Exotic Plantations Increase Risks of Flooding in Mountainous Landscapes

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Abstract

We examined the effect of land cover on stream discharge in hilly catchment streams during extreme rain events. Three years of rainfall-runoff observations, between January 2014 and December 2016, were collected in eleven neighbouring catchments. Each catchment was dominated by a different land cover, namely natural shola forests, natural grasslands and wattle (Acacia mearnsii). Rain intensities between percentiles 25-90, 90-95 and over 95 were categorised as light, heavy and extreme and were used to study stream discharge responses. Land cover significantly influenced the hydrologic response to extreme rain events. During light rains (< 38 mm/day), grassland dominated catchments showed higher discharge than shola (0.01 mm/s) and wattle (0.004 mm/s). However, during extreme rain events (> 71 mm/day) discharge was significantly higher in wattle dominated catchments when compared to the natural shola (0.033 mm/s) and grasslands (0.023 mm/s). Antecedent moisture conditions played a major role in determining peak flows along with rainfall, catchment shape and drainage density.

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Key Points:

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Stream flow is faster in wattle dominated catchments during extreme rain than native grassland or shola forest and may contribute to floods. Land cover, along with antecedent moisture and topographic conditions determine rain-runoff responses during extreme events. Hydrologic footprint of exotic invasives have consequences for ecosystems and human well being which outweigh perceived carbon benefits.

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15 Abstract

We examined the effect of land cover on stream discharge in hilly catchment streams dur-16 ing extreme rain events. Three years of rainfall-runoff observations, between January 2014 17 and December 2016, were collected in eleven neighbouring catchments. Each catchment 18 was dominated by a different land cover, namely natural shola forests, natural grasslands 19 and wattle (Acacia mearnsii). Rain intensities between percentiles 25-90, 90-95 and over 20 95 were categorised as light, heavy and extreme and were used to study stream discharge 21 responses. Land cover significantly influenced the hydrologic response to extreme rain 22 events. During light rains (< 38 mm/day), grassland dominated catchments showed higher 23 discharge than shola (0.01 mm/s) and wattle (0.004 mm/s). However, during extreme 24 rain events (> 71 mm/day) discharge was significantly higher in wattle dominated catch-25 ments when compared to the natural shola (0.033 mm/s) and grasslands (0.023 mm/s). 26 Antecedent moisture conditions played a major role in determining peak flows along with 27 rainfall, catchment shape and drainage density. 28

²⁹ Plain Language Summary

Increasing frequency of extreme rain events is a cause of concern as they often trig-30 ger floods and consequent damage. We found that catchments dominated by wattle plan-31 tations, an invasive alien species, cause significantly quicker discharge during extreme 32 rain events when compared with the natural grassland and montane forest (shola) mo-33 saics. Our study, located in the Upper Nilgiris in the Western Ghats mountains of South-34 ern India, also found that stream-flows in wattle dominated catchments were lower than 35 those of grassland dominated catchment during the dry season. We demonstrate that 36 invasive wattle plantations have a significant hydrologic footprint which alters the rainfall-37 runoff behaviour of catchments in the Western Ghats. Widespread plantations of exotic 38 species in the region, which include Eucalyptus, various acacia and pines, could have se-39 rious hydrologic consequences and could similarly alter rain-runoff response by exacer-40 bating floods during the monsoon and reducing stream-flow and hydro-power generation 41 during the dry season. 42

43 1 Introduction

Global climate change scenarios often show an increased frequency of extreme events,
 particularly rainfall, which is a major concern worldwide (Goswami et al., 2006; Guhathakurta

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et al., 2011; Mason et al., 1999; Osborn et al., 2000; Zhou et al., 2013). Such extreme 46 rain events (ERE) often lead to destructive floods causing extensive loss of lives and prop-47 erty (Fowler & Kilsby, 2003; Guhathakurta et al., 2011; Mishra & Shah, 2018; Ranger 48 et al., 2011). During heavy rains, the saturated hydrologic conductivity of soils is quickly 49 exceeded (Koutny et al., 2014) and sub-surface flow pathways get activated (Chappell 50 et al., 2017; Bonell et al., 2010), resulting in higher and quicker discharge from the basin. 51 The problem is more serious in mountainous terrain where steep slopes accelerate the 52 accumulation of stream water leading to a rapid discharge of rain water and sediments 53 downstream (Serrano-Muela et al., 2015). Understanding the relationship between rain-54 fall and discharge can help design mitigation strategies for destructive floods. Such re-55 lationships need to be studied at local scales as most water-flow enters rivers via low-56 order channels and is governed by catchment characteristics and micro-climate (Borga 57 et al., 2014). A large number of local factors such as steepness of slopes, catchment shape 58 and size, drainage networks (D'Odorico & Rigon, 2003; Rinaldo et al., 1991) and antecedent 59 moisture (Chappell et al., 2017; Haga et al., 2005; Kim et al., 2019; Song & Wang, 2019) 60 influence the rainfall-runoff relationship. 61

Vegetative cover is probably the most easily managed characteristic of the catch-62 ment which plays a significant role in mediating the rain-runoff response. Vegetation can 63 alter retention capacity and infiltration of precipitation in headwaters (Koutny et al., 64 2014). Certain vegetation adds organic matter to the soil, arresting erosion by slowing 65 down surface runoff (Bathurst et al., 2011; Koutnỳ et al., 2014; Krishnaswamy et al., 2012). 66 Replacing natural vegetation with fast growing species (Calder & Dye, 2001; Jackson et 67 al., 2005; Venkatesh et al., 2014) including those found in the Western Ghats such as eu-68 calyptus (Sikka et al., 2003; Chand et al., 2009; Sharda et al., 1988; Samraj et al., 1988), 69 wattle (Dye & Jarmain, 2004; Prinsloo & Scott, 1999; Clulow et al., 2011) greatly in-70 creases evapotranspiration and can have serious impacts on stream flow, particularly dur-71 ing the dry season. 72

Understanding the relationship between rainfall and peak discharge in this altered landscape is also critical for flood prediction. The hydrologic impact of wattle in has not been empirically established in the Nilgiris (Rangan et al., 2010) and this study tries to address this gap. Here, we compare the hydrologic response of the two native land covers, grasslands and shola forests, and the introduced invasive wattle. Woody invasives could enhance transpiration and thus reduce antecedent moisture. At the same time their

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⁷⁹ litter may not integrate into soil compared to native vegetation thereby reducing infil-

 $_{80}$ tration (Bonell et al., 2010).

1.1 Study Area

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The Nilgiris or Blue Mountains, 76°-77°15'E: 11°15'-12°15'N, are within the West-82 ern Ghats Biodiversity hot-spot (Myers et al., 2000) and are part of the Nilgiri Biosphere 83 Reserve, the first biosphere-reserve established in India in 1986 (Daniels, 1996). The Nil-84 girls are home to 15 different indigenous tribes and harbour a large number of threat-85 ened and endemic species of flora and fauna (Daniels, 1996). Elevation in the Nilgiris 86 ranges from 1000 m to 2600 m asl which has given rise to diverse vegetation types such 87 as montane-rain-forests in the valleys, locally known as sholas, interspersed by high-altitude 88 grasslands. 89



Figure 1. The Upper Nilgiris showing different catchments and the locations of water level recorders and rain-gauges. Left panel shows elevation gradient and right panel shows land-cover types in different catchments.

This study was conducted in the Nilgiris Reserve Forest, in the Western Ghats mountains in South India (Figure 1). Nilgiris forms an important catchment area for several perennial tributaries of the Cauvery (India - WRIS Project Team, n.d.) including the Bhavani, on which large human populations are dependent downstream. These streams and rivers are managed extensively for power generation, irrigation and drinking water through a chain of dams and reservoirs (*Upper Bhavani Dam D00756* -, n.d.). In the year 2018 and 2019, several parts of the Western Ghats have repeatedly witnessed heavy rains,

and subsequent floods have destroyed several villages and triggered a large number of 97 landslides leading to the loss of many human lives and livelihoods (Mishra et al., 2018; 98 Arathi Menon, 2019; Safi, 2018). Avalanche, in the Nilgiris, where the project site is lo-99 cated, was the epicentre of ERE during August 2019, which resulted in a large number 100 of landslides and floods throughout the region (Premkumar, 2019; TWC India, 2019). 101 Unfortunately, the floods washed away all the water level recorders. Floods and land-102 slides in the Western Ghats have often been attributed to changes in land-cover and land-103 use (Kumar & Bhagavanulu, 2008) over the last century. A large number of exotic plants 104 were introduced in the natural grasslands and sholas, and some of these exotics, partic-105 ularly, wattles, Acacia meansii and Acacia dealbata, have invaded natural grasslands 106 (Joshi et al., 2018). 107

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2 Materials and Methods

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2.1 Data Collection

We conducted our study in the 1000 km^2 area of the Nilgiris Forest Division which 110 covered 11 catchments dominated by three distinct land covers: sholas (2 catchments), 111 montane grasslands (1 catchment) and wattle (8 catchments) (Table 1). We collected 112 and analysed three years of rainfall-runoff data from January 2014 to December 2016. 113 Data from after 2016 was not used because a forest fire in the last weeks of February, 114 2017, completely altered the land cover (Sriramamurthy et al., 2020). Rainfall was mea-115 sured at one minute intervals from 26 tipping bucket rain gauges (RainWise, 2012) placed 116 in an approximate grid of one kilometre. Ten of these were located within the study catch-117 ments. We used the mean rainfall recorded when two or more rain gauges were present 118 in the same catchment. Water levels were measured at five minute intervals in eleven streams 119 instrumented with stilling wells and capacitance probe based water level records (WLRs) 120 (Dataflow Systems Limited, 2017), stage values were converted to discharge using the 121 velocity-area method (Shaw et al., 2010). The streams were low order (1-3) and the to-122 tal catchment area covered was 1,200ha. Digital Elevation Models obtained from SRTM 123 (NASA JPL, 2013) were used to delineate the catchments. Dominant vegetative cover 124 in each of the catchments (henceforth referred as land cover) was estimated from a su-125 pervised vegetation map generated using Landsat 8 images for the year 2017 (additional 126 description in Appendix A). 127

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Catch- ment ID	Land cover	Area (ha)	Elevation range (m)	Shape	Steep- ness	Drainage density (m/m^2)	Rainfall range (mm)
101	Wattle	28.026	2,344 - 2,588	0.4508	1.031	0.0013	1,898 - 5,165
102	Wattle	81.63	2,329 - 2,588	0.6076	0.937	0.0016	1,898 - 5,165
103	Wattle	44.33	2,292 - 2,412	0.6062	0.819	0.0015	1,900 - 5,138
104	Wattle	81	2,290 - 2,412	0.5471	0.793	0.0014	1,900 - 5,138
105	Wattle	101.14	2,325 - 2,588	0.513	0.895	0.0015	1,898 - 5,165
106	Wattle	87.29	2,281 - 2,412	0.5944	0.806	0.0018	1,473 - 4,252
107	Grassland	50.51	2,279 - 2,371	0.5349	0.729	0.0013	1,427 - 3,979
108	Shola	258	2,052 - 2,588	0.5102	1.21	0.0016	2,719 - 7,238
109	Shola	280.03	2,004 - 2,588	0.4291	1.195	0.0017	2,719 - 7,238
114	Wattle	150.97	2,282 - 2,588	0.4631	0.847	0.0016	1,790 - 4,987
115	Wattle	10.76	2,283 - 2,481	0.6051	0.978	0.0001	1,011 - 3,130

Table 1. Land cover and other morphometric characteristics of the catchments.

Note: Dominant vegetative cover of catchment; Shape: Catchment shape measured by the circulatory Index CI; Steepness: Mean steepness of slope factor.

Four different catchment morphological characteristics – shape, area, steepness of 128 the slopes, and drainage density, were derived for each of the catchments. Catchment 129 shape was measured using the circularity index (CI) or ratio which could help forecast-130 ing the flood potential of a basin. The CI is expressed as CI = Ab/Ac, where Ab is the 131 area of the basin and Ac is the area of a circle with the same length of perimeter as the 132 basin (Allaby, 2013). A layer of drainage networks was developed using r.watershed mod-133 ule in GRASS 7 (GRASS Development Team, 2018). Drainage density was expressed 134 as m/m^2 for each catchment. Steepness of the catchment slope was also obtained using 135 r.watershed module which generates slope steepness factor as defined for Universal Soil 136 Loss Equation (McCool et al., 1987). The soil type was similar across all the catchments 137 as per the National Bureau of Soil Survey and Land Use Planning (NBSS&LUP) (Sehgal 138 et al., 1987), and is described as clayey-skeletal, mixed typic haplustalfs and typic Us-139 tropepts (Table 1). 140

141 2.2 Analysis

Data processing and analysis was carried out using the R statistical software (R Core Team, 2018). Daily rainfall was grouped into three categories: i) low intensity (light rain) - rainfall values between 25th percentile ($\geq 1.3 \text{ mm/day}$) and 90th percentile (< 5.0 mm/day); ii) heavy rain - between 90th percentile ($\geq 38.0 \text{ mm/day}$) and 95th percentile (< 71.0 mm/day); and iii) extreme rainfall - 95th percentile and above ($\geq 71.0 \text{ mm/day}$). ¹⁴⁷ We excluded data from the dry season (January to April) from the analysis. Discharge ¹⁴⁸ (m^3s^{-1}) was measured for each stream as mean daily discharge and peak daily discharge ¹⁴⁹ (the maximum discharge recorded in a day). Discharge was then divided by the area of ¹⁵⁰ the corresponding catchment to obtain unit mean and unit peak daily discharges (mm ¹⁵¹ s⁻¹).

We analysed rainfall-runoff records from 8,469 days across all the catchments. An-152 tecedent moisture conditions are known to influence rainfall-runoff relationships (Chappell 153 et al., 2017; Haga et al., 2005; Kim et al., 2019; Song & Wang, 2019). We therefore de-154 rived an antecedent moisture index (AMI), which is an approximation of the moisture 155 stored within a catchment before a particular rainfall event. AMI was developed by de-156 ducting total stream flow from the cumulative rainfall over the 14 antecedent days. We 157 found a strong correlation, r=0.96, between this index and the widely used antecedent 158 precipitation index (API) (Kohler & Linsley, 1951) developed using a decay constant (k) 159 value of 0.9 and considering 14 antecedent days (more details in Appendix B). 160

Scatter plots of (log) daily rainfall versus daily unit mean discharge and daily peak 161 discharge suggested an exponential relationship (Fig 3 (a). This relationship was further 162 analysed with exponential regression models using normalised values for the data. We 163 used additive exponential regression models to test the influence of catchment land cover 164 and other morphometric characteristics on discharge, including daily rainfall and AMI. 165 Variables introducing collinearity in the model were identified using generalised variance 166 inflation factor (GVIF) (Fox & Monette, 1992). Thus, catchment area and steepness of 167 the slope were removed as they were highly correlated $(\text{GVIF}^{(1/(2^*\text{Df}))} > 2.00)$ with land 168 cover, catchment shape and drainage density. The most plausible explanatory variables 169 and the best model for describing relationship between measured runoff and rainfall was 170 selected using Akaike Information Criteria with bias adjustment (AICc) (Burnham & 171 Anderson, 2002). The rainfall only model was used as a null model to compare the re-172 sults. Model averaged coefficients were used whenever delta AICc values of less than 2 173 were obtained and relative variable importance values were used to select the most plau-174 sible explanatory variable. 175

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For anlysing peak discharge, we calculated total rainfall for 24 hours starting from the time of peak discharge up to 24 hours before the event. Similarly, an AMI was developed considering the time of peak discharge. Additive exponential regression mod-

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els, AICc, and relative variable importance values were used to select the most plausi-

- ¹⁸⁰ ble explanatory variables across different rainfall intensities. In addition we estimated
- the time lag between daily peak discharge (m^3s^{-1}) and daily maximum rainfall (mm min⁻¹)
- across all the catchments. Finally, the relationship between peak discharge and catch-
- ¹⁸³ ment characteristics using box and whisker plots and regression models.

184 **3 Results**

The Nilgiris received the bulk of its rainfall between June and September (Figure 2) with an annual average of 2150mm over the three years, a maximum of 2,740 mm during 2014 and a minimum of 1,380 mm during 2016. We analysed a total of 5,542 days of rain events (≥ 0.2 mm rainfall) across all the catchments. There were 3,607, 272 and 277 events in light, heavy, and extreme rainfall intensity categories, respectively.



Figure 2. a) Distribution of daily rainfall across the months between 2014 and 2016. b) Contribution of each month's rainfall to the total annual rainfall across all the years. Most of the rainfall was received between June and September.

As expected, an increase in discharge with an increase in daily rainfall was observed. However, this relationship showed high variations across catchments for different rainfall intensities (Figure 3 & Figure 4). Scatter-plots and fitted exponential curves between daily rainfall and unit mean and peak daily discharges suggested greater stream discharges in grasslands and sholas when compared to wattle plantations during light rain (Figure 3 (b) & Figure 4 (b), whereas trends for wattle were above grassland and shola with increase in rainfall intensities (Figure 3 (c & d), 4 (c & d)).



Figure 3. Mean daily discharge values plotted against log of daily rainfall suggest an exponential relationship between rainfall and runoff. This relationship varies with catchment land cover type and rainfall intensity. As the rainfall intensitly increases discharge from wattle changes from lowest in light rain to the highest during ERE. a) All rainfall events, discharge of wattle > grassland > shola; b) light rain (< 38 mm/day) discharge of shola \geq grasslands > wattle; c) heavy and extreme rain ($\geq 38 \text{ mm/day}$), discharge shola > wattle > grassland; d) extreme rain (> 71 mm/day) discharge of wattle > shola > grasslands. Exp(R): Exponential rainfall; p=probability; AICc: Akaike Information Criteria with bias adjustment.



Figure 4. Unit peak discharge values plotted against log of daily rainfall suggest an exponential relationship between rainfall and runoff. Daily rainfall was calculated for a 24 hour period from the time of the event to 24 hours before the event. Discharge from wattle dominated catchments was the lowest during light rains but much higher than shola or grassland dominated catchments during heavy and ERE, the threshold being about 80mm/day. a) All rainfall, discharge of wattle > shola > grassland; b) light rain (< 38 mm/day), grassland \geq shola > wattle; c) heavy rain (\geq 38 mm/day and < 71 mm/day) wattle > shola > grassland; and d) extreme rain (> 71 mm/day) wattle > shola > grasslands. Exp(R): Exponential rainfall; p=probability; AICc: Akaike Information Criteria with bias adjustment.

Exponential regression models of unit mean daily discharge suggested land-cover to be one of the important variables that influenced the rainfall-runoff relationship (Table 2). When all the rainfall events were considered for analysis, grasslands had greater discharges for a given rainfall intensity when compared to wooded shola and wattle plantations (Table 2).

Table 2. Land cover played a major role in influencing the rainfall-runoff response.

Model	Intercept	Slope	р,	r^2	AICc	Δ AICc
$\overline{\exp(\mathbf{R})^a + \mathbf{L}\mathbf{C}^b + \mathbf{C}\mathbf{I}^c + \mathbf{A}\mathbf{M}\mathbf{I}^d}$	-0.35010		< 0	.0001, 0.6962	-20433.2	0.00
Rainfall		0.3630		,		
LC: Shola		-0.0103				
LC: Wattle		-0.0042				
LC: Grassland		-				
Circulatory Index		-0.0134				
Antecedent Moisture Index		0.2089				
exp(R) + LC + CI + DD + AMI	-0.3492		< 0	.0001, 0.6962	-20432.1	1.05
Rainfall		0.3631				
LC: Shola		-0.0135				
LC: Wattle		-0.0135				
LC: Grassland		-				
Circulatory Index		-0.0135				
Drainage Density		-0.0026				
Antecedent Moisture Index		0.2091				
NULL (Rainfall only model)	-0.4458	0.4612	< 0	.0001, 0.5428	-18172.3	2260.87
	Estimate		Adj	. SE	р	RVI^{e}
Intercept	-0.3504		0.00)56	< 0.0001	-
Rainfall	0.3630		0.00	50	< 0.0001	1.00
LC: Shola	-0.0101		0.00	022	< 0.0001	1.00
LC: Wattle	-0.0042		0.00	018	< 0.05	
LC: Grassland	-		-		-	
Circulatory Index	-0.0135		0.00)17	< 0.0001	1.00
Drainage Density	-0.0026		0.00	040	> 0.05	0.37
Antecedent Moisture Index	0.2090		0.00	027	< 0.0001	1.00

Note: Exponential regression models suggest that land cover played a major role in influencing the rainfall-runoff relationshiop. Catchment shape and drainage density are other important factors. We used all rainfall events with unit mean daily discharge as the response variable. Top models with delta AICc < 2 and the null model parameters are presented. Model averaging was done for models with delta AICc < 2.0. $^{a}\exp(R)$: exponential total daily rainfall; ^bLC: Dominant Land Cover; ^cCI: Circulatory Index; ^dAMI: Antecedent Moisture Index; ^eRVI: Relative Variable Importance.

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202	Analysis at different rainfall intensities also revealed that the land cover played an
203	important role in determining rainfall-runoff relationships. However, the hydrologic re-
204	sponses at different rainfall intensities changed across land cover types with rainfall in-
205	tensities. At light and heavy rainfall intensities, grassland showed greater unit area dis-
206	charge compared to shola and wattle (Table 3). However, during extreme rainfall events

- ²⁰⁷ wattle showed significantly greater unit mean discharge values compared to natural grass-
- lands and shola forests (Table 3). In addition to land cover, we found AMI and catch-
- ²⁰⁹ ment shape to have greater effect on this influence. Drainage density also had some in-
- fluence on these trends (Table 2 & 3).

Table 3. Model averaged parameters for relationship between unit mean daily discharge and daily rainfall

Model averaged parameters	Estimate	Adjusted standard error	р	RVI^a
a) Light Rain				
Intercept	-0.0111	0.0033	< 0.001	-
Rainfall	0.0128	0.0032	< 0.001	1
LC: Shola	-0.0012	0.0002	< 0.001	1
LC: Wattle	-0.0002	0.0002	> 0.05	
LC: Grassland	-	-	-	
Antecedent Moisture Index	0.0086	0.0029	< 0.01	1
Drainage Density	0.0003	0.0001	> 0.05	0.62
Circulatory Index	-0.0009	0.0003	< 0.01	1
b) Heavy Rain				
Intercept	-0.3644	0.135	< 0.01	-
Rainfall	0.3988	0.1138	< 0.001	1
LC: Shola	-0.021	0.0146	> 0.05	0.64
LC: Wattle	-0.0273	0.0122	< 0.05	
LC: Grassland	-	-	-	
Antecedent Moisture Index	0.3111	0.0162	< 0.001	1
Drainage Density	-0.0605	0.0226	< 0.01	1
Circulatory Index	-0.0141	0.0109	> 0.05	0.45
c) Extreme Rain				
Intercept	-0.1978	0.0613	< 0.01	-
Rainfall	0.289	0.0337	< 0.001	1
LC: Shola	-0.1015	0.0342	< 0.01	1
LC: Wattle	0.0234	0.029	> 0.05	
LC: Grassland	-	-	-	
Antecedent Moisture Index	0.359	0.0358	< 0.001	1
Drainage Density	-0.0497	0.0588	> 0.05	0.33
Circulatory Index	-0.114	0.024	< 0.001	1

Note: a) light rain ($\geq 1.3 \text{ mm/day}$ - < 38.0 mm/day); b) heavy rain (38.0 mm/day-71.0 mm/day); and c) ERE (> 71 mm/day). Model averaging was done for models with Δ AICc < 2.0. Land cover influenced the rainfall-runoff relationship significantly across different rainfall intensities.^{*a*} RVI: Relative Variable Importance.

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We found that daily rainfall was the best predictor variable for modelling daily peak

flows (p < 0.001, $r^2 = 0.60$). Analysis of unit peak daily discharge did not show any in-

fluence of land cover on rainfall-runoff relationship when all rainfall values were consid-

- ered in a regression model. Peak discharge was mainly determined by antecedent mois-
- ture content of the catchment (AMI Relative variable importance 1.00) along with the
- daily rainfall (Table 4). In addition to this, drainage density (Relative variable impor-
- tance = 0.38) and catchment shapes (Relative variable importance = 0.55) also had some

- influence on these observed trends in peak flows. However, when different intensities of
- rainfall were considered, we found landcover to have an influence on peak-flow. During
- low intensity rainfall, grassland showed greater peak flows (Table 4). In some contrast,
- during extreme rainfall events, peak flows were greater in wattle compared to natural
- 222 grasslands and shola.

Model averaged parameters	Estimate	Adjusted standard error	р	RVI^a
a) All Rain Events Intercept Rainfall Antecedent Moisture Index Circulatory Index Drainage Density	$\begin{array}{c} -0.3064\\ 0.3018\\ 0.0576\\ 0.0024\\ 0.0029\end{array}$	$\begin{array}{c} 0.0052 \\ 0.0045 \\ 0.0043 \\ 0.0016 \\ 0.0028 \end{array}$	$\begin{array}{c} < 0.001 \\ < 0.001 \\ < 0.001 \\ > 0.05 \\ > 0.05 \end{array}$	$1.00 \\ 1.00 \\ 0.55 \\ 0.38$
b) Light Rain $(\geq 1.3 \text{ to } <38.0 \text{ mm/day})$ Intercept Rainfall LC: Shola LC: Wattle	-0.1565 0.1572 -0.0019 -0.0047	$\begin{array}{c} 0.0056 \\ 0.0054 \\ 0.0006 \\ 0.0005 \end{array}$	< 0.001 < 0.001 < 0.001 > 0.001	1.00 1.00
Antecedent Moisture Index Circulatory Index Drainage Density	$0.0692 \\ 0.0020 \\ 0.0029$	$0.001\bar{4} \\ 0.0004 \\ 0.0007$	$< 0.001 \\ < 0.001 \\ < 0.001 \\ < 0.001$	$1.00 \\ 1.00 \\ 1.00$
c) Heavy Rain (38.0 to 71.0 mm/day) Intercept Rainfall LC: Shola LC: Wattle	$-0.6399 \\ 0.5949 \\ 0.0098 \\ -0.0057$	$\begin{array}{c} 0.1158 \\ 0.1005 \\ 0.0124 \\ 0.0108 \end{array}$	< 0.001 < 0.001 > 0.05 > 0.05	$1.00\\0.20$
Antecedent Moisture Index Circulatory Index Drainage Density	$0.0376 \\ 1.0135 \\ 0.0119$	$0.0151 \\ 0.0103 \\ 0.0166$	$< 0.05 \\> 0.05 \\> 0.05 \\> 0.05$	$1.00 \\ 0.44 \\ 0.18$
d) Extreme Rain (> 71 mm/day) Intercept Rainfall LC: Shola LC: Wattle	-0.4036 0.3482 -0.0152 0.0328	$\begin{array}{c} 0.0428 \\ 0.0258 \\ 0.0235 \\ 0.0207 \end{array}$	< 0.001 < 0.001 > 0.05 > 0.05	1.00 1.00
Antecedent Moisture Index Circulatory Index Drainage Density	$0.0853 \\ -0.0133 \\ 0.0262$	$0.0267 \\ 0.0178 \\ 0.0451$	$< 0.01 \\> 0.05 \\> 0.05$	$1.00 \\ 0.25 \\ 0.22$

Table 4. Relationship between unit peak daily discharge and daily rainfall.

Note: Peak discharge analysed against all rainfall intensities was largely determined by antecedent moisture and daily rainfall. However, landcover (LC), influenced runoff when different rainfall intensities were considered. Model averaging was done for models with delta AICc $< 2.0.^{a}$ RVI: Relative Variable Importance.

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The median time lag between rainfall and peak stream flows was 60, 72 and 89 min-

²²⁴ utes for wattle, shola, and grasslands respectively, suggesting a more rapid stream flow

generation pathway in wattle plantations compared to natural land covers. (Figure 5).

We did not find the influence of any other co-variate on lag-time across all the catchments.



Figure 5. Time lag between peak rainfall and peak stream flow was shorter in wattle plantations compared to natural grasslands and shola forests.

²²⁷ 4 Discussion

The relationship between different rainfall intensities and mean daily discharge in 228 the Upper Nilgiris were influenced by land cover. This relationship, however, varied with 229 rain intensity. During light rain (< 38 mm/day) grasslands showed higher discharges than 230 either shola forest or wattle plantations. This may be explained by the wider spread and 231 deeper root systems of trees when compared to grasses and higher soil organic matter 232 in the form of leaf litter could help in greater rainfall infiltration and retention during 233 light rains. Vegetation cover is known to increase surface roughness and infiltration which 234 slows surface run-off (Koutnỳ et al., 2014). Canopy cover also intercepts rain and thereby 235 reduces immediate runoff (Levia et al., 2011; Cui et al., 2012; Livesley et al., 2014). 236

In contrast, wattle plantations were least able to retain rainfall during ERE, re-237 sulting in higher and more rapid stream runoff. Both mean daily discharge and peak flows 238 were higher in wattle compared to natural grasslands and shola forests during heavy and 239 extreme rainfall. In other words, ERE are more likely to generate floods from catchments 240 converted from natural grasslands to wattle plantations or those invaded by wattle. A 241 greater incidence of saturation excess overland flows (Haga et al., 2005) in wattle might 242 explain this rapid response. Soil under wattle plantations was shown to have a lower sat-243 urated hydraulic conductivity than either natural grasslands or shola forests during an 244 earlier study in these sites (Krishnaswamy et al., 2017), also see Appendix C. This also 245 explains why wattle had a shorter time lag between rainfall and peak daily discharge; 246 60, 72 and 89 minutes for wattle, shola and grasslands respectively. 247

During heavy and extreme rainfall, wattle plantations resulted in greater discharges, 248 even though antecedent moisture played a greater role in determining runoff response 249 than did land cover in these catchments. This is in contrast with other studies which found 250 that light to moderate rainfall influenced peak stream-flow, while the role of landcover 251 was minimal during extreme rains (Bathurst et al., 2011). Sub-surface flows in wattle 252 plantations would explain the low flows during light rains which increase as rain inten-253 sity increases. Prior studies in the region suggest that wattle plantations with their ex-254 tended root systems are dominated by subsurface stream flows (Krishnaswamy et al., 255 2017). The decreasing time-lag with increase in woody cover also suggests that higher 256 rain intensities trigger rapid sub-surface flows, as predicted for regions such as the Nil-257 giris which have a sub-surface dominated runoff system (Chappell et al., 2017). 258

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Antecedent moisture clearly plays an important role in controlling run-off responses 259 in the Nilgiris. This indicates that the responses are not only dependent on infiltration 260 excess (Liu et al., 2008) but are a result of complex relationships between catchment in-261 filtration rate, soil saturation, water-holding capacity, and subsurface flows, all of which 262 are influenced by land-cover. The drainage network of catchments also influences stream 263 flows; high drainage density increases peak discharges, while reduces overall runoff through 264 retaining much of the rainfall across the catchments. Similarly, stream runoff from the 265 catchment is reduced as the shape becomes more circular. 266

Natural vegetation has been found to reduce flood risks in different parts of the world (Bathurst et al., 2011; Koutnỳ et al., 2014; Krishnaswamy et al., 2012). Our results suggest that wattle plantations, when compared with native vegetation, increase flood-risk during high and extreme rainfall. The invasion of natural grasslands by wattle therefore has serious ramifications for downstream flooding during the peak rainy season.

Several studies suggest that wattle and other plantations significantly increase evap-272 otranspiration and reduce stream-flow (Dye & Jarmain, 2004; Sikka et al., 2003, 1998; 273 Chand et al., 2009; Sharma, 1984). Grasslands continue to be seen as 'degraded' areas 274 and have been targeted for afforestation historically, often with non-native plantation 275 crops (Chandran, 1997; Jha et al., 2000; Joshi et al., 2018; Arasumani et al., 2018). This 276 perception of grasslands being degraded and their conversion to plantations persists glob-277 ally (Veldman et al., 2015) and in some current policies in India (Ministry of Environ-278 ment and Forests, 2010), which seek to combat climate change through large scale af-279 forestation, disregarding the potential impact on water resources (Jackson et al., 2005). 280 Another example of this is the Atlas of Forest Landscape Restoration Opportunities (Atlas 281 of Forest and Landscape Restoration Opportunities, 2014), which identifies several of the 282 natural montane grasslands in the Western Ghats as Areas for wide-scale restoration. 283

In the Nilgiris, however, removal of wattle and restoration of the montane grasslands has been mandated by the government following a public interest litigation filed in 2014 (G.Vinod Kumar, 2014). Successful removal of wattle, however, requires sustained efforts. Once established, wattle builds substantial seed banks and rapidly regenerates after removal. This would amplify the hydrologic impact of wattle stands, which unlike the mature shola forests, are likely to have a far higher evapotranspiration rate (Cui et al., 2012).

Our results suggest that planted and invasive wattle stands have a significant hy-291 drologic footprint and might be detrimental to hydrologic services from catchments in 292 the Nilgiris, in both the dry and wet seasons. We found that wattle plants, which have 293 invaded several of the montane grassland systems in the Nilgiris (Joshi et al., 2018), in-294 crease flood risks during extreme rainfall events. This study finds that high intensity rain 295 like > 70 mm of rainfall in a day behaves differently when it is over a wattle plantation 296 compared to shola forests or grasslands. An extreme rain is more likely to generate a flood 297 in a wattle dominated catchment than one dominated by the natural shola forest and 298 grassland mosaic. We found that along with landcover, catchment morphometric char-299 acteristics such as catchment shape and drainage density also influence the rainfall-runoff 300 relationship. Therefore, future flood early warning systems and risk management strate-301 gies should consider the differential effects of land-cover including wooded vegetation and 302 catchment morphometric properties along with antecedent moisture conditions to pre-303 dict likelihood of floods at different rainfall intensities. 304

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Appendix A Land Use Land Cover Map of Nilgiris

We developed a land-cover map for the Nilgiris using LANDSAT 8 images for the year 2017. We collected a total of 3411 ground control points representing nine differ-

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ent land cover types using hand held GPS units (Garmin 20x) with a horizontal accu-323 racy of about 5 meters. These nine classes were Shola (montane evergreen forests), wa-324 ter bodies, built-up land, commercial plantation (tea), agriculture land, natural grass-325 land and forest plantation, which is a combination of exotic tree species such as wattle 326 (Acacia meansii and Acacia dealbata), blue gum (Eucalyptus globulus) and pine (Pi-327 nus patula), and scrub patches dominated by scotch broom and gorse. These point lo-328 cations were then converted into polygons by digitising around the points using Google 329 Earth imagery for the year 2017. 330

Google earth engine (http://earthengine.google.com) was used to develop land 331 use/land cover maps. The 2017 LANDSAT 8 images available at a resolution of 30 m 332 were used for the classification. A Random forest classifier algorithm (http://earthengine 333 .google.com) with a random seed of 40, was used for the classification. We used top of 334 atmosphere corrected (ToA) b1, b2, b3, b4, b5, b6 and b7 bands from Landsat, along 335 with a Normalised Difference vegetation Index (ndvi) layer and an elevation and slope 336 layer derived from Advances Land Observation Satellite (ALOS) DEM (https://www 337 .eorc.jaxa.jp/ALOS/en/aw3d30/index.htm), for the classification. The estimated kappa 338 accuracy for the developed maps was 81%. 330



Figure A1. Land-cover map of Nilgiris developed using supervised classification of Landsat images for the year 2017.

³⁴⁰ Appendix B Peak discharge across landcover types

341	We found Antecedent moisture content as an important factor that determined peak
342	discharges along with rainfall. This influence was evident across all the landcover types
343	(Supplementary Table 1). Boxplots suggested that the distribution of AMI was compa-
344	rable across all the land cover types (Figure B1). The median AMI values were 0.094, 0.097,
345	and 0.11 for grassland, wattle and shola respectively. Exponential regression models with
346	antecedent moisture index (AMI) as the only predictor variable suggested a significant
347	influence on peak discharge (p < 0.001) across all the land-cover types (Table B1). How-
348	ever, the variation in unit peak discharges explained by AMI differed across landcover
349	types; nearly 33% of variation was explained by AMI alone in grassland catchment, 22%
350	in shola catchments and only 13% in wattle (Table B2).

Landcover type	Model	AICc	Δ AICc
Grassland	1) $\exp(\mathbf{R}) + \mathbf{AMI}$	-2623.5	0
	$2) \exp(\mathbf{R})$	-2484	139.53
Wattle	1) $\exp(\mathbf{R}) + \mathbf{AMI}$	-11632.7	0
	2) $\exp(R) + AMI + DD$	-11632.3	0.42
	3) $\exp(R)$ +AMI+CI	-11631.2	1.46
	4) $\exp(R) + AMI + DD + CI$	-11630.8	1.89
	5) $\exp(\mathbf{R})$	-11552.3	80.43
Shola	1) $\exp(R) + AMI + CI$	-3708.5	0
	2) $\exp(R) + AMI + DD$	-3708.5	0
	$3) \exp(\mathbf{R})$	-3587.6	120.87

Table B1. Influence of antecedent moisture index (AMI) on runoff-rainfall relationships.

Note: Top models with delta AIC c <2 and the null model (rainfall only model) across the land cover types.

Table B2. Slope estimate and p value for linear regression analysis of unit peak discharge(response variable) and antecedent moisture index (AMI).

Landcover type	Slope	r^2 , p
Grassland	0.154373	0.3286, < 0.001
Wattle	0.183716	0.1272, < 0.001
Shola	0.156564	0.2198, < 0.001

³⁵¹ Appendix C Infiltration rates across land cover types

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We compared the observed rainfall intensity to the measured saturated hydrologic

³⁵³ conductivity across soils under wattle plantations, grasslands and shola forests. Our find-



Figure B1. Distribution of antecedent moisture across different landcover types.

- ings suggest that soils under wattle plantations have a lower infiltration rate and are more
- vulnerable to infiltration-excess overland flow compared to soil under shola forests and
- ³⁵⁶ grasslands (Figure C1).



Figure C1. Box and Whiskers plots of Infiltration under different land-cover in the Nilgiris overlayed with maximum rain intensities recorded in 2013, 2014 and 2015.

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