Inherent length scales of periodic mesoscale density structures in the solar wind over two solar cycles

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November 30, 2022

Abstract

It is now well-established through multiple event and statistical studies that the solar wind at 1 AU contains contains periodic, mesoscale (L~100-1000Mm) structures in the proton density. Composition variations observed within these structures and remote sensing observations of similar structures in the young solar wind indicate that at least some of these periodic structures originate in the solar atmosphere as a part of solar wind formation. Viall et al. [2008] analyzed 11 years of data from the Wind spacecraft near L1 and demonstrated a recurrence to the observed length scales of periodic structures in the solar wind proton density. In the time since that study, Wind has collected 14 additional years of solar wind data, new moment analysis of the Wind SWE data is available, and new methods for spectral background approximation have been developed. In this study, we analyze 25 years of Wind data collected near L1 and produce occurrence distributions of statistically significant periodic length scales in proton density. The results significantly expand upon the Viall et al. [2008] study, and further shows a possible relation of the length scales to solar "termination" events.

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Key Points: 6

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7	•	25 years of Wind solar wind data are analyzed for periodic mesoscale structures
8		in the proton density
9	•	Periodic density structures recur with particular length scales, suggesting solar for
10		mation

• The observed length scales show a potential relationship to solar termination events 11

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12 Abstract

It is now well-established through multiple event and statistical studies that the solar 13 wind at 1 AU contains contains periodic, mesoscale ($L \sim 100 - 1000$ Mm) structures 14 in the proton density. Composition variations observed within these structures and re-15 mote sensing observations of similar structures in the young solar wind indicate that at 16 least some of these periodic structures originate in the solar atmosphere as a part of so-17 lar wind formation. Viall et al. (2008) analyzed 11 years of data from the Wind space-18 craft near L1 and demonstrated a recurrence to the observed length scales of periodic 19 structures in the solar wind proton density. In the time since that study, Wind has col-20 lected 14 additional years of solar wind data, new moment analysis of the Wind SWE 21 data is available, and new methods for spectral background approximation have been 22 developed. In this study, we analyze 25 years of Wind data collected near L1 and pro-23 duce occurrence distributions of statistically significant periodic length scales in proton 24 density. The results significantly expand upon the Viall et al. (2008) study, and further 25 shows a possible relation of the length scales to solar "termination" events. 26

27 Plain Language Summary

The plasma and magnetic field in the solar atmosphere flows away from the Sun, 28 filling interplanetary space. This plasma is called the solar wind, and it constantly bom-29 bards all of the planets in the solar system. The solar wind is comprised of mesoscale 30 structures - larger than scales where particle dynamics are important, but smaller than 31 global scales - of increased density, and therefore pressure. A subgroup of mesoscale den-32 sity structures are of order the size of Earth's magnetosphere, and often quasi-periodic. 33 These periodic density structures are an important driver of dynamics in Earth's space 34 environment. In this study, we examine the statistics of the size scales of these structures 35 using 25 years, or approximately two solar cycles, of solar wind data measured by the 36 Wind spacecraft. We confirm earlier work showing a persistence of particular length scales 37 of the periodicities, and find a possible relation of the length scales to the end of a Hale 38 magnetic cycle. In addition to their driving of magnetospheric dynamics, periodic den-39 sity structures are a tracer of solar wind formation. Their lengths scales and evolution 40 are an important constraint of solar wind formation. 41

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42 **1** Introduction

The solar wind contains structures at vastly different scales, from small scale 1-2 43 minute magnetic holes (Winterhalter et al., 1994) to large scale coronal mass ejections 44 and stream interaction regions (Richardson, 2018; Kilpua et al., 2017). There is a rich 45 spectrum of structure between these two extremes, at "mesoscales", which here we de-46 fine as scale sizes $L \sim 100$ to several 1000 Mm, or equivalently time scales of $t \sim a$ few 47 minutes to several hours. Solar wind structures throughout this range of 'mesoscale' sizes 48 have been identified in Solar Terrestrial Relations Observatory (STEREO) white light 49 remote sensing data down to the resolution of the imager (DeForest et al., 2018; Viall 50 & Vourlidas, 2015). They have also been observed in situ, in the form of magnetic field 51 flux rope structures as small as 50 Mm (Murphy et al., 2020), in plasma density at scales 52 between 50 and 2000 Mm (Stansby & Horbury, 2018), and in combinations of magnetic 53 and plasma signatures (Borovsky, 2008; Sanchez-Diaz et al., 2019; Matteo et al., 2019; 54 Rouillard, Lavraud, et al., 2010; Rouillard et al., 2020). 55

A subset of mesoscale solar wind structures are quasi-periodic proton density en-56 hancements, termed periodic density structures (PDSs). They were initially discovered 57 through event studies that showed a direct correspondence between magnetospheric pul-58 sations in the mHz range (periodicities of a few minutes to a few hours) and a one-to-59 one correlation with discrete frequencies in the solar wind density observed in the up-60 stream solar wind (Kepko et al., 2002; Kepko & Spence, 2003). The apparent frequency 61 of a PDS as it flows past Earth or an in situ spacecraft is related to the radial length scale 62 of the structure as $f_{pds} = V_{sw}/L_{pds}$, where f_{pds} is simply the inverse of the ΔT between 63 each density enhancement. Numerous event studies have observed direct links between 64 the periodicities in solar wind density and periodicities in radar (Stephenson & Walker, 65 2002; Fenrich & Waters, 2008), ionospheric (Dyrud et al., 2008), and ground magnetome-66 ter (Villante et al., 2007; Villante & Tiberi, 2016) observations, at frequencies from \sim 67 4 mHz down to ~ 0.2 mHz. Viall, Kepko, and Spence (2009) identified statistically sig-68 nificant frequencies observed in 11 years of Wind proton density data near L1 and 10 69 years of dayside GOES magnetospheric B_z data. They showed that both the solar wind 70 and dayside magnetosphere contained recurrent, similar sets of observed frequencies be-71 tween ~ 0.5 -4.0 mHz, which lie in the Pc5-6 frequency range. These mHz frequencies cor-72 respond to the smaller mesoscale structures, $\sim 100-1000$ Mm, at nominal solar wind 73 speeds. 74

These length scales are on the order of the dayside magnetosphere (~ 80 Mm) and 75 larger, and therefore quasi-statically drive magnetospheric pulsations through periodic 76 dynamic pressure changes. Hence, eve the smallest mesoscale solar wind structures are 77 effective at creating a global magnetospheric response. It is this coherent, global mag-78 netospheric response to solar wind structures at this size scale that motivates our lower 79 limit definition for 'mesoscale'. We note that the equivalent frequency of an 80 Mm struc-80 ture at nominal solar wind speeds is ~ 4 mHz. At shorter length scales, the solar wind 81 structures are smaller than the dayside magnetosphere, and the interaction is no longer 82 quasi-static. Therefore, there is a general split between solar wind directly-driven oscil-83 lations at f < 4 mHz, and internally supported oscillations, such as cavity mode or field-84 line resonances, at around f > 4 mHz (Hartinger et al., 2013). 85

Since the initial papers describing the existence of periodic density structures in 86 the solar wind, there have been several attempts to identify their source. A key measure-87 ment are the occurrence distributions of statistically significant frequencies and length 88 scales observed in solar wind proton density measurements. Viall et al. (2008) found sta-89 tistically significant bands of periodic length scales, and Viall, Kepko, and Spence (2009) 90 found similar bands in frequency. These distributions of spectral peaks in solar wind den-91 sity consists of 3 sources: in situ generated structures (e.g., via turbulence); 'false pos-92 itives' at a rate determined by the chosen confidence thresholds and appropriateness of 93 the background spectral fit; and periodic density structures injected through the pro-94 cess of solar wind formation. The first two of these sources would generate a smoothly 95 varying distribution of observed periodicities, rather than the recurrent sets found by 96 Viall et al. (2008) and Viall, Kepko, and Spence (2009), while the third could produce 97 localized occurrence distribution peaks. Although it is theoretically possible that there 98 exists an MHD instability that could generate periodic structures in transit to 1 AU, for 99 example a slow mode wave (Hollweg et al., 2014), to date there has been no published 100 observations of such instabilities creating periodicities on mesoscales. Furthermore, Viall, 101 Spence, and Kasper (2009) found fewer recurrent solar wind periodicities analyzing the 102 data in time-frequency space than Viall et al. (2008) did analyzing the same data in length 103 scale-wavenumber space. This suggests structures advecting with solar wind streams, rather 104 105 than locally generated oscillations or waves at particular frequencies.

Multiple lines of evidence suggest that periodic solar wind density structures are tracers of solar wind formation. In situ observations show composition, magnetic field,

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and electron strahl changes that indicate magnetic reconnection effects that could only 108 have occurred during solar wind release and acceleration (Viall, Spence, & Kasper, 2009; 109 Kepko et al., 2016; Matteo et al., 2019). Matteo et al. (2019), using Helios data, found 110 anisotropic temperature changes within PDSs that are not observed near L1, consistent 111 with solar formation followed by temperature isotropization while in transit. Remote imag-112 ing studies using the Solar Terrestrial Relations Observatory (STEREO)/Sun Earth Con-113 nection Coronal and Heliospheric Investigation (SECCHI) white light instruments have 114 identified PDSs in the solar corona as close as 2.5 solar radii, observed as they acceler-115 ate with the surrounding solar wind (Viall et al., 2010; Viall & Vourlidas, 2015; DeFor-116 est et al., 2016, 2018). Rouillard et al. (2020) confirmed the relationship between mesoscale 117 density structures observed in images and in situ by tracking larger streamer blobs with 118 embedded ~hour long structures from STEREO SECCHI to their impact with Parker 119 Solar Probe. In short, it is now clear that the solar wind is often formed of quasi-periodic 120 mesoscale plasma density structures released as a part of solar wind formation. 121

Three factors motivate this investigation. First, while previous studies used only 122 11 years of data, 25 years of Wind solar wind data are now available, which allows an 123 examination of evolution of the recurrent length scales as a function of two complete so-124 lar cycles. Second, the Wind SWE data that the Viall et al. (2008) and Viall, Kepko, 125 and Spence (2009) statistical studies analyzed have been reprocessed (Kasper et al., 2006), 126 providing an opportunity to test the accuracy and precision of those previous results. 127 Third, recent progress on techniques used to identify statistically significant spectral peaks 128 has shown that there are limitations to using the AR(1) background assumption, and 129 suggests a different background assumption may be more appropriate (Vaughan et al., 130 2011).131

- 132 2 Methods
- 133

2.1 Data processing and quality checks

We follow the general process of data preparation and spectral analysis as the Viall et al. (2008) study. We used the proton number density and proton velocity measured by the Solar Wind Experiment (SWE) Faraday Cup onboard the Wind spacecraft (Ogilvie et al., 1995) to examine the characteristics of mesoscale periodic density structures between $\sim 80-1000$ Mm observed over the full lifetime of Wind to this point, from 1995-

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2019. In the time since the Viall et al. (2008) study, Kasper et al. (2006) developed a 139 new fitting technique to calculate separately the moments of the proton and Helium dis-140 tributions from the Wind Faraday Cup data. This new dataset, since it takes into ac-141 count the bi-maxwellian nature of the solar wind, provides a more accurate measure of 142 the proton number density and velocity. The primary impact of that reprocessing on this 143 study is that the velocity increased on average by a few percent, which increases the length-144 scales by a few percent, and the proton density decreased slightly. These changes are slightly 145 more pronounced when the velocity is high. 146

For continuity with and comparison to the Viall et al. (2008) study, we follow the 147 same processing steps prior to the spectral analysis to produce length series segments 148 9072 Mm in length, overlapping by 252 Mm. We first converted the time series of solar 149 wind proton density, n(t), to a length series, L(t), by multiplying each time step by the 150 radial velocity, $v_x(t)$. Since each step has a different velocity, this produces an irregu-151 larly sampled series that is not compatible with Fourier analysis, and must be interpo-152 lated to a fixed ΔL . Yet, due to the wide spread in solar wind speeds, resampling to a 153 single common length step would lead to oversampling at low speeds, and undersampling 154 at high speeds. We therefore produced two sets of interpolated, evenly sampled segments. 155 Segments with $\overline{v_x} \geq 550$ km/s were termed "fast" segments, with $\Delta L_f = 56.7$ Mm, 156 while those with $\overline{v_x} < 550$ km/s were termed "slow" segments, with $\Delta L_s = 35.4$ Mm. 157 For slow wind, 9072 Mm is approximately 6 hours of data at the median slow solar wind 158 speed, and the 35.4 Mm ΔL is approximately equivalent to the SWE instrument sam-159 pling rate (typically 90-100 seconds) converted to length. Similarly, for the fast wind 9072 160 Mm is approximately 4 hours, and 56.7 Mm is the equivalent sampling rate multiplied 161 by the median fast speed. While the Nyquist between between "fast" and "slow" segments 162 is different, the combination of ΔL and number of points in each segment keeps the spec-163 tral resolution the same. Note that the categorization of fast and slow data segments is 164 not an attempt at a physics-based classification of solar wind type; it is well-known that 165 speed is not the best physics-based classification (Zurbuchen et al., 2002; Roberts et al., 166 2020; Borovsky, 2012). Rather, these two categories are only the result of the effective 167 sampling rate of the data segments. 168

Figure 1 shows both a slow (panels a-c) and fast (panels d-f) segment of solar wind data comparing the original (blue) and reprocessed (red) SWE data as a time series, and both datasets converted into a length series (panels c and f). These segments are typ-

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- ical of other intervals in that they exhibit the very slight increase of a few percent in ve-
- locity in the reprocessed data. The reprocessed data also show differences in higher fre-quency variations, particularly for the fast wind (see Figure 1d).



Figure 1. Comparison of the original (blue) and reprocessed (red) solar wind data from Wind SWE for a representative segment, for both slow (left) and fast (right). Reprocessed data show slightly lower density (a and d), slightly higher velocity (b and e), and the high frequency variations are of lower amplitude than the original data.

For each data segment, we imposed data quality requirements to minimize spuri-175 ous spectral signals, and do not analyze segments that failed the data quality check. We 176 required that the Wind spacecraft be located at least 50 Earth radii (R_E) upstream of 177 Earth, to exclude any solar wind collected within or near Earth's magnetosphere, or that 178 could be contaminated with foreshock activity. This reduced the number of segments dur-179 ing the early part of the Wind mission, when it occasionally enters Earth's magnetosphere. 180 We remove single point data spikes and interpolated over them. We excluded any seg-181 ment that contained more than 10% flagged or missing data over the entire segment, or 182 3% consecutive flagged or missing data. Finally, we excluded segments that contained 183 discontinuous jumps (e.g., shocks) in the number density, since this would introduce "ring-184

ing" in the spectra. To determine a discontinuous jump, we subtracted a third order polynomial fit to the data segment, and discarded segments that contained changes in 5-point running averages that exceeded 3.7 standard deviations of the detrended median. The fraction of segments that passed these quality control checks is shown in Figure 2. There is a slight decrease in the number of segments that passed these checks using the reprocessed SWE data for the slow wind compared to the original data used by Viall et al. (2008).



Figure 2. The percentage of the slow and fast solar wind length series segments that passed the quality control checks, and that were analyzed for periodic density structures. We also include the percentage of segments that passed these same quality checks in the original Viall et al. (2008) study. The differences are due to the reprocessed Wind SWE data.

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2.2 Spectral analysis and peak detection

We perform spectral analysis on each segment that passed the quality checks. We identify statistically significant spectral speaks using an amplitude test and a harmonic F-test. For the amplitude test, we calculate the spectra, estimate the background fit, then identify statistically significant peaks above this background. We use the segments in Figure 1c and 1f to demonstrate the process, and present the results in Figure 3. Estimation of the spectra relies on the multitaper method (MTM), in which multiple, or-

thogonal Slepian tapers are convolved with the data segment to provide multiple, inde-199 pendent estimates of the spectra (Thomson, 1982). While producing a robust spectral 200 estimate, this technique reduces the effective frequency resolution of the data as a func-201 tion of the number of tapers chosen, K, to $2pf_R$, where $f_R = 1/(N\Delta L)$ is the Rayleigh 202 frequency, and p = (K + 1)/2. In this study we used 5 Slepian tapers, leading to an 203 effective resolution of $6f_R = 6.6 \times 10^{-4} \text{ Mm}^{-1}$. We zeropad the data segments by a 204 factor of 10 prior to calculating the spectral estimates. In Figure 3a and 3c we plot MTM 205 spectra for the fast and slow length series segments shown in Figure 1, for both the orig-206 inal and reprocessed data. Note that the X-axis is in units of wavenumber Mm^{-1} , and 207 we also list the equivalent length scale. Both the original and reprocessed data sets show 208 similar spectral characteristics at the longer length scales (lower wavenumbers), but dif-209 fer slightly at the smaller length scales (higher wavenumbers); the differences are more 210 pronounced in the fast wind spectra. These trends are generally persistent across all seg-211 ments, and is consistent with the reprocessed data having lower noise. 212

Viall et al. (2008), following Mann and Lees (1996), modeled the spectral background under the assumption that the observations x_i , at point t_i , followed an auto-regressive AR(1) process, such that

$$x(t_i) = ax(t_{i-1}) + \epsilon_i \tag{1}$$

where *a* is the degree of correlation between sequential data points, and ϵ is random noise with zero mean (white noise). The limit of a = 0 produces a purely white noise spectrum, while larger values of *a* produce more strongly red-noise data series. The analytical spectrum of (1) is

$$S_{AR1}(f) = S_0 \frac{1 - a^2}{1 - 2a\cos(\pi f/f_N) + a^2}$$
(2)

where $S_0 = \sigma^2/(1-a^2)$ is the average value of the power spectrum, and σ^2 is the vari-220 ance of the white-noise. We fit (2) via least-squares to the spectra computed using the 221 MTM to produce an estimation of the background under the assumption of red+white 222 noise, and confidence levels are determined relative to that background. AR(1) background 223 fits and 95% confidence levels for the original and reprocessed datasets, for the fast and 224 slow segments, are shown in Figure 3a and c, overlaid on the MTM amplitude spectra. 225 The background AR(1) fit for both the original and reprocessed data are quite similar 226 for the slow wind, with calculated values of a = 0.836 and a = 0.846, respectively. For 227 the fast wind, however, the spectra and AR(1) fits are quite different, due to reduced high 228 frequency power in the reprocessed data, with a = 0.792 and a = 0.883 for the origi-229



Figure 3. MTM spectra and F-test for both the slow (left) and fast (right) segment shown in Figure 1, and for both the original (blue) and reprocessed (red) Wind data. We have plotted an AR(1) background fit for both datasets, with the 95% confidence level. Peaks that simultaneously pass the amplitude and F-test are marked with half circles for both original (blue) and reprocessed (red) data.

nal and reprocessed data, respectively. For both fast and slow wind, the AR(1) background fits lie well above the background at shorter scales (higher wavenumber), suggesting AR(1) may not be a good background assumption. We return to this in the next section.

The determination of a significant spectral peak, in this example wavenumbers that 233 have spectral power that exceed the 95% confidence threshold, is complicated by two is-234 sues. First, by definition power spectrum and confidence levels produce false positives 235 at the rate determined by the confidence thresholds (Thomson, 1982; Mann & Lees, 1996). 236 That is, for each frequency tested for significance, for a 95% test, e.g., there is a 5% prob-237 ability of exceeding the threshold. These false positives would be randomly distributed 238 in frequency, and therefore could not produce the types of preferential occurrence dis-239 tributions identified by Viall et al. (2008). To minimize these "false positives", in addi-240

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tion to the amplitude test, we apply a second type of spectral test, the harmonic F-test, 241 which is independent of the background fit (Mann & Lees, 1996). The amplitude test 242 requires a signal to have strong power, but does not explicitly test the discrete nature 243 of the power enhancement. On the other hand, the harmonic F-test tests for phase co-244 herent signals, but does not test the power contained in those signals. As in Viall et al. 245 (2008) we require that a spectral peak pass both the narrowband (amplitude) and F-246 test simultaneously to be considered significant and counted in our statistics. The pre-247 cise value of the peak we identify is fixed to the maximum F-test frequency within the 248 spectral amplitude band that exceeds the threshold. Because a peak has to pass both, 249 independent, tests simultaneously at the 95% level, our confidence threshold in appli-250 cation is significantly higher than 95%. Assuming that the false positives from the two 251 tests are uncorrelated, requiring that a signal pass both tests is analogous to testing at 252 a 99.75% confidence threshold. The second issue in identifying significant spectral peaks 253 is that the choice of the background noise model, while not affecting the F-test, affects 254 the narrowband (amplitude) test, an issues we discuss in the next section. 255

In Figure 3b and d we show the F-test for the representative segments, and we in-256 dicate peaks that pass both the narrowband and F-test at the 95% level with half-circles. 257 Note that many peaks pass the harmonic F-test with little power, and are therefore not 258 identified as significant in this combined test. Similarly, there are several amplitude peaks 259 that exceed the amplitude threshold, but not the F-test. For example, the amplitude 260 peak at L = 200 Mm in the slow wind, while significant in terms of spectral amplitude, 261 was not considered phase coherent by the F-test, and therefore was not considered sig-262 nificant. Since the F-test is a test for phase coherence, our study likely undercounts so-263 lar wind signals that have significant power but are not precisely phase coherent. As such, 264 results that use this technique should be considered a lower bound. 265

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2.3 Background estimation

The narrowband (amplitude) spectral test is a measure of the power of a discrete signal relative to a background spectra. The AR(1) process assumption (Equation 1) is widely used, since it is reasonable to expect a physical system to have memory. However, whether that memory takes the precise form of the AR(1) in any particular segment of solar wind data is currently impossible to know *a priori*. Indeed, Figures 3a and c shows that the AR(1) does not fit the highest and lowest wavenumbers well. We find

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this to be a persistent characteristic of the AR(1) fit when applied to the solar wind number density data. In effect, this bias imposes a slightly higher or lower confidence threshold than 95% across the spectra, and indicates that the solar wind may not be modeled well as an AR(1) process for the ~ 6 hour windows we consider here.

The paleoclimatology community has studied the AR(1) background assumption extensively, where the choice of noise model impacts the ability to detect cycles in the stratigraphic record. In response to these concerns, Vaughan et al. (2011) suggest a bending power law (BPL) background spectrum fit



Figure 4. A comparison of three different background assumptions for the solar wind intervals shown in Figure 1. Shown are an AR(1) (red), a BPL (green) and power law (blue). For clarity, we have not plotted the 95% confidence levels. Peaks that simultaneously pass the amplitude and F-test at 95% are marked for the different fits. The spectral background model parameters are N = 24.33, $\beta = -0.51$, $\gamma = 1.87$, $f_b = 1.8 \times 10^{-4}$ Mm⁻¹ for BPL slow wind; p = -1.74 for PL slow wind; N = .02, $\beta = .02$, $\gamma = 2.26$, $f_b = 3.2 \times 10^{-4}$ for BPL fast wind; p = -2.1 for PL fast wind.

$$S_{BPL}(f) = \frac{Nf^{-\beta}}{1 + (f/f_b)^{\gamma - \beta}}$$
(3)

which has the AR(1) as a special case, and performs well in mixed noise spectra. Here N is the normalization, β is the spectral slope index at low frequencies, γ is the spectral slope index at high frequencies, and f_b is the frequency at which the bend occurs. For low values of f_b , the BPL reduces to a straight power law with spectral slope $-\gamma$.

The BPL fit, and the 95% confidence level, is shown in Figure 4a and b in green 285 for the same segments shown in Figure 3. Note how the BPL is a better representation 286 of the background at both the higher and lower wavenumbers compared to AR(1) (red). 287 We also plot a straight power law (blue) with spectral slope, p, for both slow and fast 288 segments for reference. There is consistency in the identified peaks using the different 289 background assumptions, with the BPL assumption producing fewer peaks in the slow 290 wind segment. This tendency for BPL to identify fewer significant peaks than AR(1), 291 particularly at lower frequencies, is a consistent feature across the entire 25-year study. 292

The BPL is flexible in that it allows for an AR(1) solution, a single power law, and a host of solutions in between. Since the BPL approximates the solar wind background spectra better than AR(1), and because it is more versatile than a straight power law, we utilize BPL as one of the two background assumptions we use for our statistical study. For consistency with Viall et al. (2008) we also run the analysis with an AR(1) background estimate.

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2.4 Occurrence Distributions

We applied the data processing and spectral analysis methods described above to 300 the reprocessed solar wind measured by the Wind spacecraft from 1995-2019. For each 301 segment we determine statistically significant peaks that pass the amplitude and F-tests 302 simultaneously, for both BPL and AR(1) background assumptions. We create separate 303 occurrence distributions (ODs) of the statistically significant lengths (inverse wavenum-304 bers) identified using the AR(1)+F-test and BPL+F-test criteria. For each set, we com-305 pute occurrence distributions over overlapping, three-year intervals, with bins of width 306 $6f_R$, the effective resolution of the MTM with our choice of K = 5, stepping by $3f_R$ 307 for each subsequent bin. The inverse of wavenumber is length, and the wavenumber res-308 olution bins of $6f_R$ corresponds to 3.5 Mm near the Nyquist and 1500 Mm near zero wavenum-309 ber. For each 3-year window, we applied the bootstrap technique (N = 500) to esti-310

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311 312 mate the uncertainty of local peaks on the histogram, and calculated a median histogram, median fit (5 point moving mean), and standard deviation from these 500 instantiations.

To demonstrate this process we show the median histograms, representing an oc-313 currence distribution, for 1995-1998 for both the fast and slow solar wind in Figure 5a 314 and b, with 2σ standard deviation bars determined via the bootstrap method (Efron & 315 Tibshirani, 1993). Visually, these histograms exhibit locally enhanced counts for partic-316 ular lengthscale bands, with strong correlation between the occurrence enhancements us-317 ing the AR(1) and BPL spectral background fits. The residuals (Figure 5b and d) high-318 light the similarity in local occurrence enhancements between the AR(1) and BPL his-319 tograms, despite the differences in the overall shape of the occurrence distributions. We 320 use the bootstrapped occurrence distributions to determine statistically significant oc-321 currence enhancements as those points that are $> 2\sigma$ above the background fit. These 322 are highlighted with circle in Figures 5b and d, and with thick lines in Figure 5a and c. 323

Importantly, although the AR(1) and BPL background models produce different 324 overall shapes of the occurrence distributions, they produce similar residuals, and sim-325 ilar occurrence enhancements are identified as statistically significant with the bootstrap 326 method for each. For the slow wind, the OD determined with the AR(1) assumption ex-327 hibits a steep slope on the short length scale (higher wavenumber) end, consistent with 328 the examples shown in Figures 3 and 4. The BPL assumption does not exhibit this bias, 329 which provides confidence for local occurrence enhancements within this region (between 330 $\sim 90 - 150$ Mm). For example, there is a local occurrence enhancement identified in 331 the BPL OD near 110 Mm as $> 2\sigma$ significant, on top of a relatively flat part of the dis-332 tribution. In the AR(1) OD, this shows up as a relatively small local enhancement, and 333 appears in the residual histogram as well, but is not significant at the 2σ level. In ad-334 dition, the ODs produced with the BPL assumption identify $\sim 50\%$ fewer significant 335 peaks than those with the AR(1) assumption. This trend is consistent throughout the 336 25-year interval, and indicates that the BPL is likely a better approximation for the so-337 lar wind background spectra, with fewer false positive detections. Despite the difference 338 between the AR(1) and BPL results in absolute counts, the relative amplitude of the en-339 hancements in the occurrence distribution are similar between the two background model 340 assumptions. 341

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Figure 5. 3-year occurrence distributions for 1995-1997 for the slow and fast solar wind calculated for both the AR(1) (red) and BPL (blue) spectral background assumptions. Vertical bars represent $\pm 2\sigma$ standard deviation. Length scales that are greater than 2σ above the median fit (dashed lines) are shown in thick lines in (a) and (c), where we have extended the significant length scale range by $f_R/2$ in either direction. The residual distributions, obtained by subtracting the median fits from the occurrence distributions, are shown in (b) for the slow and (d) for the fast wind. Circles denote points that exceed 2σ .

342 3 Results

We ran the entire 25 year Wind SWE dataset through the analysis process described 343 in Section 2. Figure 6 shows the percentage of analyzed segments that contained at least 344 1 statistically significant peak that simultaneously passed the amplitude and F-test at 345 the 95% confidence levels, for each of the AR(1) and BPL background assumptions, com-346 pared to the Viall et al. (2008) study. Viall et al. (2008), using the original Wind data, 347 showed an increasing trend with time of the fraction of segments containing ≥ 1 sta-348 tistically significant frequency, implying a trend with solar cycle. This trend does not 349 appear in the reprocessed data. Instead, there is a relatively consistent number of sig-350 nificant radial-length peaks identified in segments during the 25-year interval, with the 351 BPL background assumption producing consistently fewer statistically significant peaks 352 than AR(1). 353



Figure 6. The percentage number of segments containing ≥ 1 statistically significant length scale for the two different fits, for both fast and slow wind, compared to the results of Viall et al. (2008)

We show in Figure 7 the normalized occurrence distributions of statistically sig-354 nificant radial length scales for slow and fast wind, and for both the AR(1) and BPL back-355 ground assumptions, for all 25 years of Wind data. We computed the histograms in 3-356 year intervals, shifting by 1-year for each new histogram. We mark the occurrence en-357 hancements (i.e. the persistent length scales) that are $> 2\sigma$ above the occurrence dis-358 tribution with thick lines in Figure 7. For example, the histograms for 2017-2019 slow 359 wind in Figure 7 show in the BPL histogram 3 clear peaks below 100 Mm, and two broad 360 peaks near 130 and 160 Mm. The histogram derived from the AR(1) assumption show 361 the first 2 peaks below 100 Mm and the two broad peaks near 130 and 160 Mm, but at 362 a reduced relative amplitude compared to the BPL histogram. 363

To compare between the two background assumptions, we plot the significant length 364 scales identified in both the AR(1) and BPL derived occurrence distributions as signif-365 icant at the 2σ level as horizontal bars in Figure 8. Lengths that were identified concur-366 rently in the occurrence distributions of both model fits are shown in Figure 8 as solid 367 black bars; these same lengths are highlighted in the individual panels as darker shades 368 of red and blue. Many of the occurrence distributions exhibit local enhancements at the 369 smallest length scales, very near the Nyquist, and we shade those particular length scales 370 lighter to emphasize they may not be statistically significant. Excluding these length scales 371 below 80 Mm, we find for slow wind that 67% of the BPL lengths are contained in the 372

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Figure 7. Bootstrapped occurrence distributions of length scales identified as significant by passing both the amplitude and *F*-test at the 95% level, for both slow (left) and fast (right) wind, and both background assumptions. Local peaks that exceed the background by 2σ are considered significant and are marked with thick lines.

AR(1) distributions, and 54% of the AR(1) are in BPL. For fast wind, excluding length 373 scales < 130 Mm, these are 79% and 88% for both BPL and AR(1), respectively. The 374 primary differences are at the ends of the spectral range analyzed, and follow the gen-375 eral pattern identified in the example shown in Figure 4. At the long length-scale end 376 (low wavenumber), fewer significant peaks were identified with the BPL (blue) background 377 assumption, while at the short-length-scale end (high wavenumber), fewer peaks were 378 identified with the AR(1) (red) background assumption. Another difference occurs in 379 the early part of the mission, with AR(1) finding a band near 150 Mm in the slow wind 380 that is not apparent in the BPL results. 381

The new results are consistent with the previous results of Viall et al. (2008) that 382 covered the years 1995-2005 using the original Wind data. Figure 9 shows the concur-383 rently identified significant length scales from Figure 8 with the AR(1) derived results 384 from Viall et al. (2008). For the slow wind (Figure 9a), both studies identified signifi-385 cant lengths near 130 and 170 Mm, and an additional set near 330 Mm. The differences 386 between the original and reprocessed data occur primarily in the first 3 rows, covering 387 years 1995-2001, during the earliest portion of the Wind mission. The fast wind results 388 compare very well to the previous Viall et al. (2008) results, with 3 sets of length scales 389

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Figure 8. Bars represent statistically significant length scales identified in the occurrence distributions of Figure 7 as above the background at the 2σ level. Length scales at edge of the OD that may be affected by the Nyquist are lightly shaded. Black bars represent length scales identified simultaneously in both BPL and AR(1) distributions. We also list the equivalent frequencies using the median solar wind speed of 420 km/s for slow, and 675 for fast. For Earth's magnetosphere, or an in situ spacecraft, these length scales would appear as periodicities at these frequencies. The sunspot number cycle is shown in the middle for reference.

near 100, 300, and 400-500 Mm detected in both the original and reprocessed Wind data.
The slight shift to shorter length scales in the 80-500 Mm bands in the reprocessed data
results is due to a reduced central peak in the OD in the reprocessed data compared to
the original data.

³⁹⁴ 4 Discussion

The histograms shown in Figure 7 represent occurrence distributions of significant length scales observed in the solar wind near L1 over two solar cycles. The overall shape of these distributions exhibits a consistent pattern across the full 25 years of Wind data (Figure 7). For the slow wind, the statistically significant length scales identified using the BPL background assumption exhibit comparatively few counts at the longer length scales (> 300 Mm), and a broad peak near the center of the distribution (100-200 Mm). The AR(1)-derived histograms exhibit a steep slope at the smaller length scales, followed



Figure 9. Comparison between the statistically significant length scales identified by Viall et al. (2008) (green), using the original Wind SWE data, and this study (black), using the reprocessed Wind data, for the AR(1) background assumption.

⁴⁰² by a slow decline at the longer length scales. The histograms for the fast wind length ⁴⁰³ scales show a similar, although less pronounced, trend. Recall that at long wavenumber, ⁴⁰⁴ the bin width $(6f_R)$ becomes comparable to the length scales. Future work examining ⁴⁰⁵ longer data segments is required to understand the nature of the shape of the occurrence ⁴⁰⁶ distribution over these longer (>~ 500 Mm) length scales.

In addition to these overall trends, the occurrence distributions exhibit local en-407 hancements of length scales identified as significant. These are highlighted in the occur-408 rence distribution histograms in Figure 7 and pulled out separately in Figure 8 as bars. 409 Figures 7 and 8 together provide evidence for persistent bands of significant periodic length 410 scales. To highlight these trends we have plotted colored contour plots, along with the 411 normalized residuals from which these length scales were determined, in Figure 10a and 412 10b. The residuals here are the addition of the normalized occurrence distribution resid-413 uals from the BPL and AR(1) background assumptions. The plotted values are $(OD_{BPL} -$ 414 fit_{BPL}) + $(OD_{AR(1)} - fit_{AR(1)})$, where OD is the 3-year occurrence distribution and 415 fit is the occurrence distribution fit for the two spectral background assumptions. Length 416 scale occurrence enhancements that were detected in both occurrence distribution resid-417 uals would add together (red), while parts of the distributions that are less correlated 418 would tend to zero (green). 419

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Figure 10a and 10b shows clear patterns of periodic length scales that evolve over 420 the full 25 years of Wind SWE data. In the slow wind, $L \sim 90$ Mm (VI), $L \sim 130 -$ 421 140 Mm (III), and $L \sim 170-190$ Mm (II), are all observed for the majority of the 25-422 year dataset, with some noticeable variations we discuss below. There are two smaller 423 bands near $L \sim 210$ Mm (IV) in the middle years and between 310 and 350 Mm in the 424 later years, and a sloped band between 250 and 400 Mm (I) for the first half of the in-425 terval. An additional band appears near $L \sim 120$ Mm in the BPL-derived histograms 426 in Figure 7a, but is not apparent in the AR(1)-derived histograms, likely because this 427 region has a very strong slope; there is a similar effect with the $L \sim 90$ Mm band (see 428 Figure 7). For the fast wind there is an intermittent band between $L \sim 200-220$ Mm 429 (IV), and two bands (I and V) that are highly sloped in time, suggesting a solar cycle 430 dependence. Band I decreases from 500 Mm to 300 Mm over solar cycle 23, while Band 431 V appears at 400 Mm near the start of solar cycle 24. These bands also appear in the 432 "slow" wind results. 433

Figure 10c shows a pictorial summary of the significant length scale bands, derived 434 by examining the combined bar plots and residual contours of Figure 10a and 10b, and 435 using the additional information of the histograms in Figure 7 to provide visual guid-436 ance on persistence. Recall that the separation between fast and slow wind was math-437 ematical, for the purposes of an even sampling rate, rather than the physics of the for-438 mation of solar wind of different speeds. For this reason, we have combined the signif-439 icant length scales observed in the slow and fast wind together. We note that bands I, 440 IV and V are observed in both fast and slow wind analysis, suggesting that the creation 441 mechanism of periodic density structures is not strictly a "slow" (< 550 km/s) wind phe-442 nomenon. 443

Many characteristics of the Sun, solar corona, and solar wind are correlated with 444 solar cycle, so unraveling the specific nature of the correlation of periodic density struc-445 tures with solar cycle is a topic for future work. Here we speculate on a likely connec-446 tion. In general, the solar corona is hotter, and its magnetic topology increases in com-447 plexity, at solar maximum, as manifested in active regions and their underlying magnetic 448 concentrations, sunspots. To the right of Figures 10a-c, we show the gradual solar cy-449 cle change as measured by 3-year averages of the sunspot number, along with the more 450 abrupt "terminator" events that are the end of a Hale magnetic cycle (McIntosh et al., 451 2015, 2019). The terminator events are observed as abrupt changes in the distribution 452

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of solar EUV bright points, and occur when there is no more old cycle polarity flux left
on the solar disk. Related, Schonfeld et al. (2017) showed that the amount of hot plasma
(plasma greater than 10^{6.1}K in the solar corona) abruptly increases at the terminator,
due to an increased amount of hot plasma in active regions.

The length scale bands that we find in this paper exhibit breaks that are associ-457 ated more closely with terminators than with sunspot minimum. For example, bands I 458 and II at both ends, and band III for the termination event of solar cycle 22. Addition-459 ally, there is a gradual evolution of the characteristic length scales between termination 460 events, most pronounced in bands I and V. With data from only two, very different, so-461 lar cycles, we cannot draw definitive conclusions about the exact relationship between 462 solar wind periodic length scales and the solar cycle, but the result suggests that a re-463 lationship exists. The precise details of this relationship would likely become more clear 464 with the next solar cycle. 465

As reviewed in Section 1, there is strong evidence that periodic density structures 466 originate from the sun and are associated with magnetic reconnection of plasma from 467 closed-field regions. The evolution of periodic length scales with solar cycle could be the 468 result of changes in the nature of the interchange reconnection that releases the plasma 469 into the solar wind, due to the increase in complexity of the global magnetic topology 470 (Antiochos et al., 2011). The association with the termination event could be the result 471 of the reversal of the polarity of the leading edge of the new active regions. When the 472 leading sunspot has the opposite polarity of the surrounding coronal hole, null point topolo-473 gies can form out from decaying active regions (e.g., Mason et al. (2019)). This magnetic 474 topology is expected to have different interchange reconnection properties than when the 475 active region has a leading polarity that follows Hale's law. Alternatively, coronal tem-476 perature is correlated with solar wind speed, so it could also be that the hotter active 477 regions that occur after the terminator event accelerate solar wind, and any embedded 478 periodic density structures, differently. 479

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While this study focused specifically on mesoscale structures measured at L1 that exhibit periodicity in density, many other studies have observed mesoscale structures in the solar wind that form at the sun and advect to 1 AU. A general picture is emerging in which mesoscale structures that form through spatial structures that rotate (Borovsky, 2008, 2020) or time dynamics such as reconnection in the corona (Sanchez-Diaz et al.,

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⁴⁸⁵ 2016, 2017, 2019; M. J. Owens et al., 2018; M. Owens et al., 2020; Stansby & Horbury,

⁴⁸⁶ 2018), are an inherent part of solar wind formation (Viall & Borovsky, 2020).

In a series of papers, Rouillard, Davies, et al. (2010), Rouillard, Lavraud, et al. (2010), 487 and Rouillard et al. (2011) tracked larger mesoscale structures from their formation in 488 the corona through the inner heliosphere using SECCHI HI images, all the way to their 489 impact at the Earth. They identified the corresponding compositional and magnetic field 490 variations inherent to the structures, which were retained out to 1 AU. This set of stud-491 ies unequivocally demonstrated that large mesoscale structures created at the Sun sur-492 vive to 1 AU with identifiable in situ signatures. More recently, Rouillard et al. (2020) 493 tracked density structures through the STEREO COR2 and HI1 FOVs to their impact 494 at Parker Solar Probe, where they observed a one-to-one correlation between the $\sim 3-$ 495 4 hour density structures observed remotely and the in situ Parker measurements. They 496 showed that Parker measured additional sequences of small density peaks separated in 497 time by approximately 90-120 minutes, suggestive of the types of periodic density en-498 hancements at 90 minute timescales that have been observed in situ at L1 (Viall et al., 499 2008; Kepko & Spence, 2003), near Mercury's orbit with Helios (Matteo et al., 2019) and 500 remotely with STEREO (Viall & Vourlidas, 2015). Many of these event studies exhib-501 ited still smaller substructures at tens of minutes (Matteo et al., 2019; Kepko & Viall, 502 2019; Kepko & Spence, 2003), in the range of the structures studied here. Several stud-503 ies also found composition signatures which could only have come from formation at the 504 sun (Viall, Spence, & Kasper, 2009; Kepko & Viall, 2019). Recent work by Réville et al. 505 (2020) demonstrated that periodic density structures associated with helmet streamers 506 could be the result of the tearing mode instability at the base of the heliospheric cur-507 rent sheet. They argue that the larger, $\sim 10-20$ hour periodicities, as well as the $\sim 1-2$ hour 508 periodicities that are observed are all the result of the tearing mode. Finally, Murphy 509 et al. (2020) demonstrated a distribution of mesoscale solar wind flux ropes observed at 510 Mercury, with time scales of 2.5 minutes to 4 hours. They concluded that a portion of 511 the distribution was likely related to PDS generation. These studies together demonstrate 512 that the solar wind is often composed of mesoscale density structures, and provide am-513 ple evidence that structures of order tens of minutes timescales and longer form with the 514 515 solar wind and survive through the inner heliosphere, out to 1 AU.

Finally, this current study, which focused on the smaller end of the mesoscale range, demonstrated that at least some mesoscale structures are quasi periodic, and occur at

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repeatable sets of frequencies and/or length scales. We emphasize that these length scales 518 represent PDSs that advect with the solar wind. In the rest frame of a spacecraft or planet, 519 they would appear as a periodic density variations at a frequency determined by $f_{PDS} =$ 520 V_{sw}/L_{PDS} . Statistically, for any particular year the magnetosphere or a spacecraft would 521 see a spectrum of equivalent frequencies determined by convolving the distribution of 522 solar wind V_x with the length scales identified in Figures 8 and 10c from that year. To 523 zeroth order, we can estimate these frequencies using the median solar wind speed for 524 "fast" and "slow" solar wind. These equivalent frequencies are listed at the bottom of 525 Figure 10c. The equivalent frequencies of these structures fall in the few mHz range, which 526 for the magnetosphere is considered the Pc5-6 band. Previously, Viall, Kepko, and Spence 527 (2009) studied 11 years of Wind SWE data covering 1995-2005 for evidence of discrete 528 frequency periodicities in the solar wind number density. They found that f = 0.7, 1.3-529 1.5, 2.0-2.3, and 4.7-4.8 mHz occurred most often over that 11-year interval. Fig-530 ure 10c demonstrates that f = 1.4 mHz corresponds to Band I in the slow wind, f =531 2.0-2.3 mHz corresponds to Band IV in the slow and I in the fast, and f = 4.7-4.8532 mHz corresponds to Band VI in the slow wind. 533

Since these are periodic structures in solar wind density, they would periodically 534 compress the magnetosphere via periodic dynamic pressure changes, and we would ex-535 pect the magnetosphere to show these same sets of frequencies. In the same Viall, Kepko, 536 and Spence (2009) study, they also examined GOES magnetospheric magnetic field data 537 for intervals when GOES was near the dayside magnetopause, and found in the GOES 538 data a similar set of frequencies to those found in the solar wind. In a direct compar-539 ison between Wind and GOES, they found when a spectral peak was observed in the so-540 lar wind, that same peak was observed at GOES 54% of the time. Other statistical stud-541 ies have similarly identified persistent bands of significant mHz frequencies (e.g., Francia 542 and Villante (1997); Chisham and Orr (1997); Ziesolleck and McDiarmid (1995)). While 543 originally attributed to global cavity modes (e.g., Harrold and Samson (1992)), we now 544 know these $< \sim 4$ mHz oscillations are largely driven by solar wind periodic density struc-545 tures. Since these periodic length scales directly drive the magnetosphere, we would ex-546 pect the spectrum of discrete mHz oscillations in the magnetosphere to vary year-to-year 547 548 as the L_{PDS} vary. Since the L_{PDS} have a solar cycle dependence, this would mean the spectrum of discrete mHz waves in the magnetosphere would also have a solar cycle de-549 pendence, although the variability of the solar wind speed would produce broad, rather 550

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than narrow, enhancements. This slow year-to-year variability, and the distribution of

solar wind speeds, can explain year-to-year changes in measured frequencies. In addi-

tion, Kepko and Viall (2019), showed that ambient periodic density structures in the slow

solar wind were sometimes compressed and amplified by a faster solar wind stream from

behind, and that these amplified PDSs had an observable impact on radiation belt par-

ticles. These particular PDSs were observed with stream interaction regions, which are

⁵⁵⁷ known to be important drivers of radiation belt flux enhancements.

558 5 Conclusions

This study provides further evidence that large portions of the solar wind plasma 559 consist of periodic mesoscale structures, many of which are likely released via magnetic 560 reconnection. Using 25 years of Wind solar wind number density data observed near L1 561 we have identified bands of periodic length scales that occur more often than others. In 562 the rest frame of a spacecraft or Earth, these periodic mesoscale density structures would 563 appear at frequencies determined by the length scales of the PDSs and the solar wind 564 velocity. Each occurrence of a periodic length scale passed two independent spectral tests 565 at the 95% level, and we tested each occurrence with two different background spectral 566 models. We identify bands of occurrence enhancements that are persistent in time, and 567 are significant using both background spectral models (Figure 8c). Bands near $L \sim 130-$ 568 140 Mm and $L \sim 170-190$ Mm were evident in the slow wind, equivalent to frequen-569 cies of $f \sim 3.0$ and 2.3 mHz in the stationary frame, while bands near 230 and 300 Mm 570 were observed in both the fast and slow segments, equivalent to $f \sim 1.9$ and 1.2 mHz, 571 and $f \sim 3.1$ and 2.0 mHz, for the slow and fast wind, respectively. Longer length bands 572 were observed between 300 and 500 Mm, decreasing in length over the course of solar 573 cycles 22 and 23. The apparent frequencies of these lengths fall in the Pc5-6 pulsation 574 bands, which are known to be important for processes leading to radiation belt parti-575 cle loss, diffusion, and acceleration (Elkington & Sarris, 2016). The evolution of these 576 bands exhibited changes near solar "terminator" events marking the end of a Hale mag-577 netic cyle (Figure 10), although this is a qualitative association and requires further work. 578 Given the statistical bands of recurrent length scales in the solar wind, it may be pos-579 sible in the future to produce a statistical model for these solar-wind driven discrete os-580 cillations. 581

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Finally, while our study separated "slow" and "fast" wind based on speed, this was 582 driven by the mathematics of creating length series with a fixed sampling length. With-583 out separating the speed in this manner, length segments corresponding to fast speed 584 would have been undersampled, and slow wind segments would have been oversampled. 585 Therefore, this approach is not suited for, nor designed for, determining how the bands 586 relate to formation of different types of solar wind, nor can it determine whether differ-587 ent physical mechanisms create different bands of periodicities. Indeed, Figure 10 shows 588 that some bands in the 'fast' and 'slow' wind overlap, indicating a common mechanism 589 for those bands, independent of final wind speed. Future work includes combining our 590 event list of periodicities identified over 25 years of Wind data with clustering analysis 591 (e.g., Roberts et al. (2020); Ko et al. (2018)) and studies of compositional changes. This 592 work is already underway. 593

594 Acknowledgments

L. Kepko and N. Viall acknowledge the Heliophysics Internal Scientist Funding Model and the NASA Guest Investigator program for funding. K. Wolfinger was supported through the NASA internship program. All data used in this study were obtained from CDAWeb https://cdaweb.sci.gsfc.nasa.gov/. We thank M. Stevens for helpful discussion regarding the Wind Faraday Cup (SWE) data.

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Figure 10. The contour plots (a) and (b) are the addition of the normalized (to the peak value) residuals for both the BPL and AR(1) derived occurrence distributions for the slow and fast wind. Red indicates areas of enhancement observed in both OD residuals, blue indicates areas where both found length scales significantly below the background fit, and green indicates regions near the background or areas where BPL and AR(1) were in disagreement. The bars superimposed on (a) and (b) are from Figure 8(e) and (f), and indicate length scales that exceeded the background by 2σ . The schematic (c) is a pictorial representation of (a) and (b) combined, and includes the 3-year running average of sunspot number, and locations of the terminator