

Noise Characteristics of Operational Real-Time High-Rate GNSS Positions in a Large Aperture Network

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Abstract

Large earthquakes are difficult to model in real-time with traditional inertial seismic measurements. Several algorithms that leverage high-rate RT-GNSS positions have been proposed and it has been shown that they can supplement the earthquake monitoring effort. However, analyses of the long-term noise behavior of high-rate RT-GNSS positions, which are important to understand how the data can be used operationally by monitoring agencies, have been limited to just a few sites and to short time spans. Here we show results from an analysis of the noise characteristics of one year of positions at 213 RT-GNSS sites spanning a large geographic region from Southern California to Alaska. We characterize the behavior of noise and propose several reference noise models which can be used as baselines to compare against as technological improvements allow for higher precision solutions. We also show how to use the reference noise models to generate realistic synthetic noise that can be used in simulations of HR-GNSS waveforms. We discuss spatiotemporal variations in the noise and their potential sources and significance. We also detail how noise analysis can be used in a dynamic quality control to determine which sites should or should not contribute positions to an earthquake modeling algorithm at a particular moment in time. We posit that while there remain important improvements yet to be made, such as reducing the number of outliers in the time series, the present quality of real-time HR-GNSS waveforms is more than sufficient for monitoring large earthquakes.

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Key Points

- We study the noise behavior of 213 HR-GNSS sites from California to Alaska
- We characterize the spatiotemporal noise behavior and propose reference noise models
- The present real-time noise is low enough that GNSS can be used for monitoring earthquakes

Abstract

Large earthquakes are difficult to model in real-time with traditional inertial seismic measurements. Several algorithms that leverage high-rate RT-GNSS positions have been proposed and it has been shown that they can supplement the earthquake monitoring effort. However, analyses of the long-term noise behavior of high-rate RT-GNSS positions, which are important to understand how the data can be used operationally by monitoring agencies, have been limited to just a few sites and to short time spans. Here we show results from an analysis of the noise characteristics of one year of positions at 213 RT-GNSS sites spanning a large geographic region from Southern California to Alaska. We characterize the behavior of noise and propose several reference noise models which can be used as baselines to compare against as technological improvements allow for higher precision solutions. We also show how to use the reference noise models to generate realistic synthetic noise that can be used in simulations of HR-GNSS waveforms. We discuss spatiotemporal variations in the noise and their potential sources and significance. We also detail how noise analysis can be used in a dynamic quality control to determine which sites should or should not contribute positions to an earthquake modeling algorithm at a particular moment in time. We posit that while there remain important improvements yet to be made, such as reducing the number of outliers in the time series, the present quality of real-time HR-GNSS waveforms is more than sufficient for monitoring large earthquakes.

1. Motivation

44

45 There is broad interest in the international earthquake monitoring community in high rate
46 (HR, epoch length ≤ 1 s) real-time position estimation from Global Navigation Satellite
47 Systems (GNSS) such as the Global Positioning System (GPS) and others. It has been
48 shown that HR-GNSS displacement waveforms can supplement measurements from
49 traditional seismic networks based on inertial sensors and can be leveraged to characterize
50 moderate to large earthquakes in seconds to minutes. This interest arises because
51 algorithms that rely on inertial sensors “saturate” for large events, particularly at local and
52 regional distances (e.g. Hoshiya and Ozaki, 2014). Saturation means that large and very
53 large events look similar in inertial recordings and cannot be distinguished from one
54 another in the first minutes following a significant event. The exact causes for this are still a
55 matter of some debate but are most likely that the long period band of ground motion
56 (period >10 s) is not faithfully recorded by strong motion sensors in the near-field. In
57 contrast, this low frequency energy which distinguishes large events is recorded with fidelity
58 by HR-GNSS from the Nyquist frequency out to and including static, or permanent, offsets.
59 As a result, many researchers have studied and proposed algorithms based on HR-GNSS
60 that compute magnitude [Melgar et al., 2015], focal mechanisms (faulting style) [Crowell et
61 al. 2016, Riquelme et al. 2016], and slip distribution [Grapenthin et al. 2014, Minson et al.,
62 2014; Kawamoto et al. 2016] in real- or near real-time. Several of these algorithms have
63 been systematically evaluated with both real and simulated events and are being used to
64 complement traditional seismic approaches in earthquake and tsunami early warning
65 systems. Thorough reviews of these issues can be found in Bock & Melgar [2016] and
66 Larson [2019].

67

68 Measurements of ground motion from HR-GNSS differ from those obtained by inertial
69 seismic sensors in fundamental ways. In the electro-mechanical systems used in
70 seismometry the digitized acceleration or velocity of a proof mass inside the instrument
71 correlates directly, through a known transfer function, to the actual ground motion. HR-
72 GNSS positions are a wholly different kind of derived product. As a space-based geodetic
73 approach, calculation of HR-GNSS positions relies on measurement of the time of flight of
74 a microwave transmission between a satellite and a ground based antenna and receiver as
75 well as the phase with which the signal arrives. These measurements coupled with
76 knowledge of ancillary variables such as transmission delays through the troposphere and
77 ionosphere, knowledge of the satellite clocks and orbits, and others, are used by a
78 positioning algorithm to solve a least squares problem and produce epoch by epoch
79 *solutions* of the station coordinates in a particular reference frame. The most common
80 reference frame is the International Terrestrial Reference Frame (ITRF) which satellite
81 orbits are generally computed in [Altamimi et al., 2016]. If the GNSS antenna is firmly
82 coupled to the ground through a geodetic monument and it experiences a sudden motion,
83 such as the one produced by an earthquake, the position solutions can be used to obtain
84 displacement waveforms in local topocentric north, east, and vertical components of that
85 particular point of the surface of the Earth. While the concept behind GNSS positioning is in
86 essence simple, the estimation of the position of the antenna phase center using satellite
87 signals, the practice is complex, especially for high sample rates and in real-time. Satellite

88 orbits and clocks which determine the spatial and temporal origin of the microwave signal
89 used to solve for the position are not as well known in real-time as is necessary for precise
90 positioning. As a result, a number of external corrections must be calculated using a
91 reference regional network (e.g. Geng et al., 2013) and applied in real-time for the positions
92 to achieve the cm-level precision needed for earthquake monitoring.

93
94 In spite of this seemingly added complexity when compared to inertial sensors, real-time
95 HR-GNSS networks have proliferated to almost every tectonically active region (e.g.
96 Barrientos & Perez-Campos, 2018) and a variety of methods are employed to calculate the
97 GNSS positions. There exist proprietary software for positioning from a number of vendors
98 as well as open source academic codes (e.g. Geng et al., 2019). However, in spite of the
99 significant progress in positioning and in understanding how HR-GNSS can contribute to
100 real-time earthquake monitoring as well the rapid expansion of real-time networks one
101 important outstanding issue remains. What are the noise characteristics and long-term
102 behavior and performance of the position solutions in a real-world setting across a network
103 with large geographic aperture? Characterization of the actual real-time performance of
104 HR-GNSS has only been performed in small scale controlled settings such as shakatables
105 and on individual station to station baselines (e.g. Bock et al., 2000; Langbein & Bock,
106 2004; Genrich & Bock, 2006; Bock et al., 2011). Tests of real-time performance have also
107 been carried out in a simulated mode post-hoc for large events [e.g. Fang et al., 2013].
108 Recently Melgar et al., [2019] studied the performance of 9 HR-GNSS stations that were
109 recorded and positioned in real-time and broadcast to end users for the 2019 M6.4 and
110 M7.1 Ridgecrest, California earthquakes. When compared to post-processed solutions it
111 was found that the main features of the waveforms used for rapid source characterization,
112 the peak ground displacement (PGD) and the coseismic offsets compared favorably
113 between real-time and post-processed data. However, differences between real-time and
114 post-processed positions were also apparent. Post-processed solutions are able to
115 leverage final orbit and clock products as well as use iterative approaches and full time-
116 series filtering to compute positions whereas RT-GNSS positions can only utilize recursive
117 filters and rely heavily on phase ambiguity stability.

118
119 In this work we explore this issue further. We will study the long term noise characteristics
120 of real-time 1Hz point position time series computed in the ITRF global reference frame by
121 the Geodesy Lab at Central Washington University (CWU) for a network of 213 stations
122 (Figure 1) spanning from southern California to Alaska. These data are streamed from the
123 field site to CWU where positions are computed on the fly. The solutions are re-broadcast
124 to a number of users including the U.S. Geological Survey and the National Oceanographic
125 and Atmospheric Administration (NOAA). The data are streamed as well to the Universities
126 of Oregon and Washington where they are analyzed and archived. Here we will discuss the
127 temporal and spatial behavior of noise in these HR-GNSS solutions. We emphasize that
128 the performance we aim to characterized here is by definition a snapshot in time. GNSS
129 positioning technology is improving constantly and it is our hope that the noise models we
130 will demonstrate can be used by others to benchmark improvements and progress. These
131 proposed reference noise models can potentially be used to determine the quality of the

132 positioning solutions from a particular positioning algorithm or at a particular station of
 133 interest. Finally, we will demonstrate how to use the reference models to generate synthetic
 134 time series of noise which can be added to simulations of earthquake ground motions to
 135 more accurately represent a real-world scenario and to test rapid source estimation
 136 methods.

137

138 2. Data and Analysis Method

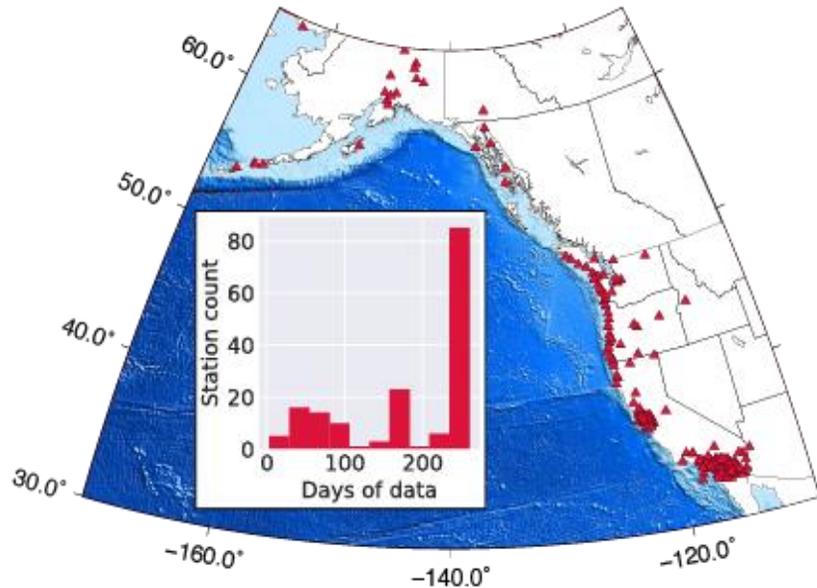
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140 2.1 Network and Positioning

141

142 Many continuous GNSS networks operate in the region spanned by this study (Figure 1,
 143 the U.S West Coast, Canada, and Alaska) and, while an exact figure on the number of
 144 available sites is hard to come by and changes frequently, it is likely on the order of ~1000
 145 stations (e.g. Blewitt et al., 2018). Of these, a subset of 213 was chosen for a
 146 demonstration project for NOAA. This agency is interested in using GNSS to supplement its
 147 local tsunami warning effort and so, starting in 2017, positions for this subset of sites began
 148 to be streamed in real-time to the Tsunami Warning Centers in Hawaii and Alaska
 149 [Melbourne et al., 2018]. In order to analyze the performance of the data, starting in
 150 October 2018 the positions are also being streamed to the University of Oregon where they
 151 archived as individual daily station files in miniSEED format for later analysis.

152



153

154 **Figure 1.** Distribution of real-time stations analyzed in this study. The inset histogram
 155 shows how many days of real-time data are available for each station.

156

157 The positions themselves are produced by Central Washington University's *FastLane*
 158 algorithm. The raw GNSS data are telemetered from the field to the central location for a
 159 particular network operator such as Boulder CO for UNAVCO Inc. sites or Berkeley CA, for
 160 UC Berkeley stations. From there the individual network operators streams the data to

161 CWU in Ellensburg, WA where Fastlane computes the epoch by epoch position solutions
162 and in turn serves them to other users such as NOAA and the UO.

163

164 The *Fastlane* positioning system [Santillan et al., 2013] produces precise point position
165 (PPP, Zumberge et al., 1997) estimates based primarily on GNSS carrier phase
166 observables (currently only from the GPS constellation) and satellite clock corrections
167 provided by the Real-Time Service (RTS) of the International GNSS Service (IGS). The
168 GPS carrier phase data is internally continuously calibrated using geometry free
169 combinations of the L1 and L2 pseudorange and phase observables. This calibration step
170 is a Kalman filter based algorithm that simultaneously estimates the best floating point
171 ambiguities while monitoring and correcting for possible cycle slips. Fastlane uses GPS
172 carrier phase based only, unlike other PPP algorithms (e.g. Kouba & Heroux 2001) that rely
173 on both phase and pseudorange. This approach to PPP relies on the fact that the
174 calibration procedure greatly mitigates the influence of code multipath that may affect the
175 estimation of the floating point ambiguities. By using well calibrated data Fastlane uses only
176 half the number of input observations therefore reducing the overall computation of the
177 position estimates which also translates into smaller latencies. This approach is far less
178 contaminated by multipath error, one of the largest sources of noise in high-rate
179 positioning. Fastlane uses a highly efficient algorithm for the resolution of carrier phase
180 initial ambiguities, which for most stations can be initially resolved in 20-30 seconds. After
181 this the positions can be efficiently determined. Positions are computed in SI units (meters)
182 in Earth Centered Earth Fixed reference frame (XYZ coordinates). Prior to streaming out
183 the solutions to users these are rotated to a more familiar topocentric local north, east, and
184 vertical reference frame. An example year-long waveform is shown in Figure 2.

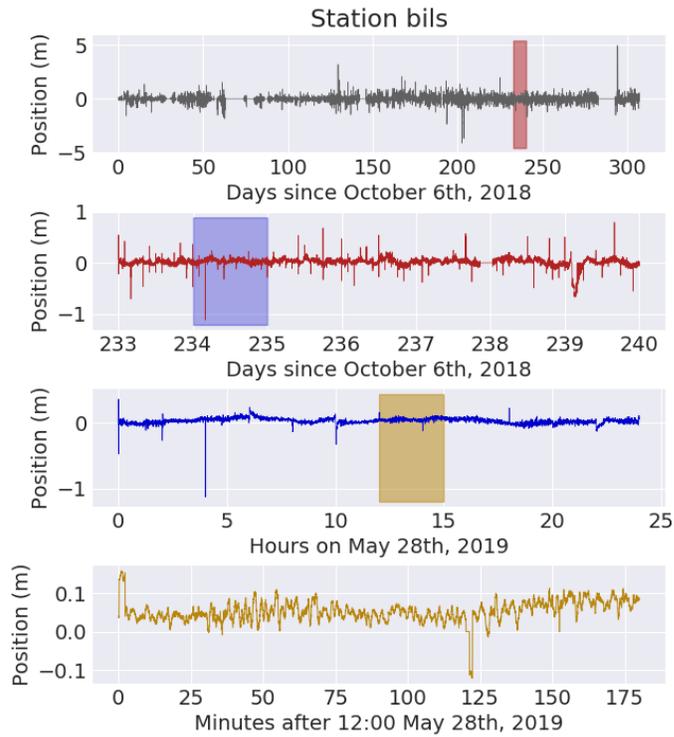
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186 2.2 Noise Analysis

187

188 First we study simple time domain features of the real-time waveforms such as the number
189 and amplitude of outliers. For every station we count how frequently displacement levels of
190 certain thresholds are exceeded in order to quantify the frequency of occurrence of the
191 large displacement excursions seen in Figure 2. However, the bulk of our analysis focuses
192 on the frequency domain. We employ the probabilistic power spectra (PPSD) technique of
193 McNamara & Buland (2004) for all sites. The PPSD method is common in seismology to
194 characterize the long term noise behavior of broadband sites. We take 20 min windows at
195 each site and for each of the three components of motions and calculate the power spectra.
196 This is repeated for every time window available for each site and an empirical probability
197 density function (PDF) of the distribution of power at each frequency is obtained for every
198 station. An example of the PPSD calculation for the same station in Figure 2 is shown in
199 Figure 3. The PPSD approach is desirable because it minimizes the need to “fix” issues
200 with the time series prior to calculating the spectra. As shown in Figure 2 there are outliers,
201 steps, and spikes, as well as gaps in the data. The PPSD will naturally deal with these. A
202 window with one of these behaviors will simply plot at a higher power. Meanwhile windows
203 without these issues, which are more frequent, will eventually illuminate the median
204 behavior as well as the lowest possible expected noise.

205



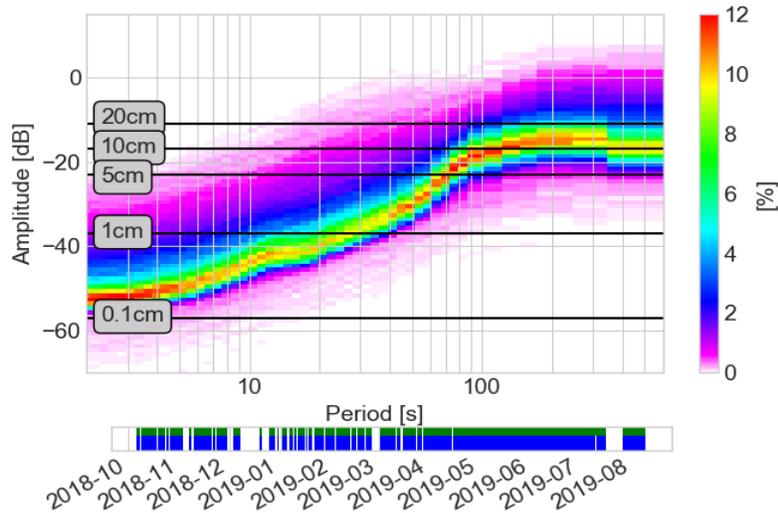
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Figure 2. Example time series for the east-west component of station BILS. Plotted are successive closeups of the data starting from the entire span and finishing with a three hour period.



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Figure 3. Example PPSD for the east-west component of displacement of station BILS (Figure 2). The black lines are used as a reference and denote the power of a Gaussian white noise time series with the specified standard deviation. The bar at the bottom denotes the time-spans covered by the data.

217 After we obtain PPSDs for each of the 213 sites in this study and for each component of
218 motion we aggregate all of them to obtain the overall behavior of the HR-GNSS noise.
219 From this regional PPSD we can extract reference noise models, for example, we select
220 the 1th percentile from the regional PPSD and term this the “low-noise” model. Similarly the
221 50th and 90th percentiles of the PPSD are use to define the “median” and “high-noise”
222 models.

223

224 2.3 Generation of synthetic noise time series

225

226 Using the regional reference noise models we demonstrate a simple method for generating
227 synthetic time series of noise that recreate the behavior observed in the real data. We
228 follow the approach first proposed by Boore [1983] and further detailed in Graves & Pitarka
229 [2010] for generating stochastic time series in seismology. The approach has three simple
230 steps, first, we create a Gaussian white noise time series with a specified sample rate (e.g.
231 1Hz) and duration. Second, we apply the Fourier transform to the white noise time series
232 and keep the random phase spectrum but replace the white noise PSD with the reference
233 noise model PSD. Finally we inverse Fourier transform to the time domain and recover a
234 time series. While we have proposed three reference models at the 1th, 50th, and 90th
235 percentiles, we have also extracted noise models for every 10th percentile. In the
236 acknowledgments we provide links to code and a tutorial that demonstrates how to
237 generate the synthetic time series.

238

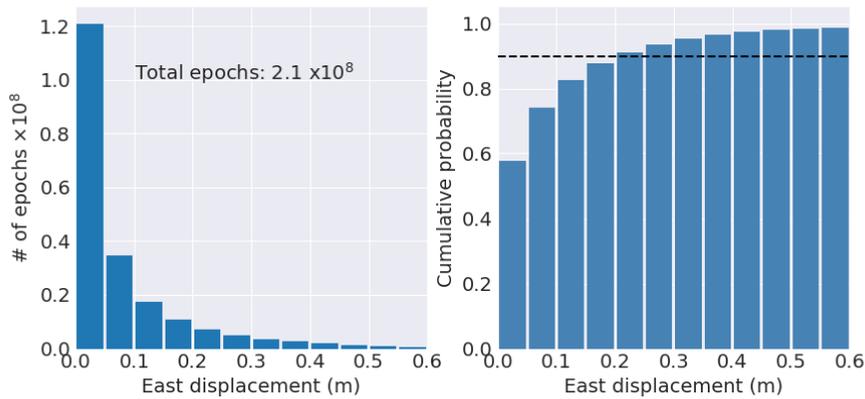
239 **3. Results**

240

241 3.1 Overall noise characteristics

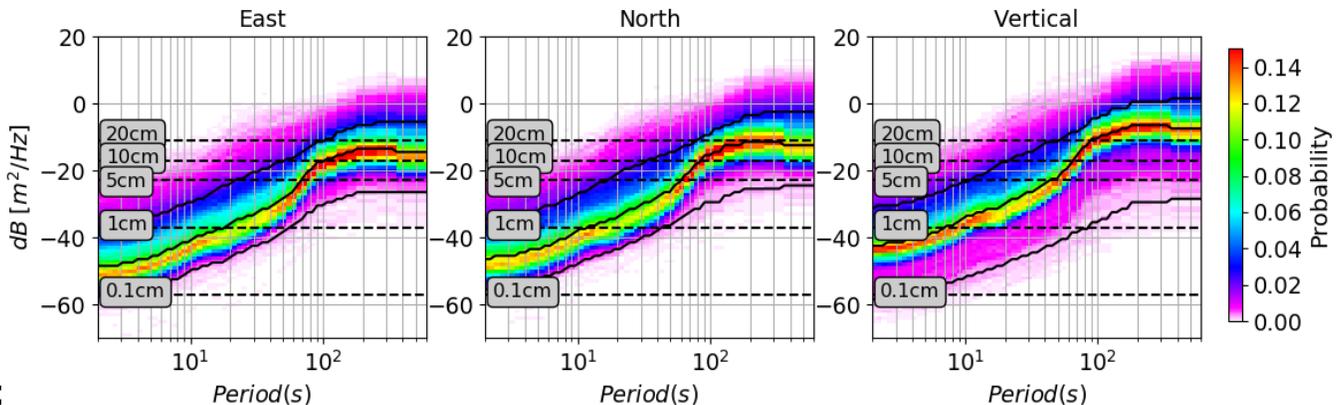
242

243 The time series in Figure 2 show seemingly meter-level accuracies in the positions, this is
244 far too high to satisfy the cm- to decimeter requirement needed to monitor large events
245 (e.g. Melgar et al., 2015; Ruhl et al, 2018), however if the data are plotted over shorter time
246 scales we can see that this is the result of outliers and that in reality over time scales of
247 minutes the data show cm-level precision. To further demonstrate this we take every 20min
248 segment, remove its mean and count how frequently it exceeds displacement thresholds of
249 certain levels. We do this for all sites and all epochs. The distribution of positions in the
250 east direction and the cumulative density function are in Figure 4 and show that in spite of
251 the outliers 90% of the data have noise smaller than 20cm.



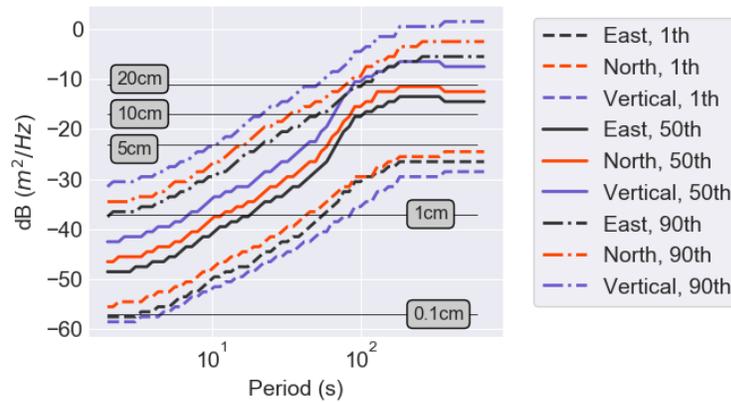
252
 253 **Figure 4.** Number of outliers in the position waveforms in the east direction for
 254 different thresholds and for all sites and all epochs. The black dashed line in the
 255 cumulative density function is the 90% level.
 256

257 Figure 5 shows the aggregate PPSD plot for all three components of motion for all stations.
 258 We note that as first described by Genrich & Bock [2006] the real time positions have
 259 roughly red noise with a plateau at periods longer than ~100s and decreasing noise levels
 260 at shorter periods. The noise is generally lowest for the east component, followed by the
 261 north component, with the highest noise levels in the vertical direction. This is consistent
 262 with what is seen in post-processed data (e.g. Bock et al., 2011, Melgar et al., 2019) and is
 263 usually attributed to the geometry of the constellation of satellites. This is more clearly seen
 264 in Figure 6 where the 1th, 50th, and 90th percentiles of the PPSDs for each component of
 265 motion are plotted together. The average difference in noise between each component of
 266 motion is about 3dB. We note that while the time domain analysis of the outliers in Figure 4
 267 suggests that noise levels in the 10-20cm range are not uncommon, the frequency domain
 268 analysis shows a more nuanced perspective. At shorter periods, shorter than 100s, which
 269 are comparable with the duration of large earthquake, noise is much closer to the ~5cm
 270 level. Meanwhile at shorter periods than that (e.g. 10s) 1cm or even sub-cm level noise is
 271 prevalent.



27:
 273 **Figure 5.** Aggregate PPSDs for the three components of motion for all stations in this
 274 study. The continuous black lines denote the 1th, 50th, and 90th percentiles. The dashed

275 lines are used as a reference and denote the power of a time series of Gaussian white
 276 noise with specified standard deviation.
 277



278
 279 **Figure 6.** Comparison plot of the 1th, 50th, and 90th percentiles of the PPSD noise
 280 distribution for all sites from Figure 5. The horizontal thin black reference lines are used as
 281 a reference and denote the power of a time series of Gaussian white noise with specified
 282 standard deviation

283

284 3.2 Spatiotemporal characteristics of noise

285

286 We are also interested in the variations of the noise distribution over long periods of time.
 287 Figure 7 shows the spectrogram for the nearly year-long time series at station BILS and for
 288 a period of 1 week. We find that, while there can be short periods of higher or lower noise
 289 overall, the general spectral shape and the behavior of the noise is somewhat stable. At a
 290 particular period the average standard deviation of the PSD throughout the year is only 5-
 291 6dB with larger excursion from this baseline behavior occurring only over short periods of
 292 time (e.g. P1 in Figures 7 and 8). The time series of power at selected periods, also shown
 293 in Figure 7, hint at regular variations in the noise behavior and also suggest that the
 294 temporal changes to the noise covary between periods. This is especially obvious in the
 295 week long time series. In Figure 9 we explore this further, we extract the time series of
 296 power spectral density at these 3 periods (2s, 60s, and 300s) for each site and calculate
 297 the spectra for each. We then stack them across all the sites to see if there are any spectral
 298 peaks that are systematically present at all sites. There are several (Figure 9), of particular
 299 prominence we note peaks at 1.96hrs, 11.38hrs and 21-23hr periods. This “spectra of
 300 spectra” should not be interpreted to suggest position signals at these periods, rather they
 301 show that with a periodicity of, for example ~2hrs, the entire spectra of the positions at all
 302 frequencies shift wholesale to higher or lower noise levels. The spectra explains some of
 303 the temporal variability in the noise behavior but we note, and will discuss further on, that
 304 the time series are punctuated by short periods of very high noise (e.g. P1 in Figures 7 and
 305 8) that occur at irregular intervals.

306

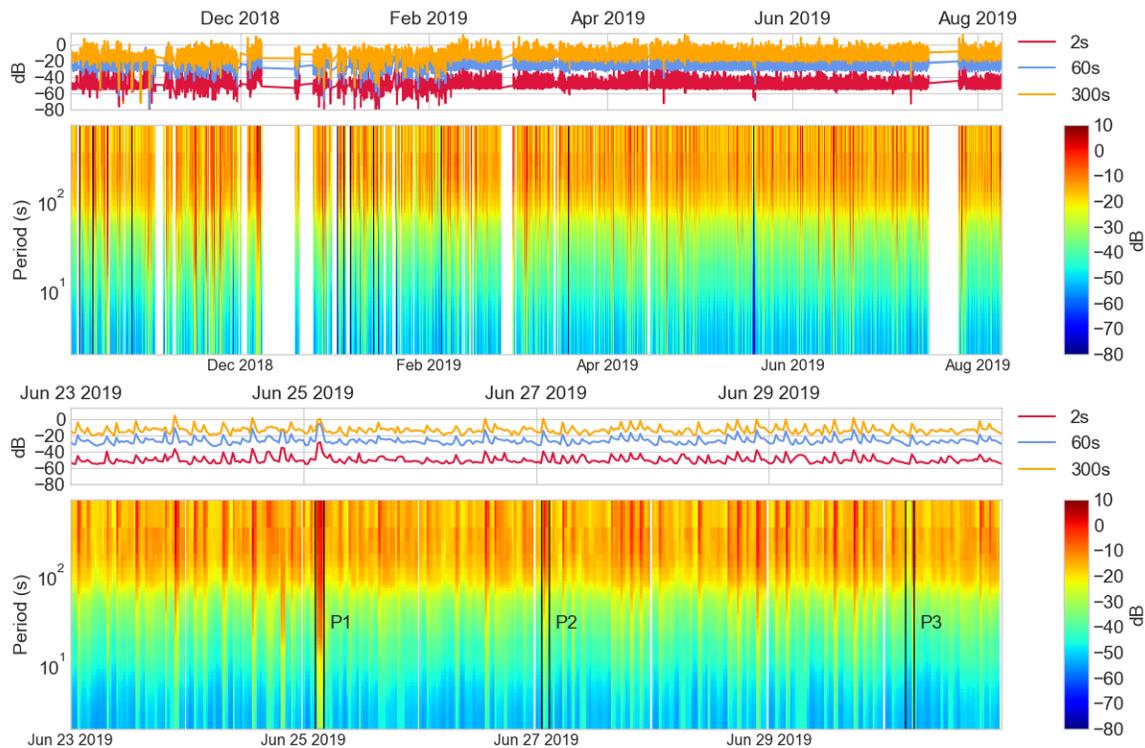
307 We also explore the spatial distribution of noise across the sites. Figure 10 is a map of the
 308 amplitude of the noise at a period of 60s across the entire network. We do not observe any
 309 strong spatial pattern with respect to preferential locations or environments for low or high
 310 noise sites. For example in the Southern California cluster there are many low noise sites
 311 (~-26dB) however, in between many of them are interspersed high noise sites with power
 312 closer to -22dB or -21dB. The same is true in the other three clusters in the Bay area, the
 313 Pacific Northwest, and Alaska. Additionally we do not observe systematically higher noise
 314 in any of the regions shown in Figure 10.

315

316 3.3 Synthetic noise time series

317

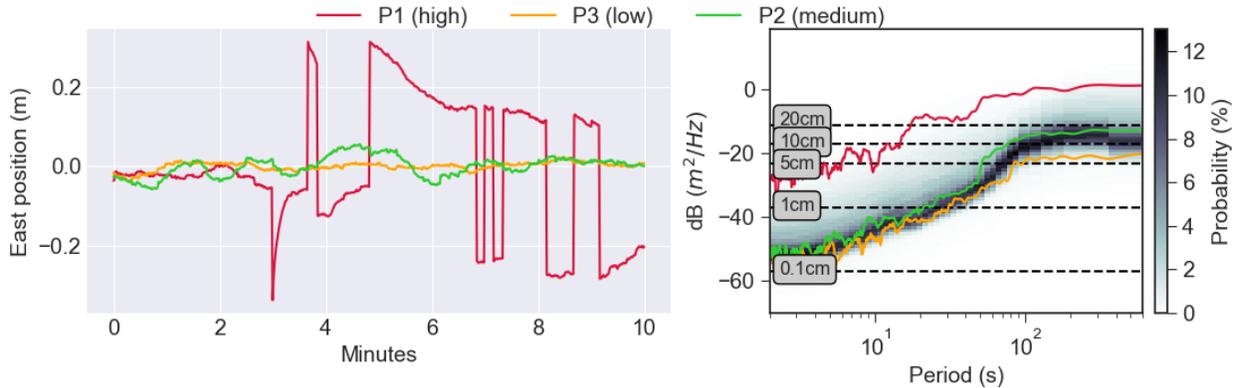
318 Figure 6 exemplifies three potential reference noise models from which to choose. We
 319 define the 1th percentile model as the “low” noise model, the 50th percentile as the
 320 “median” model, and the 90th percentile as the high noise model. These can be used to
 321 generate arbitrarily long synthetic time series of noise to be injected into simulations of
 322 earthquakes or any other potential application where high-rate positions are used or
 323 required [e.g. Melgar et al., 2016]. Figure 11 shows by way of an example a three
 324 component 20 minute time series of median synthetic noise compared to a twenty minute
 325 window from station BILS. The figure illustrates that the two are, as designed, very similar
 326 to each other.



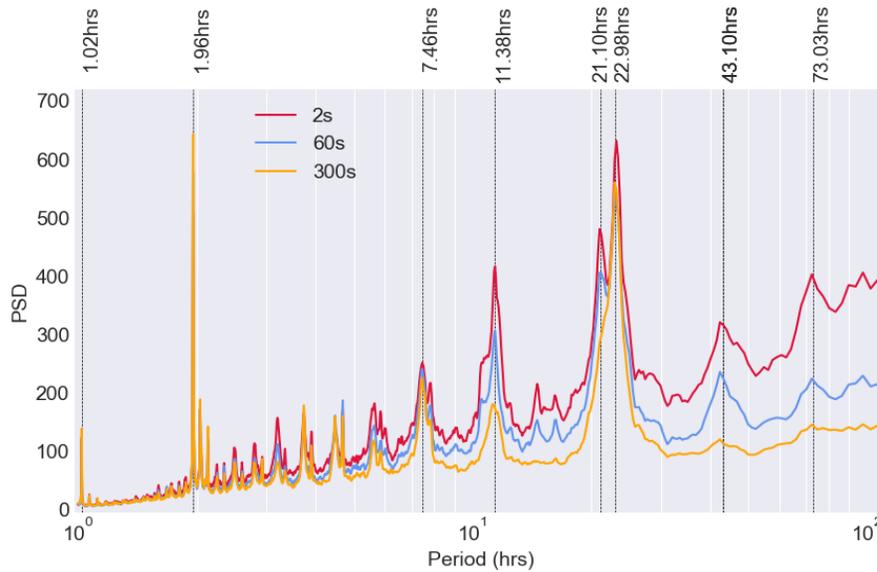
327

328 **Figure 7.** Top, year long spectrogram of positions for the east component of station
 329 BILS (Figure 2). For ease of interpretation we also plot the time series of power for 3
 330 selected periods, 2s, 60s, and 300s. Bottom, same as the top but for a shorter time
 331 span of only 1 week. The time periods labeled P1, P2, P3 enclosed in the rectangles

332 correspond to periods of high, medium, and low noise. Time series for these periods
 333 are in Figure 8.



334
 335
 336 **Figure 8.** Left: Example 10 minute long time series for periods of high, medium, and
 337 low noise for the east component of station BILS. The time periods are highlighted as
 338 P1, P2, and P3 in the spectrogram on Figure 7. Right: PPSD for station BILS, (same
 339 as Figure 3) with spectra for noise at time periods P1, P2, and P3. The dashed lines
 340 are used as a reference and denote the power of a time series of Gaussian white
 341 noise with specified standard deviation



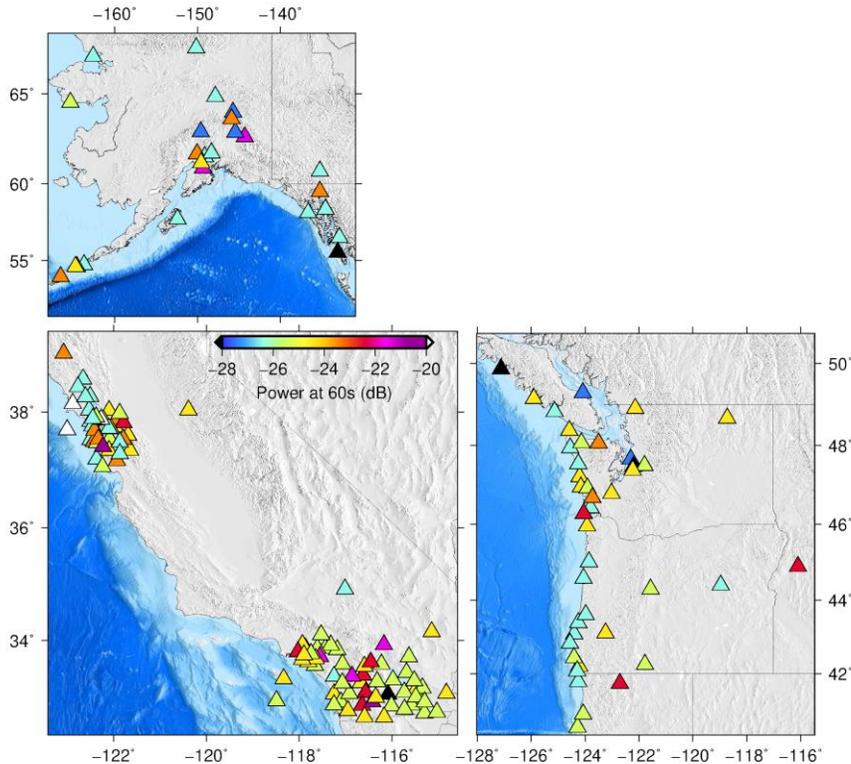
342
 343 **Figure 9.** Stacked spectra of the time series of power of the noise at 2s, 60s, and
 344 300s periods (see Figure 7). The individual spectra for each site are calculated for the
 345 entire time span and then all the sites are averaged together to create the stack.

346
 347 **4. Discussion**

348
 349 4.1 On the characteristics of the noise

350

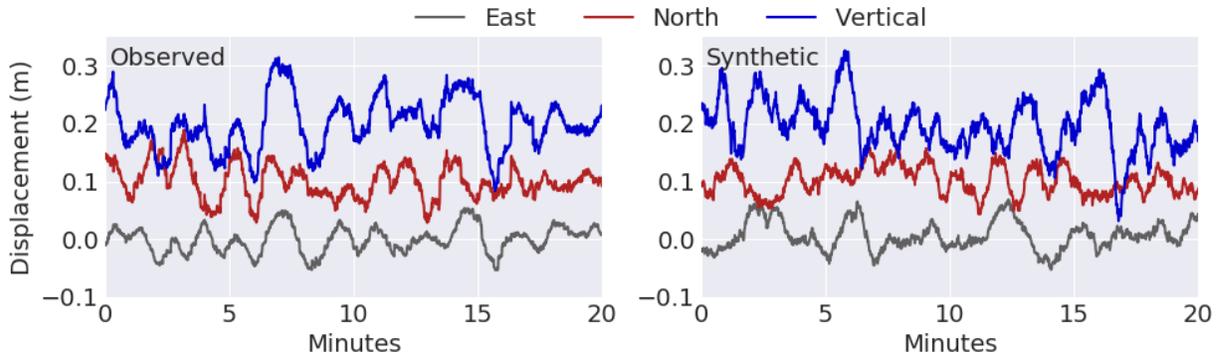
351 Langbein & Bock [2004] and Genrich & Bock [2006] first analyzed the noise behavior of
 352 HR-GNSS positions obtained from relative positioning (where the positions are with respect
 353 to a reference site). Those studies found that noise has a characteristic “dam” profile, with
 354 approximately flat power at long periods and linearly decaying (in a log-log sense) power at
 355 higher frequencies. In this study we find that positions obtained from PPP are consistent
 356 with these earlier findings. At periods longer than ~200s power is mostly flat, suggesting
 357 mostly uncorrelated positions, with power decaying with a slope of -2 at shorter periods
 358 (e.g. Figures 5 and 6). This power of 2 decay is characteristic of a random walk process
 359 (e.g. Agnew 1992). Both the slope and the location of the spectral “corner” are consistent
 360 with earlier findings from Genrich & Bock [2006] who analyzed instantaneous relative
 361 positions for three baselines in Southern California. It suggests that the primary source in
 362 the short period band up to ~200s is a combination of the troposphere and multipath. At
 363 very short periods (<5s) there is an indication that the spectra are beginning to flatten, this
 364 too would be consistent with Genrich & Bock [2006] and Bock et al. [2011] who observed
 365 mostly white noise in 50Hz sampled GNSS at periods shorter than 1-2Hz. This white noise
 366 behavior is indicative that the noise sources at these higher frequencies are uncorrelated.



367 **Figure 10.** Distribution of noise at 60s period in the California, Pacific
 368 Northwest and Alaska regions.
 369

370
 371 The long term behavior of the noise in Figure 7 is interesting. The spectrograms show a
 372 periodic variability in the noise levels which is punctuated by irregularly spaced short
 373 intervals of time where there is a wholesale increases or decrease of noise. The time series
 374 of power in Figure 7 were collected at 2s, 60s, and 300s, which alternatively correspond to

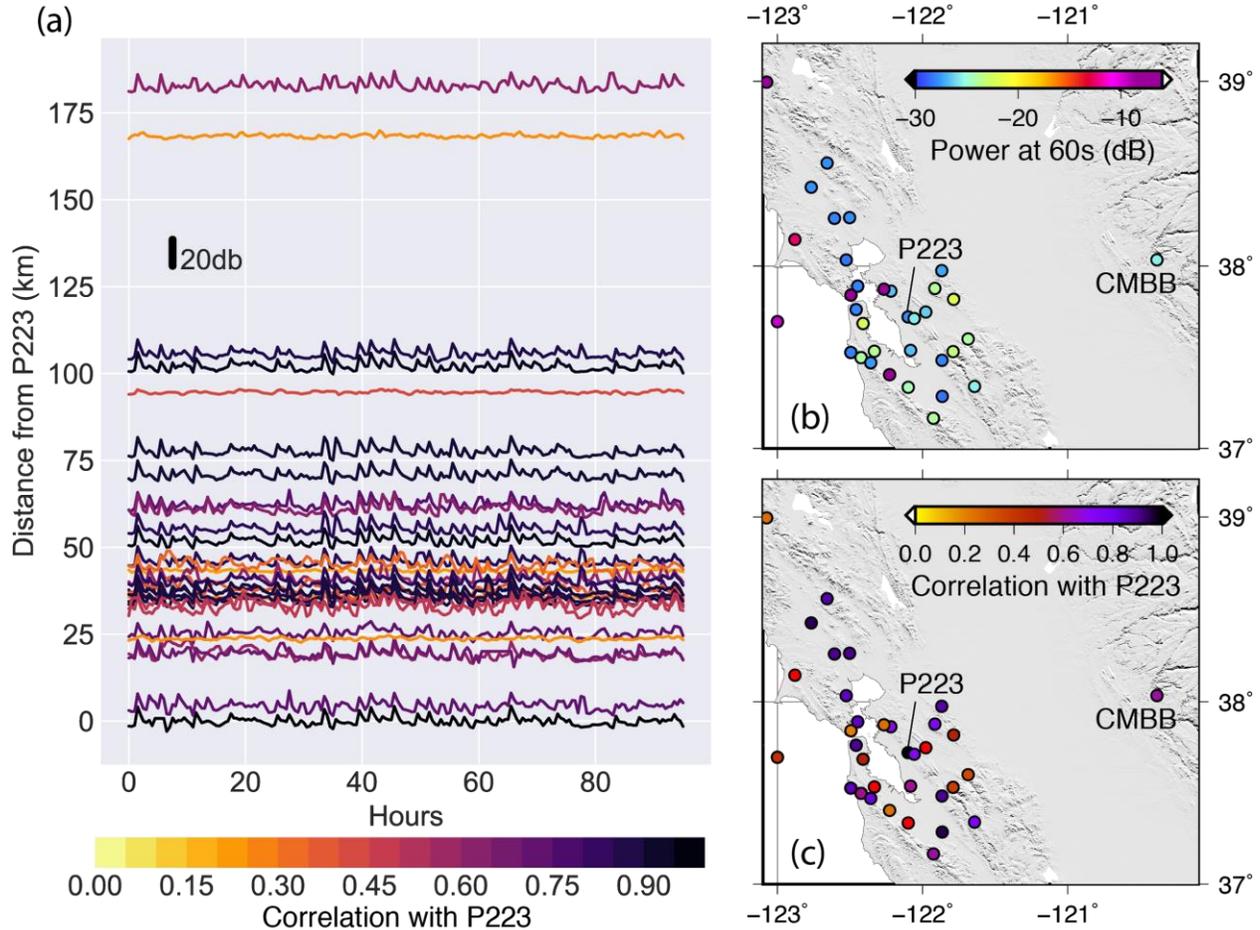
375 the short period somewhat flattened part of the spectrum, the linearly decaying part of the
 376 spectrum, and the long period approximately flat part of the spectrum. The changes in
 377 power at all of these periods co-vary, even during periods (such as P1 in Figure 7) when
 378 there are large increases in noise. This is perhaps unsurprising, Figure 8 shows that the
 379 noise increase is manifested as several step-like jumps, likely from errors in the ambiguity
 380 resolution procedure. We also see a gradual decay after each step offset which is
 381 characteristic of “re-convergence” after a cycle slip (e.g. Geng et al., 2013).
 382



383
 384 **Figure 11**, 20 minutes of noise observed at station BILS (Figure 2) and 20 minutes
 385 of synthetic noise generated using the median noise model from Figure 5. The time
 386 series for each direction of motion are offset for clarity.
 387

388
 389 The irregular distribution of station noise in Figure 10 is somewhat surprising and will
 390 warrant further study. A priori one would expect a geographic correlation between the noise
 391 levels and a number of potential candidate parameters. For example it is well known that
 392 the geometry of the constellation of GPS satellites is less favorable for positioning at higher
 393 latitudes. Similarly the ionosphere should be more active as one approaches the poles. Yet
 394 we do not observe a systematic degradation of the northern sites. Similarly, one would
 395 expect that multipath would correlate strongly with the noise performance of the stations.
 396 We also do not see a correlation between noise power and the signal to noise ratio in the
 397 L1 and L2 frequencies. We explore this further in Figure 12. Here we plot the time series of
 398 power at 60s period for a 4 day period for all sites in the San Francisco Bay Area cluster.
 399 We see clearly that in this limited geographic region over length scales of ~100km the noise
 400 at many sites is highly correlated. Both episodes of elevated and reduced noise occur close
 401 together in time between many sites. We calculate the correlation coefficient of all the time
 402 series to an arbitrary site in the middle of the cluster (P223) and indeed we find high
 403 correlation (>0.75) between many of the stations. This is true even for station CMBB which
 404 is 175km from P223 and still exhibits a high (0.7) correlation coefficient. In Figure 12 we
 405 also show the median power at 60s period for the same 4 day period. We see clearly that
 406 the stations that do not follow the same regional variation in noise are those with highest
 407 power. These noisy sites have a completely different evolution of noise with time.
 408

409 Figures 10 and 12 show that a likely explanation for the noise behavior at a particular site is
 410 a weighted sum of many factors. The strong correlations at $\sim 100\text{km}$ length scales are
 411 evidence of regional effects. This includes period of disturbances in both the ionospheric
 412 and troposphere which would affect stations over regions of this size. Similarly drifts in
 413 clocks and orbits will have a strong regional correlation. However, that the absolute level of
 414 the noise at stations analyzed in this study exhibits poor correlation with obvious
 415 geographic features strongly suggests that this is an effect local to each site. The quality of
 416 the monument, and the environment (vegetation, snow, buildings, other microwave
 417 equipment) surrounding each site is highly heterogeneous and can have an outsized effect
 418 in the positioning quality raising the noise floor substantially.



419 **Figure 12.** (a) Time series of power at 60s period for the 4 day time span from
 420 June 26th to June 30th 2019 for station in the San Francisco Bay Area cluster.
 421 The time series are colored by the correlation coefficient to a reference site
 422 inside the cluster (station P223) (b) Locations of the stations in the cluster and
 423 median power at 60s period for the selected time span. (c) Correlation between
 424 all sites to reference station P223.
 425
 426

427 Finally we note that there are other noise sources which will be specific to the positioning
 428 algorithm being used. For example the $\sim 2\text{hr}$ peak (Figure 9) is likely due to the frequency
 429 with which the orbital parameters in the broadcast ephemeris are updated by IGS. These

430 updated values are used in the FastLane processing scheme and introduce a regular
431 periodic behavior. Indeed an important point we stress is that the overall noise behavior
432 and shape of the spectra should roughly follow the “dam” profile irrespective of the
433 positioning algorithm being used. However the details of the noise behavior will be strongly
434 influenced by not just the traditional sources of noise but also by the processing strategy.
435 Deviations from the noise behavior we detail here should not be unexpected.

436

437 4.2 Implications for positioning algorithms

438

439 The noise models we propose are useful signposts which can be used to compare against
440 as improvements to GNSS positioning technologies are developed. For example broadcast
441 of new frequencies by GNSS satellites [Geng & Bock, 2013; Zhao et al., 2015], and
442 positioning strategies that harness multiple constellations (“true GNSS”) promise to provide
443 substantial improvements and reductions in noise [Odolinski et al., 2015; Geng et al., 2016;
444 Geng et al., 2019b]. As this technology is incorporated into permanent monitoring networks
445 it can be evaluated by comparison to established baselines of noise behavior.

446

447 4.3 Comparison to seismological noise

448

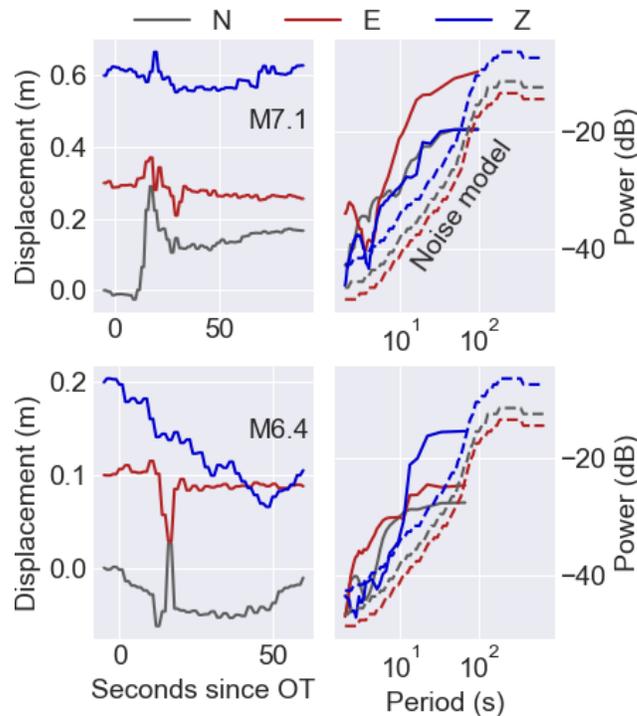
449 The concepts behind this study were inspired by the techniques proposed by McNamara &
450 Buland [2004] who carried out a similar analysis for noise at seismic sites in the continental
451 United States and established its systematic behavior. Because use of HR-GNSS is
452 becoming widespread in monitoring efforts in seismology we have attempted to establish a
453 similar baseline of behavior here. However, some critical conceptual differences warrant a
454 few comments. The background seismic noise observed at broadband sites on-land are
455 actual vibrations of the ground whose source is the interactions of the oceans with the
456 near-shore solid Earth (e.g. Longuet Higgins, 1950). This noise is of comparatively very
457 small amplitude, at periods of 1s, for example, it is expected to have a power of -170 to -
458 130 dB [Peterson, 1993]. This is several orders of magnitude below what we measure from
459 GNSS (-55dB, Figure 5). We emphasize that the source of noise in GNSS positions have
460 nothing to do with actual high-frequency motions of the ground. It is true that longer period
461 deformation of the Earth such as that induced by tides can have mm to sub-mm amplitudes
462 (e.g. Agnew, 2010) however this is outside the frequency band of interest to monitoring
463 large earthquakes. Rather, most of the noise comes from the variable delays to the
464 microwave satellite signals introduced by the troposphere and ionosphere and spurious
465 reflections (multipath) of the microwaves off of the surrounding terrain which occlude the
466 main arrival to the GNSS antenna. Another large source of noise is imperfect knowledge in
467 real-time of the satellite clocks and orbits. Long period noise induced by constellation
468 geometry is highly repeatable and can be reduced through sidereal filtering [Larson et al.,
469 2007]. Assuming that technical improvements in mitigating these noise sources are
470 possible this large gap (~70dB) between background seismic noise and current GNSS
471 noise suggests that the lower bound of what could be observed lies far beyond what is
472 possible now. There is essentially unbounded room for improvement.

473

474 4.4 Implications for seismic monitoring

475
 476 Previous studies of the relative amplitudes of ground displacements at regional distances
 477 from medium to large events [Crowell et al., 2013, Melgar et al., 2015, Ruhl et al., 2018]
 478 suggest that, in order for GNSS time series to be of use for monitoring, precision of a few
 479 centimeters is necessary. For example, if the precision were relatively poor, say 10cm, the
 480 peak ground displacement scaling laws of Melgar et al. [2015] predict that an M7
 481 earthquake would be visible to any site within 91km. That distance grows to 462km and
 482 1628km for M8 and M9 earthquakes respectively. The aggregate PPSDs in Figures 5 and 6
 483 then suggest that the current precision achieved by the real-time GNSS solutions is
 484 sufficient for monitoring large events.

485
 486 Figure 13 shows an example of the potential performance. There we plot the three
 487 component HR-GNSS displacements for station CCCC which was processed in real-time
 488 with Fastlane and recorded both the M6.4 and M7.1 Ridgecrest CA earthquakes at 35 and
 489 50km from the source [Melgar et al., 2019; Goldberg et al, 2019]. We also plot the spectra
 490 for the waveforms and the median noise models. This shows clearly that the waveforms are
 491 reliable. Melgar et al. [2019] showed that while there were some small but appreciable
 492 differences between the real-time and post-processed high-rate solutions the features of
 493 the waveforms most used in monitoring remained consistent in both sets of solutions.
 494



495
 496 **Figure 13.** Real-time three-component displacements at station CCCC during the
 497 2019 M6.4 and M7.1 Ridgecrest CA earthquakes plotted as seconds since the

498 *earthquake origin time (OT). Also shown are the spectra of the waveforms*
499 *compared to the median noise model*

500

501 Another important use of the work discussed here for monitoring is in making objective and
502 automated assessments of the station positioning quality. Figure 8 exemplifies how an
503 otherwise well-behaved station can, for limited periods of time, have elevated noise levels
504 which can have detrimental effects on any algorithm using it to model an earthquake
505 source. Monitoring agencies can use either the global noise model, or a station by station
506 noise model, and set percentile cutoffs, perhaps at a few selected periods. If the noise rises
507 above that threshold for some period of time the station can be quarantined or “black listed”
508 so that it doesn’t contribute solutions to a source modeling algorithm should an earthquake
509 occur in that time. Later as the station noise drops to an acceptable level it can be removed
510 from the black list. Similarly sites that are routinely above some threshold level will likely
511 need to be serviced or altogether removed from contribution to any real-time monitoring
512 effort.

513

514 For the Fastlane algorithm in particular one important challenge remains as it continues to
515 contribute solutions to monitoring agencies. The large outliers seen in Figure 2 are not the
516 norm (e.g. Figure 4) but they are large enough that should they occur during an earthquake
517 they could introduce significant errors into the modeling. This is problem has been noted in
518 real-time monitoring efforts elsewhere [Kawamoto et al., 2017]. This in general will not be
519 an issue for the computation of coseismic offsets, as a moving average or median filters
520 can be employed (i.e., Crowell et al. [2016]), however, for PGD scaling, significant outliers
521 or cycle slips can influence the derived magnitude estimates. During the Ridgecrest
522 earthquakes there were no occurrences of this in any of the real-time waveforms. However
523 continued effort in making the positioning strategy more robust is ongoing.

524

525 **5. Conclusions**

526

527 Large earthquakes are difficult to model in real-time with traditional inertial seismic
528 measurements. Several algorithms that leverage high-rate RT-GNSS positions have been
529 proposed and it has been shown that they can supplement the earthquake monitoring
530 effort. However, analyses of the long-term noise behavior of high-rate RT-GNSS positions,
531 which are important to understand how the data can be used operationally by monitoring
532 agencies, have been limited to just a few sites and to short time spans. Here we have
533 shown results from an analysis of the noise characteristics of one year of positions at 213
534 RT-GNSS sites spanning a large geographic region from Southern California to Alaska. We
535 have characterized the noise and proposed several reference noise models which can be
536 used as baselines to compare against as technological improvements allow for higher
537 precision solutions. We have also shown how to use the reference noise models to
538 generate realistic synthetic noise that can be used in simulations of HR-GNSS waveforms.
539 Additionally, we find that while variations in the noise have a strong spatial correlation the
540 absolute level of noise at a site does not. This is evidence that local effects
541 (monumentation, station conditions, multipath etc.) likely dominate the noise behavior.

542 Further, we have shown how this noise analysis can be used in a dynamic quality control to
543 determine which sites should or should not contribute positions to an earthquake modeling
544 algorithm at a particular moment in time. Overall, while there remain important
545 improvements yet to be made, such as reducing the number of outliers, we find that the
546 present quality of real-time HR-GNSS waveforms is more than sufficient for monitoring
547 large earthquakes.

548

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550

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562 Data Center (NCEDC), doi: 10.7932/NCEDC. The synthetic GNSS noise generation code
563 can be found at, <https://github.com/dmelgarm/MudPy> and a tutorial on how to run the code
564 is at <https://github.com/dmelgarm/MudPy/wiki/Generate-synthetic-HR-GNSS-noise>

565

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