

# Climb of jogs as a rate-limiting process of screw dislocations motion in olivine dislocation creep

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November 30, 2022

## Abstract

Dislocation recovery experiments were conducted on pre-deformed olivine single crystals at temperatures of 1,450 to 1,760 K, room pressure, and oxygen partial pressures near the Ni-NiO buffer to determine the annihilation rates constants for [001] (010) edge dislocations. The obtained rate constants were comparable with those of previously determined [001] screw dislocations. The activation energies for the motion of both dislocations are identical. This suggests that the motion of screw dislocations in olivine is not controlled by cross-slip but by the same rate-limiting process of the motion of edge dislocations, i.e. climb, at low-stress and high-temperature conditions. The diffusivity derived from dislocation climb indicates that dislocation recovery is controlled by pipe diffusion. The conventional climb controlled model for olivine can be applied to the motions of not only edge but also screw dislocations. The softness of the asthenosphere cannot be explained by the cross-slip controlled olivine dislocation creep.

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2 **dislocation creep**

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## 11 **Abstract**

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13 temperatures of 1,450 to 1,760 K, room pressure, and oxygen partial pressures near the Ni-NiO buffer  
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16 activation energies for the motion of both dislocations are identical. This suggests that the motion of  
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22 cross-slip controlled olivine dislocation creep.

23  
24 **Keywords:** dislocation recovery, dislocation creep, temperature dependence, climb controlled model,  
25 asthenosphere

## 27 **Introduction**

28 Geophysical observations regarding geoid (e.g. Hager, 1991) and post-glacial rebound (e.g.  
29 Peltier, 1998) suggested that a soft asthenosphere underlies a rigid lithosphere. Geodynamic  
30 modellings (e.g. Becker, 2017; Craig and McKenzie, 1986) also suggested the same conclusion. The  
31 reason for the soft asthenosphere is under debate. Although the simplest explanation is weakening of  
32 materials due to the high temperature, results of deformation experiments on dry peridotite implied  
33 that the high temperatures are insufficient to explain the softness of asthenosphere (Hirth and

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34 Kohlstedt, 2003). Although a popular explanation is hydrous weakening of olivine (e.g. Mackwell et  
35 al., 1985; Hirth and Kohlstedt, 2003), it is doubted by recent Si self-diffusion experiments (Fei et al.,  
36 2016; Fei et al., 2013) based on the assumption that dislocation creep is controlled by diffusion.  
37 Another possible explanation was proposed by Poirier and Vergobbi (1978). They suggested that, if  
38 cross-slip of dissociated screw dislocations controls olivine dislocation creep, the estimated upper-  
39 mantle viscosity would be one order of magnitude lower than that predicted by the climb controlled  
40 model in the stress range from 10 to 100 bar. This potentially explains the softness of the  
41 asthenosphere. However, there is no experimental study to test this hypothesis.

42       Neither diffusion nor deformation experiments can identify the rate-limiting process of motions  
43 of screw dislocations. Diffusion does not involve motions of dislocations. Although it is theoretically  
44 possible to determine the rate-limiting process of dislocation motions by examining the stress  
45 dependence of creep rates (e.g. Hirth and Kohlstedt, 2003), the stress ranges in deformation  
46 experiments are too narrow. Conventionally used stress exponent of 3.5 for dislocation creep implies  
47 a pipe diffusion controlled mechanism (Hirth and Kohlstedt, 2003, 2015). However, those  
48 experiments have a stress range only from 100 to 224 MPa. On the other hand, Kohlstedt and Goetze  
49 (1974) found that stress exponent increases with increasing stress. Poirier (1985, P.139) found that  
50 the stress exponent of olivine single crystal dislocation creep varies from 2.6 to 3.7 in different studies.

51       In this study, we conducted dislocation recovery experiments on [001](010) edge dislocations  
52 and compared the results with those of [001](010) screw dislocations from Wang et al. (2016). During  
53 recovery, dislocations move on the slip plane (glide) and out of the slip plane (climb, cross-slip)  
54 successively under the influence of internal stress. Therefore, the activation energy determined by  
55 this method represents that of the rate-limiting process of dislocation motions. Although the model

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56 from Poirier and Vergobbi (1978) was based on [100] screw dislocations, most of them have edge  
57 character at temperature less than 1350 °C (Bai and Kohlstedt, 1992; Wang et al., 2016). On the other  
58 hand, the similar density of edge and screw dislocations in [001](010) slip system (Wang et al., 2016)  
59 indicates equivalent importance of both types of dislocations in this slip system. Therefore, we focus  
60 on this slip system in this study (here after called *c*-dislocations).

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## 62 **Experimental Procedure**

63 The same Pakistan olivine and sample preparation procedure as those of Wang et al. (2016) were  
64 employed in this study. The composition of olivine was reported by Gose et al. (2010). The  
65 experimental setup is similar with that in Wang et al. (2016). The olivine single crystal was orientated  
66 by X-ray diffraction and electron backscattered diffraction (EBSD), and then placed in the cell  
67 assembly such that the [001] direction and (010) plan are parallel to the shear direction and plane,  
68 respectively.

69 Dislocations with the [001] Burgers vector on the (010) plane were produced by experimental  
70 deformation using a Kawai-type multi-anvil apparatus at University of Bayreuth. The sample  
71 assembly was first pressurized to 3 GPa with a press load of 3.6 MN, and then heated to a temperature  
72 of 1,600 K and held for 15 min to sinter crushable alumina. After that, the assembly was further  
73 compressed to a press load of 3.9 MN for 15 min to deform the sample. After the deformation, the  
74 sample was quenched by switching off the heating power, and then decompressed to room pressure  
75 for more than 16 hours. Transmission electronic microscopy (TEM) by Wang et al. (2016) found  
76 [001](010) slip system was successfully activated and dominant by this procedure. The ratio of screw  
77 to edge dislocations was 3:2, as reported by Wang et al. (2016).

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78 The deformed olivine crystals were cut into eight cubic pieces, and paired into four groups, in  
79 which the two pieces in each group shared a common (100) plane. One piece from each pair was used  
80 to determine the initial dislocation density, while the other was used to determine dislocation density  
81 after the annealing. The annealing experiments were conducted at ambient pressure and temperatures  
82 of 1,460 to 1,760 K for 35 min to 24 hours using a gas mixing furnace. The oxygen partial pressure  
83 was controlled at  $10^{-6}$ - $10^{-8}$  MPa, which is near the Ni-NiO buffer, using a CO-CO<sub>2</sub> gas mixture. Table  
84 1 summarizes conditions of the annealing experiments.

85 Dislocations were observed using oxidation decoration technique (Kohlstedt et al., 1976, Karato  
86 1987). The corresponding areas away from sub-grain boundaries on the common (100) plane in the  
87 initial and annealed pieces from the same group were observed to determine the change in dislocation  
88 densities before and after annealing. Since [001](010) edge dislocations elongate in [100] direction,  
89 these dislocations show dots contrast on (100) plane in backscattered images after decoration. The  
90 dots per unit area were counted as the dislocation density.

91 The annihilation rate constants were calculated using the second-order dislocation recovery  
92 kinetics (Karato and Ogawa, 1982; Kohlstedt et al., 1980; Wang et al., 2016)

$$93 \quad k = \frac{\frac{1}{\rho_f} - \frac{1}{\rho_i}}{t}, \quad (2)$$

94 where  $\rho_f$  and  $\rho_i$  are the dislocation densities after and before annealing, respectively, and  $t$  is the  
95 annealing time. Because of the thermally activated process, the dislocation annihilation rate constant  
96 is assumed to follow the Arrhenius relationship:

$$97 \quad k = k_0 \exp\left(-\frac{E}{RT}\right) \quad (3)$$

98 where  $k_0$  is a constant,  $E$  is the activation energy of dislocation annihilation,  $T$  is the temperature,  
99 and  $R$  is the gas constant.

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101 **Results**

102 Table 1 shows experimental results together with the annealing conditions. Dislocation density  
103 in the samples before deformation is less than  $0.0004 \mu\text{m}^{-2}$ , which is negligible in comparison with  
104 dislocation density after deformation (Table 1). Figure 1 a and b shows back-scattered electron images  
105 of decorated dislocations in the corresponding areas in the samples from the same pair before and  
106 after annealing, respectively. The *c*-screw dislocations appear as lines and the *c*-edge dislocations as  
107 dots on the (100) plane due to their geometries. Decrease in dislocation density was observed by  
108 comparing the images before and after annealing. The water contents in the samples before and after  
109 annealing were below the detection limit of infrared spectroscopy. The transmission electron  
110 microscope images of the dislocation structures after deformation were given in Fig. 4 in Wang et al.  
111 (2016).

112 Figure 1c plots the logarithmic rate constants of *c*-edge dislocations annihilation against the  
113 reciprocal temperature. The results from the previous dislocation recovery experiments on *c*-screw  
114 dislocations (Wang et al., 2016) and from other studies on dislocation recovery kinetics are also  
115 plotted in this figure. The dislocation annihilation rate constants of *c*-edge and *c*-screw dislocations  
116 are comparable, but those of the *c*-screw are about half orders of magnitude higher than those of the  
117 *c*-edge. The temperature dependences for these two dislocations are identical. Their activation  
118 energies are  $E_{c\text{-edge}} = 400 \pm 20 \text{ kJ/mol}$  and  $E_{c\text{-screw}} = 400 \pm 30 \text{ kJ/mol}$  for the *c*-edge and *c*-screw,  
119 respectively.

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121 **Discussion**

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122 The identical activation energies of annihilation rate constants of the  $c$ -edge and  $c$ -screw  
123 dislocations indicate that the motions of both dislocations are controlled by the same mechanism.  
124 Although many transport properties of olivine exhibit activation energies between 300 to 500 kJ/mol  
125 (e.g. Dohmen et al., 2002,  $529 \pm 41$  kJ/mol for silicon self diffusion,  $338 \pm 14$  kJ/mol for oxygen self  
126 diffusion), they are distinct from those determined in this study (also see the slope in Fig.3 and  
127 references there). The high accuracy in activation energies obtained in previous and this study allows  
128 us to distinguish the rate-limiting mechanisms of different processes.

129 The motion of edge dislocations is controlled by climb at high temperatures and low stresses  
130 (e.g. Hull and Bacon, 2001; Kohlstedt, 2006). However, motion of a pure screw dislocation does not  
131 involve climb because screw segments have no specific slip plane (Hull and Bacon, 2001). Since jogs  
132 in screw dislocations has edge character, we propose that the motion of screw dislocation is controlled  
133 by climb of jogs (Fig. 2). A screw dislocation can form a jog by cross-slips to overcome obstacles  
134 that it meets during glide (Fig.2A and 2B). The slip plane of the jog is defined by its dislocation line  
135 ( $\mathbf{J}$ ) and the Burgers vector ( $\mathbf{b}$ ), indicated by the yellow plane. The parent screw dislocation glides in  
136 the  $y$  direction and therefore, the jog needs to climb in the  $y$  direction to move along with their parental  
137 dislocation so that the screw dislocation can go through the obstacle (Fig. 2C). This climb of jogs  
138 should serve as the rate-limiting process of the screw dislocation motions.

139 It is noted that although the climb of edge dislocation and jog motion of screw dislocation are  
140 essentially the same, the density of climbing parts on edge dislocations and that of jogs on screw  
141 dislocations may be different, which causes difference in magnitude of rate constants. Thus, only the  
142 slope in the Arrhenius plot can be a fingerprint of the essential mechanism of rate-limiting processes  
143 in dislocation recovery experiments.

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144 Since climb is controlled by diffusion, the diffusivities derived from annihilation rate constants,  
145  $D^R$  (based on Karato and Ogawan, 1989) were compared with those of silicon and oxygen diffusion  
146 in olivine (Fig. 3). None of these data fit  $D^R$  well. Instead,  $D^R$  fall between silicon lattice and grain  
147 boundary diffusivities. This indicates that the dislocation climb in olivine may be controlled by pipe  
148 diffusion. Vacancies, dislocations and grain boundaries are 0-, 1-, 2-dimension defects, the structure  
149 distortion near these defects should increase consequently and accordingly, the associated Si  
150 diffusivity should increase consequently. In addition, the activation energy of  $D^R$  obtained in this  
151 study is between those of Si lattice (540 kJ/mol, Dohmen et al., 2002) and grain boundary diffusion  
152 (~200 kJ/mol, Fei et al., 2015). This is also consistent with the hypothesis that pipe diffusion controls  
153 dislocation climb. Although there is no data for the pipe diffusion in olivine, the fact that diffusion  
154 coefficient and activation energy of pipe diffusion are between those of lattice and grain boundary  
155 diffusion are well established in oxides (Frost and Ashby, 1983, Table 12.1). The low activation  
156 energy of oxygen lattice diffusion (~340 kJ/mol, Dohmen et al., 2002) rules out the possibility that  
157 oxygen diffusion controls dislocation climb.

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## 159 **Implications**

160 It was proposed that the softness of asthenosphere can be explained by the cross-slip controlled  
161 dislocation creep of olivine (Poirier and Vergobbi, 1978). However, this hypothesis was never tested.  
162 The present study demonstrates that motion of [001] screw dislocations is controlled by climb of jogs  
163 rather than cross-slip, suggesting that cross-slip model is not applicable for such dislocations in  
164 olivine. Although the cross-slip model is based on [100] dislocations, study of [001](010) slip system  
165 is more relevant to the asthenosphere conditions. Dislocation structure analyses indicated that most

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166 [100](010) dislocations have edge character (Bai and Kohlstedt, 1992, Wang et al., 2016), indicating  
167 that olivine dislocation creep cannot be controlled by the motions of [100] screw dislocation. On the  
168 other hand, the similar density (Wang et al., 2016) of edge and screw dislocations in [001](010) slip  
169 system indicates that both kinds of dislocations are important. Moreover, deformation experiments  
170 (e.g. Raterron et al., 2007) suggested that this slip system dominated at high pressure. Therefore, we  
171 conclude that cross-slip of screw dislocations cannot control the dislocation creep of olivine and  
172 accordingly, it cannot explain the softness of asthenosphere. The climb controlled model can be used  
173 in olivine dislocation creep regardless of dislocation characters.

174 The viscosity of asthenosphere extrapolated from dry olivine creep data using climb model is  
175 orders of magnitude higher than the estimated value from geophysical observation (Hirth and  
176 Kohlstedt 2003). Hydrous weakening of olivine was proposed to explain this discrepancy (e.g.  
177 Mackwell et al., 1985; Hirth and Kohlstedt, 2003) but it was disparaged by recent Si diffusion  
178 experiments (Fei et al., 2016; Fei et al., 2013). However, this study indicates that pipe diffusion, rather  
179 than lattice or grain boundary diffusion, may control the dislocation motions. This may explain the  
180 discrepant results between deformation and diffusion experiments. Further studies on the effect of  
181 water on the dislocation recovery or pipe diffusion in olivine are needed to identify the effect of water  
182 on the olivine rheology and to better explain the softness of the asthenosphere.

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187 **Acknowledgements**

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188 We thank H. Fischer, R. Njul in BGI for the sample and assembly preparation. This research was  
189 supported by DFG grants to TK (KA3434-3/1, KA3434-3/2, KA3434-7/1 and KA3434-8/1) and by  
190 the annual budget of BGI. All data used in this paper are in Table 1 and plotted in Figure. 1c  
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**Reference**

Bai, Q., and Kohlstedt, D. L., 1992, High-temperature creep of olivine single crystals, 2. dislocation structures: *Tectonophysics*, v. 206, no. 1–2, p. 1-29.

Becker, T. W., 2017, Superweak asthenosphere in light of upper mantle seismic anisotropy: *Geochemistry, Geophysics, Geosystems*.

Craig, C. H., and McKenzie, D., 1986, The existence of a thin low-viscosity layer beneath the lithosphere: *Earth and Planetary Science Letters*, v. 78, no. 4, p. 420-426.

Dohmen, R., Chakraborty, S., and Becker, H.-W. (2002) Si and O diffusion in olivine and implications for characterizing plastic flow in the mantle. *Geophysical Research Letters*, 29(21), 2030.

Farver, J.R., and Yund, R.A. (2000) Silicon diffusion in forsterite aggregates: Implications for diffusion accommodated creep. *Geophysical Research Letters*, 27(15), 2337-2340.

Fei, H., Koizumi, S., Sakamoto, N., Hashiguchi, M., Yurimoto, H., Marquardt, K., Miyajima, N., Yamazaki, D., and Katsura, T., 2016, New constraints on upper mantle creep mechanism inferred from silicon grain-boundary diffusion rates: *Earth and Planetary Science Letters*, v. 433, p. 350-359.

Fei, H., Wiedenbeck, M., Yamazaki, D., and Katsura, T., 2013, Small effect of water on upper-mantle rheology based on silicon self-diffusion coefficients: *Nature*, v. 498, no. 7453, p. 213.

Frost, H.J. and M.F. Ashby, *Deformation mechanism maps: the plasticity and creep of metals and ceramics*. 1982: Pergamon press.

Gose, J., Schmaedicke, E., Markowitz, M., and Beran, A., 2010, OH point defects in olivine from Pakistan: *Mineralogy and Petrology*, v. 99, no. 1-2, p. 105-111.

Hager, B. H., 1991, Mantle viscosity: A comparison of models from postglacial rebound and from the geoid, plate driving forces, and advected heat flux, *Glacial isostasy, sea-level and mantle rheology*, Springer, p. 493-513.

Hirth, G., and Kohlstedt, D., 2003, Rheology of the Upper Mantle and the Mantle Wedge: A View from the Experimentalists, *Inside the Subduction Factory*, American Geophysical Union, p. 83-105.

Hirth, G., and Kohlstedt, D., 2015, The stress dependence of olivine creep rate: Implications for extrapolation of lab data and interpretation of recrystallized grain size: *Earth and Planetary Science Letters*, v. 418, p. 20-26.

Hull, D., and Bacon, D. J., 2001, *Introduction to dislocations*, Butterworth-Heinemann. P.257

Karato, S., and Ogawa, M., 1982, High-pressure recovery of olivine: implications for creep mechanisms and creep activation volume: *Physics of the Earth and Planetary Interiors*, v. 28, no. 2, p. 102-117.

Karato, S., Scanning electron microscope observation of dislocations in olivine. *Physics and Chemistry of Minerals*, 1987. 14(3): p. 245-248.

Kohlstedt, D.L., et al., New Technique for Decorating Dislocations in Olivine. *Science*, 1976. 191(4231): p. 1045-1046.

Kohlstedt, D., Nichols, H., and Hornack, P., 1980, The effect of pressure on the rate of dislocation recovery in olivine: *Journal of Geophysical Research: Solid Earth (1978–2012)*, v. 85, no. B6, p. 3122-3130.

Kohlstedt, D. L., 2006, The Role of Water in High-Temperature Rock Deformation: Reviews in *Mineralogy and Geochemistry*, v. 62, no. 1, p. 377-396.

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- 236 Kohlstedt, D. L., and Goetze, C., 1974, Low-stress high-temperature creep in olivine single crystals:  
237 Journal of Geophysical Research, v. 79, no. 14, p. 2045-2051.
- 238 Mackwell, S.J., Kohlstedt, D.L., and Paterson, M.S. (1985) The role of water in the deformation of  
239 olivine single-crystals. Journal of Geophysical Research-Solid Earth and Planets, 90(NB13),  
240 1319-1333. Peltier, W., 1998, Postglacial variations in the level of the sea: Implications for  
241 climate dynamics and solid - earth geophysics: Reviews of Geophysics, v. 36, no. 4, p. 603-  
242 689.
- 243 Poirier, J.-P., 1985, Creep of crystals: high-temperature deformation processes in metals, ceramics  
244 and minerals, Cambridge University Press.
- 245 Poirier, J.-P., and Vergobbi, B., 1978, Splitting of dislocations in olivine, cross-slip-controlled creep  
246 and mantle rheology: Physics of the Earth and Planetary Interiors, v. 16, no. 4, p. 370-378.
- 247 Raterron, P., Chen, J., Li, L., Weidner, D., and Cordier, P. (2007) Pressure-induced slip-system  
248 transition in forsterite: Single-crystal rheological properties at mantle pressure and  
249 temperature. American Mineralogist, 92(8-9), 1436-1445.
- 250 Wang, L., Blaha, S., Pintér, Z., Farla, R., Kawazoe, T., Miyajima, N., Michibayashi, K., and Katsura,  
251 T., 2016, Temperature dependence of [100](010) and [001](010) dislocation mobility in  
252 natural olivine: Earth and Planetary Science Letters, v. 441, p. 81-90.
- 253 Weertman, J., 1955, Theory of Steady - State Creep Based on Dislocation Climb: Journal of Applied  
254 Physics, v. 26, no. 10, p. 1213-1217.
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256 **Figure and table captions**

257 Figure 1. BEIs showing the dislocation density (a) before and (b) after annealing at 1760 K for 35  
258 minutes. The images were taken on the (100) plane. Screw and edge dislocations are lines and dots,  
259 respectively, due to the geometries of their dislocation lines. The yellow scale bar represents 2  $\mu\text{m}$ .  
260 (c) Logarithmic dislocation annihilation rate constants of  $c$ -edge dislocations versus reciprocal  
261 temperature. Together plotted are the annihilation rate constants of  $c$ -screw dislocations from Wang  
262 et al., (2016). The activation energies for both dislocations are identical, i.e. 400 kJ/mol. Previous  
263 results on dislocation recovery are also plotted for comparison.

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266 Figure 2. A schematic diagram showing the jog-climb controlled motion of a screw dislocation. (a)  
267 The screw dislocation (blue line) is elongated in the  $x$  direction, which is parallel to its Burgers vector  
268  $\mathbf{b}$ , and glides in the  $y$  direction. The blue dot represents the obstacle that the screw dislocation meets  
269 during glide. (b) A jog (red segment) elongated in the  $z$  direction is produced on the screw dislocation  
270 in order to overcome the obstacle. This jog has an edge nature with the same Burgers vector  $\mathbf{b}$  as the  
271 parental screw dislocation. The yellow area indicates the glide plane of the jog, which is normal to  
272 the  $y$  direction. (c) The jog has to climb out of its glide plane in order to move along with its parental  
273 screw dislocation.

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275 Figure 3. Logarithmic diffusivity derived from dislocation annihilation rate constants of  $c$ -edge and  
276  $c$ -screw dislocations versus reciprocal temperature. Together plotted are Si and O lattice and  
277 grainboundary diffusivities.

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280 Table 1. Summary of experiment conditions and results.

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282 Table 1. Summary of experimental conditions and results\*.

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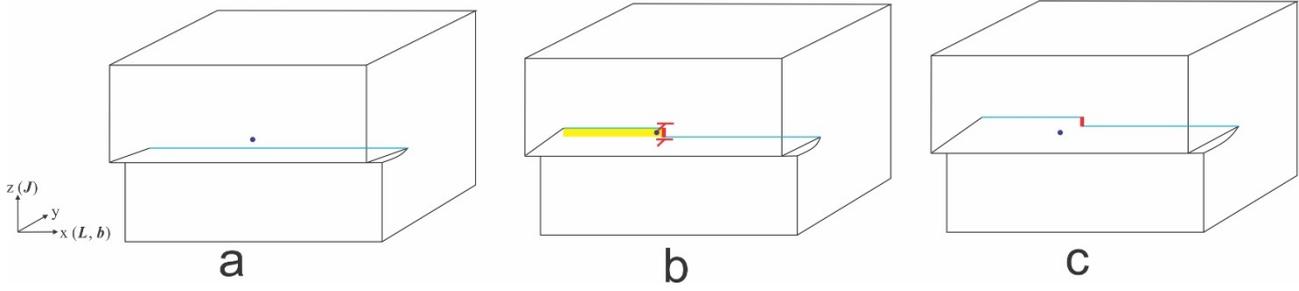
[001](010) edge dislocation						
Sample	$T$ (K)	Annealing time (h)	$\log(f\sigma_2, 10^5$ Pa)	$\rho_i$ ( $\mu\text{m}^{-2}$ )	$\rho_f$ ( $\mu\text{m}^{-2}$ )	$\log(k, \text{m}^2\text{s}^{-1})$
Z1643-1	1763	0.58	-4.9	$1.60 \pm 0.13$	$0.29 \pm 0.01$	$-14.87 \pm 0.03$
				$0.97 \pm 0.13$	$0.22 \pm 0.01$	$-14.77 \pm 0.03$
Z1643-2	1673	2.5	-5.7	$1.49 \pm 0.04$	$0.36 \pm 0.06$	$-15.63 \pm 0.09$
				$1.13 \pm 0.12$	$0.31 \pm 0.03$	$-15.58 \pm 0.05$
Z1643-3	1473	24	-7.7	$1.33 \pm 0.15$	$0.73 \pm 0.05$	$-17.14 \pm 0.09$
				$0.35 \pm 0.03$	$0.29 \pm 0.01$	$-17.22 \pm 0.27$

284 \* different  $\rho_i$  and  $\rho_f$  in each sample correspond to different areas

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291 Figure 2.



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