Quantifying the effects of EMIC wave scattering and magnetopause shadowing in the outer electron radiation belt by means of data assimilation

Juan Sebastian Cervantes Villa¹, Yuri Y Shprits², Nikita A Aseev³, and Hayley J Allison⁴

¹Helmholtz Centre Potsdam, GFZ German Research Centre for Geosciences ²Helmholtz Centre Potsdam ³GFZ German Research Centre for Geosciences ⁴GFZ

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Abstract

In this study we investigate two distinct loss mechanisms responsible for the rapid dropouts of radiation belt electrons by assimilating data from Van Allen Probes A and B and Geostationary Operational Environmental Satellites (GOES) 13 and 15 into a 3-D diffusion model. In particular, we examine the respective contribution of electromagnetic ion cyclotron (EMIC) wave scattering and magnetopause shadowing for values of the first adiabatic invariant μ ranging from 300 to 3000 MeV G. We inspect the innovation vector and perform a statistical analysis to quantitatively assess the effect of both processes as a function of various geomagnetic indices, solar wind parameters, and radial distance from the Earth. Our results are in agreement with previous studies that demonstrated the energy dependence of these two mechanisms. Loss from L^{*} = 4 to L^{*} = 4.8 is dominated by EMIC wave scattering (μ [?] 900 MeV G) and may amount to between 10%/hr to 30%/hr of the maximum value of phase space density (PSD) over all L shells for fixed first and second adiabatic invariants. Magnetopause shadowing is shown to deplete electrons across all energies, mostly between L^{*} = 5 and L^{*} = 6.6, resulting in loss from 50%/hr to 70%/hr of the maximum PSD. We also identify a boundary located between L^{*} = 3.5 and L^{*} = 5.2 clearly separating the regions where each mechanism dominates. Nevertheless, during times of enhanced geomagnetic activity, both processes can operate beyond such location and encompass the entire outer radiation belt.

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S. Cervantes^{1,2}, Y. Y. Shprits^{1,2,3}, N. A. Aseev^{1,2}, and H. J. Allison¹

5	$^1\mathrm{Helmholtz}$ Centre Potsdam, GFZ German Research Centre for Geosciences, Potsdam, Germany
6	² Institute of Physics and Astronomy, University of Potsdam, Potsdam, Germany
7	³ Department of Earth, Planetary, and Space Sciences, University of California, Los Angeles, CA, USA

8	Key Points:
9	• We present a four-year reconstruction of the outer radiation belt based on data
10	assimilation
11	• In the outer region of the inner magnetosphere, loss due to outward radial diffu-
12	sion dominates
13	• At multi-MeV energies, loss due to EMIC waves can dominate in the heart of the
14	outer radiation belt

Corresponding author: S. Cervantes, jscv@gfz-potsdam.de

15 Abstract

In this study we investigate two distinct loss mechanisms responsible for the rapid 16 dropouts of radiation belt electrons by assimilating data from Van Allen Probes A and 17 B and Geostationary Operational Environmental Satellites (GOES) 13 and 15 into a 3-18 D diffusion model. In particular, we examine the respective contribution of electromag-19 netic ion cyclotron (EMIC) wave scattering and magnetopause shadowing for values of 20 the first adiabatic invariant μ ranging from 300 to 3000 MeV G⁻¹. We inspect the in-21 novation vector and perform a statistical analysis to quantitatively assess the effect of 22 both processes as a function of various geomagnetic indices, solar wind parameters, and 23 radial distance from the Earth. Our results are in agreement with previous studies that 24 demonstrated the energy dependence of these two mechanisms. Loss from $L^* = 4$ to 25 $L^* = 4.8$ is dominated by EMIC wave scattering ($\mu \ge 900 \text{ MeV G}^{-1}$) and may amount 26 to between 10%/hr to 30%/hr of the maximum value of phase space density (PSD) over 27 all L shells for fixed first and second adiabatic invariants. Magnetopause shadowing is 28 shown to deplete electrons across all energies, mostly between $L^* = 5$ and $L^* = 6.6$, 29 resulting in loss from 50%/hr to 70%/hr of the maximum PSD. We also identify a bound-30 ary located between $L^* = 3.5$ and $L^* = 5.2$ clearly separating the regions where each 31 mechanism dominates. Nevertheless, during times of enhanced geomagnetic activity, both 32 processes can operate beyond such location and encompass the entire outer radiation belt. 33

34 1 Introduction

The physics governing the energetic electrons in the Earth's radiation belts has been 35 subject of considerable research since their discovery in 1959. The outer belt extends from 36 approximately 3 to 7 R_E , is highly dynamic, and can vary by several orders of magni-37 tude on timescales ranging from minutes to weeks. Based on an examination of 276 mod-38 erate and intense geomagnetic storms from the period 1989-2000, Reeves et al. (2003) 39 found that storms could either increase, significantly decrease, or not substantially change 40 the fluxes of relativistic electrons in the outer belt. Further studies have associated the 41 variability in the responses of the radiation belts to storms to the complex competing 42 nature between acceleration and loss (e.g., Friedel et al., 2002; Shprits, Elkington, et al., 43 2008; Shprits, Subbotin, et al., 2008; Millan & Baker, 2012; Turner, Angelopoulos, Li, 44 et al., 2014). Understanding the mechanisms responsible for the acceleration and loss 45 of electrons is indispensable for predicting the response of the radiation belts to geomag-46

⁴⁷ netic disturbances. In this study we essentially focus on the rapid loss of radiation belt
⁴⁸ electrons.

It is now widely accepted that reductions of the outer radiation belt electron flux 49 can be attributed both to adiabatic and nonadiabatic processes. Adiabatic processes (H. Kim 50 & Chan, 1997) allow electron fluxes to return to its pre-storm level in the storm recov-51 ery phase and radially transport particles in response to a change in the magnetosphere 52 to conserve the three adiabatic invariants (μ, K, Φ) . In contrast, many events associated 53 with main-phase dropouts do not recover and fluxes do not return to the original pre-54 storm values (e.g., McAdams & Reeves, 2001; Reeves et al., 2003). In such cases, the dropout 55 is a result of several different nonadiabatic processes that remove the electrons perma-56 nently. 57

One mechanism that falls into this nonadiabatic category is the loss due to pitch-58 angle scattering via resonant interaction with various types of magnetospheric waves, in-59 cluding whistler mode chorus, plasmaspheric hiss, and electromagnetic ion cyclotron (EMIC) 60 waves, which leads to electron precipitation to the atmosphere (e.g., Thorne & Kennel, 61 1971; Lyons et al., 1972; Thorne et al., 2005; Millan et al., 2007; Thorne, 2010; Turner, 62 Angelopoulos, Li, et al., 2014). Another nonadiabatic process is the loss across the mag-63 netopause, called magnetopause shadowing. This term describes the scenario in which 64 the magnetopause moves inward due to increases in solar wind dynamic pressure, result-65 ing in the depletion of electrons on open drift paths that were previously closed (e.g., 66 K. Kim et al., 2008; Ohtani et al., 2009; Morley et al., 2010; Turner et al., 2012). In ad-67 dition, the loss to the magnetopause generates a sharp gradient that further drives elec-68 tron outwards and through the magnetosphere, a process known as outward radial dif-69 fusion (Shprits et al., 2006). Nevertheless, the relative contribution of each physical pro-70 cess to electron flux dropouts still remains a fundamental puzzle. 71

Multisatellite observations provide a useful means of understanding the dominant
loss mechanisms of radiation belt dropouts. For instance, Green et al. (2004) used 52
dropout events and tested several processes that may contribute to electron flux decreases,
including adiabatic motion, magnetopause shadowing, and precipitation to the atmosphere.
Their study concluded that the most likely cause of the dropout was precipitation to the
atmosphere, although the cause of the precipitation remained uncertain. Turner et al.
(2012) analyzed data collected by several spacecraft and concluded that the sudden elec-

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tron depletion on 6 January 2011 was mainly a result of outward radial diffusion rather 79 than loss to the atmosphere. Bortnik et al. (2006) studied the relativistic electron dropout 80 on 20 November 2003 and suggested that it was caused by two separate mechanisms that 81 operate at high and low L shells. At L > 5 loss was dominated by magnetopause shad-82 owing and outward radial diffusion, whereas at L < 5 it was dominated by pitch an-83 gle scattering driven by EMIC waves. Similarly, Turner, Angelopoulos, Li, et al. (2014) 84 and Turner, Angelopoulos, Morley, et al. (2014) studied the 30 September 2012 dropout 85 event and concluded that both loss mechanisms operated, with a boundary at $L^* \sim 4$. 86 More recently, Xiang et al. (2017) investigated three distinct radiation belt dropouts ob-87 served by Van Allen Probes, subtracting the electron phase space density (PSD) versus 88 L^* profiles before and after the dropout. Their findings suggest that these events can 89 be classified in three different classes in terms of dominant loss processes: magnetopause 90 shadowing dominant, EMIC wave scattering dominant, and a combination of both mech-91 anisms. However, one limitation of in-situ data is the sparse coverage, as incomplete pro-92 files may hinder the calculation of PSD drops. 93

On the other hand, radiation belt modeling studies have also focused on the im-94 portance of loss processes in flux dropouts. For example, Shprits et al. (2006) explored 95 the viability of outward radial diffusion loss by comparing radial diffusion model sim-96 ulations with CRRES measurements. The comparison showed that nonadiabatic flux dropouts 97 near geosynchronous orbit can be effectively propagated by the outward radial diffusion 98 down to $L^* = 4$ and that magnetopause loss coupled with the radial transport can ac-99 count for the main-phase flux dropout. Su et al. (2011) examined the contribution of dif-100 ferent loss processes by comparing CRRES observations with a three dimensional (3-D) 101 radiation belt model by gradually incorporating magnetopause shadowing, adiabatic trans-102 port, radial diffusion, and plume and chorus wave-particle interactions into the code. Yu 103 et al. (2013) quantified the relative contribution of magnetopause shadowing coupled with 104 outward radial diffusion by comparing radial diffusion simulations with GPS-observed 105 total flux dropout. Their results indicated that such process accounted for 60-90%/hr 106 of the main-phase radiation belt electron dropout near geosynchronous orbit. 107

In the current study, we quantify the contribution of (1) pitch-angle scattering driven by EMIC waves and (2) magnetopause shadowing. We aim to answer the question: how much loss is caused by each mechanism? We tackle this issue with a novel approach based on the assimilation of spacecraft data in a 3-D diffusion model by means of a split-operator

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Kalman filter (KF) (Shprits, Kellerman, et al., 2013). In this way, data assimilation (DA) 112 combines spacecraft data and our model predictions in a two-way communication, such 113 that our model corrects inaccurate measurements and fills the gaps where electron PSD 114 measurements are lacking (a constraint in observational studies), and observations bring 115 our model closer to reality. We perform multiple four-year long-term runs (for the pe-116 riod 1 October 2012 to 1 October 2016 spanning different levels of geomagnetic activ-117 ity) by switching on and off in the model the above-mentioned mechanisms. We quan-118 tify their effect by means of the innovation vector, a measure on how observations and 119 model predictions differ, for various values of the adiabatic invariants μ and K. 120

The outline of this paper is as follows. A brief description of data assimilation and the methodology followed in this study are presented in section 2. We show the longterm reanalysis results of electron PSD in section 3 and the statistical analysis of the effect of scattering by EMIC waves and magnetopause shadowing employing the innovation vector in section 4. Results are discussed in Section 5 and conclusions are presented in section 6.

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2 Methodology and data

2.1 VERB code

The current study builds upon the previous work of Shprits, Kellerman, et al. (2013), 129 Kellerman et al. (2014) and Cervantes et al. (2020) and adopts the 3-D Versatile Elec-130 tron Radiation Belt Code (VERB-3D; Shprits et al., 2009; Subbotin & Shprits, 2009) 131 to assimilate spacecraft data at different locations. The VERB-3D code models the evo-132 lution of electron PSD by solving the modified 3-D Fokker-Planck diffusion equation that 133 incorporates radial diffusion, energy diffusion, pitch angle scattering, and mixed diffu-134 sion into the drift- and bounce-averaged particle PSD (Schulz & Lanzerotti, 1974). The 135 3-D Fokker-Planck equation for the evolution of PSD can be written in terms of the L136 shell, equatorial pitch angle α_0 , and relativistic momentum p, following Shprits et al. (2009) 137 and Subbotin and Shprits (2009): 138

$$\frac{\partial f}{\partial t} = L^{*^2} \frac{\partial}{\partial L^*} \Big|_{\mu,J} \left(\frac{1}{L^{*^2}} D_{L^*L^*} \frac{\partial f}{\partial L^*} \Big|_{\mu,J} \right) + \frac{1}{p^2} \frac{\partial f}{\partial p} \Big|_{\alpha_0,L^*} p^2 \left(D_{pp} \frac{\partial f}{\partial p} \Big|_{\alpha_0,L^*} + D_{\alpha_0 p} \frac{\partial f}{\partial \alpha_0} \Big|_{p,L^*} \right) \\ + \frac{1}{T(\alpha_0) \sin(2\alpha_0)} \frac{\partial}{\partial \alpha_0} \Big|_{p,L^*} T(\alpha_0) \sin(2\alpha_0) \left(D_{\alpha_0 \alpha_0} \frac{\partial f}{\partial \alpha_0} \Big|_{p,L^*} + D_{\alpha_0 p} \frac{\partial f}{\partial p} \Big|_{\alpha_0,L^*} \right) - \frac{f}{\tau} \quad (1)$$

- where f is electron PSD, μ and J are the first and second adiabatic invariants, respec-
- tively, and L^* is inversely related to the third adiabatic invariant Φ . $D_{L^*L^*}$, D_{pp} , $D_{\alpha_0\alpha_0}$,
- and $D_{\alpha_0 p}$ are the bounce-averaged radial, momentum, pitch angle, and mixed pitch angle-
- momentum diffusion coefficients, respectively. $T(\alpha_0)$ is a function related to the parti-

cle's bounce time (Lenchek et al., 1961; Schulz & Lanzerotti, 1974):

$$T(\alpha_0) = 1.3802 - 0.3198(\sin\alpha_0 + \sin^{1/2}\alpha_0)$$
⁽²⁾

The parameter τ is a loss rate assumed to be infinite outside the loss cone and equal to a quarter of the electron bounce time inside the loss cone. Readers are referred to Shprits et al. (2009) and Subbotin and Shprits (2009) for a more detailed description of the VERB-3D model.

Based on the previous findings of Drozdov, Shprits, Aseev, et al. (2017) and Wang 148 et al. (2019), who meticulously studied the sensitivity of various parameterizations of ra-149 dial diffusion, we employ the magnetic radial diffusion rates $D_{L^*L^*}$ of Brautigam and 150 Albert (2000). The parameters for dayside and nightside chorus are taken from Orlova 151 and Shprits (2014), while for hiss the parameterization of Orlova et al. (2014) is used. 152 The location of the plasmapause is calculated following Carpenter and Anderson (1992). 153 The spectral properties from Meredith et al. (2014) are used to calculate diffusion co-154 efficients for helium band EMIC waves, and they are included in the simulation when 155 the solar wind dynamic pressure is greater than or equal to 3 nPa (Drozdov, Shprits, Us-156 anova, et al., 2017). The VERB-3D code includes the Last Closed Drift Shell (LCDS) 157 as a function of time and invariant K. As in Cervantes et al. (2020), physics associated 158 with magnetopause shadowing are introduced using the LCDS. In this study, we use an 159 energy-dependent loss mechanism, since the rate of loss following a reduction in the LCDS 160 depends on the particle's drift period. We employ the Tsyganenko and Sitnov (2007) mag-161 netic field model incorporated into the IRBEM library (Boscher et al., 2012) to deter-162 mine the LCDS, and we simulate loss due to magnetopause shadowing with an exponen-163 tial decay of the electron PSD outside the LCDS, as: 164

$$f(t, L^* > \operatorname{LCDS}(t)) = f(t)e^{(-1/\tau_d)}$$
(3)

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where, τ_d is the electron drift period calculated as Walt (2005):

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$$\tau_{\rm d}(s) = C_{\rm d} \left(\frac{\mathrm{R_E}}{R_0}\right) \frac{1}{\gamma \beta^2} \left[1 - 0.333 \left(\sin\alpha_0\right)^{0.62}\right] \tag{4}$$

Here, $\beta = \frac{v}{c}$, $\gamma = (1-\beta^2)^{-1/2}$, $C_d = 1.557 \times 10^4$ for electrons, $R_E = 6.37 \times 10^3$ km, and R_0 is the distance from the center of the Earth to the equatorial crossing point of a magnetic field line. As the electron energy increases, the drift period decreases.

- The size of the computational grid is $29 \times 101 \times 91$ points along radial, energy, and pitch angle dimension, respectively. Radial grid points are distributed uniformly, whereas energy and pitch angle grid points are distributed logarithmically. The L^* grid is set from $1 R_E$ to 6.6 R_E . The energy grid is defined by a minimum of 0.01 MeV and a maximum of 10 MeV at the outer radial boundary. The pitch angle grid extends from 0.3° to 89.7°.
- For the solution of equation (1), the initial PSD is taken from the steady state so-174 lution of the radial diffusion equation. A lower radial boundary condition $(L^* = 1)$ of 175 f = 0 is used in order to simulate the loss of electrons to the atmosphere. The PSD 176 required for the upper radial boundary condition $(L^* = 6.6)$ is obtained from Geosta-177 tionary Operational Environmental Satellites (GOES) observations. The upper energy 178 boundary at 10 MeV is set equal to zero. For the lower energy boundary, the PSD is set 179 constant in time to represent a balance of convective sources and loss. The lower pitch 180 angle boundary condition is set to zero to simulate precipitation loss of electrons into 181 the loss cone in a weak diffusion regime. A zero gradient is chosen to account for the flat 182 pitch angle distribution observed at 90° (Horne et al., 2003) for the upper pitch angle 183 boundary condition. 184
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2.2 Instrumentation and data

We use simultaneous measurements of four spacecraft, the twin Van Allen Probes (renamed from Radiation Belt Storm Probes after launch) A and B, and GOES 13 and 15, covering a four-year period from 1 October 2012 to 1 October 2016. For data assimilation, observations are converted from flux to PSD in phase space coordinates (L^*, μ, K) . In-situ magnetic field measurements are employed to calculate μ , while the Tsyganenko and Sitnov (2007) model is employed to calculate K and L^* .

On board the Van Allen Probes (Mauk et al., 2012; Stratton et al., 2012), the Radiation Belt Storm Probes-Energetic particle, Composition, and Thermal plasma (RBSP-ECT) suite measures particles with energies ranging from hot to ultrarelativistic (Spence et al., 2013). In this study, we utilize measurements from the Magnetic Electron Ion Spectrometer (MagEIS) (Blake et al., 2013), which provides data in the energy range ~ 30 keV to about 4 MeV and Relativistic Electron Proton Telescope (REPT) (Baker et al., 2012) instruments, which covers energies from 2 MeV to tens of MeV. The pitch angle distribution is interpolated in a uniform grid with a step of 5°.

In addition, from satellites 13 and 15 of the multi-mission GOES spacecraft (Onsager 200 et al., 1996; Singer et al., 1996) we employ data from the MAGnetospheric Electron De-201 tector (MAGED; Hanser, 2011) and Energetic Proton, Electron, and Alpha Detector (EPEAD; 202 Onsager et al., 1996; Hanser, 2011) instruments. Nine solid-state-detector telescopes from 203 MAGED provide pitch-angle resolved in-situ electron flux measurements in five energy 204 bands: 30 - 50, 50 - 100, 100 - 200, 200 - 350, and 350 - 600 keV. Four telescopes are ori-205 ented in the north-south plane, and the other five in the east-west plane (Hanser, 2011; 206 Rodriguez, 2014a). Moreover, two EPEAD detectors (Onsager et al., 1996; Hanser, 2011) 207 on board each spacecraft measure MeV electron and solar proton fluxes in two energy 208 ranges: > 0.8 MeV and > 2 MeV. One detector is oriented westward and the other east-209 ward (Rodriguez, 2014b). MAGED and EPEAD observations at a five-minute cadence 210 are averaged over one hour. EPEAD integral fluxes are obtained by averaging the mea-211 surements over the westward and eastward telescopes, so that the resulting pitch angles 212 are averages between both directions of the two telescopes as well. Integral fluxes as a 213 function of energy are fitted to a power law which is used to interpolate between values 214 up to 1 MeV. In order to convert to differential flux, we employ the 90° pitch angle dif-215 ferential flux data from MAGED and fit the two integral channels of EPEAD to an ex-216 ponential function $f = A \exp(B * E)$, where f is the differential flux, E is the energy, 217 and A and B are positive time-dependent coefficients obtained by solving the flux in-218 tegral for averaged MAGED data. The pitch angle distribution below 500 keV is directly 219 measured by MAGED. 220

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2.3 Data assimilation and innovation vector

Data assimilation is an algorithm which aims to smoothly blend sparse and inaccurate measurements with dynamical information from a physics-based model. Several DA methods have been developed, such as the Kalman filter in its standard (Kalman, 1960), extended (Jazwinski, 1970), and ensemble versions (Evensen, 1994). The KF is a powerful sequential data assimilation method that combines a numerical model and

incomplete measurements, while minimizing mean-squared errors (Kalman, 1960). The

²²⁸ methodology of the standard KF is briefly outlined below.

A system of evolution equations may be presented in the following form:

$$\mathbf{x}_{k}^{f} = \mathbf{M}_{k-1} \mathbf{x}_{k-1}^{a} \tag{5}$$

where **x** represents a model state vector (for our model, it is the PSD on the numerical grid locations), and the model matrix **M** advances the state vector **x** in discrete time increments. The subscript k shows the time step, and superscripts f and a refer to forecast and analysis, respectively. The evolution of \mathbf{x}_{k}^{t} (superscript t refers to true), is assumed to differ from the model by a random error ϵ^{m} :

$$\mathbf{x}_{k}^{t} = \mathbf{M}_{k-1}\mathbf{x}_{k-1}^{t} + \epsilon_{k}^{m} \tag{6}$$

where ϵ_k^m is assumed to be a Gaussian white-noise sequence, with mean zero and modelerror covariance matrix **Q**.

The observations \mathbf{y}_k^o (superscript *o* refers to observed), are assumed to be contaminated by observational errors ϵ_k^o :

$$\mathbf{y}_{k}^{o} = \mathbf{H}_{k}\mathbf{x}_{k}^{t} + \epsilon_{k}^{o} \tag{7}$$

where ϵ_k^o is also assumed to be Gaussian, white in time, with mean zero and given covariance matrix **R**. The observation matrix **H**_k accounts for the fact that usually the dimension of \mathbf{y}_k^o is less than the dimension of \mathbf{x}_k^t .

During the so-called update times, when observations are available, forecast and observations are blended to yield the analysis state vector:

$$\mathbf{x}_{k}^{a} = \mathbf{x}_{k}^{f} + \mathbf{K}_{k}(\mathbf{y}_{k}^{o} - \mathbf{H}_{k}\mathbf{x}_{k}^{f})$$

$$\tag{8}$$

where the term $\mathbf{K}_k \left(\mathbf{y}_k^o - \mathbf{H}_k \mathbf{x}_k^f \right)$ is usually referred to as the innovation vector \mathbf{x}_k^i . \mathbf{K}_k is the Kalman gain matrix computed at each time step using a time-evolving forecast error covariance matrix \mathbf{P}_k^f given by:

$$\mathbf{P}_{k}^{f} = \mathbf{M}_{k-1} \mathbf{P}_{k-1}^{a} \mathbf{M}_{k-1}^{T} + \mathbf{Q}_{k-1}$$

$$\tag{9}$$

The Kalman gain matrix \mathbf{K}_k represents the optimal weights given to the observations when updating the model state vector:

$$\mathbf{K}_{k} = \mathbf{P}_{k}^{f} \mathbf{H}_{k}^{T} \left(\mathbf{H}_{k} \mathbf{P}_{k}^{f} \mathbf{H}_{k}^{T} + \mathbf{R}_{k} \right)^{-1}$$
(10)

²⁴⁸ The error covariance matrix is also updated as follows:

$$\mathbf{P}_{k}^{a} = \left(\mathbf{I} - \mathbf{K}_{k}\mathbf{H}_{k}\right)\mathbf{P}_{k}^{f} \tag{11}$$

The innovation vector \mathbf{x}_k^i warrants a more detailed discussion as this is the term where, 249 in our case, source and loss processes are effectively incorporated into the KF. The in-250 novation vector measures how much new and additional information, provided by the 251 data (hence its name), will modify the model forecast \mathbf{x}^{f} in order to produce an opti-252 mal estimate of the state of the system \mathbf{x}^{a} . The value and the sign of the innovation vec-253 tor depend on how much the modeled and observed values differ from each other, and 254 on the estimated forecast and observational errors. A perfect model would predict ex-255 actly the incoming observations, and the innovation would be zero. As the forecast er-256 ror covariance matrix \mathbf{P}_k^f approaches zero, the innovation is weighted less heavily by the 257 gain \mathbf{K}_k . In contrast, as the observational error covariance matrix \mathbf{R}_k tends to zero, the 258 Kalman gain \mathbf{K}_k weights the innovation more heavily. Shprits et al. (2007), Koller et al. 259 (2007), Daae et al. (2011), and Cervantes et al. (2020) demonstrated the usefulness of 260 the innovation vector to identify and adjust for unknown, missing physics in radiation 261 belt models in order to reduce the discrepancy between observations and model predic-262 tions. All of the above-mentioned studies employed the innovation vector to infer accel-263 eration and loss processes for short-term intervals or specific events. 264

In this paper, we perform a four-year statistical analysis of the innovation vector and employ it as a tool to quantify the loss effect of EMIC wave scattering and magnetopause shadowing on radiation belt electrons. For that purpose, we perform three data assimilation runs (Table 1). The first run includes all processes in our model (hereinafter, "full" run), and in the second and third runs, one process is neglected in each. The "full"

Table 1. Summary of data assimilation runs

Run	Processes included
1	Radial diffusion due to ULF waves + pitch angle, energy, and mixed pitch angle- energy diffusion due to chorus and hiss waves + EMIC wave scattering + magne-
	topause shadowing, i.e. "full" run
2	Radial diffusion due to ULF waves + pitch angle, energy, and mixed pitch angle- energy diffusion due to chorus and hiss waves + magnetopause shadowing
3	Radial diffusion due to ULF waves + pitch angle, energy, and mixed pitch angle-
0	energy diffusion due to chorus and hiss waves + EMIC wave scattering

simulation (number 1) accounts for: radial diffusion due to ULF waves, pitch angle, energy, and mixed pitch angle-energy diffusion due to chorus and hiss waves, EMIC wave
scattering, and magnetopause shadowing. The second run (number 2) accounts for all
processes except for scattering by EMIC waves. Finally, the third run (number 3) includes
all processes in the "full run" with the exception of magnetopause shadowing. The time
step of our VERB simulations is one hour, and assimilation of spacecraft data is performed
at the same cadence.

For each of the three runs, we calculate the hourly innovation vector \mathbf{x}_{k}^{i} at each L^{*} and normalize it by the corresponding hourly maximum value of assimilated PSD \mathbf{x}_{k}^{a} (from the "full" run) over all L^{*} . Afterwards, the difference between the absolute values of the normalized innovation of the "full" simulation and the one excluding either loss process is calculated according to the following equation:

$$\Delta \mathbf{x}_{k}^{i} = \frac{|\mathbf{x}_{1,k}^{i}| - |\mathbf{x}_{2,k}^{i}|}{\max(\mathbf{x}_{1,k}^{a})} \times 100\%$$
(12)

where subscript 1 refers to the "full" run and subscript 2 to the run lacking either EMIC wave scattering or magnetopause shadowing. Negative values of $\Delta \mathbf{x}_k^i$ indicate that the inclusion of such mechanisms provides a better agreement with the observed PSD, bringing the model prediction closer to reality. On the other hand, positive $\Delta \mathbf{x}_k^i$ suggests that the modeled effect of either process is stronger than observed, hence the ensuing loss is overestimated. In section 4 we interpret the quantity $\Delta \mathbf{x}_k^i$ as an indicator of the loss brought



Figure 1. Dependence of equatorial pitch angle α_0 (a) and electron kinetic energy E_k (b) on L shell in a dipolar magnetic field, for the four pairs of (μ, K) investigated in the present study.

by both scattering by EMIC waves and magnetopause shadowing into the dynamics ofthe outer radiation belt.

²⁹⁰ 3 Long-term reanalysis of electron PSD

In this section we present the results obtained for the radial profiles of PSD based 291 on the assimilation of the above-mentioned four-satellite measurements into the VERB-292 3D model for the four-year period starting on 1 October 2012. We mainly focus on four 293 pairs of (μ, K) and show the corresponding equatorial pitch angle α_0 and electron ki-294 netic energy E_k in a dipolar magnetic field, in Figure 1. At the heart of the outer ra-295 diation belt (L = 4.5), for the chosen values of $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$, the equatorial pitch 296 angle is approximately 52°. Electron energies at L = 4.5 are 1.53 MeV for $\mu = 700$ MeV G⁻¹, 297 2.42 MeV for $\mu = 1500 \text{ MeV G}^{-1}$, 3.25 MeV for $\mu = 2500 \text{ MeV G}^{-1}$, and 3.6 MeV 298 for $\mu = 3000 \text{ MeV G}^{-1}$. 299

Panels (a) and (c) of Figure 2 show measured Van Allen Probes and GOES hourly 300 averaged electron PSD at $\mu = 700 \text{ MeV G}^{-1}$ and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$ and $\mu = 3000 \text{ MeV G}^{-1}$ 301 and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$, respectively. The results of the "full" data assimilation run are 302 illustrated in panels (b) and (d). The assimilated PSD is consistent with the original space-303 craft data and it indicates the improvement in coverage that reanalysis provides. Pan-304 els (e) and (f) depict the solar wind dynamic pressure P_{dyn} and the geomagnetic indices 305 Kp and Dst. The data assimilation runs for electron PSD at $\mu = 1500 \text{ MeV G}^{-1}$ and 306 $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$ and $\mu = 2500 \text{ MeV G}^{-1}$ and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$ are shown in Fig-307 ure A1 in the supplementary material. 308

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Figure 2. Evolution of electron PSD as a function of L^* and time from 1 October 2012 to 1 October 2016: (a) Van Allen Probes and GOES data, and (b) assimilated radial profile of PSD for $\mu = 700 \text{ MeV G}^{-1}$ and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$; (c) and (d) same as (a) and (b) but for $\mu = 3000 \text{ MeV G}^{-1}$ and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$; (e) evolution of solar wind dynamic pressure, and (f) geomagnetic activity Kp and Dst indices. The assimilative results of the combined reanalysis of electron PSD in this figure account for 3D diffusion, mixed pitch angle-energy diffusion, scattering by EMIC waves, and magnetopause shadowing (i.e. "full" run).

The reanalysis on panels (b) and (d) exhibit sudden dropouts and buildups of PSD. Figure 2 shows that dropouts in PSD often occur in association with sharp increases of solar wind dynamic pressure (e.g., Shprits et al., 2012; Turner et al., 2012; Ni et al., 2013). It is also worth noting that during the first half of our period under study, particularly between October 2013 and October 2014, geomagnetic activity was much weaker and less PSD enhancements were apparent than during 2015 and 2016.

³¹⁵ 4 Statistical analysis of loss processes via the innovation vector

In order to understand the loss due to scattering by EMIC waves and magnetopause shadowing in the outer radiation belt, we present plots of the normalized innovation \mathbf{x}^{i} and the difference of normalized innovations $\Delta \mathbf{x}^{i}$ (equation 12) for each of our four-year runs and each of our four chosen pairs of adiabatic invariants. We first bin the hourly



Figure 3. Occurrence of (a) Kp index, (b) solar wind dynamic pressure P_{dyn} , and (c) Dst index. Note that the y-axes are logarithmic. In plots (a) and (b) the blue, black, red, and green dashed lines denote the 75th, 96th, 98th, and 99th percentiles, respectively. In plot (c) the dashed lines indicate Dst values of -100nT, -50 nT, -30 nT, and 0 nT. In plot (b) P_{dyn} is binned each 0.5 nPa, and in plot (c) Dst is binned each 10 nT.

normalized innovation vector according to the Kp index, and compute the average as 320 a function of L^* and Kp. The same procedure is then followed binning the normalized 321 innovation by solar wind dynamic pressure. Figure 3 shows the occurrence of Kp, P_{dyn} , 322 and Dst from 1 October 2012 to 1 October 2016, and the coloured lines indicate differ-323 ent thresholds of geomagnetic activity. Figures A2 and A3 (in the supplementary ma-324 terial) present the distribution of the number of measurements binned by both Kp and 325 P_{dyn} . As expected, the distribution of samples is highly skewed towards low values of 326 Kp index and solar wind dynamic pressure. 327

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4.1 Scattering by EMIC waves

The normalized innovation vector \mathbf{x}^i as a function of L^* and Kp, before (run num-329 ber 2) and after incorporating EMIC waves (run number 1, i.e. "full") into the model, 330 is shown in the first two rows of Figure 4. Negative values (blue) denote additional loss 331 missing from the radiation belt model, and thus the KF subtracts PSD in order to com-332 pensate and match the observations, i.e. our model overestimates the electron PSD. The 333 last row presents the difference $\Delta \mathbf{x}^i$ as defined by equation 12 (namely the second row 334 minus the first row) in which the blue color denotes the area in L^* and Kp where EMIC 335 wave scattering operates and effectively scatters electrons. The positive yellow bins cor-336 respond to the intervals, mostly during disturbed times, when the inclusion of EMIC waves 337

in our model brings more loss than is observed. This may indicate that the parameterization we employ based on solar wind dynamic pressure does not always perform well during periods of high geomagnetic activity. The vertical dashed lines delineating the region of EMIC induced scattering loss are drawn considering a threshold of $\Delta \mathbf{x}^{i} = 10\%/\text{hr}$.



Figure 4. First row: normalized innovation vector \mathbf{x}^i of the reanalysis without EMIC scattering (run number 2); second row: normalized innovation vector \mathbf{x}^i of the "full" run (number 1); third row: difference of innovations $\Delta \mathbf{x}^i$, where the shaded region limited by the dashed line indicates the area where EMIC scattering is effective. The results are binned by L^* and Kp. Each column indicates a different pair of adiabatic invariants (μ, K) .

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As expected, EMIC waves do not affect the $\mu = 700 \text{ MeV G}^{-1}$ population, whereas they have a much more pronounced effect for higher energy electrons (e.g., Shprits, Subbotin, et al., 2013; Kersten et al., 2014; Usanova et al., 2014; Shprits et al., 2016). The upper extent of the region of loss due to EMIC waves moves from $L^* = 4.6$ (for $\mu =$ 1500 MeV G⁻¹), to $L^* = 5.2$ (for $\mu = 2500 \text{ MeV G}^{-1}$), and further beyond to $L^* =$ 5.6 as μ increases to 3000 MeV G⁻¹. In terms of Kp, the scattering effect is evident for $Kp \geq 3$. On average, the loss brought by EMIC waves is between 15%/hr and 30%/hr of the maximum PSD, peaking at $Kp \geq 5$ and between $L^* = 4$ and $L^* = 4.8$.

We also bin \mathbf{x}^i and $\Delta \mathbf{x}^i$ by L^* and P_{dyn} as presented in Figure 5. Similar to the results from Figure 4, including EMIC waves in the model decreases the overestimation

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Figure 5. Same format as Figure 4, binning the results by L^* and P_{dyn} . Results are presented in bins of 1 nPa between $P_{dyn} = 0$ and $P_{dyn} = 20$ nPa and 5 nPa between $P_{dyn} = 20$ nPa and $P_{dyn} = 50$ nPa.



Figure 6. Difference of innovations $\Delta \mathbf{x}^i$ before and after including EMIC waves in the model for different intervals of geomagnetic activity defined by Kp index (first row) and P_{dyn} (second row) as a function of L^* and μ . The shaded region limited by the dashed line indicates the area where EMIC scattering is effective.

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- of PSD, particularly for higher values of μ between $L^* = 4.2$ and $L^* = 5.6$. The scat-
- tering effect of these waves is evident for intervals with $P_{dyn} \ge 2$ nPa, and it exceeds
- ³⁵⁴ 20% of the maximum PSD for $P_{dyn} \ge 10$ nPa and $4.2 \le L^* \le 4.8$. Our choice of bin-
- ning the innovation by solar wind dynamic pressure follows the previous works from Usanova

et al. (2008) and Usanova et al. (2012) (and references therein), which demonstrated that strong magnetospheric compressions associated with high P_{dyn} may drive EMIC waves, and that the occurrence rate of EMIC activity in the dayside outer magnetosphere is controlled to a large extent by solar wind dynamic pressure.

The top row of figure 6 shows the difference, $\Delta \mathbf{x}^i$, across a range of the first adi-360 abatic invariant extending from $\mu = 300 \text{ MeV G}^{-1}$ ($E_k = 0.87 \text{ MeV}$ at the heart of 361 the outer belt) to $\mu = 3000 \text{ MeV G}^{-1}$, for both quiet and disturbed geomagnetic con-362 ditions as defined by the Kp index. For $Kp \leq 2.7$ (corresponding to the 75th percentile, 363 see the histogram in Figure 3) EMIC waves do not contribute to loss. The next three 364 intervals, defined by the 96th, 98th, and 99th percentiles, and characterizing active times, 365 show that the effect of these waves is confined to a triangular-shaped region defined by 366 $\mu \ge 900 \text{ MeV G}^{-1}$ ($E_k = 1.78 \text{ MeV}$ at the heart of the outer belt) and extending from 367 $L^* = 3.6$ to $L^* = 6$, on average. The loss brought in by EMIC waves increases from 368 $\sim 10\%$ /hr of the maximum PSD for Kp between 4.3 and 5 to $\sim 20\%$ /hr for Kp > 5.7 369 (equivalent to the 99th percentile), between $L^* = 4.2$ and $L^* = 4.8$. A similar pat-370 tern is observed in the second row of Figure 6, where the results are plotted for differ-371 ent intervals of solar wind dynamic pressure. With increase of P_{dyn} and μ , the loss ef-372 fect due to EMIC waves is enhanced and extends in radial distance from the Earth, max-373 imizing between $L^* = 4$ and $L^* = 4.8$. 374

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4.2 Magnetopause shadowing

An important process in producing fast electron dropouts is magnetopause shad-376 owing coupled with outward radial diffusion (Shprits et al., 2006). We inspect its effect 377 in our four-year reanalysis via the difference of innovations $\Delta \mathbf{x}^i$ when including and not 378 including this process (runs number 1 and 3, respectively), binned according to Kp and 379 P_{dyn} . Figure 7 shows that loss resulting from magnetopause shadowing extends from the 380 outer boundary for Kp = 3 down to $L^* = 3.6$ for Kp > 7. Therefore, we observe a 381 statistical picture where the loss region extends to lower L^* at a rate of $\sim 0.75 R_E$ per 382 increase of 1 Kp unit. Not surprisingly, the largest values of $\Delta \mathbf{x}^i$, and accordingly, the 383 biggest loss due to magnetopause shadowing (> 60%/hr of the maximum PSD), take 384 place with $Kp \ge 5$ and at $L^* \ge 4.6$. A similar pattern is observed when binning $\Delta \mathbf{x}^i$ 385 by solar wind dynamic pressure (Figure 8). Magnetopause loss starts at $P_{dyn} = 2$ nPa, 386 and they peak (between 50 and 70%/hr of the maximum PSD) when P_{dun} exceeds 10 387

³⁸⁸ nPa at $L^* \ge 4.8$, on average. In both figures, the diagonal dashed lines that define the region of loss correspond to a threshold of $\Delta \mathbf{x}^i = 30\%/\text{hr}$.



Figure 7. First row: normalized innovation vector \mathbf{x}^i of the reanalysis without magnetopause shadowing (run number 3); second row: normalized innovation vector \mathbf{x}^i of the "full" run (number 1); third row: difference of innovations $\Delta \mathbf{x}^i$, where the shaded region indicates the region where magnetopause shadowing operates. The results are binned by L^* and Kp. Each column indicates a different pair of adiabatic invariants (μ, K) .

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Figure 9 shows that as geomagnetic activity increases from quiet to disturbed times, loss moves inward affecting all values of μ from 300 to 3000 MeV G⁻¹. The effect is more pronounced for electrons with values of the invariant $\mu \ge 1500$ MeV G⁻¹ ($\Delta \mathbf{x}^i$ between 30%/hr and 50%/hr at $L^* \ge 5$) than for those with lower μ ($\Delta \mathbf{x}^i \sim 15\%$ /hr, on average), as the former drift faster, and thus, are depleted more quickly than less energetic ones. Likewise, increases in solar wind dynamic pressure also move the loss region due to magnetopause shadowing towards low L^* .

Lastly, we analyze our results by binning $\Delta \mathbf{x}^{i}$ (Figure 10) according to the geomagnetic activity Dst index (the corresponding histogram is shown in Figure 3 and the distribution of measurements binned by Dst is presented in Figure A4). For electrons with $\mu = 700 \text{ MeV G}^{-1}$ loss due to magnetopause shadowing exceed 50%/hr of the maxi-



Figure 8. Same format as Figure 7, binning the results by L^* and P_{dyn} .



Figure 9. Difference of innovations $\Delta \mathbf{x}^i$ before and after magnetopause shadowing in the model for different intervals of geomagnetic activity defined by Kp index (first row) and P_{dyn} (second row) as a function of L^* and μ .

⁴⁰² of loss is already evident at $Dst = -75$ nT. In other words, as μ increases, less geomag- ⁴⁰³ netic activity, as described by Dst , is required to observe the same percentage loss to ⁴⁰⁴ the magnetopause. It is also worth noting that, irrespective of the particle's energy, loss ⁴⁰⁵ due to magnetopause shadowing extends down to $L^* = 4.4$ during times with -100 nT ⁴⁰⁶ $< Dst \leq -50$ nT and even below to $L^* = 3.6$ when $Dst < -100$ nT.	401	mum PSD for $Dst < -100$ nT, whereas for those with $\mu = 3000$ MeV G ⁻¹ such level
⁴⁰³ netic activity, as described by Dst , is required to observe the same percentage loss to ⁴⁰⁴ the magnetopause. It is also worth noting that, irrespective of the particle's energy, loss ⁴⁰⁵ due to magnetopause shadowing extends down to $L^* = 4.4$ during times with -100 nT ⁴⁰⁶ $< Dst \leq -50 \text{ nT}$ and even below to $L^* = 3.6$ when $Dst < -100 \text{ nT}$.	402	of loss is already evident at $Dst=-75$ nT. In other words, as μ increases, less geomag-
the magnetopause. It is also worth noting that, irrespective of the particle's energy, loss due to magnetopause shadowing extends down to $L^* = 4.4$ during times with -100 nT $< Dst \leq -50 \text{ nT}$ and even below to $L^* = 3.6$ when $Dst < -100 \text{ nT}$.	403	netic activity, as described by Dst , is required to observe the same percentage loss to
due to magnetopause shadowing extends down to $L^* = 4.4$ during times with -100 nT due to magnetopause shadowing extends down to $L^* = 4.4$ during times with -100 nT due to $Dst \le -50 \text{ nT}$ and even below to $L^* = 3.6$ when $Dst < -100 \text{ nT}$.	404	the magnetopause. It is also worth noting that, irrespective of the particle's energy, loss
$_{405}$ $< Dst \le -50$ nT and even below to $L^* = 3.6$ when $Dst < -100$ nT.	405	due to magnetopause shadowing extends down to $L^{\ast}=4.4$ during times with $-100~{\rm nT}$
	406	$< Dst \leq -50$ nT and even below to $L^* = 3.6$ when $Dst < -100$ nT.



Figure 10. Difference of innovations $\Delta \mathbf{x}^i$ binned by L^* and Dst. The dashed lines indicate thresholds of -100 nT, -50 nT, and -30 nT. Results are presented in bins of 5 nT between Dst = 0 and Dst = -100 nT and 50 nT between Dst = -100 nT and Dst = -200 nT.

4.3 Comparison of electron PSD loss mechanisms

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The previous sections have quantitatively determined via data assimilation the ef-408 fect of EMIC scattering and magnetopause shadowing in the outer radiation belt. Here 409 we analyse both processes simultaneously and compare the magnitude and the spatial 410 extent (in L^*) of the loss induced by them. Figure 11 presents the difference $\Delta \mathbf{x}^i$ as a 411 function of radial distance averaged over the following levels of geomagnetic activity dur-412 ing our four-year period under study: $-30 \text{ nT} < Dst \le 0 \text{ nT}, -50 \text{ nT} < Dst \le -30$ 413 nT, and $Dst \leq -50$ nT. The minima of these curves are interpreted as the maximum 414 loss achieved by either of the mechanisms. In accordance with the above-mentioned re-415 sults, EMIC waves bring fewer loss than magnetopause shadowing. Loss due to EMIC 416 waves is mostly seen at L^* between 3.6 and 4.6, whereas loss due to magnetopause shad-417 owing is mainly evident at higher radial distances $(L^* \ge 4.8)$. 418

The minimum values of each curve of Figure 11, as well as their corresponding L^* 419 locations, are plotted in panels (a) and (b) of Figure 12. For the lowermost geomagnetic 420 activity level, with Dst between -30 nT and 0 nT, only loss due to magnetopause shad-421 owing is apparent, fluctuating between 2%/hr and 4%/hr of the maximum PSD at $L^* =$ 422 6.4. As Dst decreases between -50 nT and -30 nT, EMIC waves scatter electrons with 423 $\mu > 1000 \text{ MeV G}^{-1}$. Such loss reaches, at most, 5%/hr for the highest μ values, and 424 is observed from $L^* = 3.6$ to $L^* = 4.8$. At the same geomagnetic activity level, mag-425 netopause shadowing depletes electrons amounting from 10%/hr to 25%/hr of the max-426 imum PSD between $L^* = 5.8$ and $L^* = 6$. For the intervals with $Dst \leq -50$ nT, the 427 maximum EMIC induced scattering ($\Delta \mathbf{x}^i \leq 10\%$ /hr) occurs at 3.4 $\leq L^* \leq 4.2$, and 428



Figure 11. Difference of innovations $\Delta \mathbf{x}^i$ binned by L^* for different intervals of geomagnetic activity defined by Dst index for the indicated pairs of adiabatic invariants (μ, K) . Blue (red) lines denote loss due to EMIC scattering (magnetopause shadowing).

it clearly intensifies with increasing μ . More dramatic loss is introduced by magnetopause shadowing, ranging on average between 20%/hr and 50%/hr, at L^* between 5.2 and 5.6.

Besides investigating the value and L^* of the maximum PSD loss, we also deter-431 mine the location at which loss due to magnetopause shadowing starts dominating over 432 that due to EMIC wave scattering, by finding the crossing between the red and blue curves 433 in Figure 11. The corresponding L^* values are plotted in panel (c) of Figure 12. This 434 intersection is clearly energy-dependent, and for Dst between -50 nT and -30 nT, it 435 extends from $L^* = 4.1 \ (\mu = 1000 \text{ MeV G}^{-1})$ to $L^* = 5.2 \ (\mu = 3000 \text{ MeV G}^{-1})$, i.e. 436 out of the two loss processes inspected, EMIC waves are the main scattering agent be-437 low such location, whereas magnetopause shadowing plays a dominant role above it. For 438 more disturbed times, with $Dst \leq -50$ nT, this boundary moves inwards and fluctu-439 ates between $L^* = 3.5$ and $L^* = 4.4$. Nevertheless, it is worth noting that EMIC waves 440 (magnetopause shadowing) may deplete electrons above (below) such location. As an 441 example, for $Dst \leq -50$ nT and $\mu = 3000$ MeV G⁻¹, EMIC waves produce loss be-442 youd the intersection at $L^* = 4.4$, extending out to $L^* = 5$. Conversely, loss due to 443 magnetopause shadowing is already seen at $L^* = 4$. 444

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Figure 12. (a) Maximum loss (as defined by $\Delta \mathbf{x}^i$) due to EMIC scattering and magnetopause shadowing for the indicated levels of geomagnetic activity; (b) L^* location corresponding to the maximum loss; (c) L^* boundary separating two distinct mechanisms of electron PSD loss.

445 5 Discussion

This work employs four-years of spacecraft data which allows us to statistically quan-446 tify the effect of both loss processes over different levels of geomagnetic activity. We show 447 that scattering by EMIC waves induces loss from $L^* = 3.6$ to $L^* = 5.6$, particularly 448 between $L^* = 4$ and $L^* = 4.8$ during the most disturbed times. The resulting deple-449 tion amounts to between 10%/hr to 30%/hr of the maximum PSD. The effect of EMIC 450 waves is seen starting from $\mu = 900 \text{ MeV G}^{-1}$, and is energy-dependent, with higher 451 energy electrons being affected the most over a broader range of L^* . Our findings are 452 consistent with previous observational and modelling studies (e.g., Usanova et al., 2014; 453 Shprits et al., 2016; Drozdov, Shprits, Usanova, et al., 2017; Xiang et al., 2017), and val-454 idate the employed wave model, since we are able to reproduce the behaviour of EMIC 455 waves and the dynamics of the ultrarelativistic electron population. 456

Loss due to magnetopause shadowing is the strongest between $L^* = 5$ and $L^* =$ 6.6. Nevertheless, the depletion of electron PSD may extend further below $L^* = 4$ and reach between 50%/hr and 70%/hr of the maximum PSD, either for large values of ge-

omagnetic indices or for enhanced solar wind dynamic pressure. This is in accordance 460 with e.g. Shprits et al. (2012), who reconstructed a depletion of the radiation belt PSD 461 down to $L^* = 3$, based on data assimilation, for a very high value of P_{dyn} around 50 462 nPa. Similar conclusions on the correlation between electron PSD dropout events and 463 solar wind dynamic pressure pulses were reached by Ni et al. (2013), based on a one-year 464 reanalysis survey of multisatellite data. Such sharp increases of P_{dyn} clearly result in the 465 compression of the magnetopause and the removal of electrons originally on closed drift 466 orbits, with the most energetic populations affected to a larger extent. Statistically, for 467 the range of μ values considered in this study, we find that magnetopause shadowing tends 468 to deplete more electrons than EMIC wave interactions during disturbed times. 469

Based on our results we identify a μ - and geomagnetic activity-dependent bound-470 ary fluctuating between $L^* = 3.5$ and $L^* = 5.2$ defining two regions in space where 471 these two distinct loss mechanisms are mostly effective. EMIC induced scattering dom-472 inates below the boundary, whereas magnetopause shadowing coupled with outward ra-473 dial diffusion is active above it. Turner, Angelopoulos, Morley, et al. (2014) suggested 474 this boundary to be located at $L^* \sim 4$. Yu et al. (2013) found it to be around $L^* \sim$ 475 5, above which more than 90%/hr of the total loss is due to magnetopause shadowing 476 together with outward radial diffusion, and below which only 60%/hr can be explained 477 by this coupled mechanism. Dropouts, however, can encompass the entire outer radia-478 tion belt, and either mechanism can induce loss beyond the above-mentioned boundary. 479 In other words, magnetopause shadowing can deplete electrons below it, and EMIC waves 480 can efficiently scatter electrons beyond it, in particular during times of enhanced geo-481 magnetic activity. A similar conclusion with a boundary identified around $L^* \sim 4$ was 482 reached by Xiang et al. (2017) based on an investigation of three dropouts as observed 483 by Van Allen Probes. 484

Our statistical study relying on four years of data has shown that, in general, loss 485 due to magnetopause shadowing tends to exceed loss produced by EMIC scattering. Nev-486 ertheless, this is not always the case, as during disturbed conditions (i.e. geomagnetic 487 storms) the effect of EMIC waves can be comparable, or even exceed, the effect of mag-488 netopause shadowing. Figures 13 and 14 show two of these events, which correspond to 489 intense storms following the classification of Gonzalez et al. (1994). The maximum de-490 pletion due to both EMIC waves and magnetopause shadowing (between 10%/hr and 491 20%/hr of the maximum PSD) is observed during the main phase of each storm, with 492

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Figure 13. First row: evolution of Dst index for the geomagnetic storm with $Dst_{min} = -108$ nT on 14 November 2012 07 UT. The initial, main, and recovery phases are highlighted in blue, red, and green, respectively. Second row: difference of innovations $\Delta \mathbf{x}^i$ binned by L^* denoting losses due to scattering by EMIC waves for the indicated pairs adiabatic invariants (μ, K) during different phases of the storm. Third row: same as second row, for magnetopause shadowing.



Figure 14. Same as Figure 13, for the storm with $Dst_{min} = -93$ nT on 20 January 2016 16 UT.

smaller contributions during the initial phase and the beginning of the recovery phase.
In these events, loss due to EMIC waves dominates in the heart of the outer radiation
belt and is within the same order of magnitude as loss produced by magnetopause shadowing, demonstrating that EMIC waves play an indispensable role in the dynamics of
the ultrarelativistic electron population.

The effects of scattering by EMIC waves and magnetopause shadowing have been 498 studied individually in the current work, i.e. only one process was excluded from the model 499 at a time. However, these two mechanisms can act simultaneously and complement each 500 other in driving the dynamics of the outer belt. Magnetopause shadowing and the con-501 sequent outward radial diffusion develop negative PSD gradients at higher L shells (e.g., 502 Turner et al., 2012), while localized and fast loss driven by EMIC waves produces deep-503 ening minimums in PSD around $L^* = 3.5$ to $L^* = 4.5$ (e.g., Aseev et al., 2017; Sh-504 prits et al., 2017), and therefore can influence the rate of outward diffusion. The com-505 bination of both processes results in efficient dropouts of radiation belt electrons, cre-506 ating several localized peaks in PSD. Moreover, EMIC wave scattering is fast at low pitch 507 angles and significantly slower at high pitch angles, (e.g., Usanova et al., 2014; Drozdov, 508 Shprits, Usanova, et al., 2017). LCDS location is also pitch angle (or K) dependent (e.g., 509 Albert et al., 2018) and magnetopause shadowing affects mainly high pitch angles (e.g., 510 Roederer, 1967; West et al., 1972). As a result, both mechanisms can remove together 511 a broad range of particles. This can irreversibly alter the content of the outer belt and 512 can lead to almost total depletion of the pre-existing electron population. Future work 513 will focus on estimating the K dependence of scattering by EMIC waves and magnetopause 514 shadowing via the analysis of the innovation vector. 515

After adding EMIC waves and magnetopause shadowing in our model (i.e. perform-516 ing the "full" run) a region of positive innovation (in yellow and red) remains at $L^* >$ 517 4.2 and Kp > 6. This underestimation of electron PSD could be due to the fact that 518 our calculation of the LCDS, based on the IRBEM library, does not account for bifur-519 cating field lines (Albert et al., 2018), thus yielding a LCDS that is too close to Earth. 520 As a result, electron PSD is depleted in excess. Another explanation for this underes-521 timation of PSD is related to the electric field induced by the compression of the mag-522 netopause. Such electric field might mitigate some of the ensuing loss by radially trans-523 porting the electron population inwards (Michael Schulz, personal communication, May 524 2019). 525

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The technique we have presented in this study can be applied to other geophys-526 ical systems where the relative contribution of specific mechanisms needs to be quan-527 tified. This comes with several caveats, however. First of all, while our technique relies 528 on spacecraft observations, our findings are not completely independent of the assump-529 tions of the model, such as the times when EMIC waves operate or the location of the 530 LCDS. Second, the metric we have introduced, $\Delta \mathbf{x}^i$, does not indicate the actual num-531 ber of electrons lost (an integration would be necessary), but rather expresses the loss 532 in each time step as a function of the hourly maximum PSD. In this regard, we have cho-533 sen our normalization factor to be the maximum value of assimilated PSD over all L^* , 534 rather than the current state at each individual L^* , to avoid division by rather small val-535 ues, which would have yielded large percentage differences at some locations. Lastly, in 536 our case, errors in the model may arise, e.g. from the employed wave parameterizations 537 or the dynamic pressure threshold used to turn on EMIC waves in the model, and in turn, 538 may affect the reconstructed electron PSD and the innovation vector. Nevertheless, the 539 difference of innovations $\Delta \mathbf{x}^i$ can be used to indicate when discrepancies between pre-540 dictions and observations arise and to pinpoint possible sources of error in the model. 541 In our current study, values of $\Delta \mathbf{x}^i$ are mainly negative and hence, indicate that loss by 542 EMIC waves and magnetopause shadowing decrease the modeled PSD and generally bring 543 the model output closer to observations. 544

545 6 Conclusions

In this paper we perform four-year reanalysis of the outer electron radiation belt 546 by assimilating Van Allen Probes and GOES electron PSD measurements into our VERB-547 3D code. We study the innovation vector to characterize the effect of two distinct pro-548 cesses, namely scattering by EMIC waves and magnetopause shadowing, identifying where 549 (in L^*) and under which conditions (as described by geomagnetic indices Kp and Dst550 as well as solar wind parameter P_{dyn}) they operate. In comparison to previous studies, 551 our novel approach accounts and corrects for limited data coverage. We quantify the loss 552 produced by these mechanisms through a comparison of the innovation before and af-553 ter their inclusion in the model, and we also explore the μ dependence (from 300 to 3000 MeV G⁻¹) 554 of both processes. 555

We find that EMIC waves mainly scatter electrons with $\mu \ge 900 \text{ MeV G}^{-1}$ between $L^* = 4$ and $L^* = 4.8$, and the ensuing depletion may reach between 10%/hr to

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⁵⁵⁸ 30%/hr of the maximum PSD under disturbed geomagnetic conditions. Magnetopause ⁵⁵⁹ shadowing is shown to be mostly effective from $L^* = 5$ to $L^* = 6.6$ and the induced ⁵⁶⁰ loss may amount up to between 50%/hr and 70%/hr of the maximum PSD, affecting all ⁵⁶¹ electrons with μ values from 300 to 3000 MeV G⁻¹. We also identify an energy- and ge-⁵⁶² omagnetic activity-dependent boundary located between $L^* = 3.5$ and $L^* = 5.2$ sep-⁵⁶³ arating both mechanisms. Scattering by EMIC waves is active below it, while magne-⁵⁶⁴ topause shadowing dominates above it.

Future studies will be aimed towards extending our DA methodology and innova-565 tion vector analysis to quantify and assess the contribution of other processes to the dy-566 namical evolution of electron PSD, such as pitch angle scattering by plasmaspheric hiss 567 or energy diffusion by chorus waves. Same methodology can be also applied to the anal-568 ysis of the ring current dynamics. Furthermore, the role of scattering by EMIC waves 569 and magnetopause shadowing will be inspected in detail for selected events, such as the 570 110 geomagnetic storms identified by Turner et al. (2019) during the Van Allen Probes 571 era, in order to determine the percentage of dropout events dominated by either mech-572 anism. Moreover, our framework can also be employed to assimilate measurements from 573 the last three years of Van Allen Probes (October 2016 to October 2019) and from on-574 going missions such as Arase (Miyoshi et al., 2018). All these efforts will be ultimately 575 directed towards achieving a better understanding of the dominant mechanisms during 576 radiation belt enhancements and dropouts. 577

578 Appendix A Supplementary material

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Figure A1. Evolution of electron PSD as a function of L^* and time from 1 October 2012 to 1 October 2016: (a) Van Allen Probes and GOES data, and (b) assimilated radial profile of PSD for $\mu = 1500 \text{ MeV G}^{-1}$ and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$; (c) and (d) same as (a) and (b) but for $\mu = 2500 \text{ MeV G}^{-1}$ and $K = 0.11 \text{ G}^{0.5} \text{ R}_{\text{E}}$; (e) evolution of solar wind dynamic pressure, and (f) geomagnetic activity Kp and Dst indices. The assimilative results of the combined reanalysis of electron PSD in this figure account for 3D diffusion, mixed pitch angle-energy diffusion, scattering by EMIC waves, and magnetopause shadowing (i.e. "full" run).



Figure A2. Distribution of the number of satellite observations employed in the reanalysis of PSD binned by L^* and Kp for the indicated pair of adiabatic invariants μ and K.



Figure A3. Distribution of the number of satellite observations employed in the reanalysis of PSD binned by L^* and P_{dyn} for the indicated pair of adiabatic invariants μ and K.



Figure A4. Distribution of the number of satellite observations employed in the reanalysis of PSD binned in L^* and *Dst* for the indicated pair of adiabatic invariants μ and *K*.

- of the dynamic evolution of the Van Allen belts using multiple satellite measurements"
- and the Helmholtz Association Recruiting Initiative.

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