

# Microlite Size Distributions and P-T-t-x (H<sub>2</sub>O) constraints of Central Plateau tephtras, New Zealand: implications for magma ascent processes of explosive eruptions

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## Abstract

Crystals within erupted volcanic rocks record geochemical and textural signatures during magmatic evolution prior to the onset of eruptions. Growth times of microlites can be derived through Crystal Size Distribution (CSD) analysis combined with well-constrained microlite growth rates, yielding petrologically-determined magma ascent timescales. Our newly developed, machine learning image processing scheme allows for the rapid generation of CSD, saving many hours of processing time, which previously involved hand-drawing the outer margins of crystals. For the present study, we examined a range of andesitic tephtras from the Tongariro Volcanic Centre, New Zealand. A total of 228 plagioclase and pyroxene microlites CSDs were generated from individual tephtra shards. All combined pyroxene and plagioclase microlite CSDs exhibit concave-up shapes, and similar intercepts and slopes at the smallest sizes. This implies similar growth durations of the smallest microlites of  $15 \pm 9$  to  $28 \pm 15$  ( $2\sigma$ ) hours, regardless of the eruptive style or source, using an orthopyroxene microlite growth rate constrained from one of the samples. The orthopyroxene thermometer and the plagioclase hygrometer reveal the magmas were erupted at  $\sim 1079$  to  $1149$  ( $\pm 39$  SEE), and H<sub>2</sub>O contents ranging from 0-0.4 to 0-1.7 wt.% (95% confidence maxima). In the absence of CO<sub>2</sub>, these results indicate shallow H<sub>2</sub>O exsolution pressures of  $< 240$  bars, using a recent H<sub>2</sub>O-CO<sub>2</sub> solubility model. Given the microlite residence times, shallow H<sub>2</sub>O exsolution driving microlite growth is inconsistent with the explosivity of the eruptions. Instead, our data suggest that the melts either carried large amounts of CO<sub>2</sub>, triggering earlier degassing of volatiles including H<sub>2</sub>O, or that microlite crystallisation began prior to degassing. Ongoing work investigates the H<sub>2</sub>O and CO<sub>2</sub> contents hosted by melt inclusions in phenocrysts and microphenocrysts in these tephtras to provide constraints on magma ascent rates, with implications for hazard characterization and mitigation.

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## Introduction

- Taupo Volcanic Zone (TVZ): southern end of the Tonga-Kermadec arc-back-arc system affected by oblique subduction, clockwise rotation and extension e.g., [1] (Fig. 1)
- Tongariro Volcanic Centre (TgVC): andesitic compound structure at the southern end of the Taupo Volcanic Zone
- Multiple dispersed eruptive centres and vents producing events of various eruptive styles
- Microlite crystallisation due to decompression-induced degassing (e.g., [3] to [6]), change in eruptive style shortly before the eruption [7]
- CSD → tool to assess ascent processes, and residence times when growth rate can be constrained (e.g. [8] to [11])
- **Aim:** use microlite textures to constrain magma ascent processes and timescales

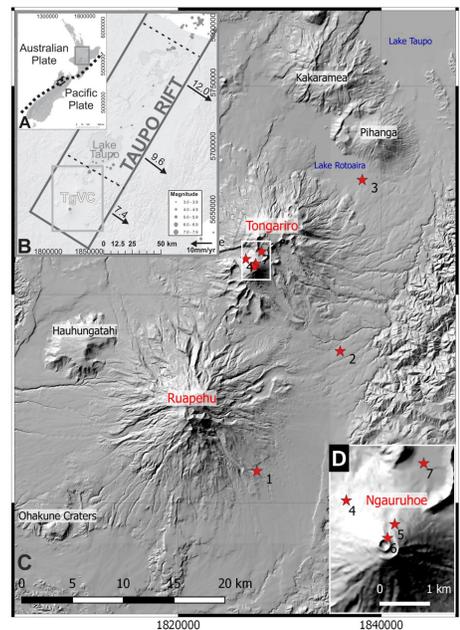


Figure 1: (A) Map of the subduction and location of the TVZ, and (B) constrain magma ascent location of the TgVC [2]. (C) (D) Map of the area with sample locations.

## Methods and Materials

- BSE images of glass shards using FE-SEM
- Semi-automatic segmentation of plg and px microlites for CSD [12]
- Generation of CSDs using Csd\_slice [13] and CSDCorrections [14]
- EPMA glass and microlite analyses
- Iteration of the plagioclase hygrometer [15], orthopyroxene thermo-barometer [16]

Table: description of tephra samples used for this research

ERUPTIVE SEQUENCE	AGE/DATE	MAIN ERUPTION STYLE	COMMENTS	REFERENCE
Ruapehu 1995 -1996	1995 -1996	Strombolian, phreatomagmatic and sub-Plinian	1995 sample: sub-Plinian 1996 sample : ash and bomb explosion	[17]
Ngauruhoe 1972 - 1975	1972 - 1975	Strombolian and Vulcanian	Steaming, block and ash ejections and flows, and eruptive columns	[18] & [19]
Tufa Trig	0 - 1.7 ka ago	Small volume Strombolian and phreatomagmatic	Focusing on Tf8, Tf13 & Tf14, Strombolian, rarely sub-Plinian	[20] to [23]
Mangatawai	1.7 - 3.5 ka ago	Small explosive up to sub-Plinian, mostly Vulcanian	Frequent, intermittent, very thin deposits containing beech leaves	[18], [22] & [24]
Mangamate	11 - 12 ka ago	Plinian	Large eruptions over a short period and from several vents along a fissure	[22] & [25]

## Results

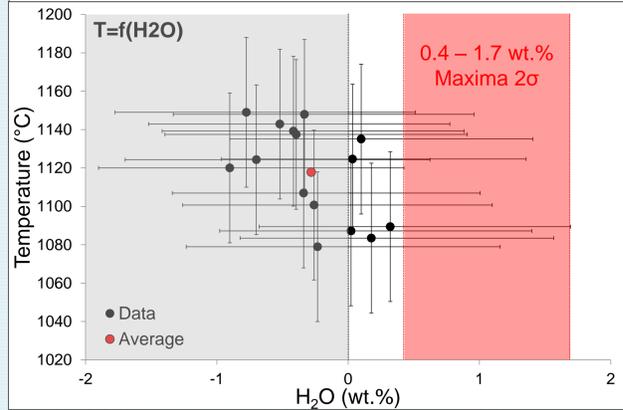
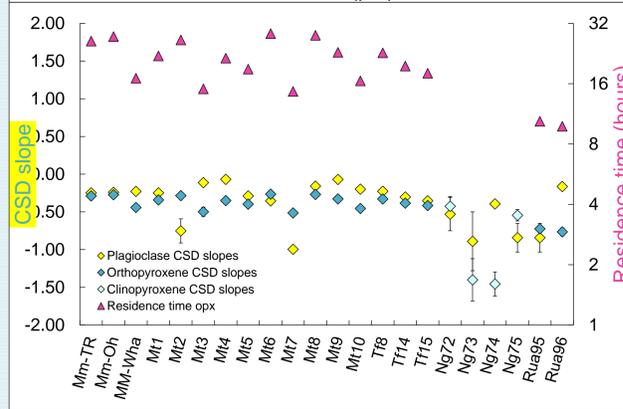
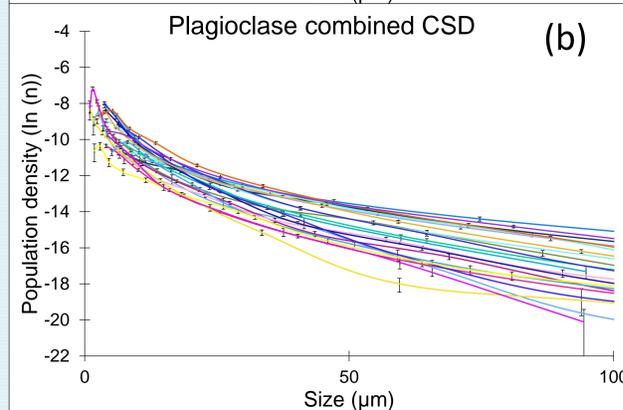
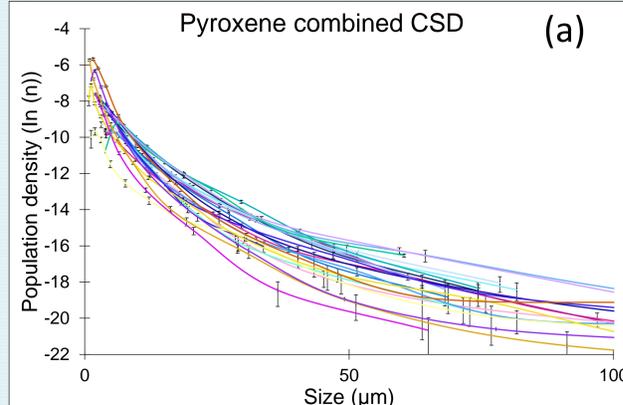


Figure 2: CSD plots of pyroxene (a) and plagioclase (b) microlites from andesitic glass shards of the TgVC tephras

A total of 228 CSDs were generated and combined when within error of each others.

A total of more than 65,000 pyroxene and plagioclase microlites are considered using the semi-automatic method for CSD generation from BSE images [13].

Concave-up CSDs → mixing of crystals populations (>50 µm micro-phenocrysts and microlites)

Figure 3: Pyroxene and plagioclase CSD steepest slopes, and opx residence time. The slope and the growth rate are combined to obtain the residence time

The orthopyroxene growth rate used was calculated for Mangatawai: 3 to 6×10<sup>-11</sup> m.s<sup>-1</sup> [26].

Figure 4: Temperature plotted against the water content for the TgVC andesitic tephras

## Conclusion

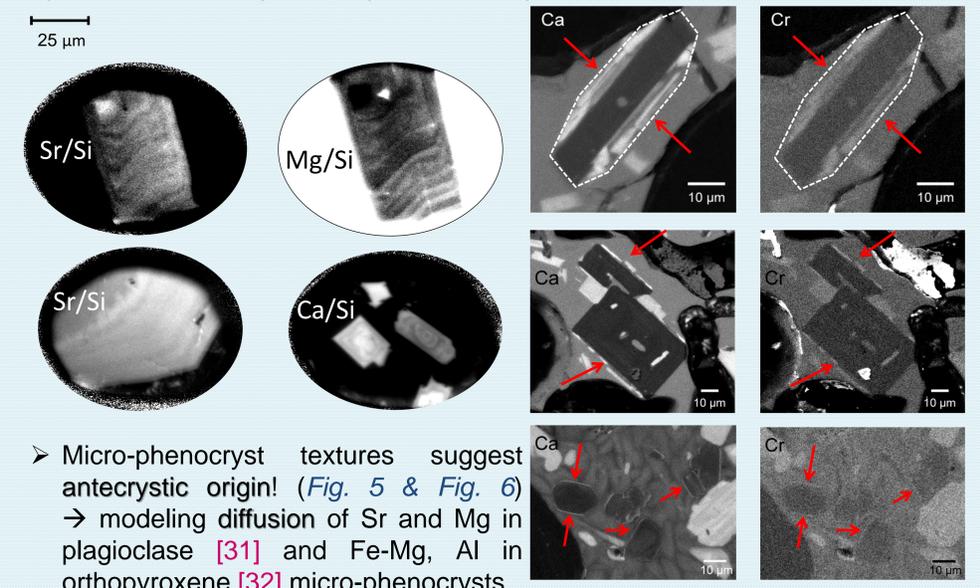
- CSD slopes at smallest sizes (Fig. 3) are similar for different eruptive sequences and eruptive styles (Table)
- Similar slopes for opx and plg → crystallised concomitantly
- Opx residence times of 10±6 to 34±19 hours (Fig. 3) → long timescales for microlites!
- Plg growth rates of 1.4×10<sup>-11</sup> to 1.9×10<sup>-10</sup> m.s<sup>-1</sup> using opx residence times → in the range of experimental growth rates determined in previous works ([3] & [27])
- Our data point to ascent of hot (T<sub>average</sub>=1117 °C), low H<sub>2</sub>O (0.4 – 1.7 wt.% H<sub>2</sub>O 2σ-maxima) magmas in line with data found in TgVC by previous workers ([28] & [29]) and very shallow degassing (i.e. <500m, [30])

Ascent rate of less than 1.4 cm/s if degassing-induced crystallisation...  
**TgVC microlites crystallised prior to H<sub>2</sub>O degassing**

## Future work

- Analyses of melt inclusions in microlites and micro-phenocrysts using 1270 ion probe equipped with SCAPS (Isotope Imaging Laboratory, Hokkaido, Japan) → better constrain on H<sub>2</sub>O and pressure

Figure 5: SCAPS images of plagioclases Figure 6: FE-SEM maps of orthopyroxenes



- Micro-phenocryst textures suggest antecrystic origin! (Fig. 5 & Fig. 6) → modeling diffusion of Sr and Mg in plagioclase [31] and Fe-Mg, Al in orthopyroxene [32] micro-phenocrysts
- Element maps of pyroxene microlites reveal cryptic phase zoning [33] in Ca and Cr layers along the long-axis of pyroxene microlites (Fig. 6)

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