Environmental modulations of nutrient conditions in the Labrador Sea reconstructed from nitrogen isotopes in a six-hundred-year-old crustose coralline alga

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Abstract

The climatological impacts on biogeochemical processes in the polar North Atlantic remain poorly understood, as there exist both biological and physical mechanisms that drive nutrient availability in the region. Here, we present nitrogen isotope measurements (δ N) from a six-hundred-year-old coralline alga to elucidate historic and modern trends in Labrador Sea nitrate utilization, defined as the degree of biological nitrate assimilation relative to nitrate supply. Prior to the Little Ice Age (LIA), periods during which utilization became complete corresponded to neutral modes of the Atlantic Multidecadal Oscillation (AMO), which we argue promoted the oceanographic conditions favorable for simultaneous phytoplankton growth and reduced nitrate input. More recently, nitrate utilization became complete during periods characterized by reduced deep-water convection in the Labrador Sea, suggesting a reduced inflow of equatorially-sourced nitrate driven by a weakening of the Labrador Current. Such nutrient rerouting may have implications for socioeconomically-important fisheries and carbon sequestration throughout the region.

1	Climate-modulated nutrient conditions along the Labrador Shelf: Evidence
2	from nitrogen isotopes in a six-hundred-year-old crustose coralline alga
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16	Key Points:
17	• Nitrogen isotopes from a crustose coralline alga are argued to record ocean circulation
18	and nutrient utilization along the Labrador Shelf
19	• Past periods of increased polar inflow waters and nutrient utilization are linked to
20	negative modes of the Atlantic multidecadal oscillation
21	• An anomalously long phase of low nutrient input since ~1870 is linked to the weakening
22	of the Atlantic meridional overturning circulation
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Abstract

The impacts of climate change on north Atlantic nutrient chemistry remain poorly understood, as 25 there exist a multitude of rapidly changing biological and physical drivers of nutrient conditions 26 throughout the region. Here, we present nitrogen isotope measurements derived from a six-27 hundred-year-old crustose coralline alga ($\delta^{15}N_{algal}$) to elucidate historical and contemporary trends 28 in nitrate utilization and circulation patterns along the Labrador Shelf. Prior to the early 1900s, we 29 argue that intervals during which utilization approached completion were controlled by reduced 30 31 nitrate advection linked to an increased proportion of nitrate-poor polar waters and subdued Atlantic influence, as expected from concurrent negative modes of the Atlantic multidecadal 32 oscillation. While nitrate conditions should have recovered in recent years, our record suggests 33 that high utilization persisted since ~1870, which we also attribute to reduced Atlantic advection, 34 likely associated with the twentieth-century anthropogenic weakening of the Atlantic meridional 35 36 overturning circulation. These results highlight the role of ongoing climate-induced circulation changes in modulating nutrient distributions throughout the subpolar north Atlantic, which may 37 38 have implications for other environmental phenomena such as fisheries and oceanic carbon storage. 39

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49 **1 Introduction**

High-latitude regions suffer the greatest impacts of modern-day climate change (Pithan & 50 Mauritsen, 2014). One effect of this change is enhanced sea-ice melt along the Labrador Shelf 51 (Halfar et al., 2013), which is driving an unprecedented increase in primary productivity relative 52 to the past several hundred years (Chan et al., 2017). While such a dramatic change in productivity 53 should intuitively lead to the greater depletion of upper-ocean nutrients, multiple simultaneous 54 environmental changes throughout the region make this relationship complex. This is because 55 nutrient availability in the subpolar north Atlantic is not only a function of biological activity, but 56 57 also of physical oceanographic processes that regulate nutrient supply.

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The Labrador Sea in the subpolar north Atlantic receives considerable scientific attention because 59 deep-water convection in the region is thought to be historically important for driving the strength 60 61 of the Atlantic meridional overturning circulation (AMOC) (Buckley & Marshall, 2016; Kuhlbrodt et al., 2007). Paleoceanographic investigations suggest that convection in the Labrador Sea has 62 63 been at an anomalously weak state over the last ~ 150 years, most likely due to anthropogenic warming and freshwater forcing (Caesar, Rahmstorf, Robinson, Feulner, & Saba, 2018; Thibodeau 64 et al., 2018; Thornalley et al., 2018). Recent observations further suggest that convection in the 65 Labrador Sea has been virtually shut off (Lozier et al., 2019), which has weakened the influence 66 67 of nitrate-rich Atlantic inflow along north-western Atlantic coast (Thibodeau et al., 2018; Thibodeau, De Vernal, Hillaire-Marcel, & Mucci, 2010). Such a reorganization of upper-ocean 68 circulation patterns may therefore also have implications for primary productivity, regionally 69 important fisheries (Stock et al., 2017) and oceanic carbon storage (Takahashi et al., 2009). Due 70 to the absence of long-term observational data, however, paleo-proxies are required to 71 contextualize these contemporary changes and their effects on nitrate delivery. 72

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Along the coastal Labrador Shelf, nitrate is supplied during winter via vertical mixing (Harrison et al., 2013; Harrison & Li, 2007; Henson, Dunne, & Sarmiento, 2009) and advection from the eastern subpolar region (Loder, Petrie, & Gawarkiewicz, 1998). Advective supply relies on the Labrador Current, which sources waters from the eastern region of the subpolar Atlantic, Baffin Bay and Hudson Strait (Figure 1a and 1b). Eastern subpolar waters are comprised of Arctic outflow, sourced from the East Greenland Current, and Atlantic inflow, sourced from the North Atlantic Current, an extension of the Gulf Stream. However, previous field surveys indicate that the nitrate derived from the Hudson Strait only impacts the northern-most tip of the shelf (Drinkwater & Harding, 2001). As such, nitrate supply along the mid-shelf region is primarily affected by vertical mixing and open-ocean advection, which today is mostly completely consumed in the upper ~20 m of the water column during the phytoplankton growing season (Figure 1c and 1d).

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87 The degree of nitrate consumption relative to supply, known as utilization (Altabet and Francois, 1994), imprints an isotopic fingerprint on biologically assimilated nitrogen. This is due to kinetic 88 fractionation processes that result in the preferential assimilation of ¹⁴N-nitrate in virtually all 89 photosynthetic organisms that have been studied (e.g., Altabet & Francois, 1994; Mariotti et al., 90 1981), including macro-benthic algae (Swart, Evans, Capo, & Altabet, 2014). During the 91 contemporary phytoplankton growing season along the Labrador Shelf, utilization approaches 92 completion and the isotopic composition of biologically assimilated nitrogen is therefore expected 93 to approximate that of the initial source (e.g., Altabet & Francois, 1994; Mariotti et al., 1981). This 94 95 utilization-induced fractionation makes the isotopic composition of geologically preserved nitrogen useful for reconstructing nutrient utilization in a variety of past marine settings (Altabet 96 97 and Francois, 1994), including the subpolar north Atlantic (Straub et al., 2013).

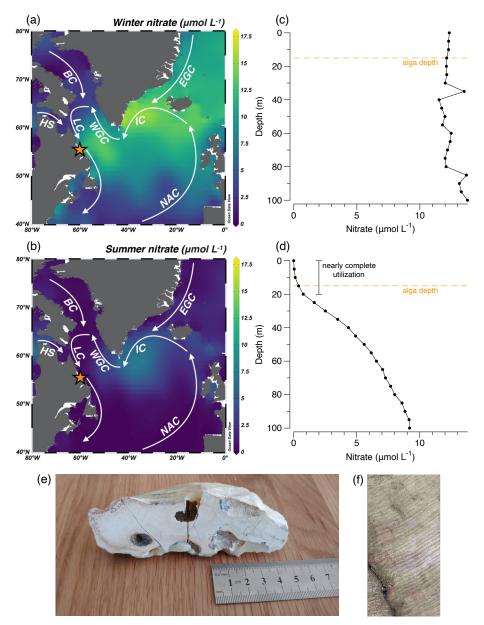
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Fig. 1. Oceanographic setting and C. compactum specimen. (a, b) Surface maps of (a) winter 105 and (b) summer nitrate distributions with a depiction of major north Atlantic circulation patterns. 106 The study site is indicated by the orange star. Abbreviations: NAC: North Atlantic Current, IC: 107 Irminger Current, EGC: East Greenland Current, WGC: West Greenland Current, LC: Labrador 108 Current, BC: Baffin Current, HS: Hudson Strait outflow. These figures were generated with the 109 help of Ocean Data View (Schlitzer, 2018) using inorganic nutrient data taken from the World 110 Ocean Atlas (Garcia et al., 2013). (c, d) Depth profiles of (c) winter and (d) summer nitrate 111 concentrations at 55.5 °N, 58.5 °W. (e) Cross-section of the Clathromorphum compactum 112 specimen sampled in this study and (f) its growth bands. 113

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115 Nitrogen is preserved in several geologically important marine organisms, such as foraminifera

116 (Ren et al., 2009), corals (Muscatine et al., 2005; Sherwood, Lehmann, Schubert, Scott, &

McCarthy, 2011; Wang et al., 2014) and bivalves (Gillikin et al., 2017). Over the last few decades, 117 crustose coralline algae have also emerged as promising paleoenvironmental archives due to their 118 119 widespread distribution, significant longevity and annual banding patterns (Figure 1e and 1f), the latter of which allows for precise and high-resolution reconstructions of oceanographic changes in 120 the recent geological past (e.g., Chan et al., 2017; Halfar et al., 2013; Moore et al., 2017). As such, 121 the $\delta^{15}N$ of nitrogen retained in the organic-rich skeletons of coralline algae may offer a unique 122 window into the history of marine nutrient dynamics on multicentennial timescales. Here, we 123 present the first ~5-year-resolved δ^{15} N record derived from a 613-year-old crustose coralline alga, 124 *Clathromorphum compactum* ($\delta^{15}N_{algal}$), and demonstrate its utility in reconstructing nitrate 125 conditions along the rapidly changing Labrador Shelf throughout the last several centuries. 126 Because a variety of factors may impact the isotopic composition of biologically assimilated 127 nitrogen, we analyze our record in the context of several other high-resolution paleoenvironmental 128 reconstructions to infer the most likely drivers of $\delta^{15}N_{algal}$ variations over the last several hundred 129 130 years.

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132 2 Materials and Methods

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2.1 Specimen Collection and Age Model Development

A living C. compactum specimen was collected off the coast of Kingitok Island, Labrador, Canada 135 (55.3983 °N; 59.8467 °W) at 15 m depth in 2011 via divers using SCUBA (Figure 1a-f). Following 136 collection, the specimen was rinsed in freshwater and cross-sectioned perpendicular to the 137 direction of growth. The cross-section was mounted on a plate using wax and polished with Allied 138 High Tech diamond lapping film and water (30, 15, and 1 µm grains) until growth bands were 139 140 clearly visible. Using high-resolution images of the specimen taken with a Nikon H600L Microscope (Figure 1f), bands were identified and assigned years to develop a growth chronology 141 in Photoshop. Previous U-Th dating indicated that the age of the specimen was ~ 630 years old 142 (Halfar et al., 2013), which was also verified by the number of growth bands counted. 143

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145 **2.2 Isotopic Analysis**

146 A high-precision, computer-driven New Wave Research Micromill Sampling System attached to 147 an x, y, and z stage was used to collect (mill) material from along the growth bands for $\delta^{15}N$

analysis. To remove external particulates before drilling, the specimen was sonicated 3 times for 148 5 minutes in Milli-Q water and then oven-dried for 24 hours. To verify the effectiveness of this 149 Milli-Q cleaning technique, $\delta^{15}N$ values and %N were measured in replicate C. compactum 150 samples exposed to our method and an oxidative cleaning technique designed for carbonate 151 samples of much lower nitrogen content (Ren et al., 2009). While %N was consistently lower 152 under oxidative cleaning, δ^{15} N values obtained from samples under both treatments remained 153 within one standard deviation of each other (Table 1). These results may indicate that the additional 154 155 nitrogen removed during oxidative cleaning was mostly derived from algal-fixed nitrogen rather than external contamination, as the removal of external contamination would be expected to also 156 alter the δ^{15} N. 157

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Region	Sample Type	Treatment	Replicates	n	$\delta^{15}N$	δ ¹⁵ N 1σ	%N	%Ν 1σ
	C. compactum	Oxidative	1	5	6.04	0.36	0.14	0.03
Qikiqtarjuaq (Baffin Bay)	C. compactum	Milli-Q	2	10	6.54	0.37	0.27	< 0.01
	Seawater NO ₃ (\geq 100 m)	na	na	3	6.51	0.51	na	na
Labrador Shelf	C. compactum	Milli-Q	1	1	5.9	na	na	na
Labrador Shell	Seawater NO ₃ (100 m)	na	3	na	6	0.28	na	na

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Table 1. Nitrogen isotope measurements from algal samples and seawater nitrate in two 161 regions of nearly complete nitrate utilization. Nitrogen isotope measurements and 162 concentrations were derived from Qikiqtarjuag replicates under Milli-Q and oxidative (Ren et al., 163 2009) cleaning treatments. Experiments were performed on powder extracted from the full cross-164 section of the same skeletal chunk. For samples cleaned only with Milli-Q, two replicate powders 165 (n = 5 for each powder) were used. Nitrogen isotope measurements of seawater nitrate for the 166 Oikigtarjuag (Baffin Bay) and Labrador Shelf regions are taken from Lehmann et al., (2019) and 167 Sherwood et al., (2011) respectively. Because both regions are subject to nearly complete nitrate 168 consumption, the agreement between the isotopic composition of seawater nitrate and algal 169 specimens supports the argument that C. compactum is likely recording marine nitrate utilization 170 (see Section 3.1 for details). 171

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Following the visible growth bands as much as possible, digitized drill paths were programmed on the Micromill screen. For each drill line, 5 to 6 150-µm-deep drill passes were made to obtain 7 mg of skeletal material. Samples were transferred and weighed in tin cups along with reference materials. The majority of samples were taken at 5-year increments with an average resolution of 3 to 6 years, except for regions of the skeleton where insufficient material required milling from more growth bands, increasing the resolution to between 7 to 20 years. Of the total 110 measurements reported here, 16 correspond to a time window of >7 years. Following the collection

of each sample, pressurized air was used to remove residual sample powder. Then, the samples 180 were analyzed using a Micromass Isoprime 100 isotope ratio mass spectrometer coupled to an 181 Elementar Vario MicroCube elemental analyzer operated in continuous-flow mode at GEOTOP 182 (Université du Québec à Montréal). Two internal reference materials (Leucine; $\delta^{15}N = -0.10 \pm$ 183 0.24‰ and DORM 2; $\delta^{15}N = +14.95 \pm 0.09$ ‰) were used to normalize the results to the AIR scale 184 based on international IAEA standards (N1, N2 and NO3). We measured a third internal standard 185 (casein; $\delta^{15}N = -0.1 \pm 0.15\%$) to assess the normalization. Results are given in delta units (δ) in 186 % vs AIR. Overall analytical uncertainty (1 σ) was calculated to be better than \pm 0.2‰, based on 187 the propagation of uncertainties of the normalization of internal reference materials and samples. 188 Chromatographs produced from replication tests suggest that the algal carbonate underwent 189 complete combustion. Specifically, an anomalous m/z 30 peak was not observed during any of the 190 analyses, which would be expected from the production of ${}^{12}C^{18}O$ (~2% of CO isotopologues) in 191 the case of incomplete combustion. In addition, N2 peak tailing was not noticeably different 192 between algal samples and organic standards, further assuring successful combustion. 193

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2.3 Spectral Analysis

Cross-wavelet coherence and phase-relationship analyses were conducted using the Cross Wavelet 196 and Wavelet Coherence toolbox in MatLab R2020a provided by A. Grinsted. For technical 197 specifications of squared-wavelet coherence and phase-direction calculations, see Grinsted et al., 198 199 2004. Significance is reported at the 95% confidence level against red noise, which was determined via 300 Monte Carlo simulations. "In-phase" relationships describe those that represent positive 200 correlations between variables and "antiphase" relationships describe negative correlations. To 201 achieve equally spaced δ^{15} N_{algal} values, five-year interpolations were calculated from the raw data 202 203 and used for spectral analysis.

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3 Results and Discussion

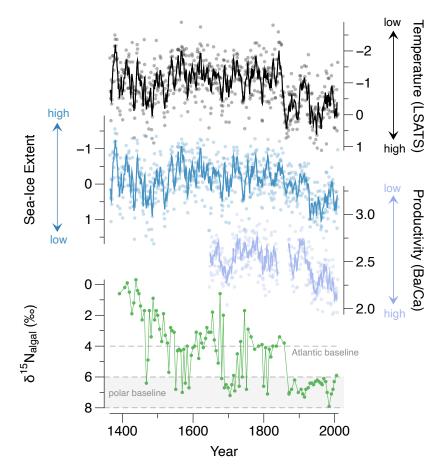
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3.1 Potential Drivers of $\delta^{15}N_{algal}$

Our record spans from 1392 to 2005. During this period, $\delta^{15}N_{algal}$ values generally followed an increasing trend from -0.3 to 7.9‰ (Figure 2). $\delta^{15}N_{algal}$ initially increased by ~7‰ from 1392 to the late 1500s prior to a transition to relatively constant average values, during which there were 4 notable periods characterized by values typically reaching maximum values of ~6 to 7‰ spaced by approximately 100 years. Beginning around 1870, $\delta^{15}N_{algal}$ increased by an average of ~2.9‰,

with high values persisting throughout the remainder of the study period. Here, we propose that

this algal record has documented changes in seawater nitrate utilization over the past six centuries.



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Fig. 2. Environmental changes in the Labrador Shelf archived in coralline algae. (top to 215 bottom) Temperature reconstructed from the Labrador Sea Algae Time Series (LSATS) reported 216 in Moore et al. (2017); historical algal-derived index of regional sea-ice extent taken from Halfar 217 et al. (2013); algal Ba/Ca ([($\mu g g^{-1}$)/($\mu g g^{-1}$) × 10⁻⁵]), negatively correlated with primary 218 productivity (Chan et al., 2017); and our coralline algal nitrogen isotope record (analytical 219 uncertainty <0.2‰). Lines for temperature anomalies, sea-ice extent and Ba/Ca represent five-220 year-smoothed means of the data. The approximate Atlantic (4‰) and polar (Baffin Bay; 6 to 8‰) 221 seawater nitrate δ^{15} N baselines are marked on the bottom panel. Phases of nearly complete 222 utilization could be consistent with $\delta^{15}N_{algal}$ values falling within this range regardless of the 223 relative proportion of Atlantic versus polar waters feeding the Labrador Current. However, we 224 225 conservatively consider only the upper-bound (polar) isotopic range as coherent with nearly complete utilization (see Section 3.1 for details). 226

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228 Recent culture experiments indicate that, like many marine autotrophs, coralline algae assimilate

229 nitrate during their growth (Hanson, Slaymark, Kamenos, & Thibodeau, 2020). Further, our most

recent $\delta^{15}N_{algal}$ measurements match the isotopic signature of ambient seawater nitrate along the 230 Labrador Shelf (Sherwood et al., 2011) (Table 1). Throughout the shelf and larger Labrador Sea 231 region, nitrate is the most critical nutrient for regulating primary productivity (Harrison & Li, 232 2007), and so the observed agreement between $\delta^{15}N_{algal}$ and the $\delta^{15}N$ of the local nitrate pool is 233 expected if driven by the degree of nitrate utilization (e.g., Altabet & Francois, 1994; Mariotti et 234 al., 1981). This agreement is also confirmed by measurements of $\delta^{15}N_{algal}$ and nearby seawater 235 nitrate at higher latitudes (Lehmann et al., 2019) (Table 1), further verifying that $\delta^{15}N_{algal}$ likely 236 tracks the isotopic composition of seawater nitrate. However, several other biological and 237 biogeochemical processes may impact the $\delta^{15}N$ of seawater nitrate and therefore might also affect 238 our record. 239

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At coastal locations such as our study site, terrestrial nutrient runoff could impact the degree of nutrient utilization and/or the baseline isotopic composition of seawater nitrogen. However, annual runoff is only 600 - 700 mm along the shelf (Rollings, 1997), which is many orders of magnitude diluted by the Labrador Current, traveling at a coastal velocity of 0.8 Sv (Lazier & Wright, 2002). While runoff could play a role in the coastal nutrient budget, its relative influence suggests a negligible contribution and is therefore not likely to account for large changes in $\delta^{15}N_{algal}$ values.

Isotopic variations driven by chemical transformations of seawater nitrate, such as removal 248 249 reactions (i.e., reduction via denitrification) and input reactions (i.e., atmospheric dinitrogen fixation), would impart isotopically heavier and lighter signals on the $\delta^{15}N$ of seawater nitrate 250 respectively. However, in-situ water-column denitrification is unlikely as near-zero oxygen 251 concentrations are required to favor such an anaerobic metabolism (Dalsgaard, Thamdrup, Farías, 252 & Revsbech, 2012) whereas high ambient oxygen concentrations (~8 mL L⁻¹) exist along the 253 Labrador Shelf (Garcia et al., 2013). While sedimentary denitrification might be possible along 254 the shelf, this process typically does not induce isotopic fractionation in the water-column nitrogen 255 pool due to complete consumption in pore waters (Lehmann, Sigman, & Berelson, 2004). One 256 exception to this has been noted along the eastern Bering Sea Shelf in the polar north Pacific, 257 where benthic denitrification coupled to the partial nitrification of ammonia yields a net increase 258 in the δ^{15} N of seawater nitrate (Granger et al., 2011). As these waters are advected to Baffin Bay, 259 such δ^{15} N variations may be relevant for our record. Yet, nitrification first requires the conversion 260

of organic nitrogen to ammonia, and so both nitrification and benthic dentrification should also be linked to productivity. If productivity had historically been driven by sea-ice extent in this region, such as it presently is throughout much of the Arctic (e.g., Arrigo, van Dijken, & Pabi, 2008), this would not be a reasonable mechanism as multicentury Arctic sea-ice reconstructions do not show the centennial-scale variability that is observed in our record. However, additional paleoenvironmental reconstructions throughout the eastern Arctic would be helpful for resolving past variations in productivity and nitrogen cycling in this region.

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Arctic cyanobacteria fix atmospheric dinitrogen at similar rates to their low-latitude counterparts, 269 which suggests that fixation may play an important role in high-latitude nitrogen cycling (Harding 270 et al., 2018). Arctic fixation could lower the isotopic composition of seawater nitrate, feed into the 271 272 East Greenland Current and Baffin Current and, eventually, enter the Labrador Shelf (Figure 1a and 1b). However, fixation should be related to temperature and sea-ice variability due to the 273 274 requirement of open-ocean conditions for both cyanobacterial growth and Arctic throughflow (Yamamoto-Kawai, Carmack, & McLaughlin, 2006) and so the absence of centennial-scale 275 variability, which is characteristic of $\delta^{15}N_{algal}$ values, in Arctic sea-ice reconstructions argues 276 against such a mechanism. Further, the lack of a clear relationship between elevated temperatures, 277 lower ice extent and lower $\delta^{15}N_{algal}$ throughout the record argues against *in-situ* fixation driving 278 our record (Figure 2). In fact, the last ~130 years in our record reflect simultaneously the warmest, 279 lowest-ice conditions and the highest $\delta^{15}N_{algal}$ values. As such, dinitrogen fixation is not a likely 280 control on $\delta^{15}N_{algal}$. 281

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Two plausible driving mechanisms of $\delta^{15}N_{algal}$ variability remain: changes in the relative 283 proportion of Atlantic- and polar-sourced waters feeding the Labrador Current and/or changes in 284 285 the relative degree of nitrate utilization. The Labrador Current is fed by waters derived from the polar Baffin Bay and from the north-eastern Atlantic (Figure 1a and 1b). Today, north-eastern 286 Atlantic waters carry upper-ocean nitrate with a $\delta^{15}N$ value of ~4‰ (Marconi, Weigand, & 287 Sigman, 2019), whereas nitrate derived from the upper ocean of the polar Baffin Bay is 288 characterized by δ^{15} N values between ~6 to 8‰ (Lehmann et al., 2019). While such baseline nitrate 289 isotopic compositions may have changed through time, GEOTRACES studies indicate that 4‰ is 290 already among the lowest value observed in the entire contemporary north-eastern Atlantic 291

(Marconi et al., 2019), with only a few patches of lighter nitrate (~ 2 to 3‰) occuring 292 predominately in some western regions (Marconi et al., 2015). As such, it is not likely that the 293 average north-eastern Atlantic nitrate inflow source could be lighter than 4‰ without global-scale 294 perturbations to the oceanic nitrogen cycle. In Baffin Bay, nitrate is mostly derived from the north 295 Pacific, which could be subject to upstream denitrification and result in isotopically heavier 296 seawater nitrate (Cline & Kaplan, 1975; Sigman et al., 2005). However, denitrification has been 297 variable over the last century in the north Pacific, first decreasing until the 1990s and then 298 subsequently increasing due to anoxia (Deutsch et al., 2014). The lack of this variability in our 299 $\delta^{15}N_{algal}$ record over the last century thus indicates that any influence of Pacific denitrification on 300 Baffin Bay nitrate is not translated to the nitrate along the Labrador Shelf. As such, it is reasonable 301 to consider 4‰ (Atlantic) and ~6 to 8‰ (polar) as the lower and upper isotopic bounds of seawater 302 303 nitrate in the source waters that feed the Labrador Current (Figure 2).

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If utilization remained complete over the last 600 years, source-water variability could therefore 305 be responsible for δ^{15} N_{algal} values within the 4 to 8‰ range, which would represent two extreme 306 307 proportioning scenarios: a Labrador Current composed of entirely Atlantic-derived waters, or of entirely polar-derived waters. As such, $\delta^{15}N_{algal}$ values outside of this range would suggest 308 309 additional fractionation mechanisms at play, likely implicating a change in nitrate utilization given the unlikelihood of the other known mechanisms described above. During each period of high 310 $\delta^{15}N_{algal}$, the isotopic signature falls within the range of source-water nitrate values (Figure 2), 311 which could be consistent with nearly complete nitrate utilization (e.g., Altabet & Francois, 1994; 312 Mariotti et al., 1981). However, here we conservatively define "nearly complete utilization" by 313 considering only the upper range of allowable $\delta^{15}N_{algal}$ values; i.e., when $\delta^{15}N_{algal}$ approximates 314 the isotopic range of polar source waters (6 to 8‰). Importantly, by our definition these intervals 315 316 would also correspond to a circulation regime dominated by polar-sourced waters, rather than Atlantic-sourced waters, along the Labrador Shelf. Yet, several intervals characterized by $\delta^{15}N_{algal}$ 317 values lower than the $\delta^{15}N$ of Atlantic nitrate suggest that nitrate utilization did not always 318 approach completion. Thus, the central questions that this study aims to address are: 1) how did 319 320 nitrate utilization come to approach completion along the modern Labrador Shelf and 2) what drove periodic phases of increased polar influence and nearly complete utilization in the past? 321

323 3.2 Drivers of Nitrate Utilization along the Labrador Shelf

Nearly complete utilization may be caused by increased biological uptake, reduced nitrate supply 324 or a combination of both. Over the last century, productivity reconstructed from algal Ba/Ca (Chan 325 et al., 2017) has been highly variable while $\delta^{15}N_{algal}$ values remained relatively constant (Figure 326 2). This implies that nitrate utilization during this interval was at least partially controlled by 327 changes in nitrate supply driven by advective and mixing processes, rather than exclusively by 328 changes in biological uptake. Because changes in the Labrador Current, an advective nitrate 329 source, are associated with changes in regional deep-water formation, the historical variability of 330 regional convection may help to inform past variations in the Labrador Current's coastal inflow 331 and its impact on nitrate supply to the Labrador Shelf. Deep-water formation is typically discussed 332 in terms of its relationship with broader climatic fluctuations; notably, the North Atlantic 333 334 Oscillation (NAO) and Atlantic Multidecadal Oscillation (AMO), which are connected to temperature and circulation patterns in the north Atlantic (e.g., Seip et al., 2019). While recent 335 336 work indicates that the AMO may not be an oscillation per se, and may be more appropriately termed "Atlantic multidecadal variability" (Mann, Steinman, & Miller, 2020), here we use the 337 338 original AMO language to be consistent with previous authors that have generated paleoreconstructions of its behavior. 339

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The NAO describes the sea-surface pressure gradient between the Icelandic Low and Azores High 341 342 systems such that positive values (NAO+) correspond to increased pressure differences and negative values (NAO-) correspond to decreased pressure differences (Hurrell, 1995). During 343 NAO- conditions, the advection of nitrate-rich Atlantic slope waters to the shelf is enhanced 344 (Petrie, 2007). The AMO describes Atlantic multidecadal sea-surface temperature (SST) 345 variations. Like the NAO, it is typically discussed in terms of an index, where positive values 346 (AMO+) represent multidecadal-scale positive SST anomalies and negative index values (AMO-) 347 represent multidecadal-scale negative SST anomalies in the north Atlantic (Schlesinger & 348 Ramankutty, 1994). In the Labrador Sea, the AMO- mode is associated with increased cooling and 349 sea-ice extent (Day, Hargreaves, Annan, & Abe-Ouchi, 2012; Miles et al., 2014) along with 350 decreased productivity (Chan et al., 2017), while the AMO+ condition is characterized by the 351 opposite. A recent synthesis of observational data, modeling studies and paleoclimate 352 reconstructions suggests that the AMOC is responsible for driving this multidecadal variability 353

throughout the Atlantic, with AMO+ conditions corresponding to a strong AMOC and AMO-354 conditions corresponding to a weak AMOC (Zhang et al., 2019). Numerical simulations 355 additionally demonstrate that this variability has been significantly associated with the strength of 356 the AMOC on 100-year periodicities throughout the last 1400 years (Knight, Allan, Folland, 357 Vellinga, & Mann, 2005), reminiscent of the approximate pacing of phases of nearly complete 358 nitrate utilization present in the $\delta^{15}N_{algal}$ record (Figure 2). Below, we illustrate the relevance of 359 the NAO and AMO in modulating nitrate conditions along the Labrador Shelf, and how the 360 twentieth-century weakening of the Labrador Current's Atlantic component has disrupted these 361 historical dynamics. 362

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3.3 Periodic Phases of Nearly Complete Nitrate Utilization

The five phases of nearly complete nitrate utilization interpreted from high $\delta^{15}N_{algal}$ values over 365 the last 600 years may either be driven by increases in primary productivity or decreases in nitrate 366 supply. Light attenuation modulated by sea-ice demise is the main driver of productivity along the 367 past and present Labrador Shelf (Chan et al., 2017; Harrison et al., 2013). As such, and because 368 369 direct productivity reconstructions are not available until after 1600, we infer indirect productivity changes from paleo-reconstructed sea-ice cover for earlier time intervals in our record (Halfar et 370 371 al., 2013). Coherence of decreased sea-ice extent (and thus, inferred increased productivity) with increased $\delta^{15}N_{algal}$ values until 1500 supports a productivity-driven change in nitrate utilization 372 (Figures 3 and 4). However, $\delta^{15}N_{algal}$ leads sea-ice extent, suggesting that productivity could not 373 have controlled nitrate utilization during this brief first phase (~ 1470). Furthermore, a significant 374 antiphase association between the sea-ice proxy and $\delta^{15}N_{algal}$ during phase 2 (~1550 - 1590) 375 indicates that increased sea-ice extent (and thus, inferred decreased productivity) occurred in 376 tangent with increased $\delta^{15}N_{algal}$ values. Because decreased productivity should normally result in 377 lower nitrate utilization, and thus lower $\delta^{15}N_{algal}$ values, such a relationship discredits productivity 378 as the main driver of nitrate utilization during this interval. 379

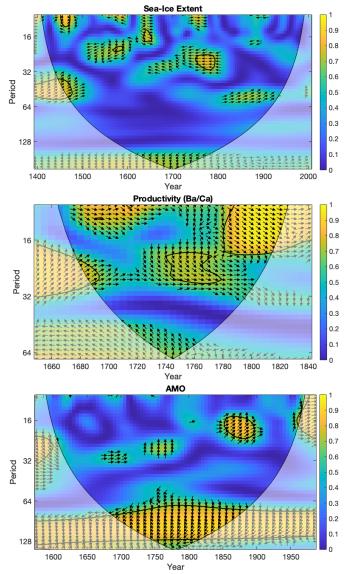
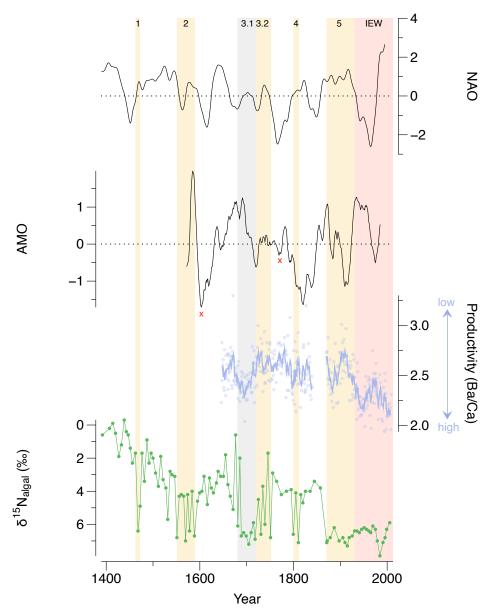


Fig. 3. Cross-wavelet coherence and phase relationships between paleo-records and $\delta^{15}N_{algal}$. 382 Significant overlaps in spectral power between signals at the 95% confidence level against red 383 noise are represented within the black contours. Arrows indicate phase relationships, where 384 leftward-pointing represents an antiphase relationship, rightward-pointing represents an in-phase 385 relationship, downward-pointing represents the titular time series leading our record and upward-386 pointing represents the opposite. The areas in which edge effects may interfere with analyses are 387 depicted by the shaded regions. All analyses were performed in MatLab R2020a using default 388 options in the Cross Wavelet and Wavelet Coherence toolbox provided by A. Grinsted. For 389 specifications of the relevant calculations, the reader is referred to the original publication 390 (Grinsted et al., 2004). Sea-ice, productivity and AMO paleo-reconstructions are from Halfar et 391 al. (2013), Chan et al. (2017) and Gray et al. (2004) respectively. 392

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395 Fig. 4. Climate variability, productivity and nitrate utilization. Phases of nearly complete utilization, corresponding to a circulation regime dominated by polar-sourced waters with 396 relatively high δ^{15} N_{nitrate} along the Labrador Shelf, are highlighted by the vertical shaded regions. 397 398 Phase 3.1 (grey interval) corresponds to an interval of increased primary productivity, whereas the 399 red interval represents the industrial-era weakening of the Labrador Current. Red x marks denote negative AMO excursions that did not trigger large increases in nitrate utilization, which are 400 hypothesized to have been offset by the concurrent negative NAO excursions (see Section 3.3). 401 NAO, AMO and productivity paleo-reconstructions are from Trouet et al., (2009), Gray et al., 402 (2004) and Chan et al., (2017) respectively. See the legend of Figure 2 for a full description of the 403 Ba/Ca data. 404

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In the 1600s, reconstructions of productivity become available from algal Ba/Ca data (Chan et al.,

407 2017). The onset of phase 3 (Phase 3.1; ~1680) occurred in concert with an increase in productivity

reconstructed from algal Ba/Ca (Figures 3 and 4). However, edge effects characteristic of the 408 wavelet analysis between Ba/Ca and $\delta^{15}N_{algal}$ prevent a robust statistical comparison of these 409 variables during this period, and sea-ice extent reconstructions do not support a significant 410 relationship between productivity and $\delta^{15}N_{algal}$ (Figure 3). Moreover, elevated $\delta^{15}N_{algal}$ values were 411 maintained well after Ba/Ca recovered to its pre-3.1 levels, until ~1750. Phase 4 (~1800 - 1810) 412 is characterized by low productivity indicated by the significant antiphase association between the 413 sea-ice proxy and $\delta^{15}N_{algal}$ (Figure 3). Further, significant relationships between Ba/Ca and $\delta^{15}N_{algal}$ 414 occurred in-phase during this interval. Since the Ba/Ca proxy is negatively correlated with 415 productivity, this also indicates that periods of low productivity corresponded to periods of high 416 nitrate utilization. Finally, during phase 5 (~1870 onward), nitrate utilization was entirely 417 uncorrelated with both sea-ice extent and Ba/Ca-reconstructed productivity (Figure 3). As such, 418 biological uptake appears to be generally unimportant for controlling utilization throughout these 419 phases, implicating changes in nitrate supply as important for driving variations in the record. 420

421

Changes in nitrate supply may be associated with ocean circulation patterns and therefore could 422 423 be sensitive to larger climatic variations. Specifically, a negative-mode AMO is linked to weak AMOC conditions (Zhang et al., 2019) and a weak Atlantic inflow, which could result in an 424 increased proportion of relatively nitrate-poor polar waters at the shelf. Indeed, phases of nearly 425 complete utilization correspond to negative excursions in the paleo-reconstructed AMO (Gray et 426 427 al., 2004) (Figure 4). Such a relationship is further supported by the significant antiphase correlation between the AMO and $\delta^{15}N_{algal}$ on periodicities of ~100 years (Figure 3), in good 428 agreement with the approximate pacing of phases of nearly complete utilization. Interestingly, the 429 relationship between the AMO and δ^{15} N_{algal} is the opposite of that previously documented between 430 the AMO and productivity at this site (Chan et al., 2017). Positive-mode AMO conditions were 431 argued to drive increased heating in the northern hemisphere and sea-ice melt along the shelf (Day 432 et al., 2012; Miles et al., 2014), which was linked to enhanced phytoplankton growth (Chan et al., 433 2017). The notion that negative AMO conditions appear to be associated with large increases in 434 nitrate utilization therefore provides further evidence that biological uptake is not a major driver 435 of our record and offers an explanation for the anticorrelation detected between $\delta^{15}N_{algal}$ values 436 and paleo-reconstructed productivity (Figure 3). 437

The antiphase relationship between the AMO and $\delta^{15}N_{algal}$ also argues against stratification as a 439 driver of increased nitrate utilization. A negative-mode AMO would promote cold conditions and 440 ice growth along coastal shelf waters (Day et al., 2012; Miles et al., 2014), which would generally 441 be expected to enhance mixing. As such, a reduction in advected nitrate modulated by the AMOC-442 induced weakening of the Labrador Current's Atlantic component is the most likely driving 443 mechanism of these five phases. However, two negative AMO excursions following phases 2 and 444 3 failed to trigger additional periods of nearly complete utilization (Figure 4). These intervals also 445 correspond to pronounced negative NAO modes (Trouet et al., 2009). Therefore, we hypothesize 446 that these negative NAO conditions facilitated the advection of higher-nitrate Atlantic slope waters 447 to the shelf (Petrie, 2007), thereby preventing large declines in nitrate supply otherwise expected 448 from the concurrent negative AMO. 449

450

Interestingly, persistently high nitrate utilization characteristic of the contemporary period (~1930 451 to present) appears to be insensitive to high-magnitude AMO fluctuations (Figure 4). During this 452 interval, the AMOC underwent an intense weakening (Caesar et al., 2018; Rahmstorf et al., 2015), 453 454 likely driven by anthropogenic climate forcing, concomitant with strong reductions in Labrador Sea convection (Thornalley et al., 2018) and a migration of nitrate-rich Atlantic waters away from 455 456 the coast (Thibodeau et al., 2018; Thibodeau et al., 2010). As such, this recent shift in the Labrador Current likely also reduced advective nitrate supply to the shelf, and is the most plausible 457 458 mechanism for the anomalously prolonged phase of increased polar-sourced waters and nearly complete utilization characteristic of the last 150 years. 459

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3.4 Environmental Implications

While previous studies have commented on the relationships between AMOC strength, nutrient 462 463 supply and productivity in the north Atlantic, such studies have historically invoked upper-ocean stratification as the main mechanism responsible for these associations (Osman et al., 2019; 464 Schmittner, 2005). However, recent modeling efforts argue that advection, rather than mixing, is 465 the dominant control on nitrate supply to the subpolar north Atlantic region (Whitt & Jansen, 466 2020). Our findings support this claim at the Labrador Shelf, further highlighting the importance 467 of upper-ocean circulation in routing nitrate away from the coast. Additionally, our study lends 468 new field evidence to the previously proposed hypothesis that biogeochemical changes along the 469

coastal north-western Atlantic may be associated with the larger-scale AMOC decline 470 characteristic of the industrial era (Claret et al., 2018). While previous nutrient reconstructions 471 based on coral compound-specific δ^{15} N measurements at the more-southern Gulf of Maine argue 472 for an enhanced supply of nitrate coincident with the onset of the industrial era (Sherwood et al., 473 2011), we show the opposite trend occurring at the Labrador Shelf. The increased nitrate supply 474 to the Gulf of Maine was attributed to an increased presence of nitrate-rich subtropical Atlantic 475 waters, which is also a direct consequence of the recent changes to the Labrador Current 476 (Sherwood et al., 2011; Thibodeau et al., 2018). However, these subtropical waters bend eastward 477 to feed the North Atlantic Current prior to recirculating and joining the Labrador Current, and 478 therefore do not directly enter the coastal Labrador Shelf (Figure 1a and 1b). Thus, the industrial-479 era changes to the Labrador Current have resulted in differential impacts along the Labrador Shelf 480 and the Gulf of Maine regions. While the Labrador Shelf has lost nitrate due to the increased 481 proportion of nitrate-poor polar waters that comprise the Labrador Current's coastal inflow, the 482 Gulf of Maine has gained nitrate due to an increased influence of nitrate-rich subtropical Atlantic 483 waters. The ecological and biogeochemical consequences of such nutrient reorganizations in the 484 485 north Atlantic are not fully understood and should be subject to further research. For example, changing nitrate distributions may alter the locations and extent of primary productivity, which 486 487 ecological models suggest strongly affect the vulnerable and socioeconomically important Atlantic cod (Ehrnsten, Bauer, & Gustafsson, 2019). In addition to the possible impacts on regional 488 489 fisheries, changes in primary productivity may have implications for the future potential of oceanic carbon storage in the region (Takahashi et al., 2009). Thus, numerous environmental challenges 490 491 may be exacerbated by ongoing anthropogenic disturbances to the north Atlantic.

492

493 **4** Conclusions

Here, we have described mechanisms that drove changes in circulation patterns and nitrate utilization along the Labrador Shelf over the last 600 years. We show that changes in nitrate uptake driven by productivity did not trigger periodic intervals of nearly complete nitrate utilization observed in our record, and rather suggest that the increased proportion of nitrate-poor polar waters associated with negative AMO excursions drove the occurrence of such intervals on approximately centennial timescales. Additionally, we suggest that the reduction in the strength of the AMOC via its weakening of Atlantic inflow likely contributed to a persistently low supply of nitrate to the Labrador Shelf, resulting in an anomalously prolonged phase of nearly complete utilization since ~ 1870 , which should have otherwise ended in the early 1900s. In contrast to past cycles of nitrate supply, nutrient conditions will not likely recover due to the ongoing reduction of Atlantic influence at the shelf. Our study thus adds to the abundant body of evidence illustrating the intense sensitivity of the high-latitude north Atlantic region to modern climate change.

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- 514

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