrodynamic and mechanical transverse dispersion by multitracer laboratory experiments, Environmental Science
 and Technology 44 (2) (2010) 688–693. doi:10.1021/es9023964.
- [11] Y. Tang, C. Werth, R. Sanford, R. Singh, K. Michelson, M. Nobu, W.-T.
 Liu, A. Valocchi, Immobilization of selenite via two parallel pathways
 during in situ bioremediation, Environmental Science and Technology
 484 49 (7) (2015) 4543-4550. doi:10.1021/es506107r.
- [12] T. Willingham, C. Werth, A. Valocchi, Evaluation of the effects of
 porous media structure on mixing-controlled reactions using pore-scale
 modeling and micromodel experiments, Environmental Science and
 Technology 42 (9) (2008) 3185–3193. doi:10.1021/es7022835.
- [13] M. Rolle, D. Hochstetler, G. Chiogna, P. K. Kitanidis, P. Grathwohl,
 Experimental investigation and pore-scale modeling interpretation of
 compound-specific transverse dispersion in porous media, Transport in

- 492 Porous Media 93 (3) (2012) 347-362. doi:https://doi.org/10.1016/
 493 j.advwatres.2020.103574.
- [14] M. Muniruzzaman, M. Rolle, Impact of multicomponent ionic transport
 on ph fronts propagation in saturated porous media, Water Resources
 Research 51 (8) (2015) 6739–6755. doi:10.1002/2015WR017134.
- ⁴⁹⁷ [15] Y. Ye, G. Chiogna, O. Cirpka, P. Grathwohl, M. Rolle, Experimental in⁴⁹⁸ vestigation of transverse mixing in porous media under helical flow con⁴⁹⁹ ditions, Physical Review E 94 (1). doi:10.1103/PhysRevE.94.013113.
- [16] J.-L. Auriault, C. Moyne, H. Souto, On the asymmetry of the dispersion
 tensor in porous media, Transport in Porous Media 85 (3) (2010) 771–
 783. doi:10.1007/s11242-010-9591-y.
- ⁵⁰³ [17] S. Pride, D. Vasco, E. Flekkoy, R. Holtzman, Dispersive transport and
 ⁵⁰⁴ symmetry of the dispersion tensor in porous media, Physical Review E
 ⁵⁰⁵ 95 (4). doi:10.1103/PhysRevE.95.043103.
- ⁵⁰⁶ [18] T. Le Borgne, D. Bolster, M. Dentz, P. Anna, A. Tartakovsky, Effective
 ⁵⁰⁷ pore-scale dispersion upscaling with a correlated continuous time ran⁵⁰⁸ dom walk approach, Water Resources Research 47 (12) (2011) W12538.
- [19] P. K. Kang, P. Anna, J. P. Nunes, B. Bijeljic, M. J. Blunt, R. Juanes,
 Pore-scale intermittent velocity structure underpinning anomalous
 transport through 3-d porous media, Geophysical Research Letters
 41 (17) (2014) 6184–6190.
- ⁵¹³ [20] A. Puyguiraud, P. Gouze, M. Dentz, Stochastic dynamics of lagrangian

- ⁵¹⁴ pore-scale velocities in three-dimensional porous media, Water Re ⁵¹⁵ sources Researchdoi:10.1029/2018WR023702.
- [21] T. Sherman, E. B. Janetti, G. R. Guédon, G. Porta, D. Bolster, Upscaling transport of a sorbing solute in disordered non periodic porous domains, Advances in Water Resources (2020) 103574doi:https://doi.
 org/10.1016/j.advwatres.2020.103574.
- [22] D. Bolster, Y. Méheust, T. Le Borgne, J. Bouquain, P. Davy, Modeling preasymptotic transport in flows with significant inertial and trapping effects-the importance of velocity correlations and a spatial markov
 model, Advances Water Resources 70 (2014) 89–103.
- ⁵²⁴ [23] T. Sherman, A. Foster, D. Bolster, K. Singha, Predicting downstream
 ⁵²⁵ concentration histories from upstream data in column experiments, Wa⁵²⁶ ter Resources Research (2018) 9684–9694.
- ⁵²⁷ [24] A. Russian, M. Dentz, P. Gouze, Time domain random walks for hydro⁵²⁸ dynamic transport in heterogeneous media, Water Resources Research
 ⁵²⁹ 52 (5) (2016) 3309–3323. doi:10.1002/2015WR018511.
- [25] E. Wright, N. Sund, D. Richter, G. Porta, D. Bolster, Upscaling mixing
 in highly heterogeneous porous media via a spatial markov model, Water
 11 (1) (2019) 53.
- [26] M. Schmuck, P. Berg, Effective macroscopic equations for species transport and reactions in porous catalyst layers, Journal of the Electrochemical Society 161 (8) (2014) E3323–E3327. doi:10.1149/2.037408jes.

- [27] H. Kim, J. Bae, D. Choi, An analysis for a molten carbonate fuel cell
 of complex geometry using three-dimensional transport equations with
 electrochemical reactions, International Journal of Hydrogen Energy
 38 (11) (2013) 4782-4791. doi:10.1016/j.ijhydene.2013.01.061.
- [28] T. Gebäck, A. Heintz, A pore scale model for osmotic flow: Homogenization and lattice boltzmann simulations, Transport in Porous Media
 126 (1) (2019) 161–176. doi:10.1007/s11242-017-0975-0.
- [29] F. Municchi, M. Icardi, Macroscopic models for filtration and heterogeneous reactions in porous media, Advances in Water Resources 141
 (2020) 103605. doi:https://doi.org/10.1016/j.advwatres.2020.
 103605.
- [30] N. L. Sund, G. M. Porta, D. Bolster, Upscaling of dilution and mixing
 using a trajectory based spatial markov random walk model in a periodic
 flow domain, Advances in Water Resources 103 (2017) 76–85.
- [31] T. Sherman, A. Paster, G. Porta, D. Bolster, A spatial markov model
 for upscaling transport of adsorbing-desorbing solutes, Journal of con taminant hydrology 222 (2019) 31–40.
- [32] P. K. Smolarkiewicz, C. L. Winter, Pores resolving simulation of Darcy
 flows, Journal of Computational Physics 229 (9) (2010) 3121–3133.
- [33] J. D. Hyman, C. L. Winter, Stochastic generation of explicit pore structures by thresholding Gaussian random fields, Journal of Computational
 Physics 277 (2014) 16–31.

- [34] OpenCFD Limited, OpenFOAM The Open Source CFD Toolbox, user
 guide version v1712, https://www.openfoam.com/, accessed: 2019-0929 (2017).
- [35] E. Hellinger, Neue begrundung der theorie quadratischer formen von un endlichvielen veränderlichen., Journal Für Die Reine Und Angewandte
 Mathematik (1909) 210–271doi:10.1515/crll.1909.136.210.
- ⁵⁶⁴ [36] T. Sherman, J. Hyman, M. Dentz, D. Bolster, Characterizing the influ-
- ence of fracture density on network scale transport, Journal of Geophys-
- ical Research: Solid Earth 125 (1) (2020) e2019JB018547.