Hydroclimatic Variability in Africa's Transboundary River Basin Using +50 Years of Extended Terrestrial Water Storage

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Abstract

Information concerning the types, patterns, and intensity of hydroclimatic variability is needed in various sectors, including water resources planning and climate change adaptation, among numerous others. In transboundary river basins, the nature of hydroclimatic variability frequently is a critical input in the treaties governing water resources allocation to competing parties and sectors. Yet, the pattern of hydroclimatic variability across Africa's Transboundary Rivers Basins (ATRB) remains poorly investigated owing primarily to lack of access to required data. To the extent that such studies exist, they often have been conducted at different times using different data sets and reference periods, making comparisons across the continent difficult. In this paper, we make use of NASA's Gravity Recovery and Climate Experiment (GRACE) satellite data to extend the Terrestrial Water Storage (TWS) +50 years prior GRACE era using the unique Generalized Additive Model for Location Scale and Shape (GAMLSS) approach on the Global Land Data Assimilation System Version 2 data (GLDAS V2). The results revealed a downward trend of the TWS over Africa with a decrease rate of 0.14 cm/yr. The spatial patterns of TWS in ATRB showed a significant decreasing trend for Nile, Niger, Chad, Volta, and Congo River Basins, compared to the insignificant trends for Zambezi, Okavango, Limpopo and Orange River Basins. The spatial trends in annual TWS is primarily related to the regional variation in the precipitation trends. The largest negative trend in precipitation were observed over West Africa and Sahel region. The dry trend over south Africa were intervened by wet records. The nature of these hydroclimatic patterns are explained by the significant reduction in the total precipitation and the increasing demands on water resources.

Hydroclimatic Variability in Africa's Transboundary River Basins Using a +50 Years of BINGHAMTO **Extended Terrestrial Water Storage** UNIVERSITY

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OBJECTIVES

Information concerning the types, patterns, and intensity of hydroclimatic variability is needed in various sectors, including water resources planning and climate change adaptation, among numerous others.

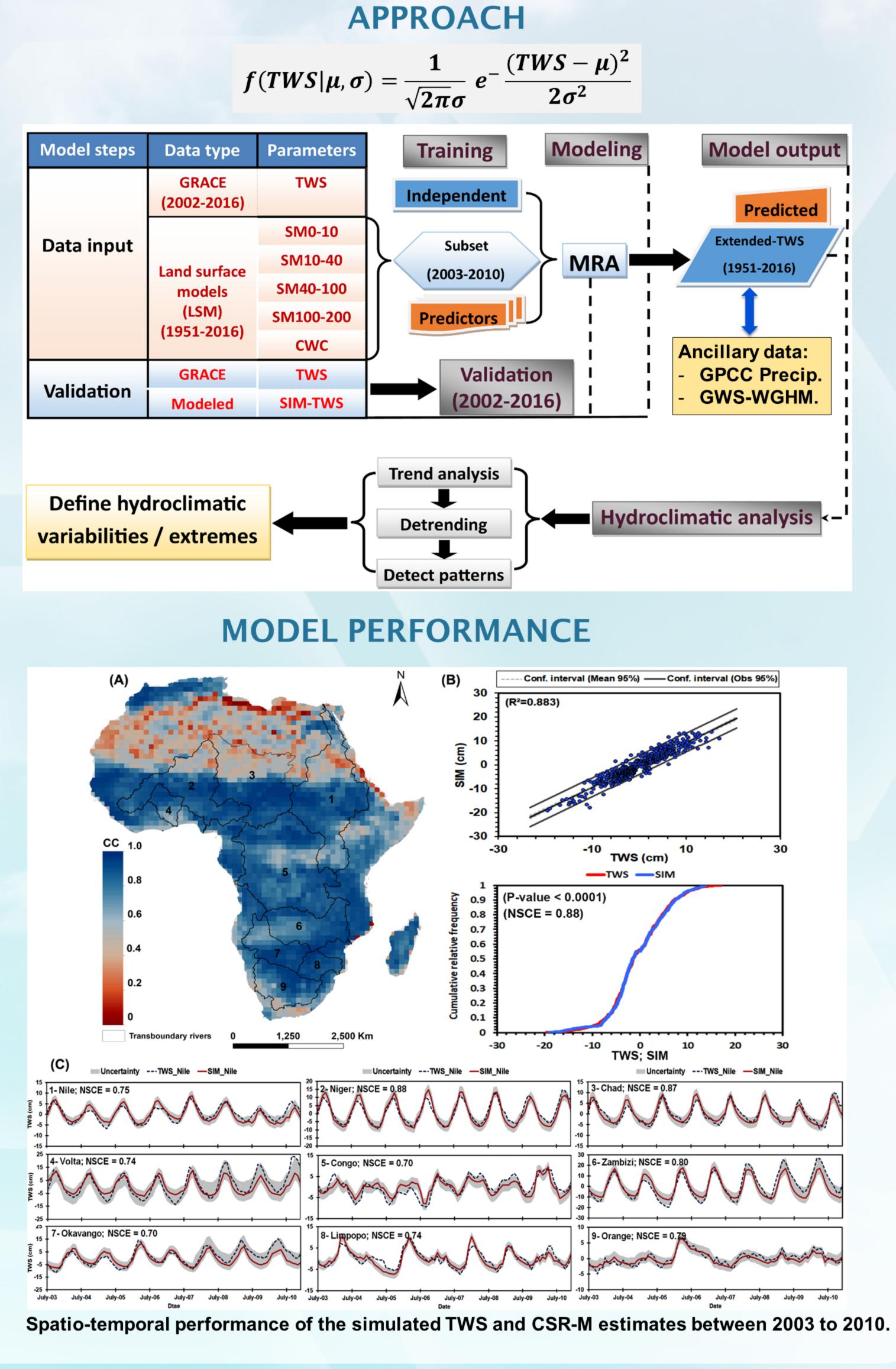
In transboundary river basins (TRBs), the nature of hydroclimatic variability frequently is a critical input in the treaties governing water resources allocation to competing parties and sectors.

In this paper, we make use of GRACE data to extend the TWS a +50 years prior GRACE era using GAMLSS approach on GLDAS-Noah V2 estimates.

The hydroclimatic nature of extended TWS across nine African TRBs (Nile, Niger, Chad, Volta, Congo, Zambezi, Okavango, Limpopo and Orange) were explained using trends, abrupt shifts, as well as the dominant modes of Basic info summary of TRBs. variability using wavelet analysis.



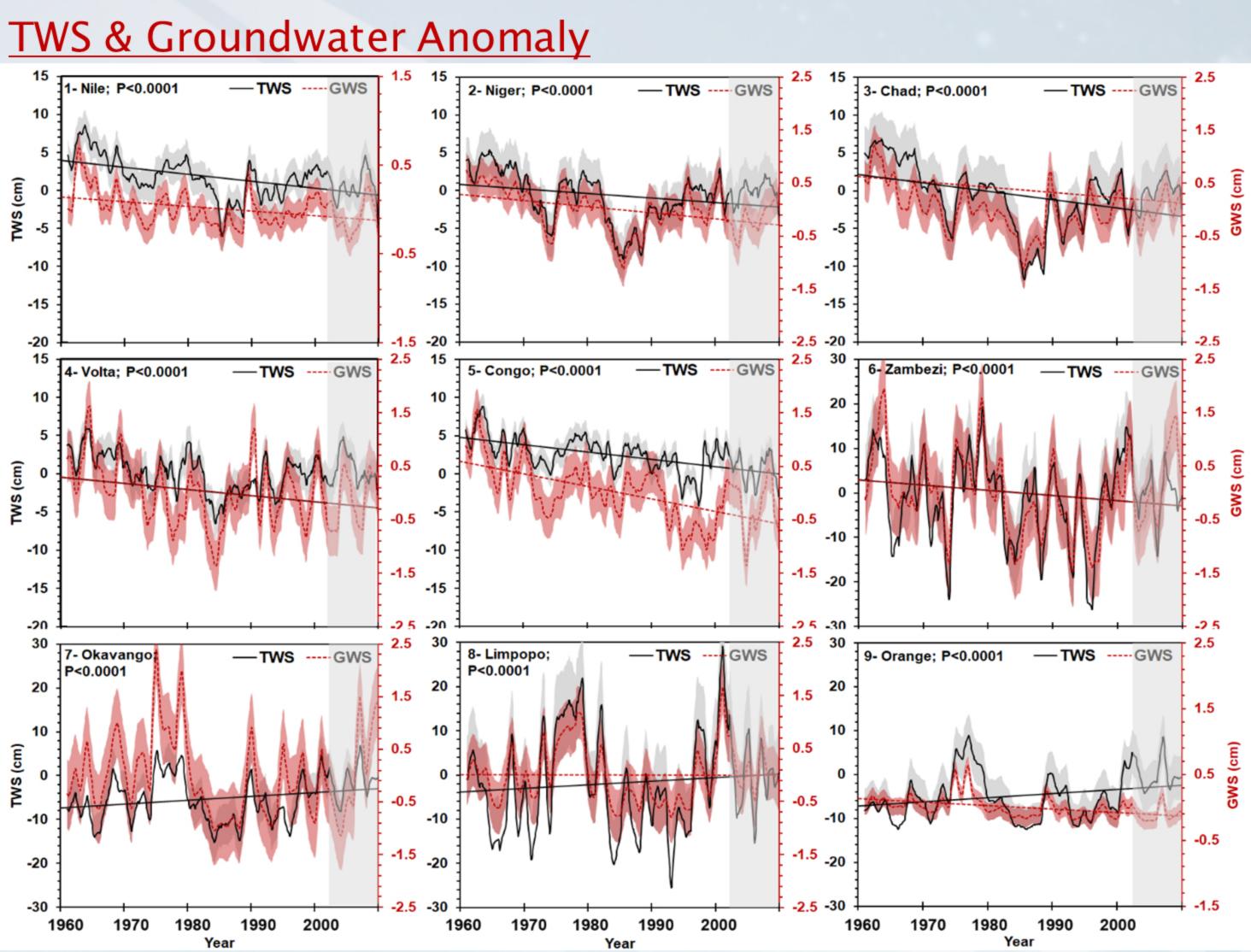
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Sele Ma Basin Nile Congo Niger Chad	Area (km ²) 3,254,555 4,014,500 2,117,700 2,434,000	Kiometres	Q (m3/s) 2,800 41,200 5,590 1,290	TRRBS Shared by Eritrea, Sudan, Ethiopia, Egypt, Tanzania, Congo DR, Rwanda, South Sudan, Uganda, Burundi, Kenya Angola, Tanzania, Central African Republic, Congo Republic, Congo DR, Rwanda, Gabon, South Sudan, Burundi, Cameroon, Malawi, Zambia Mauritania, Algeria, Chad, Guinea, Nigeria, Sierra Leone, Ivory Coast, Mali, Niger, Cameroon, Burkina Faso, Benin Chad, Libya, Sudan, Central African Republic, Algeria, Niger, Cameroon, Nigeria	(2015) 300 77 130 30	SA H SA SH
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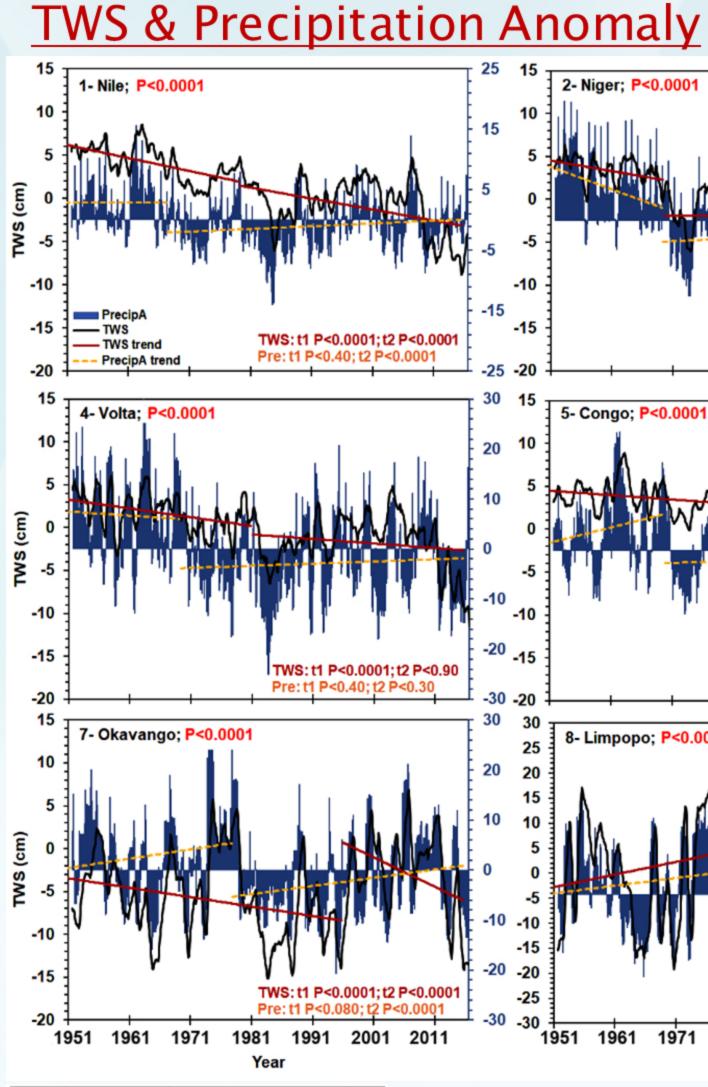
Emad Hasan^{*}, Aondover Tarhule, Pierre-Emmanuel Kirstetter, Joseph Terzungwe Zume *Department of Geography, State University of New York (SUNY) at Binghamton Binghamton, NY, 13902, USA. Email: emad.hasan@binghamton.edu

Using standard model evaluation criteria including, the Pearson correlation coefficient (R2), cumulative distribution function (CDF), the Nash Sutcliffe efficiency (NSCE) and model uncertainty, the results indicate a strong spatial agreement between the observed CSR-M and simulated TWS. Across the nine River Basins, a correlation coefficient of (R²= 0.88) is observed. A good density distribution and fitting are indicated by the cumulative density function (CDF) distribution. Time-series plots of the training period between the CSR-M and the simulated TWS between July-2003 and December-2010, the plot shows a good consistency of the model as indicated by the Nash–Sutcliffe efficiency test, the 95% uncertainty bounds are plotted.

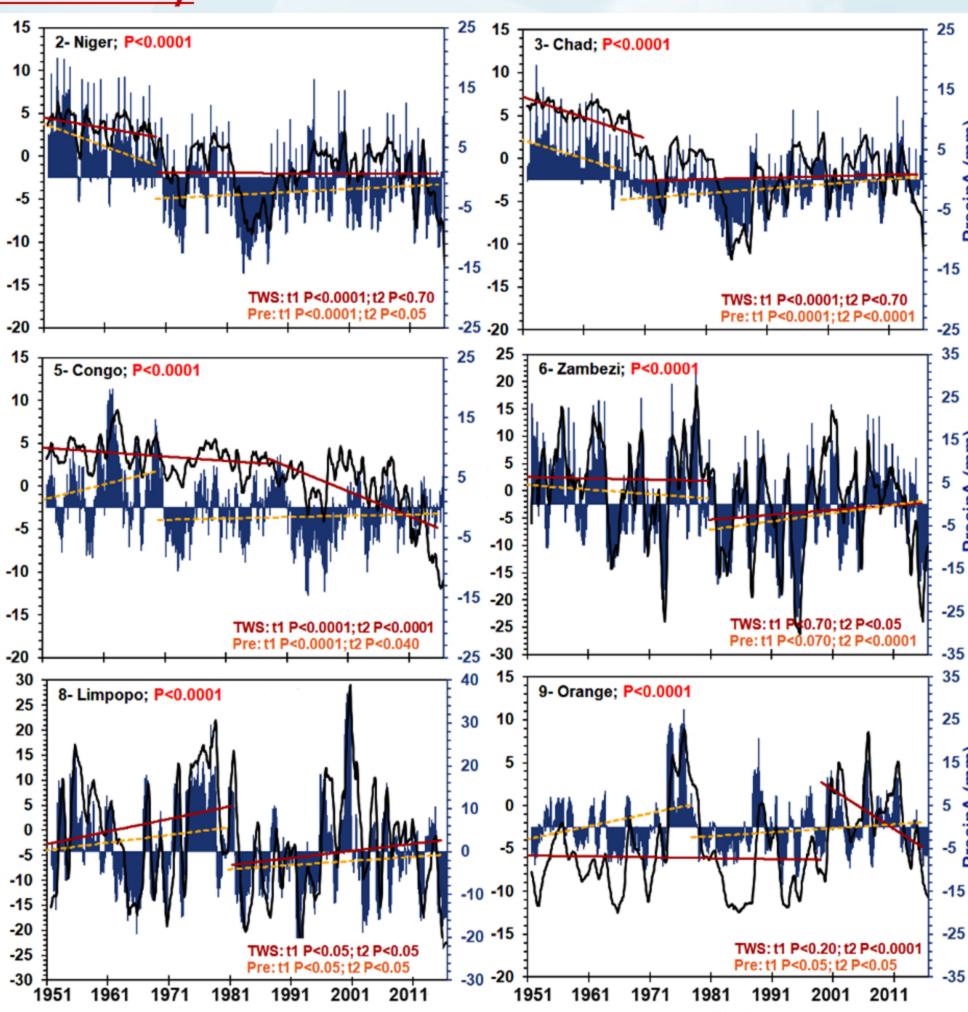
HYDROLOGICAL ESTIMATES



There is overall good agreement of ~80% correlation between the extended estimated TWS and

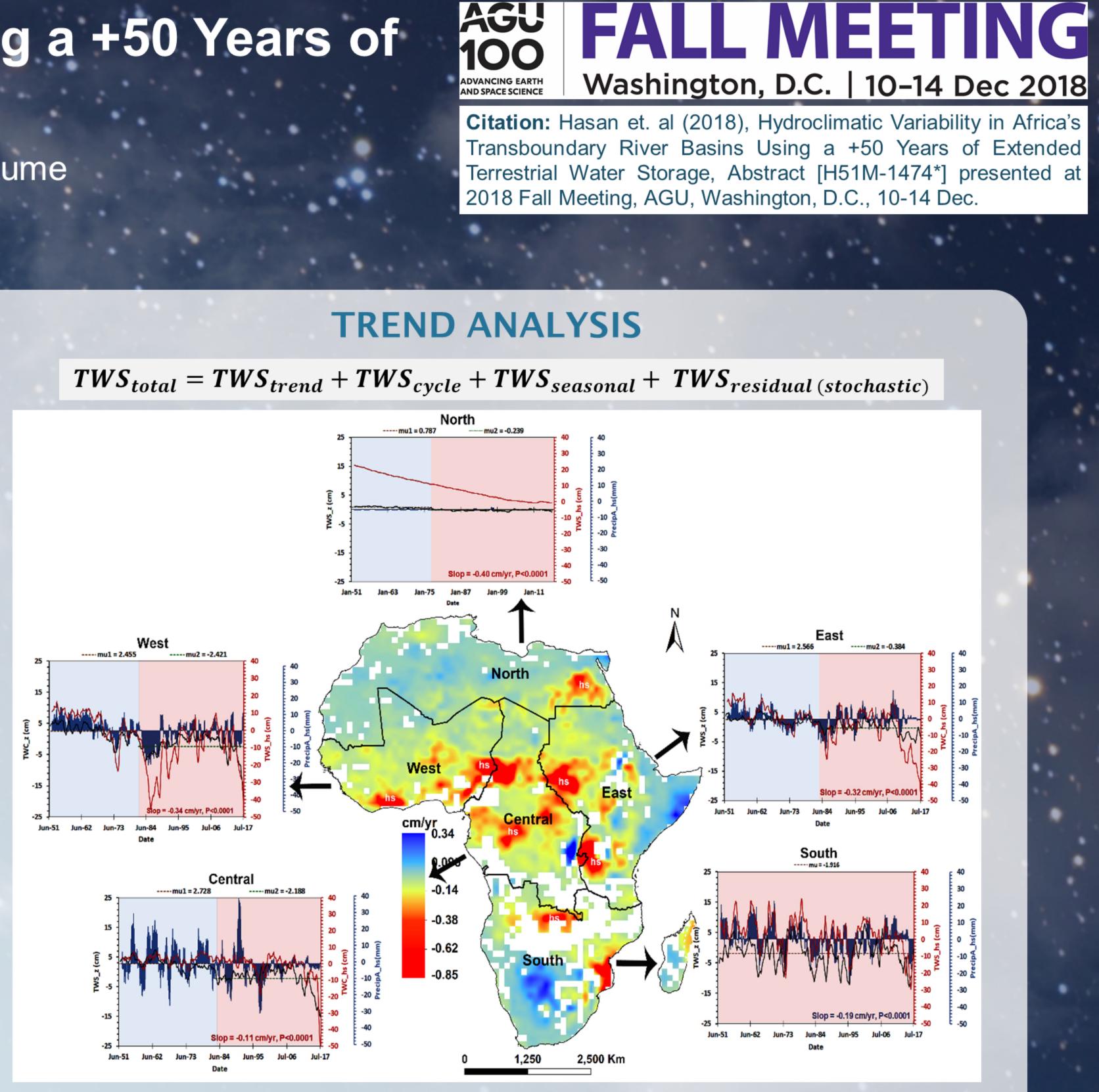


	Change point year		Trend				
Basin	PrecipA	TWS	PrecipA_t1	TWS_t1	PrecipA_t2	TWS_t2	
					٨	1	
	10.00	1000		- ↓			
Nile	1969	1980				•	
Niger	1970	1970	V		•		
					A		
Chad	1968	1972	•	•			
Volta	1970	1982		- ↓			
Tona	1770	1962		, i			
			T				
Congo	1971	1989		•		•	
Zambezi	1982	1982	•				
Lamotzi	1702	1702				- i-	
			•	i	•	i	
Okavango	1979	1997		•		•	
			٨		٨		
Timmono	1982	1982			:		
Limpopo	1982	1982		-			
			A		A		
Orange	1979	2000				•	

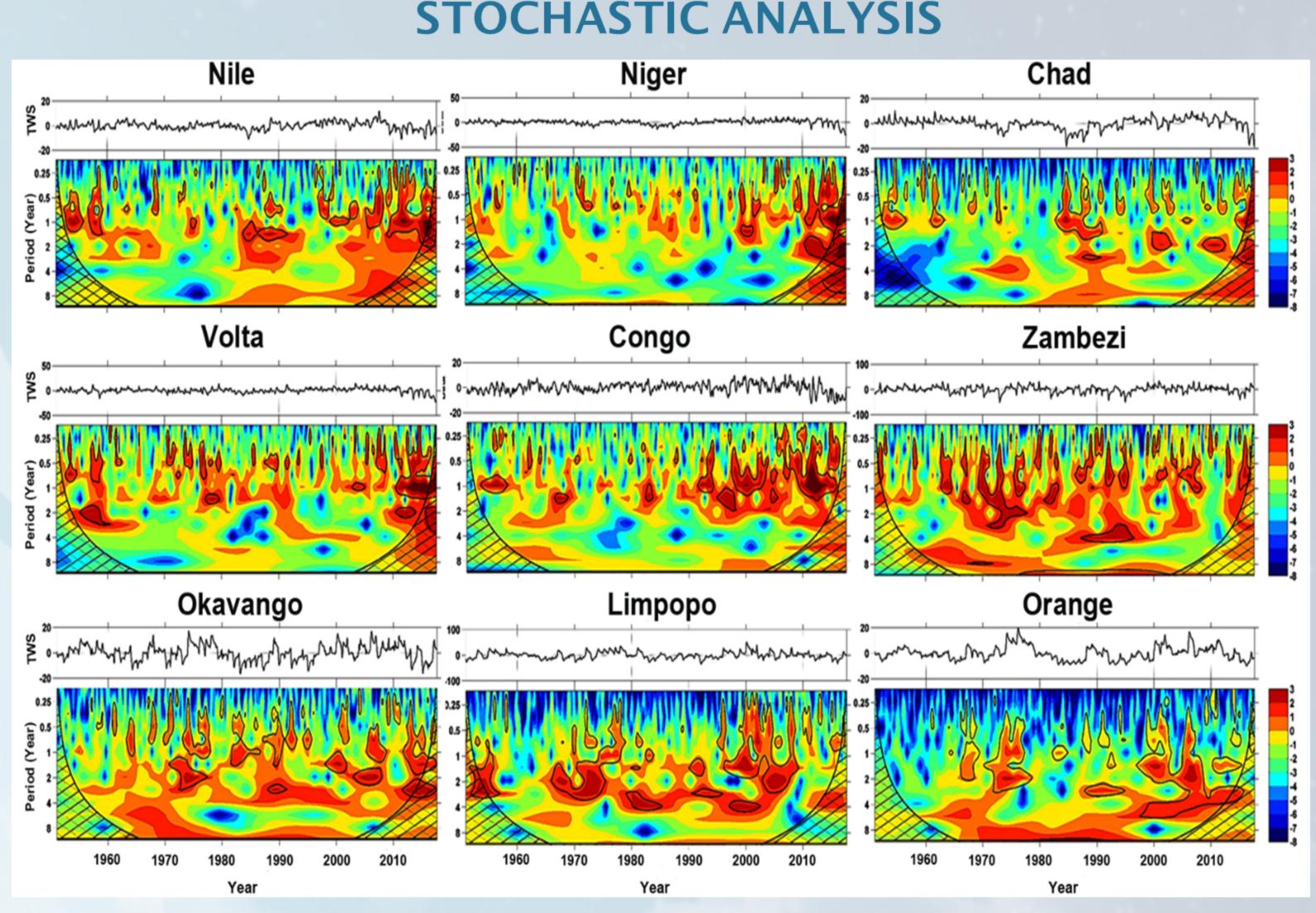


Extended-TWS and precipitation anomalies across the nine Africa's TRBs between 1951 to 2016. The coherence in the seasonality between TWS and precipitation anomalies is indicated by P-Value (in red font). The attached table is summarizing the direction of the trend before and after the change point of TWS and precipitation anomalies. The table only display the significant trend direction at a 95%.





Spatio-temporal variation of the extend-TWS trend across Africa from 1951 to 2016. The trend map shows a number of hotspots that are marked by (hs), which are experienced substantial increase/decrease in the overall TWS trend. A time series average of TWS across five major geographic areas is also indicated. The results show a regional positive TWS mean between 1951 and early 80s (blue space) for North, East, West and Central zones, then a significant negative TWS trend is observed between late 1980s to 2016.



CWT for extended-TWS for the nine river basins. The power spectrum of the Nile, Niger, Chad, Volta and Congo river basins fluctuates in cycles between 1 to 2 years from 1951 to the 1980s. After 1980s the cycles extended up to a 4-year cycles especially during the late 2000s. The Zambezi, Okavango, Limpopo and Orange river basins fluctuate in cycles from 4 to 6 year cycles, where major cycles are shown from 1960s to late 2000s.

